# **Science Needs in the Arctic**

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Our challenge remains to grasp these subtle interactions and feedback processes, to integrate our understanding and to build a comprehensive synthesis of the changing Arctic. How does permafrost thaw and degradation, and the associated changes in landscape evolution, hydrology, soil biogeochemistry and plant community dynamics, affect feedbacks to the climate system?

Can we quantify and predict landscape response to climate change?

How does permafrost degradation impact hydrological states, stocks, fluxes and pathways?

How do liquid water films and freeze-thaw processes influence microbial activity and GHG production?

How do soil carbon structure and interactions affect degradation of soil organic matter?

What are the distinguishing chemical signatures of organic matter that is degraded and mobilized in Arctic soils? How do soil C-mineral associations alter organic matter preservation, transport and degradation rates?

How do changes in microbial community structure and function affect degradation?

How does subsurface variability affect biodegradation mechanisms and rates?

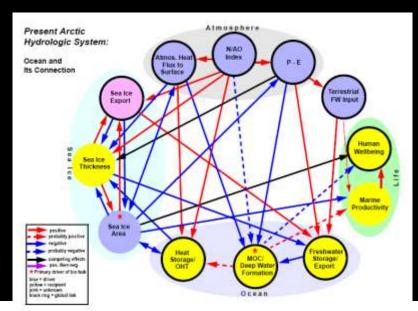
What is the residence time of organic material in permafrost soils?

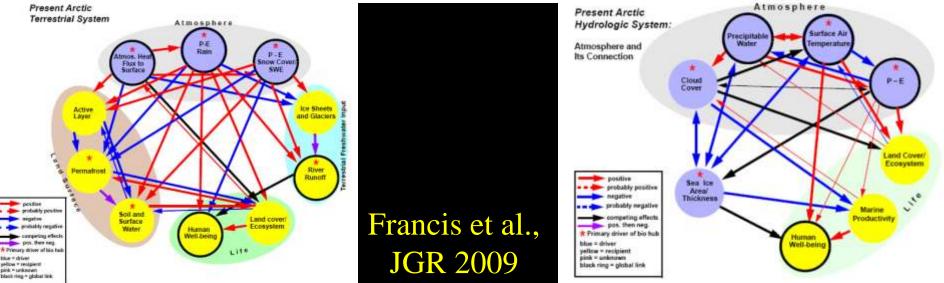
How does permafrost degradation alter energy balance at different scales?

Will shrub establishment limit the rate of albedo change?

**Arctic climate-ecosystem feedbacks: How big, how, and how fast?** 

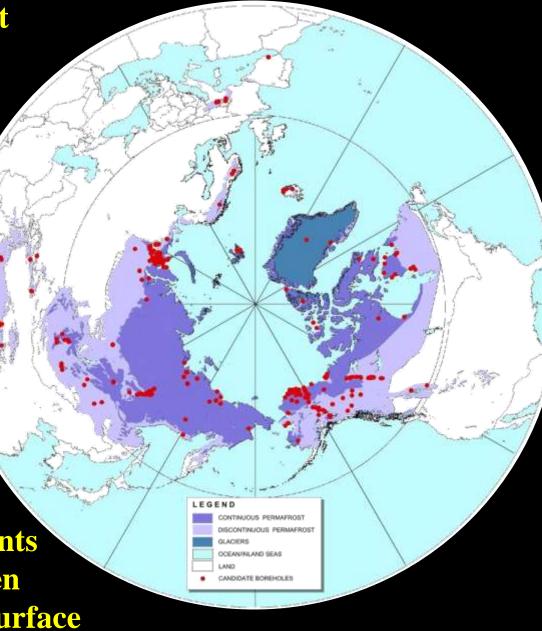
### **Resolving and quantifying the system interconnections is the first step to prediction**

