# Performance and Diagnostic Evaluation of Ozone Predictions by the Eta-Community Multiscale Air Quality Forecast System during the 2002 New England Air Quality Study

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#### **ABSTRACT**

A real-time air quality forecasting system (Eta-Community Multiscale Air Quality [CMAQ] model suite) has been developed by linking the National Centers for Environmental Estimation Eta model to the U.S. Environmental Protection Agency (EPA) CMAQ model. This work presents results from the application of the Eta-CMAQ modeling system for forecasting ozone  $(O_3)$  over the Northeastern United States during the 2002 New England Air Quality Study (NEAQS). Spatial and temporal performance of the Eta-CMAQ model for O<sub>3</sub> was evaluated by comparison with observations from the

#### **IMPLICATIONS**

Air quality forecast simulations over the Northeastern United States with the Eta-CMAQ system are evaluated against a variety of measurements from the 2002 NEAQS. The model captured the hourly variations and broad synoptic and inter-day variations seen in the observations of different gas species. In light of uncertainties in current photochemical mechanisms and specification of real-time emission estimates and prognostic meteorological fields, the model performance for O<sub>3</sub> in forecast mode can be considered to be reasonable.

EPA Air Quality System (AQS) network. This study also examines the ability of the model to simulate the processes governing the distributions of tropospheric O<sub>3</sub> on the basis of the intensive datasets obtained at the four Atmospheric Investigation, Regional Modeling, Analysis, and Estimation (AIRMAP) and Harvard Forest (HF) surface sites. The episode analysis reveals that the model captured the buildup of O<sub>3</sub> concentrations over the northeastern domain from August 11 and reproduced the spatial distributions of observed O<sub>3</sub> very well for the daytime (8:00 p.m.) of both August 8 and 12 with most of normalized mean bias (NMB) within  $\pm 20\%$ . The model reproduced 53.3% of the observed hourly O<sub>3</sub> within a factor of 1.5 with NMB of 29.7% and normalized mean error of 46.9% at the 342 AQS sites. The comparison of modeled and observed lidar O<sub>3</sub> vertical profiles shows that whereas the model reproduced the observed vertical structure, it tended to overestimate at higher altitude. The model reproduced 64–77% of observed NO2 photolysis rate values within a factor of 1.5 at the AIRMAP sites. At the HF site, comparison of modeled and observed O<sub>3</sub>/nitrogen oxide (NO<sub>x</sub>) ratios suggests that the site is mainly under strongly NO<sub>x</sub>sensitive conditions (>53%). It was found that the modeled lower limits of the O<sub>3</sub> production efficiency values (inferred from O<sub>3</sub>-CO correlation) are close to the observations.

# INTRODUCTION

 $O_3$ , a secondary pollutant, is created in part by pollution from anthropogenic and biogenic sources. The daily maximum 8-hr National Ambient Air Quality Standards (NAAQS) promulgated by U.S. Environmental Protection Agency (EPA) in 1997 to reflect more recent health-effect studies for ground-level O<sub>3</sub> are 0.08 ppm.<sup>1</sup> The standard is considered to be attained if the daily maximum 8-hr average O<sub>3</sub> concentration does not exceed 0.08 ppm more than three times in 3 years. Chameides et al.<sup>2</sup> showed that more regions in the U.S. will have daily maximum 8-hr O<sub>3</sub> concentrations that exceed the level of the revised NAAQS than the old standard, which is 0.12 ppm for the daily maximum 1-hr O<sub>3</sub>. It is desirable for local air quality agencies to accurately forecast O<sub>3</sub> concentrations to warn the public of unhealthy air and to encourage people to voluntarily reduce emissions-producing activities.

Different air pollution forecasting systems for O<sub>3</sub> have been developed and are presently in operational use in the United States and other countries, ranging from simple statistical models to comprehensive three-dimensional air quality simulation models. <sup>1</sup> EPA<sup>1</sup> reviewed the definitions, strengths, and limitations of the most commonly used forecasting methods so far, including persistence, climatology, criteria, classification and regression tree, regression equations, artificial neural networks, the phenomenological/intuition method, and three-dimensional air quality models. Three-dimensional air quality forecast models have been increasingly used in operational applications because they can forecast temporal and spatial distributions of O<sub>3</sub> and its precursors over regions where observations are sparse and help to understand the chemical-physical processes that control O<sub>3</sub> in a specific area. Kang et al.<sup>3</sup> evaluated the operational performance of three three-dimensional air quality forecast models and found that all three models tended to overestimate O<sub>3</sub> concentrations with mean biases (MBs) ranging from 0.53 ppb for maximum 8-hr forecast to 7.42 ppb for maximum 1-hr forecast. In this study, the National Weather Service operational mesoscale forecast Eta model is used to supply meteorological input to the EPA Models-3/Community Multiscale Air Quality (CMAQ) model (Eta-CMAQ model suite); the models are then used to provide estimations of  $O_3$  and related chemical species in a forecast mode. The information from the other  $O_3$ -related chemical species can help to gain an understanding of the physical and chemical processes dictating tropospheric O<sub>3</sub> distributions. The developmental testing of fine particulate matter forecast capability with the Eta-CMAQ model is under way.4 The objectives of this study are: (1) to evaluate the temporal and spatial performances of the Eta-CMAQ forecast model for O<sub>3</sub> against the observations from the Air Quality System (AQS) network over the Northeastern United States; and (2) to use a variety of diagnostic tests involving measurements from the 2002 New England Air Quality Study (NEAQS) to examine the ability of the Eta-CMAQ model in representing the physical and chemical processes dictating tropospheric O<sub>3</sub> distributions.

# DESCRIPTIONS OF THE MODEL AND OBSERVATIONAL DATA

# **Eta-CMAQ Forecast Modeling System**

The Eta-CMAQ air quality forecasting (AQF) system is based on the National Centers for Environmental Estimation (NCEP) Eta model<sup>5</sup> and the EPA CMAQ Modeling System.<sup>6</sup> Otte et al.<sup>7</sup> describe the methodology developed to link the two modeling systems. A brief summary relevant to the present study is presented here. The Eta model is used to generate the three-dimensional meteorological fields required for the chemistry and transport calculations in the CMAQ. The NCEP Product Generator software is used to perform the bilinear interpolations and nearest-neighbor mappings of the Eta postprocessor output from the Eta forecast domain to the CMAQ forecast domain. The emissions are from EPA national emission inventory (1999 NEI v1). The Carbon Bond chemical mechanism (version 4.2) (CB-IV) is used for photochemical processes.

