

Basin-Scale Assessment of Gas Hydrate Dissociation in Response to Climate Change

Understanding Subsurface, Ocean, and Atmospheric Processes
to Assess the Consequences

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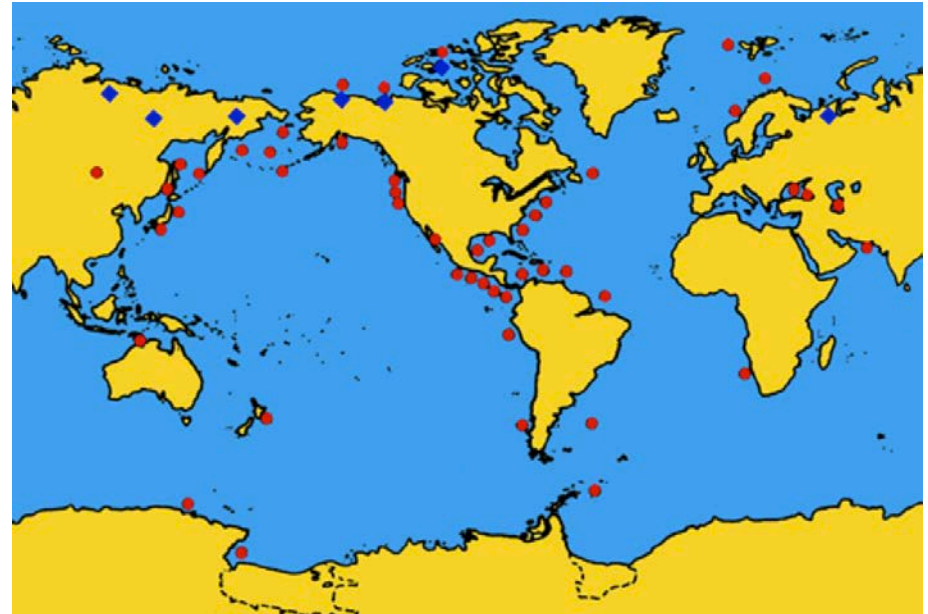
Introduction and Purpose

The goals of our research are:

1. To investigate the effect of rising water temperatures on the stability of oceanic hydrate accumulations and to determine the conditions under which methane release may occur
2. To identify geochemical effects within the water column that control the fate of hydrate-derived methane
3. To estimate the global quantity of hydrate-originating carbon that could reach the atmosphere
4. To examine, quantitatively, the possibility of hydrate-related climate feedbacks

Global Organic Carbon Distribution & Gas Hydrates

- Enormous amount of organic matter trapped in arctic systems
- **Hydrates** store **huge** amounts of methane
- Huge deposits in permafrost and in the oceans (**oceans >> permafrost**)
- Hydrate carbon significant on a global scale, and over long times (Archer, 2007; 2009)
- Recent studies suggest 3000 – 5000 Gt (Archer, 2009; Wallman, 2011)
- Much hydrate is **deep** and **low-saturation**
- **Continental margins** are the key

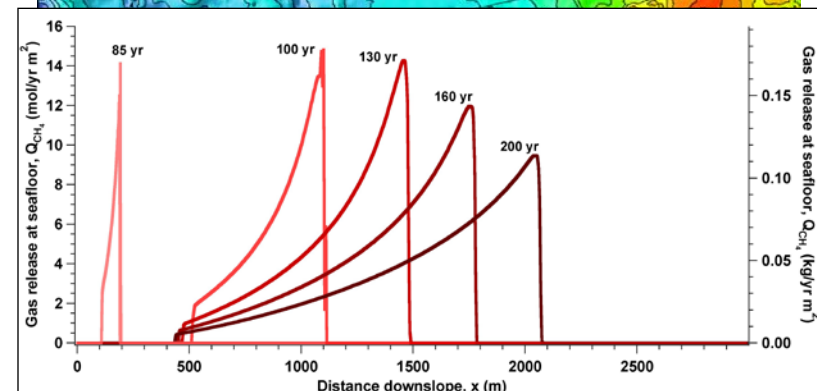
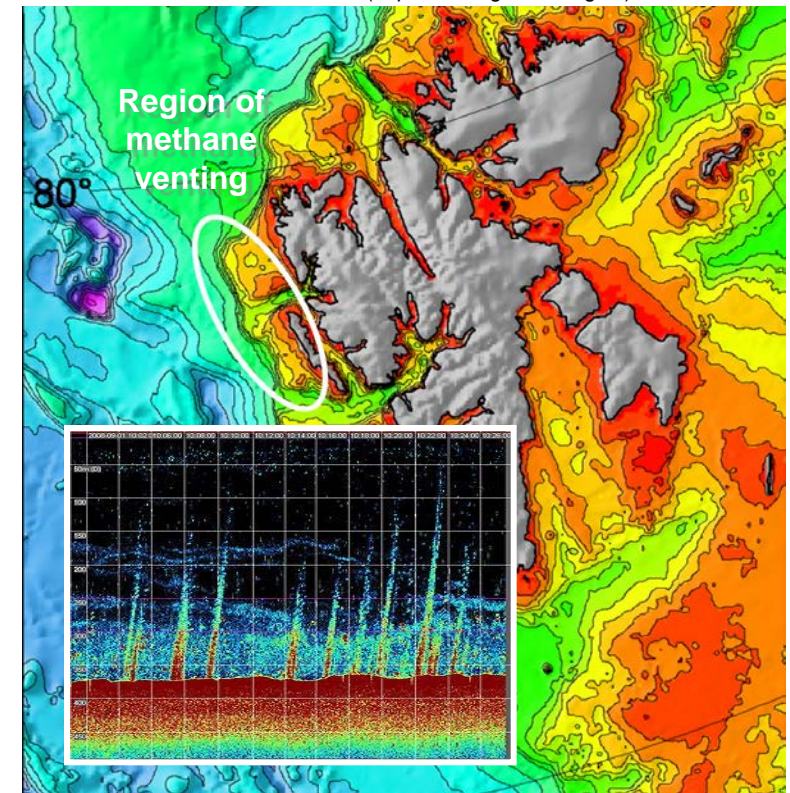


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- Much hydrate is deep and low-saturation
- Continental margins are the key
- The arctic is warming
- Methane plumes are appearing near the edge of the GHSZ in regions undergoing measured ocean warming (Westbrook et al., 2009)
- Simulations of shallow hydrates subjected to warming show plumes forming at the landward GHSZ limit¹

¹Reagan, M.T. and Moridis, G.J., "Large-Scale Simulation of Oceanic Gas Hydrate Dissociation in Response to Climate Change," *Geophysical Research Letters*, 36, L23612, 2009.

from NOAA/IBCAO (<http://www.ngdc.noaa.gov/>)

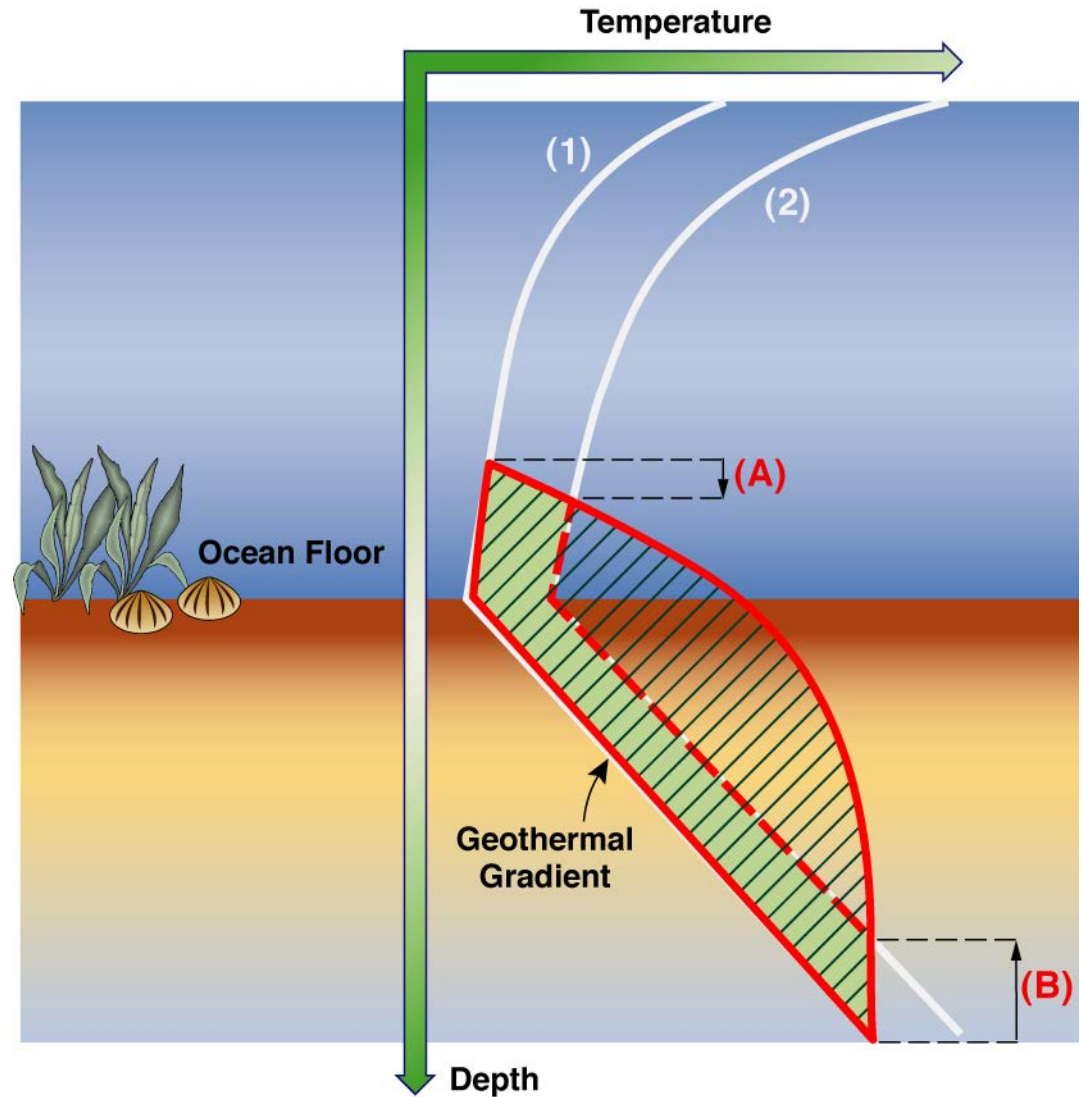


Oceanic Gas Hydrates: Dissociation

Climate change alters ocean temperature (and geothermal gradient)

