



Uranium Reactive Transport in the Hanford 300 Area Vadose Zone-Aquifer-River System

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W.R. Wiley, Environmental Molecular Sciences Laboratory at PNNL

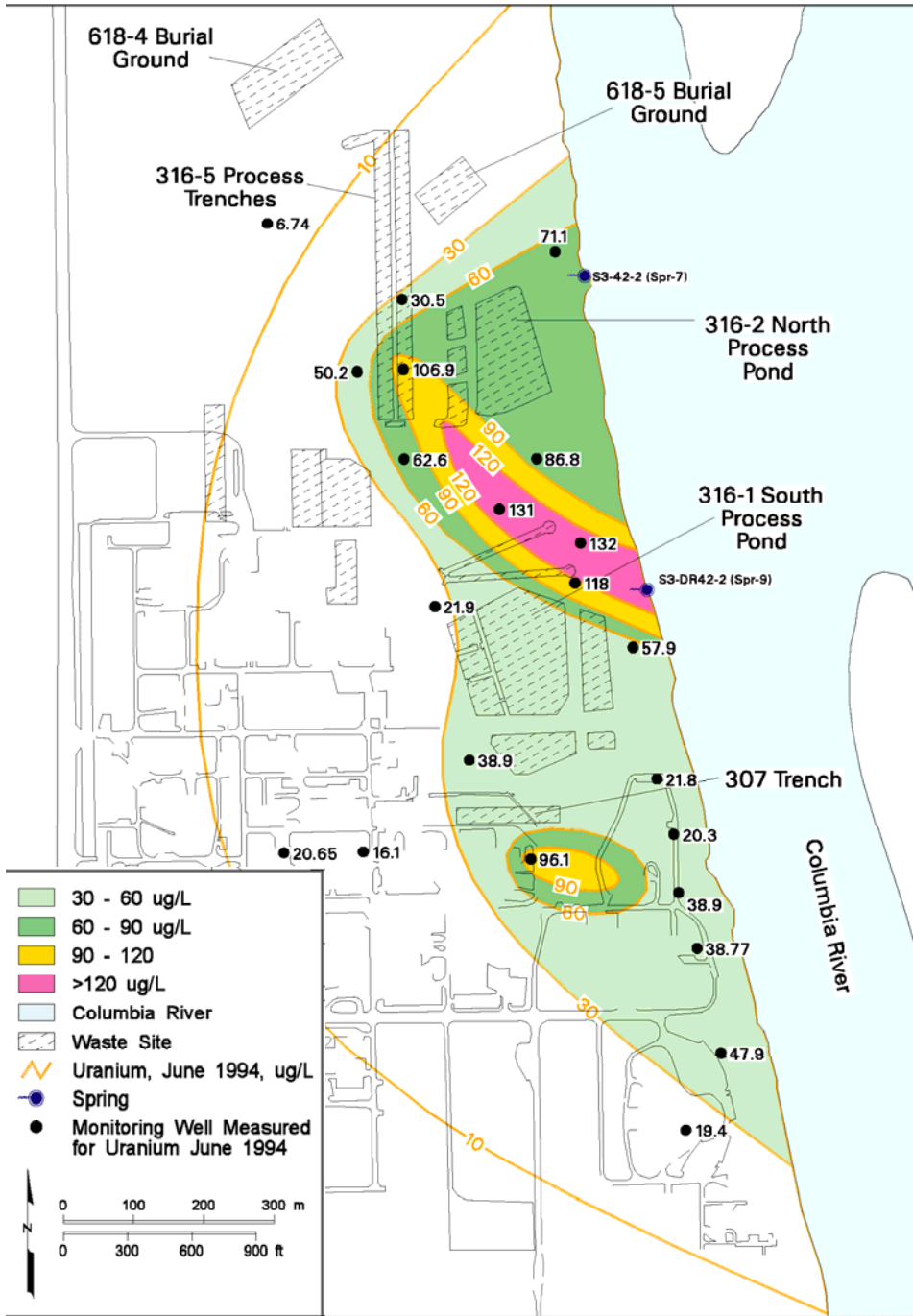
1993 Conceptual Model

Modeling Assumptions in Phase I Remedial Investigation

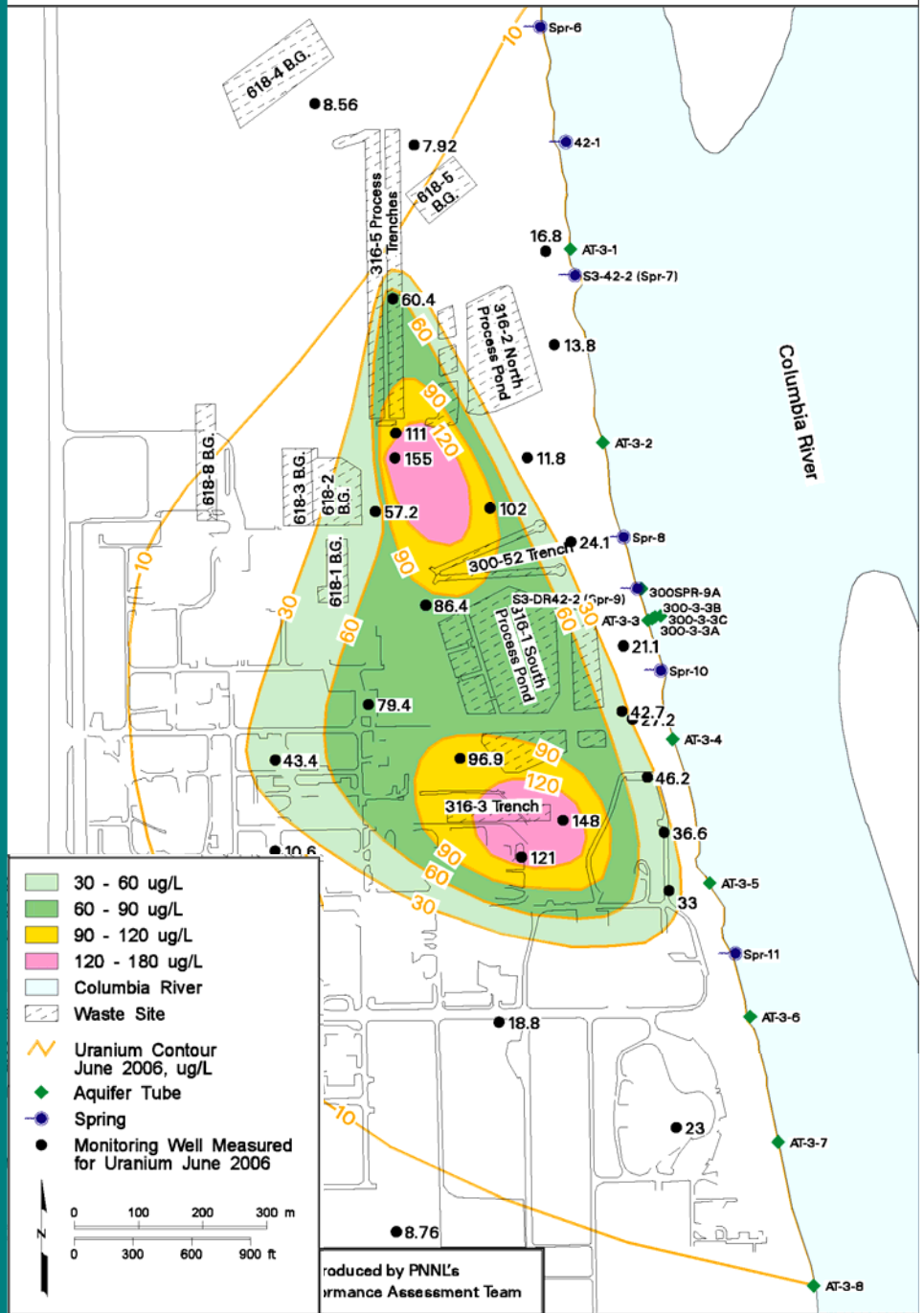
- ▶ 3-D saturated unconfined aquifer; **vadose zone not modeled**
 - Flow field driven by **monthly** changes in river stage fluctuations
 - Uranium mobility controlled by “best estimate” **constant K_d ~1-2 ml/g**
- ▶ No interaction between aquifer and river
- ▶ No interaction between aquifer and vadose zone

