## Pre-Cloud Aerosol, Cloud Droplet Concentration, and Cloud Condensation Nuclei from the VAMOS Ocean-Cloud-Atmosphere Land Study (VOCALS) Field Campaign

First Quarter 2010 ASR Program Metric Report

December 2009

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Work supported by the U.S. Department of Energy, Office of Science, Office of Biological and Environmental Research

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### 1.0 Introduction

During October and November 2008, Brookhaven National Laboratory (BNL) participated in VOCALS (VAMOS Ocean-Cloud-Atmosphere Land Study), a multi-agency, multi-national atmospheric sampling field campaign conducted over the Pacific Ocean off the coast of Arica, Chile. Support for BNL came from DOE's Atmospheric Science Program (ASP), which is now part of Atmospheric System Research (ASR) following a merger with the Department of Energy's (DOE) Atmospheric Radiation Measurement (ARM) program. A description of the VOCALS field campaign can be found at http://www.eol.ucar.edu/projects/vocals/.

Measurements made from the DOE Gulfstream-1 (G-1) aircraft are being used to assess the effects of anthropogenic and biogenic aerosol on the microphysics of marine stratus. Aerosols affect the size and lifetime of cloud droplets, thereby influencing the Earth's climate by making clouds more or less reflective and more or less long-lived. Climatic impacts resulting from interactions between aerosols and clouds have been identified by the Intergovernmental Panel on Climate Change (2007) as being highly uncertain, and it is toward the improved representation of these processes in climate models that BNL's efforts are directed.

In this, the first of a series of Program Metric Reports, we (1) describe archived data from the DOE G-1 aircraft, (2) illustrate several relations between sub-cloud aerosol, CCN, and cloud droplets pertinent to determining the effects of pollutant sources on cloud properties, and (3) post to the data archive an Excel spreadsheet that contains cloud and corresponding sub-cloud data. The Excel spreadsheet,

<u>Cloud\_Subcloud\_Comparison.xls</u>, contains the data that satisfies the 1<sup>st</sup> Quarter Metric: "Tabulate and make available a set of VAMOS Ocean-Cloud-Atmospheric Land Study showing pre-cloud aerosol size distributions, cloud condensation nuclei (CCN) number concentrations, and cloud droplet distributions." The Excel spreadsheet is derived from data in the main archive but differs in that an association has been made between pre- cloud and in-cloud observations.

### 2.0 Data Files

The parent data set from which the Excel spreadsheet has been derived is archived at: ftp://ftp.asd.bnl.gov/pub/ASP%20Field%20Programs/2008VOCALS/Processed\_Data

Data	Link	Last Modified
AMS and PILS	Aerosol Comp	5/20/2009
Cloud droplet spectra	CASPart	5/20/2009
Drizzle spectra	CIPPart	5/20/2009
Aerosol spectra	DMAPart	7/8/2009
Data summaries and	Data PDF Files	5/21/2009
Chemistry data set & CCN	General	5/21/2009
CCN: Stable conditions	CCN	12/17/2009
PCASP A spectra	PCASPPart	5/21/2009
PCASP B spectra	PCASP BPart	5/21/2009
Position information	PositionPart	5/21/2009

Table 1.Archived data.

Data are archived as ASCII files. A file contains data from one or more instruments for a single flight. The first 39 rows of each file contain metadata including heading, units, contact information, and comments. Flights are identified by yymmddL where L=a, b, or c for the first, second, or third flight in a day. In VOCALS, there were no instances of multiple flights in a day, so L=a. All measurements are referenced to a common one-second time base. The first column of each file contains date in UTC, and the second column contains time in UTC.

Data acquisition rates vary by instrument. In the case of the DMA, PILS, and AMS with acquisition times longer than one second, the instrument reading is repeated at one-second intervals from the start of a data acquisition period to the end. Each measurement is provided on a one-second time base indicated by a file name that ends in \_1.txt and averaged to a 10-second time base indicated by \_10.txt. For both times bases, "Time" is always the beginning of a measurement interval. Typical flight names are 081017a\_1.txt and 081017a\_10.txt for measurements made on the first (and only) flight of Oct. 17, 2008, put on a one-or ten-second time base.

