

JCSDA Seminar
February 24, 2010

**Assimilation of Space-Borne GPS Radio
Occultation Data in NWP**

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Outline

Part I: Data

- (1) Introduction to GPS RO Techniques
- (2) Data Processing Chain & Error Sources

Part II: Assimilation

- (3) Choices for GPS RO Data Assimilation
- (4) Highlights of Assimilation Results

Part III: Data Application

- (5) Comparison with Large-Scale Analyses in
Cloudy and Clear-Sky Conditions
- (6) Cloudy Retrieval of GPS RO Data

Future Plan

Collaborators

- **My Ph. D. students & Postdoctoral fellow:**

H. Shao, L. Lin

F. Vandenberghe, B. Wang, H. Liu, Z. Zeng

- **Ray-tracing model:**

S. Sokolovskii, M. Gorbunov

- **Non-local refractivity observation operator:**

G. Hajj

- **GPS RO data:**

R. Anthes, Y.-H. Kuo, M. Exner, D. Hunt, R. Ware,
C. Rocken, W. Schreiner, D. Feng, B. Herman

(1) Introduction to GPS RO Techniques

What Is Measured by GPS RO Technique?

- A GPS receiver onboard an LEO satellite measures the **propagation delay** of the radio signals at two L-band frequencies ($f_1=1.57542$ GHz and $f_2=1.22760$ GHz) that are transmitted from a GPS satellite
- The **propagation delay** depends on the distribution of **atmospheric refractivity** as well as on the **ionosphere electron density**
- The **ionosphere effect** can be removed by combining the two propagation delays
- The wavelengths of the GPS signals are about 20 cm, at which there is **very little effect of aerosols, clouds, and rain**

Ray Paths

In a **vacuum**, radio signals travel along **straight lines** connecting an occulting GPS satellite and LEO satellite.

In **real atmosphere**, any ray path from GPS to LEO satellite is **bended** in the ionosphere and the atmosphere. The total bending (e.g., **bending angle**) can be derived given the precise positions and velocities of both satellites.

Vertical Profiles

Due to the satellite motions, the whole atmosphere from top to surface has rays passing through it, obtaining a vertical profile of bending angles from every occurrence of radio occultation.

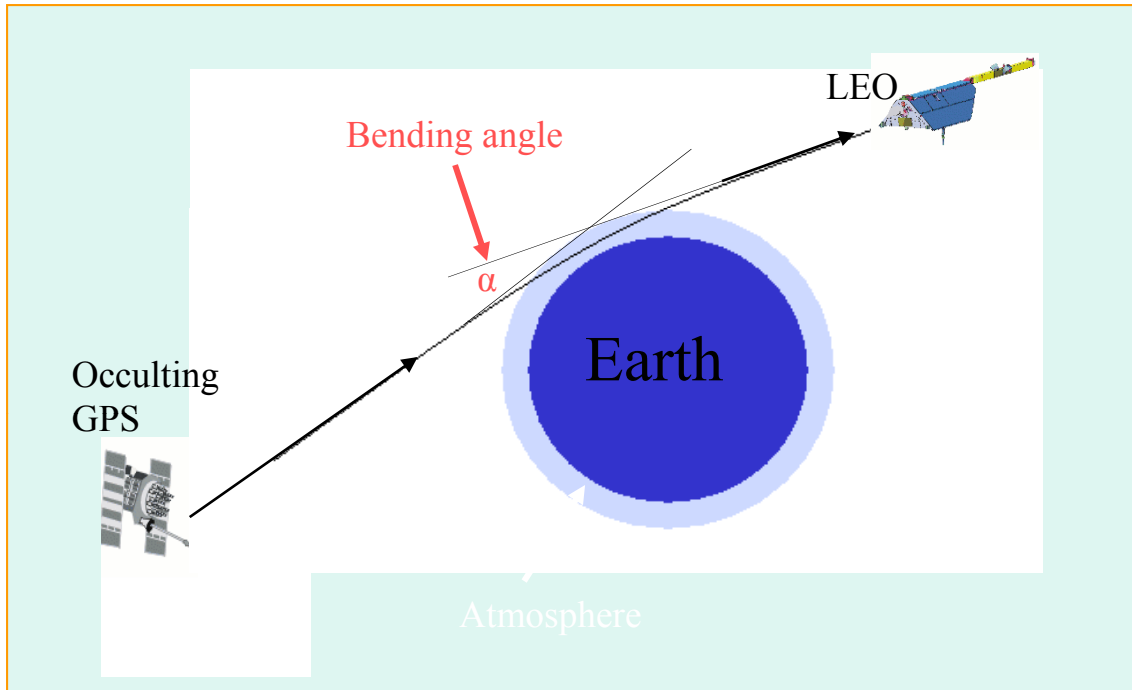
In the lowest ~100 km of the atmosphere, the LEO satellite collects data with a **50 Hz sampling rate**, resulting in **~3000 measurements** in a vertical profile and a **vertical resolution** of about **1.5 km** in the stratosphere and higher resolution of less than **0.5 km** in the lower troposphere. Such an occultation takes about **1 minute**.

Global Coverage □

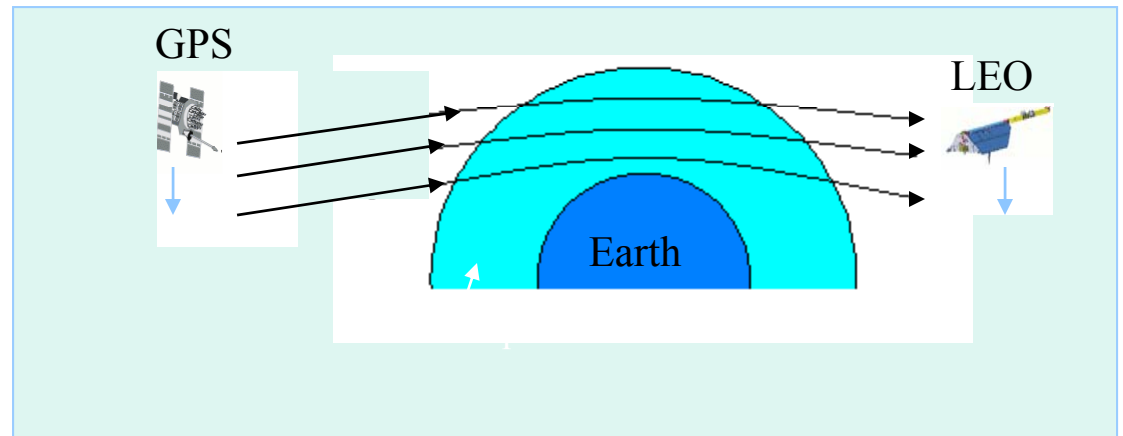
There are **24 GPS satellites** ((about 20,200 km altitude)in six orbital planes which continuously transmit electromagnetic waves at two L-band frequencies ($f_1=1.57542$ GHz and $f_2=1.22760$ GHz).

With 24 GPS satellites, a single GPS receiver in a near-polar orbit at 800 km will be observed over **500 ROs per day**, which are distributed fairly uniformly about the globe.

Bending Angle of a Ray Path



Vertical Profiling



A History of GPS RO Missions

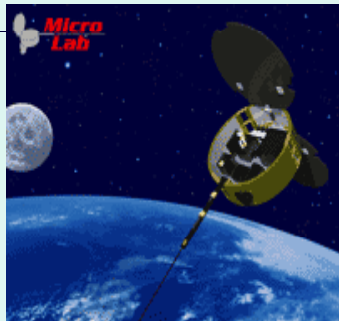
1 LEO

1 LEO

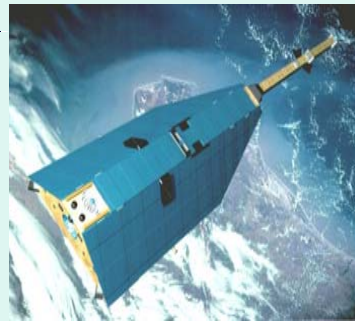
1 LEO

6 LEOs

6~12 LEOs



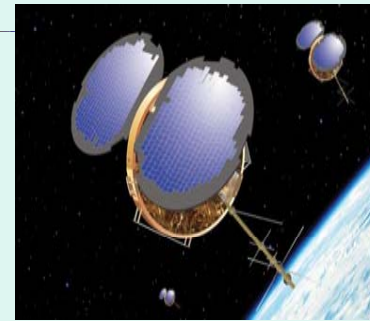
GPS/MET



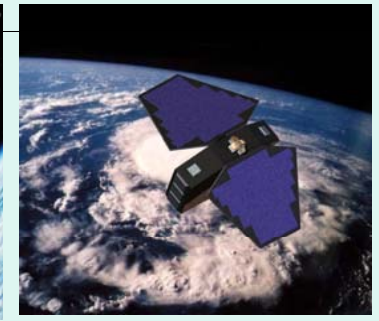
CHAMP



SAC-C



COSMIC



COSMIC II

April 1995 → Feb. 1997

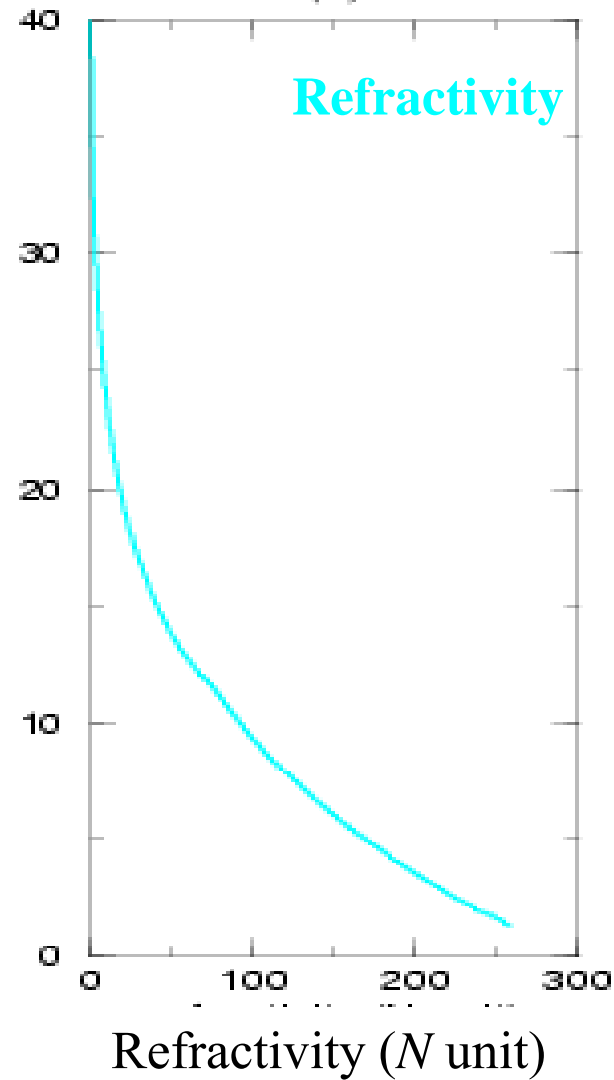
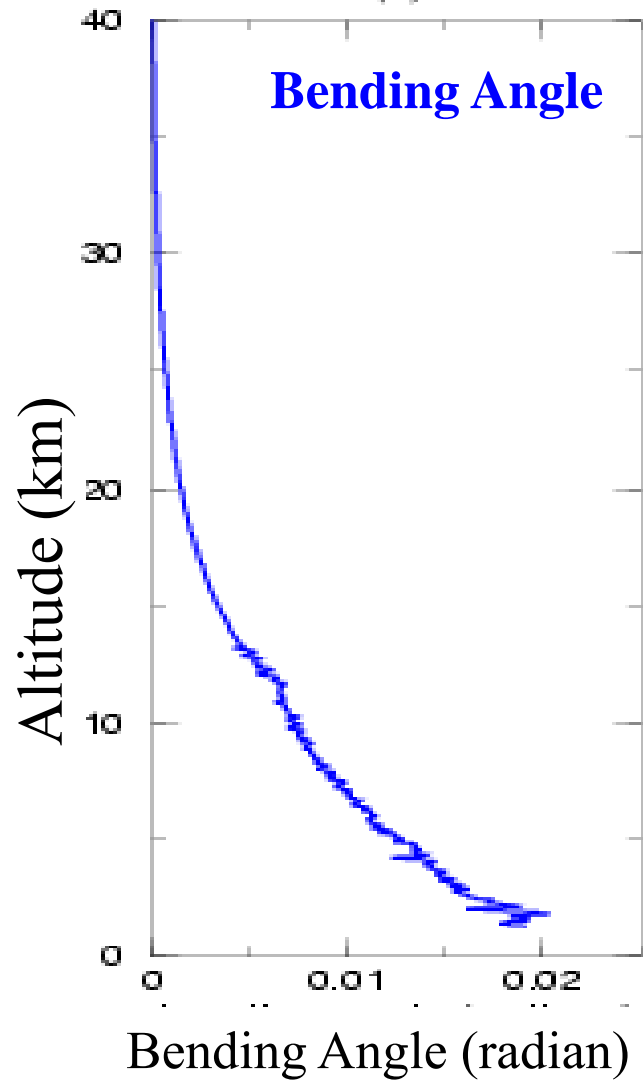
April 2000 → present