Final Prediction: $U < 20$ ug/L in **3 to 10 years** by natural flushing

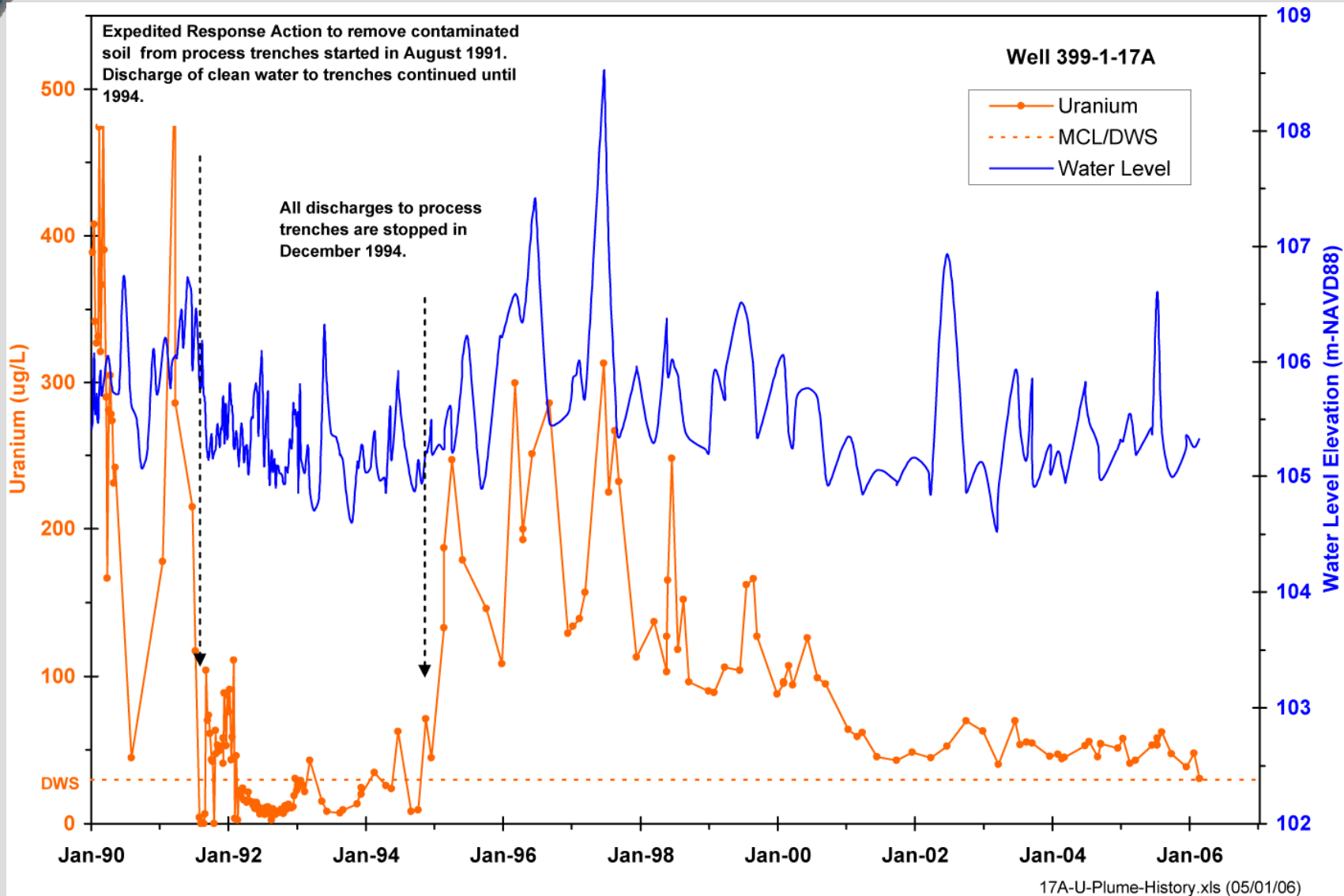
Shaded 300 Area Uranium, June 1994



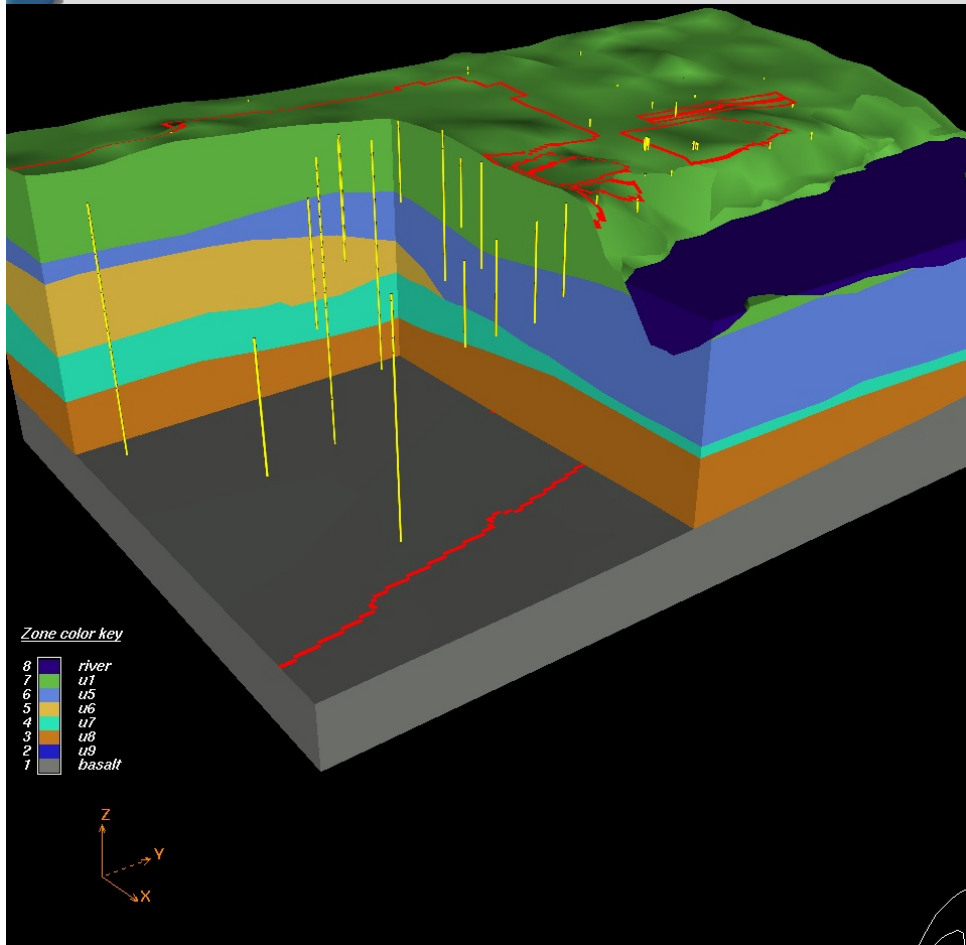
Shaded 300 Area Uranium, June 2006



Aquifer Water Levels and Uranium Concentrations

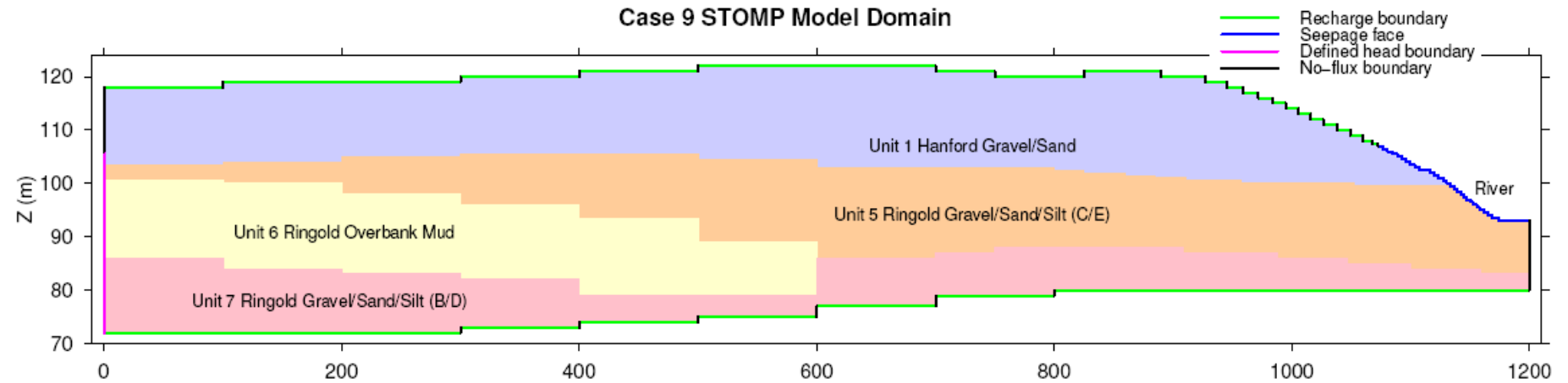


Flow and Transport: Vadose Zone – Aquifer – River System



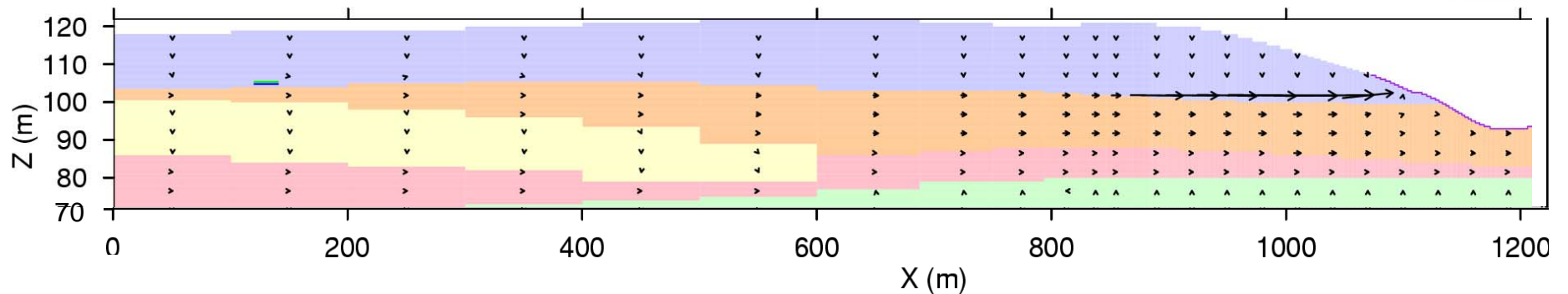
- ▶ 2-D and 3-D modeling with STOMP simulator
- ▶ Most current hydrogeology
- ▶ Flow and transport driven by hourly river stage fluctuations
- ▶ Investigate dynamics of riverbank storage and fluxes across aquifer - river interface
- ▶ Investigate release of uranium from contaminated vadose zone sediments due to water table fluctuations

