

Northwest



The Northwest's rapidly growing population, as well as its forests, mountains, rivers, and coastlines, are already experiencing human-induced climate change and its impacts¹. Regionally-averaged temperature rose about 1.5°F over the past century² (with some areas experiencing increases up to 4°F) and is projected to increase another 3 to 10°F during this century³, with higher emissions scenarios[†] resulting in the upper end of this range. Increases in winter precipitation and decreases in summer precipitation are projected by many climate models⁴, though these projections are less certain than those for temperature. Impacts related to changes in snowpack, streamflows, sea level, forests, and other important aspects of life in the Northwest are already underway, with more severe impacts expected over coming decades in response to continued and more rapid warming.

Declining springtime snowpack leads to reduced summer streamflows, straining water supplies.

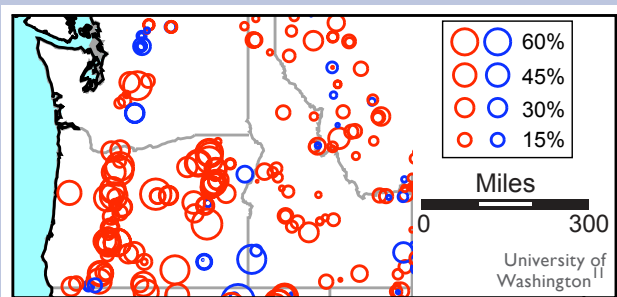
The Northwest is highly dependent on temperature-sensitive springtime snowpack to meet growing, and often competing, water demands such as municipal and industrial uses, agricultural irrigation, hydropower production, navigation, recreation, and in-stream flows that protect aquatic ecosystems including threatened and endangered species. Higher cool season (October through March) temperatures cause more precipitation to fall as rain rather than snow and contribute to earlier snowmelt. April 1 snowpack, a key indicator of natural water storage available for the warm season, has already declined substantially throughout the region. The average decline in the Cascade Mountains, for example, was about 25 percent over the past 40 to 70 years, with most of this due to the 2.5°F increase in cool season temperatures over that period^{5,6}. Further declines in Northwest snowpack are projected to result from additional warming over this century, varying with latitude, elevation, and proximity to

the coast. April 1 snowpack is projected to decline as much as 40 percent in the Cascades by the 2040s⁷. Throughout the region, earlier snowmelt will cause a reduction in the amount of water available during the warm season⁸.

In areas where it snows, a warmer climate means major changes in the timing of runoff: streamflow increases in winter and early spring, and decreases in late spring, summer, and fall. This shift in streamflow timing already has been observed over the past 50 years⁹, with the peak of spring runoff shifting from a few days earlier in some places to as much as 25 to 30 days earlier in others¹⁰.

Larger changes are expected due to increased warming, with runoff projected to shift 20 to 40 days earlier in this century¹⁰. Reductions in summer water availability will vary with midwinter temperatures experienced in different parts of the region. In relatively warm areas on the western slopes of the Cascade Mountains, for example, reductions in warm season (April through September) runoff of 30 percent or more are projected by mid-century, whereas colder areas in the Rocky Mountains are expected to see reductions on the order of 10 percent. Areas dominated by rain rather than snow are not expected to see major shifts in the timing

Trends in April 1 Snow Water Equivalent 1950-2002



April 1 snowpack (a key indicator of natural water storage available for the warm season) has declined throughout the Northwest. In the Cascade Mountains, April 1 snowpack declined by an average of 25 percent, with some areas experiencing up to 60 percent declines. On the map, decreasing trends are in red and increasing trends are in blue¹².

of runoff¹³. Extreme high and low streamflows also are expected to change with warming. Increasing winter rainfall (as opposed to snowfall) is expected to increase winter flooding in relatively warm watersheds on the west side of the Cascades. The already low flows of late summer are projected to decrease further due to both earlier snowmelt and increased evaporation and water loss from vegetation. Projected decreases in summer precipitation would exacerbate these effects. Some sensitive watersheds are projected to experience both increased flood risk in winter and increased drought risk in summer due to warming.

