

1315 **Chapter 2. A Description and Evaluation of Hydrologic**
1316 **and Climate Forecast and Data Products that Support**
1317 **Decision-Making for Water Resource Managers**

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1319 **Convening Lead Author:** Nathan Mantua, Climate Impacts Group, Univ. of Washington

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1321 **Lead Authors:** Michael D. Dettinger, U.S. Geological Survey, Scripps Institution of

1322 Oceanography; Thomas C. Pagano, National Water and Climate Center, NRCS/USDA;

1323 Andrew W. Wood, 3Tier Group / Dept. of Civil and Environmental Engineering, Univ. of

1324 Washington; Kelly Redmond, Western Regional Climate Center, Desert Research

1325 Institute

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1327 **Contributing Author:** Pedro Restrepo, NOAA

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1337 **KEY FINDINGS**

1338 There are a wide variety of climate and hydrologic data and forecast products currently
1339 available for use by decision-makers in the water resources sector. However, the use of
1340 official seasonal to interannual (SI) climate and hydrologic forecasts generated by federal
1341 agencies remains limited in the water resources sector. Forecast skill, while recognized as
1342 just one of the barriers to the use of SI climate forecast information, remains a primary
1343 concern among forecast producers and users. Simply put, there is no incentive to use SI
1344 climate forecasts when they are believed to provide little additional skill to existing
1345 hydrologic and water resource forecast approaches. Not surprisingly, there is much
1346 interest in improving the skill of hydrologic and water resources forecasts. Such
1347 improvements can be realized by pursuing several research pathways, including:

- 1348 • Improved monitoring and assimilation of real-time hydrologic observations in
1349 land surface hydrologic models that leads to improved estimates for initial
1350 hydrologic states in forecast models;
- 1351 • Increased accuracy in SI climate forecasts; and,
- 1352 • Improved bias corrections in existing forecast.

1353

1354 Another aspect of forecasts that serves to limit their use and utility is the challenge in
1355 interpreting forecast information. For example, from a forecast producer's perspective
1356 confidence levels are explicitly and quantitatively conveyed by the range of possibilities
1357 described in probabilistic forecasts. From a forecast user's perspective, probabilistic
1358 forecasts are not always well understood or correctly interpreted. Although structured
1359 user testing is known to be an effective product development tool, it is rarely done.

1360 Evaluation should be an integral part of improving forecasting efforts, but that evaluation
1361 should be extended to factors that encompass use and utility of forecast information for
1362 stakeholders. In particular, very little research is done on effective seasonal forecast
1363 communication. Instead, users are commonly engaged only near the end of the product
1364 development process.

1365

1366 Other barriers to the use of SI climate forecasts in water resources management have
1367 been identified and those that relate to institutional issues and aspects of current forecast
1368 products are discussed in chapters 3 and 4 of this report.

1369

1370 Pathways for expanding the use and improving the utility of data and forecast products to
1371 support decision-making in the water resources sector are currently being pursued at a
1372 variety of spatial and jurisdictional scales in the United States. These efforts include:

- 1373 • An increased focus on developing forecast evaluation tools that provide users
1374 with opportunities to better understand forecast products in terms of their
1375 expected skill and applicability;
- 1376 • Additional efforts to explicitly and quantitatively link SI climate forecast
1377 information with SI hydrologic and water supply forecasting efforts;
- 1378 • An increased focus on developing new internet-based tools for accessing and
1379 customizing data and forecast products to support hydrologic forecasting and
1380 water resources decision-making; and,
- 1381 • Further improvements in the skill of hydrologic and water supply forecasts.

1382

1383 Many of these pathways are currently being pursued by the federal agencies charged with
1384 producing the official climate and hydrologic forecast and data products for the United
1385 States, but there is substantial room for increasing these activities.

1386

1387 An additional important finding is that recent improvements in the use and utility of data
1388 and forecast products related to water resources decision-making have come with an
1389 increased emphasis on these issues in research funding agencies through programs like
1390 GEWEX, NOAA's RISA, SARP, TRACS and CPPA programs. Sustaining and
1391 accelerating future improvements in the use and utility of official data and forecast
1392 products in the water resources sector rests in part on sustaining and expanding federal
1393 support for programs focused on improving the skill in forecasts, increasing the access to
1394 data and forecast products, and fostering sustained interactions between forecast
1395 producers and consumers.

1396

1397 **2.1 INTRODUCTION**

1398 In the past, water resource managers relied heavily on observed hydrologic conditions
1399 such as snowpack and soil moisture to make seasonal to interannual (SI) water supply
1400 forecasts to support management decisions. Within the last decade, researchers have
1401 begun to link SI climate forecasts with hydrologic models (*e.g.*, Kim *et al.*, 2000,
1402 Kyriakidis *et al.*, 2001) or statistical distributions of hydrologic parameters (*e.g.*,
1403 Dettinger *et al.*, 1999, Sankarasubramanian and Lall 2003) to improve hydrologic and
1404 water resources forecasts. Efforts to incorporate SI climate forecasts into water resources
1405 forecasts have been prompted in part by our growing understanding of the effects of

1406 global-scale climate phenomena, like El Niño Southern Oscillation (ENSO), on U.S.
1407 climate, and the expectation that SI forecasts of hydrologically-significant climate
1408 variables like precipitation and temperature provide a basis for predictability that is not
1409 currently being exploited. To the extent that climate variables like temperature and
1410 precipitation can be forecasted seasons in advance, hydrologic and water-supply forecasts
1411 can also be made skillfully well before the end, or even beginning, of the water year¹.

1412

1413 This chapter focuses on a description and evaluation of hydrologic and climate forecast
1414 and data products that support decision-making for water resource managers. Because the
1415 focus of this CCSP product is on using SI forecasts and data for decision-support in the
1416 water resources sector, we frame this chapter around key forecast and data products that
1417 contribute towards improved hydrologic and water supply forecasts. As a result, this
1418 product does not contain a comprehensive review and assessment of the entire national SI
1419 climate and hydrologic forecasting effort. In addition, the reader should note that, even
1420 today, hydrologic and water supply forecasting efforts in many places are still not
1421 inherently linked with the SI climate forecasting enterprise.

1422

1423 Surveys identify a variety of barriers to the use of climate forecasts (Pulwarty and
1424 Redmond, 1997; Callahan *et al.*, 1999;. Hartmann *et al.*,2002), but insufficient accuracy
1425 is always mentioned as a barrier. It is also well established that an accurate forecast is, in
1426 and of itself, not sufficient to make it useful or usable for decision-making in
1427 management applications (see Table 2.1). Chapters 3 and 4 provide extensive reviews,

¹ The *water year*, or hydrologic year, is October 1st through September 30th. This reflects the natural cycle in many hydrologic parameters such as the seasonal cycle of evaporative demand, and of the snow accumulation, melt, and runoff periods in many parts of the US.

1428 case studies, and analyses that provide insights into pathways for lowering or overcoming
1429 barriers to the use of SI climate forecasts in water resources decision-making.

1430

1431 It is almost impossible to discuss the perceived value of forecasts without also discussing
1432 issues related to forecast skill. Many different criteria have been used to evaluate forecast
1433 skill (see Wilks, 1995 for a comprehensive review). Some measures focus on aspects of
1434 deterministic skill (*e.g.*, correlations between predicted and observed seasonally averaged
1435 precipitation anomalies), while many others are based on categorical forecasts (*e.g.*,
1436 Heidke skill scores for categorical forecasts of “wet,” “dry,” or “normal” conditions). The
1437 most important measures of skill vary with different perspectives. For example,
1438 Hartmann *et al.*, (2002) argue that forecast performance criteria based on “hitting” or
1439 “missing” associated observations offer users conceptually easy entry into discussions of
1440 forecast quality. In contrast, some research scientists and water supply forecasters may be
1441 more interested in correlations between the ensemble average of predictions and observed
1442 measures of water supply like seasonal runoff volume.

1443

1444 Forecast skill remains a primary concern among many forecast producers and users. Skill
1445 in hydrologic forecast systems derives from various sources, including the quality of the
1446 simulation models used in forecasting, the ability to estimate the initial hydrologic state
1447 of the system, and the ability to skillfully predict the statistics of future weather over the
1448 course of the forecast period. Despite the significant resources expended to improve SI
1449 climate forecasts over the past 15 years, few water resource related agencies have been

1450 making quantitative use of climate forecast information in their water supply forecasting
 1451 efforts (Pulwarty and Redmond 1997; Callahan *et al.*, 1999).

1452

1453

1454 **Table 2.1 Barriers to the use of climate forecasts and information for resource managers in the**
 1455 **Columbia River Basin**

1456 (Reproduced from Pulwarty and Redmond, 1997).

- a. Forecasts not “accurate” enough.
- b. Fluctuation of successive forecasts (“waffling”).
- c. The nature of what a forecast is, and what is being forecast (*e.g.*, types of El Niño and La Niña impacts, non-ENSO events, what are “normal” conditions?).
- d. Nonweather/climate factors are deemed to be more important (*e.g.*, uncertainty in other arenas, such as freshwater and ocean ecology [for salmon productivity]).
- e. Low importance is given to climate forecast information because its role is unclear or impacts are not perceived as important enough to commit resources.
- f. Other constraints deny a flexible response to the information (*e.g.*, meeting flood control or Endangered Species Act requirements).
- g. Procedures for acquiring knowledge and making and implementing decisions, which incorporate climate information, have not been clearly defined.
- h. Events forecast may be too far in the future for a discrete action to be engaged.
- i. Availability and use of locally specific information may be more relevant to a particular decision.
- j. “Value” may not have been demonstrated by a credible reliable organization or competitor.
- k. Desired information not provided (*e.g.*, number of warm days, regional detail).
- l. There may be competing forecasts or other conflicting information.
- m. Lack of “tracking” information; does the forecast appear to be verifying?
- n. History of previous forecasts not available. Validation statistics of previous forecasts not available.

1457

1458 In Section 2.2 of this chapter, we review hydrologic data and forecasts products. Section
 1459 2.3 provides a parallel discussion of the climate data and forecast products that support
 1460 hydrologic and water supply forecasting efforts in the United States. In Section 2.4, we
 1461 provide a more detailed discussion of pathways for improving the skill and utility in
 1462 hydrologic and climate forecasts and data products.

1463

1464 Section 2.5 contains a brief review of operational considerations and efforts to improve
 1465 the utility of forecast and data products through efforts to improve the forecast evaluation
 1466 and development process. These efforts include cases in which forecast providers and

1467 users have been engaged in sustained interactions to improve the use and utility of
1468 forecast and data products, and have led to many improvements and innovations in the
1469 data and forecast products generated by national centers. In recent years, a small number
1470 of water resource agencies have also developed end-to-end forecasting systems that
1471 utilize climate forecasts to directly inform hydrologic and water resources forecasts.

1472

1473 BOX 2.1: Agency Support

1474

1475 Federal support for research supporting improved hydrologic forecasts and applications through the use of
1476 climate forecasts and data has received increasing emphasis since the mid-1990s. The World Climate
1477 Research Program's Global Energy and Water Cycle Experiment (GEWEX) was among the first attempts
1478 to integrate hydrology/land surface and atmosphere models in the context of trying to improve hydrologic
1479 and climate predictability.

1480

1481 There have been two motivations behind this research: understanding scientific issues of land surface
1482 interactions with the climate system, and the development or enhancement of forecast applications, *e.g.*, for
1483 water, energy and hazard management. Early on, these efforts were dominated by the atmospheric (and
1484 related geophysical) sciences.

1485

1486 In the past, only two U.S. programs have been very relevant to hydrologic prediction: the NOAA Climate
1487 Prediction Program for the Americas (CPPA) and NOAA predecessors GEWEX Continental-scale
1488 International Project (GCIP) and GEWEX Americas Prediction Project (GAPP) and the NASA Terrestrial
1489 Hydrology Program. The hydrologic prediction and water management focus of NOAA and NASA has
1490 slowly expanded over time. Presently, the NOAA Climate Dynamics and Experimental Prediction (CDEP),
1491 Transition of Research Applications to Climate Services (TRACS) and Sectoral Applications Research
1492 Program (SARP) programs, and the Water Management program within NASA, have put a strong
1493 emphasis on the development of both techniques and community linkages for migrating scientific advances
1494 in climate and hydrologic prediction into applications by agencies and end use sectors. The longer-standing
1495 NOAA Regional Integrated Sciences and Assessments (RISA) program has also contributed to improved
1496 use and understanding of climate data and forecast products in water resources forecasting and decision-
1497 making. Likewise, the recently initiated postdoctoral fellowship program under the Predictability,
1498 Predictions, and Applications Interface (PPAI) panel of U.S. CLIVAR aims to grow the pool of scientists
1499 qualified to transfer advances in climate science and climate prediction into climate-related decision
1500 frameworks and decision tools.

1501

1502 Still, these programs are not well funded in comparison to current federally funded science-focused
1503 initiatives, and are only just beginning to make inroads into the vast arena of effectively increasing the use
1504 and utility of climate and hydrologic data and forecast products.

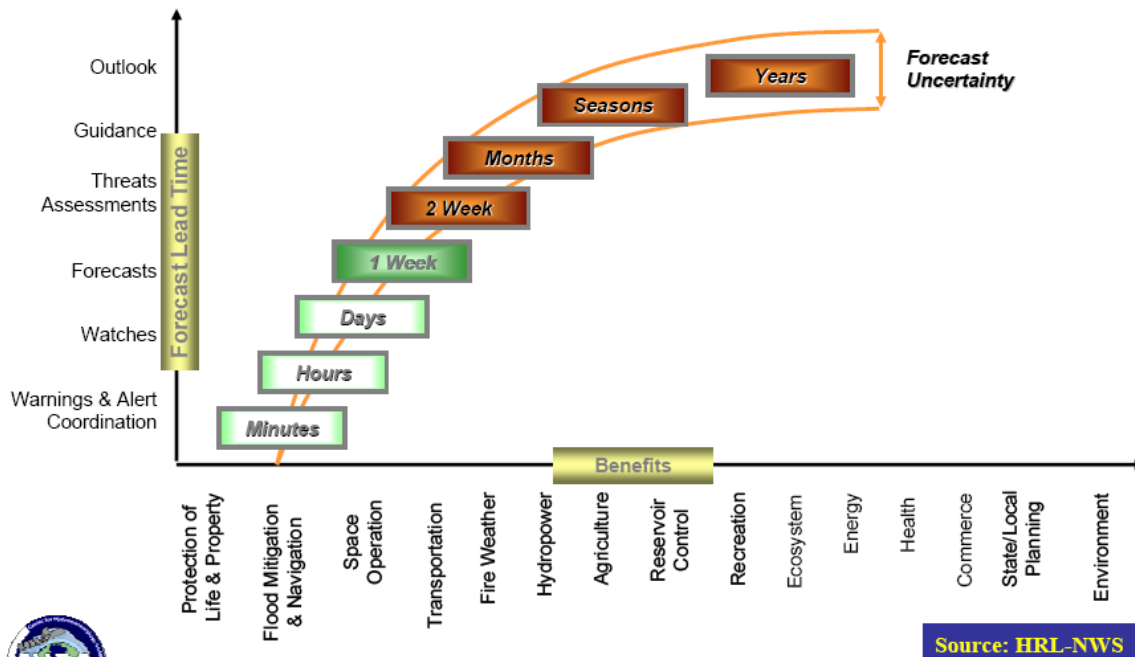
1505

1506 end BOX 2.1

1507

1508 2.2 HYDROLOGIC AND WATER RESOURCES: MONITORING AND**1509 PREDICTION**

1510 The uses of hydrologic monitoring and prediction products, and specifically those that are
 1511 relevant for water, hazard and energy management vary depending on the forecast lead
 1512 time (Figure 2.1). The shortest climate and hydrologic lead time forecasts, from minutes
 1513 to hours, are applied to such uses as warnings for floods and extreme weather, wind
 1514 power scheduling, aviation, recreation, and wild fire response management. In contrast, at
 1515 lead times of years to decades predictions are used for strategic planning purposes rather
 1516 than operational management of resources. At SI lead times, climate and hydrologic
 1517 forecast applications span a wide range that includes the management of water, fisheries,
 1518 hydropower and agricultural production, navigation and recreation. Table 2.1 lists aspects
 1519 of forecast products at these time scales that are relevant to decision-makers.
 1520



1521

Source: HRL-NWS

1522 **Figure 2.1** The correspondence of climate and hydrologic forecast lead time to user sectors in which
 1523 forecast benefits are realized (from HRL-NWS). The focus of this product is on climate and hydrologic
 1524 forecasts with lead times greater than 2 weeks and up to approximately one year.
 1525

1526 **2.2.1 Prediction Approaches**

1527 The primary climate and hydrologic prediction approaches used by operational and
1528 research centers fall into four categories: statistical, dynamical, statistical-dynamical
1529 hybrid, and consensus. The first three approaches are objective in the sense that the inputs
1530 and methods are formalized, outputs are not modified on an ad hoc basis, and the
1531 resulting forecasts are potentially reproducible by an independent forecaster using the
1532 same inputs and methods. The fourth major category of approach, which might also be
1533 termed blended knowledge, requires subjective weighting of results from the other
1534 approaches. These types of approaches are discussed in Box 2.2.

1535

1536 BOX 2.2: Forecast Approaches

1537

1538 *Dynamical:* Computer models designed to represent the physical features of the oceans, atmosphere and
1539 land surface, at least to the extent possible given computational constraints, form the basis for dynamical
1540 predictions. These models have at their core a set of physical relationships describing the interactions of the
1541 Earth's energy and moisture states. Inputs to the models include estimates of the current moisture and
1542 energy conditions needed to initialize the state variables of the model (such as the moisture content of an
1543 atmospheric or soil layer), and of any physical characteristics (called parameters -- one example is the
1544 elevation of the land surface) that must be known to implement the relationships in the model's physical
1545 core. In theory, the main advantage of dynamical models is that influence of any one model variable on
1546 another is guided by the laws of nature as we understand them. As a result, the model will correctly
1547 simulate the behavior of the earth system even under conditions that may not have occurred in the period
1548 during which the model is verified, calibrated and validated. The primary disadvantages of dynamical
1549 models, however, are that their high computational and data input demands require them to approximate
1550 characteristics of the Earth system in ways that may compromise their realism and therefore performance.
1551 For example, the finest computational grid resolution that can be practically achieved in most atmospheric
1552 models (on the order of 100~200 km per cell) is still too coarse to support a realistic representation of
1553 orographic effects on surface temperature and precipitation. Dynamical hydrologic models can be
1554 implemented at much finer resolutions (down to 10 meters per cell, for catchment-scale models) because
1555 they are typically applied to much smaller geographic domains than are atmospheric models. While there
1556 are many aspects that distinguish one model from another, only a subset of those (listed in Table 1.1) is
1557 appreciated by the forecast user, as opposed to the climate modeler, and is relevant in describing the
1558 dynamical forecast products.

1559

1560 *Statistical:* Statistical forecast models use mathematical models to relate observations of an earth system
1561 variable that is to be predicted to observations of one or more other variables (and/or of the same variable at
1562 a prior time) that serve as predictors. The variables may describe conditions at a point location (*e.g.*, flow
1563 along one reach of a river) or over a large domain, such as sea surface temperatures along the equator. The
1564 mathematical models are commonly linear relationships between the predictors and the predictand, but also
1565 may be formulated as more complex non-linear systems.

1566

1567 Statistical models are often preferred for their computational ease relative to dynamical models. In many
1568 cases, statistical models can give equal or better performance to dynamical models due in part to the
1569 inability of dynamical models to represent fully the physics of the system (often as a result of scale or data

1570 limitations), and in part to the dependence of predictability in many systems on predominantly linear
 1571 dynamics (Penland and Magorian, 1993; van den Dool, 2007). The oft-cited shortcomings of statistical
 1572 models, on the other hand, include their lack of representation of physical causes and effects, which in
 1573 theory compromise their ability to respond to unprecedented events in a fashion that is consistent with the
 1574 physical constraints of the system. In addition, statistical models may require a longer observational record
 1575 for “training” than dynamical models, which are helped by their physical structure.
 1576

1577 **Objective hybrids:** Statistical and dynamical tools can be combined using objective approaches. A primary
 1578 example is a weighted merging of the tools’ separate predictions into a single prediction (termed an
 1579 objective consolidation; van den Dool, 2007). A second example is a tool that has dynamical and statistical
 1580 subcomponents, such as a climate prediction model that links a dynamical ocean submodel to a statistical
 1581 atmospheric model. A distinguishing feature of these hybrid approaches is that an objective method exists
 1582 for linking the statistical and dynamical schemes so as to produce a set of outputs that are regarded as
 1583 “optimal” relative to the prediction goals. This objectivity is not preserved in the next consensus approach.
 1584

1585 **Blended Knowledge or Subjective consensus:** Some forecast centers release operational predictions, in
 1586 which expert judgment is subjectively applied to modify or combine outputs from prediction approaches of
 1587 one or more of the first three types, thereby correcting for perceived errors in the objective approaches to
 1588 form a prediction that has skill superior to what can be achieved by objective methods alone. The process
 1589 by which the NOAA Climate Prediction Center (CPC) and International Research Institute for Climate and
 1590 Society (IRI) constructs their monthly and seasonal outlooks for example, includes subjective weighting of
 1591 the guidance provided by different climate forecast tools. The weighting is often highly sensitive to recent
 1592 evolution and current state of the tropical ENSO, but other factors like decadal trends in precipitation and
 1593 surface temperature also have the potential to influence the final official climate forecasts.
 1594

1595 end BOX 2.2
 1596

1597 **Table 2.1 Aspects of forecast products that are relevant to users**

| Forecast Product Aspect | Description / Examples |
|--|---|
| Forecast product variables | Precipitation, temperature, humidity, windspeed, atmospheric pressure |
| Forecast product spatial resolution | Grid cell longitude by latitude, climate division |
| Domain | Watershed, river basin, regional, national, global |
| Product time step (temporal resolution) | Hourly, sub-daily, daily, monthly, seasonal |
| Range of product lead times | 1 to 15 days, 1 to 13 months |
| Frequency of forecast product update | every 12 hours, every month |
| Lag of forecast product update | The length of time from the forecast initialization time before forecast products are available: <i>e.g.</i> , 2 hours for a medium range forecast, one day for a monthly to seasonal forecast |
| Existence of historical climatology | Many users require a historical climatology showing forecast model performance to use in bias-correction, downscaling, and/or verification. |
| Deterministic or probabilistic | Deterministic forecasts have a single prediction for each future lead time. Probabilistic forecasts frame predicted values within a range of uncertainty, and consist either of an ensemble of forecast sequences spanning all lead times, or of a distinct forecast distribution for each future lead time. |
| Availability of skill / accuracy information | Published or otherwise available information about the performance of forecasts is not always available, particularly for forecasts that are steadily evolving. In principle, the spread of probabilistic forecasts contains such information about the median of the forecast; but the skill characteristics pertaining to the spread of the forecast are not usually available. |

1598

1599 Other aspects of dynamical prediction schemes related to model physical and
1600 computational structure are important in distinguishing one model or model version from
1601 another. These aspects are primary indicators of the sophistication of an evolving model,
1602 relative to other models, but are not of much interest to the forecast user community.
1603 Examples include the degree of coupling of model components, model vertical
1604 resolution, cloud microphysics package, nature of data assimilation approaches, and of
1605 the data assimilated, and the ensemble generation scheme, among many other forecast
1606 system features.