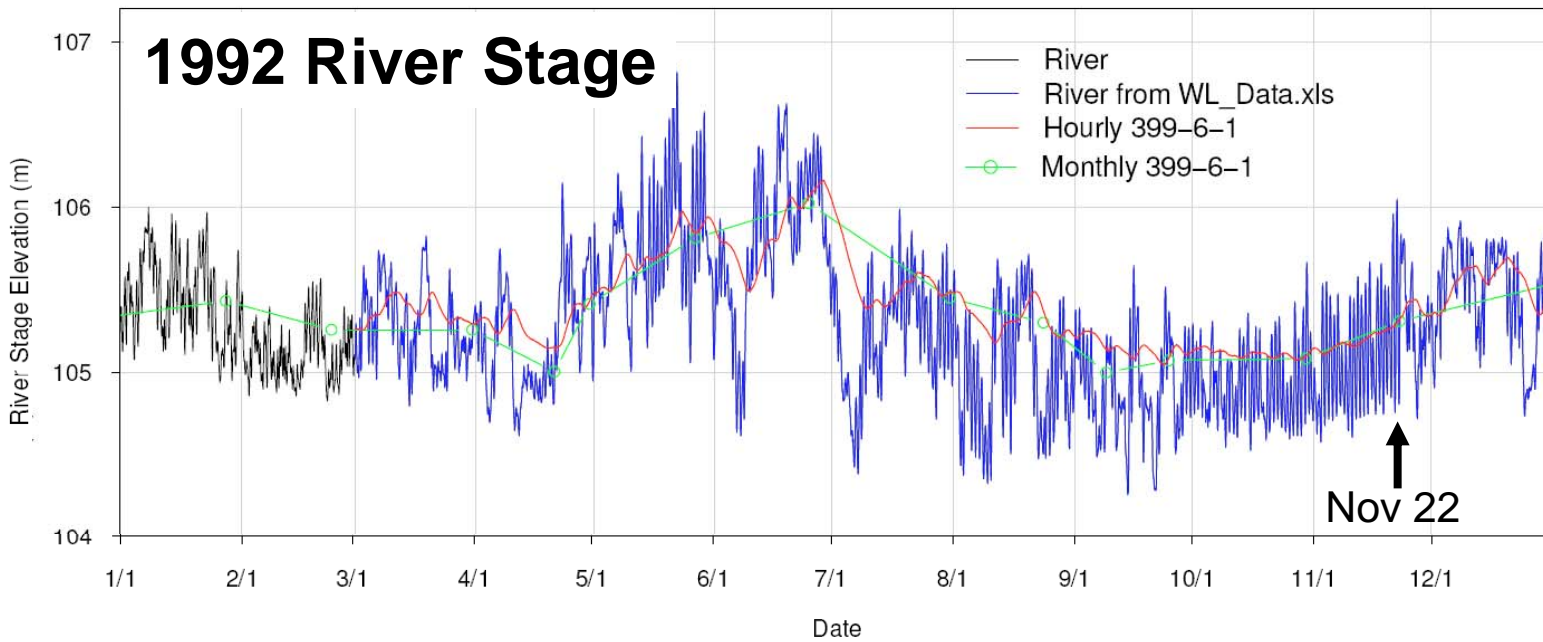
Flow and Transport: Vadose Zone – Aquifer – River System



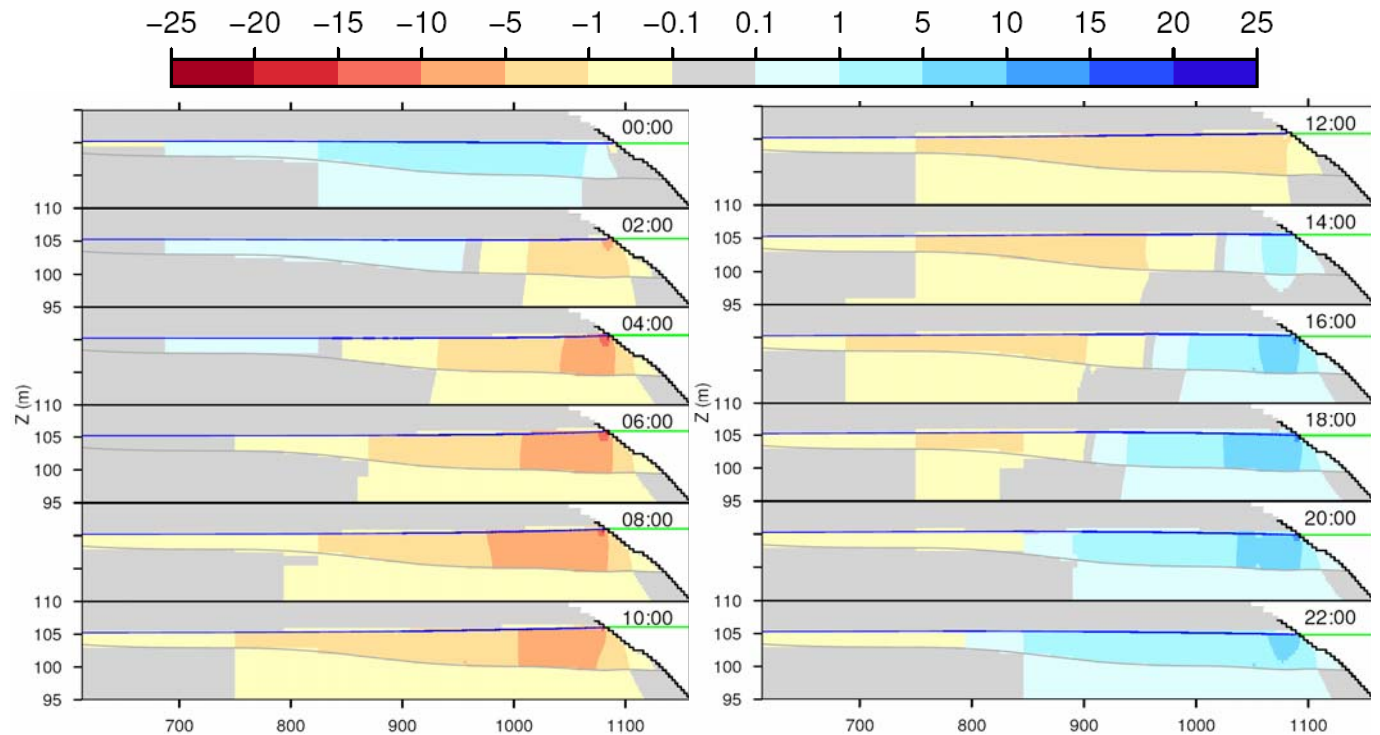
Material	$K_I^{(a)}$ (m d^{-1})	$\phi^{(b)}$	$\psi^{(c)}$ (cm)	$\lambda^{(d)}$	$\theta_r^{(e)}$
U1, Hanford Gravel/Sand	1500	0.25	23.04	0.7465	0.1471
U5, Ringold Gravel/Sand/Silt (C/E)	150	0.18	71.31	0.5193	0.1299
U6, Ringold Overbank Mud	0.01	0.18	71.31	0.5193	0.1299
U7, Ringold Gravel/Sand/Silt (B/D)	43	0.18	71.31	0.5193	0.1299
U8, Lower Ringold Mud	5e-5	0.18	71.31	0.5193	0.1299
Basalt	5e-5	0.18	71.31	0.5193	0.1299

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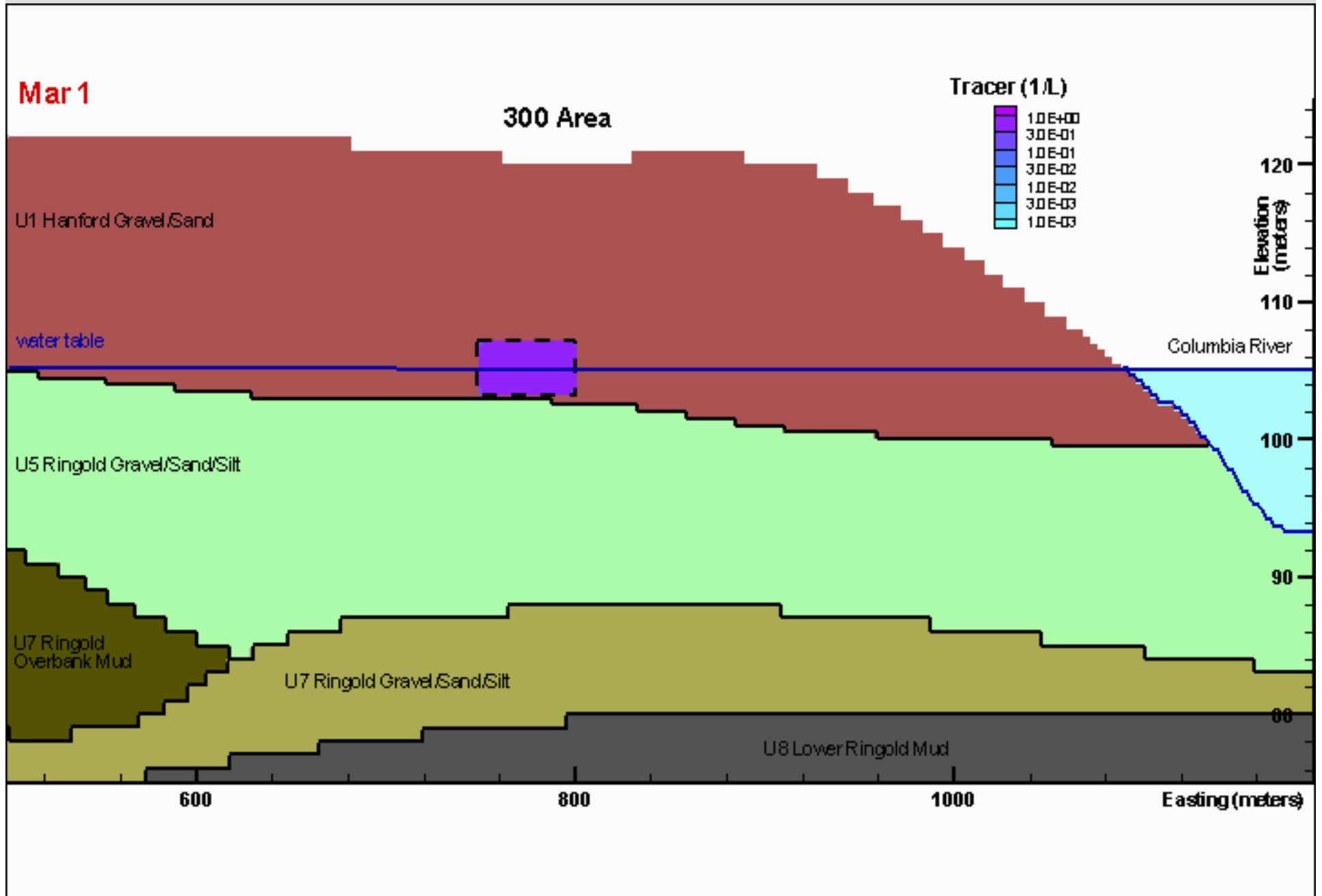




X-Direction Groundwater Flux (m/d) (Nov 22, 1992)



