

PUSHING CUMULUS PARAMETERIZATION TO HIGHER RESOLUTION

WHAT ARE THE LIMITS?

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AND
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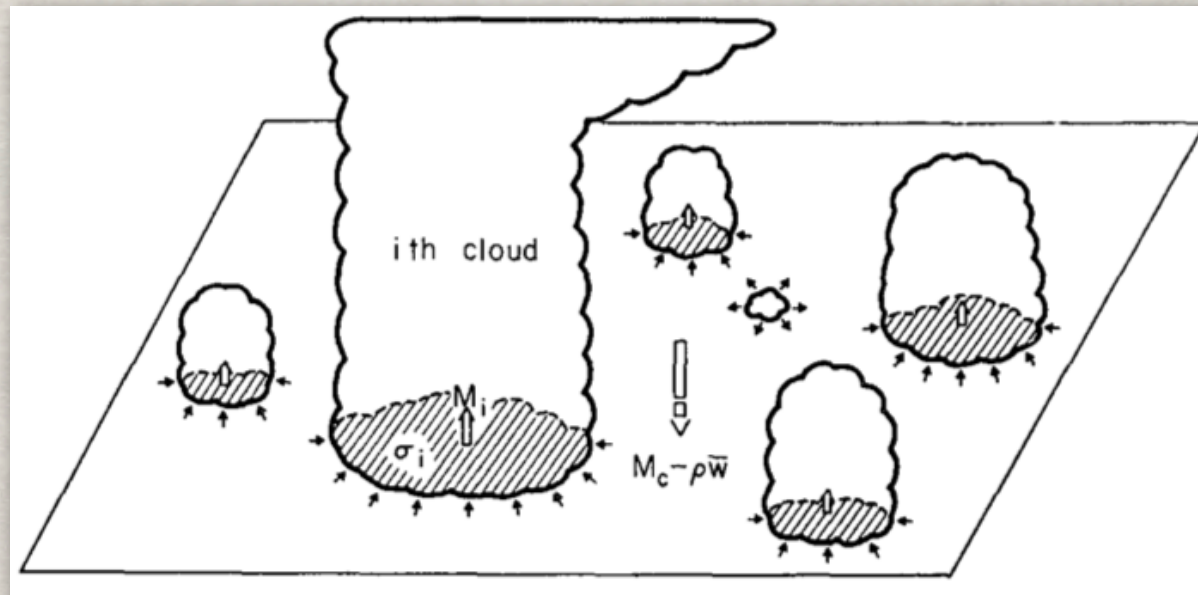
U.S. DEPARTMENT OF
ENERGY

Colorado
State
University

ASR FALL MEETING
BOULDER, CO
10/15/2010

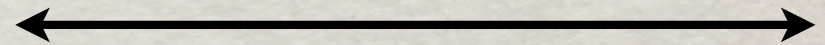
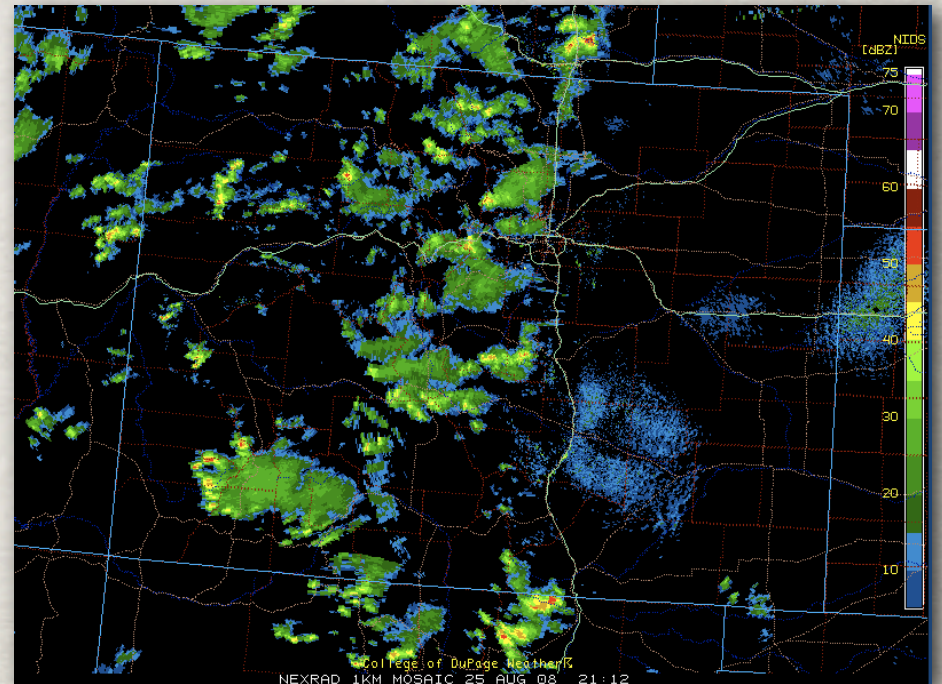
CUMULUS PARAMETERIZATION

- ✱ Describes interaction of a cumulus cloud ensemble with the large-scale
 - ✱ Often looking for the “expected values” of influence from an ensemble of clouds in quasi-equilibrium with the large-scale conditions.
 - ✱ Requires spatial and temporal scale separation.



HOW MANY THUNDERSTORMS FIT?

- ☼ With a grid spacing of 20 km or less, we definitely do not have a statistically meaningful sample of large clouds in each grid column.
- ☼ Even with a grid spacing of 200 km, the number of large clouds in a column is worryingly small.