The region's water supply infrastructure was built based on the assumption that most of the water needed for summer uses would be stored naturally in snowpack. For example, the storage capacity in Columbia Basin reservoirs is only 30 percent of the annual runoff, and many small urban water supply systems on the west side of the Cascades store less than 10 percent of their annual flow¹⁴. Besides providing water supply and managing flows for hydropower, the region's reservoirs are operated for flood-protection purposes and, as such, might have to release (rather than store) large amounts of runoff during the winter and early spring to maintain enough space for flood protection. Earlier flows would thus place more of the year's runoff into the category of hazard rather than resource. An ad-

vance in the timing of snowmelt runoff would also increase the length of the summer dry period, with important consequences for water supply, ecosystems, and wildfire management¹⁰.

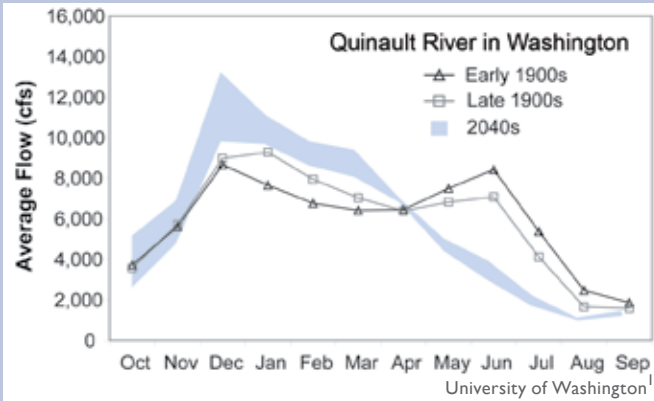
One of the largest demands on water resources in the region is hydroelectric power production. About 70 percent of the Northwest's energy needs are provided by hydropower, a far greater percentage than in any other region. Warmer summers will increase electricity demands for air conditioning and refrigeration at the same time of year that lower streamflows will lead to reduced hydropower generation. At the same time, water is needed for irrigated agriculture, protecting fish species, reservoir and river recreation, and urban uses. Conflicts between all of these water uses are expected to increase, forcing complex trade-offs between competing objectives¹⁵.

Increased insect outbreaks, wildfires, and changing species composition in forests will pose challenges for ecosystems.

Higher summer temperatures and earlier spring snowmelt are expected to increase the risk of forest fires in the Northwest by increasing summer moisture deficits; this pattern has already been observed in recent decades. Drought stress and higher temperatures will decrease tree growth in most low- and mid-elevation forests and also will increase the frequency and intensity of mountain pine beetle and other insect attacks¹⁶, further increasing fire risk and reducing timber production, an important part of the regional economy. The mountain pine beetle outbreak in British Columbia has destroyed 33 million acres of trees so far, about 40 percent of the marketable pine trees in the province. By 2018, it is projected that the infestation will have run its course and over 78 percent of the mature pines will have been killed; this will affect more than one-third of the total area of British Columbia's forests¹⁷ (see *Ecosystems* sector). Idaho's Sawtooth Mountains are also now threatened by pine beetle infestation.

In the short term, high elevation forests on the west side of the Cascade Mountains are expected to see increased growth. In the longer term, forest growth

Shift to Earlier Peak Streamflow
Quinault River (Olympic Peninsula, northern Washington)



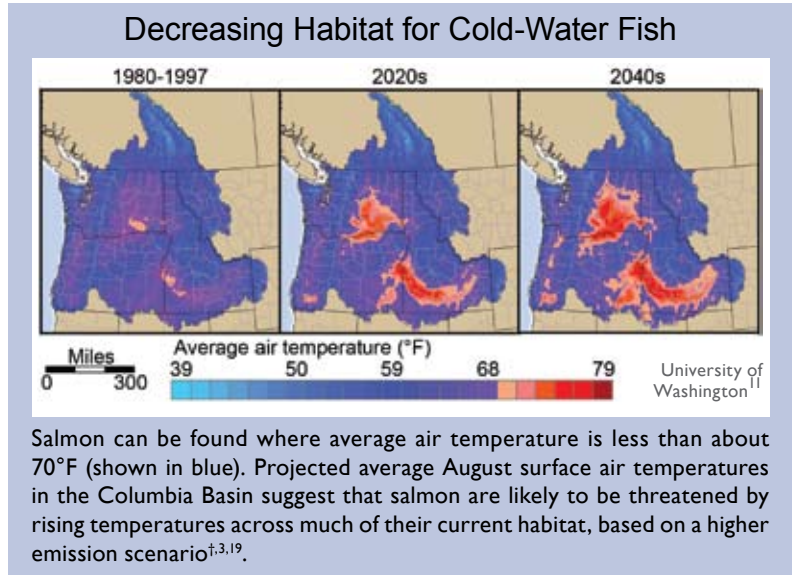
As precipitation continues to shift from snow to rain, by the 2040s, peak flow on the Quinault River is projected to occur in December, and flows in June are projected to be reduced to about half of what they were over the past century. On the graph, the blue swath represents the range of projected streamflows based on an increase in temperature of 3.6 to 5.4°F. The other lines represent streamflows in the early and late 1900s.¹⁵