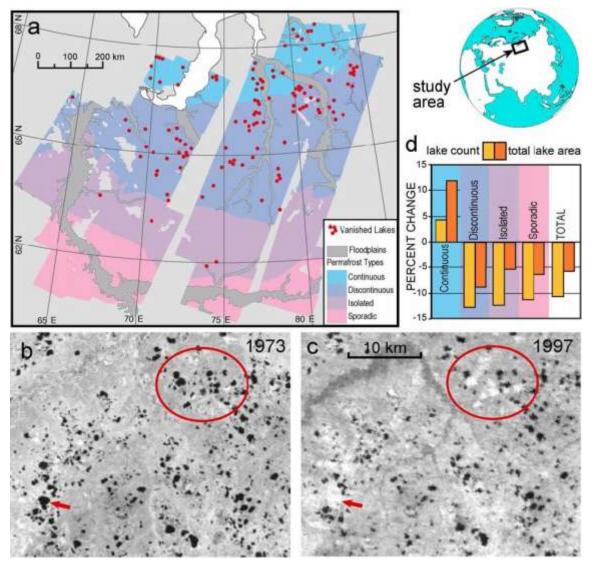




**Permafrost covers about** 25% of the terrestrial area. Degradation of permafrost in response to a warming climate will have dramatic impacts to local infrastructure, hydrology, and ecology.

Formation of a talik (unfrozen layer above permafrost) represents one of the threshold events that will initiate a sudden change in ecology and surface energy balance.





(a) Locations of Siberian lake inventories, permafrost distribution, and vanished lakes. Total lake abundance and inundation area have declined since 1973 (b), including permanent drainage and re-vegetation of former lakebeds (c). Interestingly, net *increases* in lake abundance and area have occurred in continuous permafrost (d), suggesting an initial but transitory increase in surface ponding (Larry Smith et al., 2005).

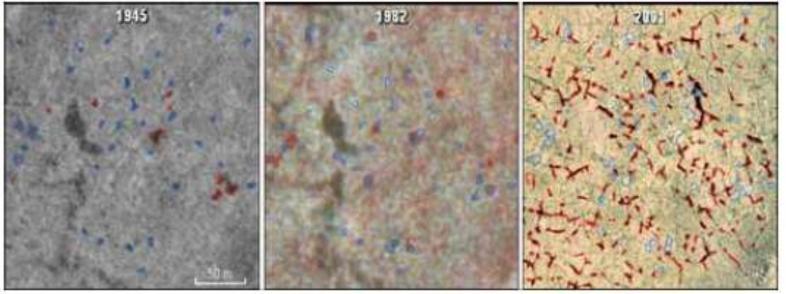


In colder permafrost, located in more northern regions of Canada, Alaska and Siberia, evidence of climate impacts to ecosystem is also apparent.