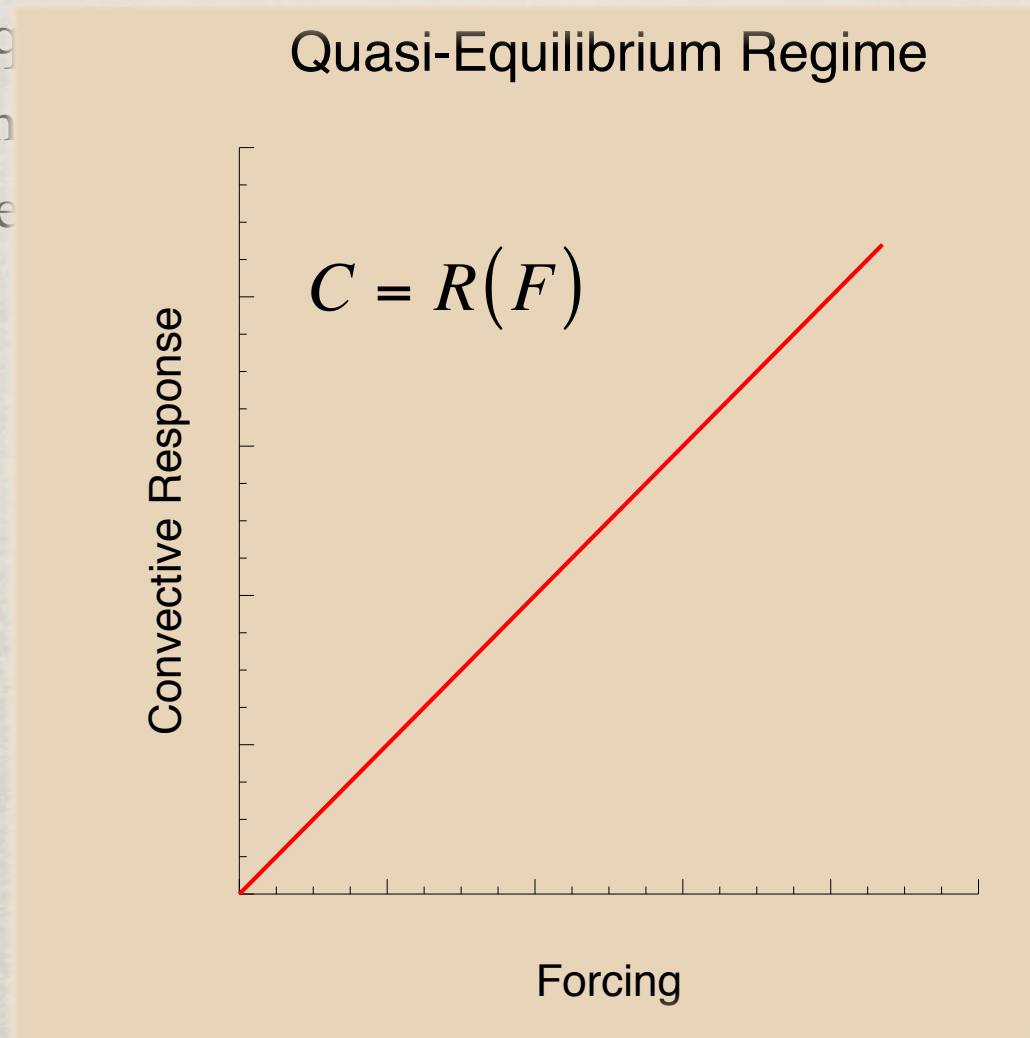
~700 km

HOW QE BREAKS DOWN

- ✱ **Quasi-equilibrium** regime
 - ✱ Forcing determines response
 - ✱ No memory
- ✱ **Non-equilibrium** regime, still deterministic
 - ✱ Out of phase, forcings change too rapidly
 - ✱ Past history of the convective statistics matters
- ✱ **Non-deterministic** regime
 - ✱ As computational areas decrease in size, the stochastic component of convection becomes significant.
 - ✱ Creates uncertainty due to small sample size -- noise from below

