

Global Climate Change Impacts in the United States

U.S. Global Change Research Program



Final Draft for Internal
Review, April 2009

Global Climate Change Impacts in the United States



A State of Knowledge Report from the
U.S. Global Change Research Program

Publication information from Cambridge will go on this page (specific formatting requirements will be typeset by Cambridge University Press).

CAMBRIDGE UNIVERSITY PRESS
Cambridge, New York, Melbourne, Madrid, Cape Town, Singapore, São Paulo, Delhi

Cambridge University Press
32 Avenue of the Americas, New York, NY 10013-2473, USA

www.cambridge.org
Information on this title: www.cambridge.org/9780521144070

This report was produced by the Synthesis Product Development Committee, an advisory committee chartered under the Federal Advisory Committee Act, for the Subcommittee on Global Change Research, and at the request of the U.S. Government. Therefore, the report is in the public domain. Some materials used in the report are copyrighted and permission was granted to the U.S. government for their publication in this report. For commercial uses that include copyrighted materials, permission for reproduction must be sought from the copyright holder: the Subcommittee on Global Change Research. In all cases, credit must be given for copyrighted materials.

First published 2009

Printed in the United States of America

A catalog record for this publication is available from the British Library.

ISBN 978-0-521-14407-0 paperback

Cambridge University Press has no responsibility for the persistence or accuracy of URLs for external or third-party Internet Web sites referred to in this publication and does not guarantee that any content on such Web sites is, or will remain, accurate or appropriate. Information regarding prices, travel timetables, and other factual information given in this work are correct at the time of first printing, but Cambridge University Press does not guarantee the accuracy of such information thereafter.

The bar at the bottom of the front cover shows the global annual average temperature from 1900-2008, see page 17.

Federal Advisory Committee Authors

Co-Chairs and Editors-in-Chief

Thomas R. Karl,
NOAA National Climatic Data Center

Jerry M. Melillo,
Marine Biological Laboratory

Thomas C. Peterson,
NOAA National Climatic Data Center

Author Team

David M. Anderson,
NOAA World Data Center for
Paleoclimatology

Donald F. Boesch,
University of Maryland Center for
Environmental Science

Virginia R. Burkett,
U.S. Geological Survey

Lynne M. Carter,
Adaptation Network

Stewart J. Cohen,
Environment Canada and University of
British Columbia

Nancy B. Grimm,
Arizona State University

Jerry L. Hatfield,
U.S. Department of Agriculture

Katharine Hayhoe,
Texas Tech University

Anthony C. Janetos,
Pacific Northwest National Laboratory

Jack A. Kaye,
National Aeronautics and Space
Administration

Jay H. Lawrimore,
NOAA National Climatic Data Center

James J. McCarthy,
Harvard University

A. David McGuire,
U.S. Geological Survey/University of
Alaska Fairbanks

Edward L. Miles,
University of Washington

Evan Mills,
Lawrence Berkeley National Laboratory

Jonathan T. Overpeck,
University of Arizona

Jonathan A. Patz,
University of Wisconsin at Madison

Roger S. Pulwarty,
NOAA Climate Program Office and Earth
System Research Laboratory

Benjamin D. Santer,
Lawrence Livermore National Laboratory

Michael J. Savonis,
U.S. Department of Transportation

H. Gerry Schwartz, Jr.,
Consultant/Transportation

Eileen L. Shea,
NOAA National Climatic Data Center/
Integrated Data and Environmental
Applications Center

John M.R. Stone,
Carleton University

Bradley H. Udall,
University of Colorado/NOAA Earth
System Research Laboratory

John E. Walsh,
University of Alaska Fairbanks

Michael F. Wehner,
Lawrence Berkeley National Laboratory

Thomas J. Wilbanks,
Oak Ridge National Laboratory

Donald J. Wuebbles,
University of Illinois

Senior Science Writer and Lead Graphic Designer

Senior Science Writer

Susan J. Hassol,
Climate Communication, LLC

Lead Graphic Designer

Sara W. Veasey,
NOAA National Climatic Data Center

Key Support Personnel

Jessica Blunden, Editorial Assistant,
STG, Inc.