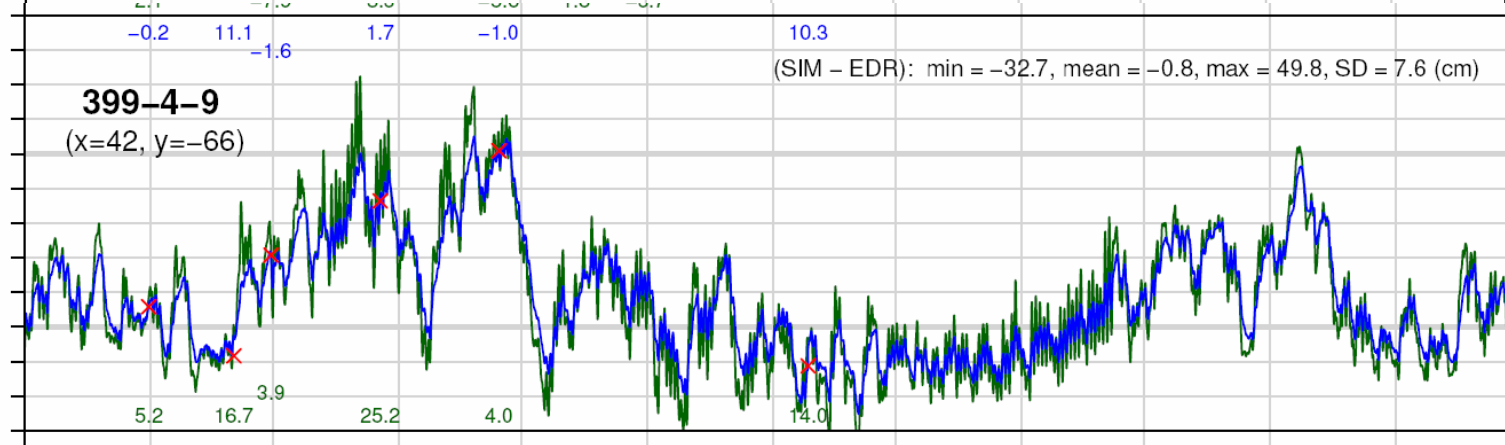
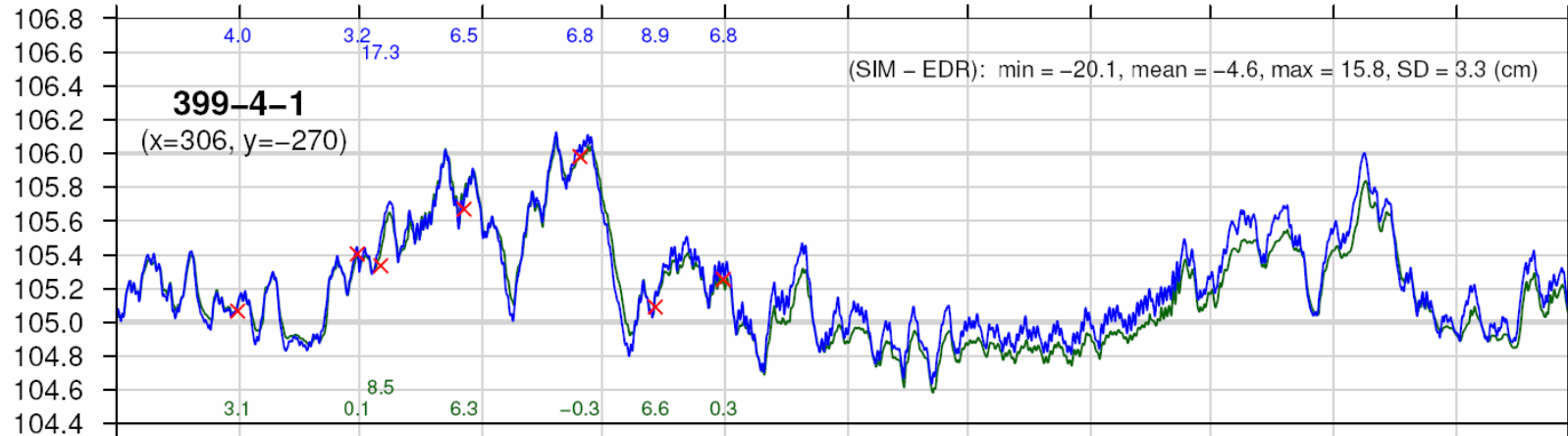
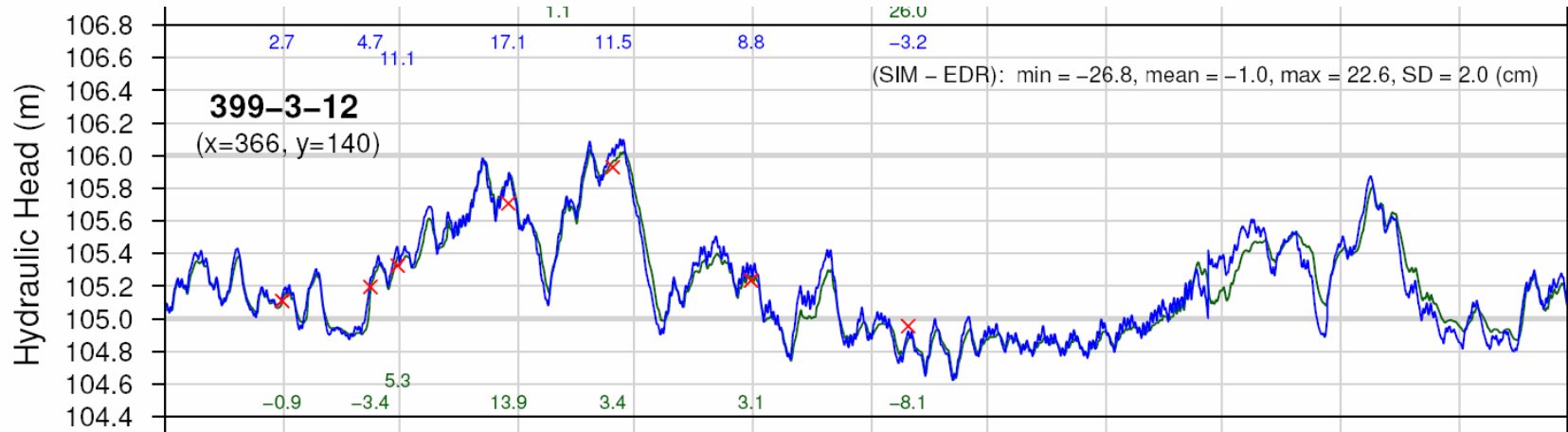
Tracer Transport