Nov. 2000 → present

April 2006 → present

2011 →

A Vertical RO Profile from CHAMP

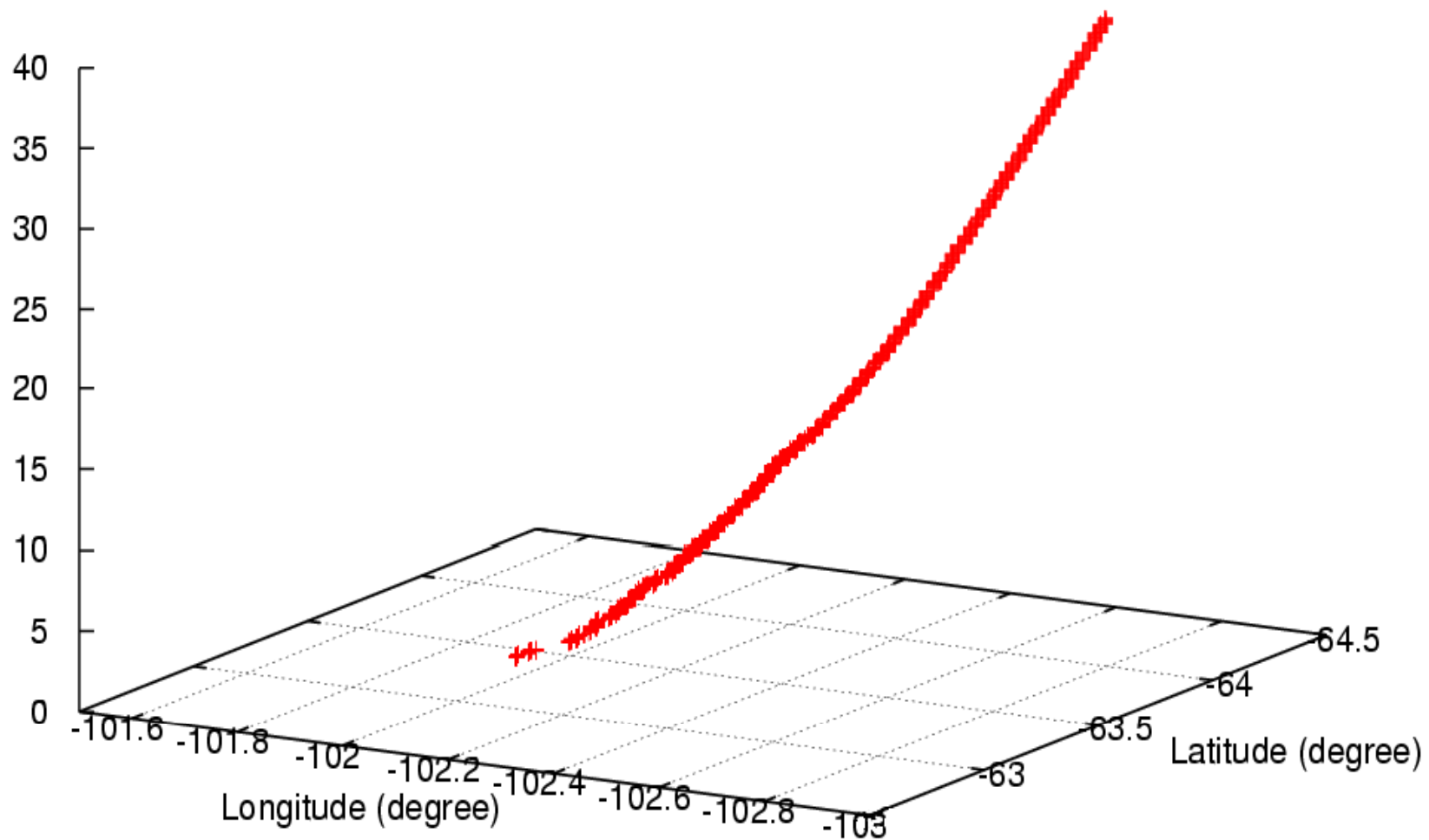


Measurement Time:
0836UTC 21 May 2002
Averaged Location:
(102.04°W, 63.40°S)

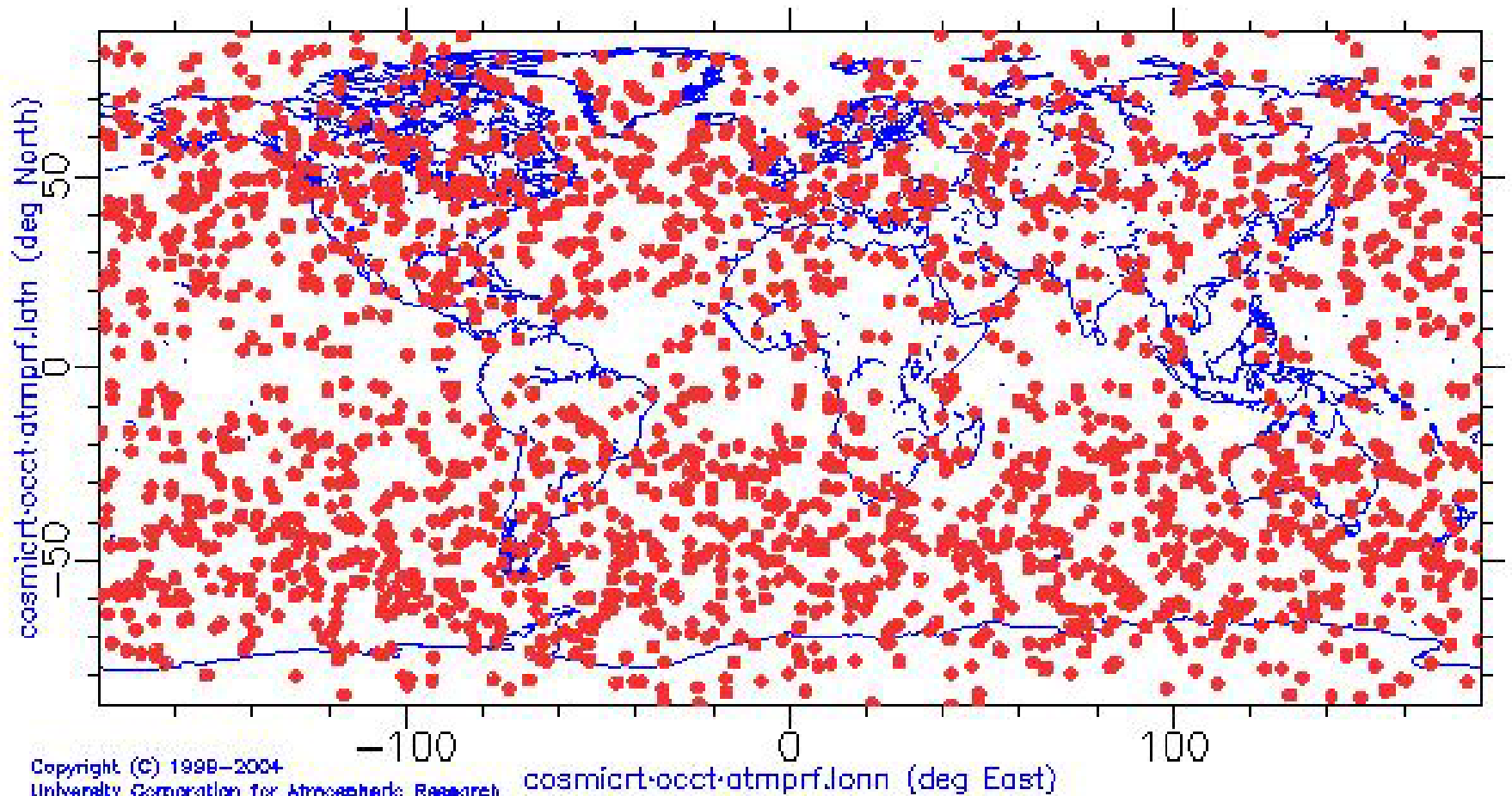
A Vertical RO Profile from CHAMP (cont.)

Locations of the tangent points

Altitude (km)



COSMIC Global Coverage on March 25, 2009

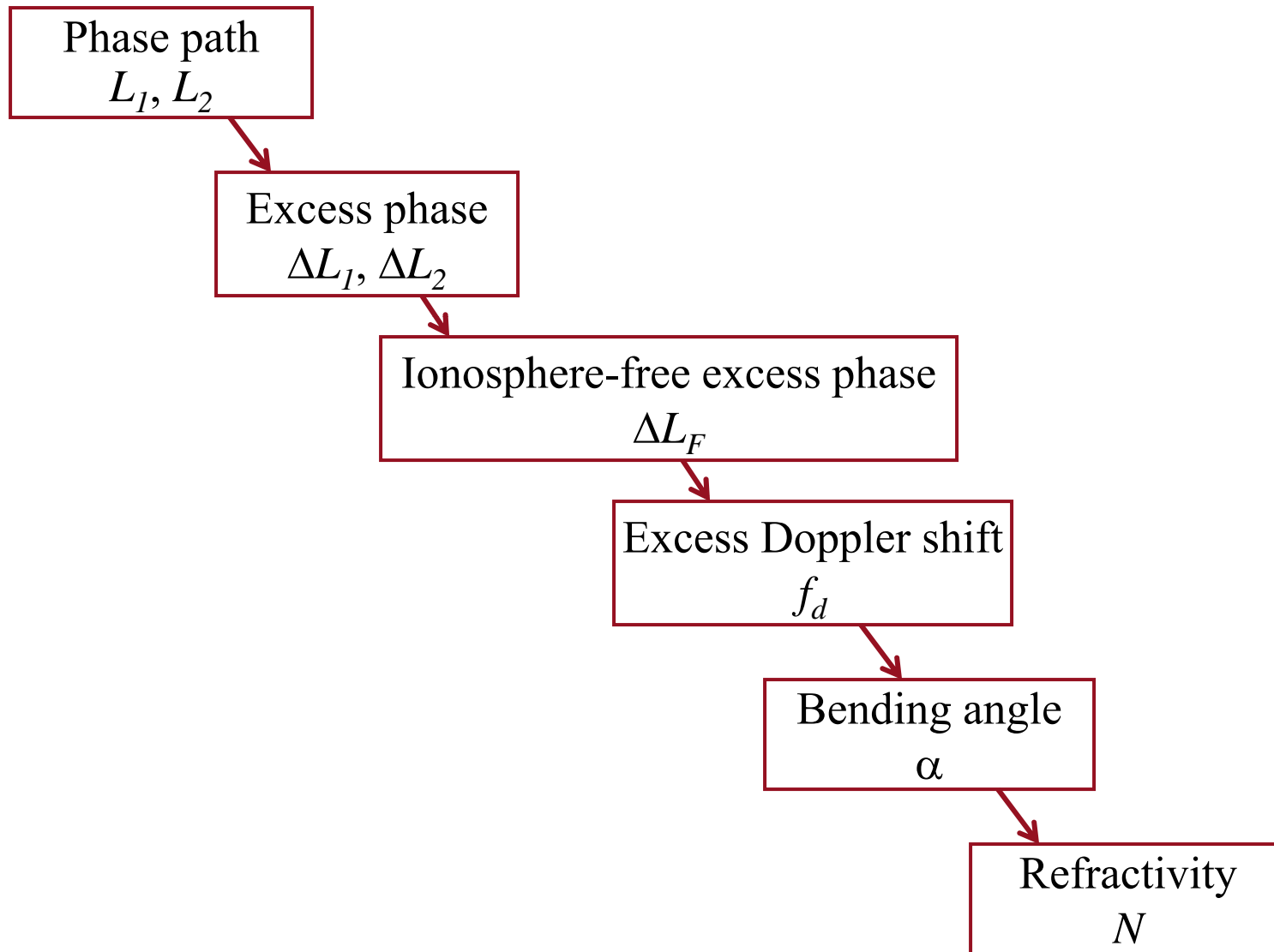


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cosmicrt-occt-atmprf.lonn (deg East)

(2) Data Processing Chain & Error Sources

Data Processing Chain



(1) Phase Path

The GPS RO observable is the **phase path**:

$$L_1 = \int_{GPS}^{LEO} n(f_1) ds, \quad L_2 = \int_{GPS}^{LEO} n(f_2) ds$$

where n is the **refractive index**:

$$n(f) = 10^{-6} \left(k_1 \frac{P_d}{T} + k_2 \frac{P_w}{T} + k_3 \frac{P_w}{T^2} - C \frac{N_e}{f^2} \right) + 1$$

refractivity \nearrow N , contribution from the neutral atmosphere \uparrow ionosphere contribution

(2) Excess Phase

The **excess phase** is calculated as the difference between the measured phase path and the “vacuum” phase path (equal to the geometric distance between GPS and LEO):

$$\Delta L_1 = L_1 - R_{GL}, \quad \Delta L_2 = L_2 - R_{GL}$$

where R_{GL} is the geometric distance between GPS and LEO.

Signs of Refractive Index and Excess Phase

For microwave frequencies,

$n < 1$ in the **ionosphere** and is frequency dependent.

$n > 1$ in the **neutral atmosphere** and is independent of frequency.

For rays with tangent heights **above the tropopause**, the excess phase is **negative** (about 10-100 m) due to the penetration of ionospheric layers on both sides of the neutral atmosphere.

The excess phase becomes **positive below the tropopause** because the effect from the neutral part of the atmosphere exceeds that of the ionosphere. The excessive phase is on the order of 1 km close to the surface.