Marta Darby, Copy Editor, STG, Inc.

David Dokken, CCSP Technical
Advisor, USGCRP

Byron Gleason, Data Analysis/Visualization,
NOAA National Climatic Data Center

Glenn M. Hyatt, Graphics Support,
NOAA National Climatic Data Center

Clare Keating, Editorial Support
Texas Tech Univ.

Staci Lewis, Technical Advisor
NOAA

Jolene McGill, Logistical Support Assistant,
NOAA National Climatic Data Center

Deborah J. Misch, Graphics Support,
STG, Inc.

William Murray, Technical Support,
STG, Inc.

Tim Owen, Logistical Support,
NOAA National Climatic Data Center

Susan Osborne, Copy Editor,
STG, Inc.

Deborah Riddle, Graphics Support,
NOAA National Climatic Data Center

Susanne Skok, Copy Editor,
STG, Inc.

Michael Squires, Cartographic Support,
NOAA National Climatic Data Center

Mara Sprain, Editorial Assistant,
STG, Inc.

Jeff VanDorn, Technical and Graphics Support,
ATMOS Research

David Wuertz, Data Analysis/Visualization,
NOAA National Climatic Data Center

Christian Zamarrá, Graphics Support,
STG, Inc.

Federal Executive Team

Acting Director, U.S. Global Change Research Program: Jack A. Kaye

Director, U.S. Global Change Research Program Office: Peter A. Schultz

Lead Agency Principal Representative to CCSP (through January 2009),
National Oceanic and Atmospheric Administration: Mary M. Glackin

Lead Agency Principal Representative to CCSP,
Product Lead, National Oceanic and Atmospheric Administration: Thomas R. Karl

Lead Agency Principal Representative to CCSP
Group Chair Synthesis and Assessment Products, Environmental Protection Agency: Michael W. Slimak

Synthesis and Assessment Product Coordinator,
U.S. Global Change Research Program Office: Fabien J.G. Laurier

Communications Advisor/Coordinator/Editor,
U.S. Global Change Research Program Office: Anne M. Waple

Special Advisor, National Oceanic and Atmospheric Administration: Chad A. McNutt

Federal Advisory Committee Designated Federal Official,
National Oceanic and Atmospheric Administration: Christopher D. Miller

Reviewers

Blue Ribbon Reviewers

Robert W. Corell,
Global Change Program, H. John Heinz III
Center for Science, Economics and the
Environment

Robert A. Duce,
Department of Atmospheric Sciences,
Texas A&M University

Kristie L. Ebi,
Independent consultant, ESS, LLC
Alexandria, VA

Christopher B. Field,
Carnegie Institution

William H. Hooke,
Atmospheric Policy Program, American
Meteorological Society

Michael C. MacCracken,
Climate Institute

Linda O. Mearns,
Environmental and Societal Impacts
Group, National Center for Atmospheric
Research

Gerald A. Meehl,
Climate and Global Dynamics Division,
National Center for Atmospheric Research

John Reilly,
Sloan School of Management,
Massachusetts Institute of Technology

Susan Solomon,
NOAA, Earth System Research
Laboratory

Steven C. Wofsy,
Harvard University

Communication Reviewers

Robert Henson,
University Corporation for Atmospheric Research

Jack W. Williams,
American Meteorological Society

This report on Global Climate Change Impacts in the United States was prepared in accordance with Section 515 of the Treasury and General Government Appropriations Act for Fiscal Year 2001 (Public Law 106-554) and the information quality act guidelines issued by the Department of Commerce and NOAA pursuant to Section 515 <<http://www.noanews.noaa.gov/stories/iq.htm>>.