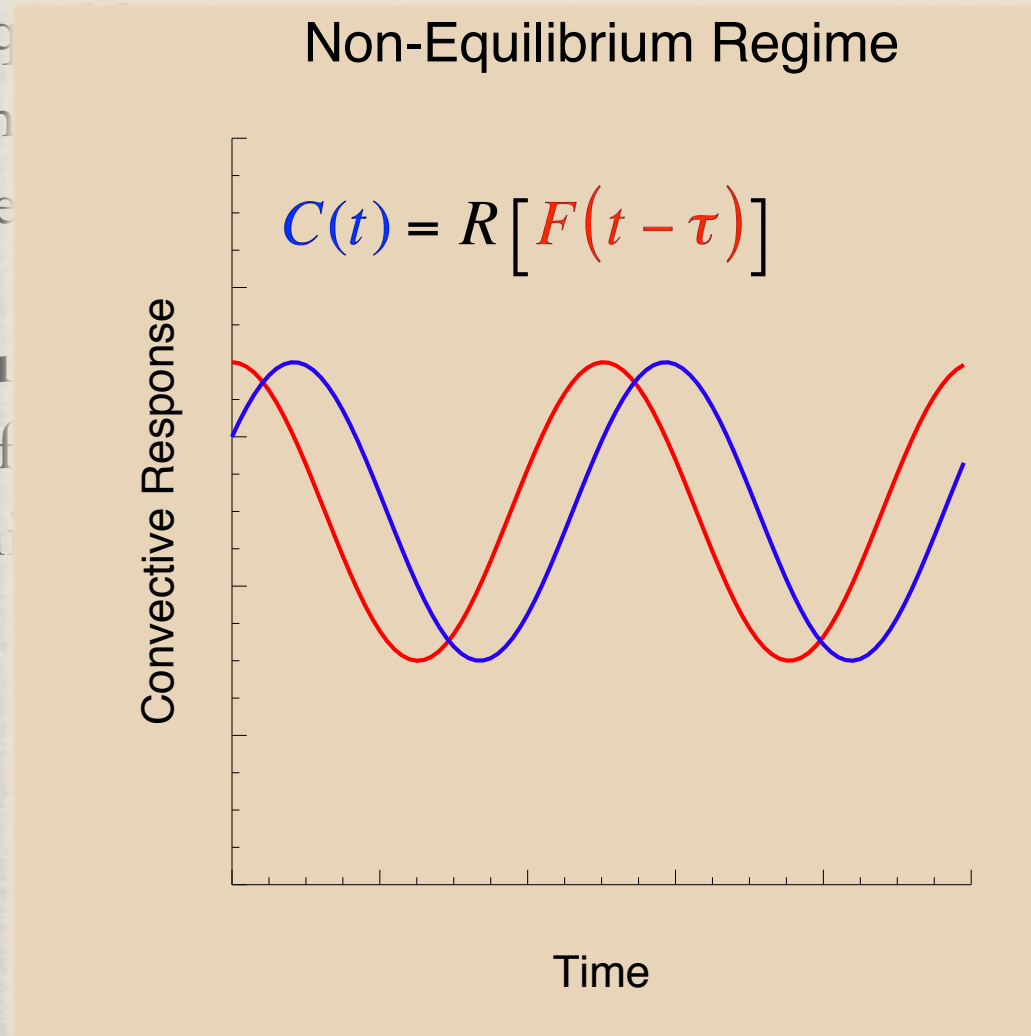
HOW QE BREAKS DOWN

- ✱ Quasi-equilibrium
- ✱ Forcing
- ✱ No memory



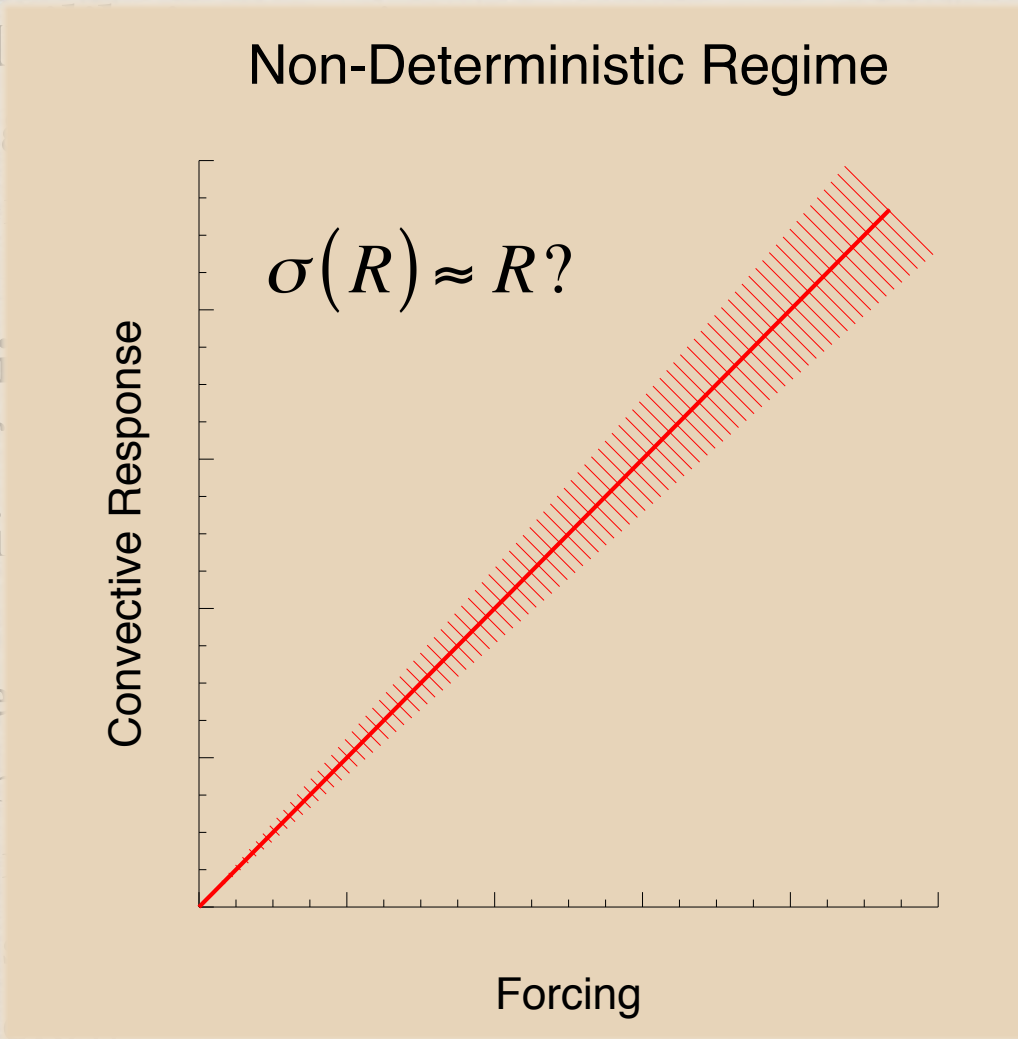
HOW QE BREAKS DOWN

- ✱ Quasi-equilibrium
- ✱ Forcing
- ✱ No memory
- ✱ Non-equilibrium
- ✱ Out of equilibrium
- ✱ Past history



HOW QE BREAKS DOWN

- ✱ Quasi-equilibrium
 - ✱ Forcing
 - ✱ No memory
- ✱ Non-equilibrium
 - ✱ Out of equilibrium
 - ✱ Past history
- ✱ Non-deterministic
 - ✱ As component
 - ✱ Component
 - ✱ Create from boundary