— STOMP, sp. interp. — Edrington × HEIS

Green numbers are (SIM - HEIS) in cm

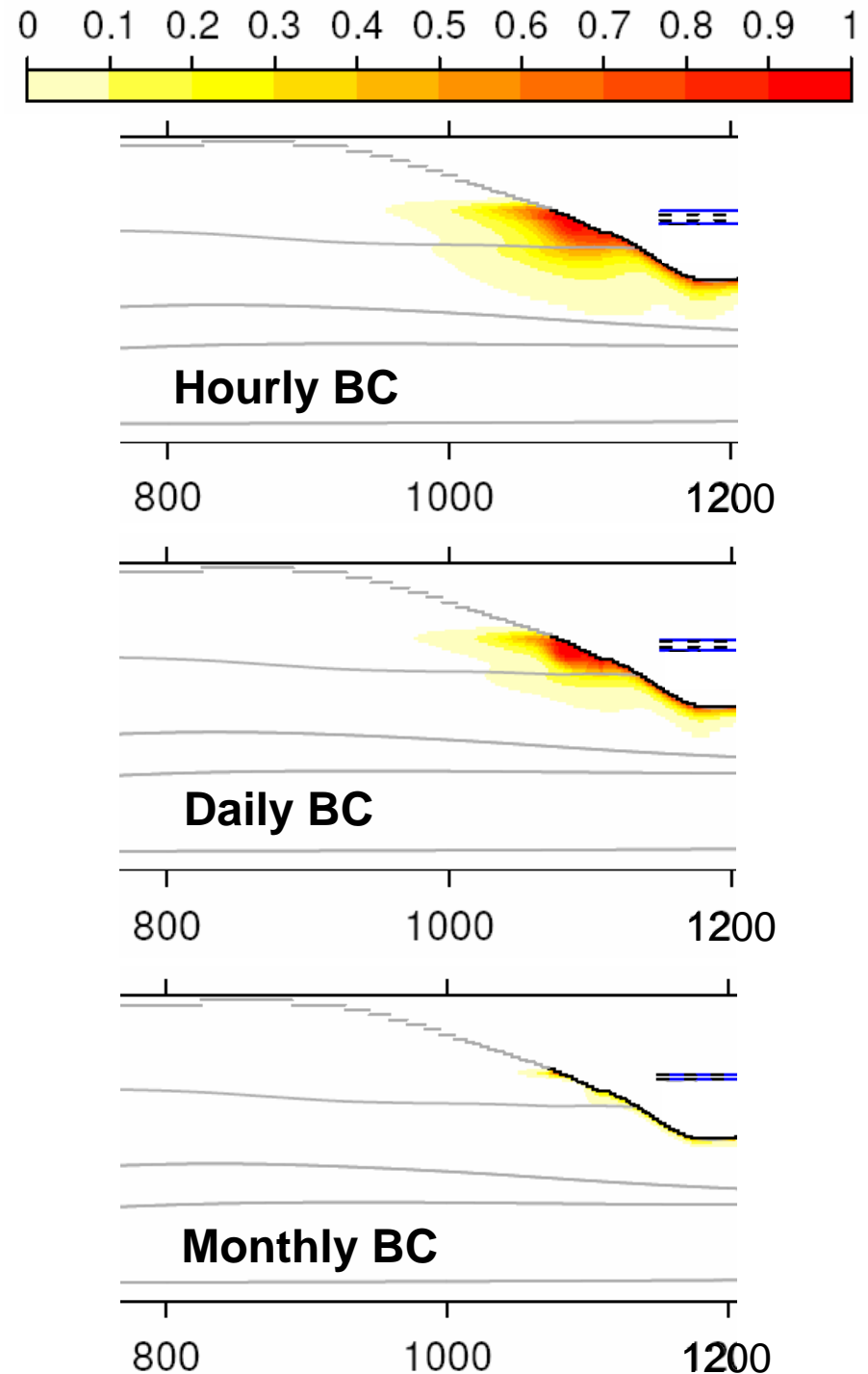
Blue numbers are (EDR - HEIS) in cm



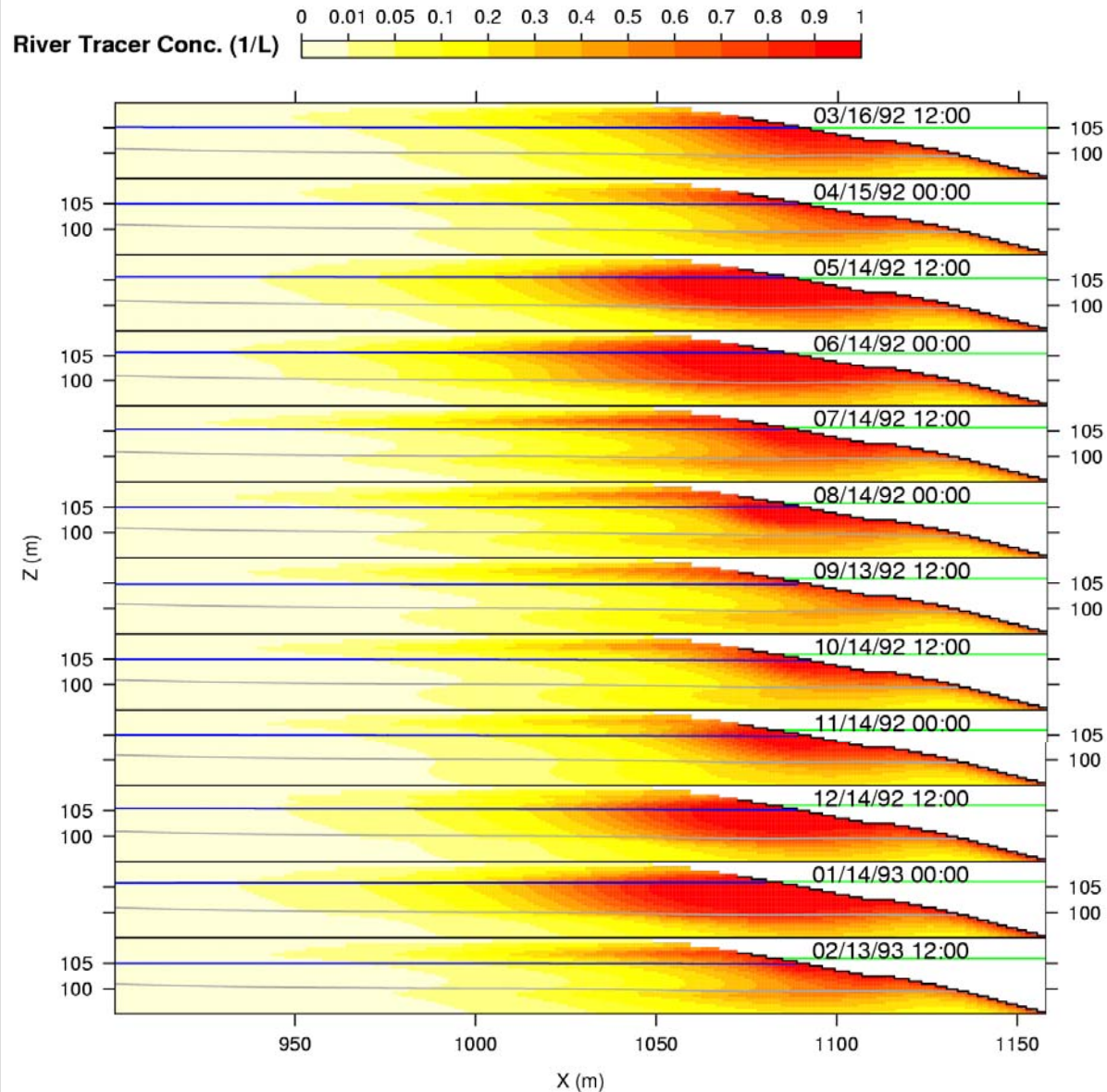
106.8
106.6
106.4
106.2
106.0
105.8
105.6
105.4
105.2
105.0
104.8
104.6
104.4

Aquifer-River Mixing Simulations

- ▶ 4-year spinup period
- ▶ River tracer mixing zone extends ~150 m inland
- ▶ Averaging river stage fluctuations over daily period reduces size of mixing zone
- ▶ Monthly average essentially eliminates mixing with river water

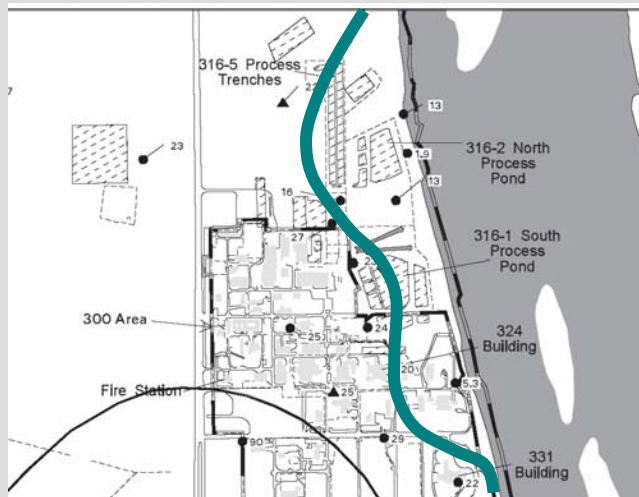


Seasonal Variation in Mixing Zone

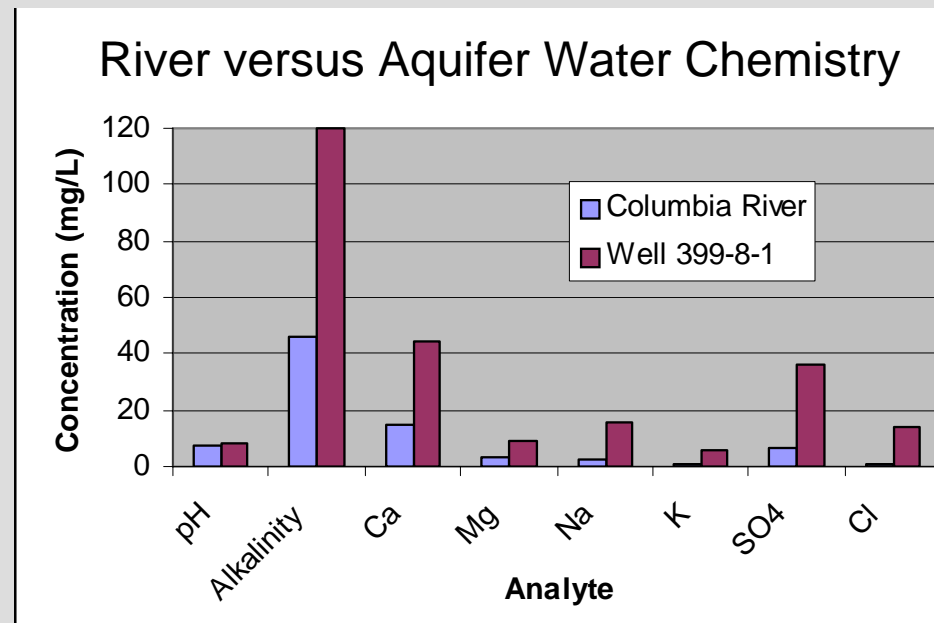


Aquifer-River Solution Chemistry

- ▶ River water influx occurs during high stage
- ▶ Prolonged seasonal high stage period allows mixing in aquifer with river water
- ▶ Significant differences in solution chemistry



2002 Nitrate Concentration

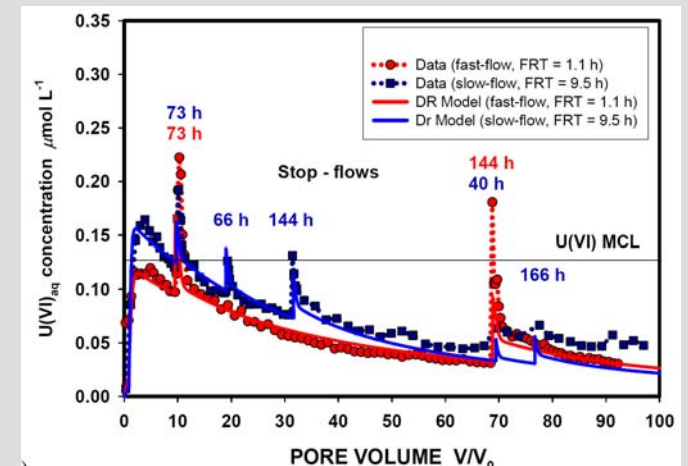
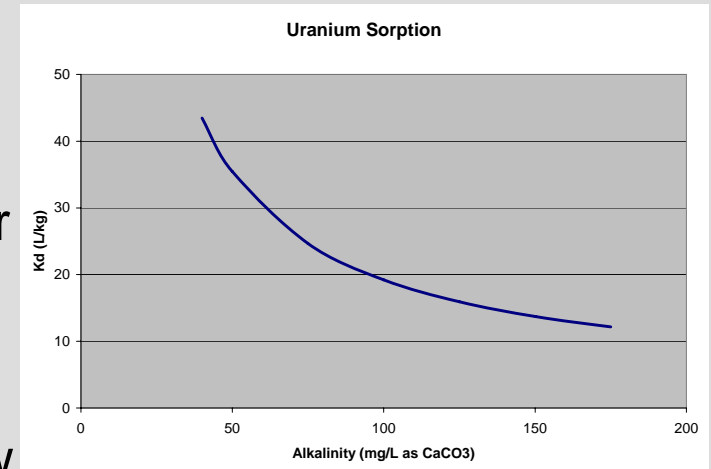


Uranium Geochemistry

- ▶ Constant K_d not consistent with experimental observations
 - Uranium sorption varies strongly with transition between aquifer and river water chemistries (e.g., U, Ca, pH, alkalinity concentrations)
 - Rate-limited uranium mass transfer identified in column experiments with flow rates consistent with field observations