Decreases hydrate stability region

Methane release to ocean by hydrate dissociation



Oceanic Gas Hydrates: Dissociation

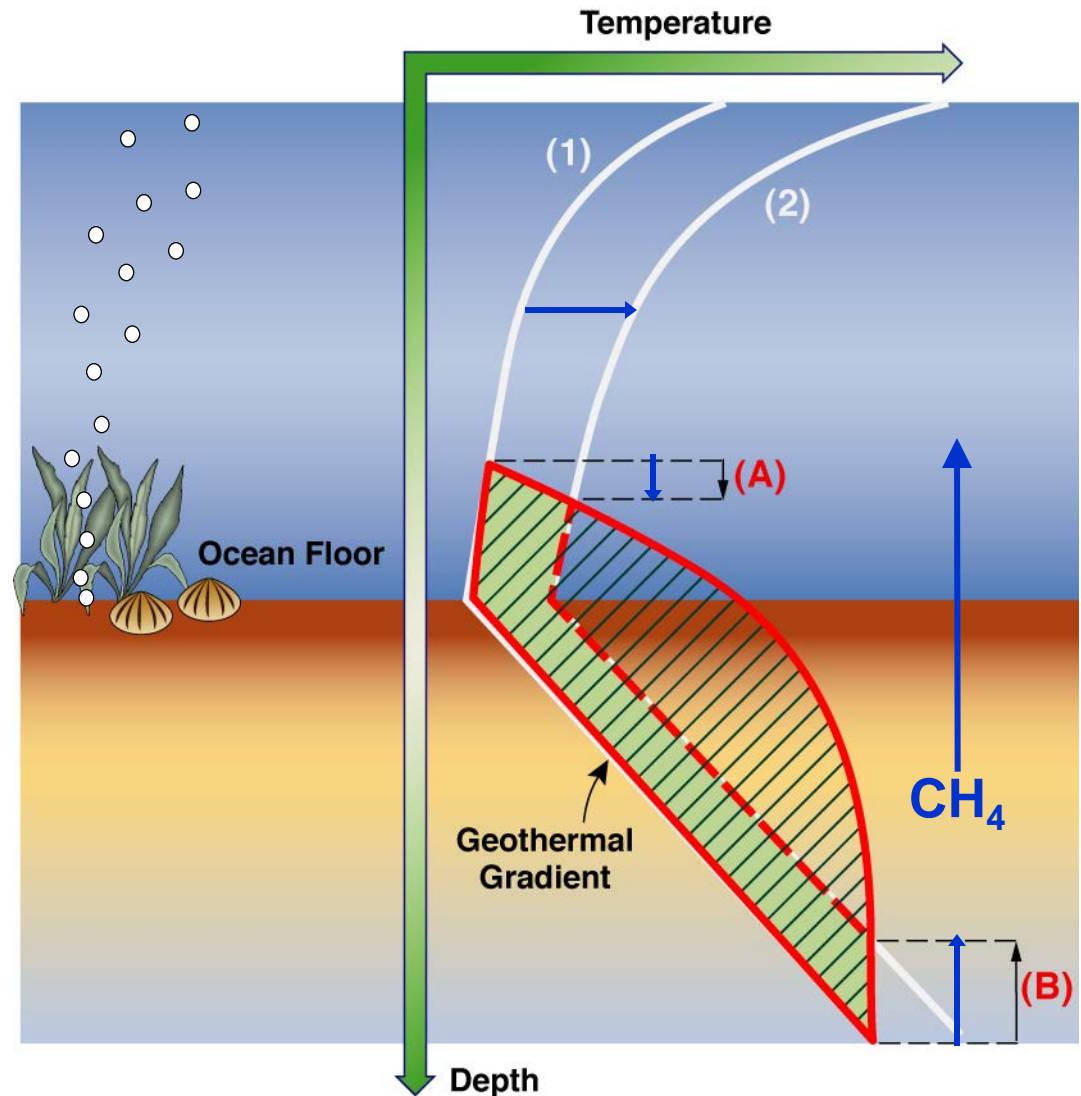
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What happens between (1) and (2)?

What is the fate of the methane?



Oceanic Gas Hydrates: Dissociation

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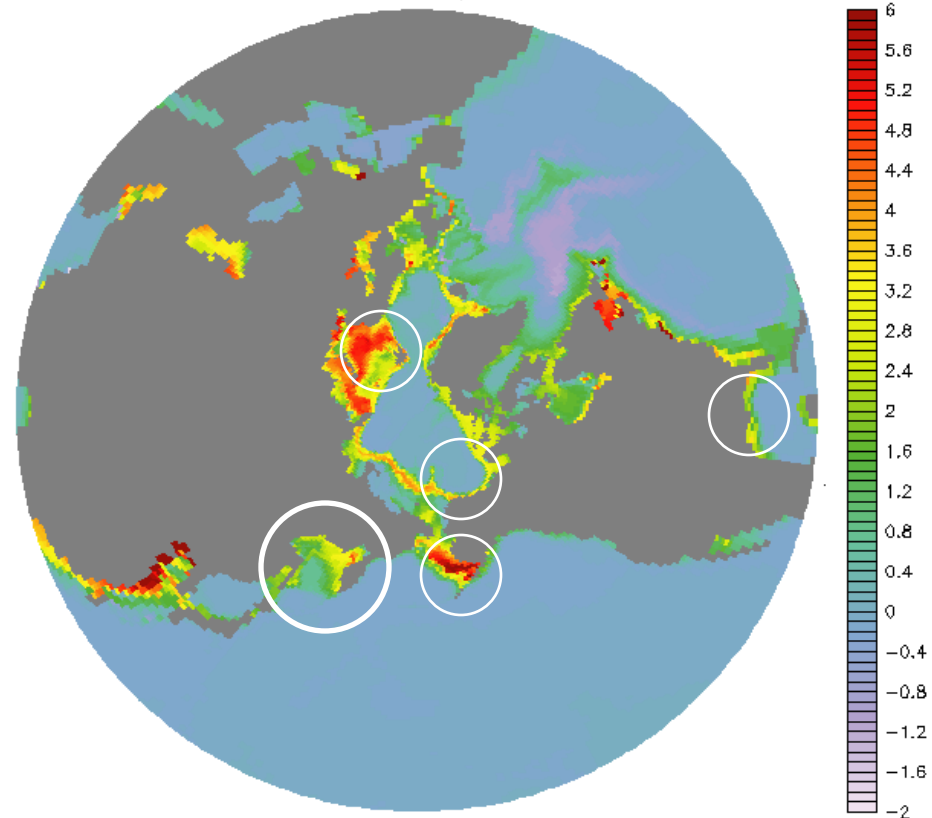
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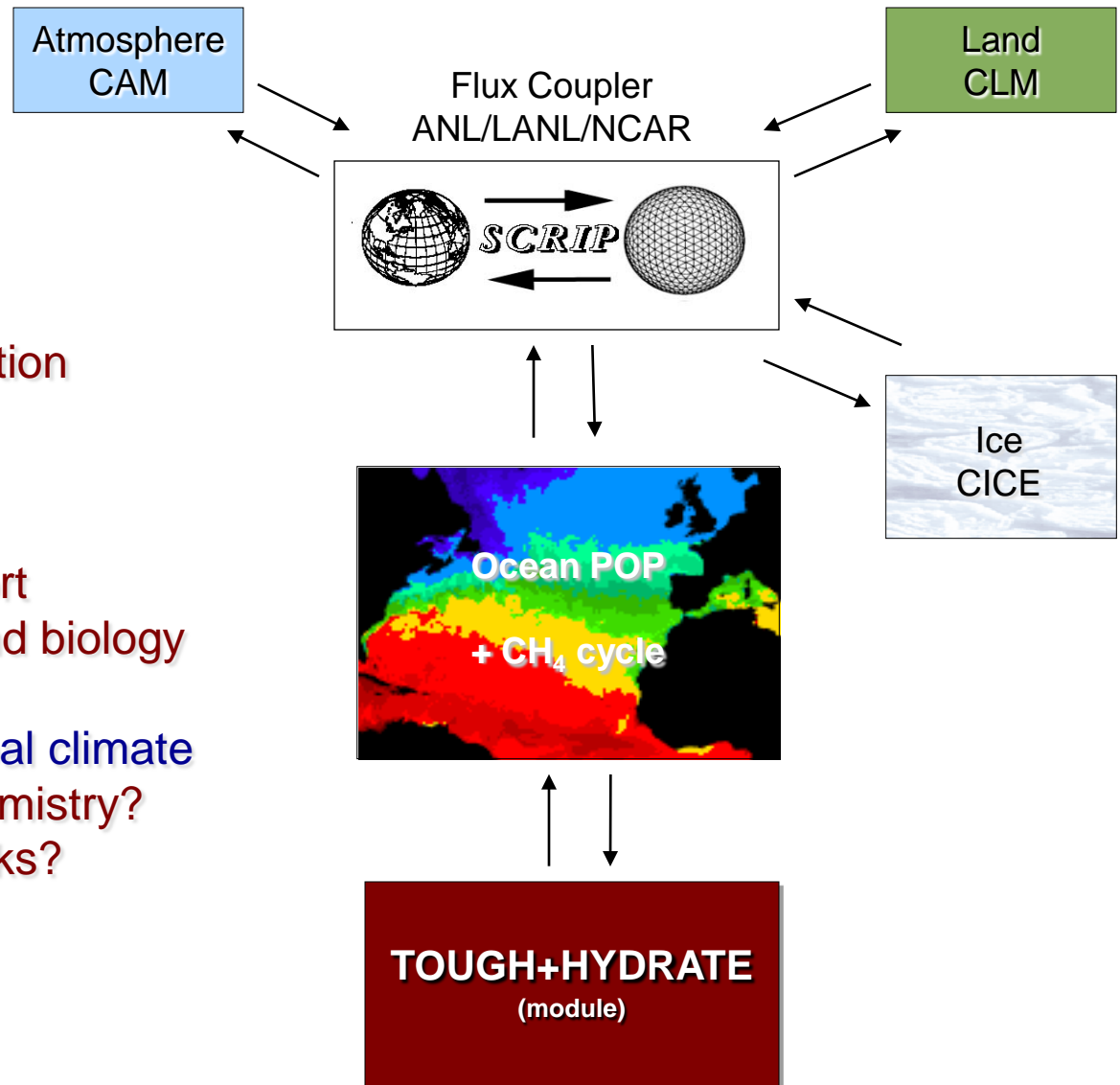
What is the fate of the methane?

change in ocean bottom temperature (degrees C)
model year 2100 – model year 2000
CCSM A1B scenario, realization 2



Outline

- 1) Sub-seafloor
 - Hydrate dissociation
 - Fluid transport
- 2) Ocean water column
 - Methane transport
 - Geochemistry and biology
- 3) Atmosphere and global climate
 - Atmospheric chemistry?
 - Positive feedbacks?



Outline

1) Sub-seafloor

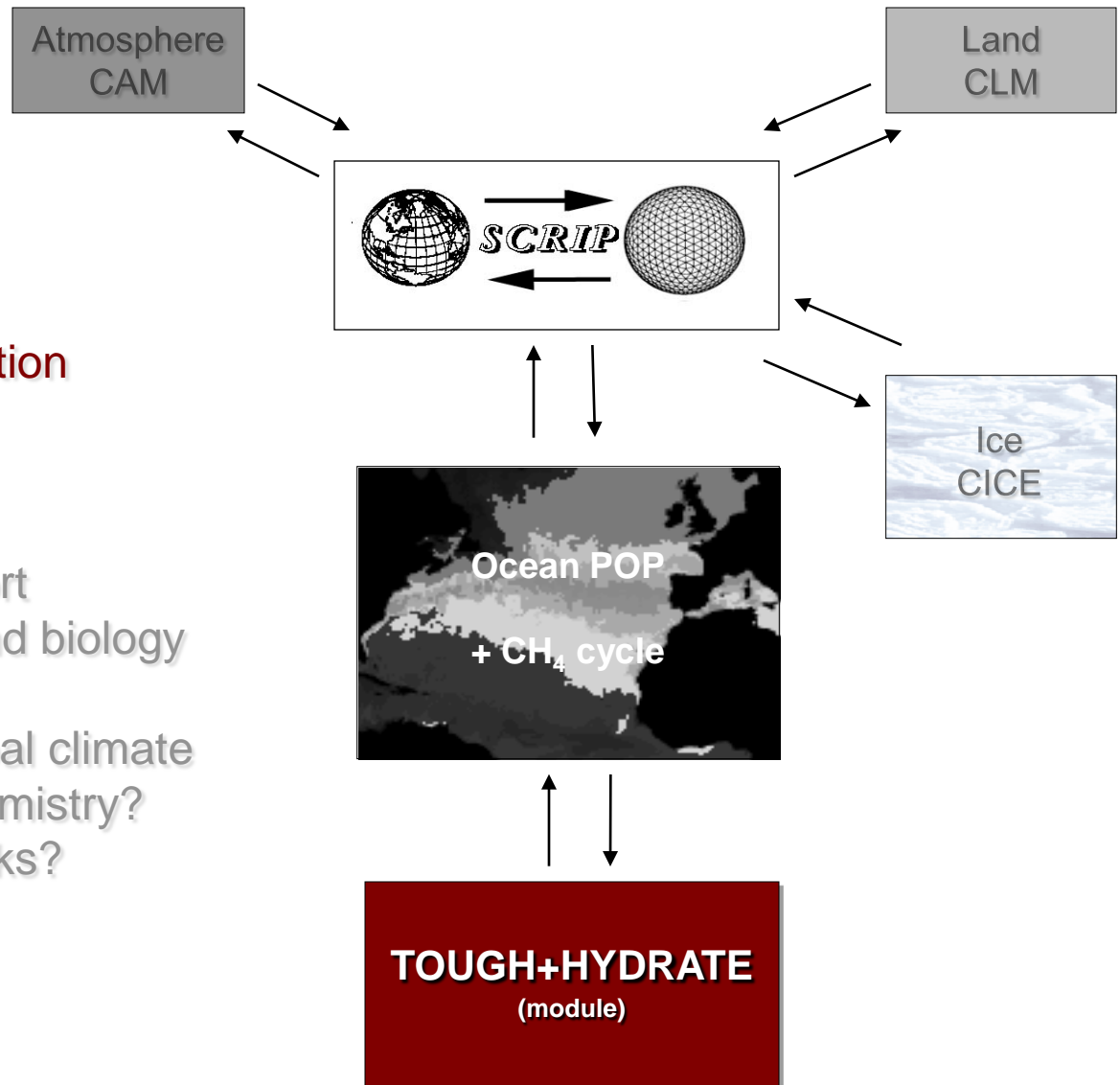
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The Numerical Model