(3) Removing the Ionosphere Effect

Since the first-order contribution of the ionosphere refractivity is inversely proportional to the frequency squared, and can thus be eliminated by applying a linear combination of L1 and L2 at the same time samples:

$$L_F = \frac{f_1^2 L_1 - f_2^2 L_2}{f_1^2 - f_2^2}, \quad \Delta L_F = \frac{f_1^2 \Delta L_1 - f_2^2 \Delta L_2}{f_1^2 - f_2^2}$$

where L_F and ΔL_F is the ionosphere free phase path and excess phase, respectively.

Magnitude of Excess Phase

The ionosphere corrected excess phase, ΔL_F is what would be measured had there been no ionosphere at all. Therefore, the ionosphere free excess phase is always positive, and is very small at the high altitude end. In fact,

$$\Delta L_F \approx \begin{cases} 1 \text{ mm at } 80 \text{ km} \\ 1\text{-}9 \text{ cm at } 60 \text{ km} \\ 1 \text{ km at the surface} \end{cases}$$

The 1-mm excess phase is less than data noise level.

(4) Excess Doppler Shift

The **excess Doppler shift** is derived as the time derivative of the **excess phase**:

$$f_d = -fc^{-1} \frac{d\Delta L}{dt}$$

where c is the light velocity in a vacuum.

(5) Bending Angle

The excess Doppler shift is related to the satellites geometry:

$$f_d = f \left(\frac{c - n_{LEO} \left(v_{LEO}^r \cos \phi_{LEO} - v_{LEO}^t \sin \phi_{LEO} \right)}{c - n_{GPS} \left(v_{GPS}^r \cos \phi_{GPS} - v_{GPS}^t \sin \phi_{GPS} \right)} - 1 \right)$$

Assuming spherical symmetry:

$$\frac{r_{GPS} n_{GPS} \sin \phi_{GPS}}{a_{GPS}} = \frac{r_{LEO} n_{LEO} \sin \phi_{LEO}}{a_{LEO}}$$

ϕ_{LEO} (ϕ_{GPS}) is the angle between LEO (GPS) satellite radius and the tangent direction of the LEO (GPS) satellite velocity in the ray plane.

(5) Bending Angle (cont.)

Bending angle:

$$\alpha(a) = \phi_{GPS} + \phi_{LEO} + \arccos\left(\frac{\mathbf{r}_{GPS} \cdot \mathbf{r}_{LEO}}{r_{GPS}r_{LEO}}\right) - \pi$$

Impact parameter:

$$a = r_{GPS} n_{GPS} \sin \phi_{GPS}$$

(6) Refractive Index

The Abel inversion:

$$n(a) = \text{EXP} \left(\frac{1}{\pi} \int_a^{\infty} \frac{\alpha(x)}{\sqrt{x^2 - a^2}} dx \right),$$

(7) Refractivity

$$N = 10^6 (n - 1)$$

Error Sources

- **Measurement Errors:**

- (i) Random errors: Clock error and thermal noise
- (ii) Systematic errors: Signal scattering in the vicinity of the antenna, position errors, velocity errors, and retrieval errors

↓
Retrieval errors:

- (1) Ionosphere calibration errors
- (2) Upper altitude boundary errors
- (3) Errors introduced by the spherical symmetry assumption
- (4) Errors induced by atmospheric multi-path propagation

(3) Choices for GPS RO Data Assimilation

What to Assimilate?

1. **Excess phase:** caused by the bending of the radio signal at two frequencies: 1227.6 MHz , 1575.4 MHz .
2. **Excess Doppler frequency shift:** estimated by the time derivative of excess phase.
3. **Bending angle and impact parameter:** derived from Doppler frequency shift based on satellite geometry (impact parameter is assumed constant at GPS and LEO).
4. **Refractivity:** calculated from bending angle through the Abel inversion (the refractivity is assumed spherically symmetric).
5. **Temperature and pressure:** retrieved from refractivity using the hydrostatic equation and neglecting water vapor content.

Pros and Cons

Obs. Variable	Obs. Error Statistics	H Complexity	H Comput. Cost	Assumption(s)
ΔL^{obs}	simplest	most complex	most expensive	no SSA
α^{obs}	simple	complex	expensive	no SSA
N^{obs}	very complex	simple (N^{loc})	inexpensive	SSA
T^{obs}, q^{obs}	most complex	simplest	Least expensive	SSA, auxiliary info.
$\Delta L^{“obs”}$	simple	simple	inexpensive	No SSA
N^{obs}	simple	simple ($N^{non-loc}$)	inexpensive	No SSA

SSA --- Spheric Symmetry Assumption

(4) Highlights of Assimilation Results

(4a) Comparison between bending angle and refractivity assimilation

(4b) Comparison between local and non-local refractivity assimilation

Observation Operator for N

$$F_I H_N(T, p, q)$$

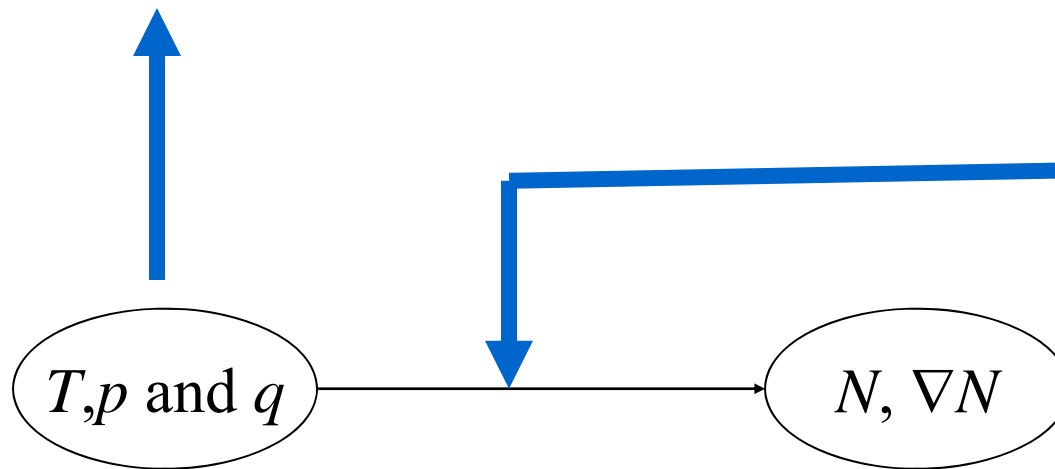
H_N

$$N = c_1 \frac{p}{T} + c_2 \frac{pq}{T^2 (0.622 + 0.378q)}$$

where $c_1=77.6$ and $c_2=3.73 \times 10^5$
Unit: T (K), q (kg/kg), p (hPa)

F_I

Interpolate the refractivity onto points along the ray-path



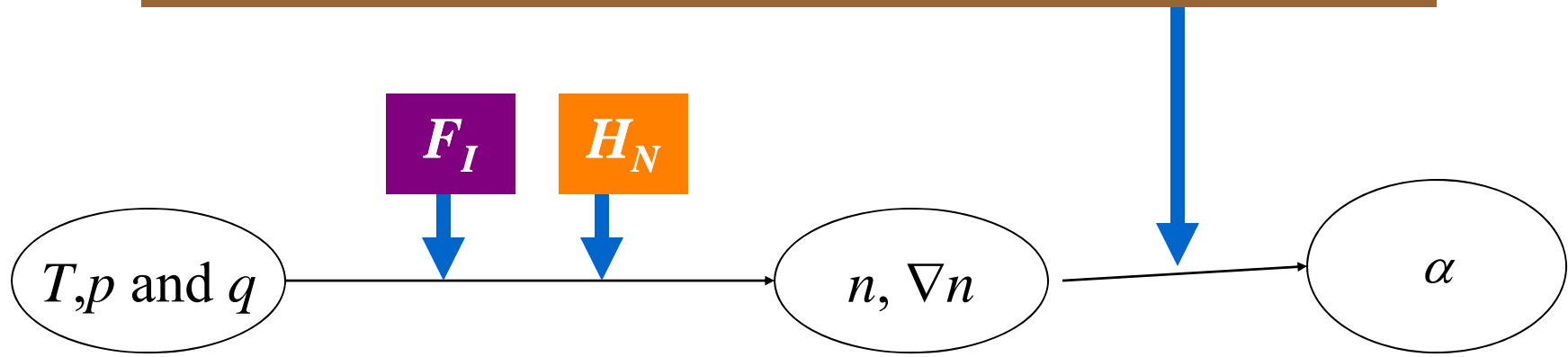
Observation Operator for α

$$\alpha = H_{\alpha} F_I H_N (T, p, q)$$

H_{α}

Ray-tracing model: $\frac{d^2 x}{d\tau^2} = n \nabla n$

where $\vec{x}(\tau)$ is the ray trajectory with $d\tau = ds/n$, s is the length of the ray, and n is the refractivity index.