▶ Key Issues

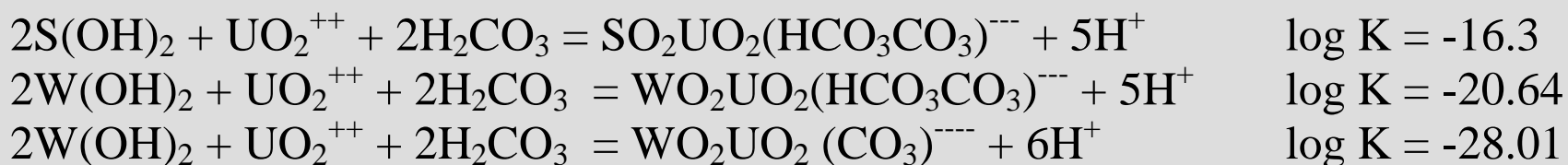
- Uranium leaching from contaminated vadose zone sediments by water table fluctuations
- Changing uranium geochemistry during mixing and exchange of river and groundwater



Uranium Geochemical Process Models

▶ Preliminary three-reaction generalized composite surface complexation model (Jim Davis, USGS)

- accounts for bicarbonate concentration, sediment surface area, and aqueous U(VI) complexation (21 reactions)
- 1 strong site and 2 weak site reactions:

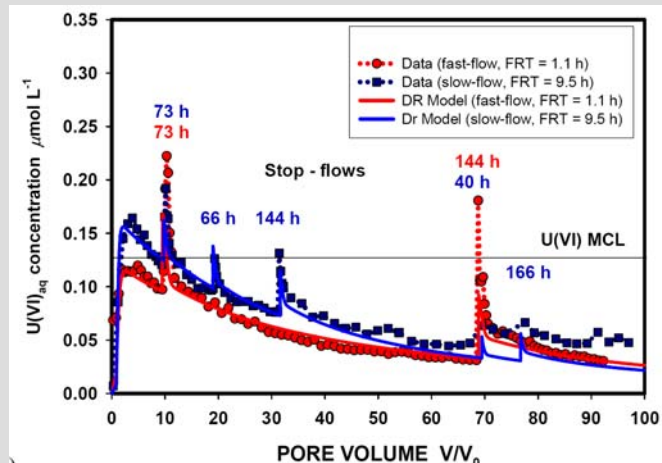


▶ Multisite model with variable uranium mass transfer kinetics (Chongxuan Liu, PNNL):

- Accounts for reaction rates and rate-limited diffusion processes
- Distributed rate parameters were assumed to follow the Gamma statistical distribution (two parameters):

$$\frac{\partial S}{\partial t} = \sum_{i=1}^N \frac{\partial S_i}{\partial t}; \quad \frac{\partial S_i}{\partial t} = \alpha_i [f_i(\alpha_i) K_d^i C - S_i]$$

$$f_i(\alpha_i) = \int_{\alpha_i}^{\alpha_i + \Delta\alpha_i} \frac{\beta^{-\eta} \tau^{\eta-1}}{\Gamma(\eta)} \exp\left(-\frac{\tau}{\beta}\right) d\tau$$



Field-Based Reactive Transport Modeling

Account for full sediment size distribution

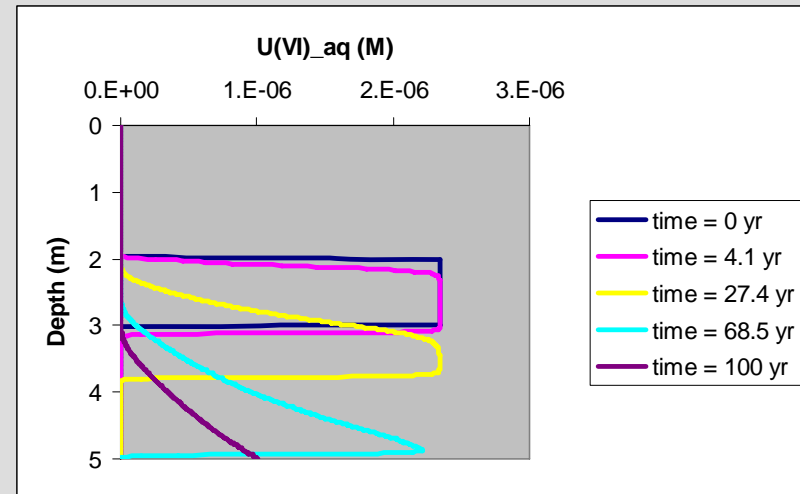
- < 2 mm size fraction in the lab studies
 - Specific surface area: 27.2 m²/g
 - 8% of total sediment
- Preliminary assumption: gravels are unreactive
 - apportion 8% of the 2.06 kg/L field bulk density for surface complexation

Size (mm)	Mass Distribution (%)
<u>Cobbles</u>	
>12.5	74.5
2.0 – 12.5	17.2
<u>Sand</u>	
1.0 – 2.0	2.64
0.5-1.0	2.34
0.25 – 0.5	0.78
0.149 – 0.25	0.33
0.106 – 0.149	0.19
0.053 – 0.106	0.20
<u>Silt + Clay</u>	
<0.053	1.78

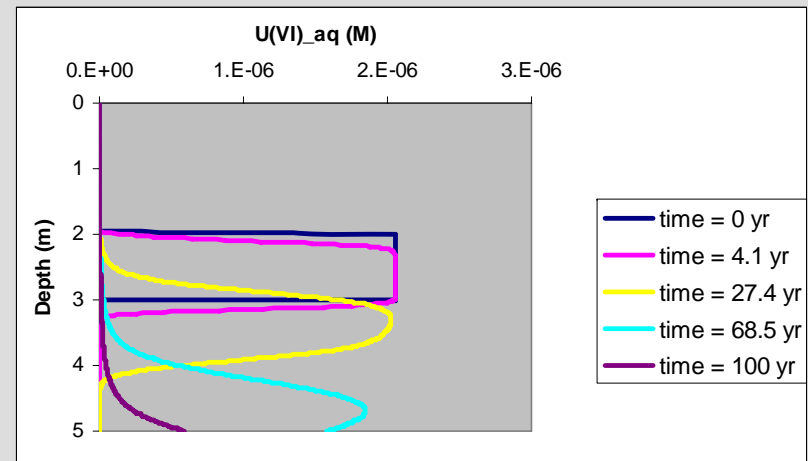
Unsaturated Flow Model Parameters	Value	Units
Horizontal Hydraulic Conductivity	1500	m/d
Vertical Hydraulic Conductivity	150	m/d
Air entry pressure	23.04	cm
Brooks-Corey λ	0.7465	
Residual Saturation	0.1471	
Relative Permeability Method	Burdine	
Porosity	0.25	
Bulk Density	2.06	Kg/L
Recharge Rate	60	mm/yr
Calculated Water Content	0.08	

1-D Unsaturated Reactive Transport Simulation

- ▶ 1-D reactive transport simulation
 - 60 mm/yr recharge results in 0.75 m/yr pore velocity
 - 5 m of vadose zone
 - 1 m of contaminated sediment in the middle
 - 30 nM/g U contaminated zone
- ▶ GC-SCM
 - Sorption front requires over 30 years to move 1 m
 - $K_d = 12.4$ L/kg for this solution chemistry
 - Lowest sediment contamination level results in U(VI) above MCL (0.126 μ M)
- ▶ Multisite kinetic model
 - Very similar to GC-SCM result
 - $K_d = 14$ similar to the GC-SCM
 - impact of kinetics largely minimized by long transport time scales



Generalized Composite SCM



Multisite Kinetic Model

Aquifer-River Interactions

Adapt GC-SCM for the situation where the solution chemistry changes from river water to groundwater

- 1.4 m/d groundwater
- 30 nM/g U-contaminated sediments
- Initial equilibrium with river water
 - 5.76E-8 M aqueous U
 - Intrinsic Kd > 500 L/kg
- After influx of groundwater
 - Aqueous U is 2.50E-6 M
 - Intrinsic Kd = 13.5 L/kg

Solution Chemistry

Components	River water (USGS 6/1/2000)	1988 Well 399-8-3
pH	7.1	7.7
HCO ₃ ⁻	9.18e-4 M	2.66e-3 M
K ⁺	1.75e-5	1.50e-4
NO ₃ ⁻	8.55e-6	1.73e-4
Sr ⁺⁺	1.23e-6	0
Na ⁺	1.00e-4	9.87e-4
Ca ⁺⁺	3.74e-4	1.10e-3
Mg ⁺⁺	1.48e-4	4.10e-4
Cl ⁻	3.10e-5	2.75e-3
SO ₄ ⁻⁻	7.08e-5	3.25e-4

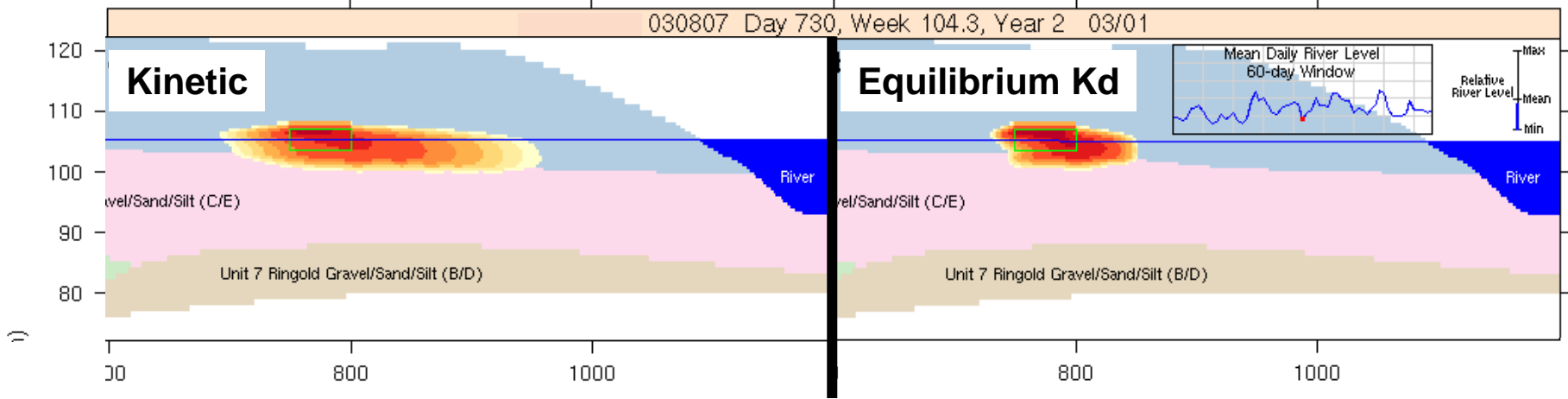
2-D Uranium Reactive Transport

- ▶ 5-year model spinup period for multicomponent reactive transport water chemistry with 0.5 h time steps
- ▶ Hypothetical source zone of U-contaminated sediments
 - 50 m long by 3 m high rectangular zone centered on mean water table elevation (105 m)
 - 700 ug U / L bulk volume
- ▶ Large Grid
 - 26,268 (20,997 active) grid cells
 - source zone 325 m from river
 - model testing: multirate kinetics, equilibrium Kd
- ▶ Small Grid
 - 8,512 (6,112 active) grid cells
 - source zone 75 m from river
 - model testing: multirate kinetics, equilibrium Kd and multicomponent
- ▶ Updated multicomponent model (Bond and Davis)
 - Model based on both North and South Process Pond sediments
 - Single site, 2 reactions