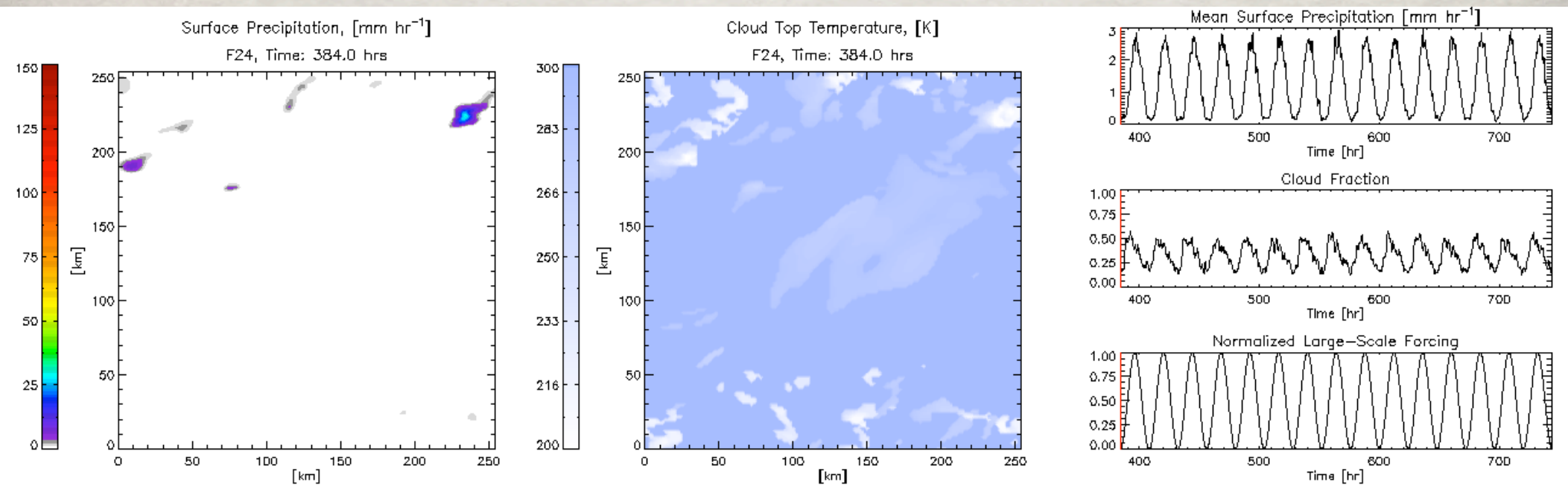


rs

stochastic
t.

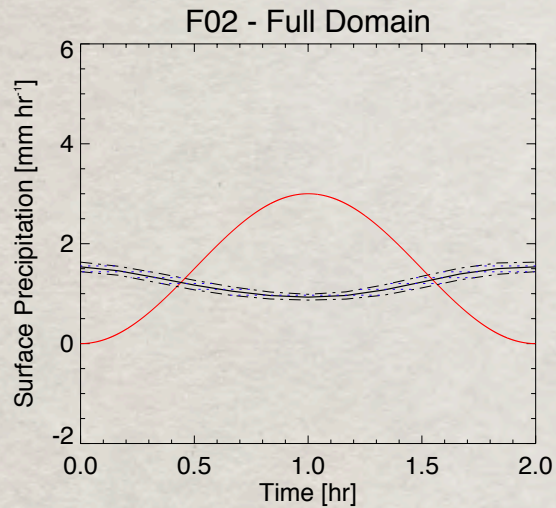
-- noise

3D NUMERICAL SIMULATION

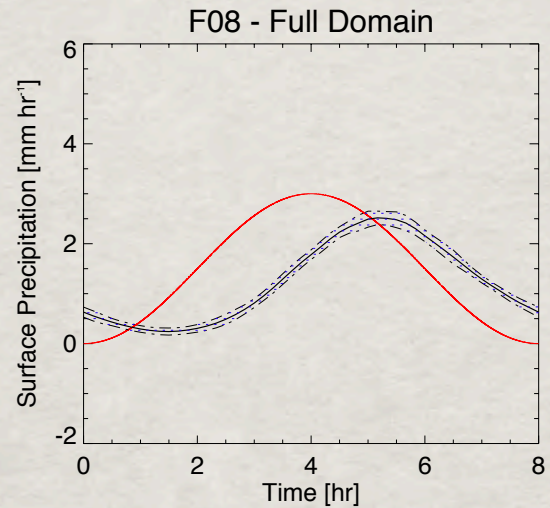


Following Xu et al. 1992

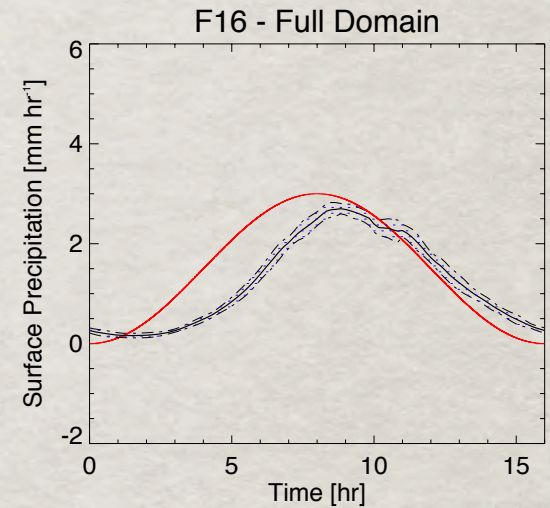
RAIN RATE



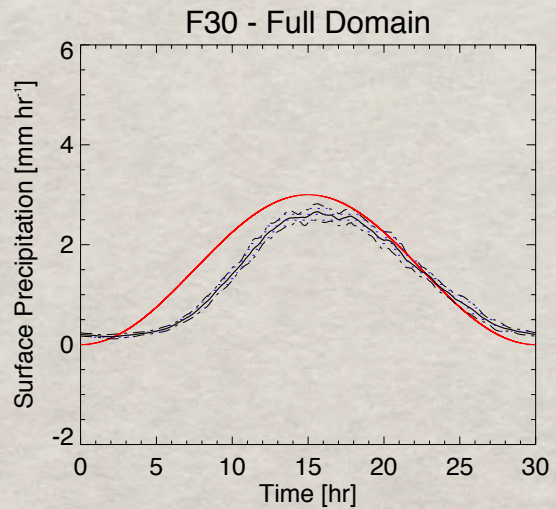
Forcing leads Precip by:
60.0 minutes (50.00 % of the forcing period)