1607

1608 **2.2.2 Forecast Producers and Products**

1609 Hydrologic forecasts are produced by many federal, regional, state, and local agencies, as
1610 well as by private sector companies such as utilities. In contrast to climate forecasts,
1611 hydrologic forecast products more directly target end use sectors -- *e.g.*, water, energy,
1612 natural resource or hazard management -- and are often region-specific. Prediction
1613 methods and forecast products vary from region to region and are governed by many
1614 factors, but depend in no small measure on the hydro-climatology, institutional traditions
1615 and sectoral concerns in each region. A representative sampling of typical forecast
1616 producers and products is given in Appendix A.1. Forecasting activities at the federal,
1617 state, regional, and local scales are discussed in the following subsections.

1618

1619 **2.2.2.1 Federal**

1620 The primary federal streamflow forecasting agencies at SI lead times are the NOAA
1621 National Weather Service (NWS) and the U.S. Department of Agriculture,(USDA)

1622 National Resource Conservation Service (NRCS) National Water and Climate Center
1623 (NWCC). The NWCC's four forecasters produce statistical forecasts of summer runoff
1624 volume in the western U.S. using multiple linear regression to estimate future streamflow
1625 from current observed snow water equivalent, accumulated water year precipitation,
1626 streamflow, and in some locations, using ENSO indicators such as the Niño3.4 index
1627 (Garen, 1992; Ref: Pagano and Garen, 2005). Snowmelt runoff is critical for a wide
1628 variety of uses (water supply, irrigation, navigation, recreation, hydropower,
1629 environmental flows) in the relatively dry summer season. The regression approach has
1630 been central in the NRCS since the mid-1930s, before which similar snow-survey based
1631 forecasting was conducted by a number of smaller groups. Forecasts are available to
1632 users both in the form of tabular summaries (Figure 2.2) that convey both the central
1633 tendency of the forecasts and estimates of uncertainty, and maps showing the median
1634 forecast anomaly for each river basin area for which the forecasts are operational (Figure
1635 2.3). Until 2006, the NWCC's forecasts were released once a month, near the first of the
1636 month, for summer flow periods such as April through July or April through September.
1637 In 2006, the NWCC began to develop automated daily updates to these forecasts, and the
1638 daily product is likely to become more prevalent as development and testing matures. The
1639 NWCC also has begun to explore the use of physically-based hydrologic models as a
1640 basis for forecasting, but this effort has barely begun.

1641

1642 NWCC water supply forecasts are coordinated subjectively with a parallel set of forecasts
1643 produced by the western U.S. NWS River Forecast Centers (RFCs), and with forecasts
1644 from Environment Canada's BC Hydro. The NRCS-NWS joint, official forecasts are of

1645 the subjective consensus type described earlier, meaning that the final forecast products
 1646 are subjective combinations of information from different sources, in this case objective
 1647 statistical tools (*i.e.*, regression-models informed by observed snow water equivalent,
 1648 accumulated water year precipitation, and streamflow) and model based forecast results
 1649 from the RFCs.
 1650

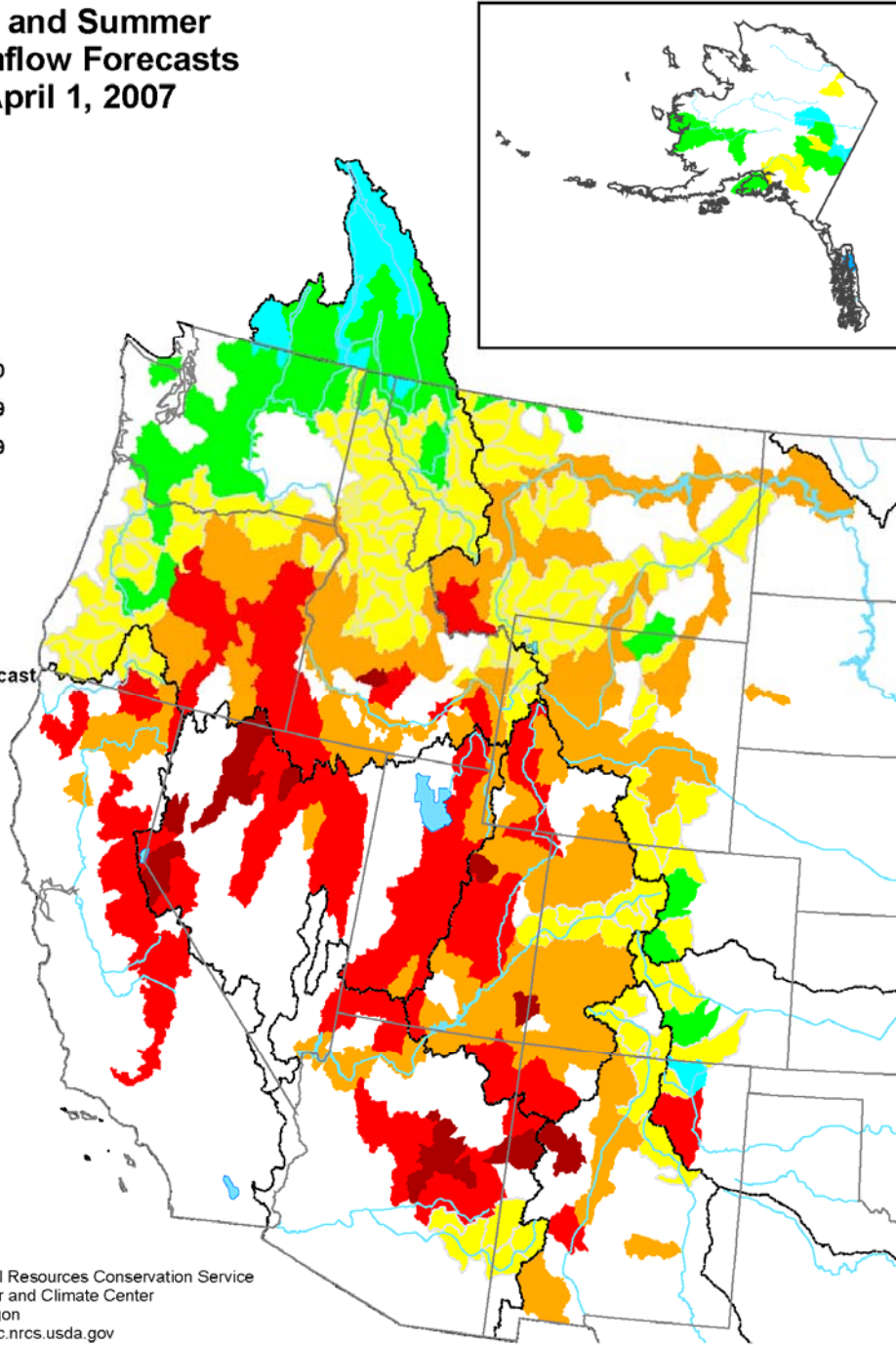
| Stream and Station | Forecast Period | Forecasts This Year | | | 30 Year | |
|---------------------------------|-----------------|---------------------|-----------------|----------|----------------------------|-------|
| | | Most Probable kaf | Reasonable %avg | Max %avg | Average '71-'00 Runoff kaf | |
| Alaska | | | | | | |
| Gulkana River | | | | | | |
| Sourdough, AK | Apr-Jul | 410 | 86 | 118 | 62 | 475 |
| Kenai River | | | | | | |
| Cooper Landing, AK | Apr-Jul | 965 | 104 | 122 | 88 | 925 |
| Ship Creek | | | | | | |
| Anchorage, AK | Apr-Jul | 45 | 78 | 102 | 57 | 58 |
| Little Susitna River | | | | | | |
| Palmer, AK | Apr-Jul | 66 | 77 | 100 | 58 | 86 |
| Talkeetna River | | | | | | |
| Talkeetna, AK | Apr-Jul | 1370 | 84 | 99 | 69 | 1630 |
| Kuskokwim River | | | | | | |
| Crooked Creek, AK | Apr-Jun | 9540 | 91 | 119 | 62 | 10500 |
| Yukon River | | | | | | |
| Eagle, AK | Apr-Jul | 38300 | 112 | 131 | 94 | 34200 |
| Stevens Village, AK | Apr-Jul | 52800 | 110 | 123 | 96 | 48200 |
| Salcha River | | | | | | |
| Salchaket, AK | Apr-Jul | 500 | 80 | 115 | 53 | 625 |
| Tanana River | | | | | | |
| Fairbanks, AK | Apr-Jul | 6900 | 97 | 112 | 84 | 7100 |
| Nenana, AK | Apr-Jul | 8290 | 92 | 107 | 77 | 9000 |
| Chena River | | | | | | |
| Two Rivers, AK | Apr-Jul | 240 | 89 | 130 | 58 | 270 |
| Little Chena River | | | | | | |
| Fairbanks, AK | Apr-Jul | 66 | 85 | 118 | 58 | 78 |
| Gold Creek | | | | | | |
| Juneau, AK | Apr-Jul | 44 | 133 | 161 | 109 | 33 |
| Saskatchewan River Basin | | | | | | |
| St. Mary River | | | | | | |
| Babb nr, MT | Apr-Sep | 400 | 89 | 103 | 74 | 450 |

1651

1652 **Figure 2.2** Example of NRCS tabular summer runoff (streamflow) volume forecast summary, showing
 1653 median (“most probable”) forecasts and probabilistic confidence intervals, as well as climatological flow
 1654 averages. Flow units are thousand-acre-feet (KAF), a runoff volume for the forecast period. This table was
 1655 downloaded from <http://www.wcc.nrcs.usda.gov/wsf/wsf.html>.
 1656

1657 The NWS surface water supply forecast program began in the 1940s in the Colorado
1658 Basin. It has since expanded to include seasonal forecasts (of volume runoff during the
1659 spring—summer snow melt period) for most of the snowmelt dominated basins important
1660 to water management in the western United States. These forecasts rely on two primary
1661 tools: Statistical Water Supply (SWS), based on multiple-linear regression, and
1662 Ensemble Streamflow Prediction (ESP), a technique based on hydrologic modeling
1663 (Schaake, 1978; Day, 1985). Results from both approaches are augmented by forecaster
1664 experience and the coordination process with other forecasting entities. In contrast to the
1665 western RFCs, RFCs in the eastern U.S. are more centrally concerned with short to
1666 medium-range flood risk and drought-related water availability out to about a three
1667 month lead time. At some eastern RFC websites, the seasonal forecast is linked only to
1668 the CPC Drought Outlook rather than an RFC-generated product (Box 2.3).
1669

**Spring and Summer
Streamflow Forecasts
as of April 1, 2007**



1670

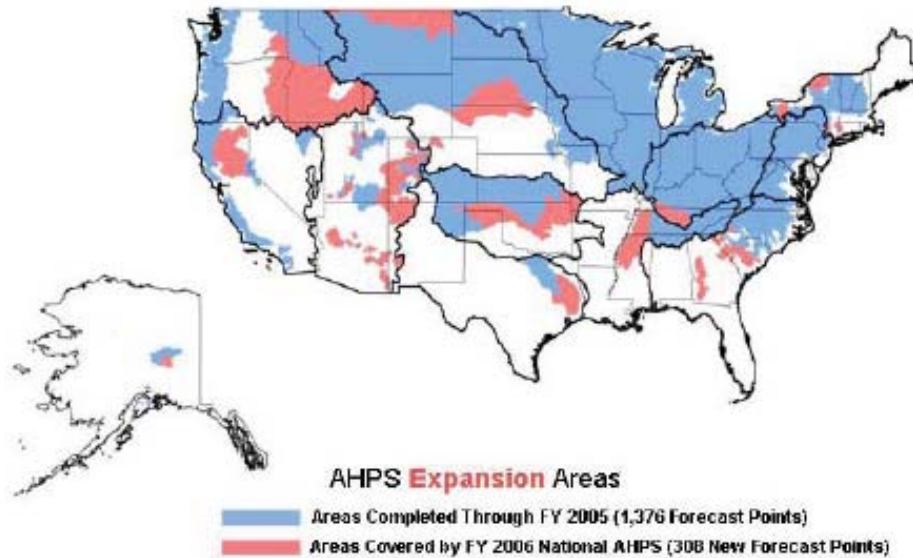
1671 **Figure 2.3** Example of NRCS spatial summer runoff (April-September streamflow) volume forecast
 1672 summary, showing median runoff forecasts as an anomaly (percent of average).
 1673

1674 The streamflow prediction services of the RFCs have a national presence, and as such are
1675 able to leverage a number of common technological elements, including models,
1676 databases and software for handling meteorological and hydrological data, and for
1677 making, assessing and disseminating forecasts; *i.e.*, website structure. Nonetheless, the
1678 RFCs themselves are regional entities with regional concerns.

1679

1680 The NWS's ESP approach warrants further discussion. In the mid 1970s, the NWS
1681 developed the hydrologic modeling, forecasting and analysis system – NWS River
1682 Forecast System (NWSRFS) – the core of which is the Sacramento soil moisture
1683 accounting scheme coupled to the Snow-17 temperature index snow model, for ESP-
1684 based prediction (Anderson, 1972, 1973; Burnash *et al.*, 1973). The ESP approach uses a
1685 deterministic simulation of the hydrologic state during a model spin-up (initialization)
1686 period leading up to the forecast start date to estimate current hydrologic conditions, and
1687 then uses an ensemble of historical meteorological sequences as model inputs (*e.g.*,
1688 temperature and precipitation) to simulate hydrology in the future (or forecast) period.
1689 Until several years ago, the RFC dissemination of ESP-based forecasts for streamflows at
1690 SI lead times was rare, and the statistical forecasts were the accepted standard. Now, as
1691 part of the NWS Advanced Hydrologic Prediction Service (AHPS) initiative, ESP
1692 forecasts are being aggressively implemented for basins across the United States (Figure
1693 2.4) at lead times from short to SI (McEnery *et al.*, 2005).

1694

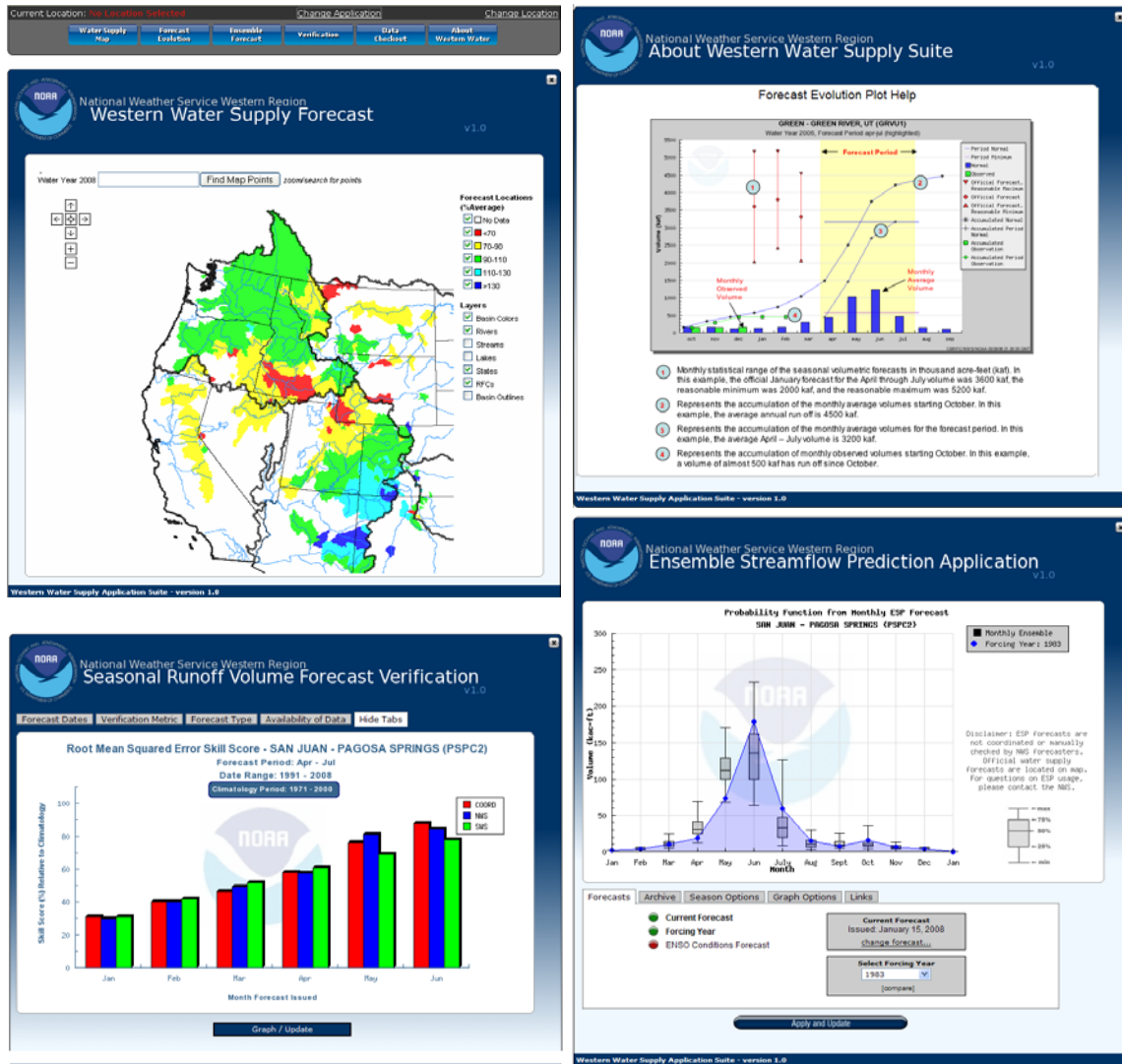


1695

1696 **Figure 2.4** Areas covered by the NWS Advanced Hydrologic Prediction Service (AHPS) initiative
 1697 (McEnery *et al.*, 2005).
 1698

1699 At the seasonal lead times, several western RFCs use graphical forecast products for the
 1700 summer period streamflow forecasts that convey the probabilistic uncertainty of the
 1701 forecasts. A unified web based suite of applications that became operational in 2008
 1702 provides forecast users with a number of avenues for exploring the RFC water supply
 1703 forecasts. For example, Figure 2.5 shows (in clockwise order from top left) (a) a western
 1704 U.S. depiction of the median water supply outlook for the RFC forecast basins, (b) a
 1705 progression of forecasts (median and bounds) during the water year together with flow
 1706 normals and observed flows; (c) monthly forecast distributions, with the option to display
 1707 individual forecast ensemble members (*i.e.*, single past years) and also select ENSO-
 1708 based categorical forecasts (ESP subsets); and (d) various skill measures, such as mean
 1709 absolute error, for the forecasts based on hindcast performance. Access to raw ensemble
 1710 member data is also provided from the same website.

