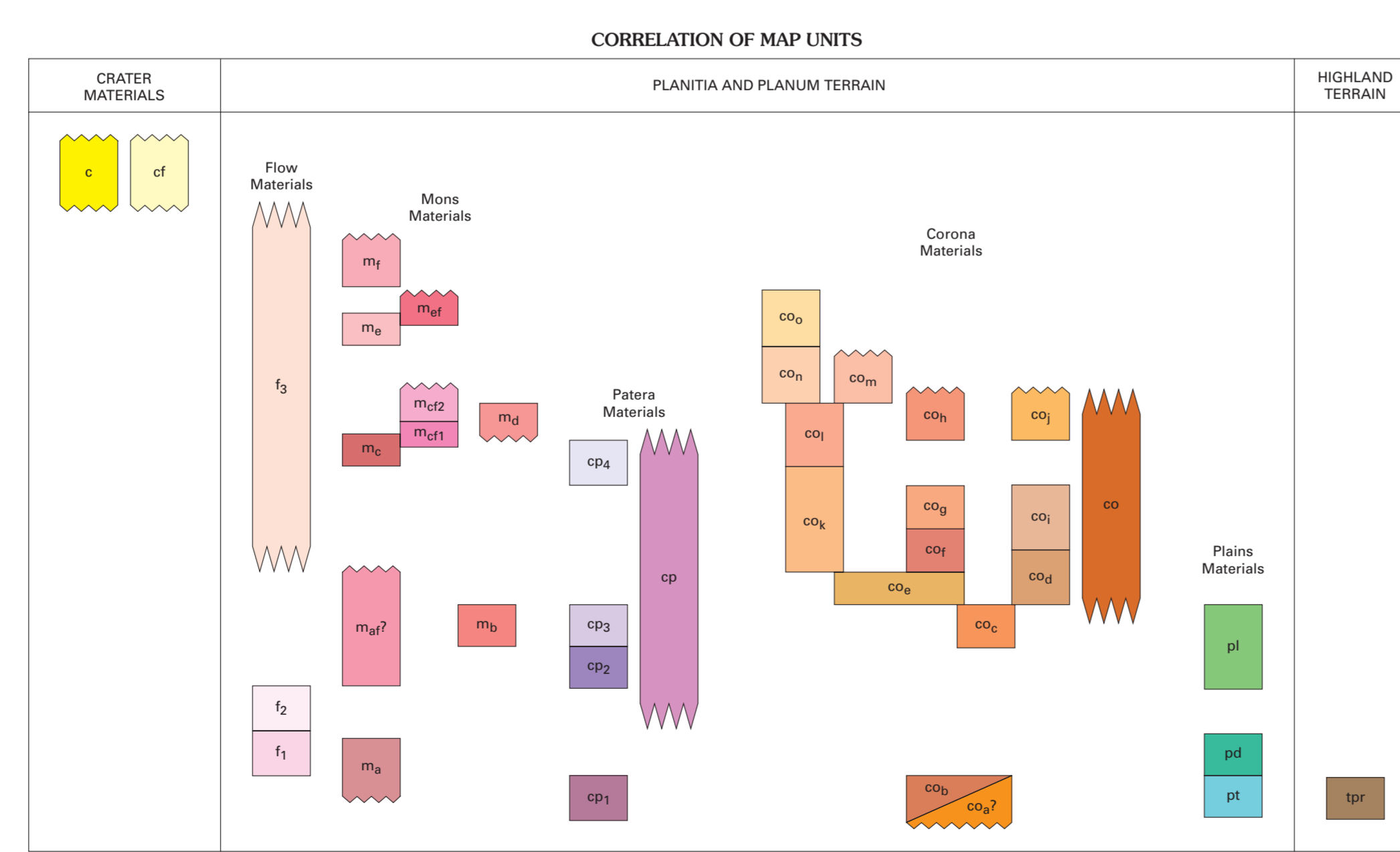


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DESCRIPTION OF MAP UNITS
Descriptions of material units from SAR images of Magellan data reflect surface scattering properties varying with polarization, incidence angle of transmitted radar waves, surface roughness, and surface physical electric properties; therefore, unit descriptions of relative brightness do not translate to those seen in other planetary surface images acquired in the visible wavelengths. Relative ages for units established by stratigraphic relations; the cumulative densities of craters are too low on Venus for dating at the 1:5,000,000 scale. Topographic measurements approximated from 500-m contour map (fig. 1B). Radar correlation and quantitative surface properties available for some units (fig. 9, table 1).

CRATER MATERIALS
Crater material—Material of radar-bright, bowl-shaped craters having complete rim crests, steep walls, ejecta well preserved and in all cases superposed on other materials. Some craters may have flat radar-dark floors; some craters may be deformed by subsequent fracturing. Interpretation: Young impact craters; impact melt present on some crater floors.
Outflow deposit—Radar-bright, digitate lobes trailing away from ejecta of impact craters. Interpretation: Impact-melt flows.