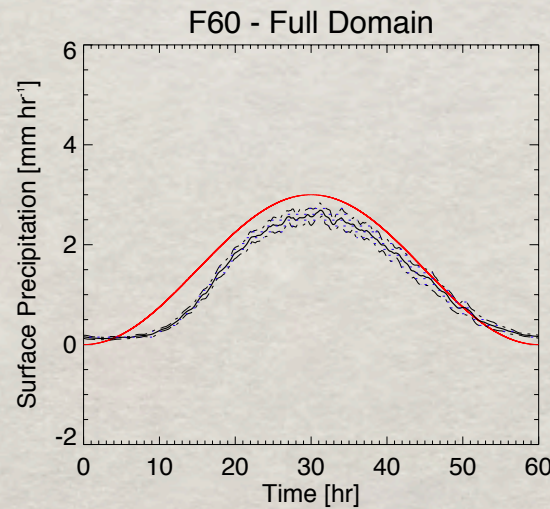
Forcing leads Precip by:
80.0 minutes (16.67 % of the forcing period)



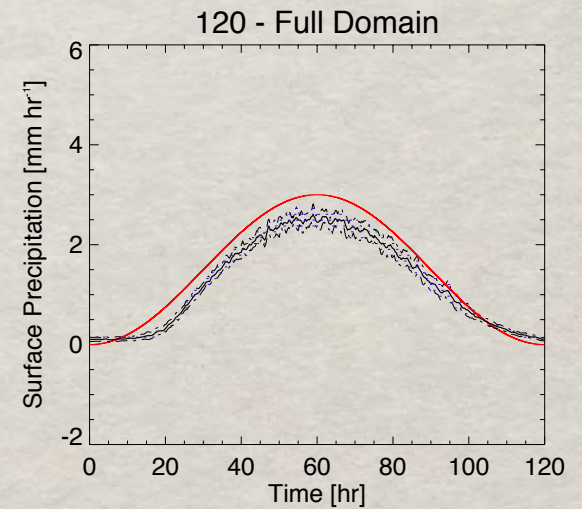
Forcing leads Precip by:
80.0 minutes (8.33 % of the forcing period)



Forcing leads Precip by:
80.0 minutes (4.44 % of the forcing period)

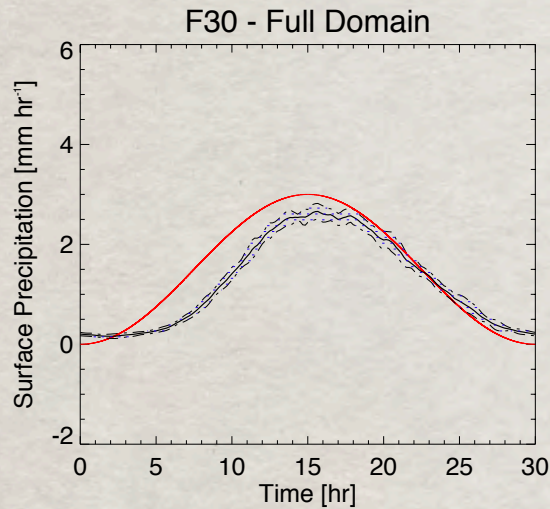


Forcing leads Precip by:
80.0 minutes (2.22 % of the forcing period)

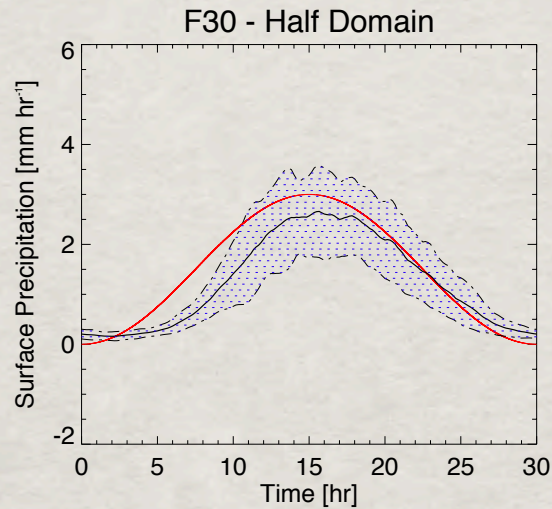


Forcing leads Precip by:
70.0 minutes (0.97 % of the forcing period)

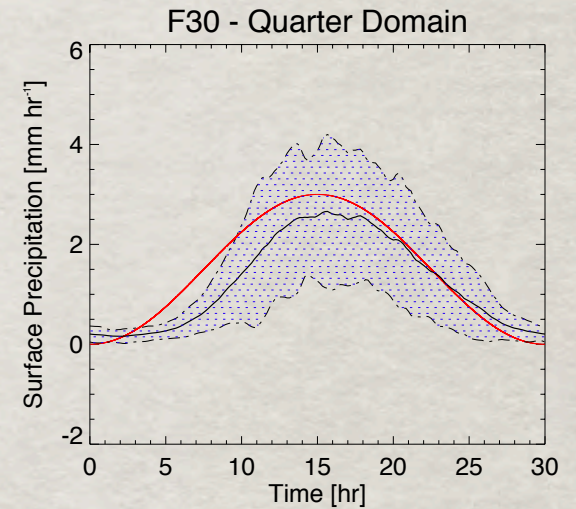
RAIN RATE



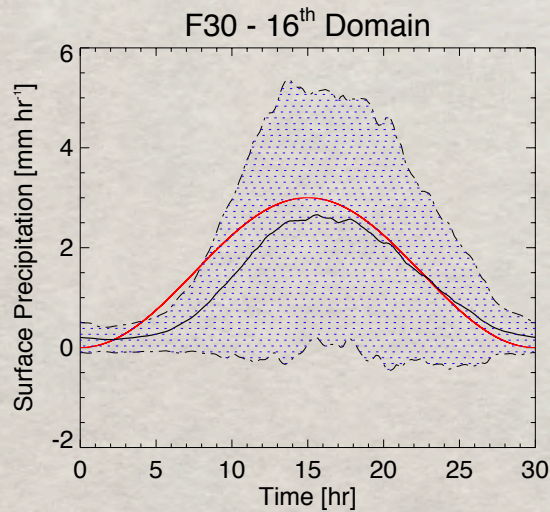
Forcing leads Precip by:
80.0 minutes (4.44 % of the forcing period)



