

Observations of Aerosol Effects on the Microphysics and Radiative Properties of Arctic Liquid-Phase Clouds



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Overview

- Aerosol indirect effects
 - Key climate system process; uncertainty
 - Requirement for studies in Arctic
- Indirect and Semi-Direct Aerosol Campaign (ISDAC)
 - Barrow, Alaska – April 2008



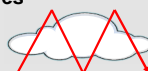
Project aircraft: National Research Council of Canada (NRC) Convair-580

Approach

- Focus on liquid-phase stratocumulus
- April 8, 26, 27: clean aerosol conditions
- April 19, 20: biomass burning (BB)
- In-situ observations
 - In-cloud:** droplet concentration, size (CDP, FSSP-100); liquid water (King probe)
 - Below-cloud:** aerosol particle concentration, size (PCASP, FSSP-300), single particle composition (SPLAT II)

Part 1: cloud microphysical and radiative properties

- Vertical profiles; porpoising legs



Part 2: cloud droplet activation

- Horizontal legs in- and below-cloud
- Droplet closure analysis



Part 1: Insight into Indirect Effects

- For comparable LWP: larger optical depth, smaller Re for polluted cases
 - Consistent with 1st indirect effect
- Precipitation inhibition: Re below ~ 10 μm (drizzle threshold) for all profiles
 - 2nd indirect effect
 - Clean cases: low LWC limits droplet growth
 - Polluted cases: high N_d limits droplet growth

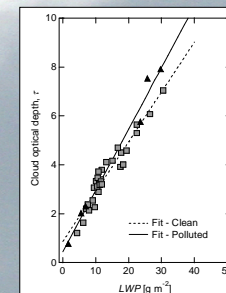


Figure 2: Optical depth as a function of LWP for clean and polluted cases with LWP < 50 g m⁻².

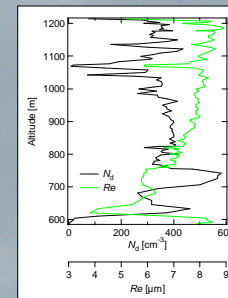


Figure 3: Vertical profiles of N_d and Re in-cloud for a representative polluted case on April 20.

Part 2: Aerosol Properties and Droplet Closure

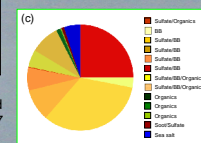
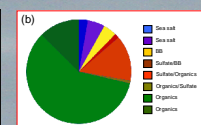
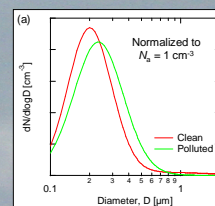


Figure 4: (a) Size distributions and chemical composition of below-cloud aerosol for (b) clean case on April 27 and (c) polluted case on April 20.

- Differences in size, composition of below-cloud aerosol
- Simulate activation in adiabatic parcel model
 - Updraft velocity from standard deviation of vertical velocity measurements in-cloud,
 - Minimize difference between measured and simulated N_d
- Polluted case: more sensitive to updraft velocity
 - Lower activated fraction, lower maximum supersaturation
 - Preferential activation of larger or more hygroscopic particles
- Different particle sizes activated in each aerosol-cloud regime
 - Implications for Re

Part 1: Average Cloud Properties

- Two distinct aerosol-cloud regimes
- Higher Re, LWP, albedo for clouds in more polluted conditions

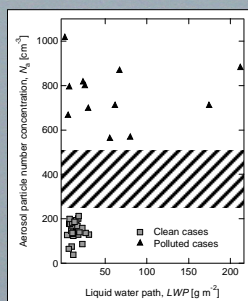


Figure 1: Aerosol concentration vs. liquid water path for all profiles.

Parameter	Clean	Polluted
Aerosol concentration, N _a [cm ⁻³]	147 ± 41	756 ± 132
Droplet concentration, N _d [cm ⁻³]	136 ± 31	304 ± 81
Activated fraction	0.96	0.41
Temperature, T [°C]	-12.9 ± 1.1	-7.5 ± 1.1
Liquid water content, LWC [g m ⁻³]	0.07 ± 0.02	0.16 ± 0.11
Cloud thickness, H _c [m]	180 ± 43	296 ± 64
Liquid water path, LWP [g m ⁻²]	13.4 ± 6.1	61.9 ± 66.8
Droplet effective radius, Re [μm]	5.4 ± 0.7	5.7 ± 1.2
Cloud optical thickness, τ	3.60 ± 0.30	14.13 ± 13.64
Cloud albedo, A	0.34 ± 0.08	0.55 ± 0.25

Table 1: Average cloud and aerosol parameters and standard deviations for all profiles under clean and polluted aerosol conditions.

Case	Hygroscopicity	Updraft velocity cm s ⁻¹	% Difference N _d
Clean	0.3	0.6 – 1	8 %
Polluted	0.3	0.5	3 %

Table 2: Parcel model simulation results. All simulations assume an internally-mixed aerosol and condensation coefficient of 1.

Summary and Future Work

- Two distinct aerosol-cloud regimes observed in Arctic springtime clouds (liquid-phase)
- Assessment of indirect effects complicated by variations in LWP, droplet activation
- Precipitation suppression (2nd indirect effect) observed in both clean and polluted cases
- Future Work:** extend analysis to additional cases; incorporate updrafts from LES