1711



1712

1713 **Figure 2.5** A graphical forecast product from the NWS River Forecast Centers, showing a forecast of
 1714 summer (April—July) period streamflow on the Colorado River, Colorado-Arizona. These figures were
 1715 obtained from <http://www.nwrhc.noaa.gov/westernwater>.
 1716

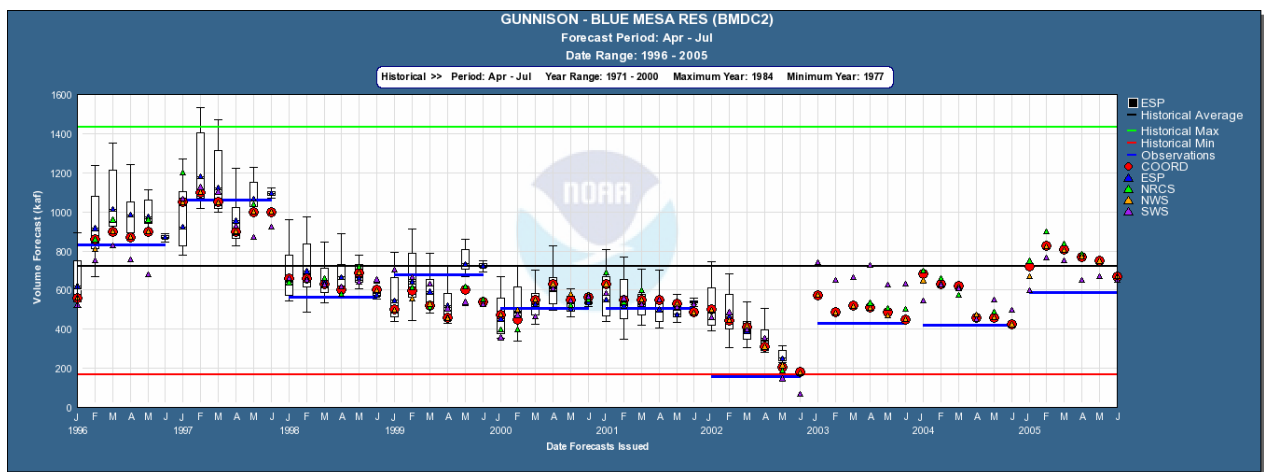
1717 The provision of a service which assists hydrologic forecast users in either customizing a
 1718 selection of ESP traces to reflect, perhaps, the users interest in past years that they
 1719 perceive as analogues to the current year, or the current ENSO state, is a notable advance
 1720 from the use of “climatological” ESP (*i.e.*, using all traces from a historical period) in the
 1721 prior ESP-related seasonal forecast products. Some western RFCs have also
 1722 experimented with using the CPC seasonal climate outlooks as a basis for adjusting the

1723 precipitation and temperature forcings used in climatological ESP, but found that the
1724 CPC outlook anomalies were generally too small to produce a distinct forecast from the
1725 climatological ESP (Hartmann *et al.*, 2002). In some RFCs, NWS statistical water supply
1726 forecasts have also provided perspective (albeit more limited) on the effect of future
1727 climate assumptions on future runoff by including results from projecting 50, 75, 100,
1728 125 and 150 percent of normal precipitation in the remaining water year. At times, the
1729 official NWS statistical forecasts have adopted such assumptions, *e.g.*, that the first
1730 month following the forecast date would contain other than 100% of expected
1731 precipitation – based on forecaster judgment and consideration of a range of factors,
1732 including ENSO state and CPC climate predictions.

1733

1734 Figure 2.6 shows the performance of summer streamflow volume forecasts from both the
1735 NWS and NRCS over a recent 10-year period; this example is also part of the suite of
1736 forecast products that the western RFC designed to improve the communication of
1737 forecast performance and provide verification information. Despite recent literature
1738 (Welles *et al.*, 2007) that has underscored a general scarcity of such information from
1739 hydrologic forecast providers, the NWS has recently codified verification approaches and
1740 developed verification tools, and is in the process of disbursing them throughout the RFC
1741 organization (NWS, 2005, “River Forecast Verification Plan”). The existence in digitized
1742 form of the retrospective archive of seasonal forecasts is critical for the verification of
1743 forecast skill. The 10-year record shown in Figure 2.6, which is longer than the record
1744 available (internally or to the public) for many public agency forecast variables, is of
1745 inadequate length for some types of statistical assessment, but is an undeniable advance

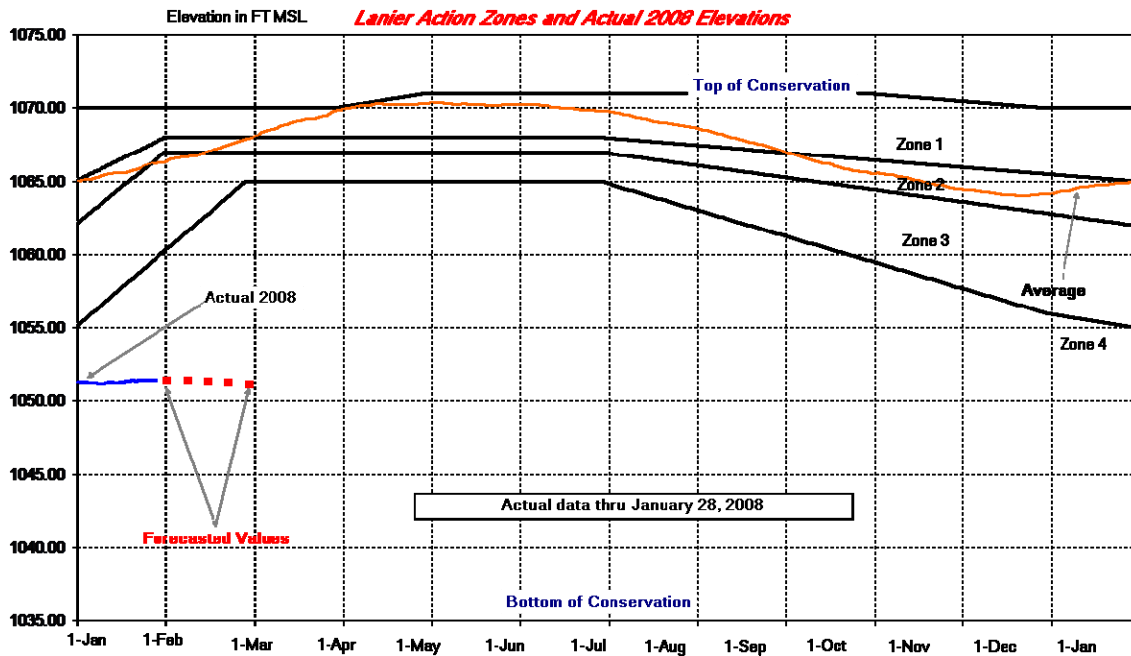
1746 in forecast communication relative to the services that were available previously. Future
 1747 development priorities include a climate change scenario application, which would
 1748 leverage climate change scenarios from IPCC or similar to produce inputs for future
 1749 water supply planning exercises. In addition, forecast calibration procedures (*e.g.*, Seo *et*
 1750 *al.*, 2006; Wood and Schaake, 2008) are being developed for the ensemble forecasts to
 1751 remove forecast biases. The current NOAA/NWS web service Internet web address is:
 1752 (<http://www.nwrfc.noaa.gov/westernwater>)
 1753



1754
 1755 **Figure 2.6** Comparing ESP and statistical forecasts from the NRCS and NWS for a recent 10-year period.
 1756 The forecasts are for summer (April—July) period streamflow on the Gunnison River, Colorado.
 1757

1758 A contrast to these probabilistic forecasts is the deterministic 5-week forecast of lake
 1759 elevation in Lake Lanier, GA, produced by the U.S. Army Corps of Engineers (USACE)
 1760 based on probabilistic inflow forecasts from the NWS southeastern RFC. Given that the
 1761 lake is a managed system and the forecast has a subseasonal lead time, the single-valued
 1762 outlook may be justified by the planned management strategy. In such a case, the lake
 1763 level is a constraint that requires transferring uncertainty in lake inflows to a different
 1764 variable in the reservoir system, such as lake outflow. Alternatively, the deterministic

1765 depiction may result from an effort to simplify probabilistic information in the
 1766 communication of the lake outlook to the public.



1767

1768 Figure 2.7 A deterministic 5-week forecast of reservoir levels in Lake Lanier, Georgia, produced by
 1769 USACE. <http://water.sam.usace.army.mil/lanfc.htm>.

1770

1771 **2.2.2.2 State and Regional**

1772 Regionally-focused agencies such as the U.S. Bureau of Reclamation (USBR), the
 1773 Bonneville Power Administration (BPA), the Tennessee Valley Authority (TVA), and the
 1774 Great Lakes Environmental Research Laboratory (GLERL) also produce forecasts
 1775 targeting specific sectors within their priority areas. Figure 2.7 shows an example of an SI
 1776 lead forecast of lake levels produced by GLERL. GLERL was among the first major
 1777 public agency to incorporate climate forecast information into operational forecasts
 1778 hydrologic and water management variables. Forecasters use coarse-scale climate
 1779 forecast information to adjust climatological probability distribution functions (PDFs) of
 1780 precipitation and temperature that are the basis for generating synthetic ensemble inputs

1781 to hydrologic and water management models, the outputs of which include lake level as
1782 shown in the figure. In this case, the climate forecast information is from the CPC
1783 seasonal outlooks (method described in Croley, 1996).

1784

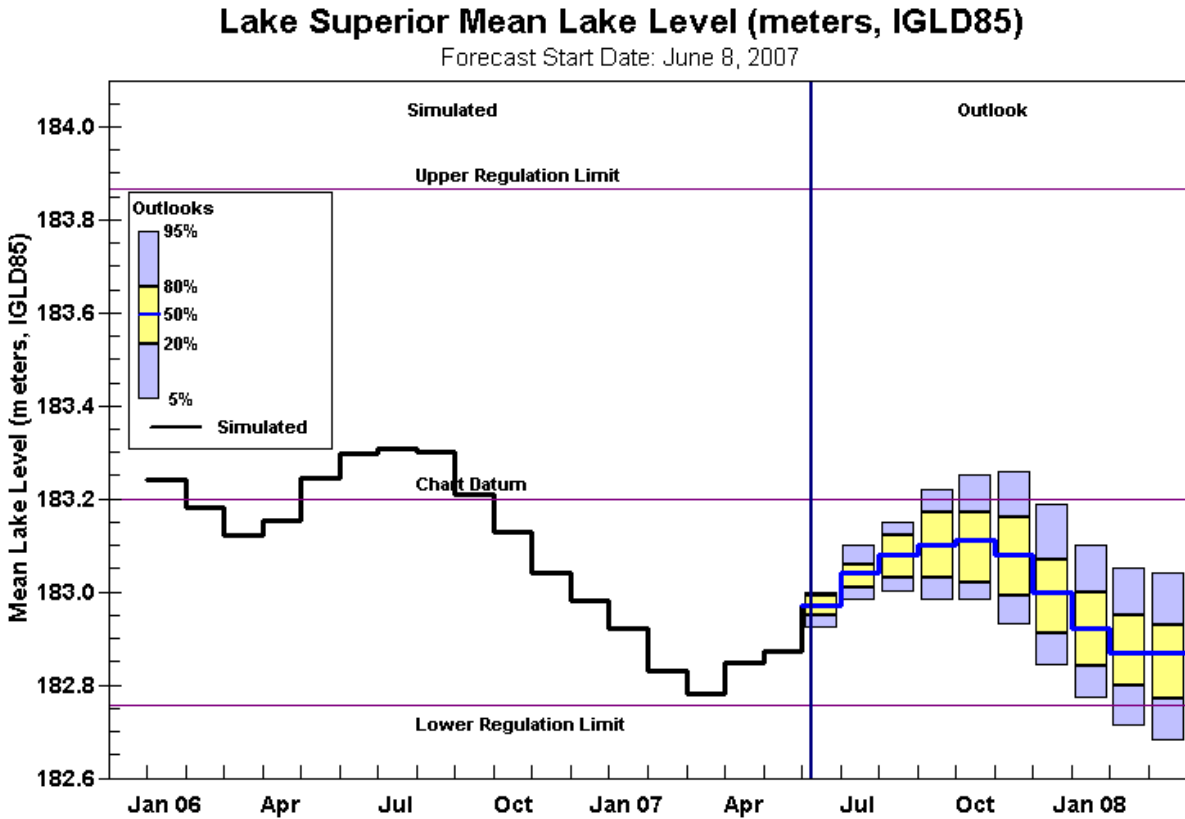
1785 The Bonneville Power Administration, which helps manage and market power from the
1786 Columbia River reservoir system, is both a consumer and producer of hydrologic forecast
1787 products. The BPA generates their own ENSO-state conditioned ESP forecasts of
1788 reservoir system inflows as input to management decisions, a practice supported by
1789 research into the benefits of ENSO information for water management (Hamlet and
1790 Lettenmaier, 1999).

1791

1792 A number of state agencies responsible for releasing hydrologic and water resources
1793 forecasts also make use of climate forecasts in the process of producing their own
1794 hydrologic forecasts. The South Florida Water Management District (SFWMD) predicts
1795 lake (*e.g.*, Okeechobee) and canal stages, and makes drought assessments, using a
1796 decision tree in which the CPC seasonal outlooks play a role. SFWMD follows GLERL's
1797 lead in using the Croley (1996) method for translating the CPC seasonal outlooks to
1798 variables of interest for their system.

1799

1800



1801

1802 **Figure 2.7** Probabilistic forecasts of future lake levels disseminated by GLERL (from:
 1803 <http://www.glerl.noaa.gov/wr/ahps/curfcst/>).
 1804

1805 **2.2.2.3 Local**

1806 At an even smaller scale, some local agencies and private utilities may also produce
 1807 forecasts or at least derive applications-targeted forecasts from the more general climate
 1808 or hydrology forecasts generated at larger agencies or centers. Seattle Public Utilities
 1809 (SPU; see CASE STUDY IN Chapter 4) for example, operates a number of reservoirs for
 1810 use primarily in municipal water supply. SPU makes SI reservoir inflow forecasts using
 1811 statistical methods based on observed conditions in their watersheds (*i.e.*, snow and
 1812 accumulated precipitation), and on the current ENSO state, in addition to consulting the

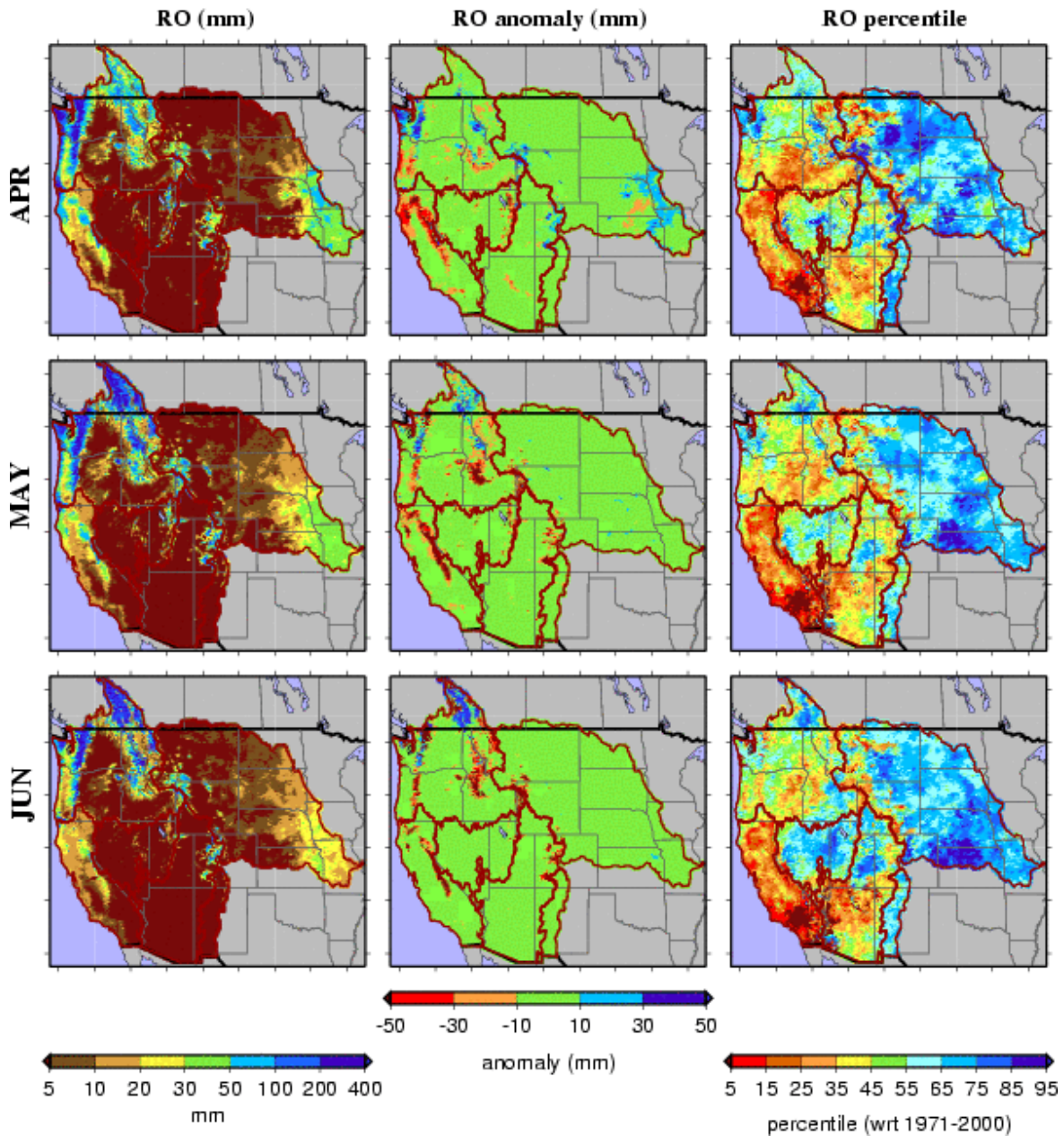
1813 NWRFC volume runoff forecasts. The SPU forecasts are made and used internally rather
1814 than disseminated to the public.

1815

1816 **2.2.2.4 Research**

1817 Research institutions such as universities also produce hydrologic forecasts of a more
1818 experimental nature. A prime example is the Integrated Forecast and Reservoir
1819 Management (INFORM) project housed at the Hydrologic Research Center (HRC),
1820 which produces not only streamflow forecasts in the state of California, but also reservoir
1821 system forecasts; this project is discussed at greater length in Chapter 4 (Georgakakos *et*
1822 *al.*, 2005). At the University of Washington and Princeton University, approximately five
1823 years ago, researchers launched an effort to produce operational hydrologic and
1824 streamflow predictions using distributed land surface models that were developed by an
1825 interagency effort called the Land Data Assimilation System (LDAS) project (Mitchell *et*
1826 *al.*, 2004; Wood and Lettenmaier, 2006); Figure 2.8 shows an example that is based on
1827 the use of CPC climate outlooks. In addition to generating SI streamflow forecasts in the
1828 western and eastern United States, the project also generates forecasts for land surface
1829 variables such as runoff, soil moisture, and snow water equivalent. These forecasts, like
1830 the NWS ESP predictions, are also physically-based, dynamical and objective. The effort
1831 is supported primarily by NOAA, and like the INFORM project collaborates with public
1832 forecast agencies in developing research-level prediction products. The federal funding is
1833 provided with the intent of migrating operational forecasting advances that arise in the
1834 course of these efforts into the public agencies, a topic discussed briefly in Section 2.1.

Runoff (RO) Forecasts (April 1, 2007)



1835

1836 **Figure 2.8** Ensemble median forecasts of monthly runoff from an experimental hydrologic model based on
 1837 CPC climate outlooks. The hydrologic prediction project has run operationally since 2004 at the University
 1838 of Washington, and has a partner effort at Princeton University. Other variables, not shown, include soil
 1839 moisture, snow water equivalent and streamflow. This map was obtained from
 1840 <http://cses.washington.edu/cig/fpt/waterfc/weststreamflowfc.shtml>.
 1841

1842 2.2.3 Skill in SI Hydrologic and Water Resource Forecasts

1843 This section focuses on the skill of hydrologic forecasts; section 2.5 includes a discussion
1844 of forecast utility. Forecasts are statements about events expected to occur at specific
1845 times and places in the future. They can be either deterministic, single-valued predictions
1846 about specific outcomes, or probabilistic descriptions of likely outcomes that typically
1847 take the form of ensembles, distributions, or weighted scenarios.

1848

1849 The hydrologic and water resources forecasts made for water resources management
1850 reflect three components of predictability: the seasonality of the hydrologic cycle,
1851 predictability associated with large-scale climate teleconnections, and persistence of
1852 anomalies in hydrologic initial conditions. Evapotranspiration, runoff (*e.g.*, Pagano *et al.*,
1853 2004) and ground-water recharge (*e.g.*, Earman *et al.*, 2006) all depend on soil moisture
1854 and (where relevant) snowpack conditions one or two seasons prior to the forecast
1855 windows, so that these moisture conditions, directly or indirectly, are key predictors to
1856 many hydrologic forecasts with lead times up to six months. Although hydrologic initial
1857 conditions impart only a few months of predictability to hydrologic systems, during their
1858 peak months of predictability, the skill that they contribute is often paramount. This is
1859 particularly true in the western U.S., where much of the year's precipitation falls during
1860 the cool season, as snow, and then accumulates in relatively easily observed form, as
1861 snowpack, until it predictably melts and runs off in the warm-season months later.

1862 Information about large-scale climatic influences, like the current and projected state of
1863 ENSO, are valued because some of the predictability that they confer on water resources
1864 has influence even before snow begins to accumulate or soil-recharging fall storms

1865 arrive. ENSO, in particular, is strongly synchronized with the annual cycle, so that, in
1866 many instances, the first signs of an impending warm (El Niño) or cold (La Niña) ENSO
1867 event may be discerned toward the end of the summer before the fluctuation reaches its
1868 maturity and peak of influence on the U.S. climate, in winter. This advanced warning for
1869 important aspects of water year climate allows forecasters, in some locations, to
1870 incorporate the expected ENSO influences into hydrologic forecasts before or near the
1871 beginning of the water year (*e.g.*, Hamlet and Lettenmaier, 1999).