Detailed description of the governing equations, physical parameterizations, and numerical techniques used in the CMAQ model can be found in Byun and Ching.6 In this study, the modeling system is deployed over a domain covering the Northeast United States (Figure 1). A Lambert Conformal map projection is used, and the horizontal domain is discretized using grids of 12-km resolution. Twenty-two layers of variable thickness set on a  $\sigma$  coordinate are used to resolve the vertical extent from the surface to ~100 hPa. The lateral boundary conditions are set using horizontally constant and typically "clean" continental profiles for O<sub>3</sub> and other trace gases with some vertical variations based on climatology. The initial conditions for modeled chemical species are set from the previous forecast cycle. The Eta 12 UTC and 06 UTC cycles are used for the forecast cycle.7 The primary Eta-CMAQ model forecast for next-day surface-layer O<sub>3</sub> is based on the 12 UTC Eta cycle of the current day, and products are issued daily no later than 1:30 p.m. The target forecast period is local midnight through local midnight (04 UTC to 03 UTC for the Northeast United States), so in this work, the first 16-hr results of the CMAQ forecast based on the 12-UTC run are discarded. An additional 8 hr is required beyond midnight to calculate peak 8-hr average O<sub>3</sub> concentrations. Therefore, a 48-hr Eta-CMAQ forecast is needed on the basis of the 12-UTC initialization to obtain the desired 24-hr forecast period. The model performance from August 6 to August 17, 2002, based on the 12-UTC run for the target forecast period is evaluated in this study.

# **Observational Databases**

Hourly  $\rm O_3$  data at 342 sites in the Northeast United States are available from EPA AQS network (Figure 1). The cell location of a site is used for matching the model estimations with observations. For cases in which the model grid cells contained more than one monitor, the average concentration from the monitors was used. Four Atmospheric Investigation, Regional Modeling, Analysis, and Estimation (AIRMAP) sites and Harvard Forest ([HF] 42.64 $^{\circ}$  N, 72.17 $^{\circ}$  W) provided continuous measurements of  $\rm O_3$  and related photochemical species, as well as meteorological parameters during the 2002 NEAQS. The four AIRMAP

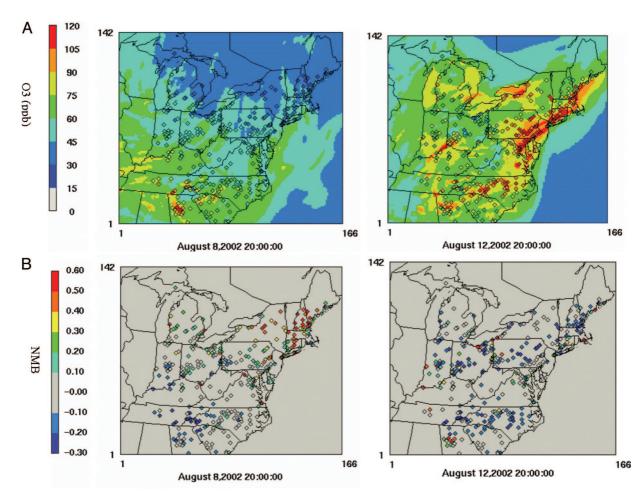


Figure 1. (A) AQF system simulation results for O<sub>3</sub> concentration (ppb) with AQS observed data overlaid (♦) and (B) the normalized MB (NMB = [model - observed]/observed) over the Northeastern United States at 20:00 UTC (3:00 p.m.) August 8 and 12, 2002.

sites include Castle Springs ([CS] 43.73° N, 71.33° W), Isle of Schoals ([IS] 42.99° N, 69.33° W), Mount Washington Observatory ([MWO] 44.27° N, 71.30° W), and Thompson Farm ([TF] 43.11° N, 70.95° W) sites. Both CS and TF sites are characterized by mixed hardwood/pine forest, and IS is an uninhabited ocean site (see airmap.unh.edu/). MWO is the highest mountain (1916 m) in the Northeastern United States. To compare the observations with the model estimations, the hourly averages for the observations were calculated if >50% of the 1-min observations in that hour were available. O3 lidar vertical profiles obtained from the lidar onboard the NOAA ship Ronald H. Brown during the NEAQS 2002 (www.etl.noaa.gov/et2/ data/data\_pages/neaqs/opal/) were used to evaluate the model performance in the vertical. The lidar measured O<sub>3</sub> concentrations from 295 m to 1975 m with 30-m vertical resolution and 5-min time resolution; to compare with the model estimations, the hourly mean  $O_3$  vertical profiles are calculated.

# **RESULTS AND DISCUSSION** O<sub>3</sub> Episode Analysis during the **Simulation Period**

Because the mobile sources (or urban plumes) are rich in CO and nitrogen oxide (NO<sub>x</sub>), and point source emissions from power plants are often rich in SO2 and NOx, the

 $SO_2$ /reactive odd nitrogen ([NO<sub>v</sub>] NOy = NO + NO<sub>2</sub> +  $NO_3 + 2*N_2O_5 + HONO + HNO_3 + PNA + peroxyacetyl$ nitrate [PAN] + NTR [an inert organic nitrogen terminal product in the CB-IV chemical mechanism]) ratios can be used to discriminate between air masses dominated by mobile sources ( $SO_2/NO_v < 1$ ) or by point sources ( $SO_2/NO_v < 1$ )  $NO_v > 1$ ). Figure 2 shows the time series of modeled and observed  $SO_2$ , CO, and  $(SO_2)/(NO_v)$  ratios at the CS and TF sites. As can be seen, these sites were significantly influenced by the polluted plumes from both mobile and point sources after approximately August 10, because observed CO, SO<sub>2</sub>, and (SO<sub>2</sub>)/(NO<sub>v</sub>) ratios increased significantly. The model captured the buildup of SO<sub>2</sub> after August 10 at the CS site, whereas the model systematically overestimated the observed SO<sub>2</sub> mixing ratios at the CS and TF sites most of time as shown in Figure 2. One of the possible reasons for the overestimation of SO<sub>2</sub> comes from nonrepresentative locations and elevations of surface observation sites, because most of SO<sub>2</sub> is emitted from stacks (point sources) above local shallow inversion layers, with the observation stations located close to the surface below the inversions.8 Consequently, more rapid mixing in the model relative to reality may result in overestimation of SO<sub>2</sub> concentrations at these sites. On the other hand, the model captured the temporal variations of observed CO at these two sites much better than SO<sub>2</sub> as shown in Figure