What is in these files is given by the name of the folder that they came from and by the metadata in the files. Files that contain data from instruments that give a size distribution (e.g. CASPart, CIPPart, DMAPart, PCASPPart, and PCASP\_BPart) will have a series of columns giving particle number concentration for each size bin, followed by columns that give ancillary information such as total particle concentration, surface area, and volume. Files in the folders Aerosol\_Comp, General, and PositionPart contain measurements from multiple instruments. CCN data can be found in two places. The "General" files contain a complete record, while the CCN file picks out time periods in which the instruments were run under stable conditions. The use of a common time base facilitates combining data from different files.

Although specifically addressed in the metadata, past experience is that there are two points of confusion. First, all aerosol measurements are presented at ambient conditions of temperature and pressure. Second, AMS concentrations are given at a collection efficiency of one.

Table 2 indicates data availability on 17 flights. All results presented below are obtained from the archived data. Data from flight 081018a were not used because of problems with research power.

Data from Table 1 were used to create an Excel spreadsheet <u>Cloud\_Subcloud\_Comparison.xls</u> containing cloud droplet and sub- cloud (i.e., pre-cloud) aerosol concentration in a form useful for examining the fraction of aerosol that gets activated and thereby contributes to the climatically important first aerosol indirect effect. This file has been archived and is described in Section 3.2 below.



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Time Landing (UTC, hhmm:ss)		18:53:40	16:50:14	15:16:12	19:57:40	16/25:09	17:07 39	16:36:43	17:18:04	19:33:08	18:57:59	16:51:40	16:02:35	16:21:05	16:31:51	16:50:05	16:55:07	16
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#### 3.0 Results

An objective of BNL's participation in VOCALS is to understand the effect that sub-cloud aerosol has on cloud droplet number concentration (CDNC). Activation depends on aerosol composition and size distribution. Cloud dynamics is the other important variable, but this report stops short of that consideration.

#### 3.1 Aerosol Composition and Size Distribution

As an introduction to the VOCALS aerosol data set, composition and size distributions are presented in Figures 1 and 2. Data for these figures are obtained from <u>Aerosol\_Comp</u> and <u>DMAPart</u>, respectively, and encompass all of the below-cloud-level portions of 16 flights.

The chemical composition of sub-micrometer aerosol measured below cloud level was remarkably uniform as a function of longitude as indicated by AMS data presented in Figure 1. Within a longitude grouping there are only small fluctuations (not shown) around an average composition of 25% neutralized  $H_2SO_4$  ( $H_{1.5}(NH_4)_{0.5}SO_4$ ) plus a 10% admixture of organics. The amount and longitudinal gradient of aerosol sulfate, and a consideration of the locations of Cu smelters and power plants in Chile, strongly suggest that the sub-micron aerosol is dominated by anthropogenic emissions.



**Figure 1**. Below-cloud-level aerosol composition determined from AMS measurements averaged into two-degree longitude bins. Aerosol particles measured by the AMS have a maximum diameter of ~0.5  $\mu$ m. Collection efficiency for the AMS equals one. Within the size range of the AMS, chloride concentrations were very low and are not resolved in the pie diagrams. Sea salt and nitrate make a significant contribution to larger particles as indicated by data from the PILS (Particle into Liquid Sampler), which measured the composition of particles smaller than ~ 2.5  $\mu$ m (Lee et al. 2009).

Aerosol size distributions shown in Figure 2 have a two-mode structure characteristic of marine aerosol. This structure has been shown to develop because of in-cloud chemistry forming aerosol sulfate from gas phase precursors (Hoppel et al. 1986). A fraction of aerosol is activated to form cloud droplets. The aerosol that is formed upon cloud droplet evaporation is larger than the initial aerosol because of mass gained by in-cloud chemistry. Thus the accumulation mode centered at  $\sim$ 170 nm represents aerosol that has been activated at some point its history, while the Aitken size particles centered at 40–60 nm represent aerosol that has not been activated.



**Figure 2**. Size distribution of below-cloud-level aerosol as a function of longitude, which is very nearly proportional to distance from the Chilean coast.

#### 3.2 Sub-Cloud Aerosol and Comparison with Cloud Droplet Number Concentration

Our approach is to identify a time period with unbroken cloud and compare the number concentration of cloud droplets to the number concentration of aerosol particles below the cloud in air that will eventually be circulated into the cloud. In the VOCALS environment, below-cloud aerosol concentrations are relatively homogeneous on horizontal spatial scales of 10 seconds of km, through the depth of the boundary layer, so an exact equivalence between a cloud location and precloud aerosol is not required. In so far as there are inhomogenieties in the spatial distribution of aerosol, these will be greatest near the coastline.

