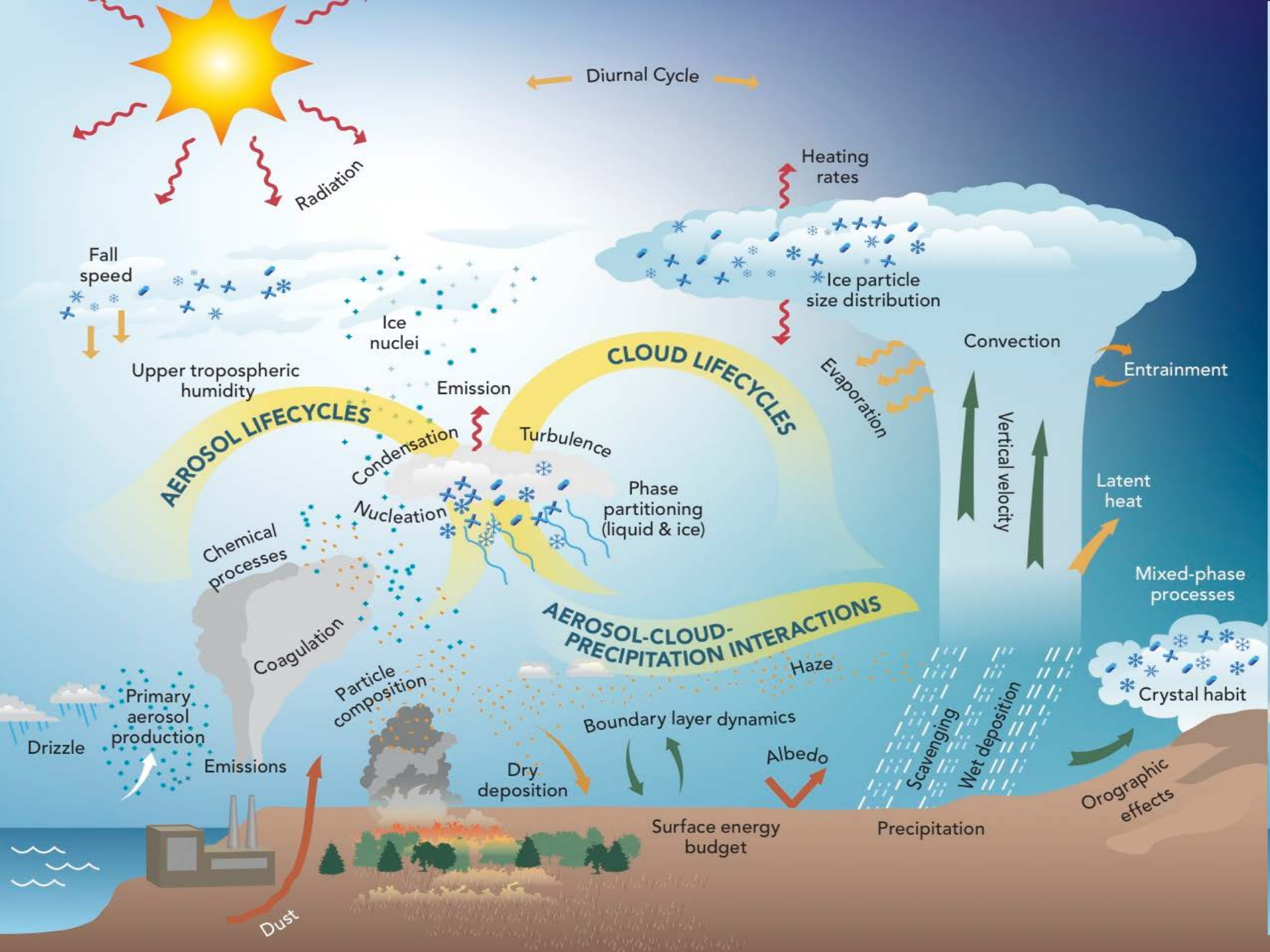


# Ice Cloud Microphysics Issues

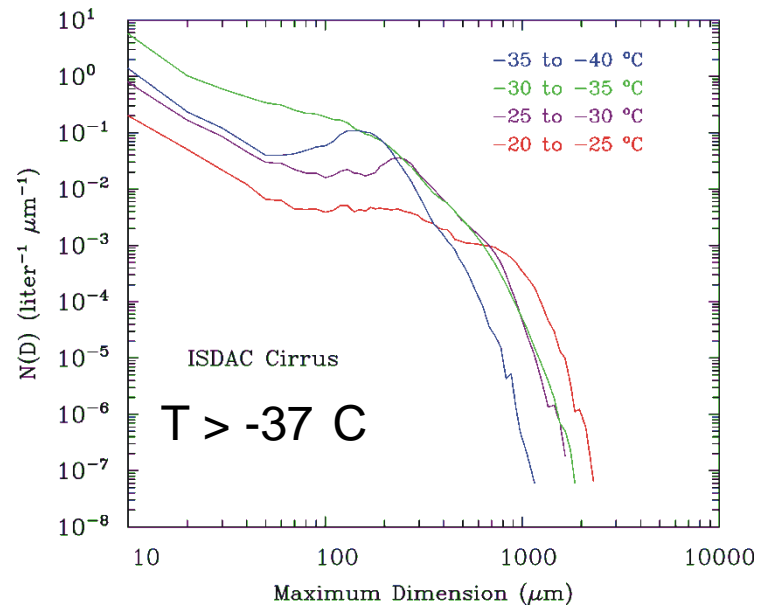
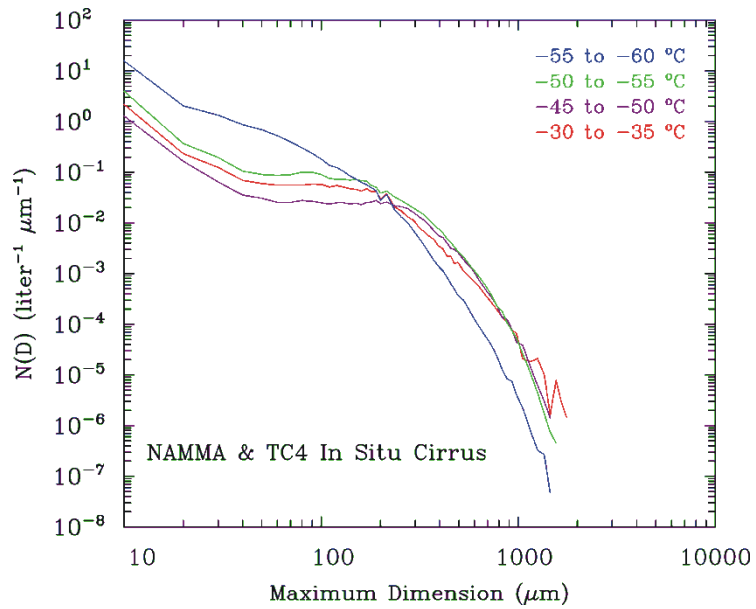
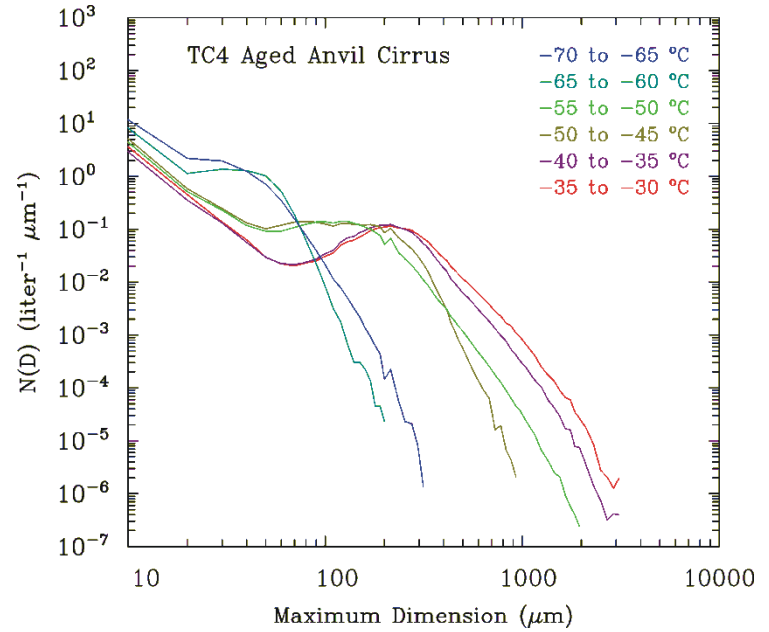
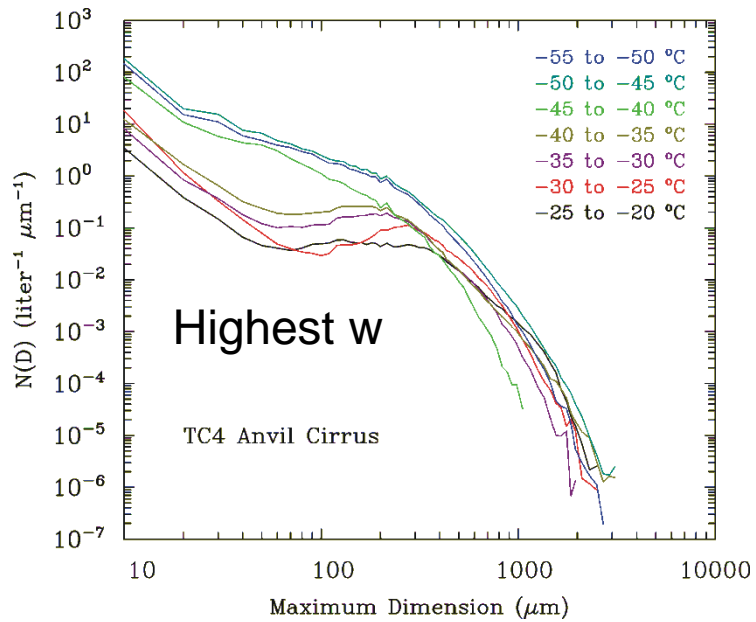
**David L. Mitchell and Subhashree Mishra**  
*Desert Research Institute, Reno, Nevada*