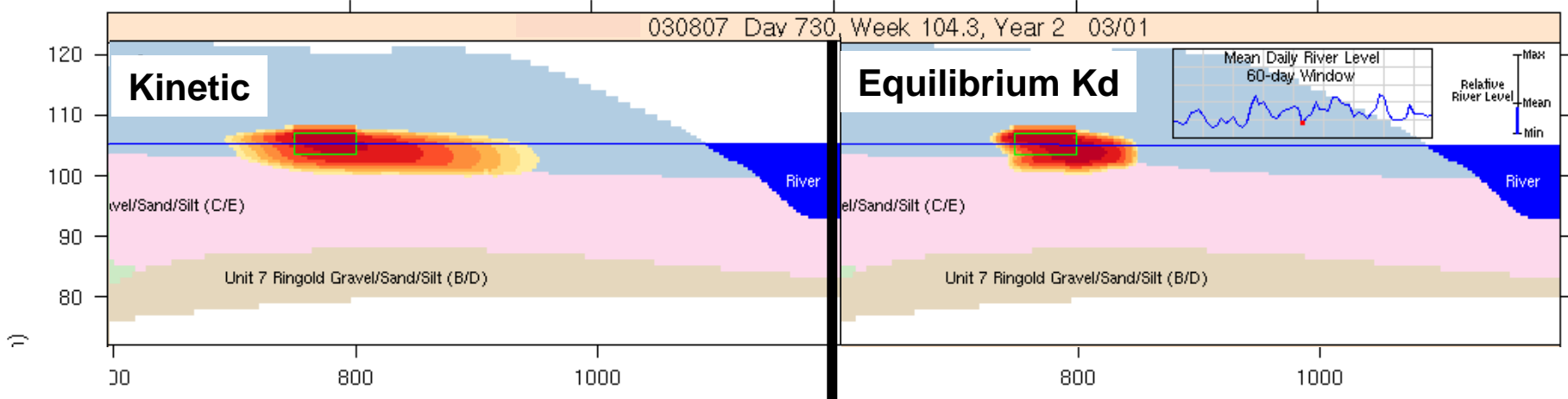
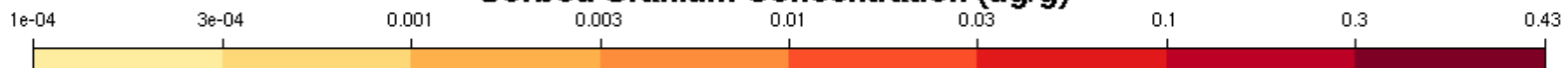
$\text{SOH} + \text{UO}_2^{2+} + \text{H}_2\text{O} = \text{SOUO}_2\text{OH} + 2\text{H}^+$	-4.548
$\text{SOH} + \text{UO}_2^{2+} + 2\text{H}_2\text{CO}_3 = \text{SOUO}_2(\text{HCO}_3)_2^- + 3\text{H}^+$	-4.047

Large Grid Simulations

Aqueous Uranium Concentration (ug/L)

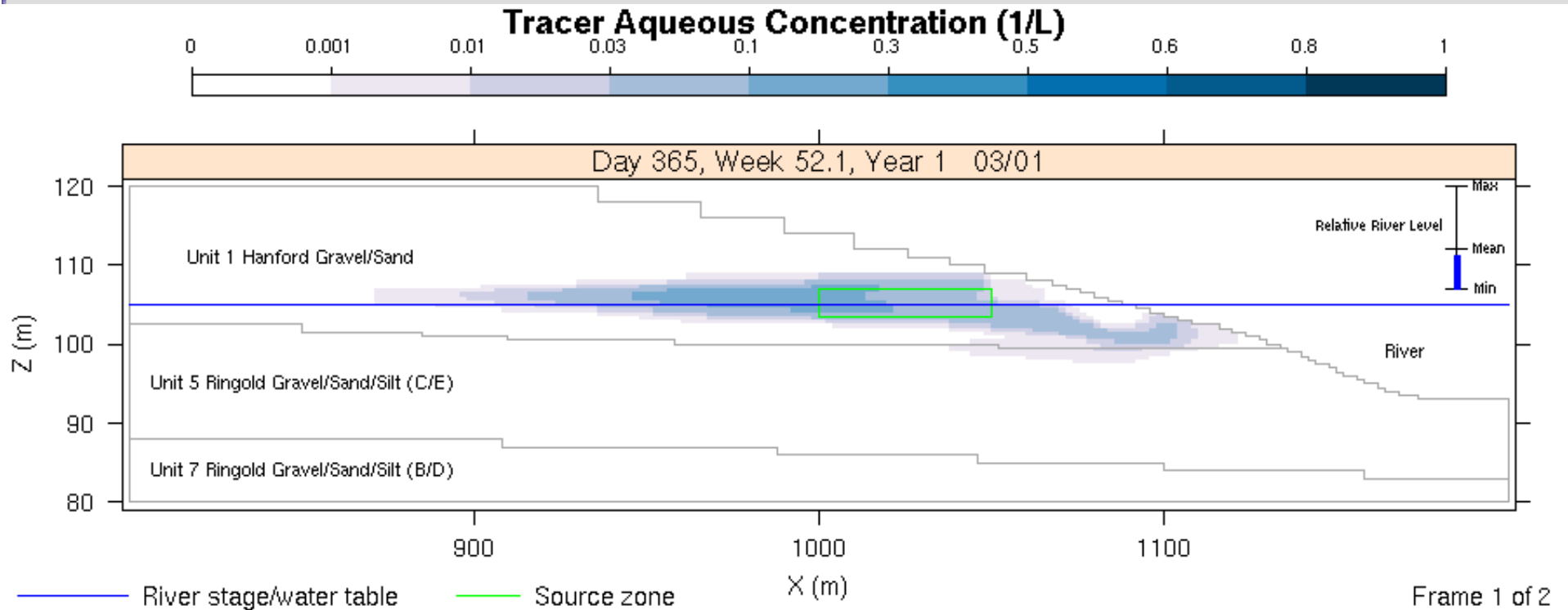


Sorbed Uranium Concentration (ug/g)

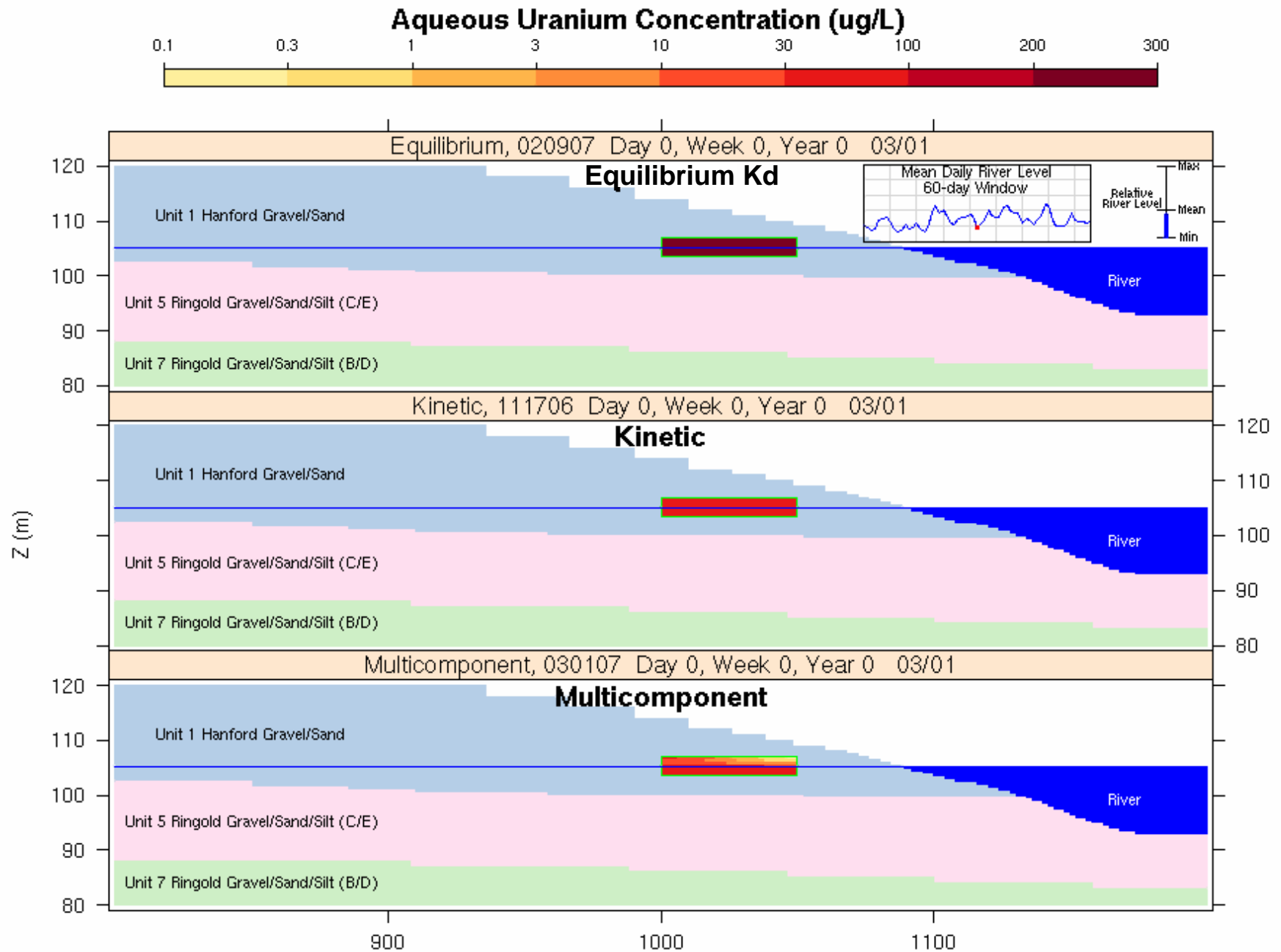


Tracer Transport Near River

- ▶ Net flux above water table is directed inland
- ▶ Net flux below water table is toward river