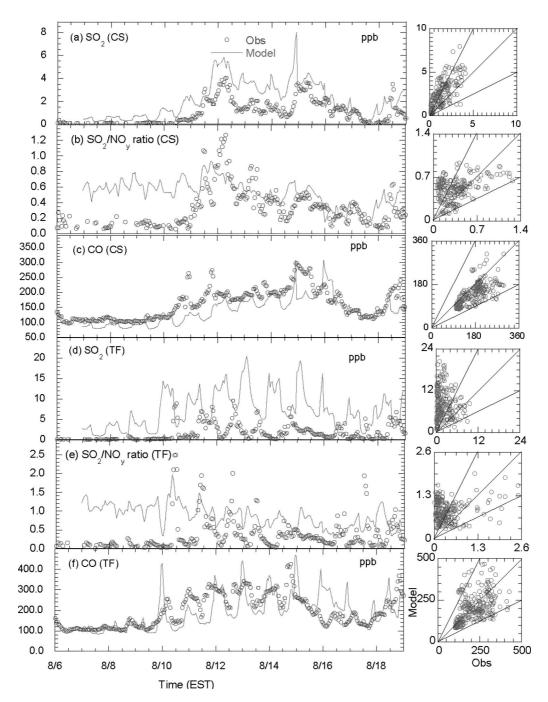


Figure 2. Time series of modeled and observed SO<sub>2</sub>, CO, and [SO<sub>2</sub>]/[NO<sub>y</sub>] ratios and their scatter plots (the 2:1, 1:1, and 1:2 lines are shown for reference) at CS and TF.

2 (also see Table 2). This indicates that the model simulated the impact at these sites of mobile sources (or urban plumes) better than that from power plant point sources during this period. To contrast the regional  $O_3$  distributions before and during the  $O_3$  episode, Figure 1 shows comparisons of the model estimations and observations at the AQS sites for  $O_3$  at 3:00 p.m. on August 8 and 12. The model reproduced the spatial distributions of observed  $O_3$  very well for the daytime (3:00 p.m.) of both August 8 and 12 with most of NMB less than  $\pm 20\%$ . The differences in meteorological conditions between August 8 and 12 can explain why  $O_3$  concentrations in the Northeast increased from near background concentrations on

August 8 to concentrations >100 ppb by August 12, because the emission forcings do not change significantly between these 2 days. Examination of the meteorological conditions impacting the Northeast during the August 8–13 period reveals a classic "transitional anticyclone" scenario<sup>9</sup> in which a clean, cold-core continental polar air mass transitions, through continued subsidence, into a warm-core, mixing-limiting air mass that is conducive to the formation of O<sub>3</sub>. At the surface, both the temperature and dew point also increase (between 5 and 10 °C) as initially north/northeasterly winds veer into the south, advecting warmer, precursor-laden air from the industrialized Midwest and Ohio River Valley. The dominant SW

Table 1. Operational evaluation on the basis of the AQS data over the Northeastern United States.

Data	RMSE (ppb)	MB (ppb)	NMB (%)	NME (%)	r
Hourly	24.23	12.55	29.77	46.87	0.63
Max 1-hr	17.65	2.32	3.24	20.05	0.70
Max 8-hr	17.03	5.74	9.02	21.86	0.70

wind on August 12 placed the Northeast United States downwind of concentrated NO<sub>x</sub> sources in the industrialized Midwest and Ohio River Valley. The buildup of the observed SO<sub>2</sub>, CO, and (SO<sub>2</sub>)/(NO<sub>v</sub>) at the CS and TF sites after August 10 in Figure 2 indicates the effects of the southwesterly flow. The air masses originating from the north and influencing these sites are not affected by large anthropogenic emission sources.<sup>10</sup>

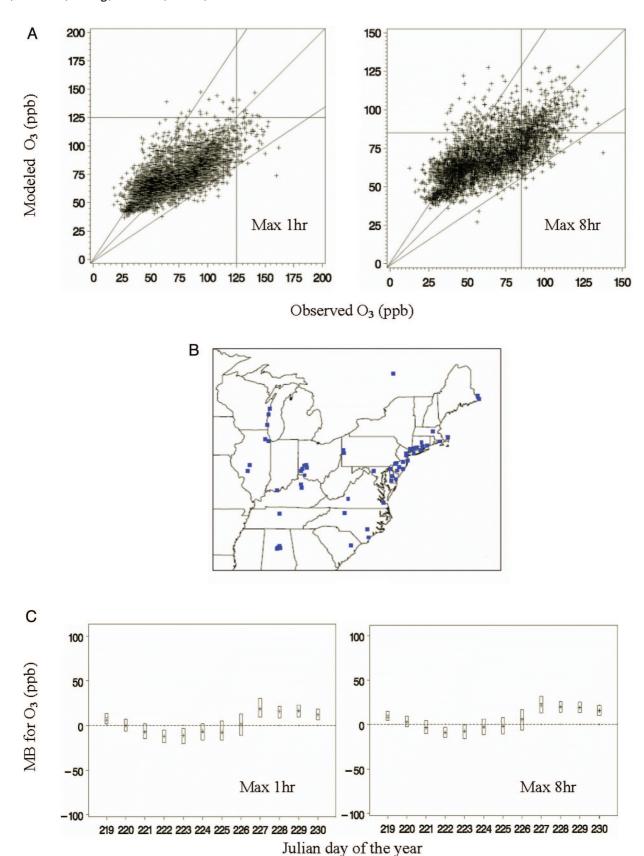
# Performance Evaluation over the Northeast U.S. **Domain at the AQS Sites**