TOUGH+HYDRATE code (Moridis *et al.*, 2008):

- Describes fully coupled non-isothermal hydrate dissociation/formation, CH₄ flow, and phase behavior in porous media
- Descendent of the TOUGH family of codes (YMP, oil & gas, CO₂ sequestration, hydrology)
- Validated by 1) analytical solutions, 2) gas production from permafrost deposits in the field, 3) dissociation and thermal behavior in lab experiments

Components

- (1) H₂O
- (2) CH₄
- (3) Hydrate
- (4) Salt
- (5) Inhibitor
- (6) Heat

Phases

(1) Aqueous
H₂O, CH₄, salt, I

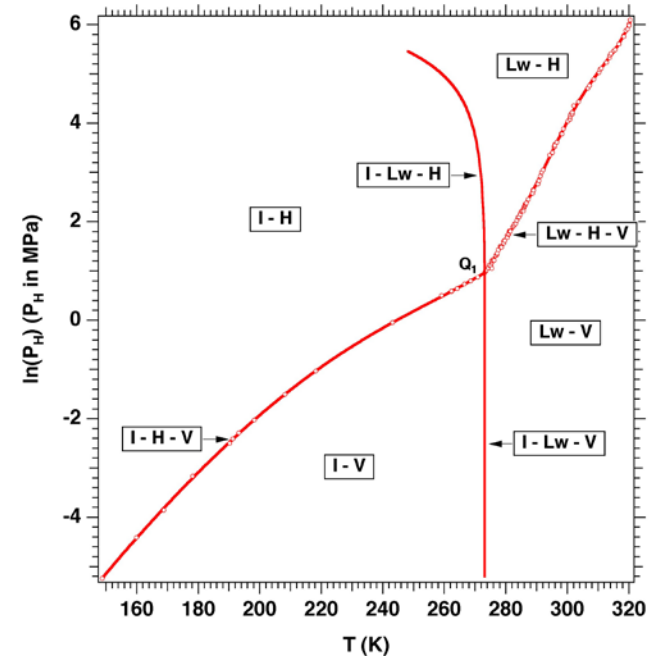
(2) Gas
CH₄, H₂O, I

(3) Solid-Hydrate
CH₄·N_mH₂O

(4) Solid-Ice
H₂O

30 possible phase combinations

Latest hydration *P-T* relationships

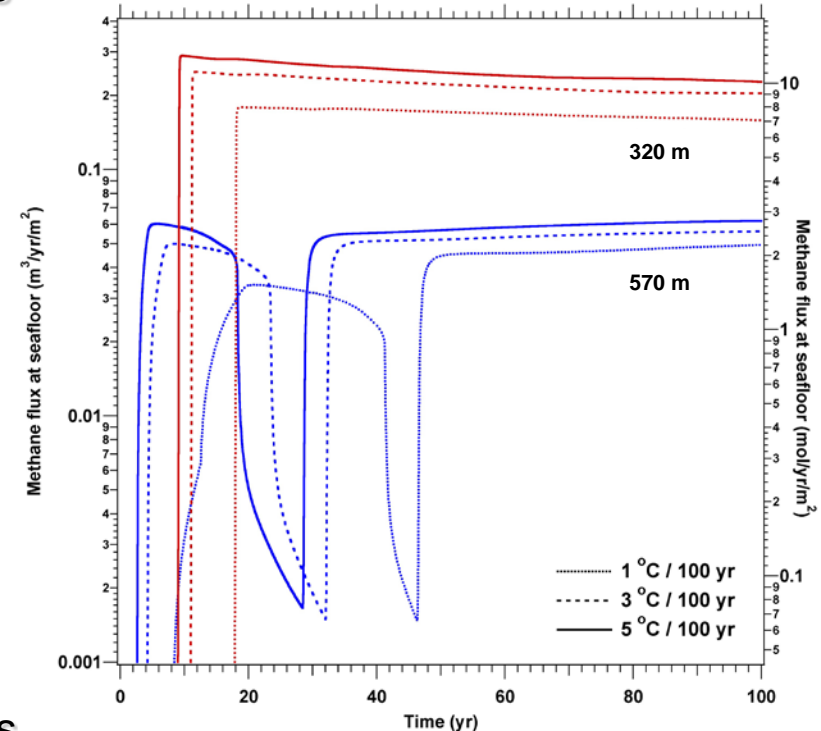


Hydrate Stability/Dissociation

- Previous work^{1,2,3,4} studied **Class 4** (disperse) deposits, $S_H \square = 0.03-0.10$, 1 °C to 5 °C warming
- Provided view of GHSZ extent, localized dissociation estimates, localized fluxes, sensitivity
- 1D columns^{2,3} and 2D continental slope deposits^{1,4}

Multiple scenarios tested^{1,2,3,4}:

- Deep ocean, 1000 m: **stable**
- GoM, 570 m: **unstable?**
- Arctic shelf, 320 m: **unstable**
- Barents Sea, 390 m: **unstable**
- Releases are sustained, but not explosive^{1,2}
- Constrained by thermal effects and sediment transport properties^{3,4}
- Sea level rise (10m) **not mitigating**³
- Results can be generated for any z/T /geological location via linkable subroutines



²Reagan, M.T. and G.J. Moridis, oceanic gas hydrate instability and dissociation under climate change scenarios, *Geophys. Res. Lett.*, **34**, L22709, doi: 10.1029/2007GL031671, 2007.

³Reagan, M.T. and G.J. Moridis, The dynamic response of oceanic hydrate deposits to ocean temperature change, *J. Geophys. Res. Oceans*, **113**, C12023, doi:10.1029/2008JC004938, 2008.

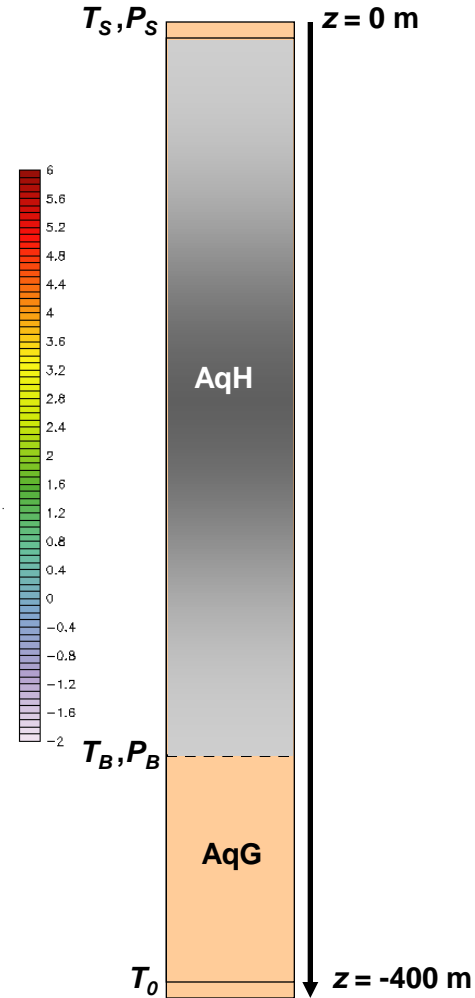
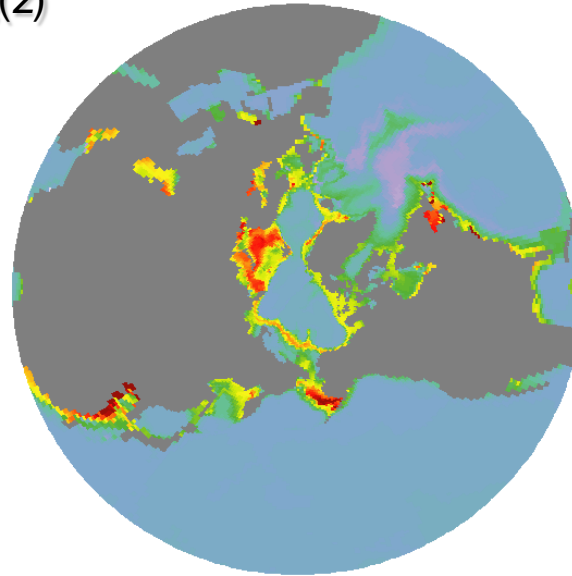
⁴Reagan, M.T., Moridis, G.J., Elliott, S.M., and Maltrud, M., "Contributions of Oceanic Gas Hydrate Dissociation to the Formation of Arctic Ocean Methane Plumes," *J. Geophys. Res. Oceans*, 2011JC007189, in press.

Integration: Basin-Scale Assessments

- 1D sediment column; thermal, chemical, hydrostatic equilibrium, fine discretization: $dz = 10 \text{ cm} - 1 \text{ m}$
- $k < 1 \text{ mD}$, $\phi = 0.45 - 0.55$, consolidated sediments
- Hydrate exists from $z = -1 \text{ m}$ to $T_B(z)$, $P_B(z)$
- $S_{H\Box} = 0.01-0.10$ (~0.03%-3 %vol)
- Okhotsk: $T_s = 0 \text{ }^\circ\text{C} - 2 \text{ }^\circ\text{C}$
- Arctic: $T_s = 0.5 \text{ }^\circ\text{C} - 1 \text{ }^\circ\text{C}$