Final Draft for Internal Review, April 2009

Do Not Cite Or Quote

TRANSMITTAL LETTER HERE

DRAFT



About this Report 7



Executive Summary 9



Global Climate Change 13



National Climate Change 27

Climate Change Impacts by Sector



Water Resources 41



Energy Supply and Use 53



Transportation 61



Agriculture 71



Ecosystems 79







Human Health 89



Society 99



Regional Climate Change Impacts

	Northeast.....	107
	Southeast.....	111
	Midwest.....	117
	Great Plains.....	123
	Southwest.....	129
	Northwest.....	135
	Alaska.....	139
	Islands.....	145
	Coasts.....	149



An Agenda for Climate Impacts Science.....	153
---	------------



Concluding Thoughts.....	157
---------------------------------	------------

Author Team Biographies.....	159
-------------------------------------	------------

Primary Sources of Information.....	161
--	------------

Acronyms.....	164
----------------------	------------

References.....	165
------------------------	------------

Photography Credits.....	202
---------------------------------	------------

Recommended citation to be placed here.
(Maybe a photo)

DRAFT

About this Report

What is this report?

This report summarizes the science of climate change and the impacts of climate change on the United States, now and in the future. It is largely based on results of the U.S. Climate Change Science Program (CCSP),^a and integrates those results with related research from around the world. This report discusses climate-related impacts for various societal and environmental sectors and regions across the nation. It is an authoritative scientific report written in plain language, with the goal of better informing public and private decision making at all levels.

Who called for it, who wrote it, and who approved it?

The CCSP called for this report. An expert team of scientists operating under the authority of the Federal Advisory Committee Act, assisted by communication specialists, wrote the document. The report was extensively reviewed and revised based on comments from experts and the public. The report was approved by its lead CCSP Agency, the National Oceanic and Atmospheric Administration, the other CCSP agencies, and the Committee on the Environment and Natural Resources on behalf of the National Science and Technology Council.^b This report meets all Federal requirements associated with the Information Quality Act, including those pertaining to public comment and transparency.

What are its sources?

The report draws from a large body of scientific information. The foundation of this report is the set of 21 Synthesis and Assessment Products of the CCSP, which were designed to address key policy-relevant issues in climate science (see page 161); several of these were also summarized in the *Scientific Assessment of the Effects of Climate Change on the United States* published by the CCSP in 2008. In addition, other peer-reviewed scientific assessments were used, including those of the Intergovernmental Panel on Climate Change, the U.S. National Assessment of the Consequences of Climate Variability and Change, the Arctic Climate Impact Assessment, the National Research Council's Transportation Research Board report on the Potential Impacts of Climate Change and U.S. Transportation, and a variety of regional climate impact assessments. These assessments were augmented with government statistics as necessary (such as population census and energy usage) as well as publicly available observations and peer-reviewed research published through the end of 2008. This new work was carefully selected by the author team with advice from expert reviewers to update key aspects of climate change science relevant to this report. The icons on the bottom of this page represent some of the major sources drawn upon for this synthesis report.

On the first page of each major section, the sources primarily drawn upon for that section are shown using these icons. Endnotes, indicated by superscript numbers and compiled at the end of the book, are used for specific references throughout the report.



^a In this report, the Climate Change Science Program (CCSP) is used interchangeably with the U.S. Global Change Research Program which was established in 1990 by the Global Change Research Act.

^b A description of the National Science and Technology Council (NSTC) can be found at www.ostp.gov/cs/nstc.



Does this report deal with options for responding to climate change?

L1 While the primary focus of this report is on the
 L2 impacts of climate change in the United States,
 L3 it also deals with some of the actions society is
 L4 already taking or can take to respond to the climate
 L5 challenge. Responses to climate change fall into
 L6 two broad categories. The first involves “mitiga-
 L7 tion” measures to reduce climate change by reduc-
 L8 ing emissions of heat-trapping gases and particles,
 L9 or increasing removal of heat-trapping gases from
 L10 the atmosphere. The second involves “adapta-
 L11 tion” measures to improve our ability to cope with
 L12 or avoid harmful impacts and take advantage of
 L13 beneficial ones, now and in the future. Both of these
 L14 are necessary elements of an effective response
 L15 strategy. These two types of responses are linked in
 L16 that more effective mitigation measures reduce the
 L17 amount of climate change, and therefore the need
 L18 for adaptation.
 L19