Small Grid: Near-River Simulations



Modeling Summary

- ▶ Interaction between hourly river stage dynamics, highly transmissive and heterogeneous sediments, and spatially variable uranium create field situation more complex than 1993 conceptual model
 - Lower vadose zone uranium accessed by high river stage
 - Diurnal cycling of high pore velocities
 - Mixing zone of aquifer and river water chemistries
 - dictated by river forcing and hydraulic conductivity
 - sensitive to temporal resolution
 - Equilibrium constant K_d is not consistent with lab results
- ▶ Uranium distribution and mobility can be significantly impacted by
 - rate-limited mass transfer
 - near-river mixing of groundwater and river water
- ▶ Work in progress
 - Ongoing limited field investigation (LFI): sediment cores for detailed analysis, geophysical logging to map uranium distribution
 - Laboratory studies provide framework for understanding uranium mobility
 - Solution chemistry
 - Kinetics
 - Field-scale studies identify large-scale transport context for understanding uranium fate

Model Description

- ▶ Variably-saturated flow with hourly river-stage boundary condition
- ▶ Dynamic seepage face boundary condition for bank storage effects
- ▶ Recharge 60 mm/yr at surface, 0.9 mm/yr upward leakage from basalt
- ▶ Variable grid spacing
 - 2.0 to 100 m in the horizontal
 - 0.5 to 2.0 m in the vertical
 - large grid: 26,268 (20,997 active) grid cells
 - small grid: 8,512 (6,112 active) grid cells
- ▶ 30 minute time steps

Uranium Reactions

Reaction	log K (I = 0)
$\text{UO}_2^{2+} + \text{H}_2\text{O} = \text{UO}_2\text{OH}^+ + \text{H}^+$	-5.25
$\text{UO}_2^{2+} + 2\text{H}_2\text{O} = \text{UO}_2(\text{OH})_{2,\text{aq}} + 2\text{H}^+$	-12.15
$\text{UO}_2^{2+} + 3\text{H}_2\text{O} = \text{UO}_2(\text{OH})_3^- + 3\text{H}^+$	-20.25
$\text{UO}_2^{2+} + 4\text{H}_2\text{O} = \text{UO}_2(\text{OH})_4^{2-} + 4\text{H}^+$	-32.4
$2\text{UO}_2^{2+} + \text{H}_2\text{O} = (\text{UO}_2)_2\text{OH}^{3+} + \text{H}^+$	-2.70
$2\text{UO}_2^{2+} + 2\text{H}_2\text{O} = (\text{UO}_2)_2(\text{OH})_2^{2+} + 2\text{H}^+$	-5.62
$3\text{UO}_2^{2+} + 4\text{H}_2\text{O} = (\text{UO}_2)_3(\text{OH})_4^{2+} + 4\text{H}^+$	-11.90
$3\text{UO}_2^{2+} + 5\text{H}_2\text{O} = (\text{UO}_2)_3(\text{OH})_5^+ + 5\text{H}^+$	-15.55
$3\text{UO}_2^{2+} + 7\text{H}_2\text{O} = (\text{UO}_2)_3(\text{OH})_7^- + 7\text{H}^+$	-32.20
$4\text{UO}_2^{2+} + 7\text{H}_2\text{O} = (\text{UO}_2)_4(\text{OH})_7^+ + 7\text{H}^+$	-21.9
$\text{UO}_2^{2+} + \text{CO}_3^{2-} = \text{UO}_2\text{CO}_3(\text{aq})$	9.94
$\text{UO}_2^{2+} + 2\text{CO}_3^{2-} = \text{UO}_2(\text{CO}_3)_2^{2-}$	16.61
$\text{UO}_2^{2+} + 3\text{CO}_3^{2-} = \text{UO}_2(\text{CO}_3)_3^{4-}$	21.84
$2\text{UO}_2^{2+} + \text{CO}_3^{2-} + 3\text{H}_2\text{O} = (\text{UO}_2)_2\text{CO}_3(\text{OH})_3^- + 3\text{H}^+$	-0.855
$\text{Ca}^{2+} + \text{UO}_2^{2+} + 3\text{CO}_3^{2-} = \text{CaUO}_2(\text{CO}_3)_3^{2-}$	25.4
$2\text{Ca}^{2+} + \text{UO}_2^{2+} + 3\text{CO}_3^{2-} = \text{Ca}_2\text{UO}_2(\text{CO}_3)_3(\text{aq})$	30.55
$\text{UO}_2^{2+} + \text{NO}_3^- = \text{UO}_2\text{NO}_3^+$	0.3
$\text{UO}_2^{2+} + \text{Cl}^- = \text{UO}_2\text{Cl}^+$	0.17
$\text{UO}_2^{2+} + 2\text{Cl}^- = \text{UO}_2\text{Cl}_2(\text{aq})$	-1.1
$\text{UO}_2^{2+} + \text{SO}_4^{2-} = \text{UO}_2\text{SO}_4(\text{aq})$	3.15
$\text{UO}_2^{2+} + 2\text{SO}_4^{2-} = \text{UO}_2(\text{SO}_4)_2^{2-}$	4.14
$\text{SOH} + \text{UO}_2^{2+} + \text{H}_2\text{O} = \text{SOUO}_2\text{OH} + 2\text{H}^+$	-4.548
$\text{SOH} + \text{UO}_2^{2+} + 2\text{H}_2\text{CO}_3 = \text{SOUO}_2(\text{HCO}_3)_2^- + 3\text{H}^+$	-4.047

Other Reactions

Reaction	Log K
$\text{H}^+ + \text{HCO}_3^- = \text{H}_2\text{CO}_3$	6.3414
$\text{HCO}_3^- = \text{CO}_3^{2-} + \text{H}^+$	-10.3249
$\text{Ca}^{2+} + \text{HCO}_3^- = \text{CaCO}_3(\text{aq}) + \text{H}^+$	-7.0088
$\text{Ca}^{2+} + \text{Cl}^- = \text{CaCl}^+$	-0.7004
$\text{Ca}^{2+} + 2\text{Cl}^- = \text{CaCl}_2(\text{aq})$	-0.6535
$\text{Ca}^{2+} + \text{HCO}_3^- = \text{CaHCO}_3^+$	1.0420
$\text{Ca}^{2+} + \text{NO}_3^- = \text{CaNO}_3^+$	1.3
$\text{Mg}^{2+} + \text{NO}_3^- = \text{MgNO}_3^+$	1.3
$\text{Ca}^{2+} + \text{H}_2\text{O} = \text{CaOH}^+ + \text{H}^+$	-12.85
$\text{Ca}^{2+} + \text{SO}_4^{2-} = \text{CaSO}_4(\text{aq})$	2.1004
$\text{H}^+ + \text{Cl}^- = \text{HCl}(\text{aq})$	0.6999
$\text{H}^+ + \text{NO}_3^- = \text{HNO}_3(\text{aq})$	-1.3081
$\text{K}^+ + \text{Cl}^- = \text{KCl}(\text{aq})$	-1.5004
$\text{K}^+ + \text{SO}_4^{2-} = \text{KSO}_4^-$	0.875
$\text{Mg}^{2+} + \text{HCO}_3^- = \text{MgCO}_3(\text{aq}) + \text{H}^+$	-7.3562
$\text{Mg}^{2+} + \text{Cl}^- = \text{MgCl}^+$	-0.1386
$\text{Mg}^{2+} + \text{HCO}_3^- = \text{MgHCO}_3^+$	1.0329
$\text{Mg}^{2+} + \text{SO}_4^{2-} = \text{MgSO}_4(\text{aq})$	2.4125
$\text{Na}^+ + \text{HCO}_3^- = \text{NaCO}_3^- + \text{H}^+$	-9.8156
$\text{Na}^+ + \text{Cl}^- = \text{NaCl}$	-0.7821
$\text{Na}^+ + \text{HCO}_3^- = \text{NaHCO}_3(\text{aq})$	0.1557
$\text{Na}^+ + \text{H}_2\text{O} = \text{NaOH}(\text{aq}) + \text{H}^+$	-14.7986
$\text{Na}^+ + \text{SO}_4^{2-} = \text{NaSO}_4^-$	0.82
$\text{H}^+ + \text{OH}^- = \text{H}_2\text{O}$	13.9911
$\text{Sr}^{2+} + \text{HCO}_3^- = \text{SrCO}_3(\text{aq}) + \text{H}^+$	-7.4703
$\text{Sr}^{2+} + \text{Cl}^- = \text{SrCl}^+$	-0.2533
$\text{Sr}^{2+} + \text{NO}_3^- = \text{SrNO}_3^+$	0.8
$\text{Sr}^{2+} + \text{H}_2\text{O} = \text{SrOH}^+ + \text{H}^+$	-13.29
$\text{Sr}^{2+} + \text{SO}_4^{2-} = \text{SrSO}_4(\text{aq})$	2.3
$\text{Na}^+ + \text{NO}_3^- = \text{NaNO}_3(\text{aq})$	-0.2564
$\text{Ca}^{2+} + \text{HCO}_3^- = \text{Calcite}(\text{s}) + \text{H}^+$	-1.8542