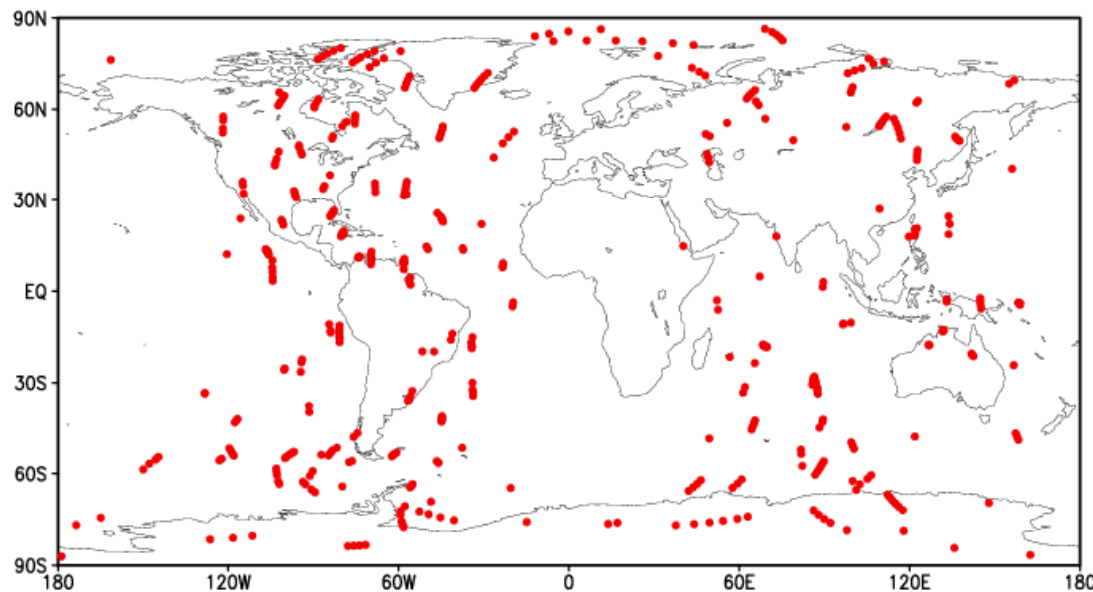


Experiment Design

NOGPS: Control run

BA: Bending angle added to data used in NOGPS

REF: Refractivity added to data used in NOGPS



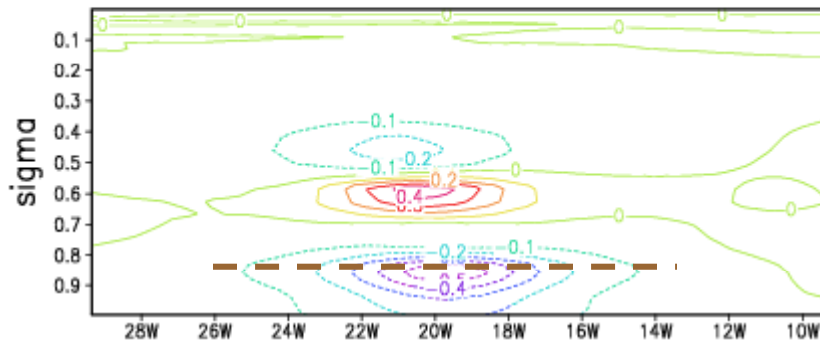
Total 434 ROs

System: NCEP/SSI T170L42

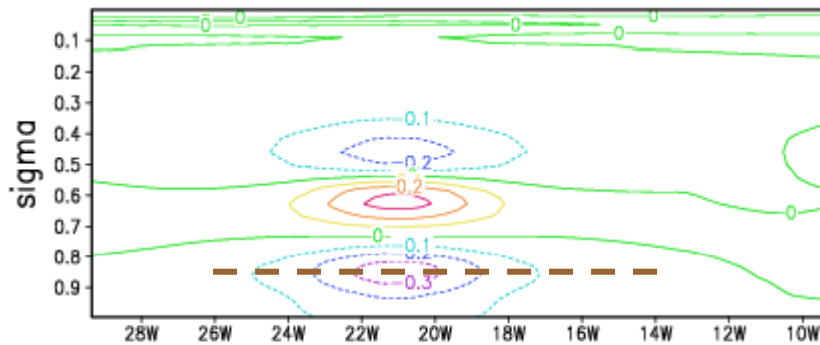
Data: 03-09 UTC from 21-31 May 2002

Differences in Moisture Analysis

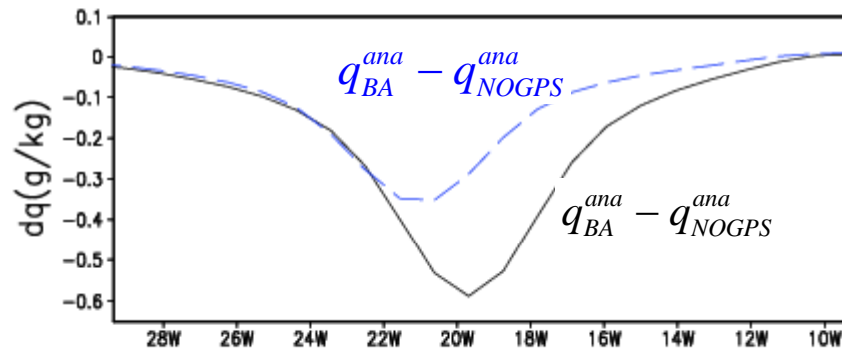
$$q_{BA}^{ana} - q_{NOGPS}^{ana}$$



$$q_{REF}^{ana} - q_{NOGPS}^{ana}$$



$$\sigma = 0.8585$$

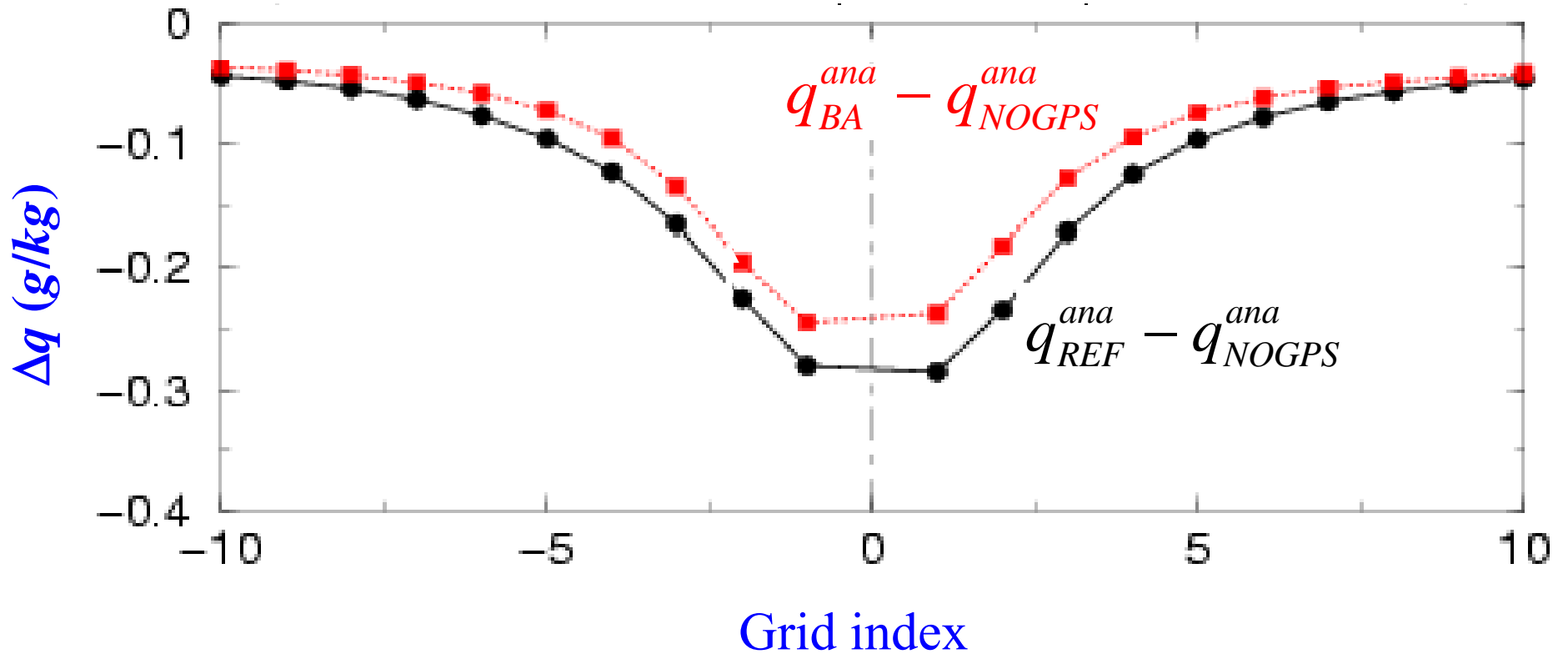


Cross-section passing through
a CHAMP RO observed at 06
UTC 21 May 2002

RO location:
(52.64°N 19.36°W)

Differences in Moisture Analysis (cont.)

Zonal variations averaged over all 434 ROs



Comparing Three Analyses with CHAMP Refractivity Observations

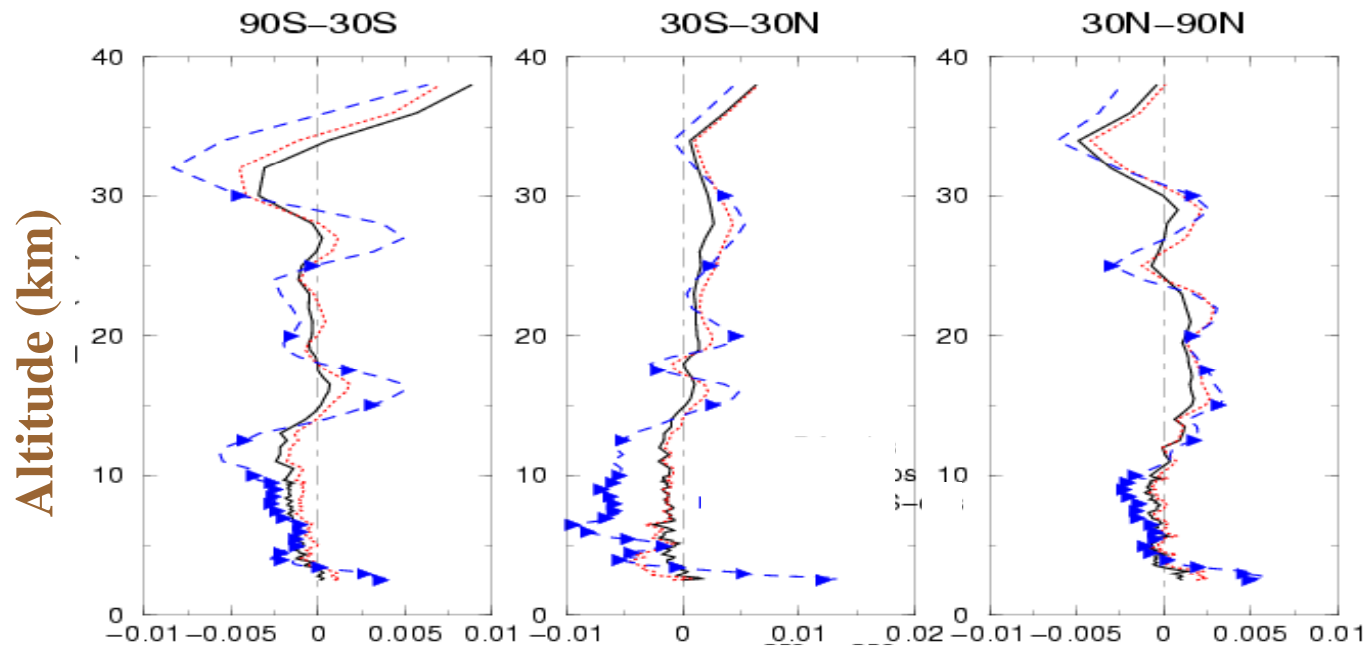
First, $N^{non-loc}$ is calculated based on analyses from three data assimilation experiments in the same way as GPS refractivity is derived, i.e., ray-tracing + Abel inversion

Second, calculate the mean and standard deviation of the analysis-modeling refractivity from observations, i.e.,

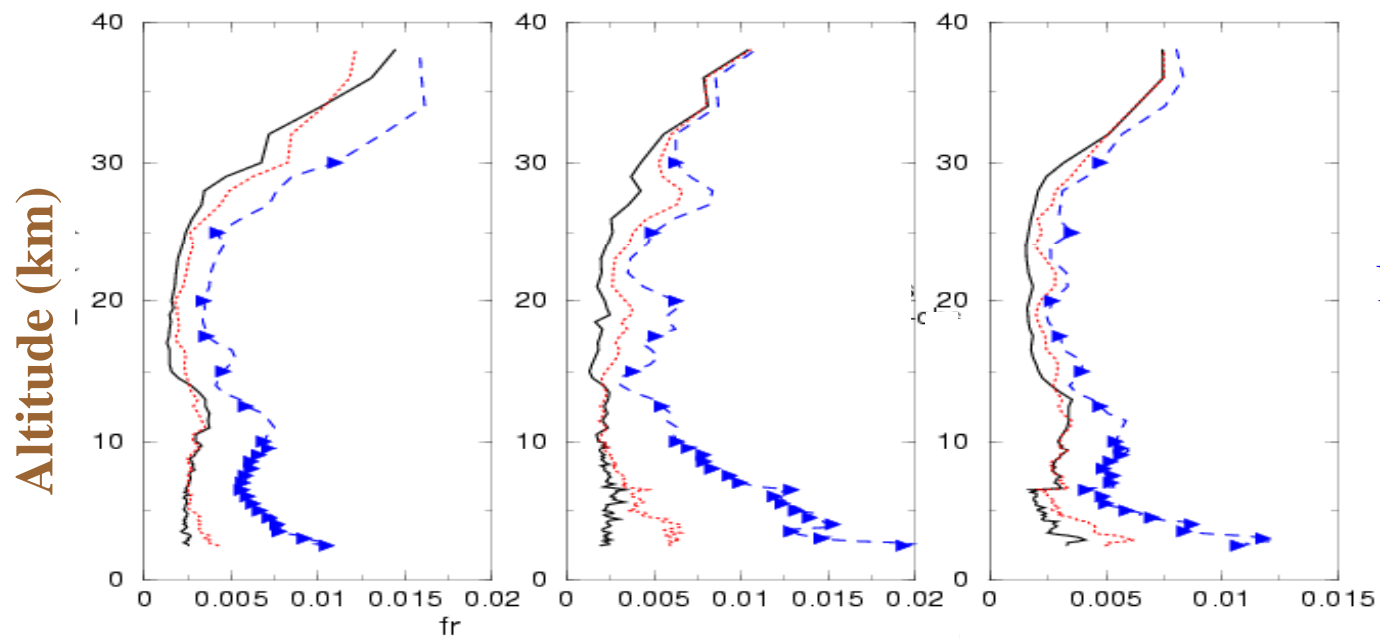
$$N_{NOGPS}^{non-loc} - N_{CHAMP}^{obs}$$