L20 This report underscores the importance of mitiga-
 L21 tion by comparing impacts resulting from higher
 L22 versus lower emissions scenarios. The report shows
 L23 that choices made about emissions in the next few
 L24 decades will have far-reaching consequences for
 L25 climate change impacts. Over the long term, lower
 L26 emissions will lessen both the magnitude of climate
 L27 change impacts and the rate at which they appear.
 L28

L29 While the report underscores the importance of
 L30 mitigation as an essential part of the Nation’s cli-
 L31 mate change strategy, it does not evaluate mitigation
 L32 technologies or undertake an analysis of the effec-
 L33 tiveness of various approaches. These issues are the
 L34 subject of ongoing studies by the U.S. Government’s
 L35 Climate Change Technology Program and sev-
 L36 eral federal agencies including the Department of
 L37 Energy, Environmental Protection Agency, National
 L38 Oceanic and Atmospheric Administration, and
 L39 Department of Agriculture. The range of mitiga-
 L40 tion responses being studied includes more efficient
 L41 production and use of energy, increased use of non-
 L42 carbon-emitting energy sources, and carbon capture
 L43 and storage.
 L44

L45 Adaptation options also have the potential to moder-
 L46 ate harmful impacts of current and future climate
 L47 variability and change. While this report does ad-
 L48
 L49
 L50
 L51

R1 dress adaptation, it does not do so comprehensively.
 R2 Rather, in the context of impacts, this report identi-
 R3 fies examples of actions currently being pursued
 R4 in various sectors and regions to address climate
 R5 change, as well as other environmental problems
 R6 that could be exacerbated by climate change such as
 R7 urban air pollution and heat waves. In most cases,
 R8 there is currently insufficient peer-reviewed infor-
 R9 mation to evaluate the practicality, effectiveness,
 R10 costs, or benefits of these measures, highlighting a
 R11 need for research in this area. Thus, the discussion
 R12 of various public and private adaptation examples
 R13 should not be viewed as an endorsement of any
 R14 particular option, but rather as illustrative examples
 R15 of approaches being tried.
 R16

How is the likelihood of various outcomes expressed given that the future is not certain?

R17 With regard to expressing the range of possible
 R18 outcomes and identifying the likelihood of par-
 R19 ticular impacts, this report takes a plain-language
 R20 approach to expressing the expert judgment of the
 R21 author team based on the best available evidence.
 R22 For example, an outcome termed “likely” has at
 R23 least a two-thirds chance of occurring; an outcome
 R24 termed “very likely,” at least a 90 percent chance.¹
 R25 In using these terms, the Federal Advisory Com-
 R26 mittee has taken into consideration a wide range of
 R27 information, including the strength and consistency
 R28 of the observed evidence, the range and consistency
 R29 of model projections, the reliability of particular
 R30 models as tested by various methods, and most
 R31 importantly, the body of work addressed in earlier
 R32 synthesis and assessment reports. Key sources of
 R33 information used to develop these characterizations
 R34 of uncertainty are referenced in endnotes.
 R35

How does this report address incomplete scientific understanding?

R36 This assessment identifies areas in which scientific
 R37 uncertainty limits our ability to estimate future
 R38 climate change and its impacts. The section on *An*
 R39 *Agenda for Climate Impacts Science* at the end of
 R40 this report highlights some of these areas.
 R41
 R42
 R43
 R44
 R45
 R46
 R47
 R48
 R49
 R50
 R51

Executive Summary



L1
L2
L3
L4
L5
L6
L7
L8
L9
L10
L11
L12
L13
L14
L15
L16
L17
L18
L19
L20
L21
L22
L23
L24
L25
L26
L27
L28
L29
L30
L31
L32
L33
L34
L35
L36
L37
L38
L39
L40
L41
L42
L43
L44
L45

Observations show that warming of the climate is unequivocal. The global warming observed over the past 50 years is due primarily to human-induced emissions of heat-trapping gases. These emissions come mainly from the burning of fossil fuels (coal, oil, and gas), with additional contributions from the clearing of forests and agricultural activities.