For the model performance evaluation, we calculated summary and regression statistics along with two measures of bias, the MB and the normalized MB (NMB), and two measures of error, the root mean square error (RMSE) and normalized mean error (NME).3 Table 1 summarizes the evaluation results for the hourly, daily maximum 1-hr and maximum 8-hr O<sub>3</sub> concentrations. The recommended performance criteria for O<sub>3</sub> by EPA are: mean normalized bias  $\pm 5$  to  $\pm 15\%$ ; mean normalized gross error 30–35%; and unpaired peak estimation accuracy:  $\pm 15\%$  to  $\pm 20\%$ . The NMB (29.8%) and NME (46.9%) values for the hourly O<sub>3</sub> are higher than those performance criteria. The NMB and NME values for maximum 1-hr (maximum 8-hr) O<sub>3</sub> are 3.2% (9%) and 20.1% (21.9%), respectively, close to the performance criteria for the unpaired peak O<sub>3</sub>. Scatter plot of Figure 3A indicates that the model reproduced a majority of the observed maximum 1-hr (88.6%) and maximum 8-hr  $O_3$  (83%) within a factor of 1.5. The model generally overestimated at the low-observed O<sub>3</sub> concentration ranges but underestimated at the high-observed O<sub>3</sub> concentration ranges (see Figure 3A). The overestimations at the low  $O_3$  concentration range are, in part, because of the assumed lateral boundary conditions of 40 ppb that result in relatively higher simulated background O<sub>3</sub> levels. A closer inspection reveals that most of sites with MB >40 ppb and estimated (or observed)  $O_3$  concentration <50 ppb for the maximum 8-hr O<sub>3</sub> in Figure 3A are located within the metropolitan regions along the Washington, DC/New York, NY/Boston, MA urban corridor as shown in Figure 3B. This means that the model did not titrate  $O_3$  enough at the urban sites, causing the overestimations at the low O<sub>3</sub> concentration range. It was found that correlation coefficients (r) are between 0.5 and 0.75 at a majority of the sites with higher correlations for the maximum 1-hr and maximum 8-hr compared with the hourly estimations. In general, at sites with higher errors and bias, the model also shows poorer correlation with the observations. The time series of comparisons (data not shown) indicate that the model captured the domain mean observed hourly O<sub>3</sub> concentrations very well during the

daytime (7:00 a.m. to 4:00 p.m.) but consistently overestimated the observations during the other times (5:00 p.m. to 6:00 a.m.). Daily variations of modeled bias in Figure 3b reveal that the model overestimated (hourly, maximum 1-hr and 8-hr) at the beginning and ending dates of the study period with better performance during the middle part of the simulation. One of the reasons for this is that the  $O_3$  concentrations in the middle part of the simulation (from August 11 to August 16) were much higher than at the beginning and ending dates.

# Diagnostic Evaluation during the 2002 NEAQS

Time Series Comparison at the AIRMAP and Harvard Forest Sites. As examples, Figures 4–6 show time series comparisons and scatter plots of the model estimations and observations for O<sub>3</sub>, NO, NO<sub>2</sub>, CO, NO<sub>v</sub>, PAN, NO<sub>2</sub> photolysis rates  $(J_{NO2})$ , and meteorological parameters at the CS and HF sites (time series of CO and SO<sub>2</sub> at CS and TF sites are shown in Figure 2). Table 2 summarizes the statistical results for the model evaluations at the AIRMAP and HF sites. Following Yu et al.,8 the percentages of the comparison points where the model results are within a factor of 1.5 and 2, respectively, of the observations are listed in Table 2. Here, factor is defined as ratio of model estimation to observation if the model estimation is higher than the observation, whereas it is defined as ratio of observation to model estimation if the observation is higher than the model estimation. The model captured the hourly variations and broad synoptic changes seen in the observations of different gas species (O<sub>3</sub>, CO, NO<sub>v</sub>, and PAN; correlation coefficient >0.50; see Table 2) except NO and  $SO_2$  at each site. For  $O_3$ , the model reproduced the general temporal variations most of the time at all sites (>58% within a factor of 1.5 and >75% within a factor of 2; r >0.69; see Table 2) with better performance at the HF and MWO sites. The model estimations for NO are noticeably worse with underestimations of the observations most of the time at the CS and MWO sites, in part, possibly reflecting the inherent subgrid variability in NO emissions and concentrations that are not adequately captured by the model grid resolution. The model performance for CO at all of the sites is very good, with >90% of the comparisons within a factor of 2. For NO<sub>v</sub>, the model reproduced 76.7%, 74.1%, and 51.6% of observations within a factor of 2 at the CS, HF, and TF sites, respectively, but the model mean NO<sub>v</sub> concentrations are systematically ~50% higher than the observations at all three of the sites (see Table 2). A closer inspection of Figures 4 and 6 reveals that overestimations of peak NO<sub>v</sub> at the CS and HF sites occur during the nighttime with the largest contribution by HNO<sub>3</sub> (~20-60%) and followed by NTR ( $\sim 10-40\%$ ). In the CB-IV chemical mechanism, the species NTR represents an inert organic nitrogen terminal product. There are several reasons for these nighttime NO<sub>v</sub> overestimations, including: (1) the relatively coarse vertical resolution of the model that cannot adequately resolve sharp nocturnal gradients near the surface; (2) uncertainties associated with atmospheric sinks for the modeled organic nitrate species represented by NTR; and (3) overestimation of the gas-phase hydrolysis reaction of N<sub>2</sub>O<sub>5</sub> with water, which produces too much



**Figure 3.** (A) Scatter plots between the models and observations for maximum 1-hr and maximum 8-hr  $O_3$  concentrations with exceedance thresholds and 1:1.5, 1:1, and 1.5:1 lines indicated at the AQS sites. (B) Locations with maximum 8-hr  $O_3$  concentrations <50 ppb and MB >40 ppb. (C) Boxplots (denoting 75th, 50th, and 25th percentiles) for daily MB (MB = Model-Obs) for maximum 1-hr and maximum 8-hr  $O_3$  concentrations over the domain.