PLANTIA AND PLANUM TERRAINS
Unit 3—Occurs locally in patches superposed on material units as old as corona unit 1 and flow unit 2.
Unit 2—Occurs over much of the map area underlying corona units c-o, shows one embayment relation with flow unit 1 at lat 20.5° S, long 264.5° E.
Unit 1—Embays Phobos Regio tessera material and tessera-embaying plains material on east part of map area. Overlain by flow unit 2 and other mons, cones, and patera units.
Mons materials—Consist of interdigitate tongues of varied radar reflectance bound by lobate edges that can be traced to cone- or dome-shaped mountains. Interpretation: Flank and radial lavas associated with and composing shield volcanoes.

CRATER MATERIALS
Unit f—Traceable to 2-km-high cone (lat 13° S, long 262°) containing central pit. Flows embay Phobos Regio tessera material and corona units a and o.
Unit e—Occurs on two cone-shaped mounds (near center of map area) having heights between 0.5 and 2 km. Superposed on corona units e, l, and n.
Unit d—Forms a fractured, 1.4-km-high mound at lat 16° S, long 253°; embayed by corona material n.

The Magellan Mission
The Magellan spacecraft orbited Venus from August 10, 1990, until it plunged into the venusian atmosphere on October 12, 1994. Magellan had the objectives of (1) improving knowledge of the geologic processes, surface properties, and geologic history of Venus by analysis of radar characteristics, topography, and morphology and (2) improving knowledge of the geophysics of Venus by analysis of venusian gravity.
The Magellan spacecraft carries a 12.6-m radar system to map the surface of Venus. The transmitter and receiver systems were used to collect three datasets: synthetic aperture radar (SAR) images of the surface, passive microwave thermal emission observations, and measurements of the backscattered power at small angles of incidence, which were processed to yield altimetric data. Radar imaging and altimetric and radiometric mapping of the venusian surface were done in mission cycles 1, 2, and 3, from September 1990 until September 1992. Ninety-eight percent of the surface was mapped with radar resolution of approximately 120 meters. The SAR observations were projected to a 75-m nominal horizontal resolution; these full-resolution data compose the image base used in geologic mapping. The primary polarization mode was horizontal transmit, horizontal receive (HH), but additional data for selected areas were collected for the vertical polarization sense. Incidence angles varied from about 20° to 45°.
High-resolution Doppler tracking of the spacecraft was done from September 1992 through October 1994 (mission cycles 4, 5, 6). High-resolution gravity observations from about 950 orbits were obtained between September 1992 and May 1993, while Magellan was in an elliptical orbit with a periastron near 175 kilometers and an apoastron near 8,000 kilometers. Observations from an additional 1,500 orbits were obtained following orbit circularization in mid-1993. These data exist as a 75 × 75 harmonic field.
Magellan Radar Data
Radar backscatter power is determined by the morphology of the surface at a broad range of scales and by the intrinsic reflectivity, or dielectric constant, of the material. Topography at scales of several meters and larger can produce quasi-specular echoes, with the strength of the return greatest when the local surface is perpendicular to the incident beam. This type of scattering is most important at very small angles of incidence, because natural surface generally have low slope tilted facets at high angles. The exception is in areas of steep slopes, such as ridges or rift zones, where favorably tilted terrain can produce very bright signatures in the radar image. For most other areas, diffuse echoes from roughness at scales comparable to the radar wavelength are responsible for variations in the SAR return. In either case, the echo strength is also modulated by the reflectivity of the surface material. The density of the upper few wavelengths of the surface can have a significant effect. Low-density layers, such as crater ejecta or volcanic ash, can absorb the incident energy and produce a lower observed echo. On Venus, a rapid increase in reflectivity exists at a certain critical elevation, above which high-dielectric minerals or coatings are thermodynamically stable. This effect leads to very bright SAR echoes from virtually all areas above that critical elevation.
The measurements of passive thermal emission from Venus, though of much lower spatial resolution than the SAR data, are more sensitive to changes in the dielectric constant of the surface than to roughness. As such, they can be used to augment studies of the surface and to discriminate between roughness and reflectivity effects. Observations of the near-radar backscatter power, collected using a separate smaller antenna on the spacecraft, were modeled using the Hagfors expression for echoes from gently undulating surfaces to yield estimates of planetary radius, Fresnel reflectivity, and root-mean-square (rms) slope. The topographic data produced by this technique have horizontal footprint sizes of about 10 km near periastron and a vertical resolution of approximately 100 m. The Fresnel reflectivity data provide a comparison to the emissivity maps, and the rms slope parameter is an indicator of the surface tilt, which contributes to the quasi-specular scattering component.

GEOLOGIC/GEOMORPHIC MAP OF THE GALINDO QUADRANGLE (V-40), VENUS

By
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