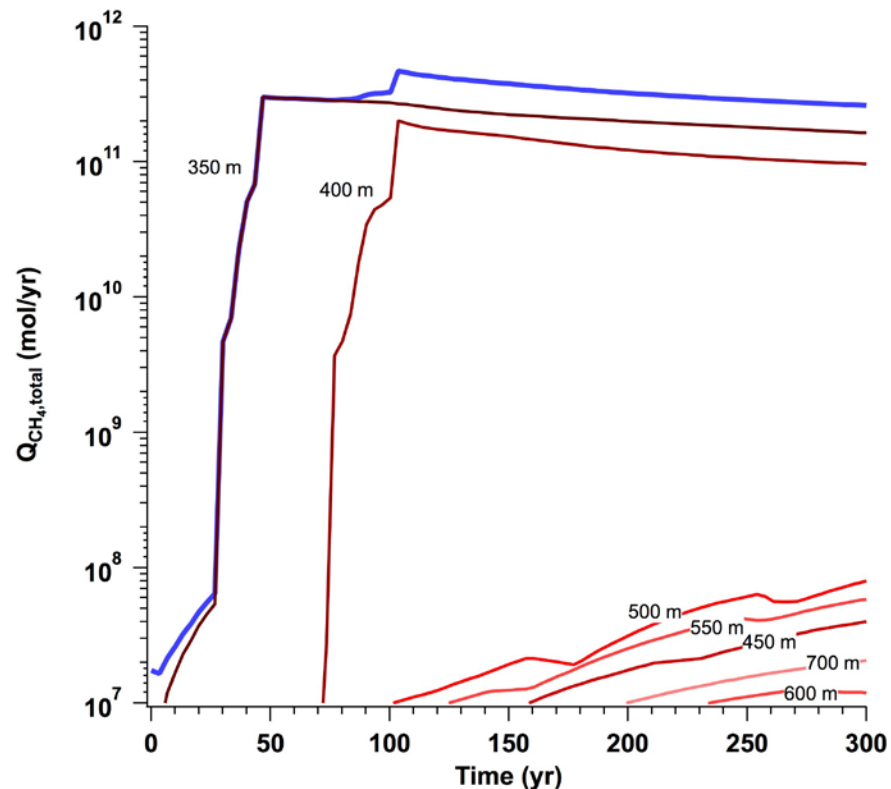
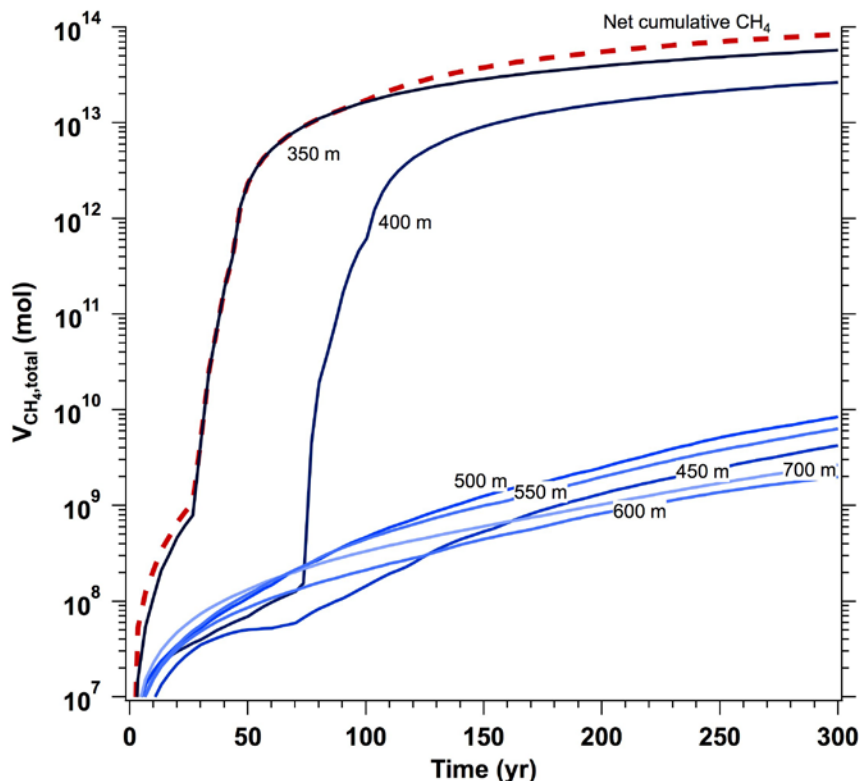
- $\Delta T = 1, 3, 5, 7 \text{ }^\circ\text{C}/100 \text{ yr}$, or $\Delta T = f(z)$
- 100 yr warming, then $T_s = \text{constant}$

- Simulate CH_4 flux vs. t at each location
- Fluxes a function of depth, T_s , ΔT
- Perform 1-D release simulations on 40-min grid, (ETOPO2) at 25 m depth increments
- Integrate total methane release/flux over bathymetry
- Estimate basin-scale CH_4 emission



Integration: Sea of Okhotsk

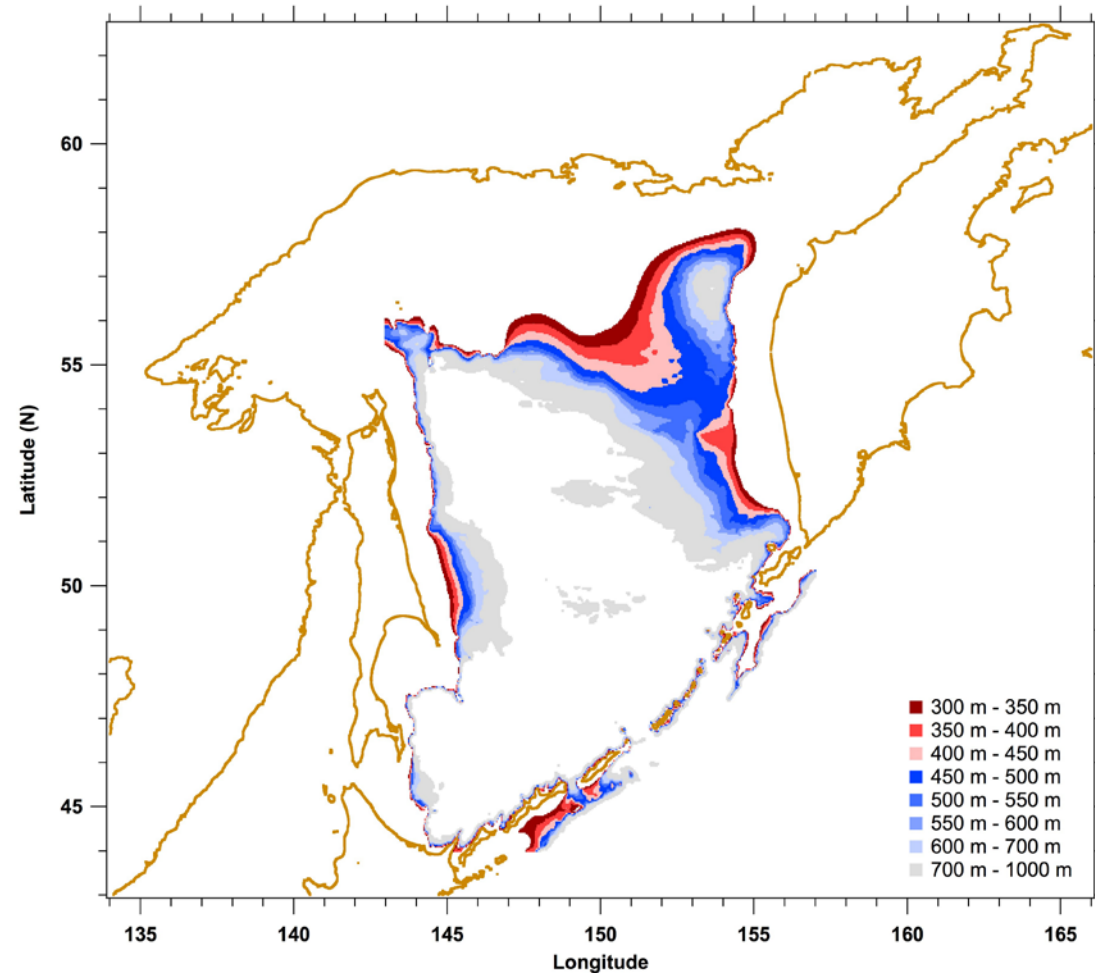
Integrated cumulative release (V) and net flux (Q)



- Adjust for 80% BSR occurrence (Ludmann & Wong, 2003)
- Instability confined to a narrow band near the top of the GHSZ
- 240 Tg released in the first century, fluxes < 5 Tg/yr

Integration: Sea of Okhotsk

Depth contours, coloring suggesting areas of potential destabilization (reds) vs. low-flux areas (blues)

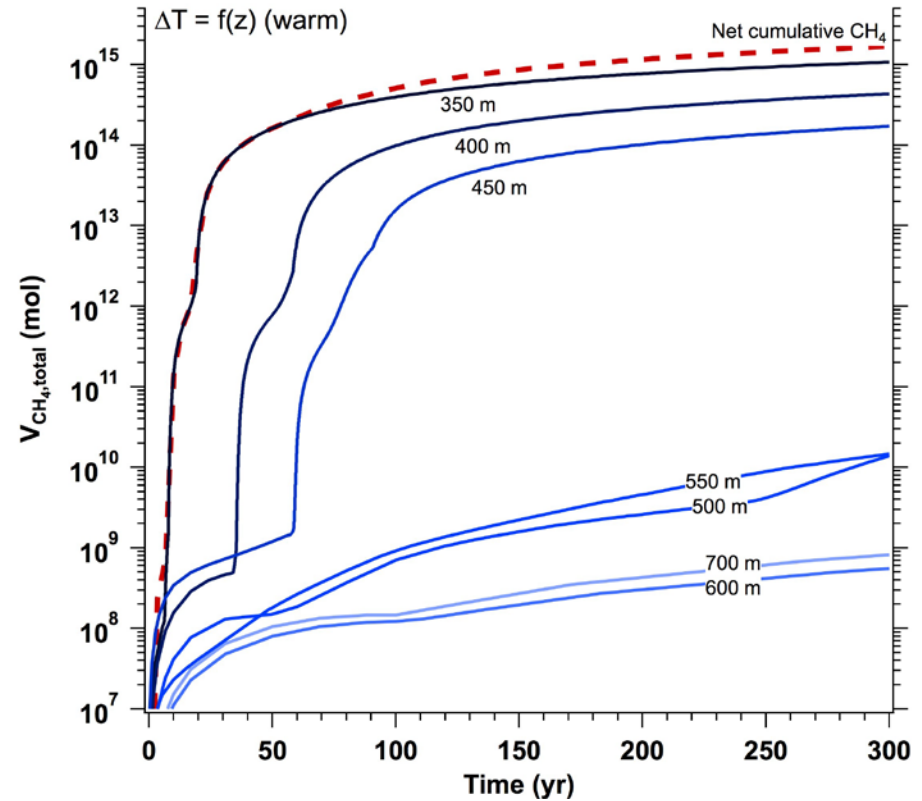
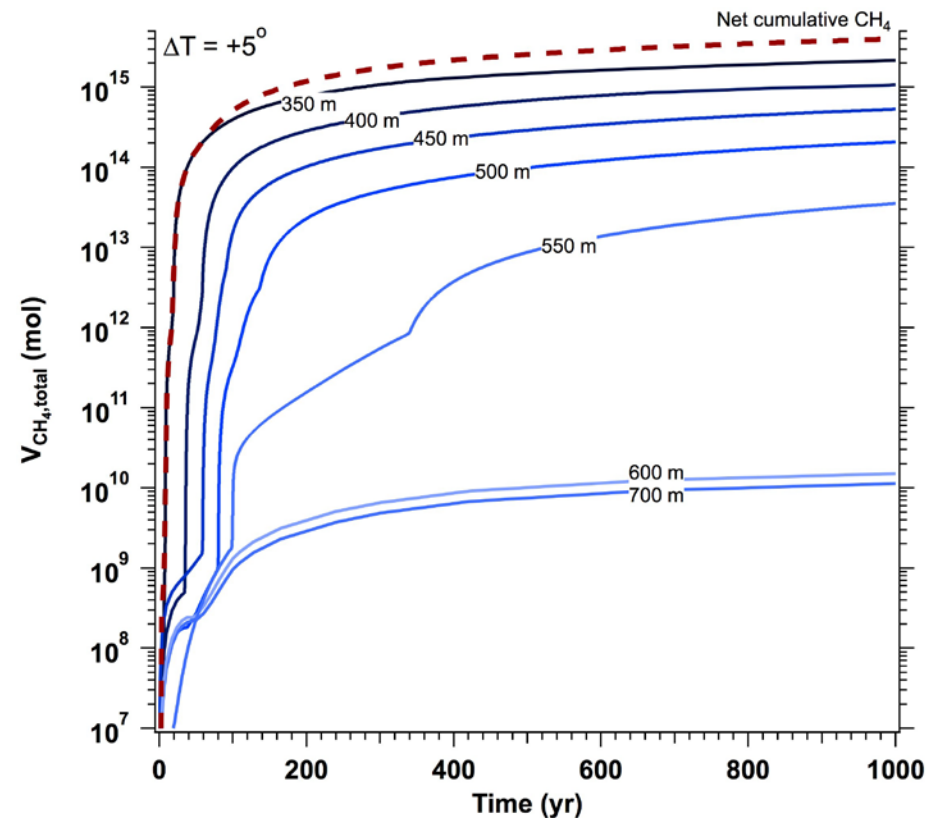