**R. Paul Lawson and Brad Baker**  
*SPEC, Inc., Boulder, Colorado*

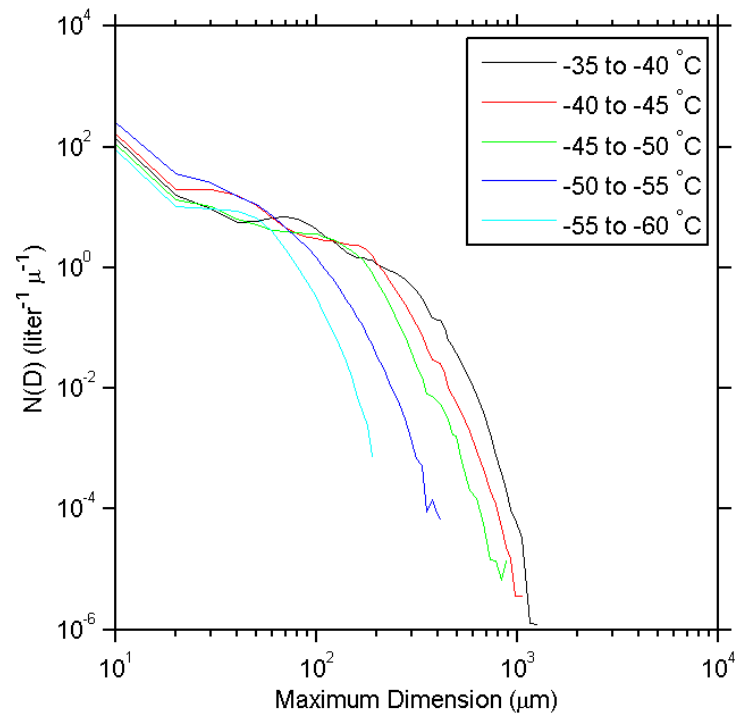
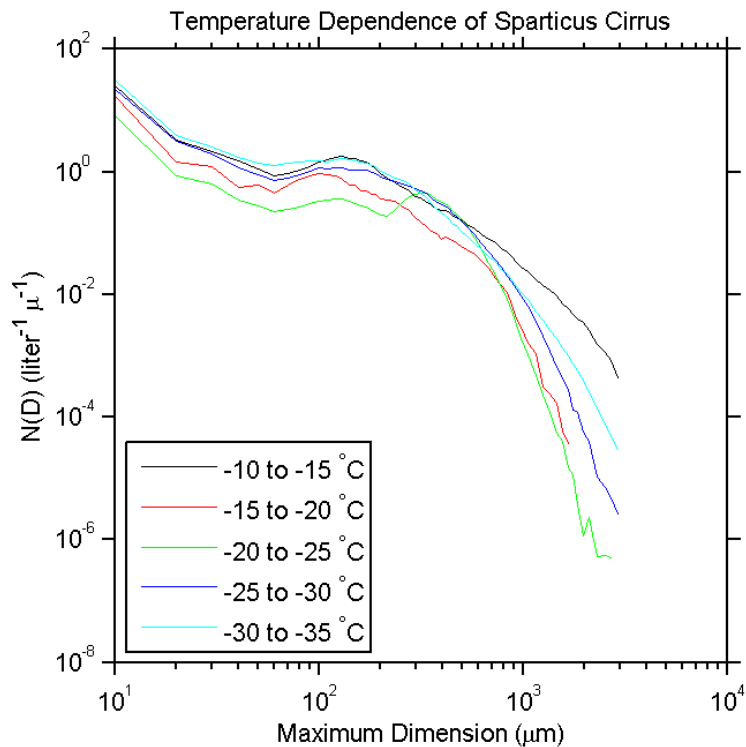




