Linking the uncertainty of low frequency variability in tropical forcing to regional climate change

Chris E. Forest and Wei Li, Penn State Joe Barsugli, CIRES, U. Colorado DOE CESM PI Meeting 19-22 Sept 2011 Funded by: DOE Regional and Global Climate Modeling Program Thanks: Prashant Sardeshmukh, Lisa Murphy

Research Goals

- Explore model uncertainty using regional climate change from atmospheric GCM response to SST patterns (i.e., ignore local forcing)
- 2. Define Global Teleconnection Operator, **GTO**, to compare model responses at multiple scales
- 3. Introduce **Random Perturbation Method** as an alternative to the **Patch Method** to estimate GTO

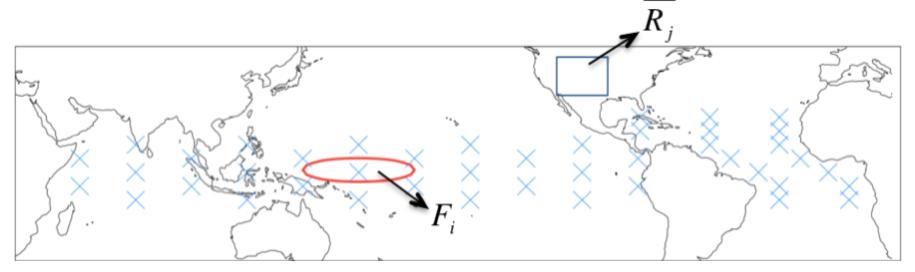
4. Identify model structural differences

that lead to different GTO patterns (convection, resolution, etc.) and influence predictability

Can we quantify uncertainty in model response at <u>regional scales</u>?

- Create *idealized* experiments with "known" surface forcings to provide metrics (or framework) for comparing model response <u>at regional scales</u>
- Two options:
 - Patch Experiments (indiv. tropical SST anomalies)
 - Random SST anomalies applied across both tropical and extratropical oceans (Random Patch Method, **RPM**)

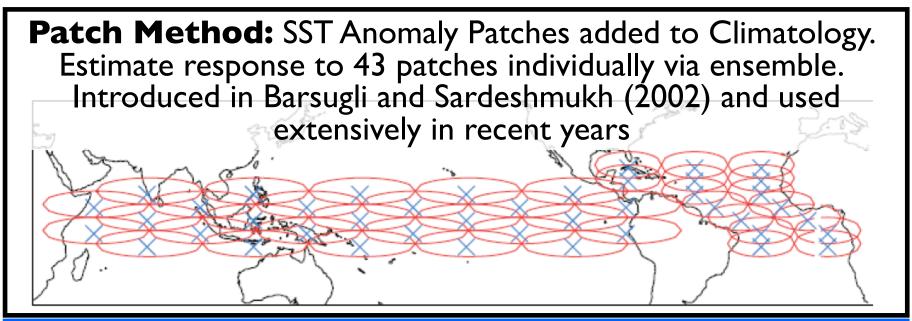
Basic Method: Estimate the ensemble-mean response, R_j , to the Δ SST forcing, F_i



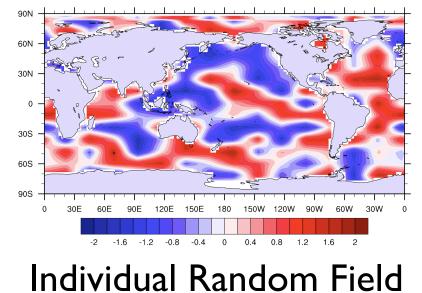
Estimate Global Teleconnection Operator, K_{ii} , from:

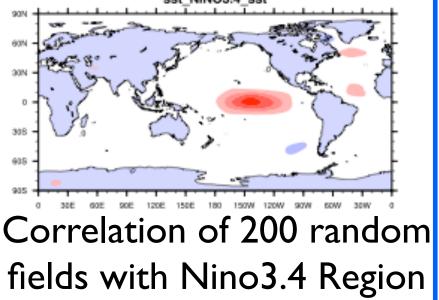
$$\overline{R_j} = K_{ij} \cdot F_i + arepsilon$$

