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The distribution of ice wedges was determined from the literature, from mapping permafrost distribution based on climate and surficial geology is polygonal patterns evident on remote sensing imagery, and from our field experience. Ice wedges actively form mainly in the continuous permafrost zone, and are inactive to weakly active in the discontinuous zone (Péwé 1975). Holocene ice wedges, which are limited to the top 3-5 m of permafrost, circum-arctic permafrost map (Brown et al. 1997), which made minor are smaller than large, deep (up to 35 m) syngenetic ice wedges formed during modifications to the initial map by Ferrians (1965). To map permafrost, we the Late Pleistocene. Symbols for abundant ice wedges denote general developed a rule-based model (see color-coded table) that incorporated mean locations, whereas, symbols for sparse Holocene and Late Pleistocene wedges indicate specific areas, though distribution remains poorly known.

Thermokarst landforms are abundant in all permafrost zones (Jorgenson permafrost relationships developed by Kreig and Reger (1982) and our et al. 2008). They are varied, due to differences in temperature, ground ice volume, soil texture, slope, and hydrologic conditions. Abundance of thermokarst is difficult to map because of the wide range in size of features from small pits to large lakes, and similar landforms may have different

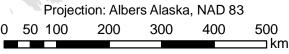
The permafrost zones underlie 80% of Alaska, including continuous (32%), discontinuous (31%), sporadic (8%), and isolated (10%) permafrost. Glaciers and ice sheets occupy 4% of the area.

Many improvements are needed for a better permafrost map, including: a soil texture, permafrost extent, ground ice volume, and primary thermokarst surficial geology map with updated information and better spatial accuracy; more information of terrain/ground ice/temperature/permafrost relationships, more temperature boreholes, and improved spatial models.

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