# Homogeneous freezing nucleation at higher updrafts?

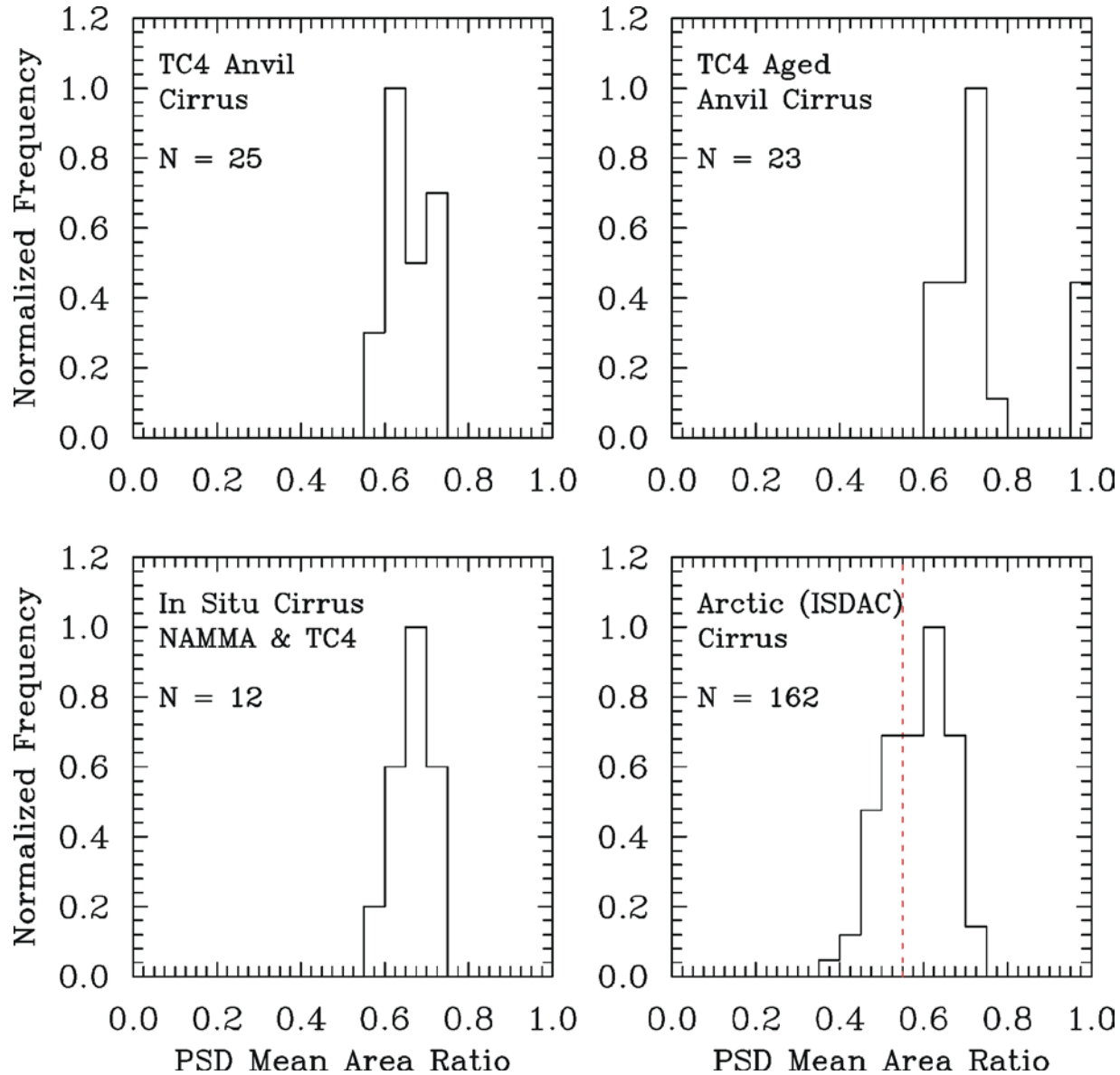


# Homogeneous freezing nucleation in synoptic midlatitude cirrus?



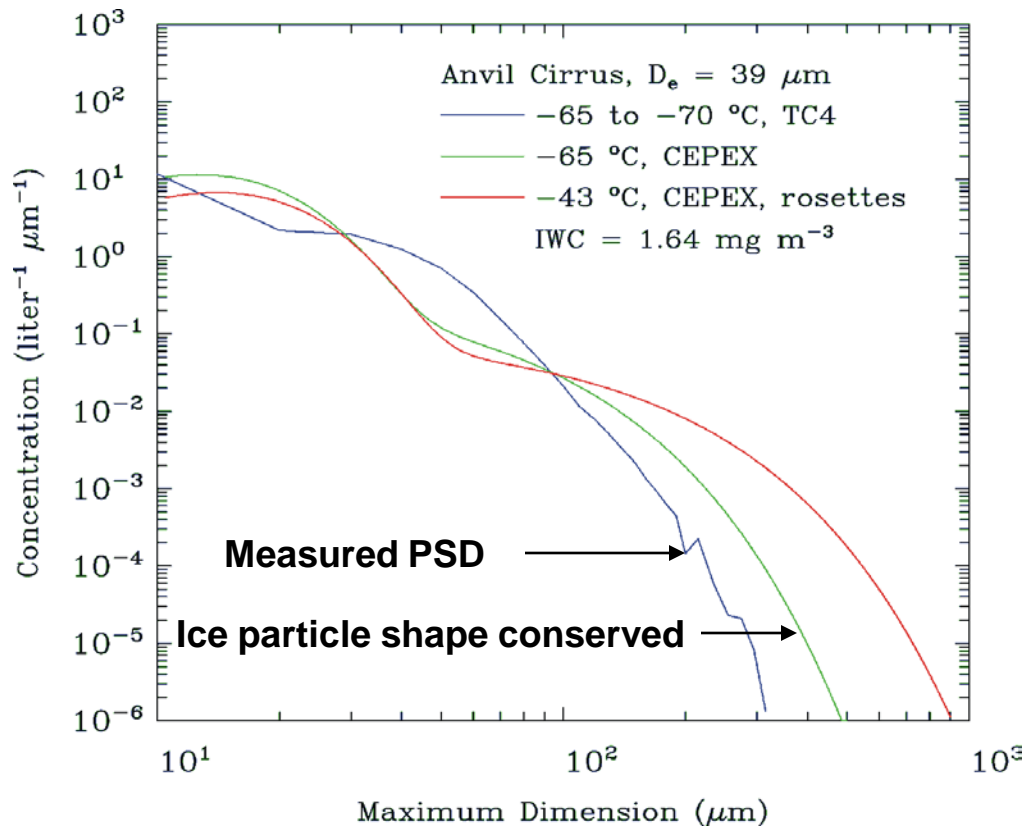


# Arctic cirrus crystals have different area ratios

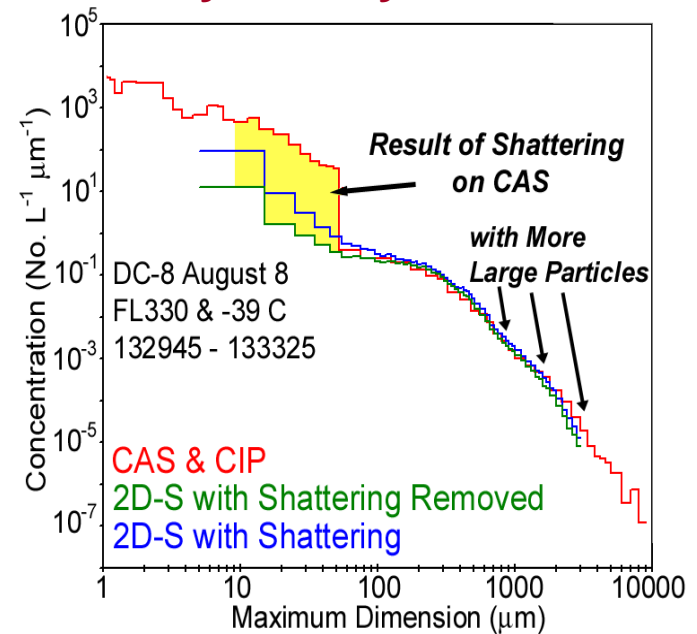


Conventional wisdom is that for a given  $D_e$  and IWC, ice cloud optical properties are not sensitive to ice crystal shape and PSD shape. Often this is not true.

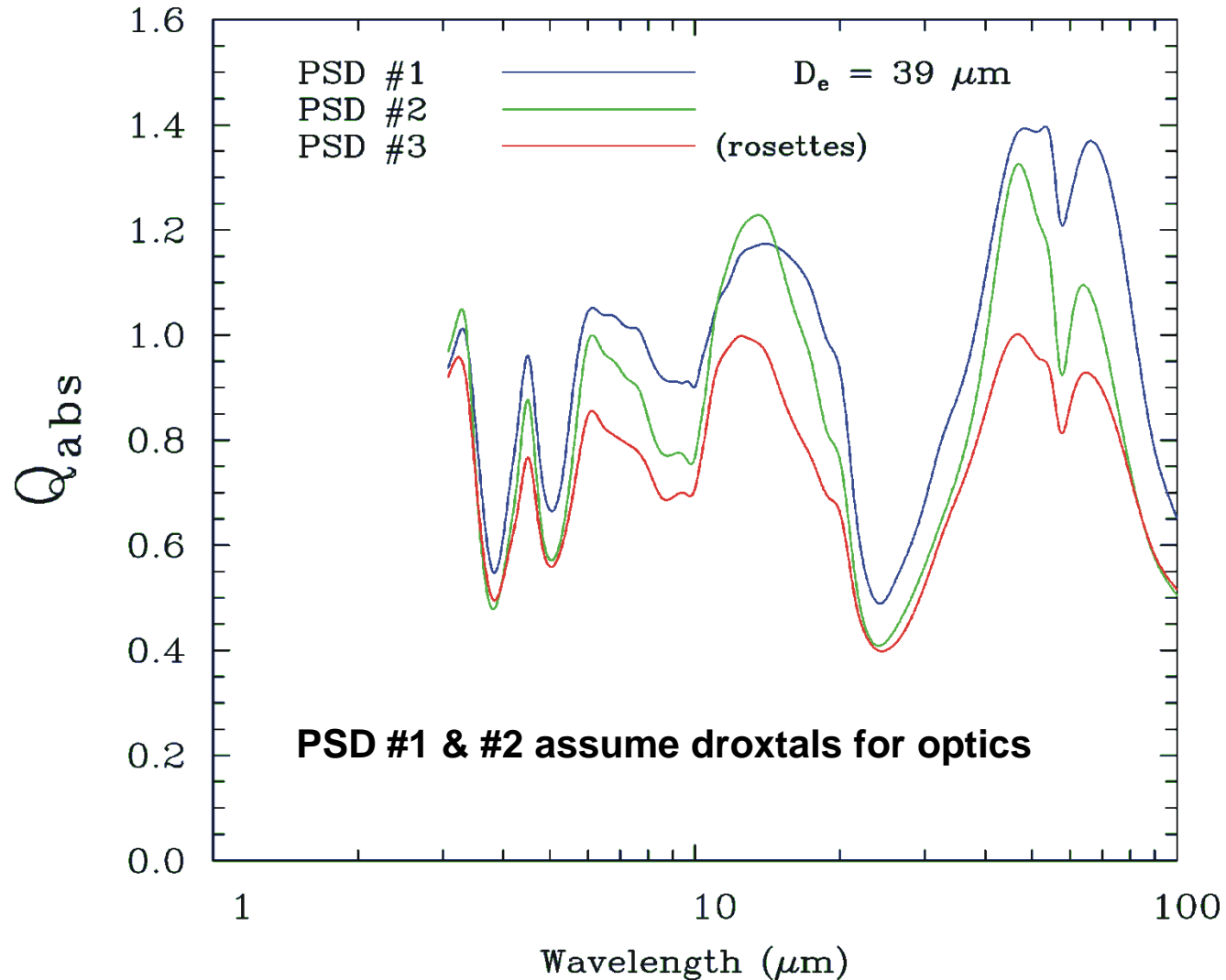
## Changing the ice crystal shape changes the PSD for constant $D_e$ & IWC (red curve).