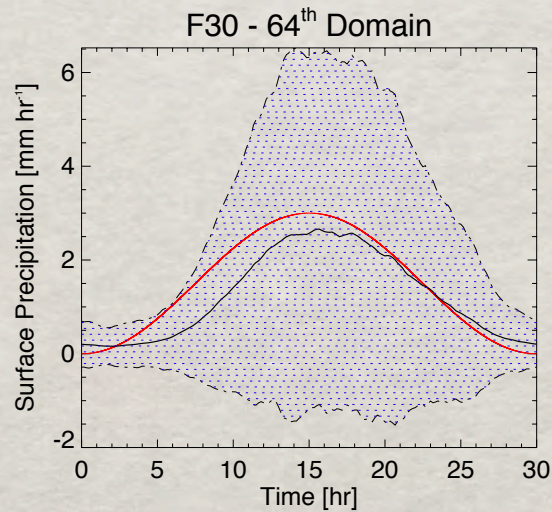
Forcing leads Precip by:
85.0 minutes (4.72 % of the forcing period)



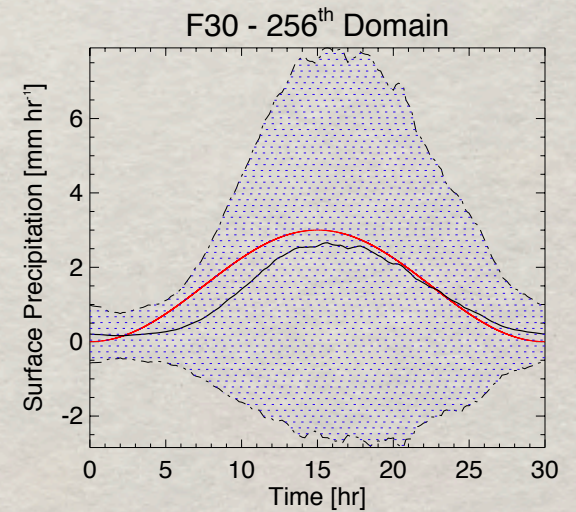
Forcing leads Precip by:
82.5 minutes (4.58 % of the forcing period)



Forcing leads Precip by:
86.9 minutes (4.83 % of the forcing period)

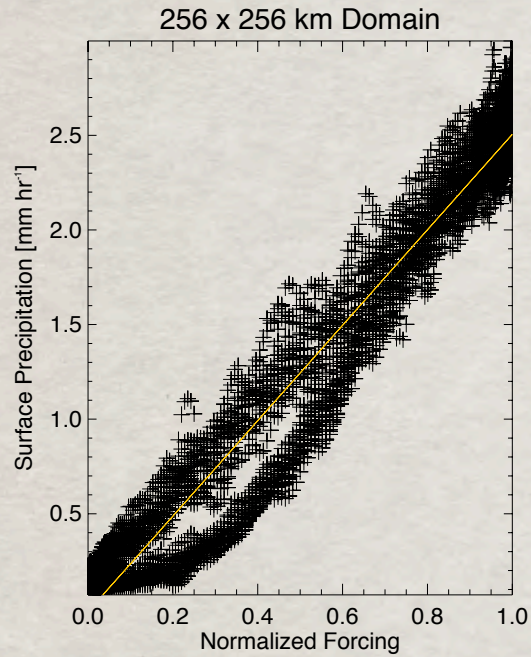


Forcing leads Precip by:
113.8 minutes (6.32 % of the forcing period)

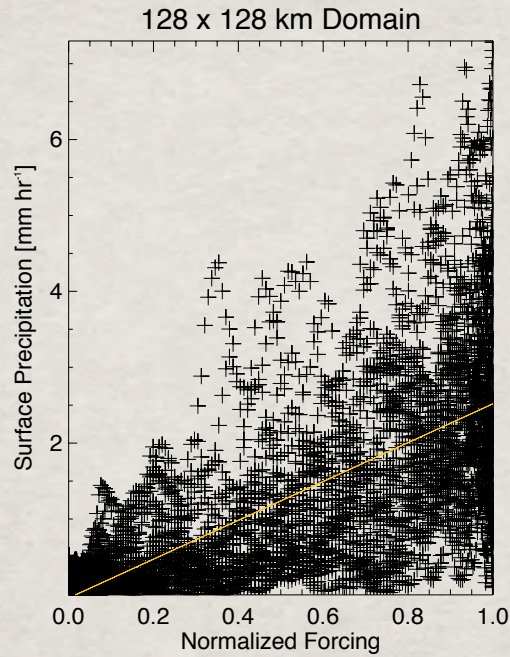


Forcing leads Precip by:
149.0 minutes (8.28 % of the forcing period)