$$N_{BA}^{non-loc} - N_{CHAMP}^{obs}$$

$$N_{REF}^{non-loc} - N_{CHAMP}^{obs}$$



Fractional Mean

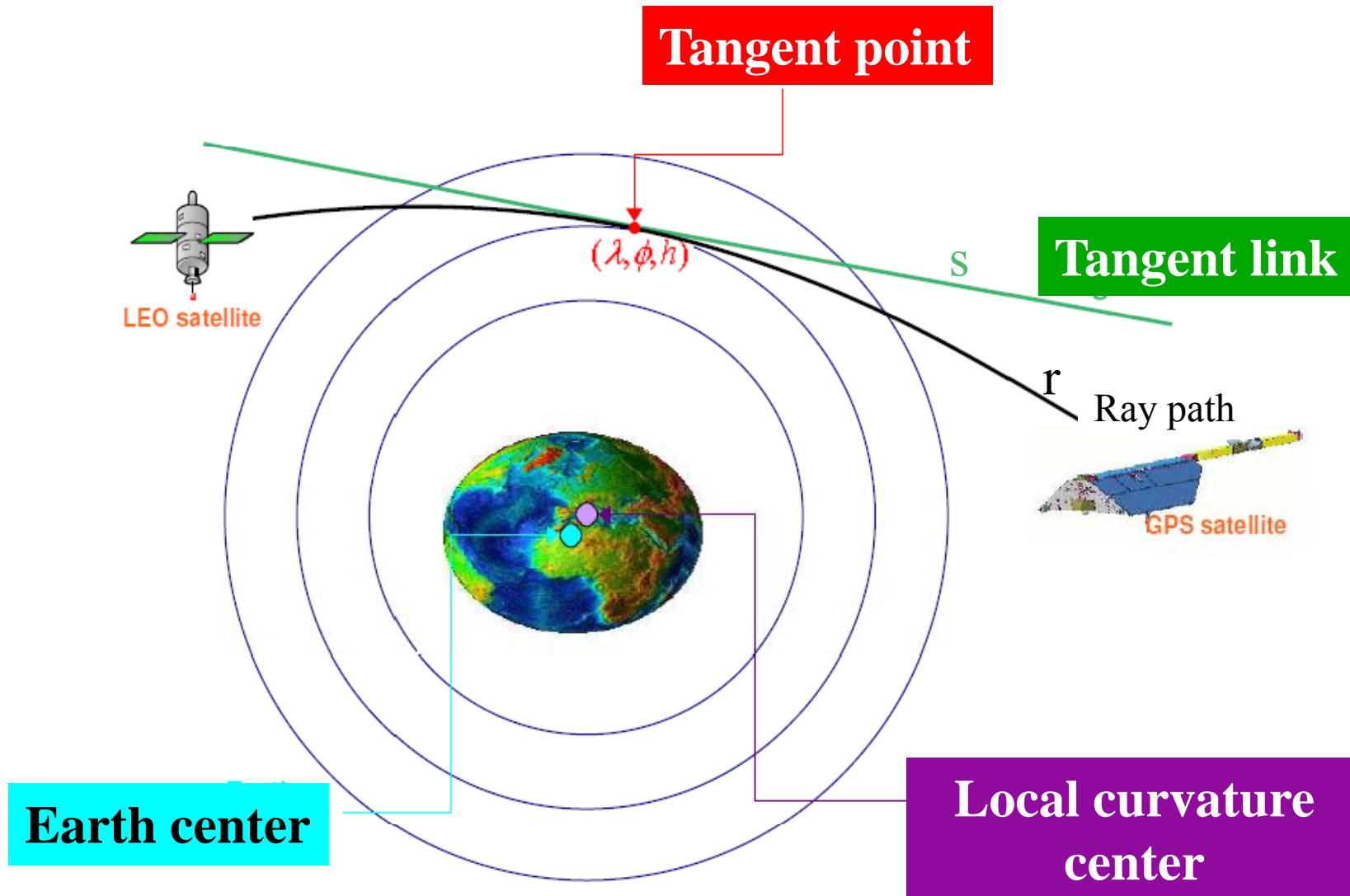


Fractional STD

NOGAP - CHAMP
 BA - CHAMP
 REF - CHAMP

(4b) Comparison between local and non-local refractivity assimilation

Schematic Illustration of Non-Local Operator



Mathematical Expression of Non-Local Operator

$$\int_r N^{obs} dr = \int_r N^{LOC} dr$$

Approximate the integration along a ray path by its tangent link

$$\int_s N^{obs} ds = \int_s N^{LOC} ds$$

Numerical discretization of the two integrals

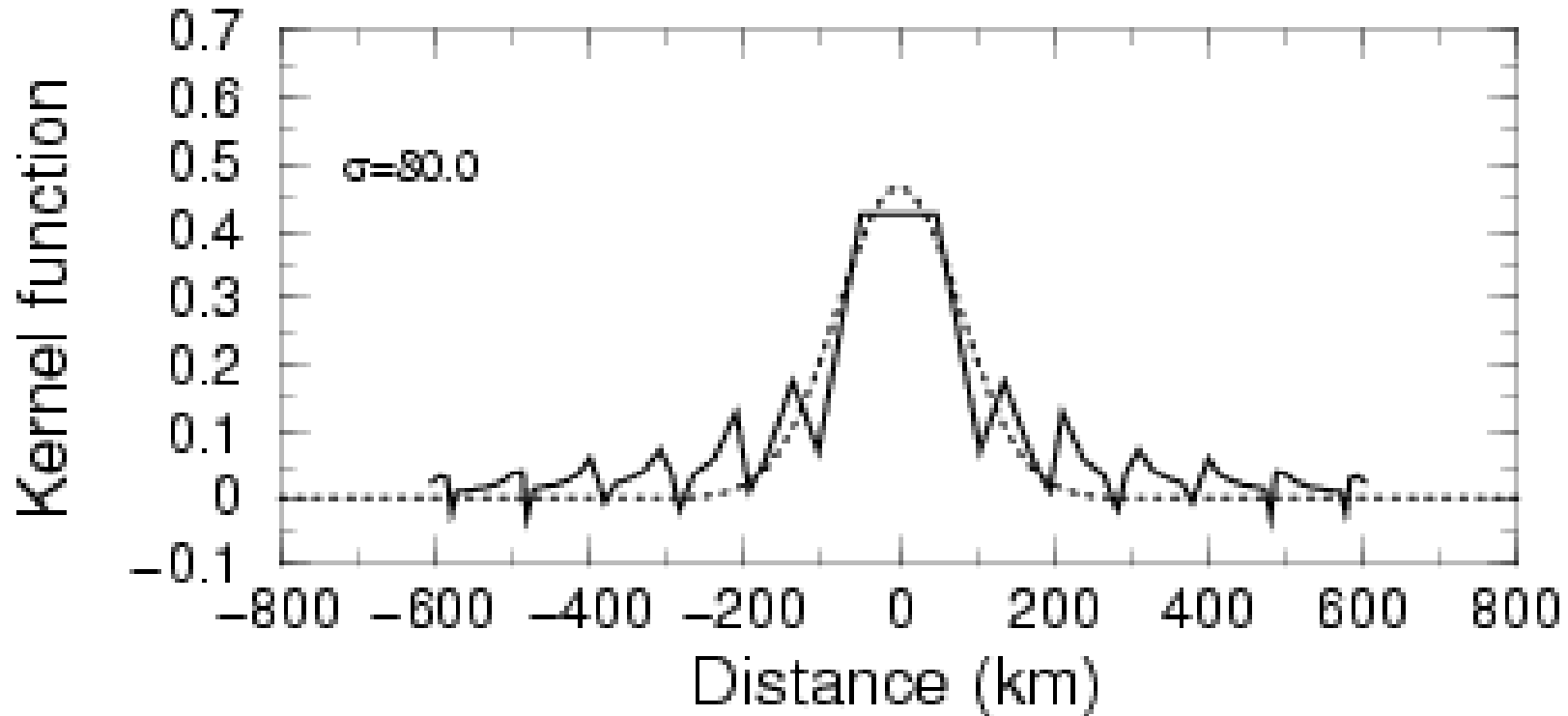
$$\mathbf{B}N^{obs} = \mathbf{A}N^{LOC}$$

$$N^{GPS} = \mathbf{K}N^{LOC}$$

$$\mathbf{K} = \mathbf{B}^{-1}\mathbf{A}$$

GPS refractivity (N^{GPS}) is a weighted sum of local refractivity (N^{LOC}) along the tangent link. \mathbf{K} is the kernel matrix.

An Example of K Distribution



Vertical resolution is 1 *km*. Horizontal resolution is 1°. The tangent link is from west to east. The tangent point is located at (0, 30°N).

Practical Implementation

GPS Observations: $L^{obs} = \mathbf{B}N^{obs}$

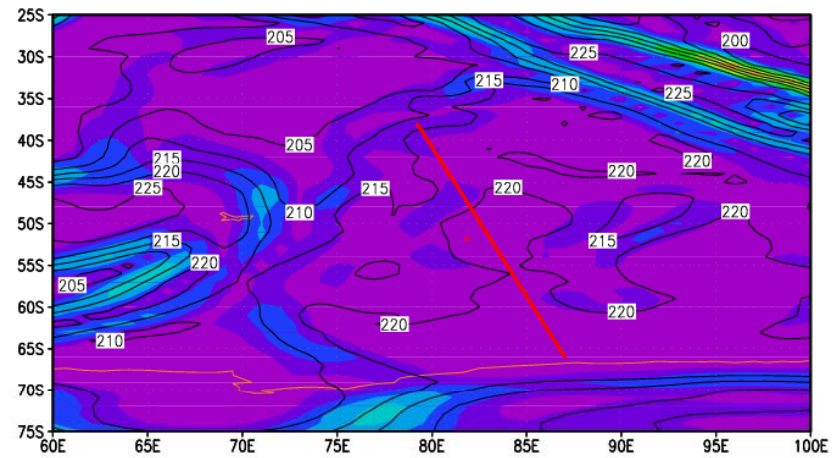
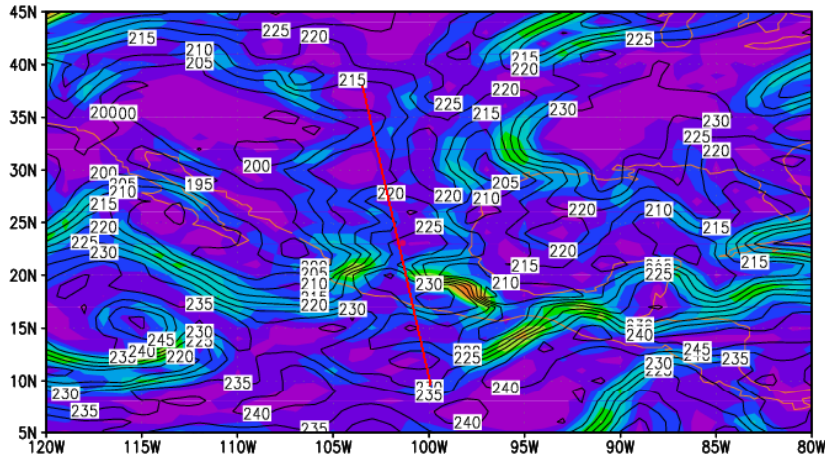
Non-local operator: $L = \mathbf{A}N^{LOC}$

$$J = \mathbf{L} + \left(L - L^{obs} \right)^T \mathbf{O}^{-1} \left(L - L^{obs} \right)$$

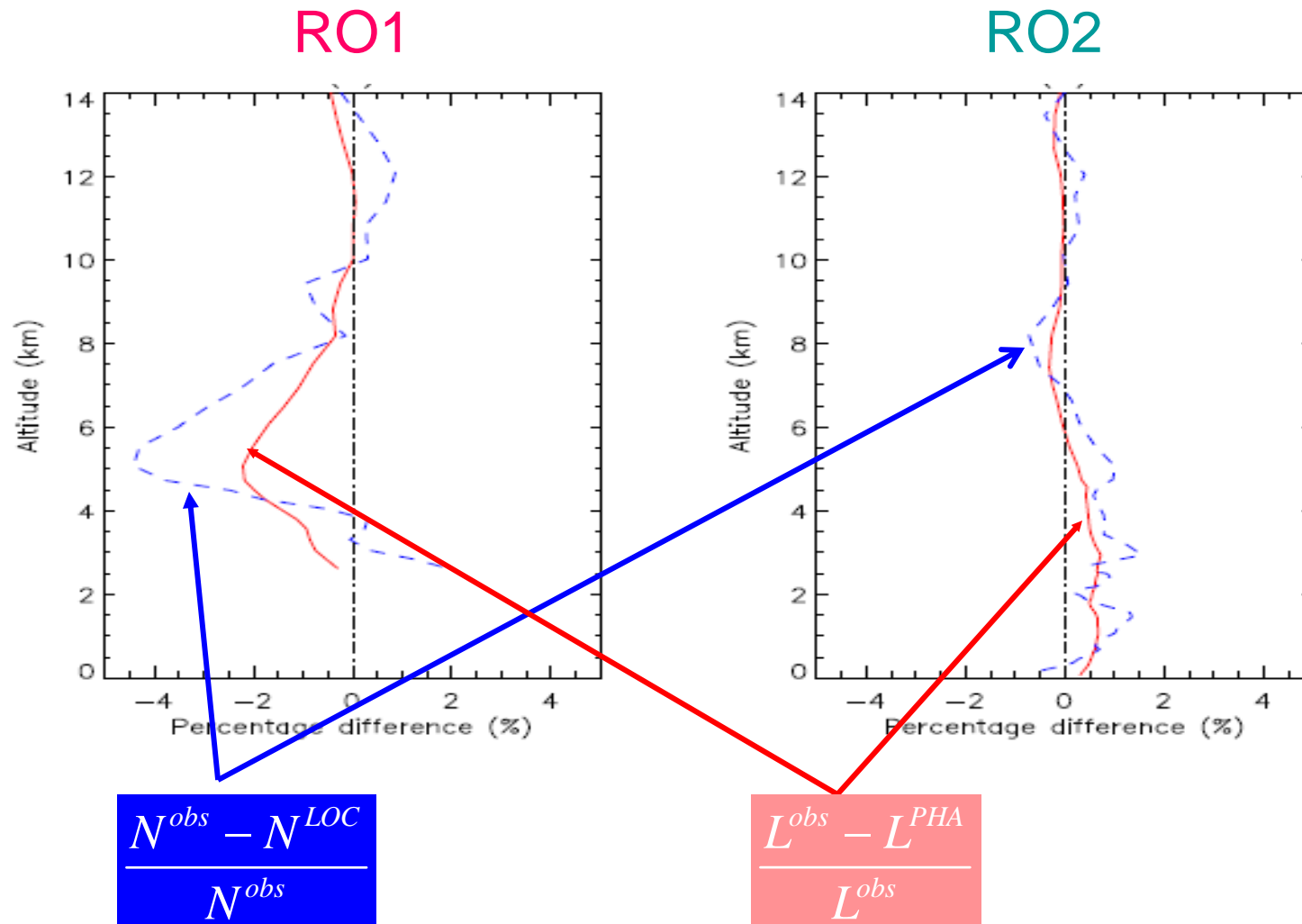
700-hPa Refractivity (contour) and Horizontal Gradients (shaded)