Repeat this estimate for all SST anomaly locations. Follows: Barsugli and Sardeshmukh (2002, J. Climate)



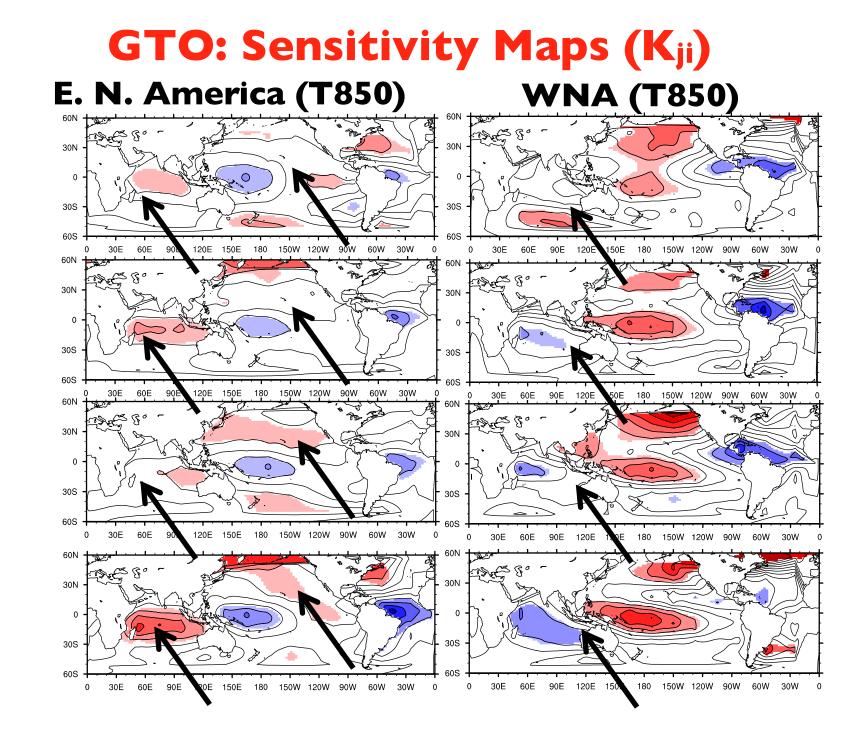
Random Perturbation Method (RPM) Ensemble response to random field (~10x more efficient)





Experiments

	Models							
S		NCAR CAM3					GFDL AM2	HadAM3
olutions		CAM3.I	CAM3.5	CAM4.0	CAM5.0		FV2.0x2.5	
	T42	Х					0	0
	Т85	Х						
esol	FV1.9×2.5	Х	Х	Х	Х			
	FV4x5				0			
		x: done o: To be done						
SST Scenario		SST patterns				Source		
C20C		CMIP3 Ensemble –mean AOGCM PC response to Climate of the 20 th century runs				PC	MDI CMIP3	archive
future		CMIP3 Ensemble-mean SST(SRES AIB) PCN					MDI CMIP3	archive

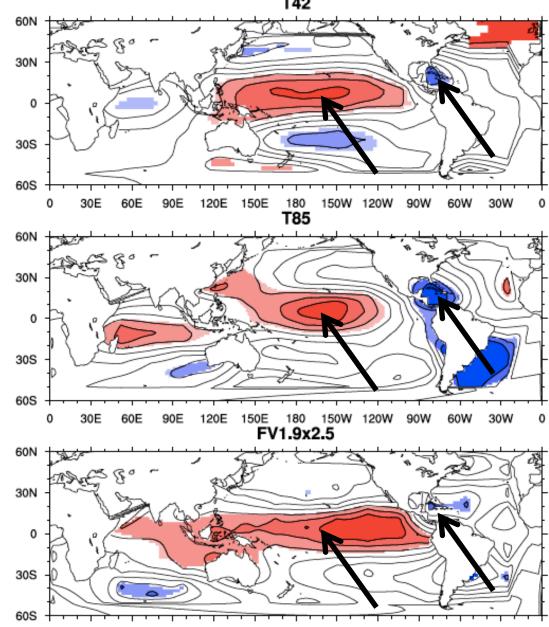


CAM3 **CAM3.5 CAM4.0 CAM5.0**

GTO: Sensitivity Maps (K_{ji})