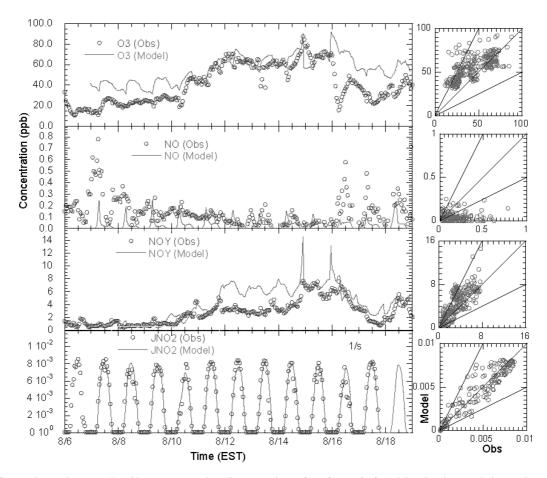


Figure 4. Time series and scatter plots (the 2:1, 1:1, and 1:2 lines are shown for reference) of model estimations and observations at the CS site.

nighttime  ${\rm HNO_3}$  in the CB-IV chemical mechanism. Although it is generally believed that the reaction of  ${\rm N_2O_5}$  with water

$$(N_2O_5 + H_2O \xrightarrow{k_1} 2HNO_3)$$
 (1)

is a major sink of  $\mathrm{NO_x}$  in the troposphere during the nighttime, the importance of this reaction is still a subject of scientific debate. Atkinson et al., ^11 for instance, suspected that this  $\mathrm{N_2O_5}$  homogeneous hydrolysis reaction does not exist at all. The CB-IV  $\mathrm{k_1}$  value of  $1.3 \times 10^{-21}$  cm³ molec. ^1 sec ^1 is close to upper limits of 1.3–1.5  $\times$ 

 $10^{-21}~\text{cm}^3~\text{molec.}^{-1}~\text{sec}^{-1}$  obtained from smog chamber studies  $^{11,12}$  but up to a factor of 5 larger than those (0.3–  $1\times10^{-21}~\text{cm}^3~\text{molec.}^{-1}~\text{sec}^{-1})$  obtained by atmospheric NO $_3$  measurements.  $^{13}$  The laboratory measurements of Wahner et al.  $^{14}$  found that the  $k_1$  value was  $2.5\times10^{-22}~\text{cm}^3~\text{molec.}^{-1}~\text{sec}^{-1}$  with a third-order reaction:

$$(N_2O_5 + 2H_2O \xrightarrow{k_2 = 1.8 \times 10^{-39} \text{ cm}^6 \text{molec.}^{-2} \text{ sec}^{-1}} 2HNO_3 + H_2O)$$
(2)

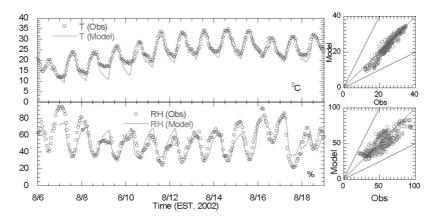


Figure 5. Same as Figure 4 but for meteorological conditions at the CS site. The 2:1, 1:1, and 1:2 lines are also shown for reference in the scatter plots.

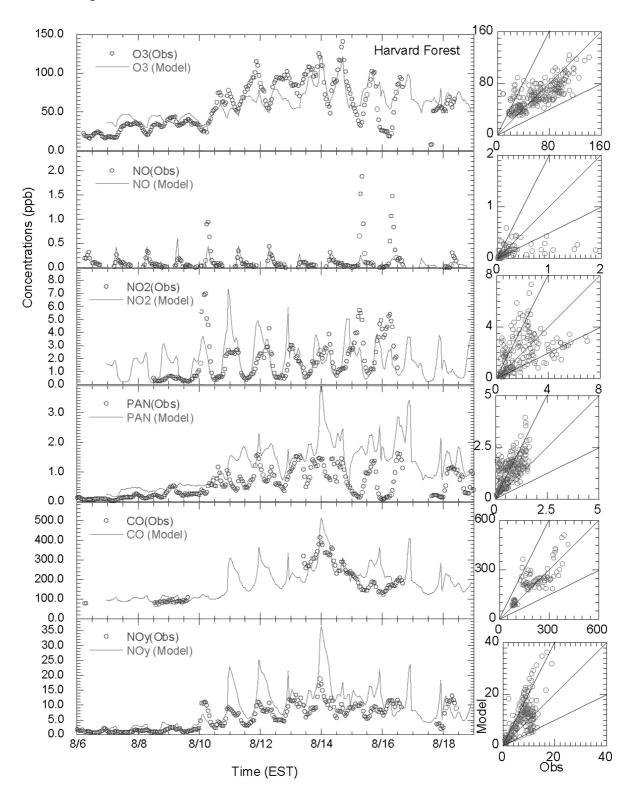


Figure 6. Same as Figure 4 but at the HF site. The 2:1, 1:1, and 1:2 lines are also shown for reference in the scatter plots.

Clearly, the high value of  $k_1$  in the model results in overestimation of nighttime  ${\rm HNO_3}$  formation, mainly causing the overprediction of nighttime  ${\rm NO_y}$  shown in Figures 4 and 6. The photolysis rates of  ${\rm NO_2}$ 

$$(NO_2 + hv \xrightarrow{J_{NO_2}} NO + O(^3P))$$
 (3)

at the CS, MWO, and TF sites were measured during the 2002 NEAQS. Following Thornton et al., 15 we focus our

analysis on daylight hours by excluding data where  $J_{NO2}$  <5 × 10<sup>-5</sup> sec<sup>-1</sup>. As shown in Figure 4 and Table 2, the model reproduced the diurnal variations of observed  $J_{NO2}$  at each site very well, with r > 0.92. Table 2 indicates that the model reproduced 77.1%, 64%, and 70.9% of observed  $J_{NO2}$  values within a factor of 1.5 at the CS, MWO, and TF sites, respectively. Demore et al. <sup>16</sup> suggest that in computing  $J_{NO2}$  values, as much as

Table 2. Statistical summaries of the comparisons of the model results with the observations at the different sites during the 2002 NEAQS.