L1 is expected to decrease as summertime soil
 L2 moisture deficits limit forest productivity, with
 L3 low-elevation forests experiencing these changes
 L4 first. The extent and species composition of
 L5 forests also are expected to change as tree spe-
 L6 cies respond to climatic changes. There is also
 L7 the potential for extinction of local populations
 L8 and loss of biological diversity if environmental
 L9 changes outpace species' ability to shift their
 L10 ranges and form successful new ecosystems.

L11
 L12 Agriculture, especially production of tree fruit
 L13 such as apples, is also an important part of the
 L14 regional economy. Decreasing irrigation supplies
 L15 and increased competition from weeds, pests,
 L16 and disease are likely to have negative effects on
 L17 agricultural production.

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 L20 **Salmon and other cold-water species**
 L21 **will experience additional stresses as a**
 L22 **result of rising water temperatures and**
 L23 **declining summer streamflows.**
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L25 Northwest salmon populations are at historically
 L26 low levels due to stresses imposed by a variety of
 L27 human activities including dam building, logging,
 L28 pollution, and over-fishing. Climate change affects
 L29 salmon throughout their life stages and poses an
 L30 additional stress. As more winter precipitation falls
 L31 as rain rather than snow, higher winter stream-
 L32 flows scour streambeds, damaging spawning nests
 L33 and washing away incubating eggs. Earlier peak
 L34 streamflows flush young salmon from rivers to
 L35 estuaries before they are physically mature enough
 L36 for the transition, increasing a variety of stresses
 L37 including the risk of being eaten by predators.
 L38 Lower summer streamflows and warmer water
 L39 temperatures create less favorable summer stream
 L40 conditions for salmon and other cold-water fish
 L41 species in many parts of the Northwest. In addition,
 L42 diseases and parasites that infect salmon tend to
 L43 flourish in warmer water. Climate change also im-
 L44 pacts the ocean environment, where salmon spend
 L45 several years of their lives. Historically, warm
 L46 periods in the coastal ocean have coincided with
 L47 relatively low abundances of salmon, while cooler
 L48 ocean periods have coincided with relatively high
 L49 salmon numbers.
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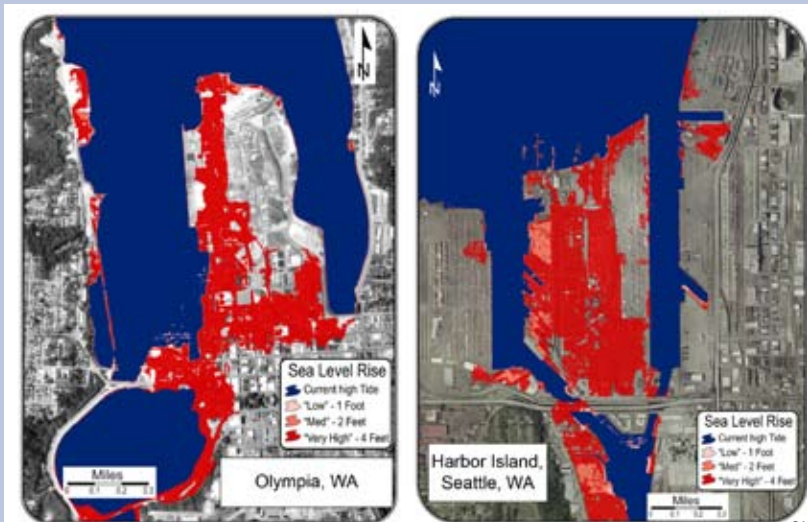