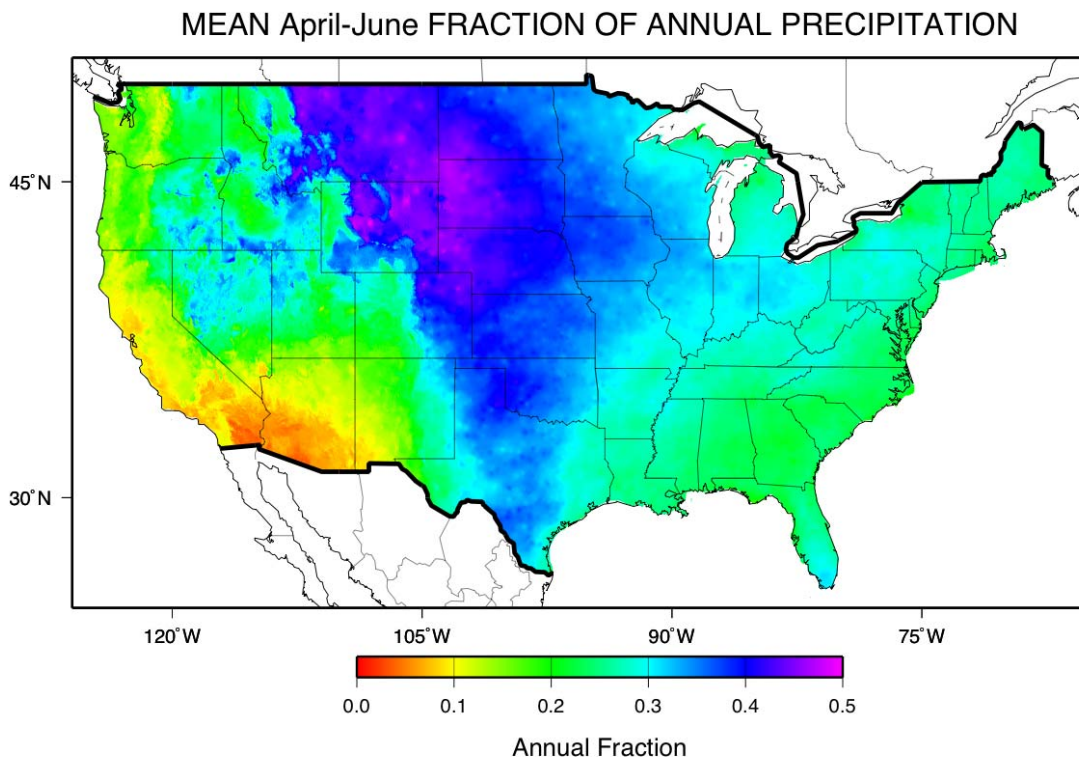
1872

1873 These large-scale climatic influences, however, rarely provide the high level of skill that
1874 can commonly be derived later in the water year from estimates of land surface moisture
1875 state, *i.e.*, from precipitation accumulated during the water year, snow water equivalent or
1876 soil moisture, as estimated indirectly from streamflow. Finally, the unpredictable, random
1877 component of variability remains to limit the skill of all real-world forecasts. The
1878 unpredictable component reflects a mix of uncertainties and errors in the observations
1879 used to initialize forecast models, and errors in the models, and the chaotic complexities
1880 in forecast model dynamics and in the real world.

1881

1882 Many studies have shown that the single greatest source of forecast error is unknown
1883 precipitation after the forecast issue date. Schaake and Peck (1985) estimate that for the
1884 1947-1984 forecasts for inflow to Lake Powell, almost 80% of the January 1st forecast
1885 error is due to unknown future precipitation; by April 1st, Schaake and Peck find that
1886 future precipitation still accounts for 50% of the forecast error. Forecasts can perform
1887 poorly specifically in years with extreme spring precipitation (*e.g.*, 1983 above), or

1888 generally, they can do poorly if spring precipitation is normally a significant component
1889 of the annual cycle. For example, in California, the bulk of the moisture falls from
1890 January-March and rarely does it rain in spring, meaning that April 1 forecasts of spring-
1891 summer streamflow are generally very accurate. In comparison (see Figure 2.9), in
1892 eastern Wyoming and the front range of Colorado, April-through-June is the wettest time
1893 of year and by April 1 the forecaster can only guess at future precipitation events because
1894 of an inability to skillfully forecast springtime precipitation in this region one season in
1895 advance.



1896

1897 **Figure 2.9** Mean percentages of annual precipitation that fall from April through June, 1971-2000 (based
1898 on 4-km PRISM climatologies). This figure was obtained from <http://www.prism.oregonstate.edu/>.
1899

1900 Pagano *et al.* (2004) discovered that the second greatest factor influencing skill is how
1901 much influence snowmelt has on the hydrology of the basin and how warm it is during

1902 the winter. For example, in basins high in the mountains of Colorado, the temperature
1903 remains below freezing for most of the winter. Streamflow is generally low through April
1904 until temperatures rise and the snow starts to melt. The stream then receives a major pulse
1905 of snowmelt over the course of several weeks. Spring precipitation may supplement the
1906 streamflow, but any snow that falls in January is likely to remain in the basin until April
1907 when the forecast target season starts. In comparison, in western Oregon, warm rain-
1908 producing storms can be interspersed with snow-producing winter storms. Most of the
1909 runoff occurs during the winter and it is possible for a large snowpack in February to be
1910 wasted away by March rains. For the forecaster, attempting to predict April-to-July
1911 streamflow is difficult to anticipate, particularly the quantity of water is going to “escape”
1912 before the target season begins.

1913

1914 Some element of forecast accuracy depends on the variability of the river itself. It would
1915 be easy to incur a 100% forecast error on, for example, the San Francisco River in
1916 Arizona, whose observations vary between 17% of average to over 750% of average. It
1917 would be much more difficult to do so on a river such as the Stehekin River in
1918 Washington, where the streamflow ranges only between 60% and 150% of average. A
1919 user may be interested in this aspect of accuracy (*e.g.*, percent of normal error), but most
1920 forecasters use skill scores (*e.g.*, correlation) that would normalize for this effect and
1921 make the results from these two basins more comparable. As noted by Hartmann *et al.*
1922 (2002), consumers of forecast information may be more interested in measures of
1923 forecast skill other than correlations.

1924

1925 **2.2.3.1 Skill of current seasonal hydrologic and water-supply forecasts**

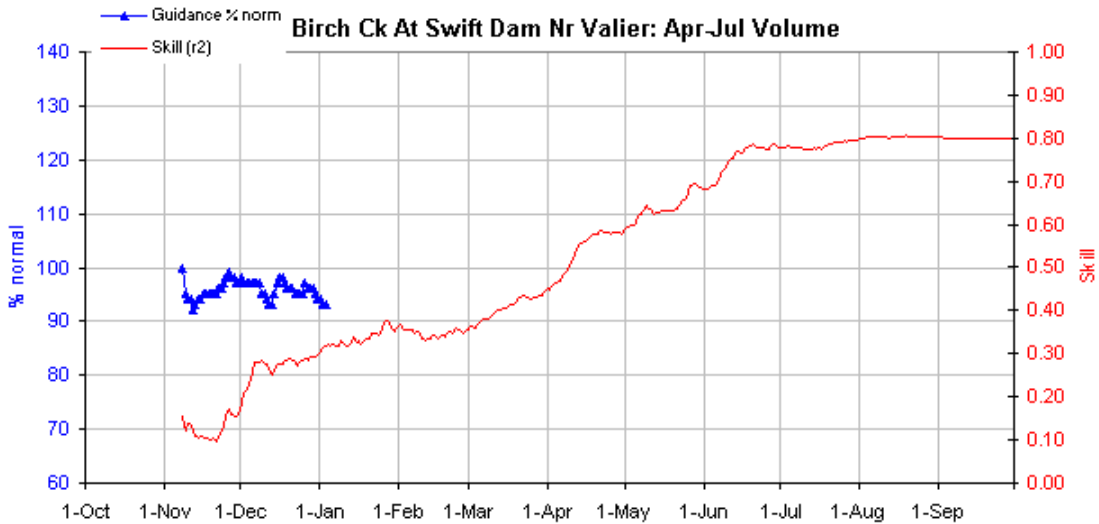
1926 As previously indicated, hydrologic and streamflow forecasts that extend to a 9 -month
1927 lead time are made for western U.S. rivers, primarily during the winter and spring,
1928 whereas in other parts of the United States, where seasonality of precipitation is less
1929 pronounced, the forecasts either link to CPC drought products, are qualitative (the NWS
1930 Southeastern RFC, for instance, provides water supply related briefings from their
1931 website) or in other regards are less amenable to skill evaluation. For this reason, the
1932 following discussion of water supply forecast skill focused mostly on western U.S.
1933 streamflow forecasting, and in particular water supply (*i.e.*, runoff volume) forecasts, for
1934 which most published material relating to SI forecasts exists.

1935

1936 In the western U.S., the skill of operational forecasts generally improves progressively
1937 during the winter and spring months leading up to the period being forecasted, as
1938 increasing information about the year's land surface water budget are observable (*i.e.*,
1939 reflected in snowpack, soil moisture, streamflow and the like). An example of the long-
1940 term average seasonal evolution of NWCC operational forecast skill at a particular stream
1941 gage is shown in Figure 2.10. The flow rates that are judged to have a 50% chance of not
1942 being exceeded (*i.e.*, the 50th percentile or median) are shown by the blue curve for the
1943 early part of 2007. The red curve shows that early in the water year, the April-July
1944 forecast has little skill, measured by the regression coefficient of determination (r^2 or
1945 correlation squared), with only about 10% of historical variance captured by the forecast
1946 equations. By about April 1, the forecast equations predict about 45% of the historical
1947 variance, and at the end of the season, the variance explained is about 80%. This measure

1948 of skill does not reach 100% because the observations available for use as predictors do
 1949 not fully explain the observed hydrologic variation.

1950



1951

1952 **Figure 2.10** Recent operational NWCC forecasts of April-July 2007 streamflow volume in Birch Creek at
 1953 Swift Dam near Valier, showing daily median-forecast values of percentages of long-term average
 1954 streamflow total for summer 2007 (blue) and the long-term estimates of correlation-based forecast skill
 1955 corresponding to each day of the year. (Figure obtained from the National Water and Climate Center
 1956 (NWCC) -- <http://www.wcc.nrcs.usda.gov/>).

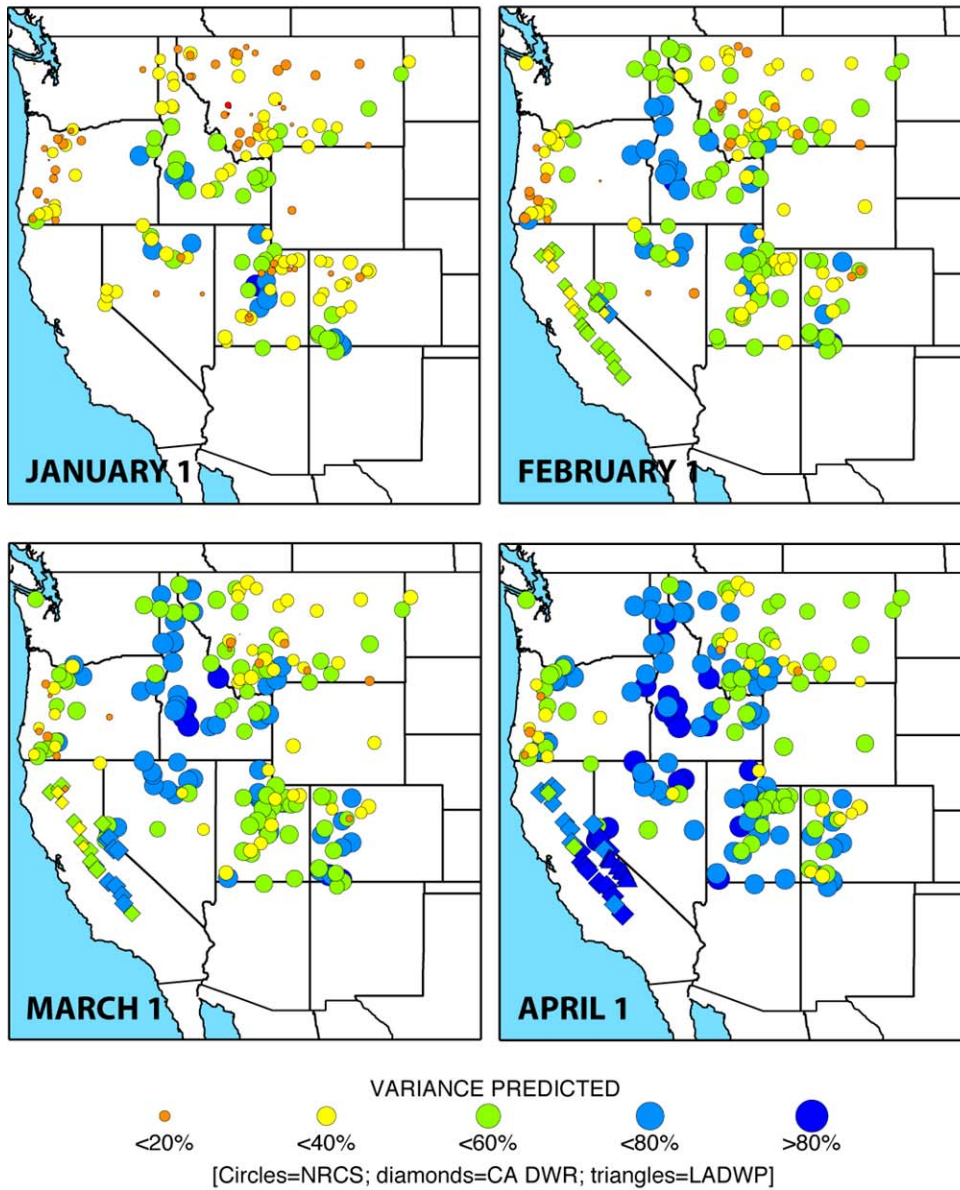
1957

1958 Comparisons of “hindcasts”—seasonal flow estimates generated by applying the
 1959 operational forecast equations to a few decades (lengths of records differ from site to site)
 1960 of historical input variables at each location with observed flows provide estimates of the
 1961 expected skill of current operational forecasts. The actual skill of the forecast equations
 1962 that are operationally used at as many as 226 western stream gages are illustrated in
 1963 Figure 2.11, in which skill is measured by correlation of hindcast median with observed
 1964 values.

1965

1966 The symbols in the various panels of Figure 2.11 become larger and bluer in hue as the
1967 hindcast dates approach the start of the April-July seasons being forecasted. They begin
1968 with largely unskillful beginnings each year in the January 1 forecast; by April 1 the
1969 forecasts are highly skillful by the correlation measures (predicting as much as 80% of
1970 the year-to-year fluctuations) for most of the California, Nevada, and Idaho rivers and
1971 many stations in Utah and Colorado.

**HISTORICAL CORRELATION SKILLS
FOR APRIL-JULY FLOW VOLUMES**



1972

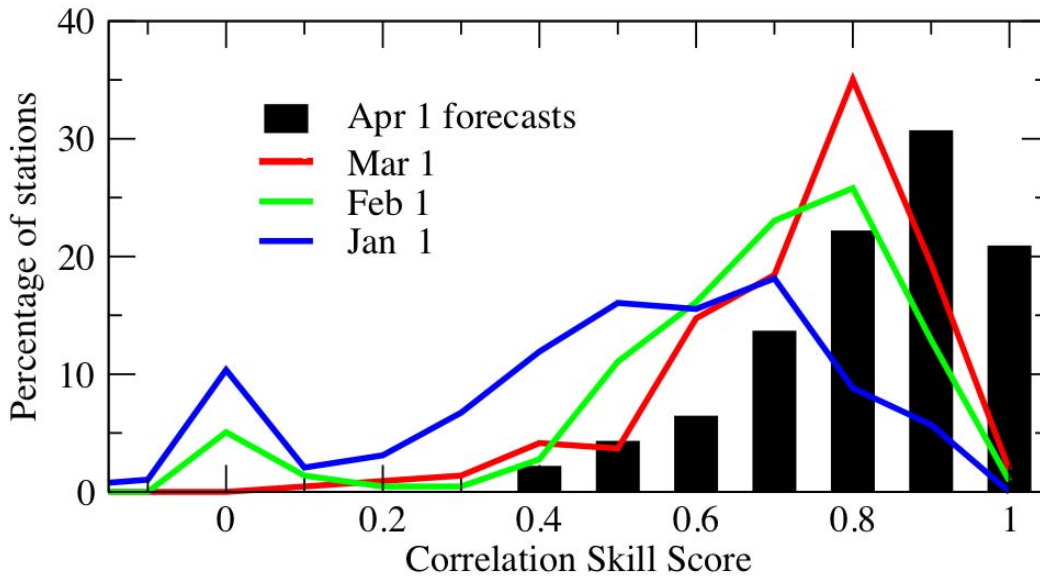
1973 **Figure 2.11** Skills of forecast equations used operationally by NRCS, California Department of Water
 1974 Resources, and Los Angeles Department of Water and Power, for predicting April-July water supplies
 1975 (streamflow volumes) on selected western rivers, as measured by correlations between observed and
 1976 hindcasted flow totals over each station's period of forecast records. Figure provided by Tom Pagano,
 1977 USDA NRCS.

1978

1979 The general increases in skill and thus in numbers of stations with high (correlation) skill

1980 scores as the April 1 start of the forecast period approaches is shown in Figure 2.12.

1981



1982

1983 **Figure 2.12** Percentages of stations with various correlation skill scores in the various panels (forecast
 1984 dates) of Figure 2.11.

1985

1986 A question not addressed in this report relates to the probabilistic skill of the forecasts.
 1987 That is, how reliable are the confidence limits around the median forecasts that are
 1988 provided by the published forecast quantiles (10th and 90th percentiles, for example). In
 1989 a reliable forecast, the frequencies with which the observations fall between various sets
 1990 of confidence bounds matches the probability interval set by those bounds. That is, 80%
 1991 of the time, the observed values fall between the 10th and 90th percentiles of the forecast.
 1992 Among the few analyses that have been published focusing on the probabilistic
 1993 performance of U.S. operational streamflow forecasts, Franz *et al.* (2003) evaluated
 1994 Colorado River basin ESP forecasts using a number of probabilistic measures and found
 1995 reliability deficiencies for many of the streamflow locations considered.

1996

1997 **2.2.3.2 The implications of decadal variability and long term change in climate for**
1998 **seasonal hydrologic prediction skill**

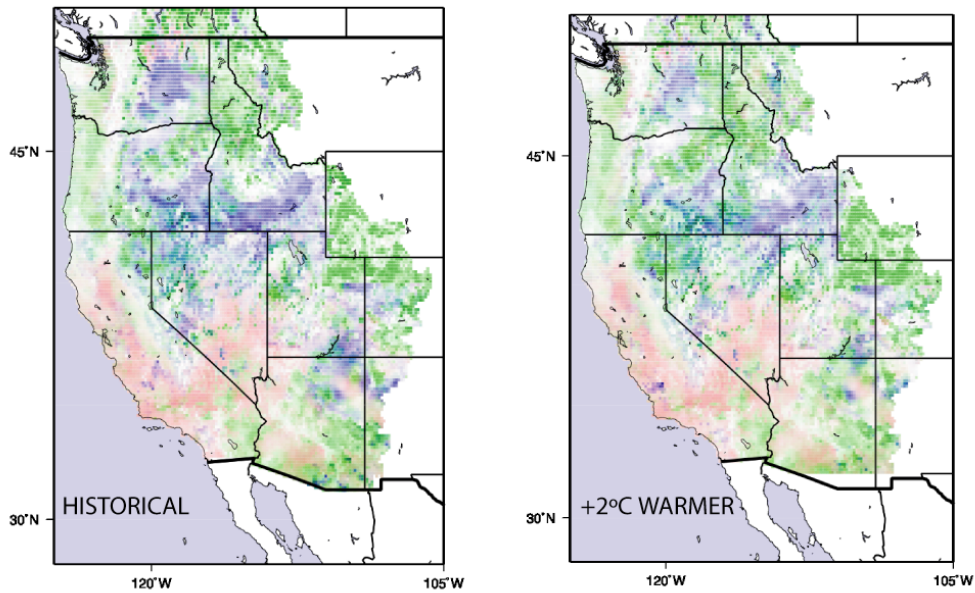
1999 In the earlier discussion of sources of water-supply forecast skill, we highlighted the
2000 amounts and sources of skill provided by snow, soil moisture, antecedent runoff
2001 influences. IPCC projections of global and regional warming, with its expected strong
2002 effects on western U.S. snowpacks (Stewart *et al.*, 2004; Barnett *et al.*, 2008) raises the
2003 concern that prediction methods such as regression that depend on a consistent
2004 relationship between these predictors and future runoff may not perform as expected if
2005 the current climate system is being altered in ways that then alters these hydro-climatic
2006 relationships. Decadal climate variability, particularly in precipitation (*e.g.*, Mantua *et al.*,
2007 1997; McCabe and Dettinger, 1999), may also represent a challenge to such methods,
2008 although some researchers suggest that knowledge of decadal variability can be
2009 beneficial for streamflow forecasting (*e.g.*, Hamlet and Lettenmaier, 1999). One view
2010 voiced in the literature (*e.g.*, Wood and Lettenmaier, 2006) is that hydrologic model-
2011 based forecasting may be more robust to the effects of climate change and variability due
2012 to the physical constraints of the land surface models, but this thesis has not been
2013 comprehensively explored.