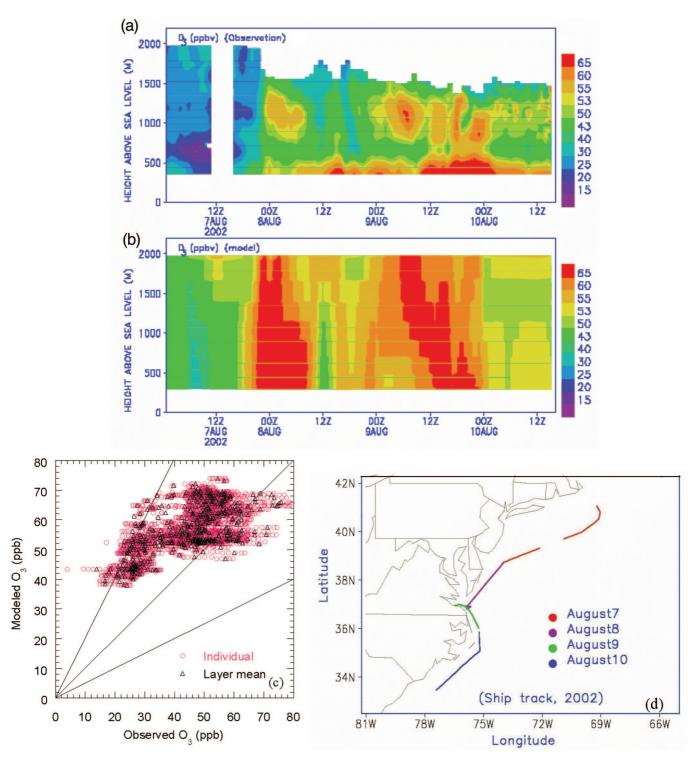
	< <b>C</b> > <sup>a</sup>				
Parameters	Observed	Model	r	% Within a Factor of 1.5 <sup>b</sup>	% Within a Factor of 2 <sup>b</sup>
CS (N = 288)					
$0_3$	42.94	55.21	0.695	58.7	86.5
NO	0.15	0.03	0.172	9.3	18.1
CO	166.31	137.96	0.724	86.5	99.0
$NO_Y$	2.91	4.15	0.800	46.5	76.7
SO <sub>2</sub>	1.29	2.46	0.751	23.4	44.7
$J_{NO2}$ (1/sec)	$4.37 \times 10^{-3}$	$4.40 \times 10^{-3}$	0.971	77.1	91.4
Temperature (C)	24.61	23.59	0.962	99.0	100.0
RH (%)	53.22	51.23	0.843	98.6	100.0
HF ( $N = 288$ )					
03	60.01	60.26	0.773	80.8	92.0
NO	0.16	0.09	0.467	36.8	55.6
$NO_2$	1.85	1.97	0.443	38.4	60.5
CO	197.75	220.55	0.804	87.8	98.0
$NO_{Y}$	6.10	9.20	0.766	40.1	74.1
PAN	0.62	1.23	0.690	24.2	47.2
RH (%)	74.08	60.65	0.882	96.6	100.0
IS $(N = 288)$					
03	56.33	69.00	0.733	65.3	86.1
CO	237.93	168.00	0.631	48.1	90.6
MWO ( $N = 288$ )					
03	60.65	53.11	0.827	86.3	99.3
NO	0.24	0.02	0.066	12.4	23.2
CO	154.59	119.68	0.639	82.7	96.0
$SO_2$	1.23	1.52	0.662	33.2	48.4
J <sub>NO2</sub> (1/sec)	$4.01 \times 10^{-3}$	$4.72 \times 10^{-3}$	0.917	64.0	77.7
TF ( $N = 288$ )					
03	44.39	56.74	0.819	59.0	76.0
NO	0.16	0.19	0.636	23.7	42.0
CO	203.46	194.24	0.550	80.5	92.7
$NO_{Y}$	4.60	9.51	0.505	25.1	51.6
SO <sub>2</sub>	1.94	7.64	0.028	13.8	18.5
J <sub>NO2</sub> (1/sec)	$4.32 \times 10^{-3}$	$4.24 \times 10^{-3}$	0.944	70.9	86.3
Temperature (C)	24.09	24.40	0.942	100.0	100.0
RH (%)	64.26	55.38	0.834	92.0	100.0

Notes: r is correlation coefficient between the model predictions and observations;  $^{a}$  <C> is the mean concentration (ppb); b Percentages (%) are the percentages of the comparison points at which model results are within a factor of 1.5 and 2 of the observations.

±20% uncertainty can arise because of uncertainties in the cross-section and quantum yield data. Additional uncertainties in the model simulations can also arise from uncertainties and errors associated with the spatial and temporal representation of cloud fields in the model and their subsequent effects on photolysis attenuation.

Figure 5 shows that the model reproduced the temporal variations of the observed temperatures and relative humidity (RH) at the CS site very well. The modeled mean temperatures (RH) at the CS and TF are 23.59 °C (51.23%) and 24.40 °C (55.4%), respectively, very close to the observations with  $\sim \pm 5\%$  errors for temperature and  $\sim \pm 10\%$ for RH as shown Table 2.

Evaluation of O<sub>3</sub> Vertical Profiles against Lidar Data. Diurnal variations in surface O<sub>3</sub> concentrations have been found to be associated with residual O<sub>3</sub> aloft.<sup>17,18</sup> Berkowitz et al.17 showed that these elevated photochemically aged layers frequently formed over the Northeastern United States during the summer and that turbulent mixing and transport led to the formation of these layers. Comparisons of modeled and ship-based lidar measurements of vertical profiles provide an assessment of the ability of the model to represent vertical structure of O<sub>3</sub> distributions. Figure 7, a and b, presents comparisons of modeled and observed time-height variations in O<sub>3</sub> structure along the ship tracks (see Figure 7d) during the period of August 7-10, 2002. The scatter plots between the observations and model estimations for individual values and model-layer means are shown in Figure 7c. As can be seen, the model captured most of the observations (70%) within a factor of 1.5, especially for the layer means (71%). The model reproduced the observed O<sub>3</sub> concentrations reasonably well at the lower altitudes between 355 and 500 m, especially after August 9, as shown in Figure 7. The mean model estimation at these altitudes (56.9  $\pm$  8.3



**Figure 7.** Vertical O<sub>3</sub> profiles for the (a) lidar observations and (b) model during the period from August 7 to 10, 2002. (c) is a scatter plot between the observations and model estimations for individual value and model-layer means (the 2:1, 1:1, and 1:2 lines are shown for reference), and (d) shows ship tracks.

ppb) is in good agreement with the observation (55.0  $\pm$  13.7 ppb). However, relative to the complex observed structures, the modeled vertical distributions of  $\rm O_3$  are more uniform with a tendency to overestimate at higher altitudes as shown in Figure 7.