RO1 in large gradient area

RO2 in small gradient area



Differences between Model Simulations and GPS CHAMP Observations



**(5) Comparison with Analyses in
Cloudy and Clear-Sky Conditions**

CloudSat

Instrument: 94-GHz profiling radar

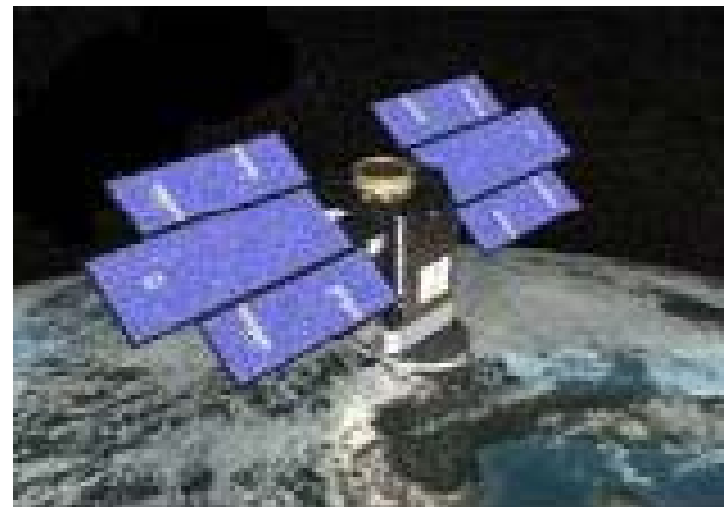
Launch time: April 28, 2006

One orbital time: ~1.5 hours

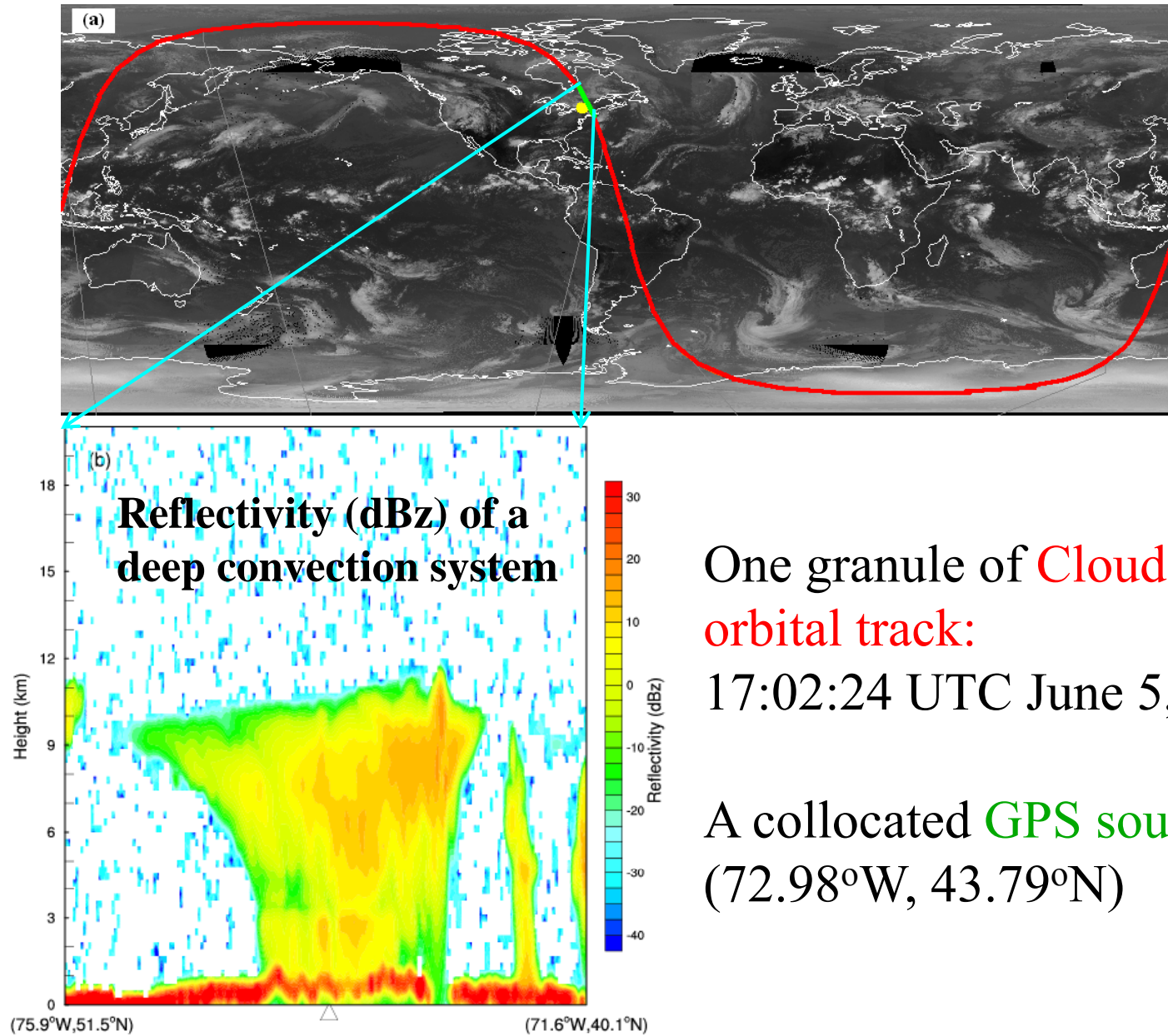
Along-track resolution: ~1.1 km

Track width: ~1.4 km

Observed variables: {
reflectivity
liquid/ice water content
cloud top height
cloud base height
cloud types



A CloudSat Orbital Track and a Collocated GPS RO

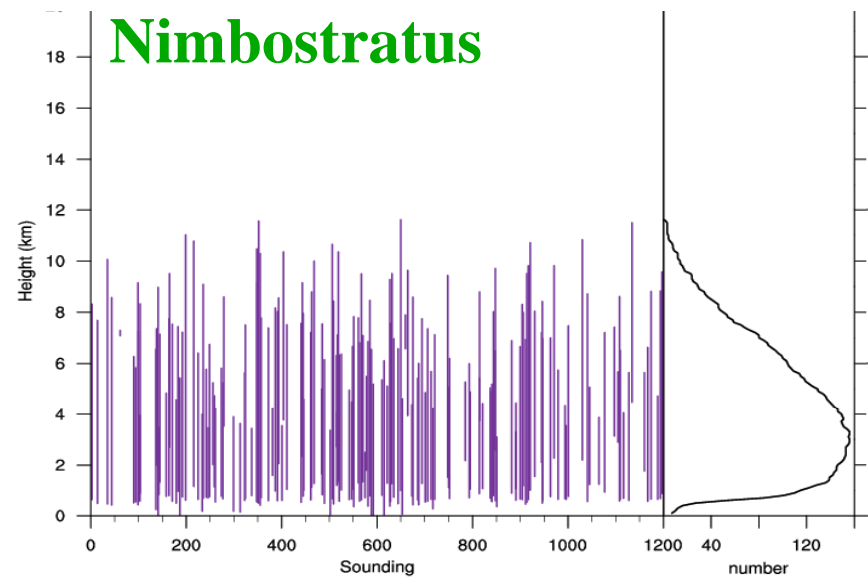
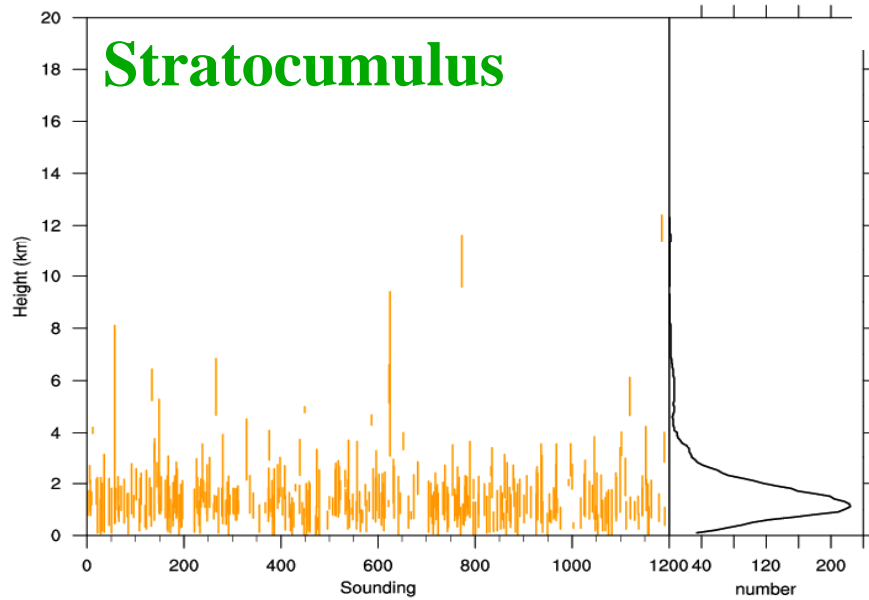
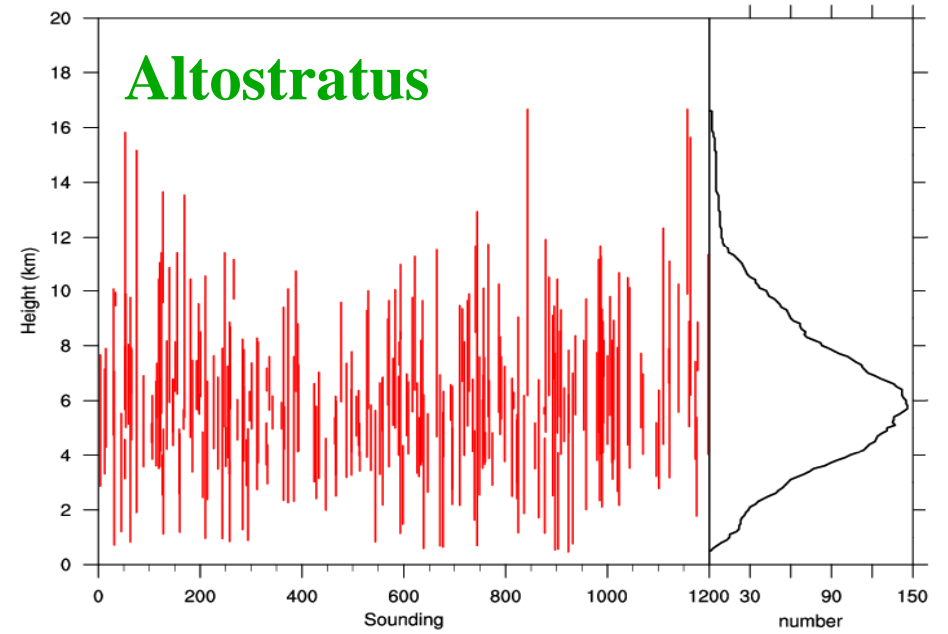
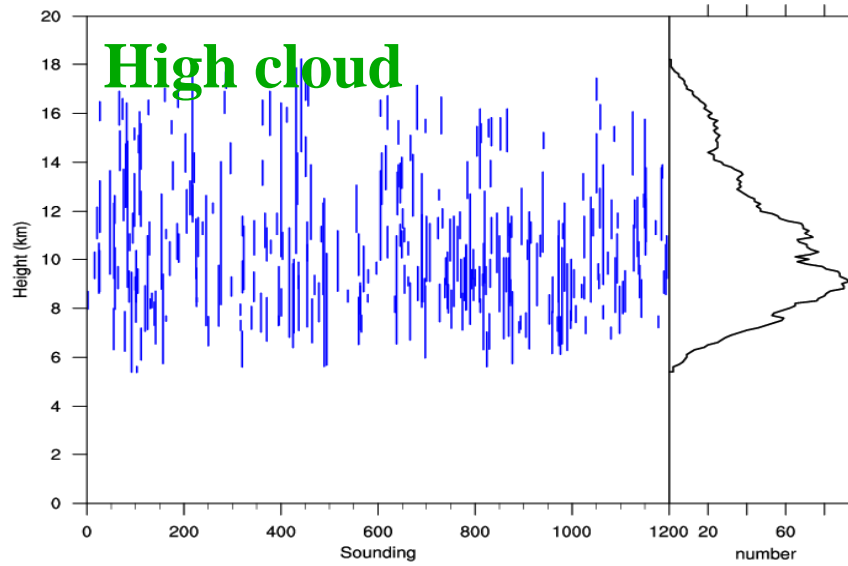


One granule of **CloudSat** orbital track:

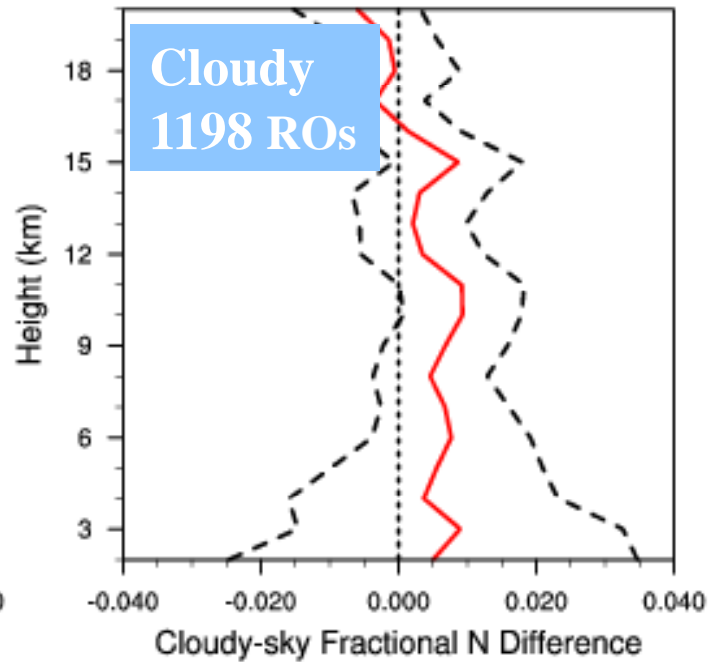
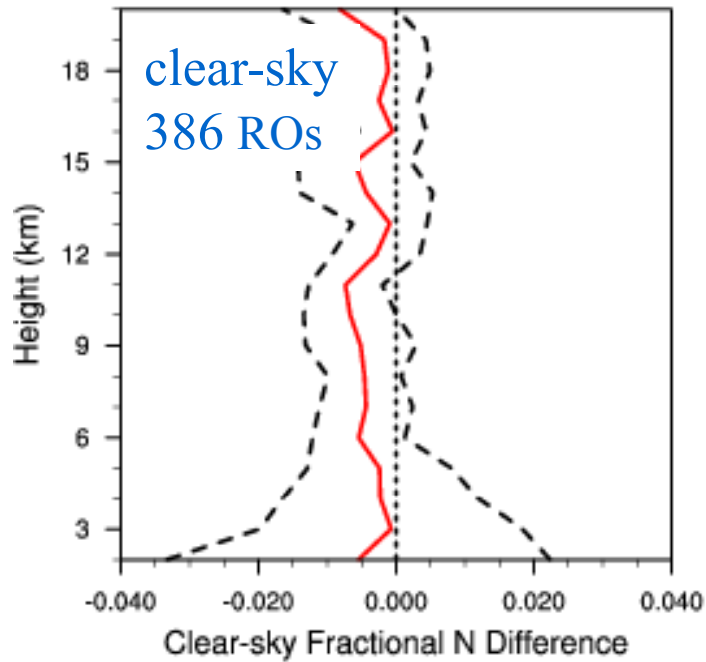
17:02:24 UTC June 5, 2007