2014
2015 The maps shown in Figure 2.13 are based on hydrologic simulations of a physically-
2016 based hydrologic model, the Variable Infiltration Capacity (VIC) model (Liang *et al.*,
2017 1994), in which historical temperatures are uniformly increased by +2°C. These figures
2018 show that the losses of snowpack and the tendencies for more precipitation to fall as rain
2019 rather than snow in a warmer world reduce overall forecast skill, shrinking the areas

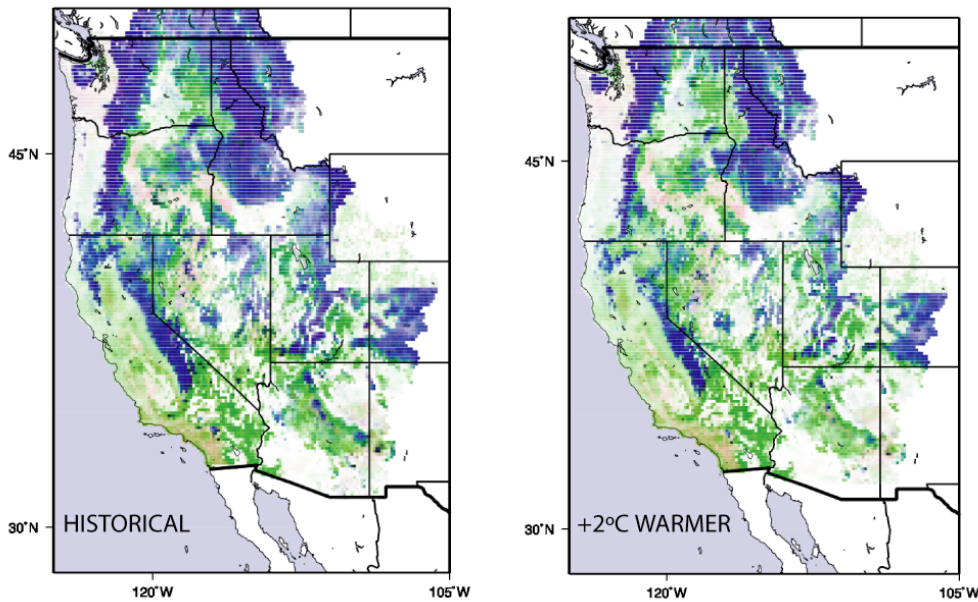
2020 where snowpack contributes strong predictability and also making antecedent runoff a
2021 less reliable predictor. Thus many areas where warm-season runoff volumes are
2022 accurately predicted historically are likely to lose some forecast skill along with their
2023 snowpacks. Overall, the average skill declines by about 2% (out of a historical average of
2024 35%) for the January-March volumes and by about 4% out of a historical average of 53%
2025 for April-July. More importantly, though, are the declines in skill at grid cells where
2026 historical skills are greatest, nearly halving the occurrence of high-end (>0.8) January-to-
2027 March skills and reducing high-end April-to-July skills by about 15% (Figure 2.14).
2028

CHANGES IN CONTRIBUTIONS OF FORECAST SKILL FOR SEASONAL RUNOFF IN RESPONSE TO +2°C WARMING

JANUARY-MARCH RUNOFF FROM DECEMBER PREDICTORS



APRIL-JULY RUNOFF FROM MARCH PREDICTORS



2029

2030

2031

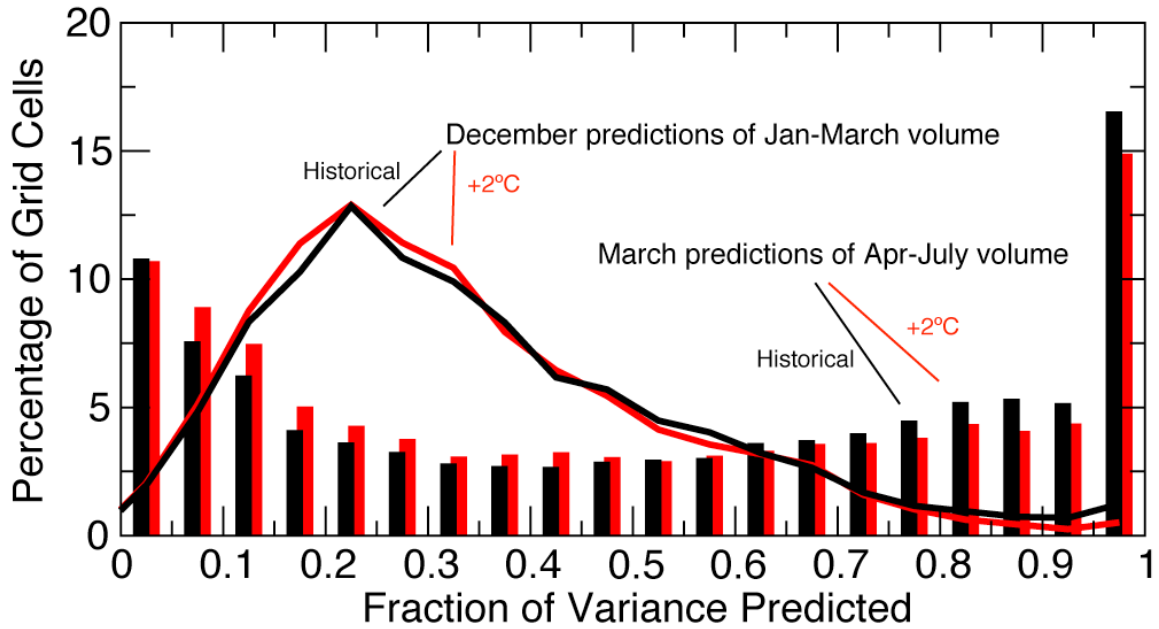
2032

2033

2034

2035

Figure 2.13 Potential contributions of antecedent snowpack conditions, runoff, and Niño 3.4 sea-surface temperatures to seasonal forecast skills in hydrologic simulations under historical, 1950-99, meteorological conditions (left panels) and under those same conditions but with a +2°C uniform warming imposed. (Dettinger, 2007)



2036

2037 **Figure 2.14** Distributions of overall fractions of variance predicted, in Fig. 2.13, of January-March
 2038 (curves) and April-July (histograms) runoff volumes under historical (black) and +2°C warmer conditions.
 2039 (Dettinger, 2007)

2040

2041 This enhanced loss among the most skillful grid cells reflects the strong reliance of those
 2042 grid cells on historical snowpacks for the greater part of their skill, snowpacks which
 2043 decline under the imposed +2°C warmer conditions. Overall, skills associated with
 2044 antecedent runoff are more strongly reduced for the April-to-July runoff volumes, with
 2045 reductions from an average contribution of 24% of variance predicted (by antecedent
 2046 runoff) historically to 21% under the +2°C warm conditions; for the January-to-March
 2047 volumes, skill contributed by antecedent runoff only declines from 18.6% to 18.2% under
 2048 the imposed warmer conditions. The relative declines in the contributions from snowpack
 2049 and antecedent runoff make antecedent runoff (or, more directly, soil moisture, for which
 2050 antecedent runoff is serving as a proxy here) a more important predictor to monitor in the
 2051 future.

2052

2053 It is worth noting that the changes in skill contributions illustrated in Figure 2.13 are best-
2054 case scenarios. The skills shown are skills that would be provided by a complete
2055 recalibration of forecast equations to the new (imposed) warmer conditions, based on 50
2056 years of runoff history. In reality, the runoff and forecast conditions are projected to
2057 gradually and continually trend towards increasingly warm conditions, and fitting new,
2058 appropriate forecast equations (and models) will always be limited by having only a brief
2059 reservoir of experience with each new degree of warming. Consequently, we must expect
2060 that regression-based forecast equations will tend to be increasingly and perennially out
2061 of date in a world with strong warming trends. This problem with the statistics of forecast
2062 skill in a changing world suggests development and deployment of more physically
2063 based, less statistically based forecast models should be a priority in the foreseeable
2064 future.

2065

2066 **2.2.3.3 Skill of climate forecast-driven hydrologic forecasts**

2067 The extent to which the ability to forecast United States precipitation and temperature
2068 seasons in advance can be translated into long-lead hydrologic forecasting has been
2069 evaluated by Wood *et al.* (2005). That evaluation compared hydrologic variables in the
2070 major river basins of the western conterminous U.S. as simulated by the VIC hydrologic
2071 model (Liang *et al.*, 1994), forced by two different sources of temperature and
2072 precipitation data: (1) observed historical meteorology (1979-1999); and (2) by hindcast
2073 climate-model-derived 6-month-lead climate forecasts.

2074

2075 The Wood *et al.* (2005) assessment quantified and reinforced an important aspect of the
2076 hydrologic forecasting community’s intuition about the current levels of hydrologic
2077 forecast skill using long-lead climate forecasts generated from various sources. The
2078 analysis first underscored the conclusions that, depending on the season, knowledge of
2079 initial hydrologic conditions conveys substantial forecast skill. A second finding was that
2080 the additional skill available from incorporating current (at the time) long-lead climate
2081 model forecasts into hydrologic prediction is limited when all years are considered, but
2082 can improve streamflow forecasts relative to climatological ESP forecasts in extreme
2083 ENSO years. If performance in all years is considered, the skill of current climate
2084 forecasts (particularly, of precipitation) is inadequate to provide readily extracted
2085 hydrologic-forecast skill at monthly to seasonal lead times. This result is consistent with
2086 findings for North American climate predictability (Saha *et al.*, 2006). During El Niño
2087 years, however, the climate forecasts have high enough skill for temperatures, and mixed
2088 skill for precipitation, so that hydrologic forecasts for some seasons and some basins
2089 (especially California, the Pacific Northwest and the Great Basin) provide measurable
2090 improvements over the ESP alternative.

2091

2092 The authors of that assessment concluded “climate model forecasts presently suffer from
2093 a general lack of skill, [but] there may be locations, times of year and conditions (*e.g.*,
2094 during El Niño or La Niña) for which they improve hydrologic forecasts relative to ESP”
2095 (Wood *et al.*, 2005). However, their conclusion was that improvements to hydrologic
2096 forecasts based on other forms of climate forecasts, *e.g.*, statistical or hybrid methods that

2097 are not completely reliant on a single climate model may prove more useful in the near
2098 term, presumably until pure climate-model forecasts have improved considerably.

2099

2100 **2.3 CLIMATE DATA AND FORECAST PRODUCTS**

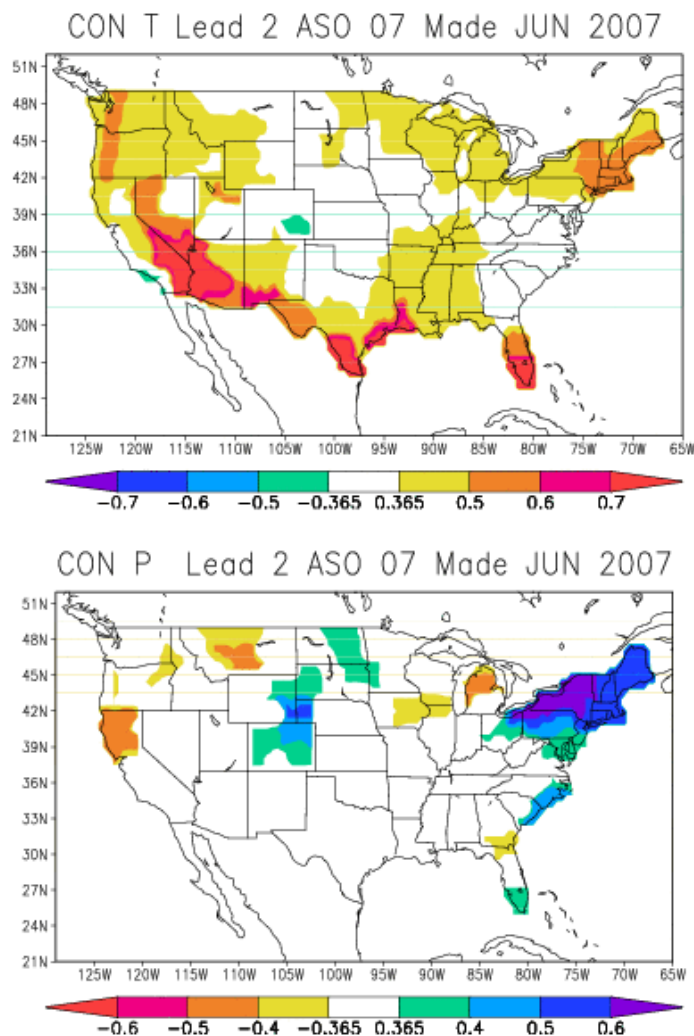
2101 **2.3.1 A Sampling of SI Climate Forecast Products of Interest to Water Resource** 2102 **Managers**

2103 At SI lead times, a wide array of dynamical prediction products exists. A representative
2104 sample of SI climate forecast products is listed in Appendix A.1. The current dynamical
2105 prediction scheme used by NCEP, for example, is a system of models comprising
2106 individual models of the oceans, global atmosphere and continental land surfaces. These
2107 models were developed and originally run for operational forecast purposes in an
2108 uncoupled, sequential mode, an example of which is the so-called “Tier 2” framework in
2109 which the ocean model runs first, producing ocean surface boundary conditions that are
2110 prescribed as inputs for subsequent atmospheric model runs. Since 2004, a “Tier 1”
2111 scheme was introduced in which the models, together called the Coupled Forecast
2112 System (CFS; Saha *et al.*, 2006), were fully coupled to allow dynamic exchanges of
2113 moisture and energy across the interfaces of the model components.

2114

2115 At NCEP, the dynamical tool, CFS, is complemented by a number of statistical forecast
2116 tools, three of which, Screening Multiple Linear Regression (SMLR), Optimal Climate
2117 Normals (OCN), and Canonical Correlation Analysis (CCA), are merged with the CFS to
2118 form an objective consolidation forecast product (Figure 2.15). While the consolidated
2119 forecast exceeds the skill of the individual tools, the official seasonal forecast from CPC

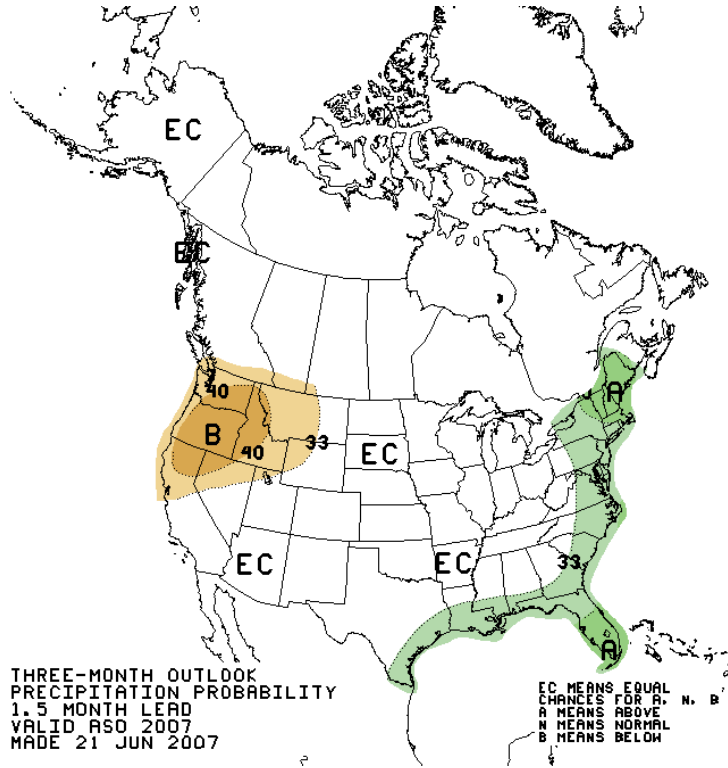
2120 involves a subjective merging of it with forecast and nowcast information sources from a
 2121 number of different sources, all accessible to the public at CPC’s monthly briefing. The
 2122 briefing materials comprise 40 different inputs regarding the past, present and expected
 2123 future state of the land, oceans and atmosphere from sources both internal and external to
 2124 CPC, that are posted online at:
 2125 (<http://www.cpc.ncep.noaa.gov/products/predictions/90day/tools/briefing/>).



2126

2127 **Figure 2.15** CPC objective consolidation forecast for precipitation and temperature for the three month
 2128 period Aug-Sep-Oct 2007, made June 2007 (lead 2 months). Figure obtained from
 2129 <http://www.cpc.ncep.noaa.gov>.
 2130

2131 The resulting official forecast briefing has CPC’s primary presentation of climate forecast
2132 information each month. Forecast products are accessible directly from CPC’s root level
2133 home page in the form of maps of the probability anomalies for precipitation and
2134 temperature in three categories, or “terciles”, representing below-normal, normal and
2135 above-normal values; a two-category scheme (above and below normal) is also available.
2136 This framework is used for the longer lead outlooks (Figure 2.16). The seasonal forecasts
2137 are also available in the form of maps of climate anomalies in degrees Celsius for
2138 temperature and inches for precipitation (Figure 2.17). The forecasts are released
2139 monthly, have a time-step of three months, and have a spatial unit of the climate division
2140 (Figure 2.18). For users desiring more information about the probabilistic forecast than is
2141 given in the map products, a probability of exceedence (POE) plot, with associated
2142 parametric information, is also available for each climate division (Figure 2.19). The
2143 POE plot shows the shift of the forecast probability distribution from the climatological
2144 distribution for each lead-time of the forecast.
2145



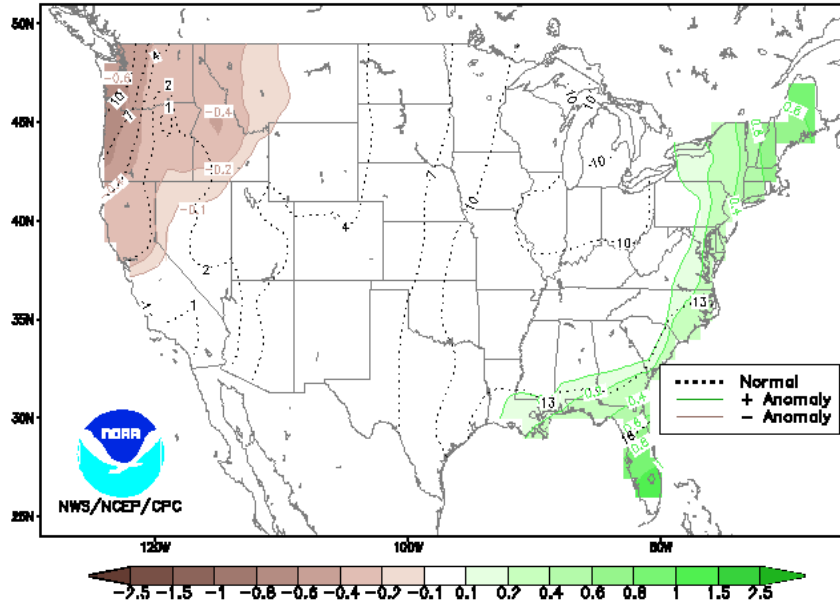
2146

2147

2148 **Figure 2.15** NCEP CPC seasonal outlook for precipitation also shown as a tercile probability map. Figure
2149 obtained from
2150 http://www.cpc.ncep.noaa.gov/products/predictions/multi_season/13_seasonal_outlooks/color/page2.gif.
2151

Anomaly (Inches) of the Mid-value of the 3-Month Precipitation Outlook Distribution for ASO 2007

Dashed lines are the median 3-month precipitation (inches) based on observations from 1971–2000. Shaded areas indicate whether the anomaly of the mid-value is positive (green) or negative (brown) compared to the 1971–2000 average. Non-shaded regions indicate that the absolute value of the anomaly of the mid-value is less than 0.1. For a given location, the mid-value of the outlook may be found by adding the anomaly value to the 1971–2000 average. There is an equal 50–50 chance that actual conditions will be above or below the mid-value. Please note that this product is a limited representation of the official forecast, showing the anomaly of the mid-value, but not the width of the range of possibilities. For more comprehensive forecast information, please see our additional forecast products.