 $O_3$  and CO Relationships in the Photochemically Aged Air at Each Site. Because CO is a long-lived tracer of human activity with well-known sources from combustion, industry, mobile, and oxidation of hydrocarbons,  $O_3$ -CO correlations have in the past been used to diagnose

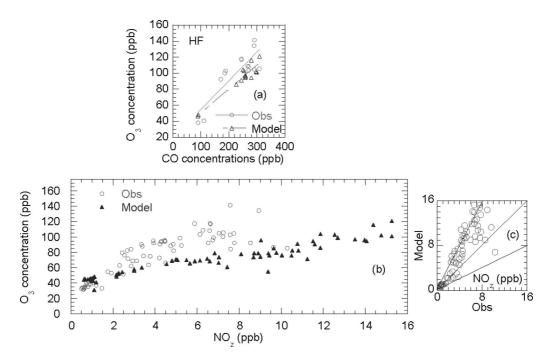


Figure 8. (a) O<sub>3</sub>-CO correlations for the model estimations and observations for the daytime period from 1:00 to 5:00 p.m.; (b) O<sub>3</sub> as a function of NO<sub>z</sub> for the NO<sub>x</sub>-limited conditions indicated by the observational data with [O<sub>3</sub>]/[NO<sub>x</sub>]>46 at the HF site and scatter plot between model and observation for NO<sub>z</sub>.

pollution influence of anthropogenic sources on O<sub>3</sub>.<sup>19</sup> Following Chin et al., 19 we only used the observed data with  $NO_x/NO_v < 0.3$  (photochemically aged rural air) and between the period from 1:00 to 5:00 p.m., when surface air is most likely representative of the boundary layer. With this selection criterion, only data for the period August 10-15 when the sites were mainly influenced by the southwesterly flow as analyzed in the previous section were left. As shown in Figure 8a and Table 3, there are strong correlations (correlation coefficient >0.86) between O<sub>3</sub> and CO for both model estimations and observations at the HF site. The slope  $(\Delta O_3/\Delta CO)$  is 0.29 for the model estimation, close to

that from the observation (0.36; see Table 3). This is in agreement with Chin et al., 19 who found that  $\Delta O_3$  $\Delta CO \approx 0.3$  was a uniform characteristic of boundary layer air over eastern North America in summer. The analyses of O<sub>3</sub> and CO at TF and CS over 2 years (2001– 1003) by Mao and Talbot<sup>20</sup> reveal that in summer during the local afternoon (1:00-6:00 p.m.) a well-defined positive  $O_3$ -CO correlation with a slope  $\sim 0.37$  existed in the air masses from the southeasterly and westerly wind sectors. Our results are also in agreement with their findings. As analyzed by Chin et al.,19 a lower limit for the net  $O_3$  production efficiency ( $\varepsilon_N$  defined as the net number of O<sub>3</sub> molecules produced per molecule

Table 3. Summary of results at the HF site.

Sites	$0_3\text{-CO}$ Correlation Equations	Correlation Coefficient	
$D_3$ -CO correlation equations ( $N = 14$ ) <sup>a</sup>			
Observation	$0_3 = 18.69 + 0.3600$	r = 0.857	
Model	$0_3 = 20.76 + 0.2900$	r = 0.950	
$O_3$ -NO <sub>z</sub> correlation equations ( $N = 82$ ) <sup>b</sup>	-		
Observation	$[0_3] = 9.1[N0_7] + 39.5$	r = 0.849	
Model	$[0_3] = 4.2[N0_7] + 41.9$	r = 0.942	
No. of hours for the different O <sub>3</sub> /NO <sub>x</sub> bins <sup>c</sup>			
0–14	28 (21) <sup>d</sup>	16 (12) <sup>e</sup>	
15–25	11 (8) <sup>d</sup>	26 (20) <sup>e</sup>	
26–45	11 (8) <sup>d</sup>	20 (15) <sup>e</sup>	
>46	83 (62) <sup>d</sup>	71 (53) <sup>e</sup>	
Total hours	133 (100) <sup>d</sup>	133 (100) <sup>e</sup>	

Notes:  ${}^{a}O_{3}$ -CO correlations in photochemically aged air (as defined by the observed  $NO_{v}/NO_{v} < 0.3$ ) for the daytime period from 1:00 to 5:00 p.m.;  ${}^{b}Correlations$ between  $O_3$  and  $NO_2$  for the  $NO_x$ -limited conditions indicated by the observational data with  $[O_3]/[NO_x] > 46$  (aged air masses); "Statistical summary of number of hours for response surface indicator ratios  $(0_3/N0_x)$  for model and observations for all days (observed-limited hours; the values in parentheses are the percentages) at the HF site during the period of August 6-17, 2002; dObservations; eModel.

of  $NO_x$  consumed) can be estimated by scaling the slope  $\Delta O_3/\Delta CO$  in the photochemically aged air to a  $CO/NO_x$  source ratio. Such an estimate is a lower limit because of deposition of  $O_3$ . According to EPA emission inventory used in the model, the mean value of the  $CO/NO_x$  emission ratio over the Midwest (the upwind region of the site) is  $8.5 \pm 2.9$ , yielding estimated  $\varepsilon_N$  lower limit values of  $2.5 \pm 0.8$  for the model and  $3.1 \pm 1.0$  for the observation.

Comparison of Modeled and Observed NO<sub>x</sub>-Sensitive Chemical Regimes, Air Mass Photochemical Age, and O<sub>3</sub> Production Efficiencies at the HF Site. Analyses of indicator ratios (i.e.,  $[O_3]/[NO_x]$ ,  $[NO_z]/[NO_v]$ , and  $[O_3]/[NO_z]$ ) can be used to evaluate the model performance in determining NO<sub>x</sub>sensitive and volatile organic compound (VOC)-sensitive chemical regimes, air mass photochemical ages, and O<sub>3</sub> production efficiency. Following Tonnesen and Dennis<sup>21</sup> and Arnold et al.,<sup>22</sup> the total hours spent in each extreme region and nearer to the  $[O_3]$  ridgeline according to the  $[O_3]/[NO_x]$  values are calculated and listed in Table 3 for the HF site.  $[O_3]/[NO_x]$  values >46 indicate strong  $NO_x$ sensitive conditions, whereas values <14 indicate VOCsensitive conditions, which are often prevalent at the surface during peak morning commutes or in fresh power plant plumes.<sup>22</sup> Table 3 reveals that for the most part, the model correctly reproduced the temporal variations in the observed  $[O_3]/[NO_x]$  ratios across the different conditions represented at the HF site. For example, both model and observations show that the site is mainly under strongly  $NO_x$ -sensitive conditions ( $\geq 53\%$ ; see Table 3).