CAM3.I Central North America (Precip)

Three Resolutions: T42, T85, FVI.9x2.5



Summary of Results so far

• Estimated GTO for multiple CAM models

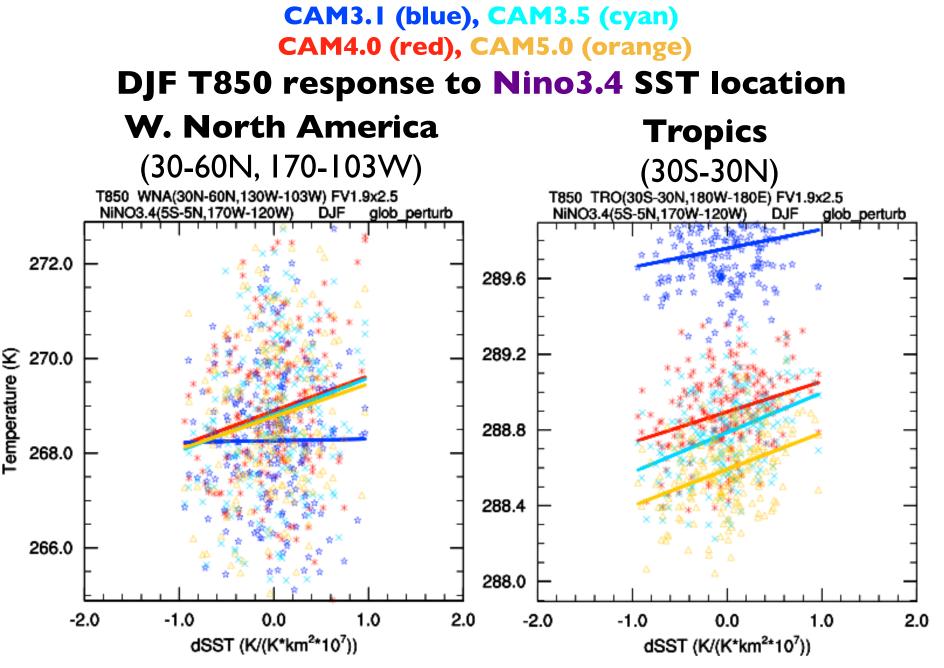
- **Structure**: CAM3.1, CAM3.5, CAM4.0 and CAM5.0 (FV1.9x2.5)
- **Resolution**: CAM3.1 @ T42, T85, and FV1.9x2.5

• Sensitivity of larger scales is more robust

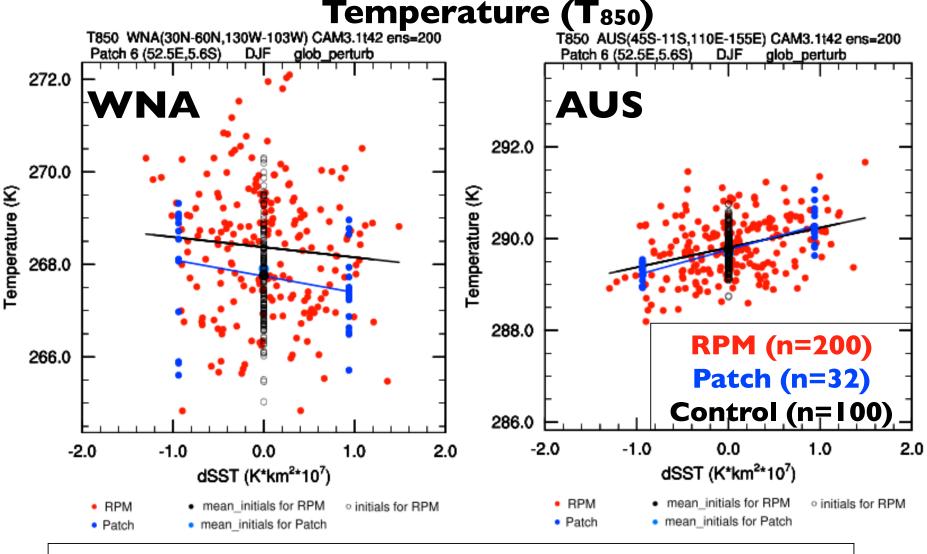
- Global, Tropical, and Extra-tropical targets
- Unforced variability contributes significantly at regional scales and more at high latitude
- Verified with linear reconstruction via AMIP runs and observed SST
- Resolution & Grid <u>may</u> be more important than model physics