- The Okhotsk basin contains extensive hydrate deposits and high methane concentrations (Obzhirov; KOMEX)
- Hydrates and free gas imaged (Ludmann & Wong, 2003; Wallmann et al., 2006)
- **T+H: ~ 96 Tg** (± 80) after 30 yr release ($t = 83$ yr)
- **Previous estimate***: 94.4 Tg (by $t = 30$ yr)
- Only 0.1%-1% of the estimated Okhotsk methane reservoir
- Continued release may increase V_{CH_4} $5\times$

* Elliott et al., 2010

Integration: Arctic Ocean

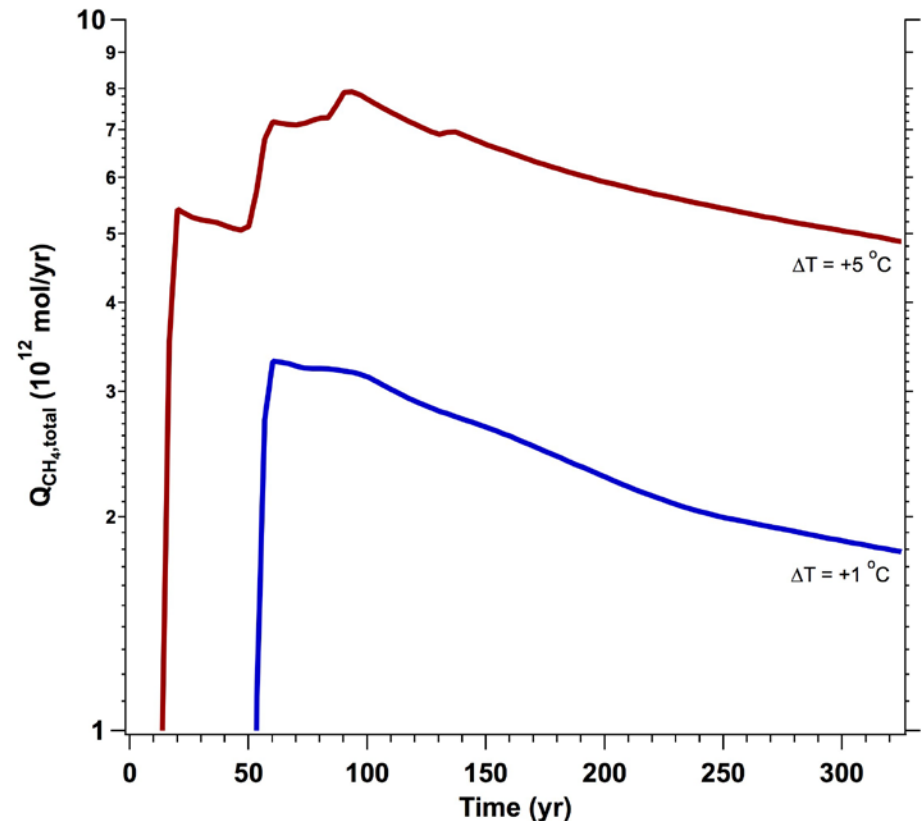
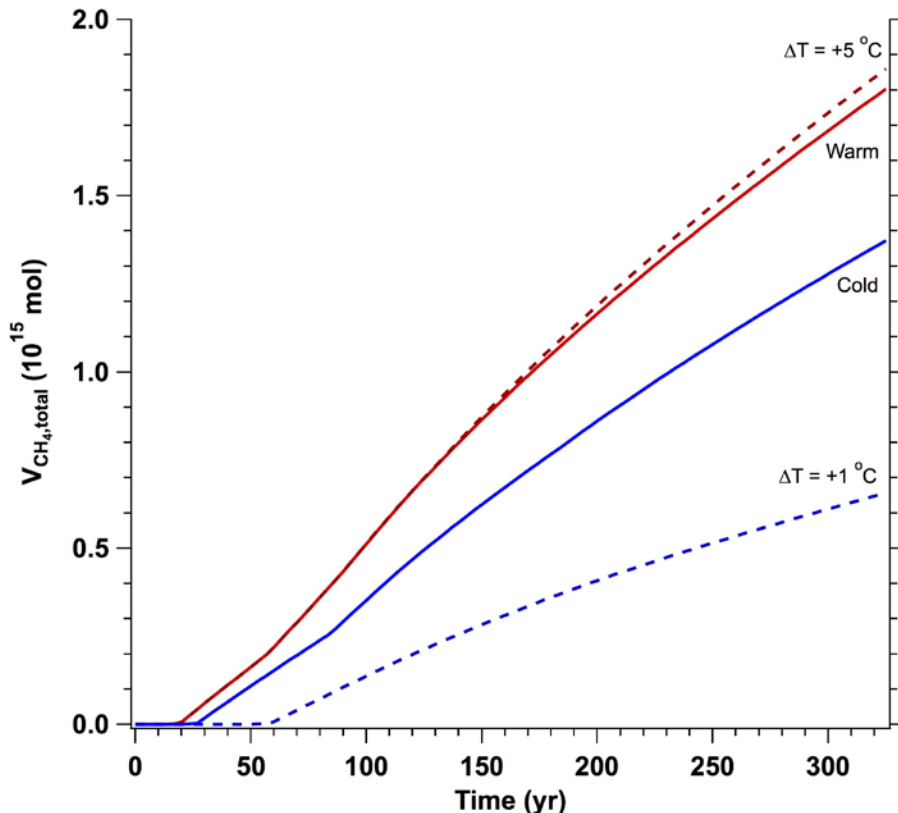
Cumulative release (V_{CH_4}) vs. depth



- Instability confined to a narrow band near the top of the GHSZ
- Up to 6,400 Tg CH_4 at $t = 100$ yr

Integration: Arctic Ocean

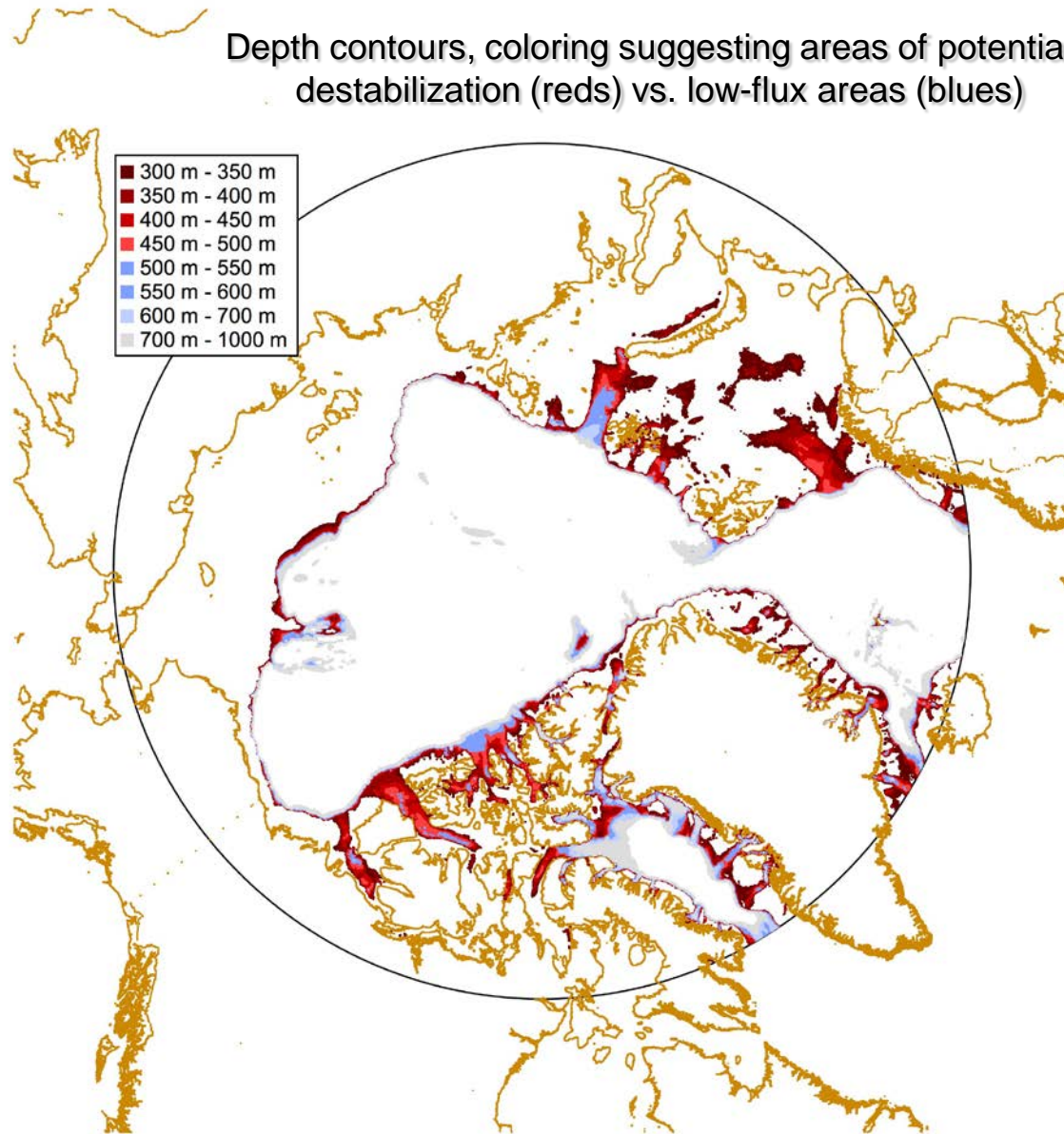
Cumulative release (V) and net flux (Q)



- Instability confined to a narrow band near the top of the GHSZ
- Scenarios: +5°C, +3°C, warm (1°C - 5°C as $f(z)$), cold (1°C - 3°C as $f(z)$)
- Short-term: 1600 - 3200 Tg CH₄ @ 30 yr (previous assumption: 240 Tmol)
- 60 – 120 Tg/yr peak fluxes