The fraction of  $NO_y$  converted to  $NO_z$  can be used to represent the air mass photochemical age.  $[NO_z]/[NO_y]$  values <0.6 indicate a fresher  $NO_x$  plume with an increased potential for  $O_3$  production or loss, conditional on radical availability in the system, whereas higher  $[NO_z]/[NO_y]$  values indicate an aged air mass with less potential for change in  $O_3$ .<sup>22</sup> The air mass photochemical age approaches unity as  $NO_x$  is completely oxidized. It is found that the percentage of the daytime (6:00 a.m. to 6:00 p.m.) hours with air mass photochemical age values  $[NO_z]/[NO_y] > 0.6$  for the model is 87% at the HF site, close to the corresponding observed values of 82% at the HF site. These findings are similar to those reported by Olszyna et al.,<sup>23</sup> who found that >70% of the midday ratios were >0.6 at a rural site in the Eastern United States.

The  $O_3$  production efficiency  $(\varepsilon_N)$ , that is, the number of O<sub>3</sub> molecules produced for each NO<sub>x</sub> molecule processed to NO<sub>z</sub>, can be estimated by the slope of the O<sub>3</sub>-NO<sub>z</sub> correlation. Following Arnold et al.,<sup>22</sup> both modeled and observed O<sub>3</sub>-NO<sub>z</sub> slopes are obtained for only observational data with  $[O_3]/[NOx] > 46$  during the daytime (6:00 a.m. to 6:00 p.m.) at the HF site to ensure that the system is well out of the radical-sensitive region of the response surface. Table 3 indicates that there is significant correlation between O<sub>3</sub> and NO<sub>z</sub> for both model estimations and observations (r > 0.77; also see Figure 8). Note that the data during the nighttime corresponding to the NO<sub>v</sub> overestimations shown in Figure 6 are automatically not included in the O<sub>3</sub>-NO<sub>2</sub> analysis for Figure 8 and Table 3 when the above-mentioned screening criteria are used. The modeled  $\varepsilon_N$  value (4.2) at the HF site is close (~16% lower) to the lower bound of the estimated range (5-10) of other investigators<sup>23</sup> at rural sites in the Eastern United States. In contrast, the observed  $\varepsilon_N$  value (9.1) is close to the higher bound of the estimated range of other investigators. The modeled intercept is slightly higher than the observations. Chin et al.<sup>20</sup> suggest that the  $\varepsilon_N$  values estimated by the O<sub>3</sub>-NO<sub>z</sub> slopes are upper limits, because NO<sub>z</sub> species (primarily HNO<sub>3</sub>) are removed from the atmosphere more rapidly than O<sub>3</sub>. Figure 8 shows that compared with the observations, the model produced less O<sub>3</sub> at the high NO<sub>2</sub> regime. The scatter plots of Figure 8 also show that the modeled NO<sub>z</sub> concentrations were higher than the observations, indicating that the model chemistry produces more terminal oxidized nitrogen products than inferred from observations, thereby contributing to the noted underestimation of  $\varepsilon_N$ .

## **CONCLUSIONS**

The Eta-CMAQ AQF system has been developed and applied to forecast O<sub>3</sub> over the Northeastern United States during the 2002 NEAQS. The episode analysis reveals that the model captured the buildup of  $O_3$  concentrations over the NE domain from August 10 and reproduced the spatial distributions of observed O<sub>3</sub> very well for the daytime (3:00 p.m.) for both regional low (August 8) and high (August 12) with NMB values at most sites within  $\pm 20\%$ . On the basis of the evaluation at the 342 AQS sites, it is found that the model reproduced 53% of observed hourly O<sub>3</sub> within a factor of 1.5 with domain-wide NMB of 29.7% and NME of 46.9%. The comparison of modeled and lidar-based observed O<sub>3</sub> vertical profiles shows that whereas the model reproduced the observed O<sub>3</sub> concentrations well at the lower altitudes between 355 and 500 m, it tended to overestimate at higher altitudes. On the basis of the evaluation results at the four AIRMAP and HF sites, it was found that the model captured the hourly variations and broad synoptic and interday variations seen in the observations of different gas species (O<sub>3</sub>, NO<sub>2</sub>, CO, NO<sub>y</sub>, PAN, and SO<sub>2</sub>). The model reproduced 77.1%, 64%, and 70.9% of observed  $J_{NO2}$  values within a factor of 1.5 at the CS, MWO, and TF sites, respectively. On the basis of results at the HF site, it is found that both models and observations show that the site is mainly under strongly NO<sub>x</sub>-sensitive conditions (>53%). The percentage of daytime (6:00 a.m. to 6:00 p.m.) hours when the HF site was influenced by relatively chemically mature plumes ( $[NO_z]/[NO_v] > 0.6$ ) is close to that of the observations. The modeled lower limits of the  $O_3$  production efficiency  $(\varepsilon_N)$  value (2.9) are slightly lower than the observation (3.6) at the HF site estimated on the basis of relationship between  $O_3$  and CO. However, the modeled upper limit (4.2) of the  $\varepsilon_N$  values estimated by the  $O_3$ -NO<sub>z</sub> slopes is about half of the observations (9.1). There are uncertainties in the photochemical mechanism, emission inventories, and prognostic model forecasts of meteorological fields for real time. In light of these uncertainties and difficulties, the performance of the Eta-CMAQ forecast model for  $O_3$  over the Northeastern U.S. domain can be considered to be reasonable. The performance of the Eta-CMAQ forecast model over the past 3 summers is also being continuously evaluated. Detailed analyses of model estimations against extensive datasets collected during

the 2004 NEAQS campaign are under way and will provide additional evaluation of the model's estimative capability for tropospheric  $O_3$  distributions.

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