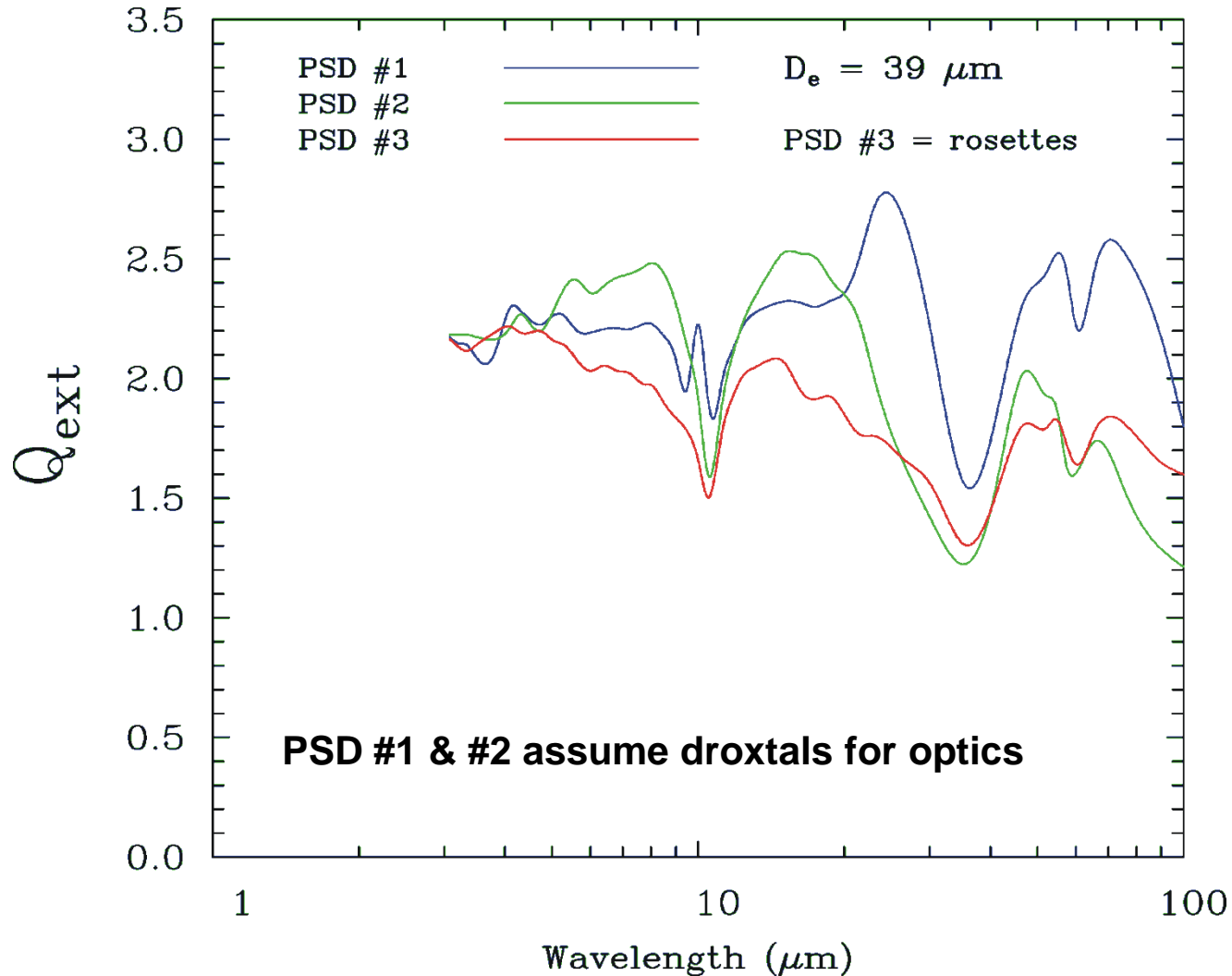
**Shattering often increases small crystals by  $10^2$**



Calculate optical properties for (1) measured PSD; (2) altered PSD with same particle shapes; (3) altered PSD with bullet Rosettes. All use ice optics database of Yang et al. (2005).

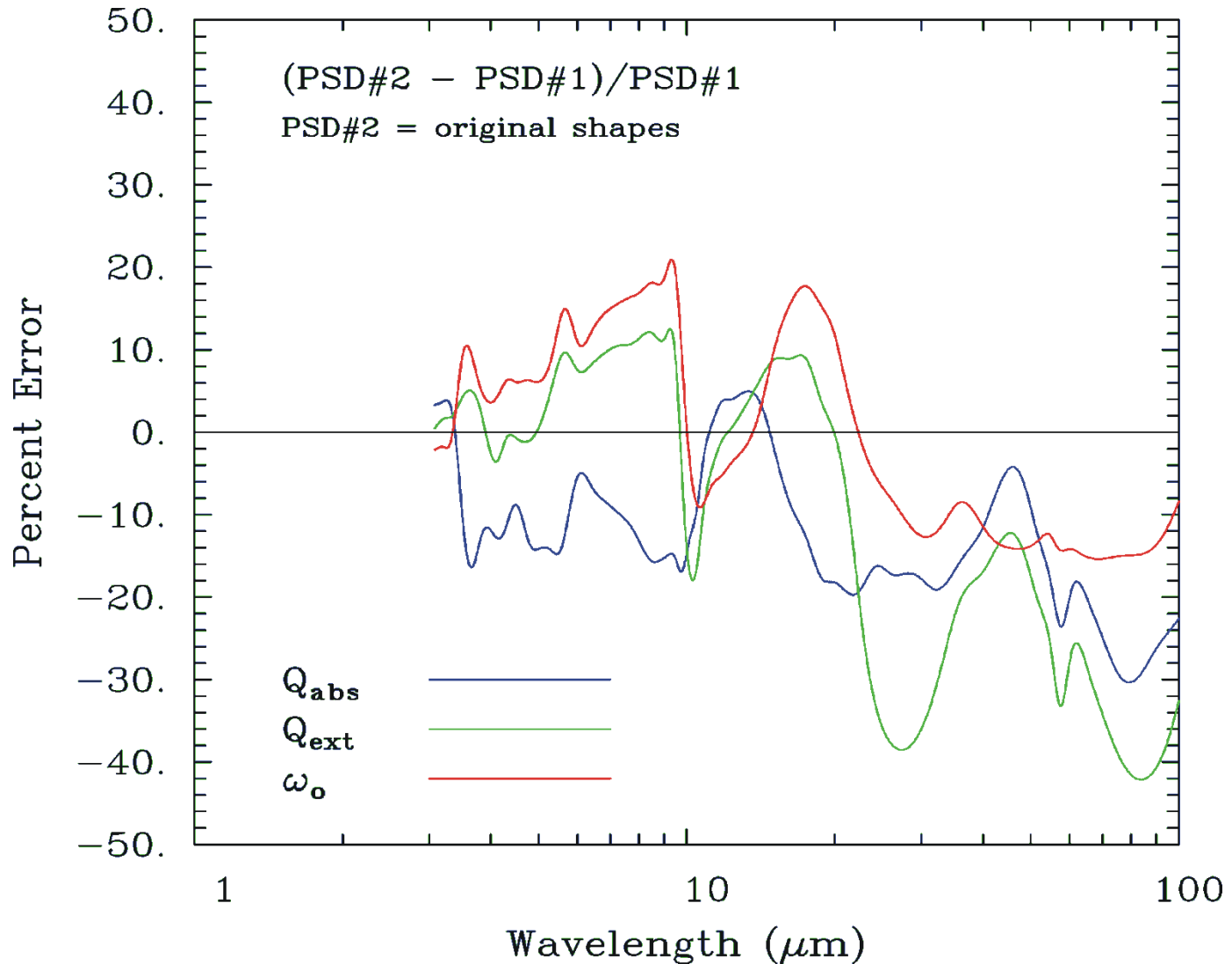


Extinction efficiencies for (1) measured PSD; (2) altered PSD with same particle shapes; (3) altered PSD with bullet rosettes.