A collocated **GPS sounding**:
(72.98°W, 43.79°N)

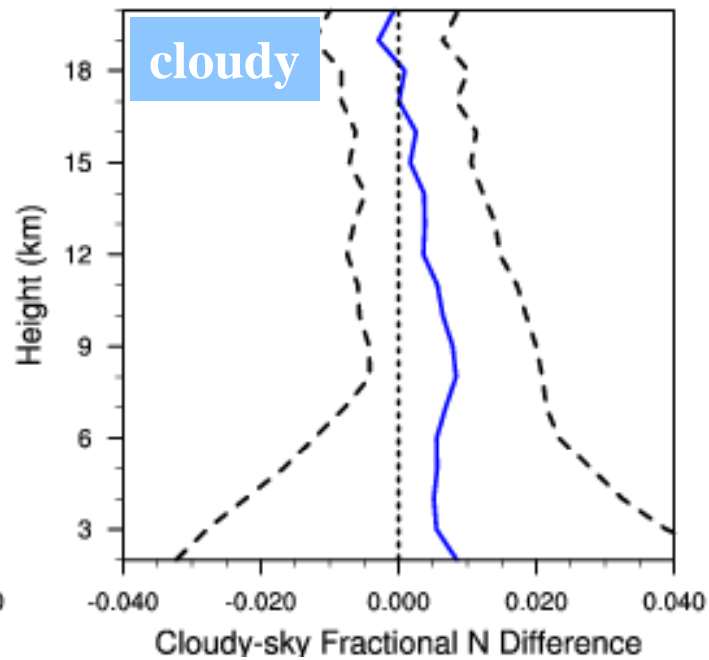
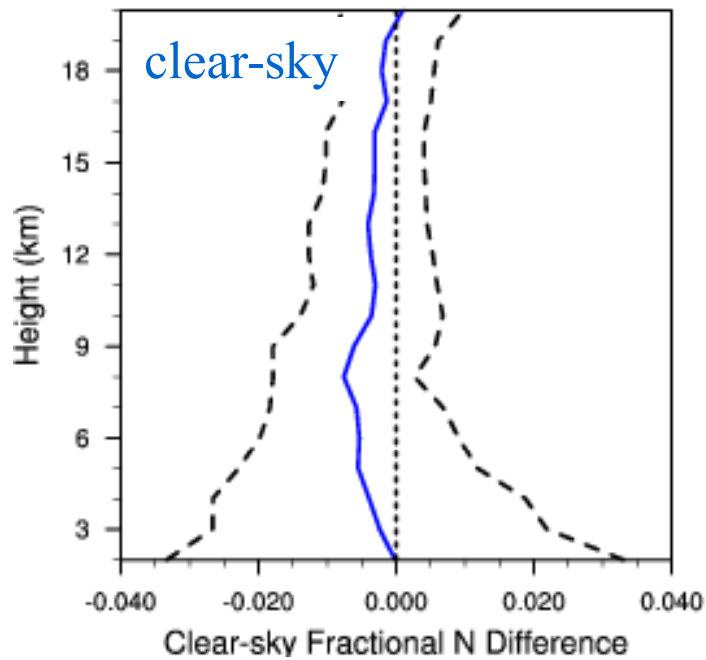
Vertical Extent of Different Cloud Types



Mean/RMS of Fractional N Differences

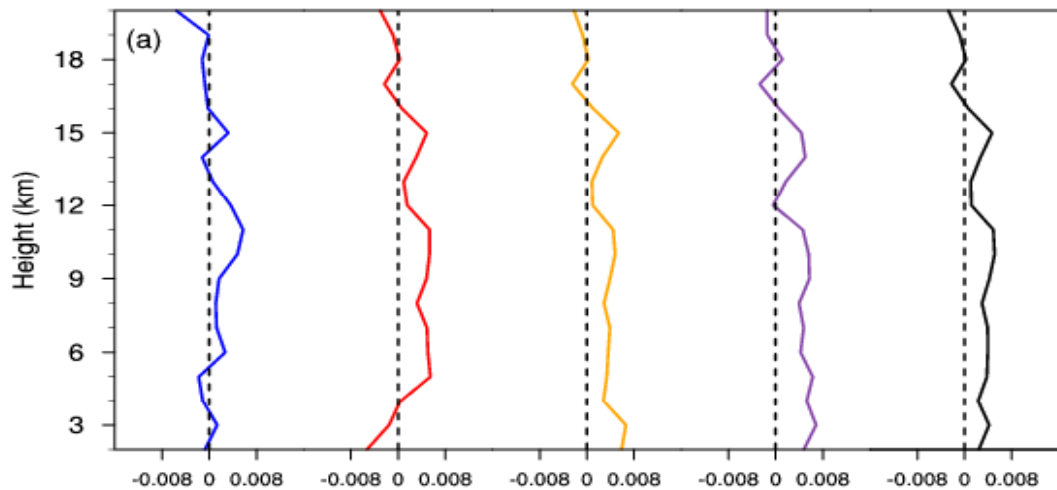


ECMWF



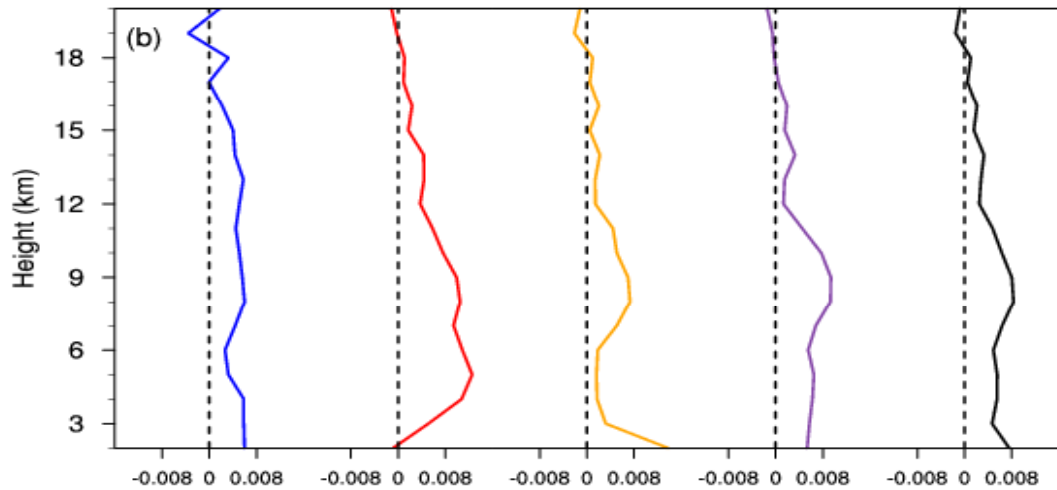
NCEP

09/2007-08/2008



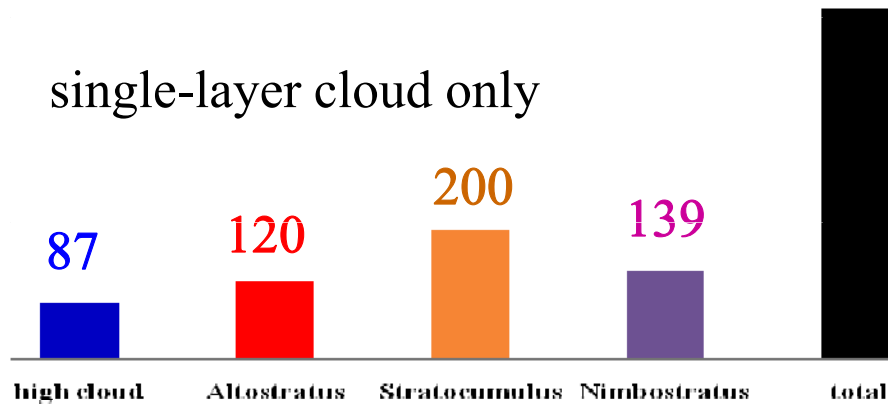
ECMWF

Dependence of Positive *N* Bias on Cloud Type



NCEP

single-layer cloud only



High cloud
 Altostratus
 Stratocumulus
 Nimbostratus

(6) Cloudy Retrieval of GPS RO Data

GPS Cloudy Retrieval Algorithm

Assumption: Cloudy air is saturated.

Atmospheric refractivity for cloudy air

GPS observation

$$N = 77.6 \frac{P}{T} + 3.73 \times 10^5 \frac{e_s(T)}{T^2} + 1.45W = N(T, P)$$

dry term

wet term

liquid water
term

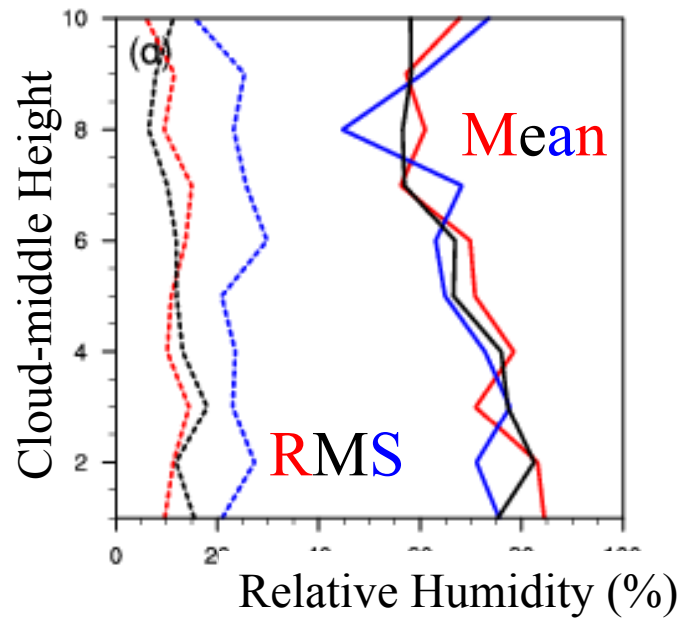
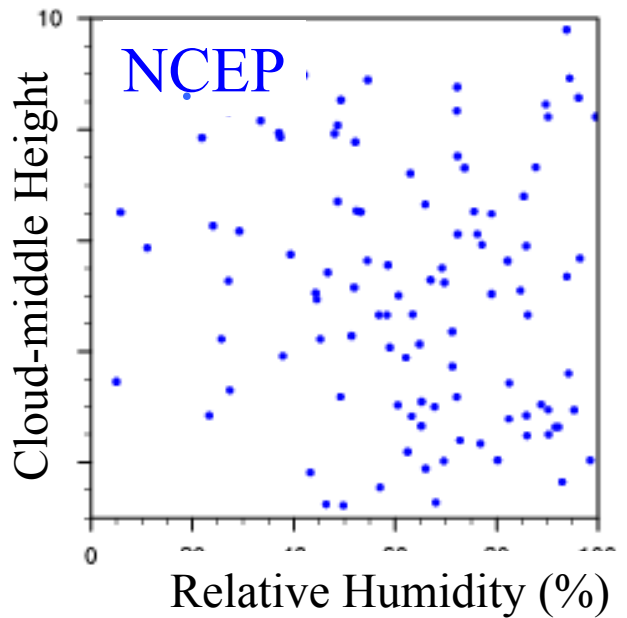
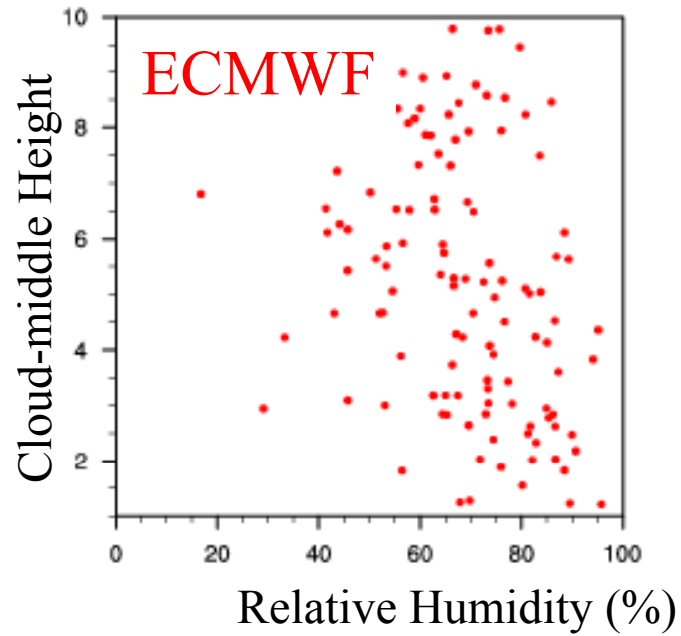
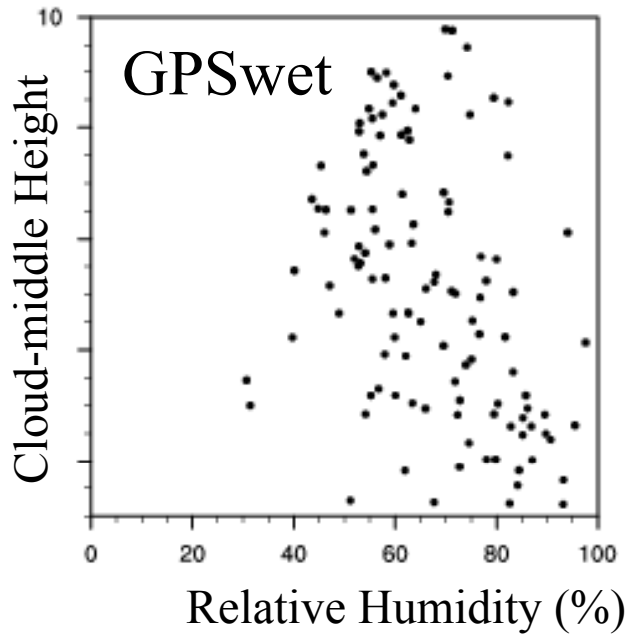
Hydrostatic equation:

$$\frac{1}{P} \frac{dP}{dz} = - \frac{g}{R_d T (1 + 0.61 q_s)}$$

We have two equations for two unknown variables T and P .