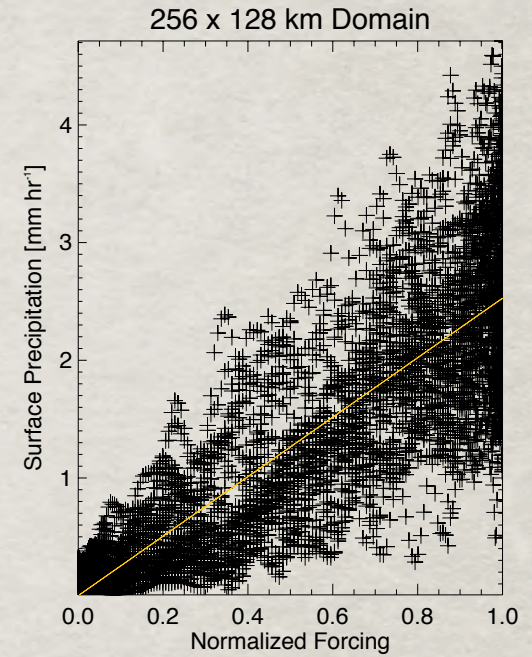
F60



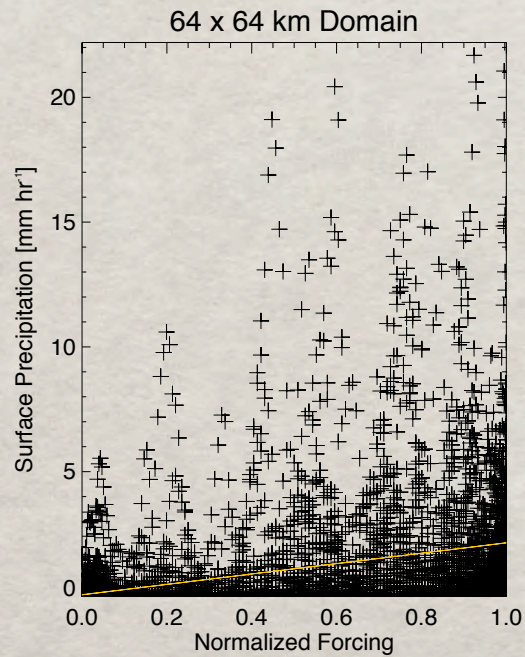
$r^2 = 0.948$
Slope = 2.522



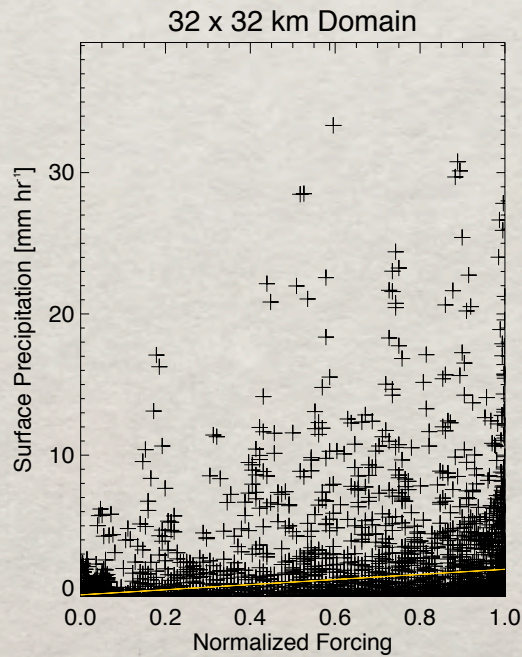
$r^2 = 0.473$
Slope = 2.558



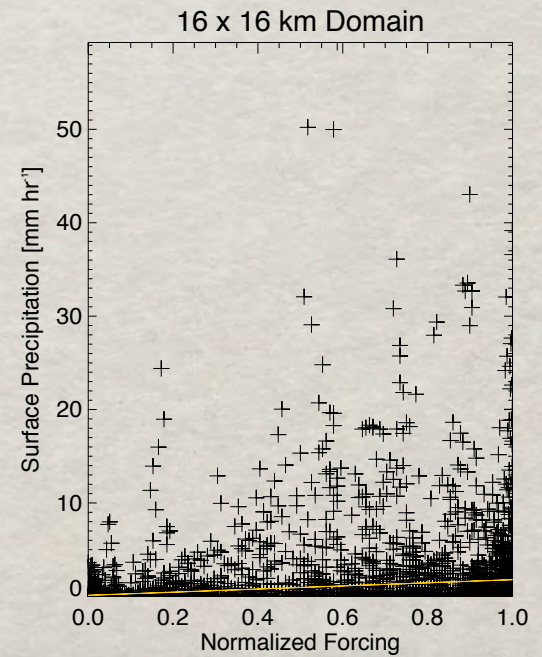
$r^2 = 0.723$
Slope = 2.527



$r^2 = 0.095$
Slope = 2.082



$r^2 = 0.049$
Slope = 1.809

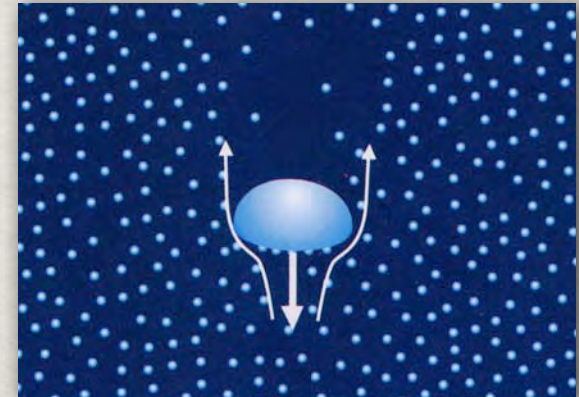
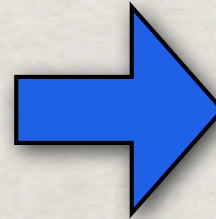
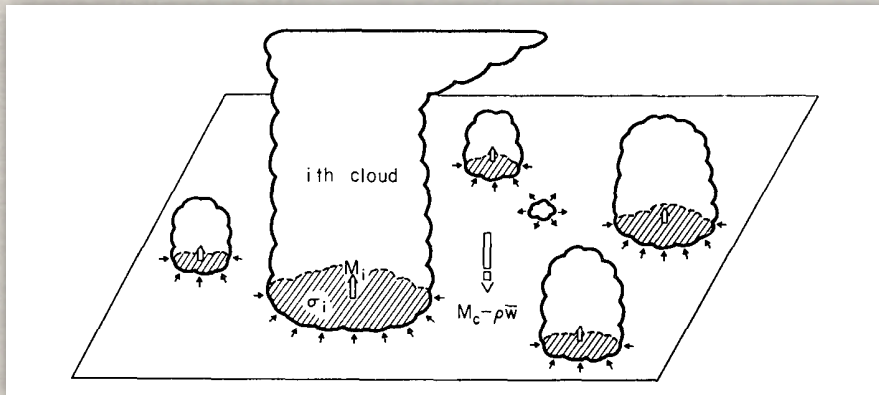


$r^2 = 0.030$
Slope = 1.673

LIMITS OF EQUILIBRIUM

- ✱ QE purports to describe “expected values” of Q_1 and Q_2 .
- ✱ A cloud-resolving model generates a “realization” of Q_1 and Q_2 .
- ✱ A stochastic parameterization aims to generate a “realization” of Q_1 and Q_2 .
- ✱ For weather prediction, we will need ensembles of realizations.
- ✱ We can imagine a generalization of QE that predicts pdfs of Q_1 and Q_2 . This is an interesting but difficult challenge.