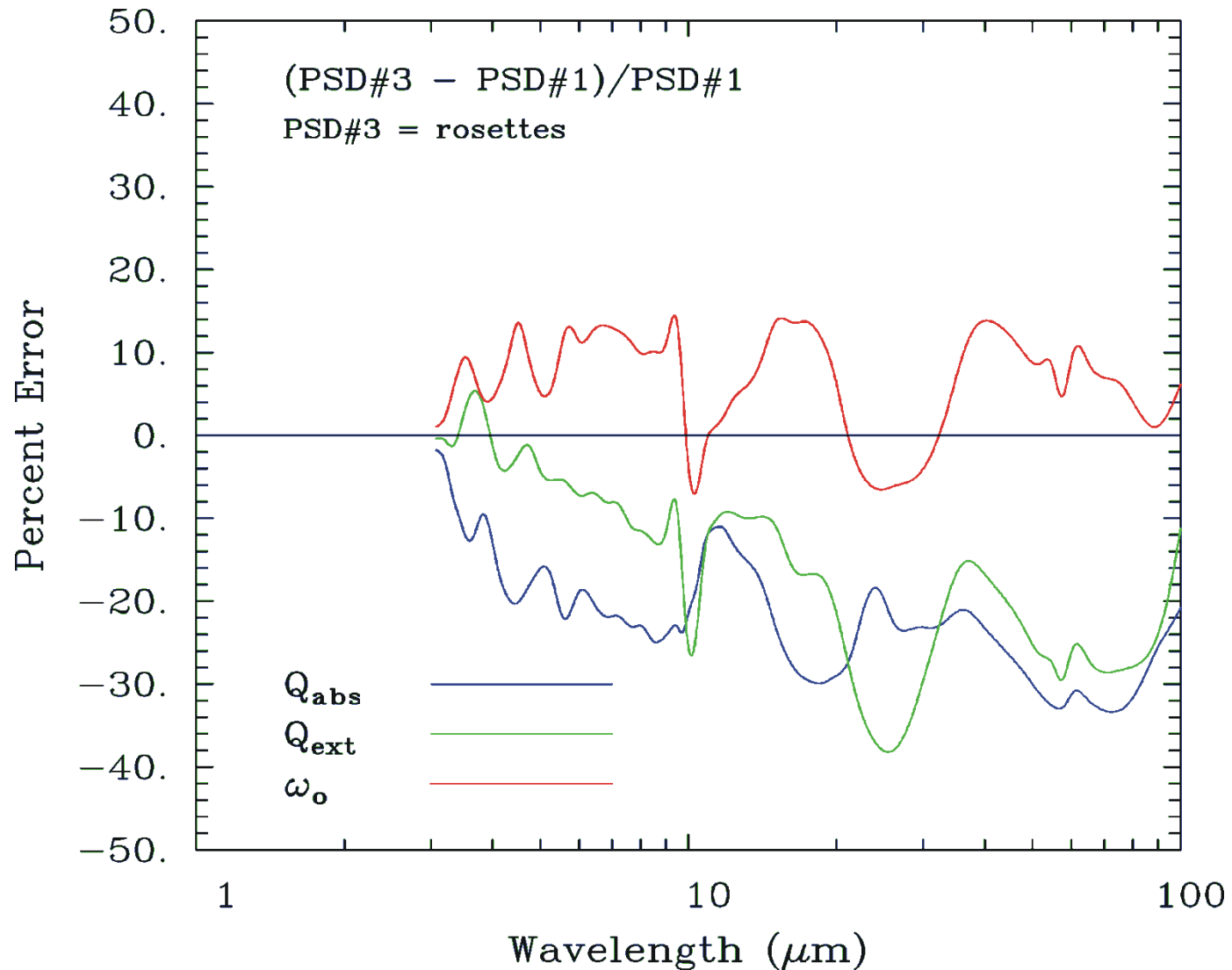




# Errors when only PSD shape changes and ice particle shapes remain constant. Implications for cloud radar.



# Errors when both PSD shape and ice particle shape change



## ICE OPTICS CONCLUSIONS

1. Changes in PSD shape alone (while holding  $D_e$  and particle shape constant) substantially affects IR ice optical properties.
2. For constant  $D_e$ , changing the ice particle shape assumption further changes the PSD shape, which further changes optical properties relative to the reference (i.e. measured) PSD.
3. Climate modeling challenge can be met by using the Modified Anomalous Diffraction Approximation (MADA) that is formulated in terms of PSD parameters and ice particle shape attributes.

## Parameterizing the Ice Fall-speed from In Situ Measurements: General Approach

1. The size resolved 2D-S measurements of number, projected area and mass concentration appear reasonable.
  - Ice artifacts from shattering greatly reduced
  - Good agreement between 2D-S and CVI IWC during TC4

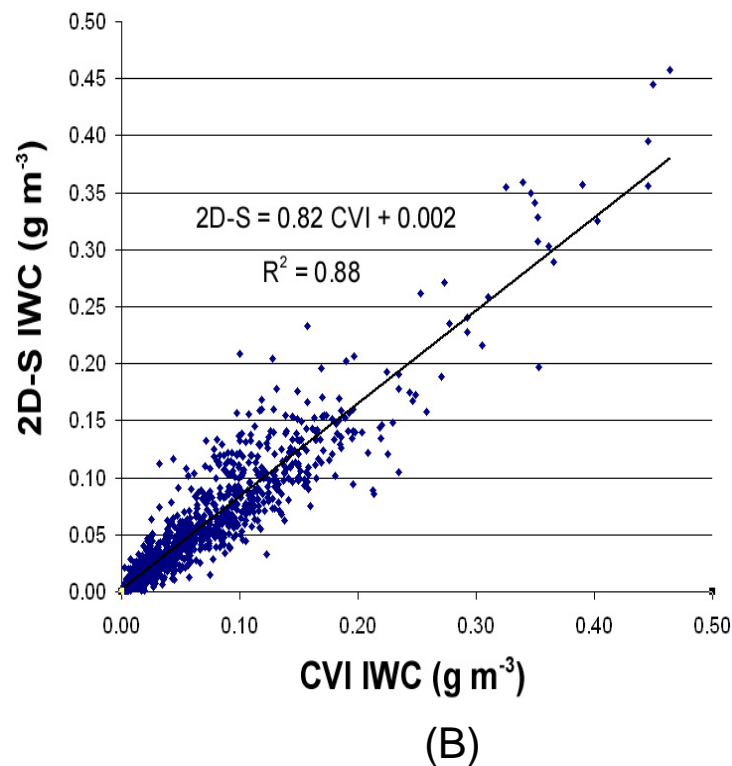
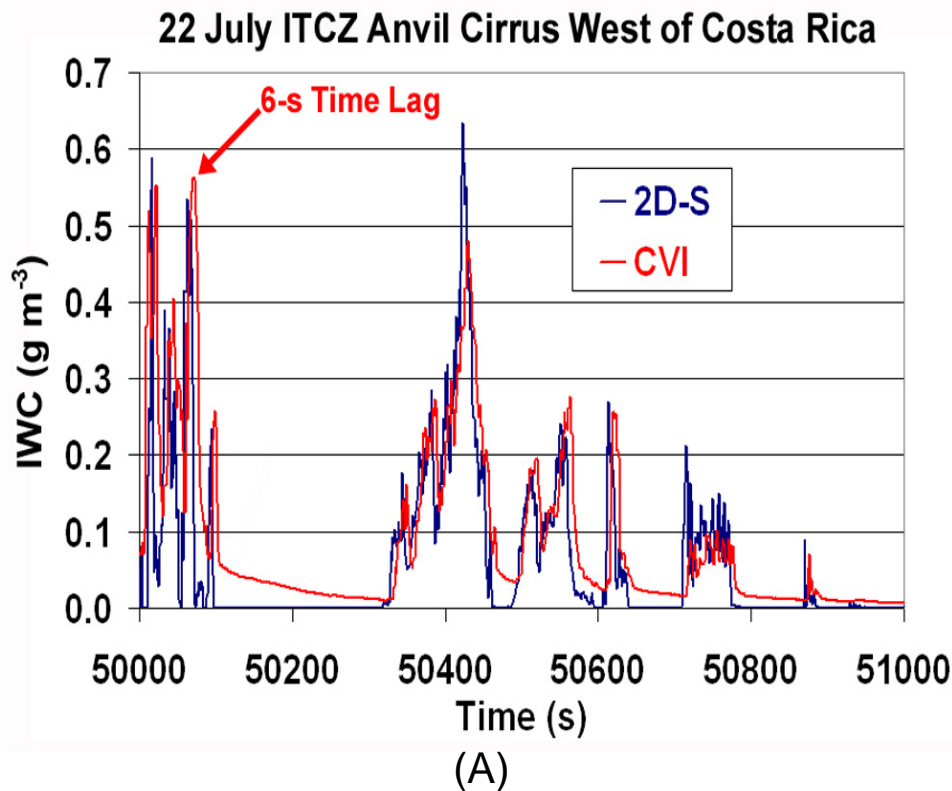
2. Therefore calculate  $V_m$  and  $D_e$  **directly** from these measurements:

$$V_m = \sum v(D) m(D) N(D) \Delta D / \sum m(D) N(D) \Delta D$$

$$D_e = (3/2) \sum m(D) N(D) \Delta D / (\rho_i \sum A(D) N(D) \Delta D)$$

- $m(D)$  &  $A(D)$  are bin mass or bin area concentration / bin number conc.
3. Relate  $V_m$  and  $D_e$  to T and IWC for model validation purposes
  4. Relate  $V_m$  to  $D_e$  to predict  $V_m$  from the model microphysics scheme

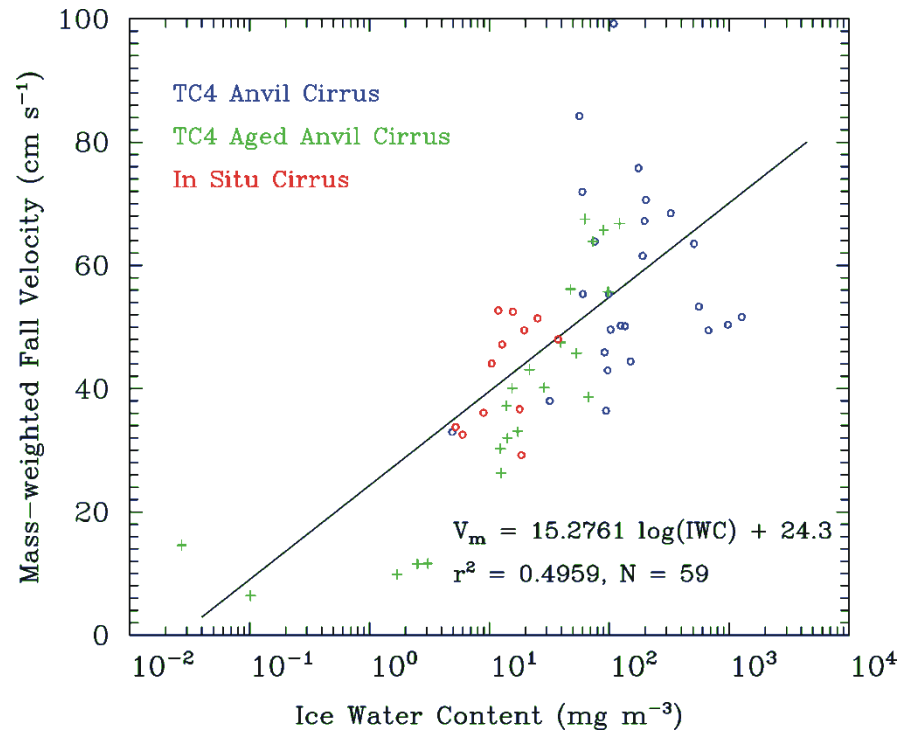
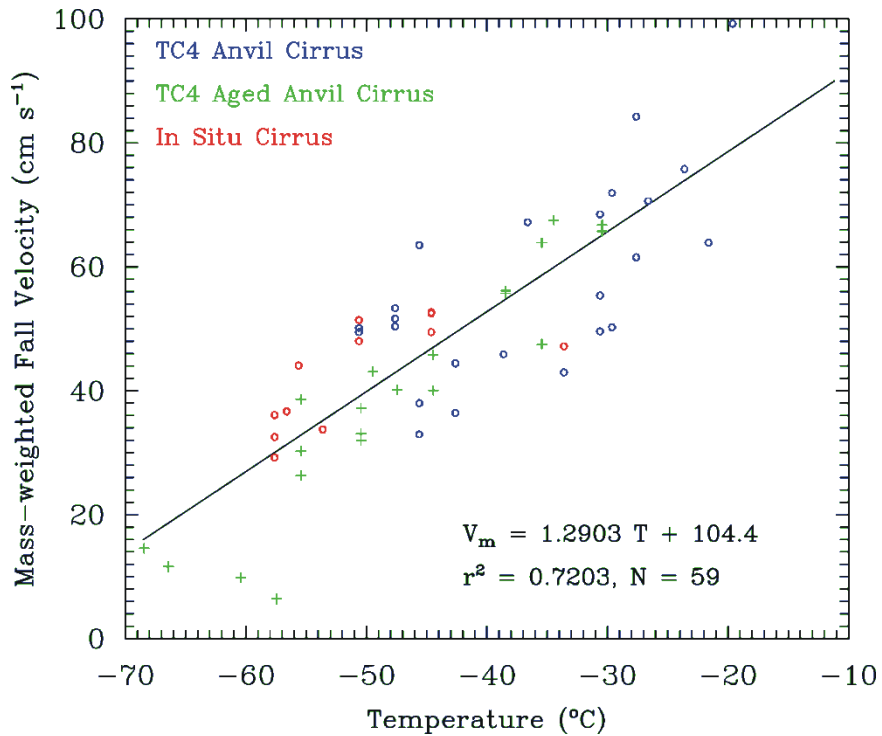
# COMPARISON OF 2D-S AND CVI IWCs DURING TC4



A: Time series of the 2D-S and CVI IWC for a TC4 case study. CVI response time lagged 6 seconds behind 2D-S measurements, producing a slight offset.

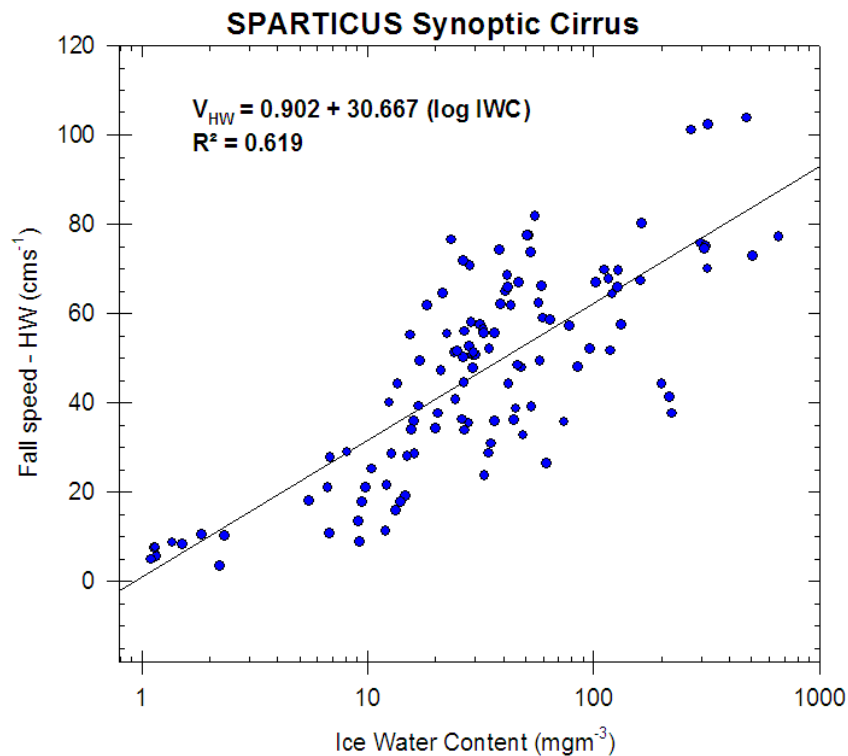
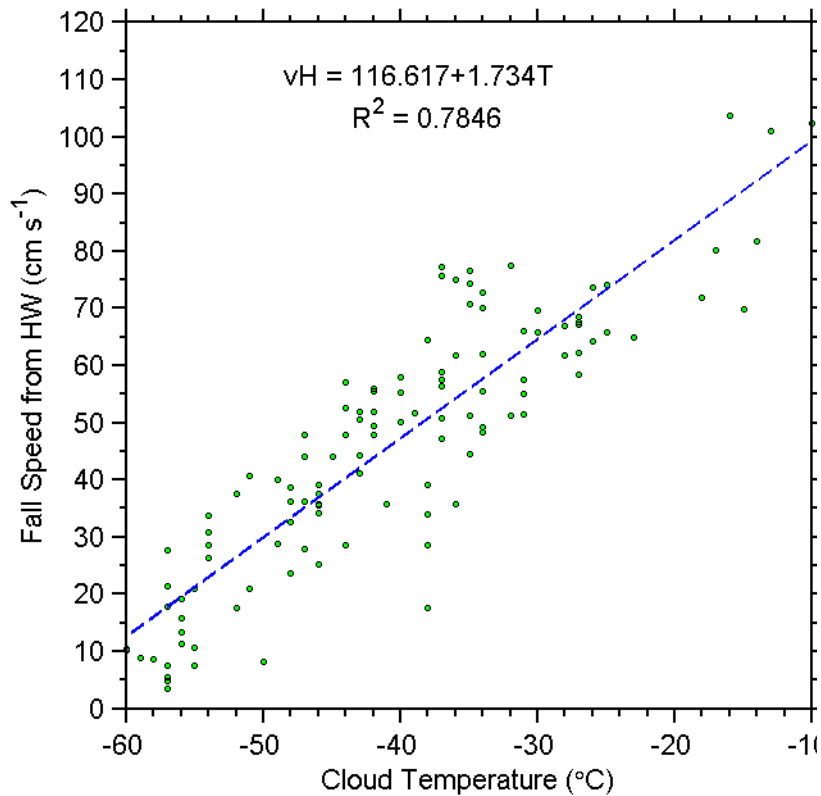
B: 2D-S IWCs compared with CVI IWCs for 12,000 1-Hz measurements (averaged over 10-s) in TC4 anvils cirrus.