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 R17 Most wild Pacific salmon populations are extinct or
 R18 imperiled in 56 percent of their historical range in
 R19 the Northwest and California¹⁸, and populations are
 R20 down more than 90 percent in the Columbia River
 R21 system. Many species are listed as either threat-
 R22 ened or endangered under the Federal Endangered
 R23 Species Act. Studies suggest that about one-third of
 R24 the current habitat for the Northwest's salmon and
 R25 other cold-water fish will no longer be suitable for
 R26 them by the end of this century as key temperature
 R27 thresholds are exceeded. Because climate change
 R28 impacts on their habitat are projected to be nega-
 R29 tive, climate change is expected to hamper efforts
 R30 to restore depleted salmon populations.

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 R33 **Sea-level rise will result in increased**
 R34 **erosion along vulnerable coastlines.**
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R36 Climate change is projected to exacerbate many
 R37 of the stresses and hazards currently facing the
 R38 coastal zone. Sea-level rise will increase erosion of
 R39 the Northwest coast and cause the loss of beaches
 R40 and significant coastal land areas. Among the most
 R41 vulnerable parts of the coast are the heavily popu-
 R42 lated south Puget Sound region, which includes
 R43 the cities of Olympia, Tacoma, and Seattle, Wash-
 R44 ington. Some climate models project changes in
 R45 atmospheric pressure patterns that suggest a more
 R46 southwesterly direction of future winter winds.
 R47 Combined with higher sea levels, this would accel-
 R48 erate coastal erosion all along the Pacific Coast.
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Northwest Cities at Risk to Sea-Level Rise



Highly populated coastal areas throughout Puget Sound, Washington, are vulnerable to sea-level rise. The maps show regions of Olympia and Harbor Island (both located in Puget Sound) that are likely to be lost to sea-level rise by the end of this century based on moderate and high estimates.

Sea-level rise in the Northwest (as elsewhere) is determined by global rates of sea-level rise, changes in coastal elevation associated with local vertical movement of the land, and atmospheric dynamics that influence wind-driven “pile up” of sea level along the coast. A mid-range estimate of relative sea-level rise for the Puget Sound basin is about 13 inches by 2100. However, higher levels of up to 50 inches by 2100 in more rapidly subsiding portions of the basin are also possible given the large uncertainties about accelerating rates of ice melt from Greenland and Antarctica in recent years²⁰.

An additional concern is landslides on coastal bluffs. The projected heavier winter rainfall suggests an increase in saturated soils and, therefore, an increased number of landslides. Increased frequency and/or severity of landslides is expected to be especially problematic in areas where there has been intensive development on unstable slopes. Within Puget Sound, the cycle of beach erosion and bluff landslides will be exacerbated by sea-level rise, increasing beach erosion, and decreasing slope stability.

Adaptation: Improved Planning to Cope with Future Changes

States, counties, and cities in the Northwest are beginning to develop strategies to adapt to climate change. In 2007, Washington State convened stakeholders to develop adaptation strategies for water, agriculture, forests, coasts, infrastructure, and human health. Recommendations included improved drought planning, improved monitoring of diseases and pests, incorporating sea-level rise in coastal planning, and public education. An implementation strategy is under development.

In response to concerns about increasing flood risk, King County, Washington, approved plans in 2007 to fund repairs to the county’s aging levee system. The county also will replace more than 57 “short-span” bridges with wider span structures that allow more debris and floodwater to pass underneath rather than backing up and causing the river to flood. The county has begun incorporating porous concrete and rain gardens into road projects to manage the effects of stormwater runoff during heavy rains, which are increasing as climate changes. King County also has published an adaptation guidebook that is becoming a model that other local governments can refer to in order to organize adaptation actions within their municipal planning processes.

Concern about sea-level rise in Olympia, Washington, contributed to the city’s decision to relocate its primary drinking water source from a low-lying surface water source to wells on higher ground. The city adjusted its plans for construction of a new City Hall to locate the building in an area less vulnerable to sea-level rise than the original proposed location. The building’s foundation also was raised by 1 foot.