HEATING AND DRYING ON FINE AND COARSE GRIDS



- ☼ Parameterizations for low-resolution models are designed to describe the collective effects of ensembles of clouds.

- ☼ Parameterizations for high-resolution models are designed to describe what happens inside individual clouds.

Scale-dependence of heating & drying

$$Q_1 \equiv LC - \frac{1}{\rho} \frac{\partial}{\partial z} (\rho \overline{w' s'}) - \frac{1}{\rho} \nabla_H \cdot (\rho \overline{\mathbf{V}_H' s'}) + Q_R$$
$$Q_2 \equiv LC + \frac{L}{\rho} \frac{\partial}{\partial z} (\rho \overline{w' q_v'}) + \frac{L}{\rho} \nabla_H \cdot (\rho \overline{\mathbf{V}_H' q_v'})$$

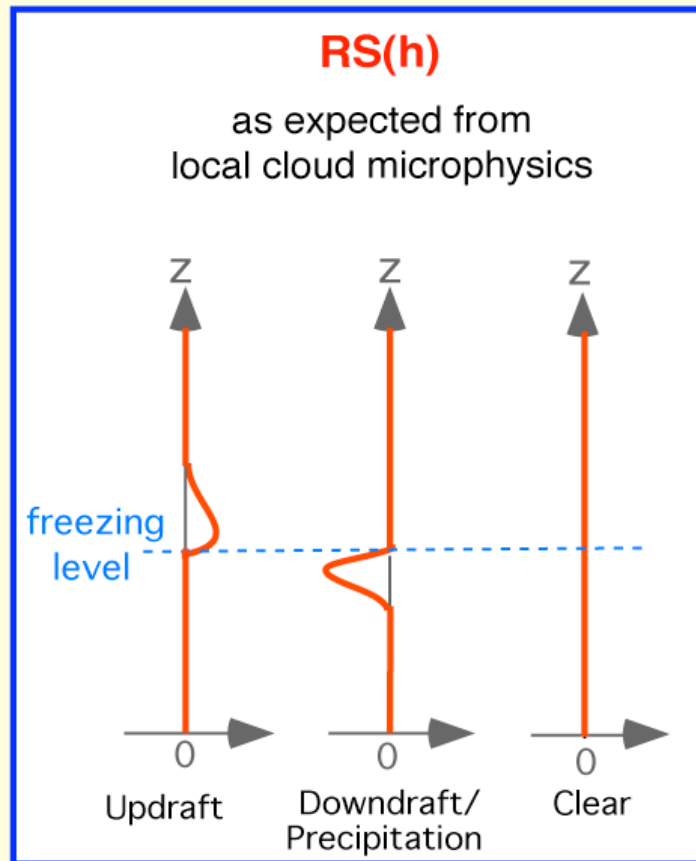
*These quantities are **defined** in terms of spatial averages.*

As the averaging length becomes smaller:

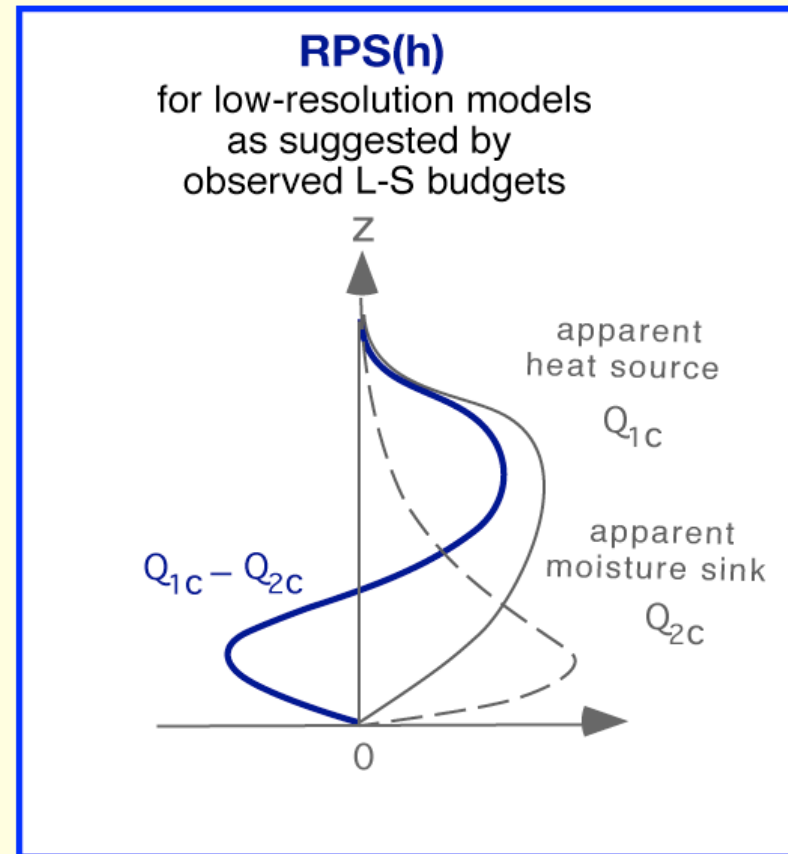
- ◆ **The vertical transport terms become less important. Later horizontal averaging does not change this.**
- ◆ **The horizontal transport terms become more important locally. Horizontal averaging kills them, though.**
- ◆ **The phase-change terms become dominant.**

Typical vertical profiles of the apparent moist static energy source due to convective activity

Fine mesh



Coarse mesh



Any space/time/ensemble averages of the profiles in the left panel do NOT give the profile in the right panel.

Three problems with parameterizations at high resolution:

- ◆ **The sample size is too small.**
- ◆ **The “resolved-scale forcing” varies too quickly.**
- ◆ **Convective transports should give way to microphysics, but we have no quantitative theory for this transition.**

**Expected values --> Individual realizations
--> Global CRMs**

You can run a high-resolution model with a cumulus parameterization.

- ◆ **The model won't blow up.**
- ◆ **Numbers will come out.**
- ◆ **You can make a plot.**
- ◆ **You can publish a paper.**
- ◆ **You can get a grant.**
- ◆ **You won't be arrested.**