Mass-weighted fall velocity was related to both temperature and IWC during TC4 and SPARTICUS but not during ISDAC. Why?

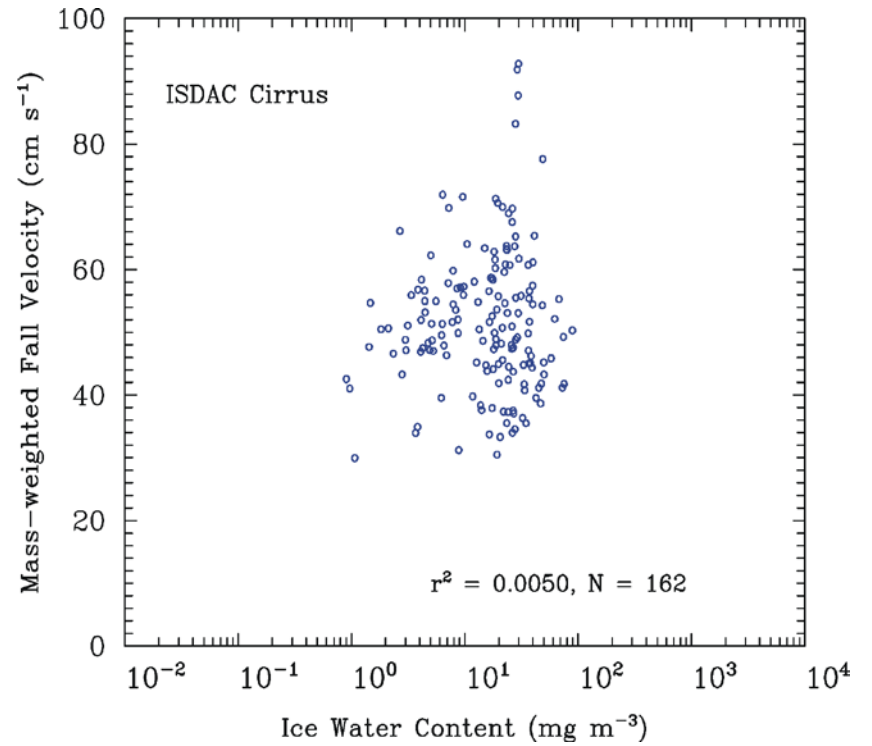
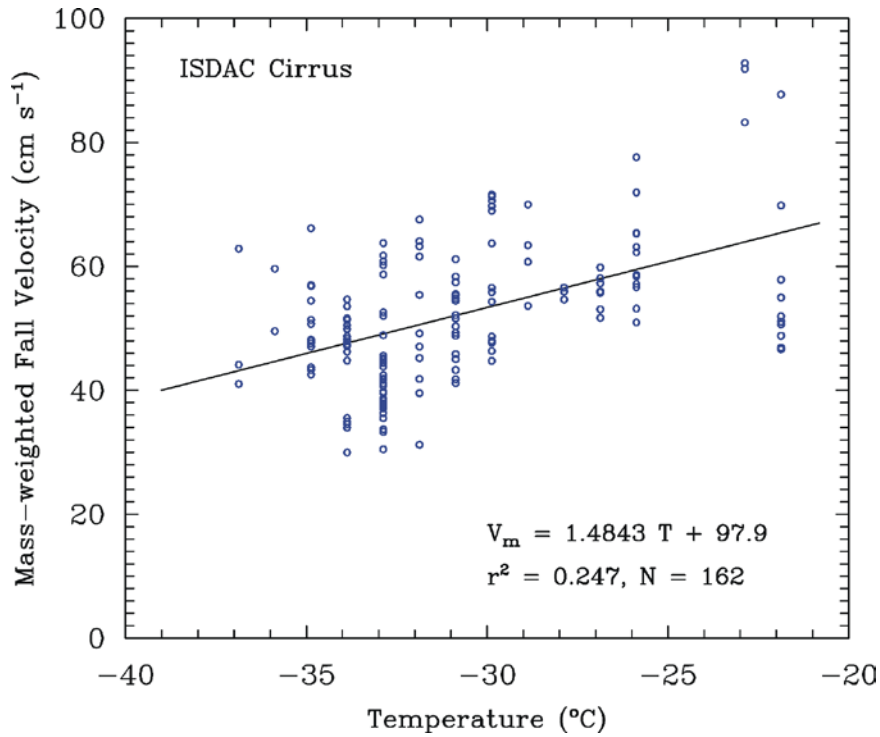