R1
R2
R3
R4
R5
R6
R7
R8
R9
R10
R11
R12
R13
R14
R15
R16
R17
R18
R19
R20
R21
R22
R23
R24
R25
R26
R27
R28
R29
R30
R31
R32
R33
R34
R35
R36
R37
R38
R39
R40
R41
R42
R43
R44
R45

Warming over this century is projected to be considerably greater than over the last century. The global average temperature since 1900 has risen by about 1.5°F. By 2100, it is projected to rise another 2 to 10°F. The U.S. average temperature has risen by a comparable amount and is very likely to rise more than the global average over this century, with some variation from place to place. Several factors will determine future temperature increases. Increases at the lower end of this range are more likely if global heat-trapping gas emissions are cut substantially. If emissions continue to rise at or near current rates, temperature increases are more likely to be near the upper end of the range. Volcanic eruptions or other natural variations

could temporarily mask some of the human-induced warming, but these effects would only last a few years.

Reducing emissions of carbon dioxide would lessen warming over this century and beyond. Sizeable early cuts in emissions would significantly reduce the pace and the overall amount of climate change. Earlier cuts in emissions would have a greater effect in reducing climate change than comparable reductions made later. In addition, reducing emissions of some shorter-lived heat-trapping gases, such as methane, and some types of particles, such as soot, would begin to reduce warming within weeks to decades.

Climate-related changes have already been observed globally and in the United States. These include increases in air and water temperatures, reduced frost days, increased frequency and intensity of heavy downpours, a rise in sea level, and reduced snow cover, glaciers, permafrost, and sea ice. A longer ice-free period on lakes and rivers, lengthening of the growing season, and increased water vapor in the atmosphere have also been observed. Over the past 30 years, temperatures have risen faster in winter than in any other season, with winter temperatures in the Midwest and northern Great Plains increasing more than 7°F. Some of the changes have been faster than previous assessments had suggested.

These climate-related changes are expected to continue while new ones develop. Likely future changes for the United States and surrounding coastal waters include more intense hurricanes with related increases in wind, rain, and storm surges (but not necessarily an increase in the number of these storms that make landfall), as well as drier conditions in the Southwest and Caribbean. These changes will impact human health, water supply, agriculture, coastal areas, and many other aspects of society and the natural environment.

This report synthesizes information from a wide variety of scientific assessments (see page 7) and recently published research to summarize what is known about the observed and projected consequences of climate change on the United States. It combines analysis of impacts on various sectors



L1 such as energy, water, and transportation at the
 L2 national level with an assessment of key impacts on
 L3 specific regions of the United States. For example,
 L4 sea-level rise will increase risks of erosion, storm
 L5 surge damage, and flooding for coastal communi-
 L6 ties, especially in the Southeast and parts of Alaska.
 L7 Reduced snowpack will alter the timing and amount
 L8 of water supplies, exacerbating water scarcity in the
 L9 West.

L10
 L11 Society and ecosystems can adjust to some climatic
 L12 changes, but this takes time. The projected rapid
 L13 rate and large amount of climate change over this
 L14 century will challenge the ability of society and
 L15 natural systems to adapt. For example, it is difficult
 L16 and expensive to alter or replace infrastructure
 L17 designed to last for decades (such as buildings,
 L18 bridges, roads, airports, reservoirs, and ports) in re-
 L19 sponse to continuous and/or abrupt climate change.
 L20 Impacts are expected to become increasingly severe
 L21 for more people and places as the amount of warm-
 L22 ing increases. Rapid rates of warming would lead
 L23 to particularly large impacts on natural ecosystems
 L24 and the benefits they provide to humanity. Some of
 L25 the impacts of climate change will be irreversible,
 L26 such as species extinctions and coastal land lost to
 L27 rising seas.