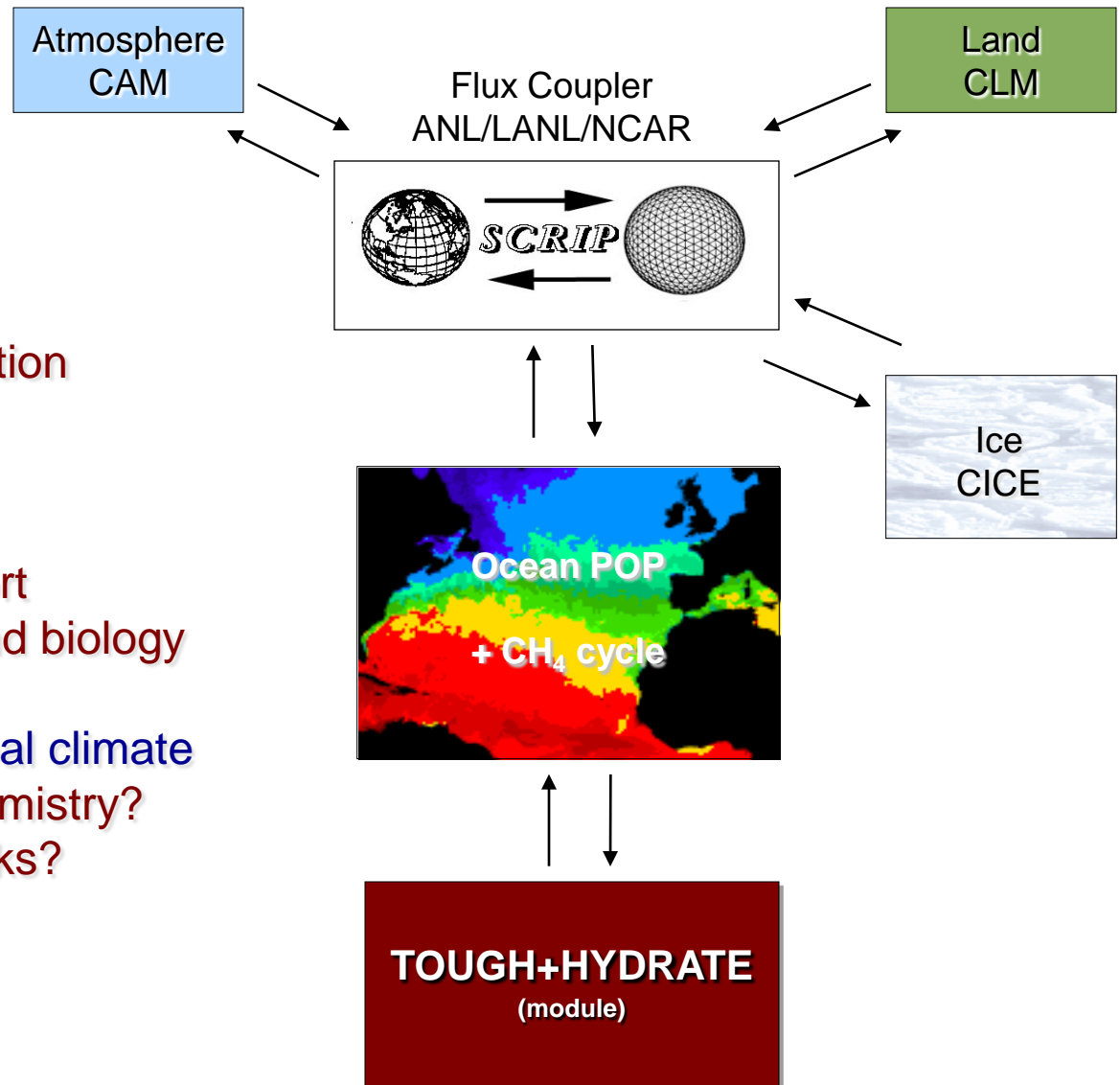
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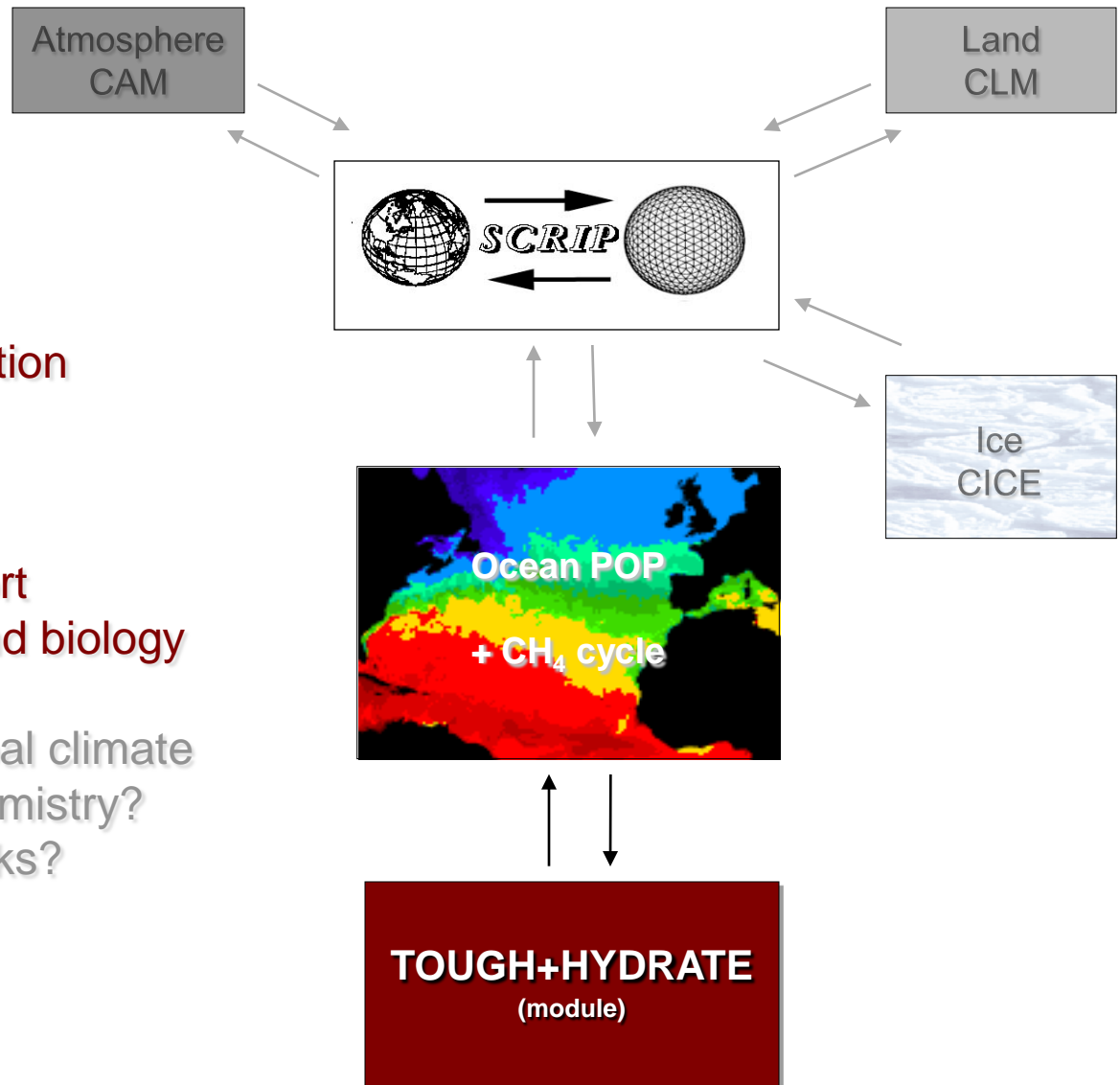
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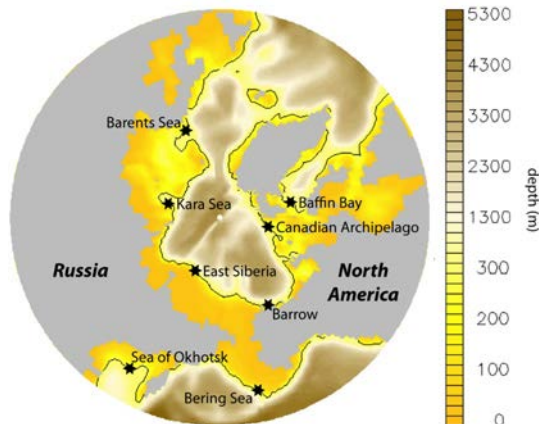
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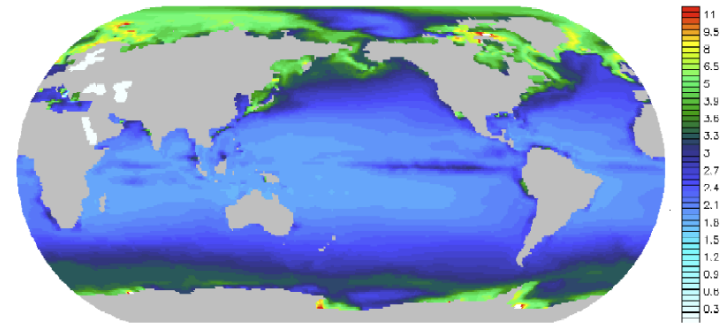
Coupled Ocean Modeling & Biochemistry

- POP extended to consider methane biogeochemistry in the water column^{5,6}
- A newly generated background methane cycle creates a baseline for methane release calculations⁶
- Localized methane releases from hydrates inserted into POP⁶ (fully dissolved*)

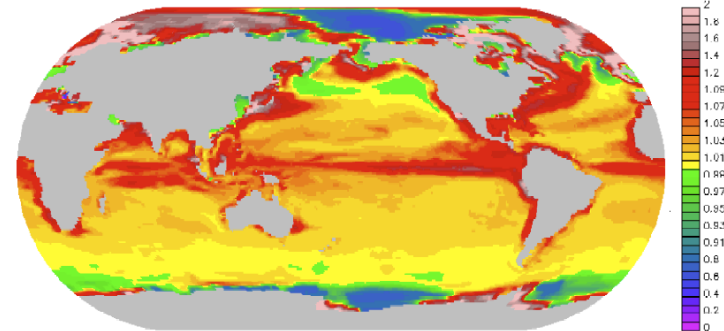
*Plume physics forthcoming



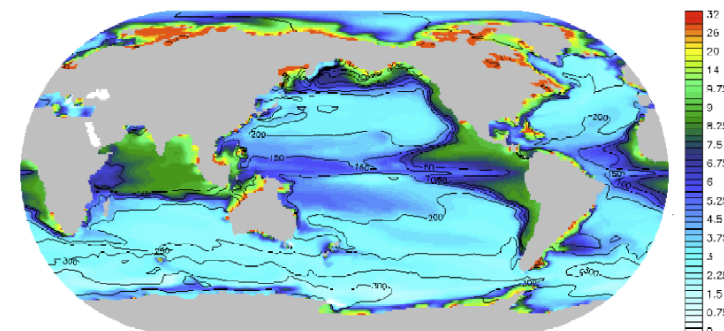
CH₄ Surface Concentration (nanomolar)



CH₄ Surface Saturation Ratio



CH₄ 150m Concentration (nanomolar) with O₂ contours

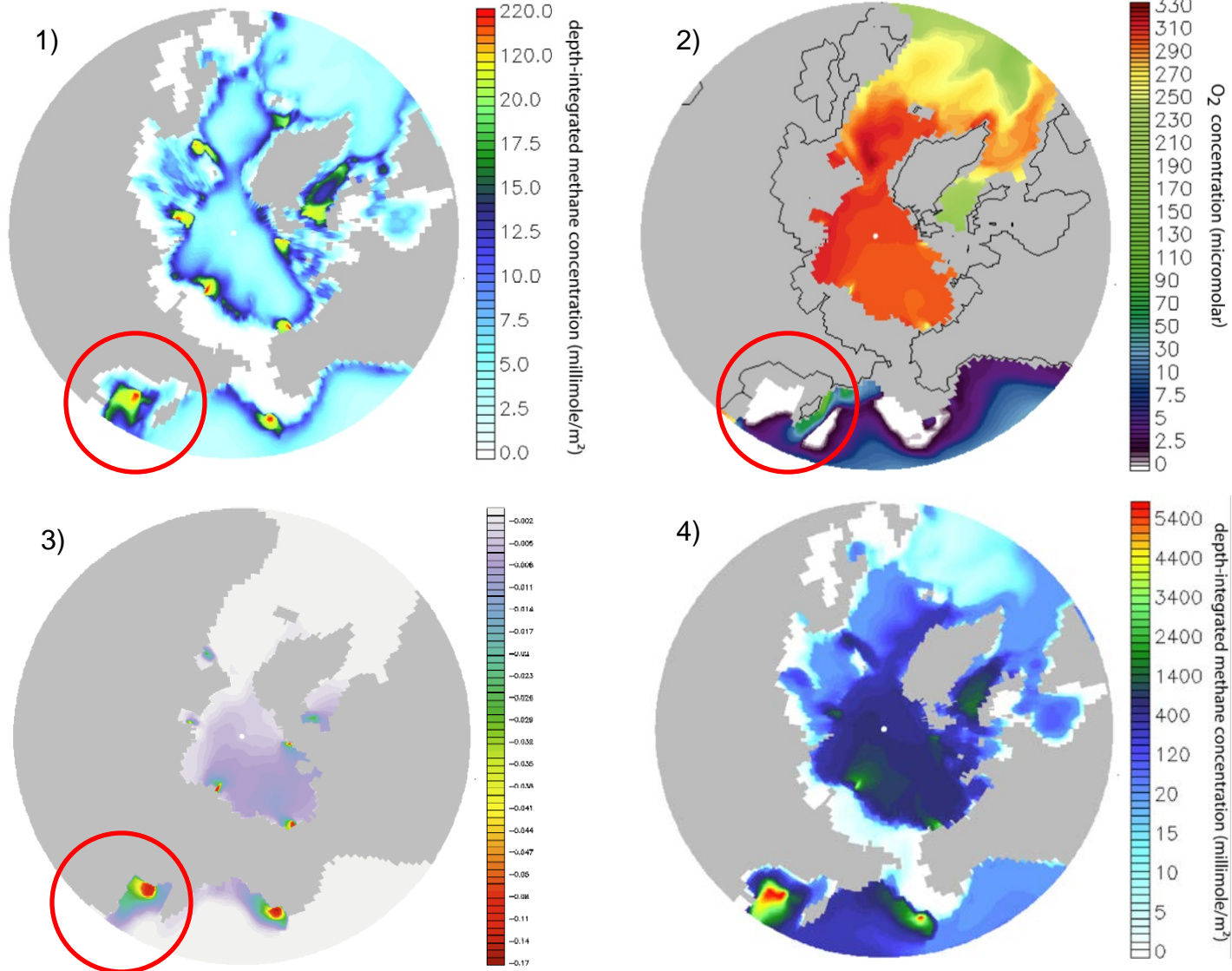


⁵Elliott, S.M., Maltrud, M., Reagan, M.T., Moridis, G.J., Cameron-Smith, P.J., "Geochemistry of Clathrate-Derived Methane in Arctic Ocean Waters," *Geophys. Res. Lett.*, **37**, L12607, 2010.