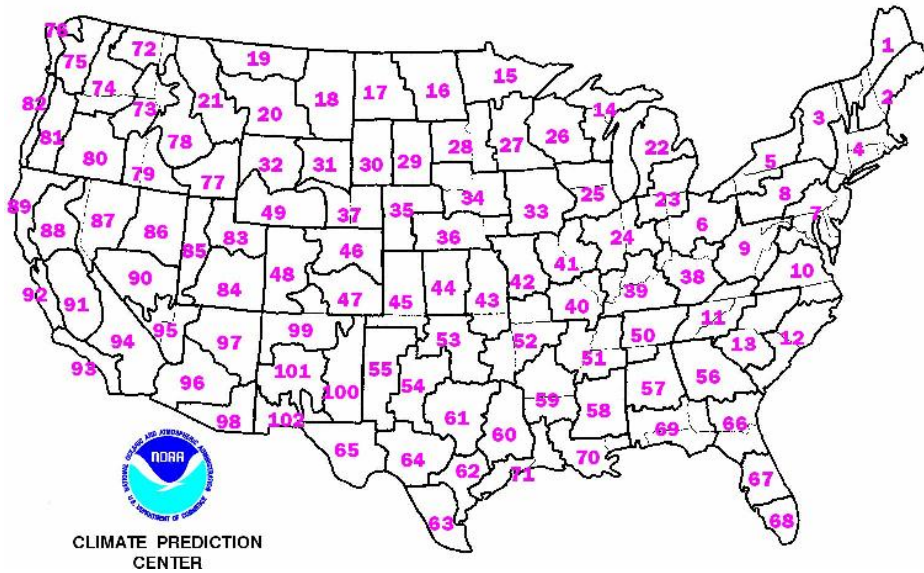


2152

2153 **Figure 2.16** The NCEP CPC seasonal outlook for precipitation from Figure 2.18, but shown as an anomaly
 2154 in inches of total precipitation for the 3-month target period. Figure obtained from
 2155 http://www.cpc.ncep.noaa.gov/products/predictions/long_range/poe_index.php?lead=3&var=p
 2156

2157

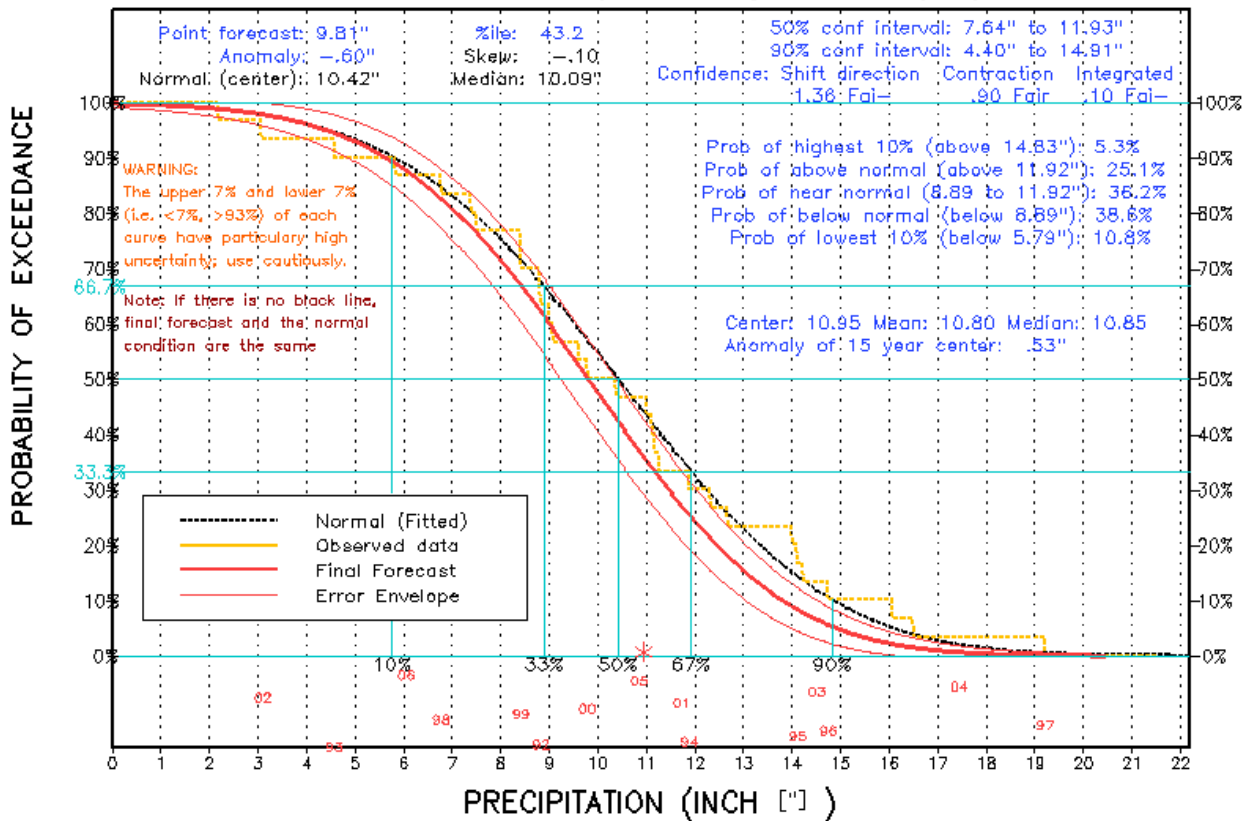
2158



2159

2160 **Figure 2.17** The CPC climate division spatial unit on which the official seasonal forecasts are based.
2161 Figure obtained from
2162 http://www.cpc.ncep.noaa.gov/products/predictions/long_range/poe_index.php?lead=3&var=p.
2163

PRECIPITATION OUTLOOK FOR ASO 2007
1.5 MONTH LEAD OUTLOOK – MADE June 21 2007
Climate Division 75 (Seattle Region, Washington)



2164

2165 **Figure 2.18** The NCEP CPC seasonal outlook for precipitation from Figure 2.17 but shown as an anomaly
 2166 in inches of total precipitation for the 3-month target period.
 2167 http://www.cpc.ncep.noaa.gov/products/predictions/long_range/poe_graph_index.php?lead=3&climdiv=75
 2168 &var=p.
 2169

2170 In addition to NCEP, a few other centers, (e.g., the International Research Institute for
 2171 Climate and Society (IRI)) produce similar consensus forecasts and use a similar map-
 2172 based, tercile-focused framework for exhibiting their results. A larger number of centers
 2173 run dynamical forecast tools, and the NOAA Climate Diagnostics Center, which
 2174 produces monthly climate outlooks internally using statistical tools, also provides
 2175 summaries of climate forecasts from a number of major sources, both in terms of
 2176 probabilities or anomalies, for selected surface and atmospheric variables. The

2177 Experimental Climate Prediction Center (ECPC) at Scripps Institute provides monthly
2178 and seasonal time step forecasts of both climate and land surface variables at a national
2179 and global scale, from dynamical models. Using these model outputs, ECPC also
2180 generates forecasts for derived variables that target wildfire management – *e.g.*, soil
2181 moisture, the Fireweather Index (See Chapter 4 for a more detailed description of Water
2182 Resource Issues in Fire-Prone U.S. Forests and the use of this index) . The CPC has
2183 similar efforts in the form of the Hazards Assessment, a short to medium range map
2184 summary of hazards related to extreme weather (such as flooding and wildfires), and the
2185 CPC Drought Outlook (Box 2.3), a subjective consensus product focusing on the
2186 evolution of large-scale droughts, that is released once a month, conveying expectations
2187 for a 3-month outlook period.

2188

2189 The foregoing is a brief survey of climate forecast products from major centers in the
2190 United States, and as such is far from a comprehensive presentation of the available
2191 sources. It does, however, provide examples from which the following observations about
2192 the general nature of climate prediction in the U.S. may be drawn. First, that operational
2193 SI climate forecasting is conducted at a relatively small number of federally-funded
2194 centers, and forecast products are national to global in scale. These products tend to have
2195 a coarse resolution in space and time, and are typically for basic earth system variables
2196 (*e.g.*, temperature, precipitation, atmospheric and surface pressure) that are of general
2197 interest to many sectors. Forecasts are nearly always probabilistic, and the major products
2198 attempt to convey the inherent uncertainty via maps or data detailing forecast

2199 probabilities, although deterministic reductions (such as forecast variable anomalies) are
2200 also available.

2201

2202 **2.3.2 Sources of Climate-Forecast Skill**

2203 Much as with hydrologic forecasts, the skill of forecasts of climate variables (notably,
2204 temperature and precipitation) varies from region to region, varies with forecast season
2205 and lead time, is limited by the chaotic and uncertain character of the climate system, and
2206 derives from a variety of sources. While initial conditions are an important source for
2207 skill in SI hydrologic forecasts, the initial conditions of an atmospheric forecast are
2208 effectively forgotten after about 8-10 days and have no influence on SI climate forecast
2209 skill (Molteni *et al.*, 1996). SI forecasts are actually forecasts of those variations of the
2210 climate system that reflect predictable changes in boundary conditions, like sea-surface
2211 temperatures (SSTs), or in external ‘forcings’, disturbances in the radiative energy budget
2212 of the Earth’s climate system. At time scales of decades to centuries, potential skill rests
2213 in predictions for slowly varying components of the climate system like the atmospheric
2214 concentrations of CO₂ that influence the greenhouse effect, or slowly evolving changes
2215 in ocean circulation that can alter SSTs and thereby change the boundary conditions for
2216 the atmosphere. Not all possible sources of SI climate-forecast skill have been identified
2217 or exploited, but contributors that have been proposed and pursued include a variety of
2218 large-scale air-sea connections (*e.g.*, Redmond and Koch, 1991; Cayan and Webb, 1992;
2219 Mantua *et al.*, 1997; Enfield *et al.*, 2001; Hoerling and Kumar, 2003), snow and sea ice
2220 patterns (*e.g.*, Cohen and Entekhabi, 1999; Clark and Serreze, 2000; Lo and Clark, 2002;

2221 Liu *et al.*, 2004), and soil moisture and vegetation regimes (*e.g.*, Koster and Suarez, 1995,
2222 2000; Ni-Meister *et al.*, 2005).

2223

2224 In operational practice, however, most of the forecast skill provided by current forecast
2225 systems (especially, including climate models) derives from our ability to predict the
2226 evolution of ENSO events on time scales of 6 to 12 months, coupled with the

2227 “teleconnections” from the events in the tropical Pacific to many areas of the globe.

2228 Barnston *et al.* (1994), in their explanation of the advent of the first operational long-lead
2229 forecasts from the NOAA Climate Prediction Center, stated that “while some

2230 extratropical processes probably develop independently of the Tropics..., much of the

2231 skill of the forecasts for the extratropics comes from anomalies of ENSO-related tropical

2232 sea-surface temperatures.” Except for the changes associated with diurnal cycles,

2233 seasonal cycles, and possibly the (30-60 day) Madden-Julian Oscillation of the tropical

2234 ocean-atmosphere system, “ENSO is the most predictable climate fluctuation on the

2235 planet” (McPhaden *et al.*, 2006). Diurnal cycles and seasonal cycles are predictable on

2236 time scales of hours-to-days and months-to-years, respectively, whereas ENSO mostly

2237 provides predictability on SI time scales (*e.g.*, Figure 2.19b, from a potential

2238 predictability study by Collins 2002). Notice, in Figure 2.19a, that temperatures over the

2239 tropical oceans and lands, and extratropical oceans are much more correlated from season

2240 to season than are conditions on the extratropical continents. To the extent that they can

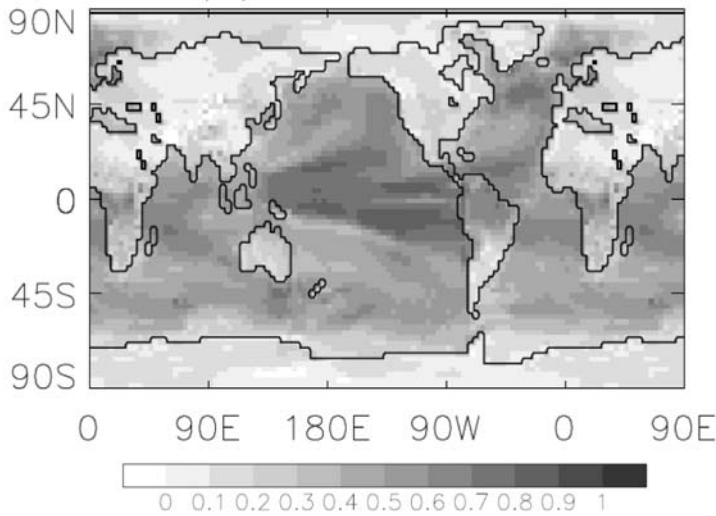
2241 anticipate the slow evolution of the tropical oceans, indicated by these correlations, SCFs

2242 in the extratropics that harken to the tropical oceans are provided a basis for prediction

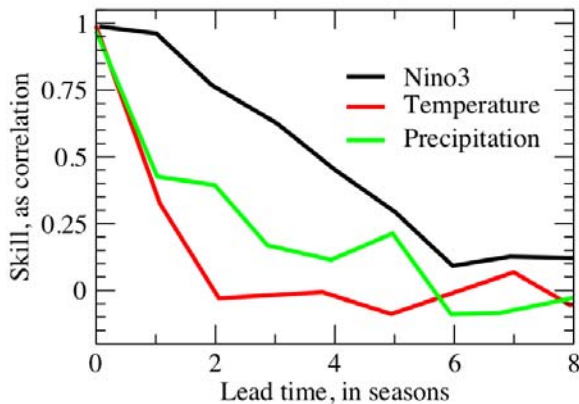
2243 skill; to the extent that the multiseasonal long-term potential predictability of the ENSO

2244 episodes (Figure 2.19b) can be drawn upon in certain regions at certain times of year, the
 2245 relatively meager predictabilities of North American temperatures and precipitation can
 2246 be extended.

2247



2248



2249

2250 **Figure 2.19** (a) Map of correlations between surface-air temperatures in each season and the following
 2251 season in 600 years of historical climate simulation by the HadCM3 model (Collins 2002); (b) Potential
 2252 predictability of a common ENSO index (Niño3 SST, the average of SSTs between 150°W and 90W, 5°S
 2253 and 5°N), average temperatures over the United States and Canada, and average precipitation over the
 2254 United States and Canada, with skill measured by anomaly correlations and plotted against the forecast lead
 2255 times; results extracted from Collins (2002), who estimated these skills from the reproducibility among
 2256 multiple simulations of 30yrs of climate by the HadCM3 coupled ocean-atmosphere model. Correlations
 2257 below about 0.3 are not statistically significant at 95% level.
 2258

2259 The scattered times between ENSO events drastically limits skillful prediction of events
2260 until, at least, the first faltering steps towards the initiation of an ENSO event have been
2261 observed. ENSO events, however, are frequently (but not always) phase-locked
2262 (synchronized) with aspects of the seasonal cycle (Neelin *et al.*, 2000), so that (a)
2263 forecasters know when to look most diligently for those “first faltering steps” and (b) the
2264 first signs of the initiation of an event are often witnessed 6-9 months prior to ENSO’s
2265 largest expressions in the tropics and Northern Hemisphere (*e.g.*, Penland and
2266 Sardeshmukh, 1995). Thus ENSO influences, however irregular and unpredictable they
2267 are on multiyear time scales, regularly provide the basis for SI climate forecasts over
2268 North America. ENSO events generally begin their evolution sometime in late (northern)
2269 spring or early summer, growing and maturing until they most often reach full strength
2270 (measured by either their SST expressions in the tropical Pacific or by their influences on
2271 the Northern Hemisphere) by about December – March (*e.g.*, Chen and van den Dool
2272 1997). An ENSO event’s evolution in the tropical ocean and atmosphere during the
2273 interim period is reproducible enough that relatively simple climate indices that track
2274 ENSO-related SST and atmospheric pressure patterns in the tropical Pacific provide
2275 predictability for North American precipitation patterns as much as two seasons in
2276 advance. Late summer values of the Southern Oscillation Index (SOI), for instance, are
2277 significantly correlated with a north-south see-saw pattern of wintertime precipitation
2278 variability in western North America (Redmond and Koch 1991).

2279

2280 **2.4 IMPROVING WATER RESOURCES FORECAST SKILL AND PRODUCTS**

2281 Although forecast skill is only one measure of the value that forecasts provide to water
2282 resources managers and the public, it is an important measure and current forecasts are
2283 generally understood to fall short of the maximum possible skill on SI time scales (*e.g.*,
2284 http://www.clivar.org/organization/wgsip/spw/spw_position.php). Schaake *et al.* (2007)
2285 describe the SI hydrologic prediction process for model-based prediction in terms of
2286 several components: (i) development, calibration and/or downscaling of SI climate
2287 forecasts; (ii) estimation of hydrologic initial conditions, with or without data
2288 assimilation; (iii) SI hydrologic forecasting models and methods; and (iv) calibration of
2289 the resulting forecasts. Notable opportunities for forecast skill improvement in each area
2290 are discussed here.

2291

2292 **2.4.1 Improving SI Climate Forecast Use for Hydrologic Prediction**

2293 SI climate forecast skill is a function of the skill of climate system models, the efficacy of
2294 model combination strategies if multiple models are used, the accuracy of climate system
2295 conditions from which the forecasts are initiated, and the performance of post processing
2296 approaches applied to correct systematic errors in numerical model outputs.

2297 Improvements are sought in all of these areas.

2298

2299 **2.4.1.1 Climate forecast use**

2300 Many researchers have found that SI climate forecasts must be downscaled,
2301 disaggregated and statistically calibrated to be suitable as inputs for applied purposes
2302 (*e.g.*, hydrologic prediction, as in Wood *et al.*, 2002). Downscaling is the process of
2303 bridging the spatial scale gap between the climate forecast resolution and the

2304 application’s climate input resolution, if they are not the same. If the climate forecasts are
2305 from climate models, for instance, they are likely to be at a grid resolution of several 100
2306 km, whereas the application may require climate information at a point (*e.g.*, station
2307 location). Disaggregation is similar to downscaling, but in the temporal dimension – *e.g.*,
2308 seasonal climate forecasts may need to be translated into daily or subdaily temperature
2309 and precipitation inputs for a given application (as described in Kumar, 2008). Forecast
2310 calibration is a process by which the statistical properties (such as bias and spread errors)
2311 of a probabilistic forecast are corrected to match their observed error statistics (*e.g.*,
2312 Atger, 2003; Hamill *et al.*, 2006). These procedures may be distinct from each other, or
2313 they may be inherent parts of a single approach (such as the analogue techniques of
2314 Hamill *et al.*, 2006). These steps do not necessarily improve the signal to noise ratio of
2315 the climate forecast, but done properly, they do correct bias and reliability problems that
2316 would otherwise render impossible their use in applications. For shorter lead predictions,
2317 corrections to forecast outputs have long been made based on (past) model output
2318 statistics (MOS; Glahn and Lowry, 1972). MOS are sets of statistical relations (*e.g.*,
2319 multiple linear regression (MLR)) that effectively convert numerical model outputs into
2320 unbiased, best climate predictions for selected areas or stations, where “best” relates to
2321 past performance of the model in reproducing observations. MOS corrections are widely
2322 used in weather prediction (Dallavalle and Glahn 2005). Corrections may be as simple as
2323 removal of mean biases indicated by historical runs of the model, with the resulting
2324 forecasted anomalies superimposed on station climatology. More complex methods
2325 specifically address spatial patterns in climate forecasts based on specific inadequacies of

2326 the models in reproducing key teleconnection patterns or topographic features (*e.g.*,
2327 Landman and Goddard 2002, Tippett *et al.*, 2003).

2328

2329 A primary limitation on calibrating SI forecasts is the relatively small numbers of
2330 retrospective forecasts available for identifying biases. Weather predictions are made
2331 every day and thus even a few years' of forecasts provide a large number of examples
2332 from which to learn. SI forecasts, in contrast, are comparatively infrequent and even
2333 several decades' worth may not provide an adequate resource with which to develop
2334 model-output corrections (Kumar, 2007). This limitation is exacerbated when the
2335 predictability and biases themselves vary between years and states of the global climate
2336 system. Thus there is a clear need to expand current "reforecast" practices for fixed SI
2337 climate models over long historical periods to provide both for quantification (and
2338 verification) of the evolution of SI climate forecast skills and for post-processing
2339 calibrations to those forecasts.

2340

2341 **2.4.1.2 Development of objective multi-model ensemble approaches**

2342 The accuracy of SI climate forecasts has been shown to increase when forecasts from
2343 groups of models are combined into multi-model ensembles (*e.g.*, Krishnamurti *et al.*,
2344 2000; Palmer *et al.*, 2004; Tippett *et al.*, 2007). Multi-model forecast ensembles yield
2345 greater overall skill than do any of the individual forecasts included, in principle, as a
2346 result of cancellation of errors between ensemble members. Best results thus appear to
2347 accrue when the individual models are of similar skill and when they exhibit errors and
2348 biases that differ from model to model. In part, these requirements reflect the current

2349 uncertainties about the best strategies for choosing among models for inclusion in the
2350 ensembles used and, especially for weighting and combining the model forecasts within
2351 the ensembles. Many methods have been proposed and implemented (*e.g.*, Rajagopalan *et*
2352 *al.*, 2002; Yun *et al.*, 2005), but strategies for weighting and combining ensemble
2353 members are still an area of active research (*e.g.*, Doblas-Reyes *et al.*, 2005; Coelho *et*
2354 *al.*, 2004). Multi-model ensemble forecast programs are underway in Europe
2355 (DEMETER, Palmer *et al.*, 2004) and in Korea (APEC; *e.g.*, Kang and Park, 2007). In
2356 the United States, IRI forms an experimental multi-model ensemble forecast, updating
2357 monthly, from seasonal forecast ensembles run separately at 7 centers, a 'simple multi-
2358 model' approach that compares well with centrally organized efforts such as DEMETER
2359 (Doblas-Reyes *et al.*, 2005). The NOAA Climate Test Bed Science Plan also envisions
2360 such a capability for NOAA (Higgins *et al.*, 2006).

2361

2362 **2.4.1.3 Improving climate models, initial conditions, and attributions**

2363 Improvements to climate models used in SI forecasting efforts should be a high priority.
2364 Several groups of climate forecasters have identified the lack of key aspects of the
2365 climate system in current forecast models as important weaknesses, including
2366 underrepresented linkages between the stratosphere and troposphere (Baldwin and
2367 Dunkerton 1999), limited processes and initial conditions at land surfaces (Beljaars *et al.*,
2368 1996; Dirmeyer *et al.*, 2006; Ferranti and Viterbo, 2006), and lack of key biogeochemical
2369 cycles like carbon dioxide.