1945



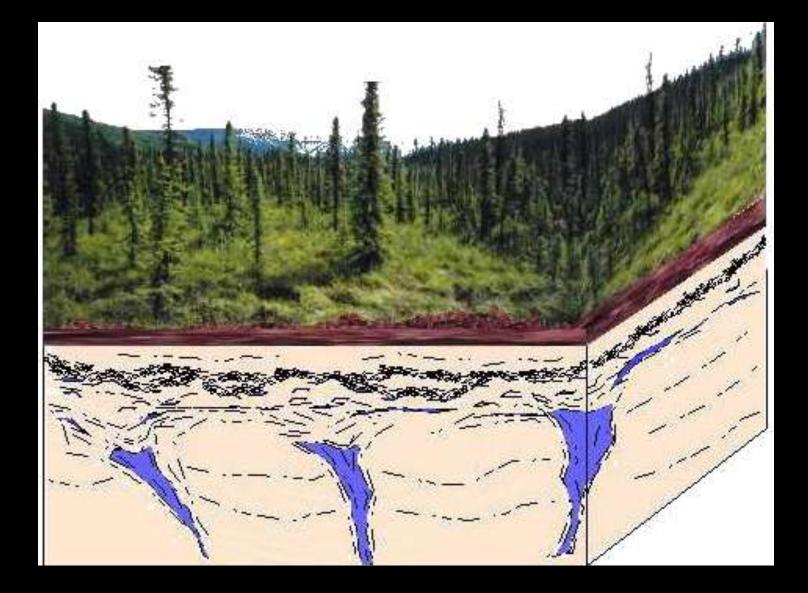
#### 2001

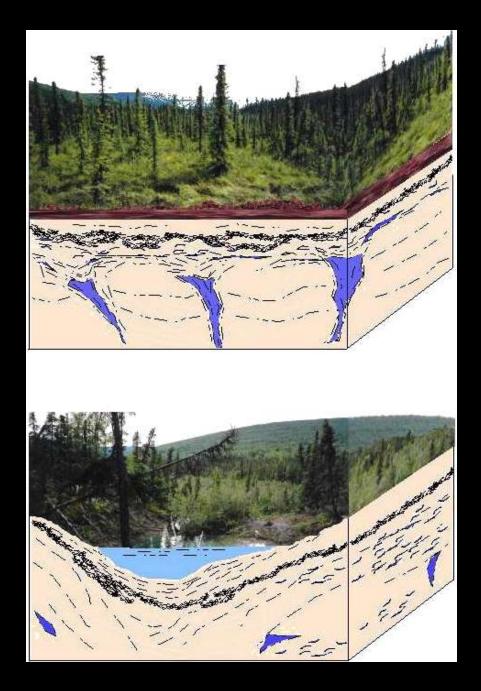


Abundance of degrading (advanced stage only in red) and stabilizing (initial and advanced stages in blue) thermokarst pits within the 0.6 km<sup>2</sup> intensive study area.

Jorgenson et al., 2006. GRL









#### 1949



**Chandler River, 50 miles S. of Umiat: Sturm, Racine and Tape: Fifty Years of Change in Arctic Alaskan Shrub Abundance** 

#### 2001



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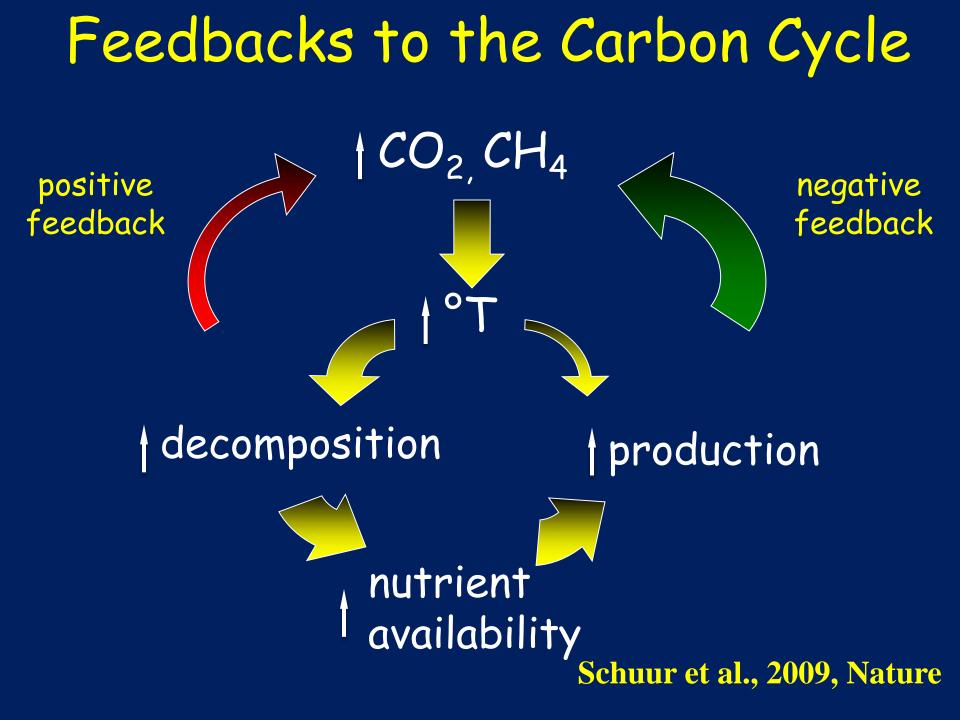
# Global Carbon PoolsGlobal Vegetation C6Global Soil C (1m)1Atmosphere7