⁶Elliott, S.M., Maltrud, M., Reagan, M.T., Moridis, G.J., Cameron-Smith, P.J., "Marine Methane Cycle Simulations for the Period of Early Global Warming," *J. Geophys. Res. Biog.*, **116**, G01010, 2010.

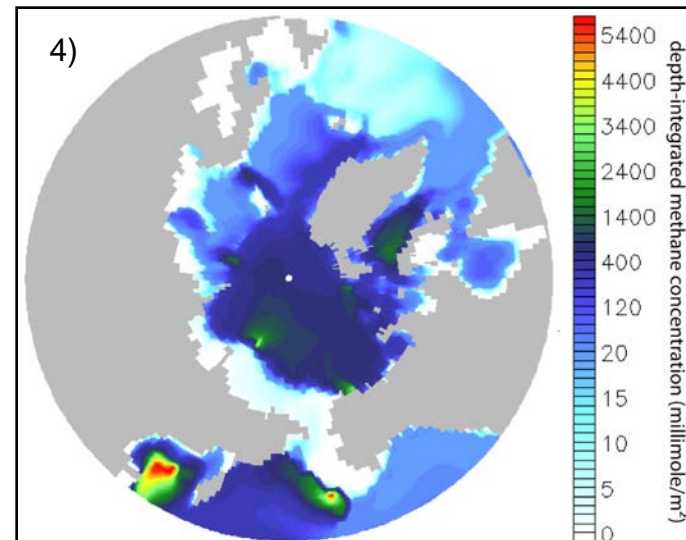
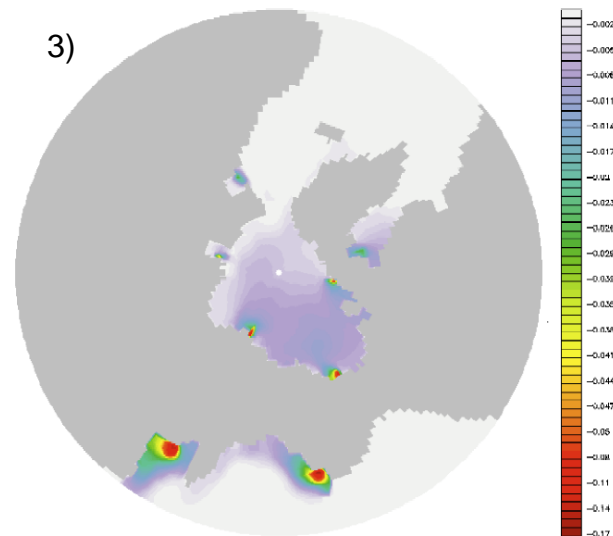
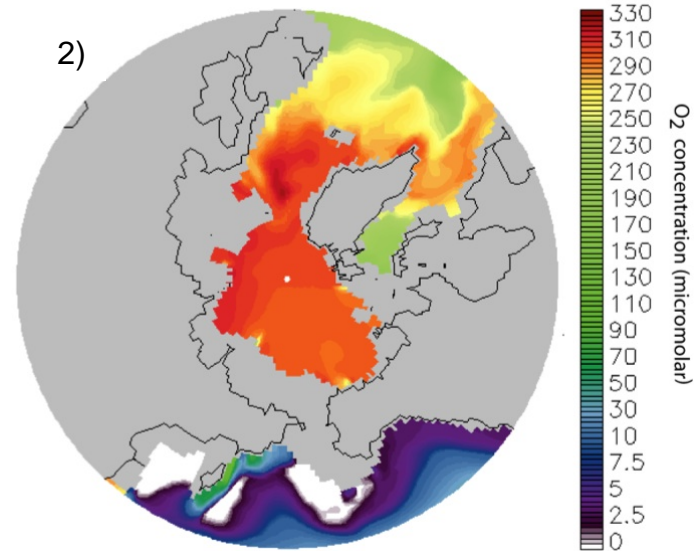
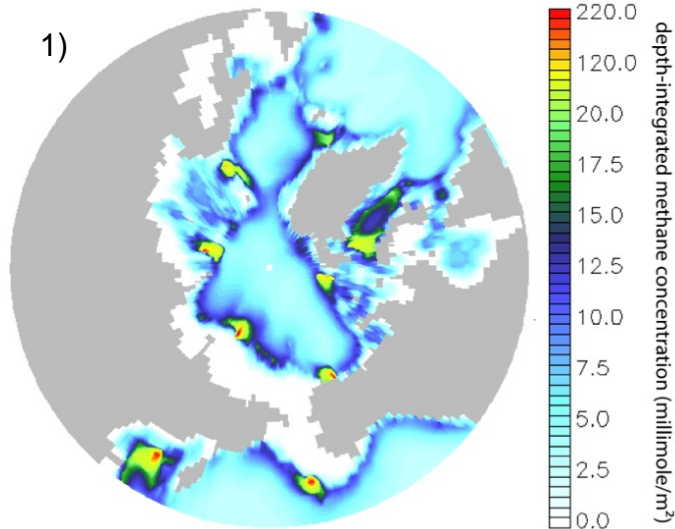
Coupled Ocean Modeling & Biochemistry

$t = 30$ yr:



Coupled Ocean Modeling & Biochemistry

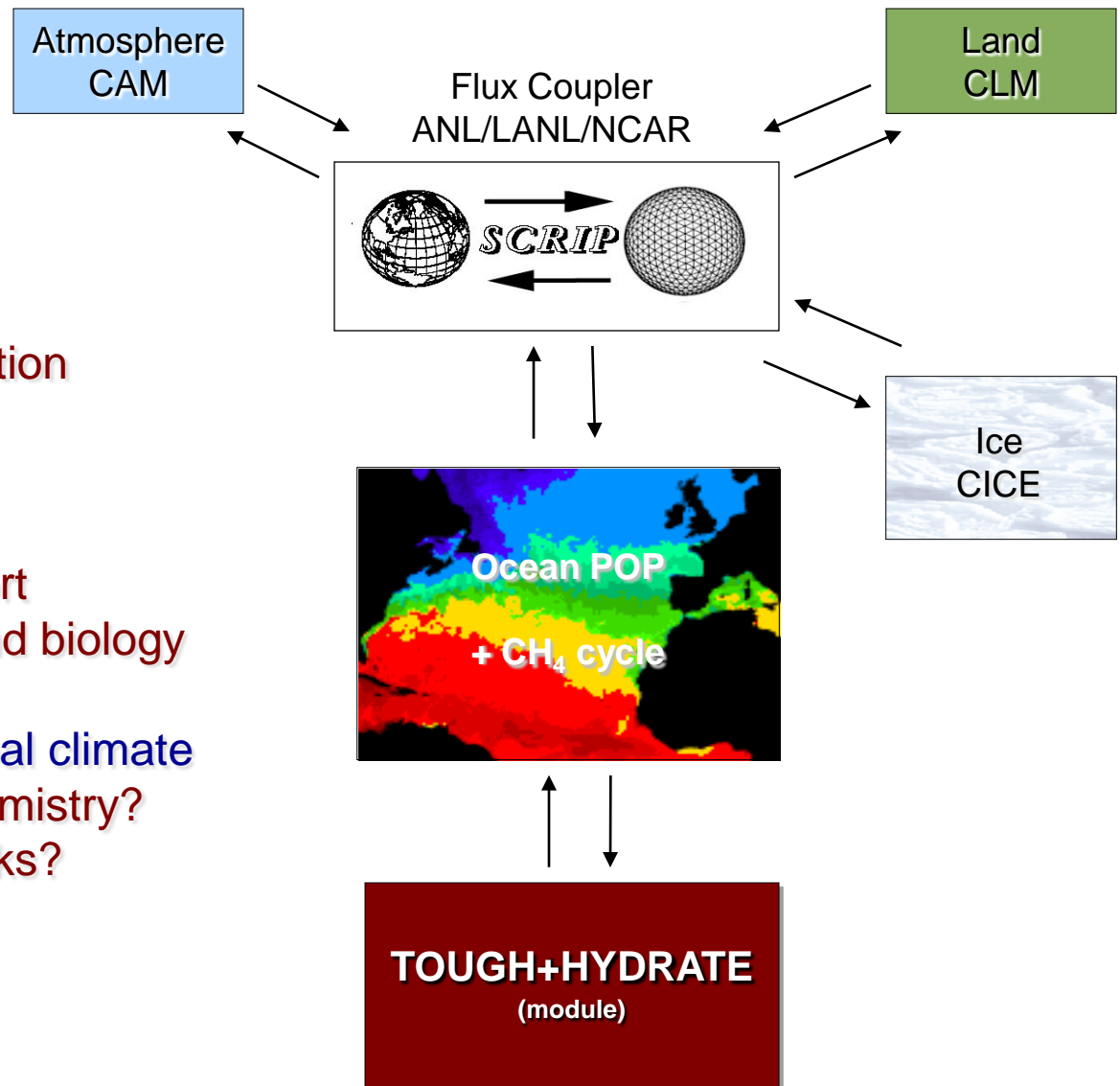
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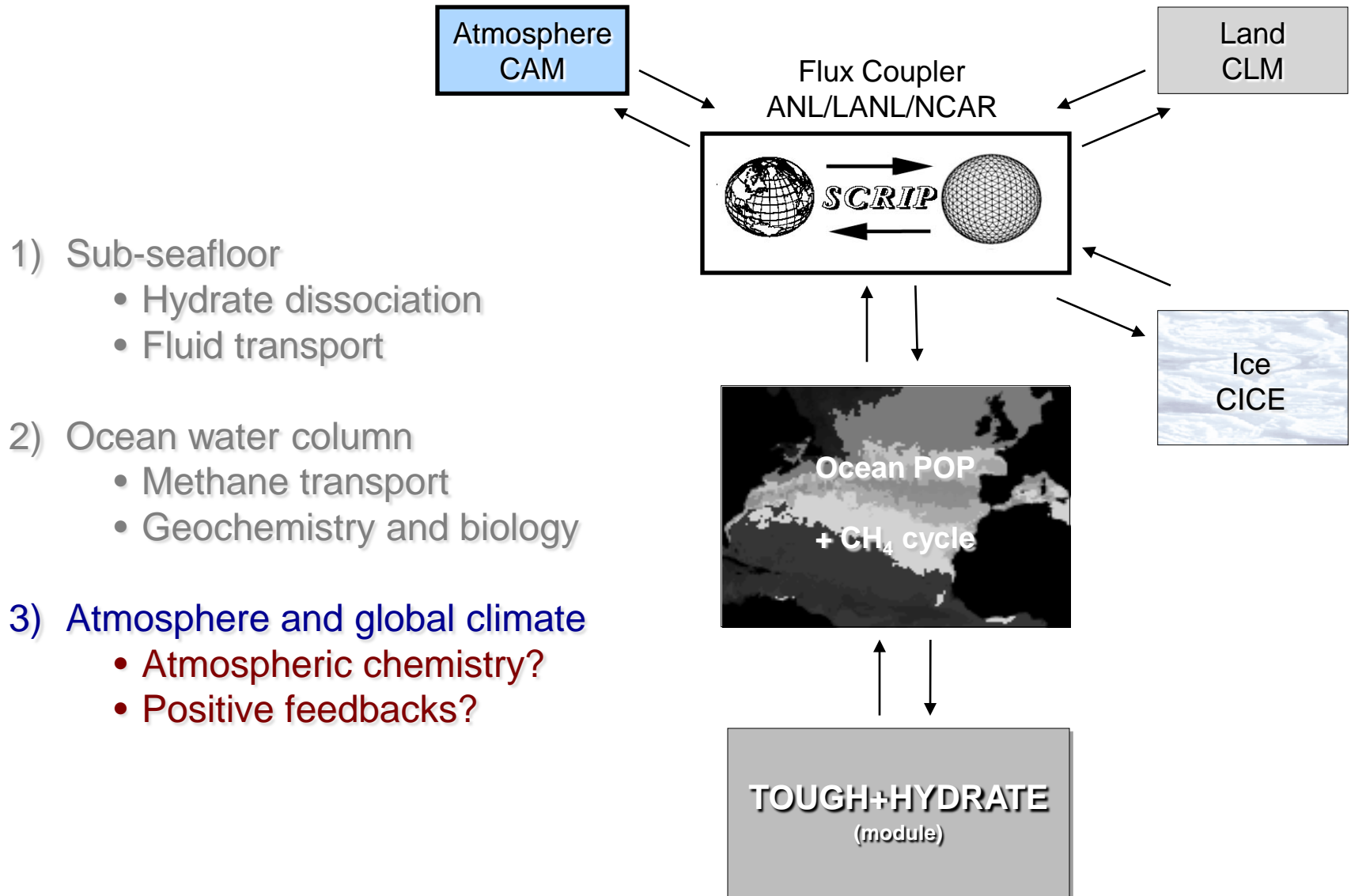
**~60%
Transfer to
atmosphere!**