2370

2371 Because climate prediction is, by most definitions, a problem determined by boundary
2372 condition rather than an initial condition, specification of atmospheric initial conditions is
2373 not the problem for SI forecasts that it is for weather forecasts. However, SI climate
2374 forecast skill for most regions comes from knowledge of current SSTs or predictions of
2375 future SSTs, especially those in the tropics (Shukla *et al.*, 2000; Goddard and Dilley,
2376 2005; Rosati *et al.*, 1997). Indeed, forecast skill over land (worldwide) increases directly
2377 with the strength of an ENSO event (Goddard and Dilley, 2005). Thus an important
2378 determinant of recent improvements in SI forecast skill has been the quality and
2379 placement of tropical ocean observations, like the TOGA/TAO network of buoys that
2380 monitors the conditions that lead up to and culminate in El Niño and La Niña events
2381 (Trenberth *et al.*, 1998; McPhaden *et al.*, 1998; Morss and Batistti, 2004). More
2382 improvements in all of the world's oceans are expected from the broader Array for Real-
2383 time Geostrophic Oceanography (ARGO) upper-ocean monitoring arrays and Global
2384 Ocean Observing System (GOOS) programs (Nowlin *et al.*, 2001). In many cases, and
2385 especially with the new widespread ARGO ocean observations, ocean-data assimilation
2386 has improved forecast skill (*e.g.*, Zheng *et al.*, 2006). Data assimilation into coupled
2387 ocean-atmosphere-land models is a difficult and unresolved problem that is an area of
2388 active research (*e.g.*; Ploshay, 2002; Zheng *et al.*, 2006). Land-surface and cryospheric
2389 conditions also can influence the seasonal scale dynamics that lend predictability to SI
2390 climate forecasting, but incorporation of these initial boundary conditions into SI climate
2391 forecasts is in an early stage of development (Koster and Suarez, 2001; Lu and Mitchell,
2392 2004; Mitchell *et al.*, 2004). Both improved observations and improved avenues for

2393 including these conditions into SI climate models, especially with coupled ocean-
2394 atmosphere-land models, are needed.
2395
2396 Finally, a long-standing but little explored approach to improving the value of SI climate
2397 forecasts is the attribution of the causes of climate variations. The rationale for an
2398 attribution effort is that forecasts have greater value if we know why the forecasted event
2399 happened, either before or after the event, and why a forecast succeeded or failed, after
2400 the event. The need to distinguish natural from human-caused trends, and trends from
2401 fluctuations, is likely to become more and more important as climate change progresses.
2402 SI forecasts are always likely to fail from time to time, or to realize less probable ranges
2403 of probabilistic forecasts; knowing that forecasters understand the failures (in hindsight)
2404 and have learned from them will help to build increasing confidence through time among
2405 users. Attempts to attribute causes to important climate events began as long ago as the
2406 requests from Congress to explain the 1930s Dust Bowl. Recently NOAA has initiated a
2407 Climate Attribution Service (<http://www.cdc.noaa.gov/CSI/>) that will combine historical
2408 records, climatic observations, and many climate model simulations to infer the principle
2409 causes of important climate events of the past and present. Forecasters can benefit from
2410 knowledge of causes and effects of specific climatic events as well as improved
2411 feedbacks as to what parts of their forecasts succeed or fail. Users will also benefit from
2412 knowing the reasons for prediction successes and failures.

2413

2414 **2.4.2 Improving Initial Hydrologic Conditions for Hydrologic and Water Resource**
2415 **Forecasts**

2416 Operational hydrologic and water resource forecasts at SI time scales derive much of
2417 their skill from hydrologic initial conditions, with the particular sources of skill
2418 depending on seasons and locations. Thus better estimation of hydrologic initial
2419 conditions will in some seasons lead to improvements in SI hydrologic and consequently
2420 water resources forecast skill. The four main avenues for progress in this area are: (1)
2421 augmentation of climate and hydrologic observing networks; (2) improvements in
2422 hydrologic models (*i.e.*, physics and resolution); (3) improvements in hydrologic model
2423 calibration approaches; and (4) data assimilation.

2424

2425 **2.4.2.1 Hydrologic observing networks**

2426 As discussed previously (in section 2.2), hydrologic and hydroclimatic monitoring
2427 networks provide crucial inputs to hydrologic and water resource forecasting models at SI
2428 time scales. Continuous or regular measurements of streamflow, precipitation and snow
2429 water contents provide important indications of the amount of water that entered and left
2430 river basins prior to the forecasts and thus provide directly or indirectly the initial
2431 conditions for model forecasts.

2432

2433 Observed snow water contents are particularly important sources of predictability in most
2434 of the western half of the United States, and have been measured regularly at networks of
2435 snow courses since the 1920s and continually at SNOTELs (automated and telemetered
2436 snow instrumentation sites) since the 1950s. Snow measurements can contribute as much
2437 as 3/4 of the skill achieved by warm-season water supply forecasts in the West. However,
2438 recent studies have shown that measurements made at most SNOTELs are not

2439 representative of overall basin water budgets, so that their value is primarily as indexes of
2440 water availability rather than as true monitors of the overall water budgets (Molotch and
2441 Bales 2005). The discrepancy arises because most SNOTELs are located in clearings, on
2442 flat terrain, and at moderate altitudes, rather than (historically) sampling snow conditions
2443 throughout the complex terrains and micrometeorological conditions found in most river
2444 basins. The discrepancies limit some of the usefulness of SNOTEL measurements as the
2445 field of hydrologic forecasting moves more and more towards physically-based, rather
2446 than empirical-statistical models. To remedy this situation and to provide the sorts of
2447 more diverse and more widespread inputs required by most physically-based models,
2448 combinations of remotely sensed snow conditions (to provide complete areal coverage)
2449 and extensions of at least some SNOTELs to include more types of measurements and
2450 measurements at more nearby locations will likely be required (Bales *et al.*, 2006).

2451

2452 Ground-water level measurements are made at thousands of locations around the country,
2453 but only recently have they been made available for widespread use in near-real time
2454 (<http://ogw01.er.usgs.gov/USGSGWNetworks.asp>). Few operational surface-water
2455 resource forecasts have been designed to use ground-water measurements. Similarly
2456 climate-driven SI ground-water resource forecasts are rarely made, if at all. However,
2457 surface-water and groundwater are interlinked in nearly all cases and, in truth, constitute
2458 a single resource (Winter *et al.*, 1998). Thus, with the growing availability of real-time
2459 groundwater data dissemination, opportunities for improving water resource forecasts by
2460 better integration and use of surface- and ground-water data resources may develop.

2461 Groundwater level networks already are contributing to drought monitors and response
2462 plans in many states.
2463
2464 Similarly, long-term soil-moisture measurements have been relatively uncommon until
2465 recently. Soil moisture is an important control on the partitioning of water between
2466 evapotranspiration, groundwater recharge and runoff, and thus plays an important (but
2467 largely unaddressed) role in the quantities addressed by water resource forecasts. Soil
2468 moisture varies rapidly from place to place (Vinnikov *et al.*, 1996; Western *et al.*, 2004)
2469 so that networks that will provide representative measurements have always been
2470 difficult to design (Wilson *et al.*, 2004). Nonetheless, the Illinois State Water Survey has
2471 monitored soil moisture at about 20 sites in Illinois for many years
2472 (<http://www.sws.uiuc.edu/warm/soilmoist/ISWSSoilMoistureSummary.pdf>), but for most
2473 of that time was alone in monitoring soil moisture at the state scale. As the technologies
2474 for monitoring soil moisture have become less troublesome, more reliable, and less
2475 expensive in recent years, more and more agencies are beginning to install soil-moisture
2476 monitoring stations (*e.g.*, the NRCS is augmenting many of its SNOTELs with soil-
2477 moisture monitors and has established a national Soil Climate Analysis Network (SCAN;
2478 <http://www.wcc.nrcs.usda.gov/scan/SCAN-brochure.pdf>); Oklahoma's Mesonet
2479 micrometeorological network includes soil-moisture measurements at its sites; California
2480 is on the verge of implementing a state-scale network at both high and low altitudes).
2481 With the advent of regular remote sensing of soil-moisture conditions (Wagner *et al.*,
2482 2007), many of these *in situ* networks will be provided context so that their geographic
2483 representativeness can be assessed and calibrated (Famligietti *et al.*, 1999). As with

2484 ground water, soil moisture has not often been an input to water resource forecasts on the
2485 SI time scale, instead, if anything, being simulated rather than measured, where values
2486 were required. Increased monitoring of soil moisture, both remotely and *in situ*, will
2487 provide important checks on the models of soil-moisture reservoirs that underlie nearly
2488 all of our water resources and water resource forecasts, making hydrological model
2489 improvements possible.

2490

2491 Augmentation of real-time stream gauging networks is also a priority, a subject discussed
2492 in SAP 4.3 (CCSP, 2008).

2493

2494 **2.4.2.2 Improvements in hydrologic modeling techniques**

2495 Efforts to improve hydrologic simulation techniques have been pursued in many areas
2496 since the inception of hydrologic modeling in the 1960s and 1970s when the Stanford
2497 Watershed Model (Crawford and Linsley, 1966), the Sacramento Model (Burnash *et al.*,
2498 1973) and others were created. More recently, physically-based, distributed and semi-
2499 distributed hydrologic models have been developed, both at the watershed scale (*e.g.*,
2500 Wigmosta *et al.*, 1994; Boyle *et al.*, 2000) to account for terrain and climate
2501 inhomogeneity, and at the regional scale (Liang *et al.*, 1994 among others). The latter
2502 category, macroscale models, were motivated in part by the need to improve land surface
2503 representation in climate system modeling approaches (Mitchell *et al.*, 2004), but these
2504 models have also been found useful for hydrologic applications related to water
2505 management (*e.g.*, Hamlet and Lettenmaier, 1999; Maurer and Lettenmaier, 2004; Wood
2506 and Lettenmaier, 2006). The NOAA North American Land Data Assimilation Project

2507 (Mitchell *et al.*, 2004 and NASA Land Information System (Kumar *et al.*, 2006) projects
2508 are leading agency-sponsored research efforts that are focused on advancing the
2509 development and operational deployments of the regional, physically based models.
2510 These efforts include research to improve the estimation of observed parameters (*e.g.*, use
2511 of satellite remote sensing for vegetation properties and distribution), the accuracy of
2512 meteorological forcings, model algorithms and computational approaches. Progress in
2513 these areas has the potential to improve the ability of hydrologic models to characterize
2514 land surface conditions for forecast initialization, and to translate future meteorology and
2515 climate into future hydrologic response.

2516

2517 Aside from improving hydrologic models and inputs, strategies for hydrologic model
2518 implementation are also important. Model calibration – *i.e.*, the identification of optimal
2519 parameter sets for simulating particular types of hydrologic output (single or multiple) –
2520 has arguably been the most extensive area of research toward improving hydrologic
2521 modeling techniques (Wagener and Gupta, 2005 is but one article from a broad
2522 literature). This body of work has yielded advances in the understanding of the model
2523 calibration problem from both practical and theoretical perspectives. The work has been
2524 conducted using models at the watershed scale to a greater extent than the regional scale,
2525 and the potential for applying these techniques to the regional scale models not been
2526 much explored.

2527

2528 Data assimilation is also an area of active research (*e.g.*, Andreadis and Lettenmaier
2529 2006; Reichle *et al.*, 2002; Vrugt *et al.*, 2005; Seo *et al.*, 2006). Data assimilation is a

2530 process in which verifying observations of model state or output variables are used to
2531 adjust the model variables as the model is running, thereby correcting simulation errors
2532 on the fly. The primary types of observations that can be assimilated include snow water
2533 equivalent and snow covered area, land surface skin temperature, remotely sensed or *in*
2534 *situ* soil moisture, and streamflow. NWSRFS has the capability to do objective data
2535 assimilation; in practice NWS (and other agencies) perform a qualitative data
2536 assimilation, in which forecaster judgment is used to adjust model states and inputs to
2537 reproduce variables such as streamflow, snow line elevation and snow water equivalent
2538 prior to initializing an ensemble forecast.

2539

2540 **2.4.3 Calibration of Hydrologic Model Forecasts**

2541 Even the best real-world hydrologic models have biases and errors when applied to
2542 specific gages or locations. Statistical models often are tuned well enough so that their
2543 biases are relatively small, but physically-based models often exhibit significant biases.
2544 In either case, further improvements in forecast skill can be obtained, in principle, by
2545 post-processing model forecasts to remove or reduce any remaining systematic errors, as
2546 detected in the performance of the models in hindcasts. Very little research has been
2547 performed on the best methods for such post processing (Schaake *et al.*, 2007), which is
2548 closely related to the calibration corrections regularly made to weather forecasts. Seo *et*
2549 *al.* (2006), however, describe an effort being undertaken by the National Weather Service
2550 for short lead hydrologic forecasts, a practice that is more common than for longer lead
2551 hydrologic forecasts. Other examples include work by Hashino *et al.* (2007) and
2552 Krzysztofowicz (1999). At least one example of an application for SI hydrologic

2553 forecasts is given in Wood & Schaake (2008); but as noted earlier, a major limitation for
2554 such approaches is the limited sample sizes available for developing statistical
2555 corrections.

2556

2557 2.5 Improving Products: Forecast and related information Packaging and delivery

2558 The value of SI forecasts can depend on more than their forecast skill. The context that is
2559 provided for understanding or using forecasts can contribute as much or more to their
2560 value to forecast users. Several avenues for re-packaging and providing context for SI
2561 forecasts are discussed in the following paragraphs.

2562

2563 Probabilistic hydrologic forecasts typically represent summaries of collections of
2564 forecasts, forecasts that differ from each other due to various representations of the
2565 uncertainties at the time of forecast or likely levels of climate variation after the forecast
2566 is made, or both (Schaake *et al.*, 2007). For example, the “ensemble streamflow
2567 prediction” methodology begins its forecasts (generally) from a single best estimate of
2568 the initial conditions from which the forecasted quantity will evolve, driven by copies of
2569 the historical meteorological variations from each year in the past (Franz *et al.*, 2003).

2570 This provides ensembles of as many forecasts as there are past years of appropriate
2571 meteorological records, with the ensemble scatter representing likely ranges of weather
2572 variations during the forecast season. Sometimes deterministic forecasts are extended to
2573 represent ranges of possibilities by directly adding various measures of past hydrologic or
2574 climatic variability. More modern probabilistic methods are based on multiple climate
2575 forecasts, multiple initial conditions or multiple parameterization (including multiple

2576 downscalings) (Clark *et al.*, 2004; Schaake *et al.*, 2007). However accomplished, having
2577 made numerous forecasts that represent ranges of uncertainty or variability, the
2578 probabilistic forecaster summarizes the results in terms of statistics of the forecast
2579 ensemble and presents the probabilistic forecast in terms of selected statistics, like
2580 probabilities of being more or less than normal.

2581

2582 In most applications, it is up to the forecast user to interpret these statistical descriptions
2583 in terms of their particular data needs, which frequently entails (1) application of various
2584 corrections to make them more representative of their local setting and (2), in some
2585 applications, essentially a deconvolution of the reported probabilities into plausible
2586 examples that might arise during the future described by those probabilities. Forecast
2587 users in some cases may be better served by provision of historical analogs that closely
2588 resemble the forecasted conditions, so that they can analyze their own histories of the
2589 results during the analogous (historical) weather conditions. Alternatively, some forecast
2590 users may find that elements from the original ensembles of forecasts would provide
2591 useful examples that could be analyzed or modeled in order to more clearly represent the
2592 probabilistic forecast in concrete terms. The original forecast ensemble members are the
2593 primary source of the probabilistic forecasts and can offer clear and definite examples of
2594 what the forecasted future COULD look like (but not specifically what it WILL look
2595 like). Thus, along with the finished forecasts—which should remain the primary forecast
2596 products, other representations of what the forecasts are and how they would appear in
2597 the real world could be a useful and more accessible complements for some users, and
2598 would be a desirable addition to the current array of forecast products.

2599

2600 Another approach to providing context (and, potentially, examples) for the SI water
2601 resource forecasts involves placing the SI forecasts in context of paleo-climate
2602 reconstructions. The 20th century has, by and large, been climatically benign in much of
2603 the nation, compared to previous centuries (Hughes and Brown, 1992; Cook *et al.*, 1999).
2604 As a consequence, the true likelihood of various forecasted, naturally occurring climate
2605 and water resource anomalies may best be understood in the context of longer records,
2606 which paleoclimatic reconstructions can provide. At present, approaches to incorporating
2607 paleoclimatic information into responses to SI forecasts are uncommon and only
2608 beginning to develop, but eventually they may provide a clearer framework for
2609 understanding and perfecting probabilistic SI water resource forecasts. One approach that
2610 is being investigated is the statistical synthesis of examples (scenarios) that reflect both
2611 the long-term climate variability identified in paleorecords AND time-series-based
2612 deterministic long-lead forecasts (Kwon *et al.*, 2007).

2613

2614 **2.5 THE EVOLUTION OF PROTOTYPES TO PRODUCTS AND THE ROLE OF**

2615 **EVALUATION IN PRODUCT DEVELOPMENT**

2616 Studies of what makes forecasts useful have identified a number of common
2617 characteristics in the process by which forecasts are generated, developed, and taught to
2618 and disseminated among users (Cash and Buizer, 2005). These characteristics include:
2619 ensuring that the problems that forecasters address are themselves driven by forecast
2620 users; making certain that knowledge-to-action networks (the process of interaction
2621 between scientists and users which produces forecasts) are end-to-end inclusive;

2622 employing “boundary organizations” (groups or other entities that bridge the
2623 communication void between experts and users) to perform translation and mediation
2624 functions between the producers and consumers of forecasts; fostering a social learning
2625 environment between producers and users (*i.e.*, emphasizing adaptation); and providing
2626 stable funding and other support to keep networks of users and scientists working
2627 together.

2628

2629 This section begins by providing a review of recent processes used to take a prototype
2630 into an operational product, with specific examples from the NWS. The section then
2631 reviews a few examples of interactions between forecast producers and users that have
2632 lead to new forecast products, and concludes by describing a vision of how user-centric
2633 forecast evaluation could play a role in setting priorities for improving data and forecast
2634 products in the future.

2635

2636 **2.5.1 Transitioning Prototypes to Products**

2637 During testimony for this report, heads of federal operational forecast groups all painted a
2638 relatively consistent picture of how most in-house innovations currently begin and
2639 evolve. Although formal and quantitative innovation planning methodologies exist (see
2640 Appendix A.3: TRANSITIONING NWS RESEARCH INTO OPERATIONS and How
2641 the Weather Service Prioritizes the Development of Improved Hydrologic Forecasts), for
2642 the most part, the operational practice is often relatively ad-hoc and unstructured except
2643 for the larger and longer-term projects. The Seasonal Drought Outlook is an example of a

2644 product that was developed under a less formal process than that used by the NWS (Box
 2645 2.3).

2646

2647 **BOX 2.3: The CPC Seasonal Drought Outlook**

2648

2649 The CPC Drought Outlook (DO) is a categorical prediction of drought evolution for the 3 months forward
 2650 from the forecast date. The product, which is updated once per month, comprises a map that is
 2651 accompanied by a text discussion of the rationale for the categories depicted on the map.

2652

2653 The starting conditions for the DO are given by the current Drought Monitor (DM) (a United States map
 2654 that is updated weekly showing the status of drought nationwide located:
 2655 <http://www.drought.unl.edu/DM/monitor.html>), and the DO shows likely changes in and adjacent to the
 2656 current DM drought areas. The DO is a subjective consensus forecast that is assembled each month by a
 2657 single author (rotating between CPC and NDMC) with feedback from a panel of geographically distributed
 2658 agency and academic experts. The basis for estimating future drought evolution includes a myriad of
 2659 operational climate forecast products: from short and medium range weather forecasts to seasonal
 2660 predictions from the CPC climate outlooks and the NCEP CFS outputs; consideration of climate tendencies
 2661 for current ENSO state; regional hydroclimatology; and medium range to seasonal soil moisture and runoff
 2662 forecasts from a variety of sources.

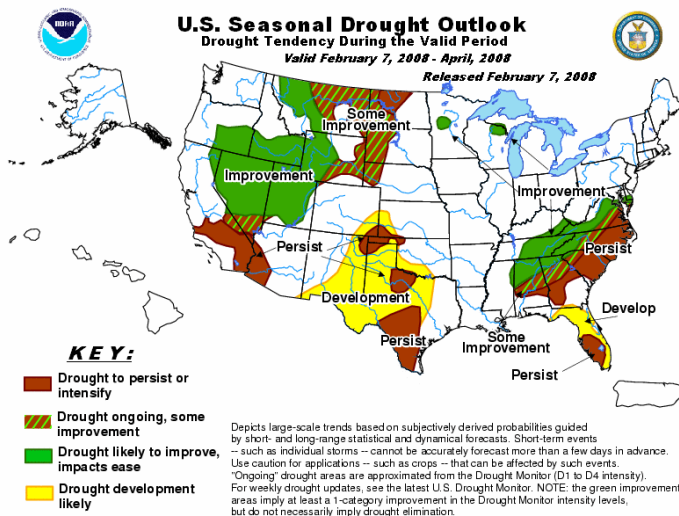
2663

2664 The DO thus makes use of the most advanced objective climate and hydrologic prediction products
 2665 currently available, including not only operational, but experimental products, although the merging of the
 2666 different inputs is based on expert judgment rather than an objective system. The DO is verified by
 2667 comparing the DM drought assessments at the start and end of the DO forecast period; verification skill
 2668 scores have been tracked for the last 7 years. The DO is the primary drought-related agency forecast
 2669 produced in the United States, and is widely used by the drought management and response community
 2670 from local to regional scales.

2671

2672 The DO was developed in the context of new drought assessment partnerships between the CPC, USDA
 2673 and the National Drought Mitigation Center following the passage of the National Drought Policy Act of
 2674 1998. The DM had been released as an official product in August, 1999, with the expectation that a weekly

2675



2693

2694 was informal and lasted about six months. In November 2000, the first Drought Monitor Forum was held,
 2695 at which producers and users (agency, state, private, academic) came together to evaluate the DM in its first

or seasonal drought forecast capacity would be added in the future. A drought on the eastern seaboard in the fall of 1999 required briefings for the press and the U.S. administration; internal discussions between DM participants at the CPC led to the formation of the first version of the DO (maps and text) for these briefings. These were released informally to local, state and federal agency personnel throughout the winter of 1999-2000, and received positive feedback.