Extra slides

NCAR CAM Inter-Comparison:



GTO: RPM versus Patch



Patch Region: 52.5E, 5.6S (Indian Ocean) Response: WNA (left), Australia (right)

Define metric: GTO Global Teleconnection Operator

• Estimate **K**_{ji} from the following linear relation:

$$R_j = K_{ji} F_i + b_{0j} + \varepsilon$$

- where \mathbf{R}_{j} is the model response of targeted region of interest, $\mathbf{F}_{i} = (\Delta SST)_{i}$ is the SST anomaly at a given location, \mathbf{b}_{0j} is the intercept and $\boldsymbol{\varepsilon}$ is the error.
- K is the linear operator that estimates the local climate change response to an SST anomaly over a specified region.

We call **K** the Global Teleconnection Operator, **GTO**.

 The GTO can be used to estimate the response to observed SST patterns via K*SST(obs). (i.e., an empirical Green's function)

"Giorgi" Regions

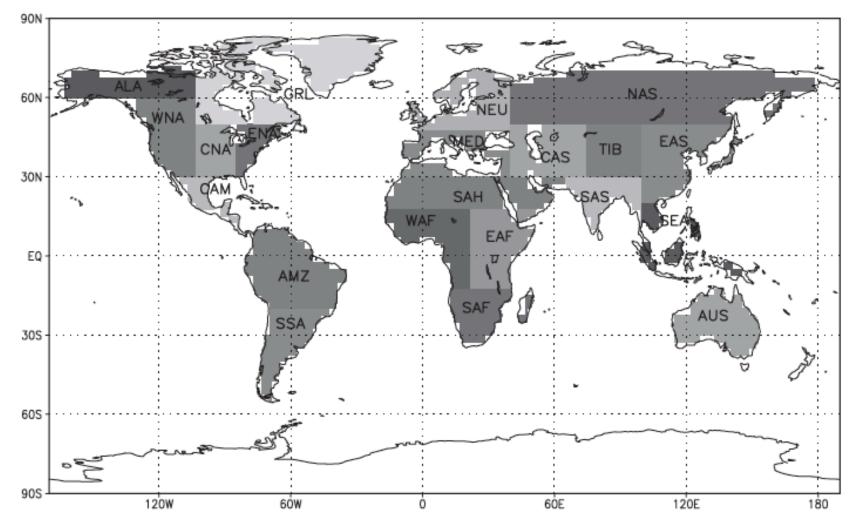
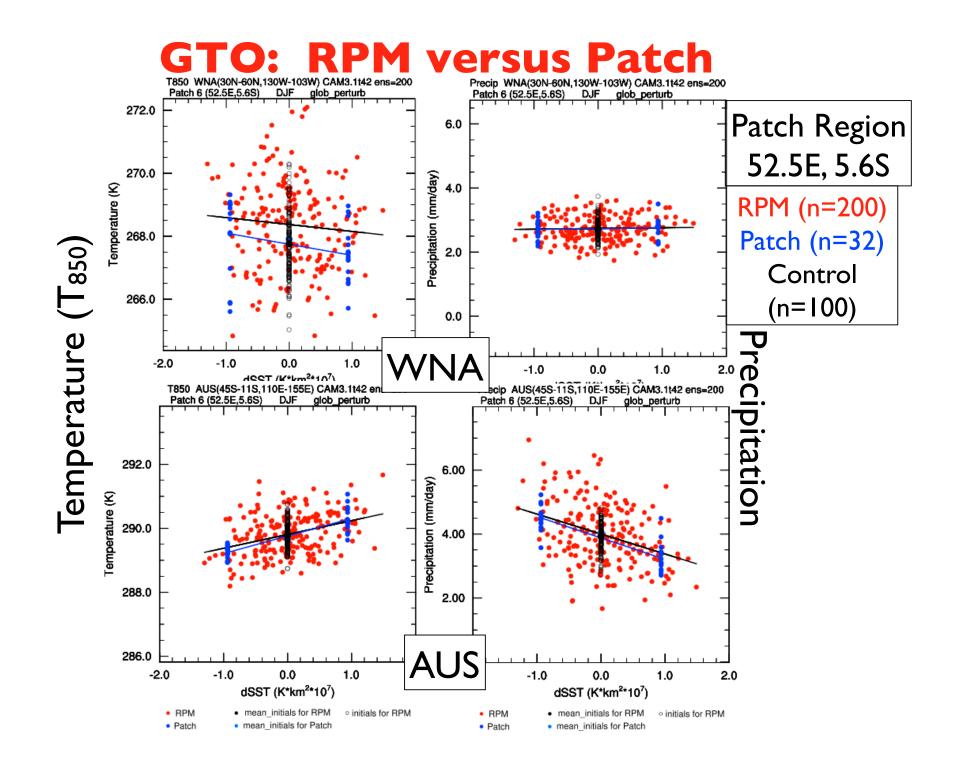
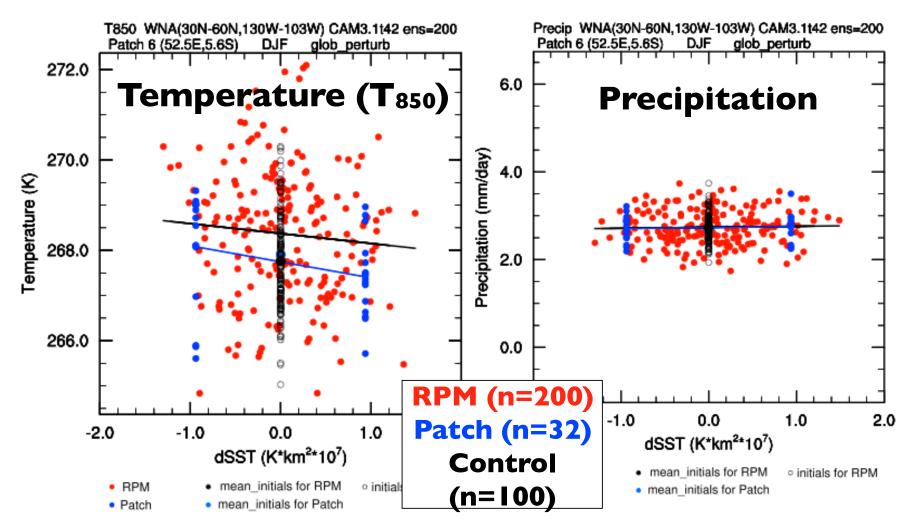