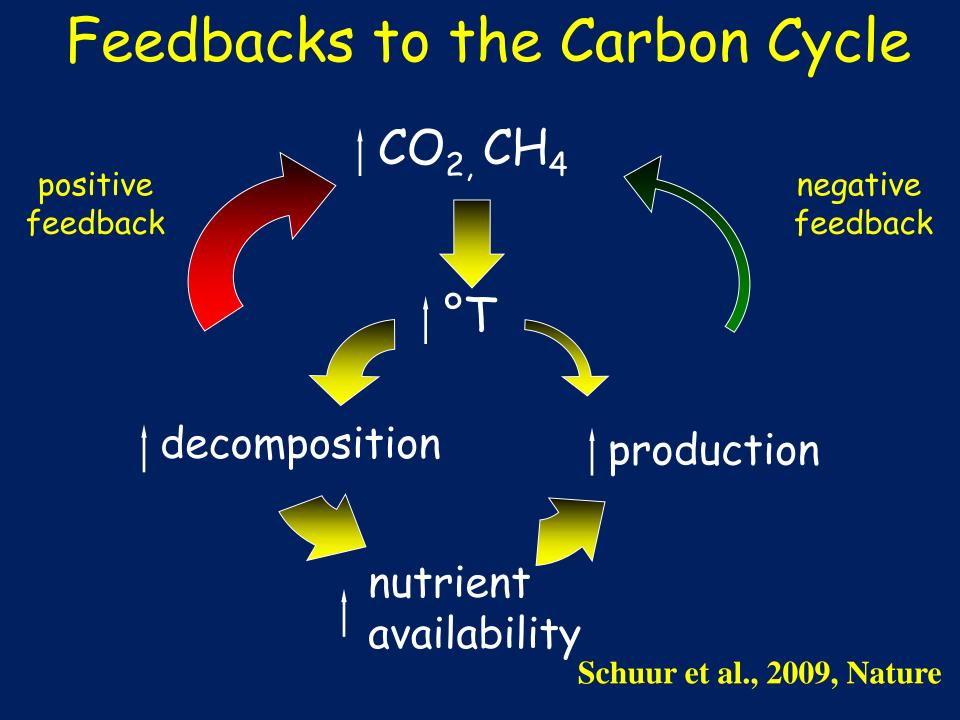
650 Pg 1500 Pg 777+ Pg

Permafrost Zone Soil C Peatlands (several m) Mineral Soil (3m) Siberian Deep C (~25m) Alluvial Deep C (~25m)

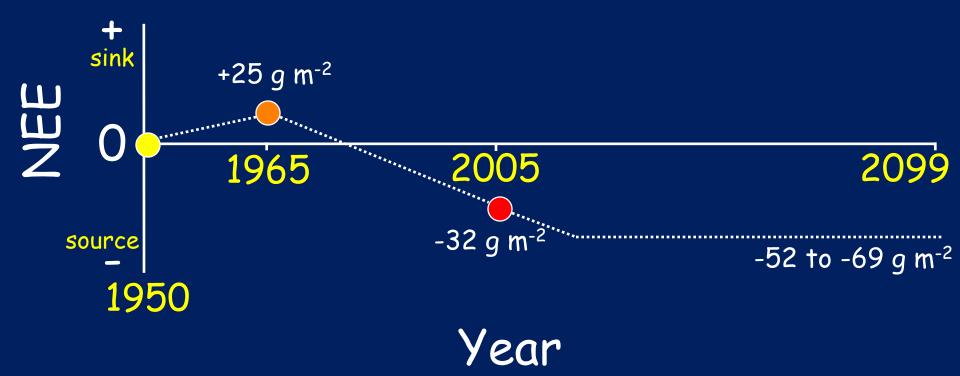
277 Pg 747 Pg 407 Pg 241 Pg 1672 Pg

Jobaggy et al. 2000, Field et al. 2007, Zimov et al. 2006, Tarnocai, in press, Schuur et al. 2008]





# Net Carbon Exchange Projections



Carbon loss by 2099: 4.4-6.0 kg/m<sup>-2</sup> (9.4-12.9%)

Equilibrium C loss much higher over centuries

[Schuur et al. , 2009]