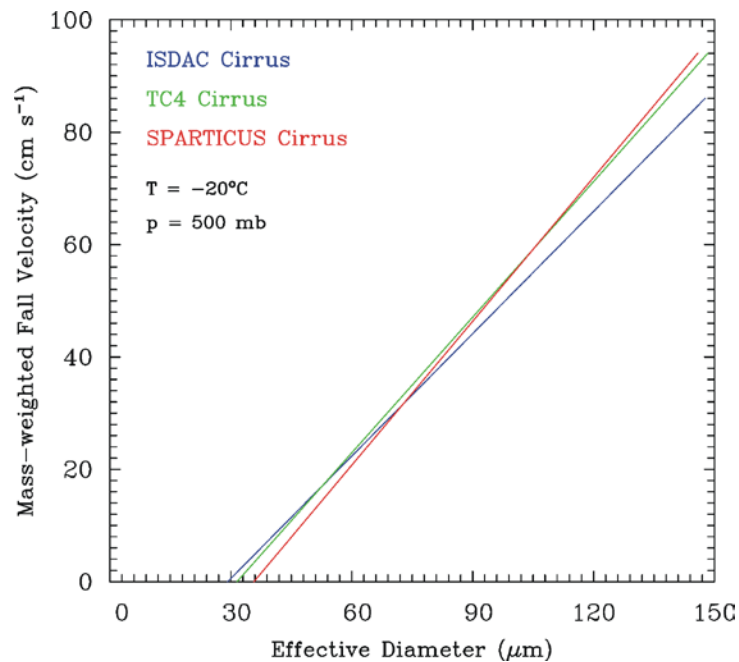
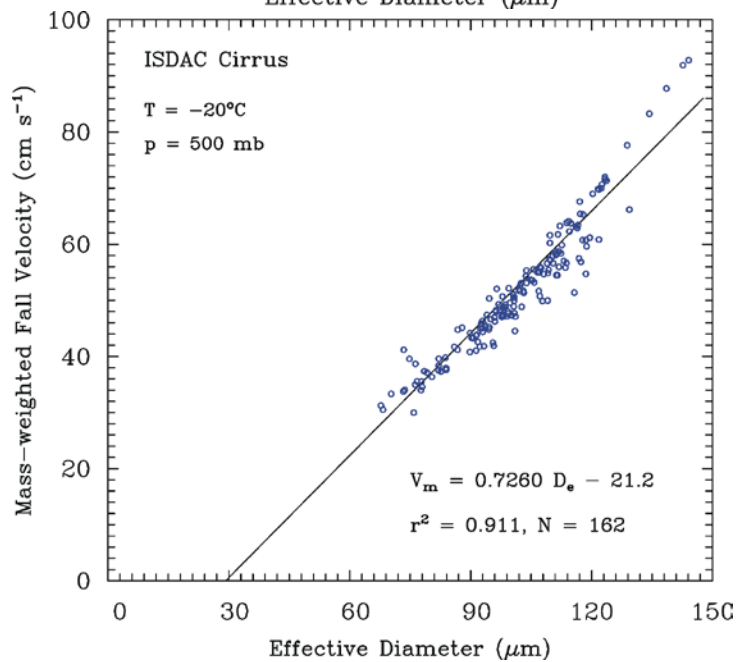
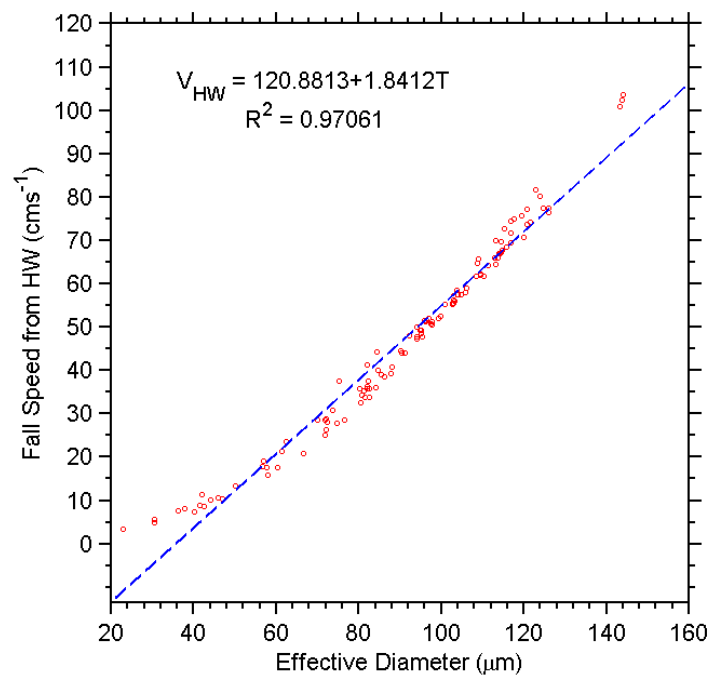
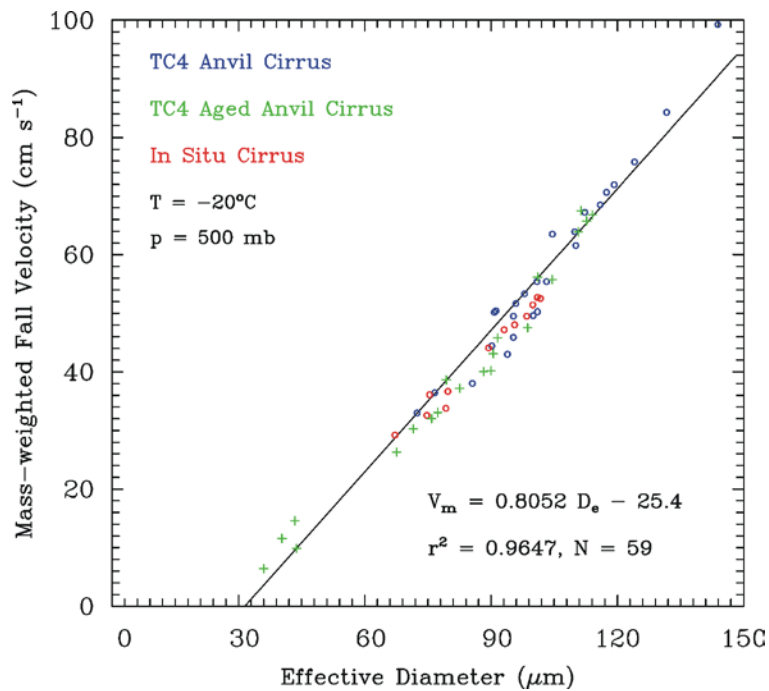


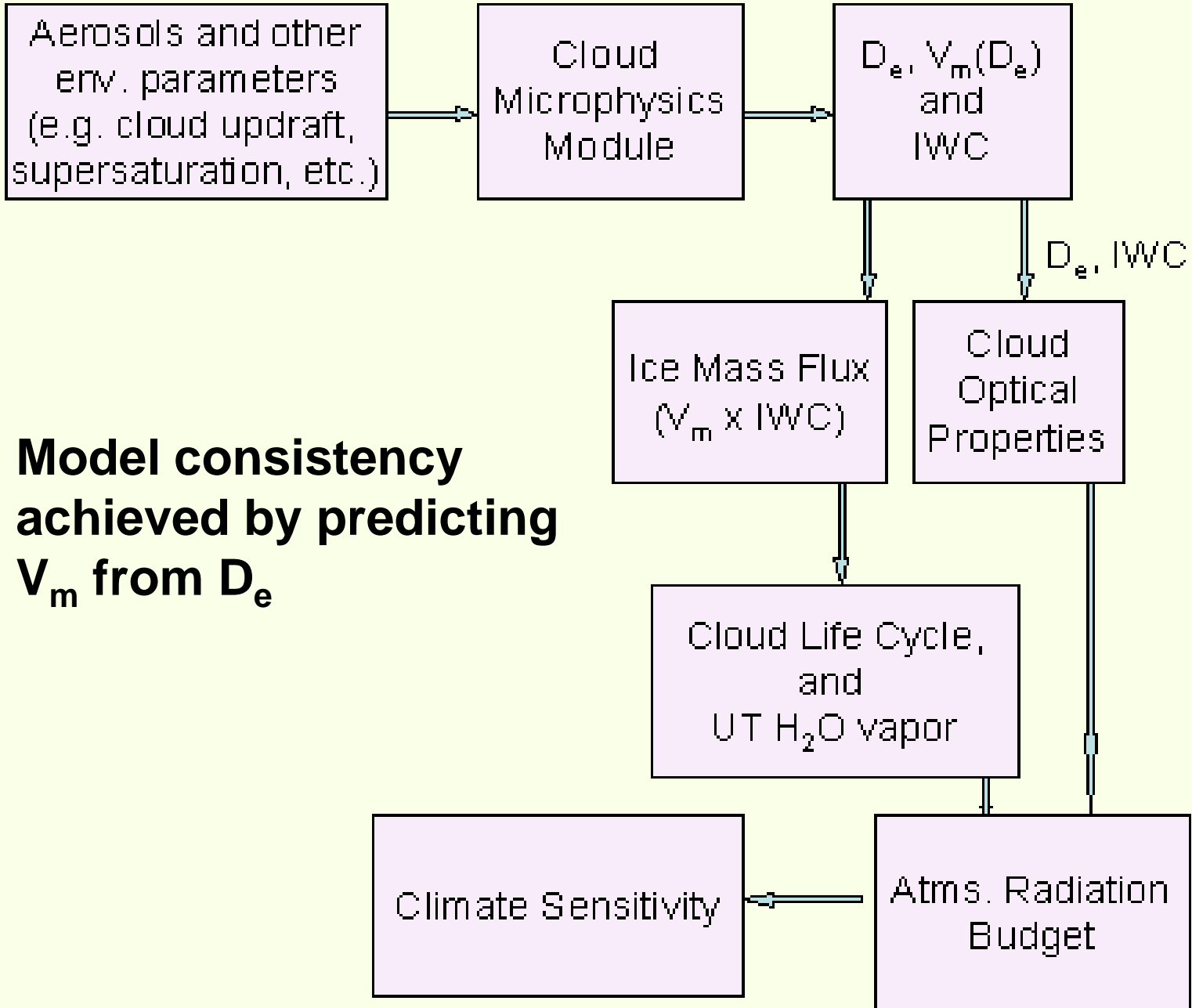
# SPARTICUS SYNOPTIC CIRRUS: $V_m$ vs. T and $V_m$ vs. IWC



# ISDAC FIELD CAMPAIGN







**EXTRA SLIDES**