In-cloud profiles of T and p can be uniquely determined from GPS ROs given initial conditions at the cloud top.

Mean Relative Humidity within Clouds



GPS Refractivity within Cloud

Cloud occupies only a fraction of an analysis grid box.

Atmospheric refractivity for cloudy air

$$N^{obs} = N^{dry} + (1 - \alpha) \cdot N^{wet} + \alpha \cdot N^{sat}$$

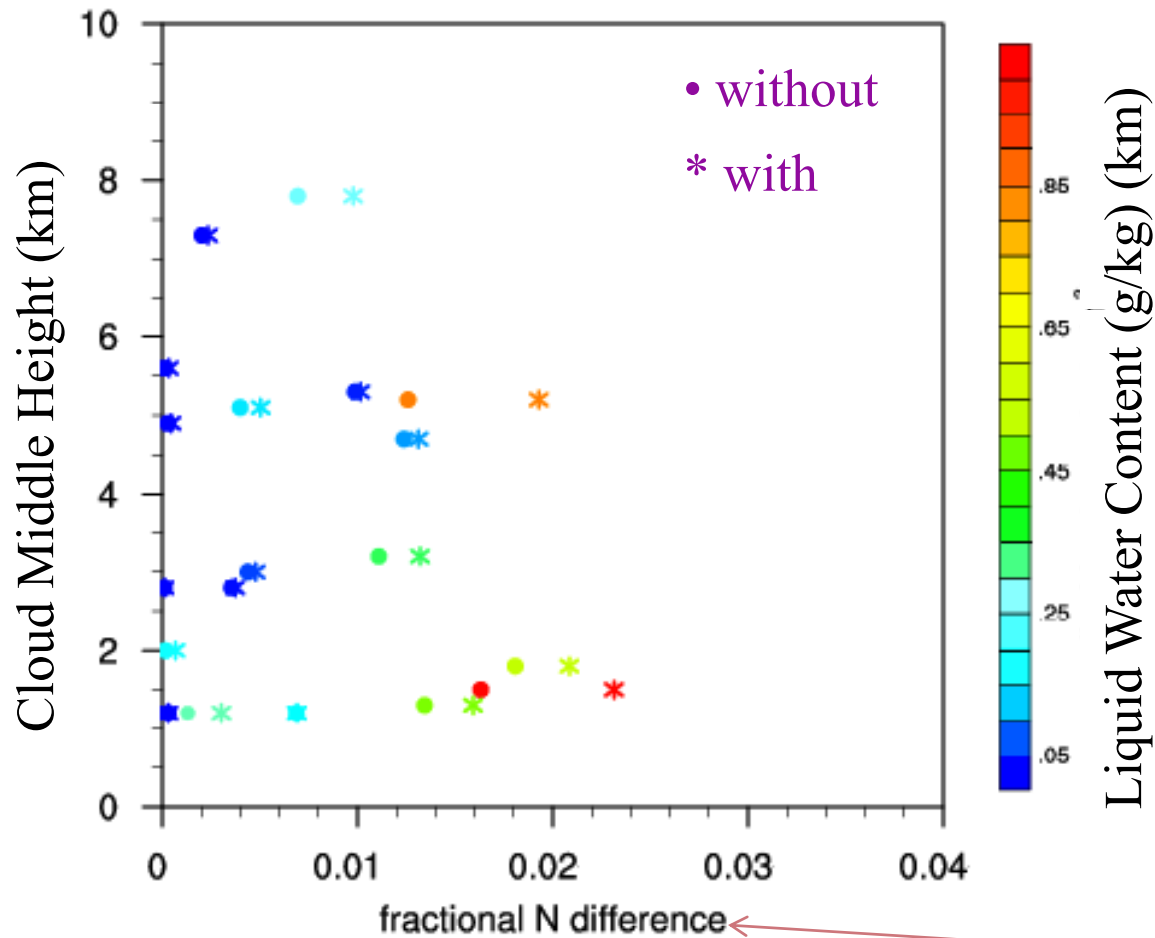
↑
relative humidity parameter

where

$$\alpha = \begin{cases} 5.273 \times IWC + 0.6849, & \text{if } IWC \leq 0.05975 \text{ gm}^{-3} \\ 1, & \text{if } IWC > 0.05975 \text{ gm}^{-3} \\ 0.85, & \text{for liquid water cloud} \end{cases}$$

Importance of Liquid Water

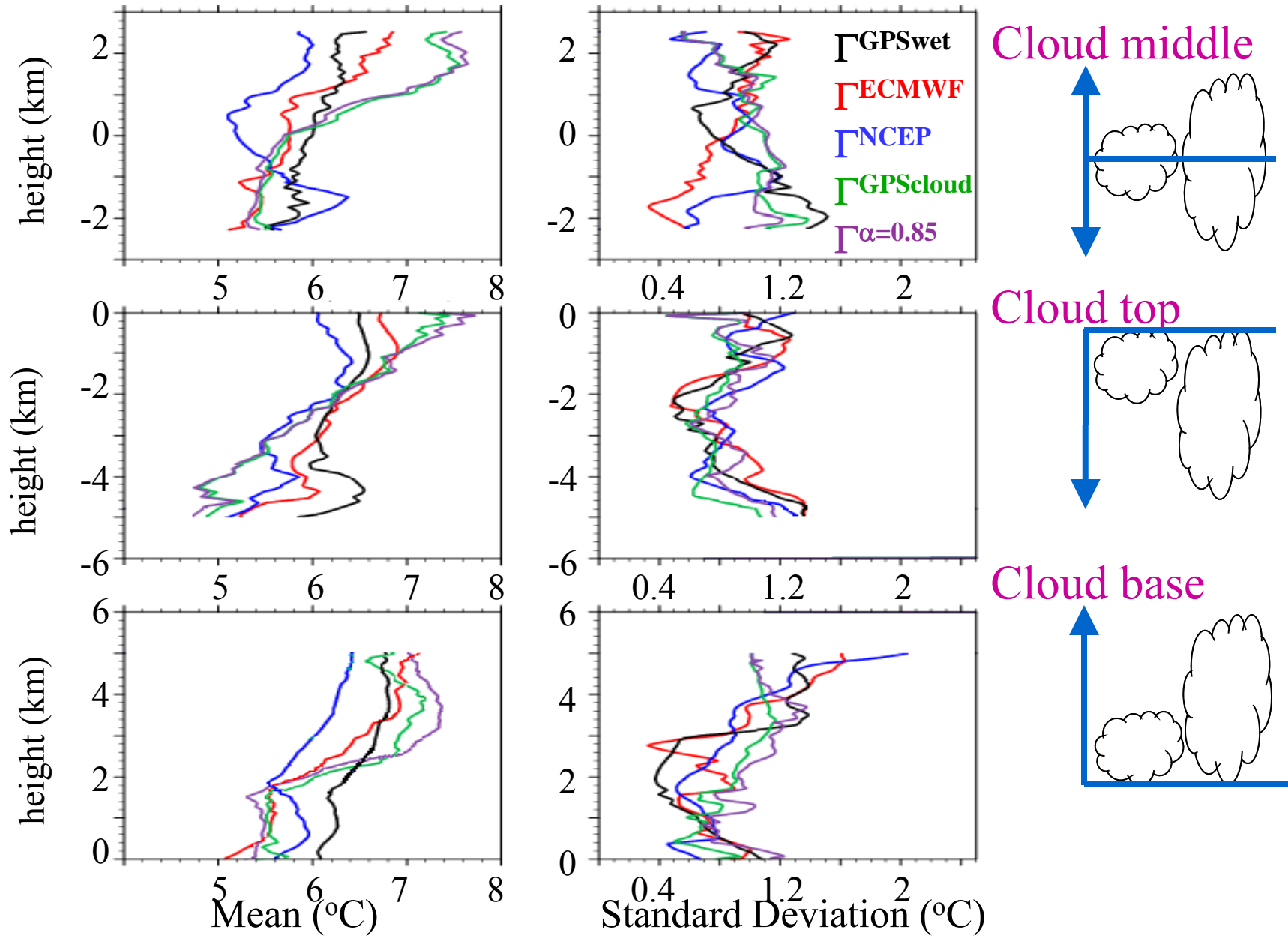
(with and without including cloud liquid water term)



(19 water clouds identified by CloudSAT)

$$\frac{N^{GPS} - N^{ECMWF}}{N^{ECMWF}}$$

Lapse Rates within Cloud



Future Plans

- 1. Global cloud climatology derivable from GPS ROs**
- 2. Optimal mix of microwave satellite observations with GPS RO data for cloud data assimilation**
- 3. Validation of moisture physical parameterization schemes using GPS RO data**

(7) Publications on space-borne GPS RO data assimilation

- (1) Zou, X., Y.-H. Kuo, and Y.-R. Guo, 1995: Assimilation of atmospheric radio refractivity using a nonhydrostatic adjoint model. *Mon. Wea. Rev.*, **123**, 2229-2249.
- (2) Ware, R., M. Exner, D. Feng, M. Gorbunov, K. Hardy, B. Herman, Y.-H. Kuo, T. Meehan, W. Melbourne, C. Rocken, W. Schreiner, S. Sokolovski, F. Solheim, X. Zou, R. Anthes, and S. Businger, 1996: GPS sounding of the atmosphere from low earth orbit: Preliminary results. *Bull. Am. Meteor.Soc.*, **77**, 19-40.
- (3) Rocken, C., R. Anthes, M. Exner, D. Hunt, S. Sokolovskiy, R. Ware, M. Gorbunov, W. Schreiner, D. Feng, B. Herman, Y.-H. Kuo, and X. Zou, 1997: Analysis and validation of GPS/MET data in the neutral atmosphere. *J. Geophys. Res.*, **102**, 29,849-29,866.
- (4) Kuo, Y.-H., X. Zou, and W. Huang, 1998: The impact of GPS data on the prediction of an extratropical cyclone: An observing system simulation experiment, *Dyn. Atmos. Ocean.*, **27**, 439-470.
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(7) Publications (cont.)

- (6) Zou, X., F. Vandenberghe, B. Wang, M. E. Gorbunov, Y.-H. Kuo, S. Sokolovskiy, J. C. Chang, J. G. Sela, and R. Anthes, 1999: A raytracing operator and its adjoint for the use of GPS/MET refraction angle measurements. *J. Geoph. Res.*, **104**, 22,301-22,318.
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- (8) Liu, H., X. Zou, R. A. Anthes, J. C. Chang, J.-H. Tseng, and B. Wang, 2001: The Impact of 837 GPS/MET bending angle profiles on assimilation and forecasts for the period June 20-30, 1995, *J. Geoph. Res.*, **106**, 31771-31786.
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- (10) Shao H., and X. Zou, 2002: On the observational weighting and its impact on GPS/MET bending angle assimilation. *J. Geoph. Res.*, **107**, ACL 19, 1-28.

(7) Publications (cont.)

- (11) Liu, H., and X. Zou, 2003: Improvements to a forward GPS raytracing model and their impacts on assimilation of bending angle, *J. Geoph. Res.*, **108**, D17, 4548.
- (12) Zou, X., H. Liu, R. A. Anthes, H. Shao, J. C. Chang, and Y.-J. Zhu, 2004: Impact of CHAMP occultation observations on global analysis and forecasts in the absence of AMSU radiance data. *Journal of the Meteorological Society of Japan*, **82**, 533-549.
- (13) Zou, X. and Z. Zeng, 2006: A quality control procedure for GPS RO data. *J. Geoph. Res.*, **111**, D02112, doi:10.1029/2005JD005846.
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- (16) Lin, L., X. Zou, R. Anthes and Y.-H. Kuo, 2010: COSMIC GPS RO temperatures profiles in clouds. *Mon. Wea. Rev.*, doi:10.1175/2009MWR2986.1.