Because we are interested in the size distributions of particles smaller than 100 nm, we rely on the DMA (size range 15–440 nm) and consider average values over the nominal 1-minute time period for the DMA to complete a size scan. Having identified a cloud penetration (criteria that liquid water content > 0.05 g/m<sup>3</sup>) we identify below-cloud DMA scans that satisfy screening criteria including the specification that no one-second period contains liquid water greater than 0.01 g/m<sup>3</sup> and that the below-cloud point be within 15 km of the cloud point. In order to eliminate clouds with holes, cloud droplet concentrations are calculated for that fraction of the 1-minute DMA scan that has one- second LWC (liquid water content) > 0.05 g/m<sup>3</sup>. This procedure yielded 543 in-cloud data points (i.e., almost 8 hours of data) in which a cloud and corresponding sub-cloud DMA scan could be identified. In the usual case that more than one sub-cloud point corresponded to a cloud penetration, results were averaged over the several below cloud points.

The Excel spreadsheet <u>Cloud\_Subcloud\_Comparison.xls</u> added to the data archive as part of this report contains the time and position of the in-cloud DMA scans. On the cloud side, LWC, CDNC, and a size spectrum ( $dN/dLogD_p$ ) of cloud droplets are provided. For the corresponding sub-cloud aerosol, number concentration, size spectra ( $dN/dLogD_p$ ), and CCN at 0.2% supersaturation are given.

Cloud droplet number concentration as a function of pre-cloud aerosol is shown in Figure 3. Results are within the range of values obtained at different locations as depicted in Figure 5 of Ramanathan et al. (2001). The common tendency is for CDNC to increase linearly with sub-cloud aerosol ( $N_a$ ) at low aerosol concentration and then for a less steep increase at higher values of  $N_a$ . The two sets of outliers with CDNC below the main groupings come from measurements where the aerosol was unusually non-homogeneous.



**Figure 3**. Comparison of cloud droplet number concentration with sub-cloud aerosol for the entire VOCALS campaign. Data points that are connected are sequential in time. Breaks appear because aircraft flight patterns consisted of transects and vertical profiles in and out of cloud.

If the only independent variable in our data set is sub-cloud aerosol concentration (i.e., meteorological variability can be ignored), then Figure 3 provides an important result on how clouds respond to a perturbation in anthropogenic aerosol, which with known liquid water path can be used to determine changes in cloud albedo. The color coding by LWC qualitatively suggests that in this case the relation between aerosol and cloud droplets is not strongly affected by LWC.

#### 3.3 Aerosol Activation

Figure 4 shows the number concentration of aerosol activated at ~ 0.2% supersaturation as a function of the total aerosol number concentration.



**Figure 4**. Number concentration of aerosols activated at a nominal supersaturation of 0.2% versus the total aerosol number concentration measured by DMA.

Aerosols with a composition illustrated in Figure 1 are expected to have a CCN forming potential comparable to  $(NH_4)_2SO_4$ , which is among the most easily activated, atmospherically relevant aerosols. An  $(NH_4)_2SO_4$  particle with a dry diameter of 90 nm activates at a supersaturation of 0.2% (Seinfeld and Pandis, 1998). We have compared the CCN measurements at 0.2% supersaturation with DMA spectra to determine a "critical diameter" above which 100% activation accounts for the CCN observations. For the data points in Figure 4, we obtain an average critical diameter of 97 nm (median = 91 nm) with a standard deviation of 44 nm, very close to that predicted for  $(NH_4)_2SO_4$ , albeit with considerable scatter. The scatter is probably not due to composition differences. In contrast to other locations where aerosols have large and variable insoluble or sparingly soluble components (Wang et al. 2008), aerosol chemical composition in VOCALS is relatively constant (Figure 1).

Given that a Hoppel minimum is formed by preferential activation of large particles, it is relevant that the size distributions in Figure 2 show a separation between accumulation mode and the smaller, less easily activated Aitken mode between 70 and 100 nm.

If CDNC data from Figure 3 were added to Figure 4, one would see that CDNC is generally less than CCN at 0.2% supersaturation. This could be because the controlling supersaturation is less than 0.2% or it could have a cloud dynamic explanation. Both possibilities are being investigated.

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