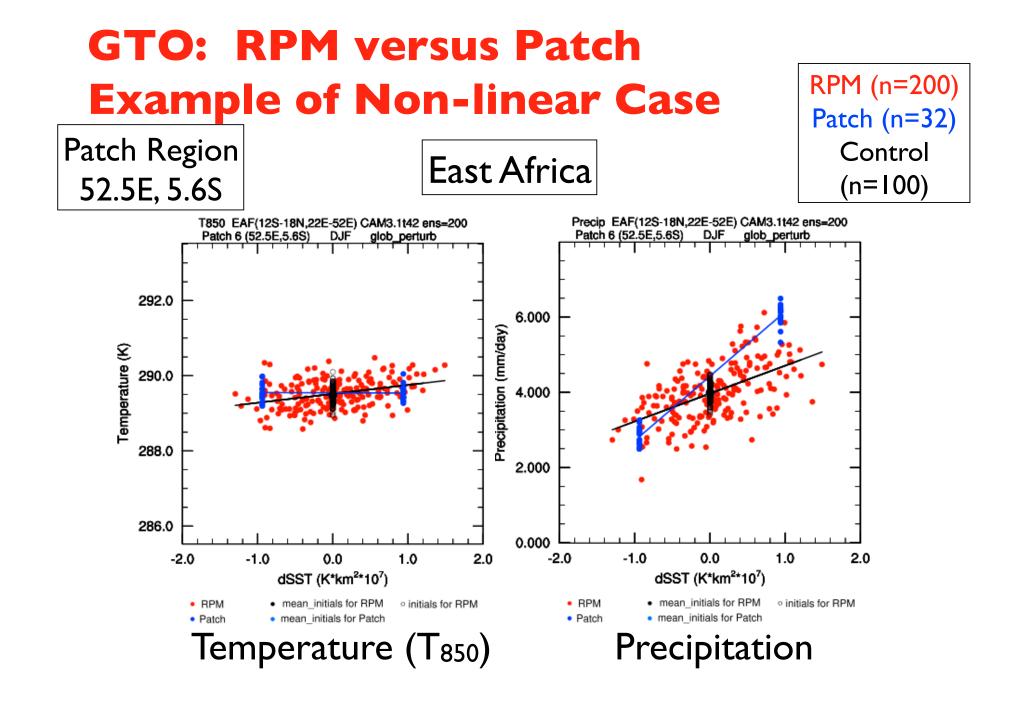
Fig. 1 Regions used in the analysis presented in this work (see Table 2 and text)

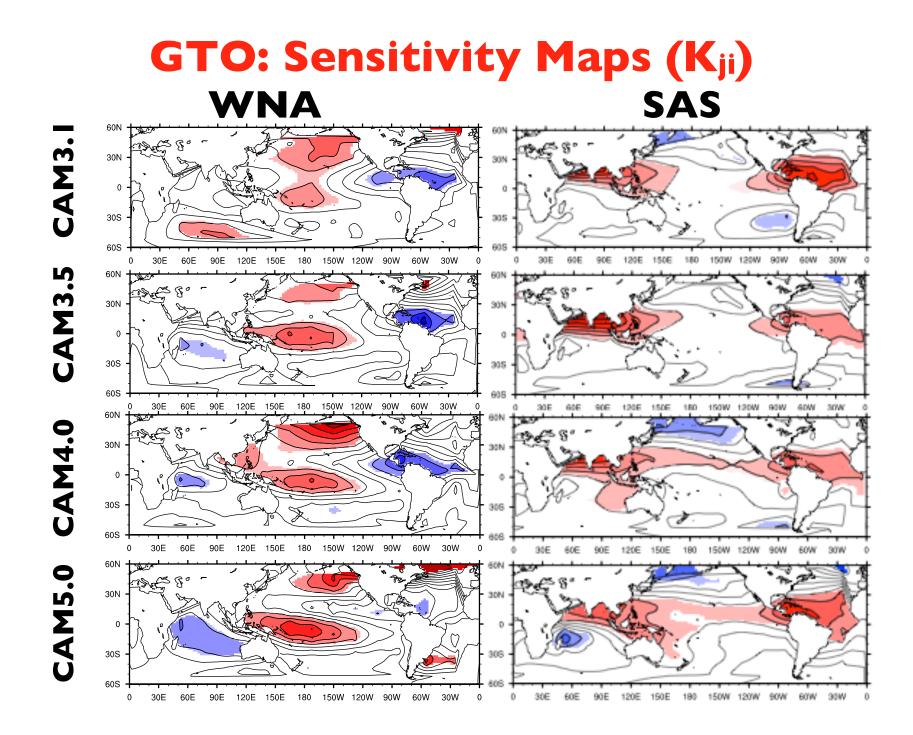


GTO: RPM versus Patch



Patch Region: 52.5E, 5.6S (Indian Ocean) Response: Giorgi Region: W. North America





GTO: Sensitivity Maps (K_{ji}) N. America EAS