The CPC decided to make the products official, provided public statements and developed product specifications, and made the product operational in March 2000. The initial development process

2696 year and plan for its second, providing in addition a venue for discussion of the DO. This forum still meets
2697 bi-annually, focusing on both DM and DO-relevant issues. Developmental efforts for the DO are internal at
2698 CPC or within NCEP, and the primary avenues for feedback are the website and at presentations by DO
2699 authors at workshops and conferences. The DO authors also interact with research efforts funded by the
2700 NOAA Climate Program Office and other agency funding sources, and with NOAA research group efforts
2701 (such as at NCEP), as part of the ongoing development effort. (URL:
2702 http://www.cpc.noaa.gov/products/expert_assessment/drought_assessment.shtml)
2703

2704 **end BOX 2.3*******
2705

2706 Climate and water resource forecasters are often aware of small “fixes” or tweaks to
2707 forecasts that would make their jobs easier; these are often referred to as “forecasts of
2708 opportunity.” A forecaster may be aware of a new dataset or method or product that
2709 he/she believes could be useful. Based on past experience, production of the forecast may
2710 seem feasible and it could be potentially skillful. Especially in climate forecasting, where
2711 there is very high uncertainty in the forecasts themselves and there is marginal user
2712 adoption of existing products, the operational community often focuses more on potential
2713 forecast skill than likely current use. The belief is that if a product is skillful, a user base
2714 could be cultivated. If there is no skill, even if user demand exists, forecasting would be
2715 futile.

2716

2717 Attractive projects may also develop when a new method comes into use by a colleague
2718 of the forecaster (someone from another agency, alumni, friend or prior collaborator on
2719 other projects). For example, Redmond and Koch (1991) published the first major study
2720 of the impacts of ENSO on western U.S. streamflow. At the time the study was being
2721 done, a NRCS operational forecaster was one of Koch’s graduate students. The student
2722 put Koch's research to operational practice at the NRCS after realizing that forecast skill
2723 could be improved.

2724

2725 Efficiency is also often the inspiration for an innovation. A forecaster may be looking for
2726 a way to streamline or otherwise automate an existing process. For example, users
2727 frequently call the forecaster with a particular question; if it is possible to automate the
2728 answering of that question with a new Internet-based product, the forecaster's time may
2729 be freed up to work on other tasks. While most forecasters can readily list several
2730 bottlenecks in the production process, this knowledge often comes more from personal
2731 experience than any kind of structured system review.

2732

2733 At this stage, many ideas exist for possible innovations, although only some small subset
2734 of them will be pursued. The winnowing process continues with the forecaster and/or
2735 peers evaluating the feasibility of the innovation: Is the method scientifically defensible?
2736 Are the data reliably available to support the product? Are the computers powerful
2737 enough to complete the process in a reasonable time? Can this be done with existing
2738 resources, would it free up more resources than it consumes, or is the added value worth
2739 the added operational expense? In other words, is the total value of the advance worth the
2740 effort? Is it achievable and compatible with legacy systems or better than the total worth
2741 of the technology, installed base and complementary products?

2742

2743 If it is expected to be valuable, some additional questions may be raised by the forecaster
2744 or by management about the appropriateness of the solution. Would it conflict with or
2745 detract from another product, especially the official suite (*i.e.*, destroy competency)?
2746 Would it violate an agency policy? For example, a potential product may be technically
2747 feasible but not allowed to exist because the agency's webpage does not permit

2748 interactivity because of increasingly stringent congressionally-mandated cyber-security
2749 regulations. In this case, to the agency as a whole, the cost of reduced security is greater
2750 than the benefit of increased interactivity. It is important to note that if security and
2751 interactivity in general are not at odds, the issue may be that a particular form of
2752 interactivity is not compatible with the existing security architecture. If a different
2753 security architecture is adopted or a different form of interactivity used (*e.g.*, written in a
2754 different computer language), then both may function together, assuming one has the
2755 flexibility and ability to change.

2756

2757 Additionally, an agency policy issue can sometimes be of broader, multi-organizational
2758 scope and would require policy decisions to settle. For example, currently no agency
2759 produces water quality forecasts. Which agency should be responsible for this? The
2760 USDA, Environmental Protection Agency, USGS or NWS? What of soil moisture
2761 forecasts? Should it be the first agency to develop the technical proficiency to make such
2762 forecasts? Or should it be established by a more deliberative process to prevent “mission
2763 creep”? Agencies are also concerned about whether innovations interfere with the
2764 services provided by the private sector.

2765

2766 If appropriate, the forecaster may then move to implement the solution on a limited test
2767 basis, iteratively developing and adapting to any unforeseen challenges. After a
2768 successful functional prototype is developed, it is tested in-house using field personnel
2769 and/or an inner circle of sophisticated customers and gradually made more public as
2770 confidence in the product increases. In these early stages, many of the “kinks” of the

2771 process are smoothed out, developing the product format and look and feel, adapting to
2772 initial feedback (*e.g.*, “please make the map labels larger”) but for the most part the initial
2773 vision remains intact.

2774

2775 There is no consistent formal procedure across agencies for certifying a new method or
2776 making a new product official. A product may be run and labeled “experimental” for 1-2
2777 years in an evaluation period. The objectives and duration of the evaluation period are
2778 sometimes not formalized and one must just assume that if a product has been running for
2779 an extended period of time with no obvious problems, then it succeeds and the
2780 experimental label removed. Creating documentation of the product and process is often
2781 part of the transition from experimental to official, either in the form of an internal
2782 technical memo, conference proceedings or peer-reviewed journal article, if appropriate.

2783

2784 If the innovation involves using a tool or technique that supplements the standard suite of
2785 tools, some of the evaluation may involve running both tools in parallel and comparing
2786 their performance. Presumably ease of use and low demand on resources are criteria for
2787 success (although the task of running models in parallel can, by itself, be a heavy demand
2788 on resources). Sometimes an agency may temporarily stretch its resources to
2789 accommodate the product for the evaluation period and if additional resources are not
2790 acquired by the end of the evaluation (for one of a number of reasons, some of which
2791 may not be related to the product but rather due to variability in budgets), the product
2792 may be discontinued.

2793

2794 Sometimes skill is used to judge success, but this can be a very inefficient measure. This
2795 is because seasonal forecast skill varies greatly from year to year, primarily due to the
2796 variability of nature. Likewise, individual tools may perform better than other tools in
2797 some years but not others. In the 1-2 years of an evaluation period the new tool may be
2798 lucky (or unlucky) and artificially appear better (or worse) than the existing practice.

2799

2800 If the agency recognizes that a tool has not had a fair evaluation, more emphasis is placed
2801 on “hindcasting,” using the new tool to objectively and retrospectively generate realistic
2802 “forecasts” for the last 20-30 years and comparing the results to hindcasts of the existing
2803 system and/or official published forecasts. The comparison is much more realistic and
2804 effective, although hindcasting has its own challenges. It can be very operationally
2805 demanding to produce the actual forecasts each month (*e.g.*, the agency may have to
2806 compete for the use of several hours of an extremely powerful computer to run a model),
2807 much less do the equivalent of 30 years worth at once. These hindcast datasets, however,
2808 have their own uses and have proven to be very valuable (*e.g.*, Hamill *et al.*, 2006 for
2809 medium range weather forecasting and Franz *et al.*, 2003 for seasonal hydrologic
2810 forecasting). Often times, testbeds are better suited for operationally realistic hindcasting
2811 experiments (Box 2.4).

2812

2813 **BOX 2.4: What Role Can a “Testbed” Play in Innovation?**

2814

2815 For an innovation to be deemed valuable, it must be able to stand on its own and be better than the entire
2816 existing system, or marginally better than the existing technology if it is compatible with the rest of the
2817 framework of the existing system. If the innovation is not proven or believed likely to succeed, its adoption
2818 is less likely to be attempted. However, who conducts the experiments to measure this value? And who has
2819 the resources to ensure backwards-compatibility of the new tools in an old system?

2820

2821 Later sections of this report will describe in more detail what is sometimes referred to as the “loading dock”
2822 model of forecast delivery (*i.e.*, the producer creates something, leaves it on the loading dock where the

2823 user seeks it out, picks it up, drives off and uses it; if this process fails, the loading dock mostly comes to
2824 serve as a metaphorical storage facility). This model lacks any direct communication between user and
2825 producer and leaves out the necessary support structure to help users make the most of the product (Cash *et*
2826 *al.*, 2006). Similarly, testbeds are designed as an alternative to the “loading dock” model of transferring
2827 research to operations.

2828
2829 Previously, a researcher may get a short-term grant to develop a methodology, and conduct an idealized,
2830 focused study of marginal operational realism. The results may be presented at research conferences or
2831 published in the scientific literature. While a researcher's career may have a unifying theme, for the most
2832 part, this specific project may be finished when publication is accomplished and the grant finishes.
2833 Meanwhile, the operational forecaster is expected to seek out the methodology and attempt to implement it,
2834 although often times the forecaster does not have the time, resources or expertise to use the results. Indeed,
2835 the forecaster may not be convinced of the incremental advantage of the technique over existing practices if
2836 it has not endured a realistic operational test and been compared to the results of the official system.

2837
2838 Testbeds are intermediate activities, a hybrid mix of research and operations, serving as a conduit between
2839 the operational, academic and research communities. A testbed activity may have its own resources to
2840 develop a realistic operational environment. However, the testbed would not have real-time operational
2841 responsibilities and instead, would be focused on introducing new ideas and data to the existing system and
2842 analyzing the results through experimentation and demonstration. The old and new system may be run in
2843 parallel and the differences quantified. The operational system may even be deconstructed to identify the
2844 greatest sources of error and use that as the motivation to drive new research to find solutions to operations-
2845 relevant problems. The solutions are designed to be directly integrated into the mock-operational system
2846 and therefore should be much easier to directly transfer to actual production.

2847
2848 NOAA has many testbeds currently in operation: Hydrometeorological (floods), Hazardous Weather
2849 (thunderstorms and tornadoes), Aviation Weather (turbulence and icing for airplanes), Climate (ENSO,
2850 seasonal precipitation and temperature) and Hurricanes. The Joint Center for Satellite Data Assimilation is
2851 also designed to facilitate the operational use of new satellite data. A testbed for seasonal streamflow
2852 forecasting does not exist. Generally, satisfaction with testbeds has been high, rewarding for operational
2853 and research participants alike.

2854
2855 **end BOX 2.4 *******
2856

2857 During the evaluation period, the agency may also attempt to increasingly
2858 “institutionalize” a process by identifying and fixing aspects of a product or process that
2859 do not conform to agency guidelines. For example, if a forecasting model is demonstrated
2860 as promising but the operating system or the computer language it is written in does not
2861 match the language chosen by the agency, a team of contract programmers may rewrite
2862 the model and otherwise develop interfaces that make the product more user-friendly for
2863 operational work. A team of agency personnel may also be assembled to help transfer the
2864 research idea to full operations, from prototype to project. For large projects, many
2865 people may be involved, including external researchers from several other agencies.

2866

2867 During this process of institutionalization, the original innovation may change in
2868 character. There may be uncertainty at the outset and the development team may
2869 consciously postpone certain decisions until more information is available. Similarly,
2870 certain aspects of the original design may not be feasible and an alternative solution must
2871 be found. Occasionally, poor communication between the inventor and the developers
2872 may cause the final product to be different than the original vision. Davidson *et al.* (2002)
2873 found success in developing a hydrologic database using structured, iterative
2874 development involving close communication between users and developers throughout
2875 the life of the project. This model is in direct contrast to that of the inventor generating a
2876 ponderous requirements document at the outset, which is then passed on to a separate
2877 team of developers who execute the plan in isolation until completion.

2878

2879 **2.5.2 Evaluation of Forecast Utility**

2880 As mentioned in Section 2.1, there are many ways to assess the usefulness of forecasts,
2881 one of which is forecast skill. While there are inherent limitations to skill (due to the
2882 chaotic nature of the atmosphere), existing operational systems also fall short of their
2883 potential maximum skill for a variety of reasons. Section 2.4 highlights ways to improve
2884 operational skill, such as by having better models of the natural system or denser and
2885 more detailed climate and hydrologic monitoring networks. Other factors, such as
2886 improved forecaster training or better visualization tools, also play a role. This section
2887 addresses the role of forecast evaluation in driving the technology development agenda.

2888

2889 Understanding the current skill of forecast products is a key component to ensuring the
2890 effectiveness of programs to improve the skill of these products. There are several
2891 motivations for verifying forecasts including administrative, scientific and economic
2892 (Brier and Allen, 1951). Evaluation of very recent forecasts can also play a role in
2893 helping operational forecasters make mid-course adjustments to different components of
2894 the forecast system before issuing an official product.

2895

2896 Of particular interest to forecasting agencies is administrative evaluation because of its
2897 ability to describe the overall skill and efficiency of the forecast service in order to
2898 inform and guide decisions about resource allocation, research directions and
2899 implementation strategies (Welles 2005). For example, the development of numerical
2900 weather prediction (NWP) forecasting models is conducted by numerous, unaffiliated
2901 groups following different approaches, with the results compared through objective
2902 measures of performance. In other words, the forecasts are verified, and the research is
2903 driven, not by ad hoc opinions postulated by subject matter experts, but by the actual
2904 performance of the forecasts as determined with objective measures (Welles *et al.*, 2007).

2905 The most important sources of error are identified quantitatively and systematically and
2906 are paired with objective measures of the likely improvement resulting from an
2907 innovation in the system.

2908

2909 Recently the NWS adopted a broad national-scale administrative initiative of hydrologic
2910 forecast evaluation. This program defines a standard set of evaluation measures,
2911 establishes a formal framework for forecast archival and builds flexible tools for access

2912 to results. It is designed to provide feedback to local forecasters and users on the
2913 performance of the regional results, but also to provide an end-to-end assessment of the
2914 elements of the entire system (HVSRT, 2006). Welles *et al.* add that these activities
2915 would be best served by cultivating a new discipline of “hydrologic forecast science” that
2916 engages the research community to focus on operational-forecast-specific issues.

2917

2918 While administrative evaluation is an important tool for directing agency resources,
2919 ultimately innovation should be guided by the anticipated benefit to forecast users. Some
2920 hydrologists would prefer not to issue a forecast that they suspect the user could not use
2921 or would misinterpret (Pielke Jr, 1999). Additionally, these evaluations should be
2922 available and understandable to users. Uncertainty about the accuracy of forecasts
2923 precludes users from making more effective use of them (Hartmann *et al.*, 2002). Users
2924 want to know how good the forecasts are so they know how much confidence to place in
2925 them. Agencies want to focus on the aspects of the forecast that are most important to
2926 users. Forecast evaluation should be more broadly defined than skill, it should also
2927 include measures of communication and understandability, relevance and so on. In
2928 determining these critical aspects, Agencies must make a determination of the key
2929 priorities to address given the number and varied interest of potential forecast users; the
2930 Agencies can not satisfy all users. The Advanced Hydrologic Prediction System (AHPS)
2931 of the NWS provides a nice case study of product development and refinement in
2932 response to user-driven feedback (Box 2.5).

2933

2934 **BOX 2.5: The Advanced Hydrologic Prediction Service**
2935

2936 Short to medium range forecasts (those with lead times of hours to days) of floods are a critical component
2937 of NWS hydrological operations and these services generate nearly \$2 billion of benefits annually (NHWC,
2938 2002). In 1997 the NWS Office of Hydrologic Development began the Advanced Hydrologic Prediction
2939 Service (AHPS) program to advance technology for hydrologic products and forecasts. This 16-year multi-
2940 million dollar program seeks to enhance the agency's ability to issue and deliver specific, timely, and
2941 accurate flood forecasts. One of its main foci is the delivery of probabilistic and visual information through
2942 an Internet based interface. One of its seven stated goals is also to "Expand outreach and engage partners
2943 and customers in all aspects of hydrologic product development." (NWS, 2004)
2944

2945 Starting in 2004, the National Research Council reviewed the AHPS program and also analyzed the extent
2946 that users were actually playing in the development of products and setting of the research agenda
2947 (National Research Council, 2006). The study found that AHPS had largely a top-down structure with
2948 technology being developed at a national center to be delivered to regional and local offices. Although
2949 there was a wide range of awareness, understanding and acceptance of AHPS products inside and outside
2950 the NWS, little to no research was being done in early 2004 on effective communication of information,
2951 and some of the needs of primary customers were not being addressed. From the time the NRC team
2952 carried out its interviews, the NWS started acting on the perceived deficiencies, so that, by the time the
2953 report was issued in late 2006, the NWS had already made some measurable progress. This progress
2954 included a rigorous survey process in the form of focus groups, but also a more engaged suite of outreach,
2955 training, and educational activities that have included presentations at the national floodplain and
2956 hydrologic manager's conferences, the development of closer partnerships with key users, committing
2957 personnel to education activities, conducting local training workshops, and awarding a research grant to
2958 social scientists to determine the most effective way to communicate probabilistic forecasts to emergency
2959 and floodplain managers.

2960
2961 **end BOX 2.5**
2962

2963 There is another component to forecast skill beyond the assessment of how the forecast
2964 quantities are better (or worse) than a reference forecast. Thinking of forecast assessment
2965 more broadly, the forecasts should be evaluated for their 'skill' communicating their
2966 information content in ways that can be correctly interpreted both easily and reliably --
2967 *i.e.*, no matter what the quantity (*e.g.*, wet, dry, or neutral tercile) in the forecast is, the
2968 user can still correctly interpret it (Hartmann *et al.*, 2002).

2969
2970 Finally, it seems important to stress that agencies should provide for user-centric forecast
2971 assessment as part of the process for moving prototypes to official products. That would
2972 include access to user tools for assessing forecast skill (*i.e.*, the Forecast Evaluation Tool,
2973 which is linked to by the NWS Local 3-month Temperature Outlook (Box 2.6), and field
2974 testing of the communication effectiveness of the prototype products. Just as new types of

2975 forecasts should show (at least) no degradation in predictive skill, they should also show
2976 no degradation in their communication effectiveness.

2977

2978 **BOX 2.6: NWS Local 3-Month Outlooks for Temperature and Precipitation**

2979

2980 In January 2007, the NWS made operational the first component of a new set of climate forecast products
2981 called Local 3-Month Outlooks (L3MO). Accessible from the NWS Weather Forecast Offices (WFO),
2982 River Forecast Centers (RFC) and other NWS offices, the Local 3-Month Temperature Outlook (L3MTO)
2983 is designed to clarify and downscale the national-scale CPC Climate Outlook temperature forecast product.
2984 The corresponding local product for precipitation is still in development as of the writing of this report.
2985 The local outlooks were motivated by ongoing NOAA NWS activities focusing on establishing a dialog
2986 with NWS climate product users (<http://www.nws.noaa.gov/directives/>),. In particular, a 2004 NWS
2987 climate product survey (conducted by Claes Fornell International for the NOAA Climate Services Division)
2988 found that a lack of climate product clarity lowered customer satisfaction with NWS CPC climate outlook
2989 products; and presentations and interactions at the annual Climate Prediction Application Science
2990 Workshop (CPASW) highlighted the need for localized CPC climate outlooks in numerous and diverse
2991 applications.

2992 In response to these user-identified issues, CSD collaborated with the NWS Western Region Headquarters,
2993 CPC and the National Climatic Data Center (NCDC) to develop localized outlook products. The
2994 collaboration between the four groups, which linked several line offices of NOAA (*e.g.*, NCDC, NWS),
2995 took place in the context of an effort that began in 2003 to build a climate services infrastructure within
2996 NOAA. The organizations together embarked on a structured process that began with a prototype
2997 development stage, which included identifying resources, identifying and testing methodologies, and
2998 defining the product delivery method. To downscale the CPC climate outlooks (which are at the climate
2999 division scale) to local stations, the CSD and WR development team assessed and built on internal, prior
3000 experimentation at CPC that focused on a limited number of stations. To increase product clarity, the team
3001 added interpretation, background information, and a variety of forecast displays providing different levels
3002 of data density. A NWS products and services team made product mockups that were reviewed by all 102
3003 WFOs, CPC and CSD representatives and a small number of non-agency reviewers. After product
3004 adjustments based on the reviews, CSD moved toward an experimental production stage by obtaining union
3005 approval, providing NWS staff with training and guidelines, releasing a public statement about the product
3006 and writing product description documentation. Feedback was solicited via the experimental product
3007 website beginning in August 2006, and the products were again adjusted. Finally, the products were
3008 finalized, the product directive was drafted and the product moved to an operational stage with official
3009 release. User feedback continues via links on the official product website
3010 (<http://www.weather.gov/climate/l3mto.php>).
3011

3012 In general, the L3MO development process exhibited a number of strengths. Several avenues existed for
3013 user needs to reach developers, and user-specified needs determined the objectives of the product
3014 development effort. The development team spanning several parts of the agency then drew on internal
3015 expertise and resources to propose and to demonstrate tentative products responding to those needs. The
3016 first review stage of the process gave mostly internal (*i.e.*, agency) reviewers an early opportunity for
3017 feedback, but this was followed by an opportunity for a larger group of users in the experimental stage,
3018 leading to the final product. An avenue for continued review is built into the product dissemination
3019 approach.

3020

3021 end BOX 2.6*****

3022

3023

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