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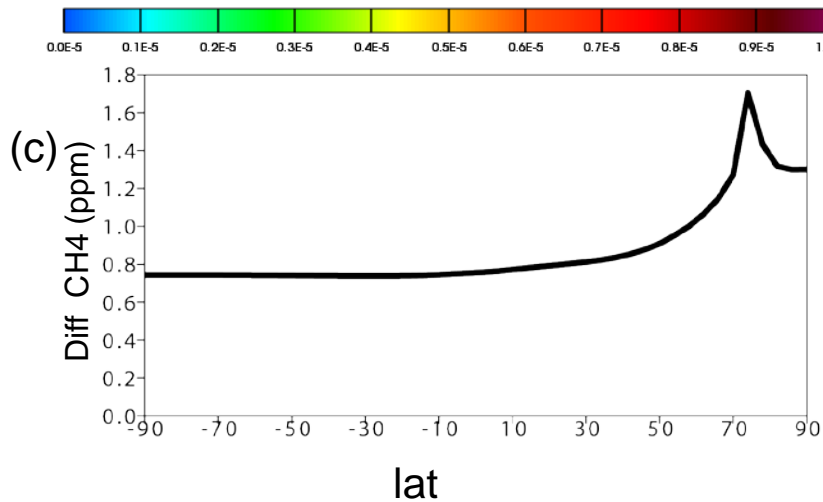
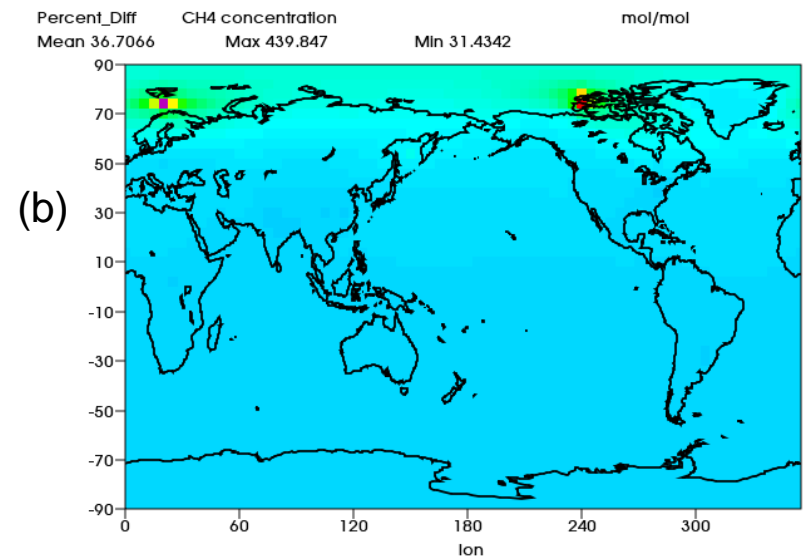
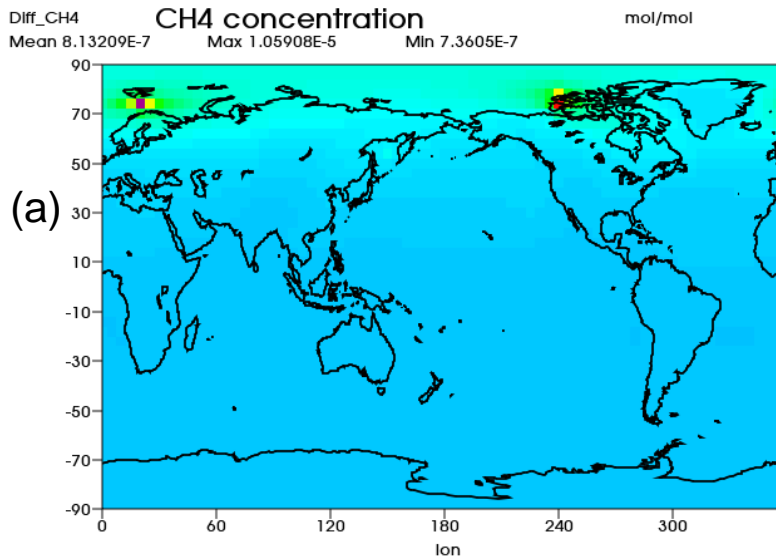
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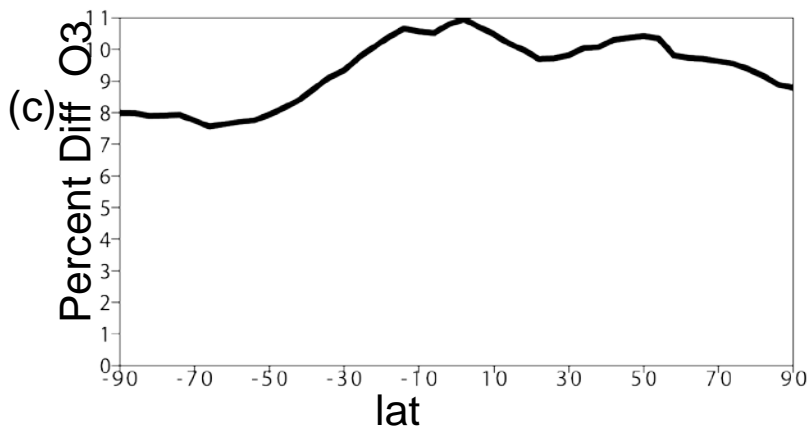
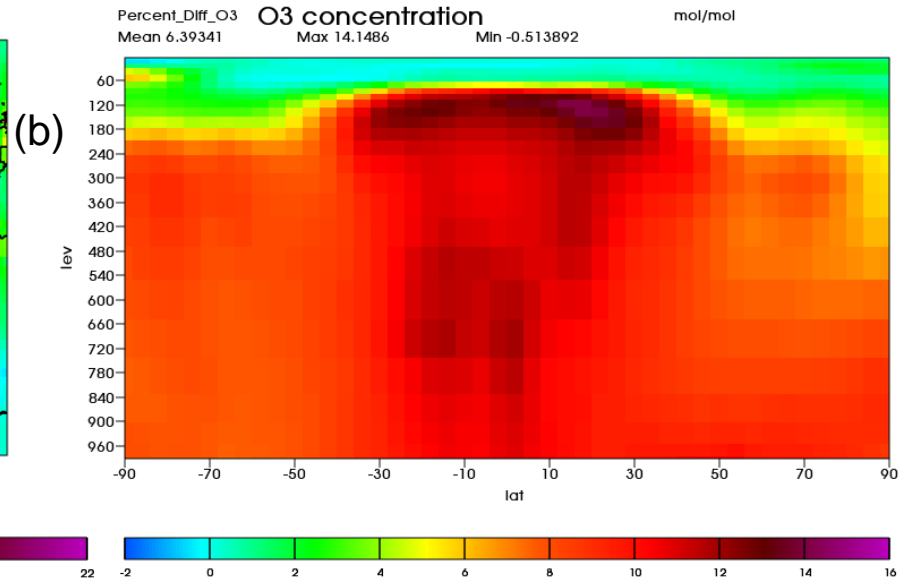
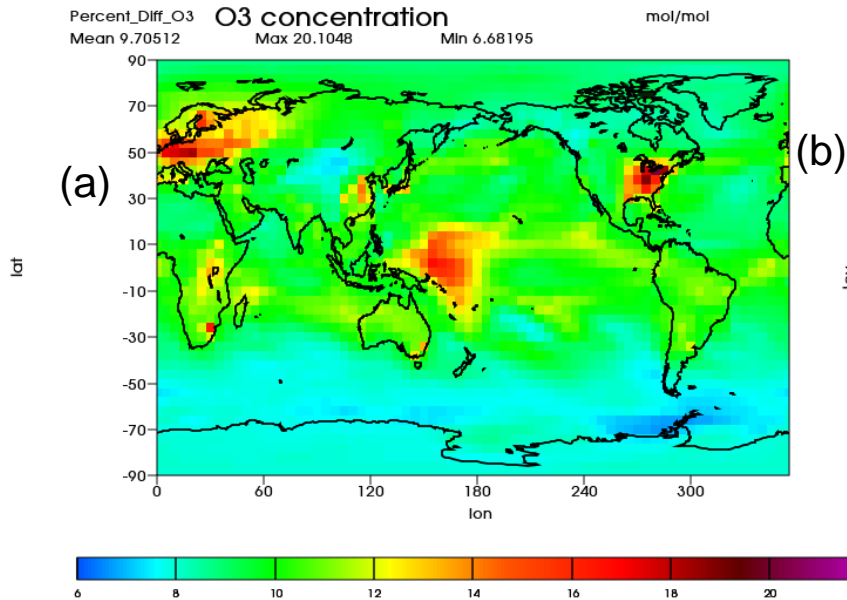


Increase in global CH₄ concentration



- (a) Difference in surface CH₄ concentration relative to control
- (b) Percent Difference in CH₄ concentration relative to Control
- (c) Difference in surface CH₄, zonal mean

Significant Increases in Ozone



- (a) Percent difference in surface ozone relative to control
- (b) Percent difference in atmospheric ozone (zonal mean)
- (c) Percent difference in surface ozone (zonal mean)

Conclusions



- Shallow hydrates can release significant methane rapidly, with significant methane fluxes regulated by coupled thermo-hydrological processes
- Methane is relevant to ocean (and atmospheric!) chemistry, not just as a contributor to total atmospheric CO₂
- 1-D models averaged over depth/temperature/area can estimate basin-scale release potential
- The vast majority of deep hydrates are stable, in the short term, but the methane release potential is still large
- Limited instability/release can feed biochemical/chemical changes in the ocean and atmosphere, before climate effects are considered
- Resource limitations overturn assumptions about methane oxidation
- New coupled seafloor-ocean-atmosphere calculations under way (with plume physics, extended biochemistry, higher resolution) leading to a coupled global model... and better estimates

<http://esdtools.lbl.gov/info/hydrate-publications/climate/>