L28
 L29 Unanticipated impacts of increasing carbon dioxide
 L30 and climate change have already occurred and
 L31 more are possible in the future. For example, it has
 L32 recently been observed that the increase in atmo-
 L33 spheric carbon dioxide concentration is causing an
 L34 increase in ocean acidity. This reduces the ability of
 L35 corals and other sea life to build shells and skeletons
 L36 out of calcium carbonate. Additional impacts in the
 L37 future might stem from unforeseen changes in the
 L38 climate system, such as major alterations in oceans,
 L39 ice, or storms; and unexpected consequences of
 L40 ecological changes, such as massive dislocations
 L41 of species or pest outbreaks. Unexpected social or
 L42 economic changes, including major shifts in wealth,
 L43 technology, or societal priorities would also affect
 L44 our ability to respond to climate change. Both
 L45 anticipated and unanticipated impacts become more
 L46 challenging with increased warming.

L47
 L48 Projections of future climate change come from
 L49 careful analyses of outputs from global climate
 L50

models run on the world’s most advanced comput-
 ers. The model simulations analyzed in this report
 used plausible scenarios of human activity that
 generally lead to further increases in heat-trapping
 emissions. None of the scenarios used in this report
 assumes adoption of policies explicitly designed to
 address climate change. However, the level of emis-
 sions varies among scenarios because of differences
 in assumptions about population, economic activity,
 and choice of energy technologies. Scenarios cover
 a range of emissions of heat-trapping gases, and
 the associated climate projections illustrate that
 lower emissions result in less climate change and
 thus reduced impacts over this century and beyond.
 Under all scenarios considered in this report,
 however, relatively large and sustained changes in
 many aspects of climate are projected by the middle
 of this century, with even larger changes by the end
 of this century, especially under higher emission
 scenarios.

In projecting future conditions, there is always
 some level of uncertainty. For example, there is a
 high degree of confidence in projections that future
 temperature increases will be greatest in the Arctic
 and in the middle of continents. For precipitation,
 there is high confidence in projections of continued
 increases in the Arctic and sub-Arctic (including
 Alaska) and decreases in the regions just outside
 the tropics, but the precise location of the transition
 between these is less certain. At local to regional
 scales and on time frames up to a few years, natural
 climate variations can be relatively large and can
 temporarily mask the progressive nature of global
 climate change. However, the science of making
 skillful projections at these scales has progressed
 considerably, allowing useful information to be
 drawn from regional climate studies such as those
 highlighted in this report.

This report focuses on observed and projected
 climate change and its impacts on the United States.
 However, a discussion of these issues would be
 incomplete without mentioning some of the actions
 society can take to respond to the climate chal-
 lenge. The two major categories are “mitigation”
 and “adaptation.” Mitigation refers to options for
 reducing heat-trapping emissions such as carbon
 dioxide, methane, nitrous oxide, and halocarbons.

R1
 R2
 R3
 R4
 R5
 R6
 R7
 R8
 R9
 R10
 R11
 R12
 R13
 R14
 R15
 R16
 R17
 R18
 R19
 R20
 R21
 R22
 R23
 R24
 R25
 R26
 R27
 R28
 R29
 R30
 R31
 R32
 R33
 R34
 R35
 R36
 R37
 R38
 R39
 R40
 R41
 R42
 R43
 R44
 R45
 R46
 R47
 R48
 R49
 R50

L1 Adaptation refers to changes made to better respond
L2 to present or future climatic and other environmen-
L3 tal conditions, thereby reducing harm or taking
L4 advantage of opportunity. Effective mitigation
L5 measures reduce the need for adaptation. Mitigation
L6 and adaptation are both essential parts of a compre-
L7 hensive climate change response strategy.
L8

L9 Carbon dioxide emissions are a primary focus of
L10 mitigation strategies. These include improving
L11 energy efficiency, using energy sources that do not
L12 produce carbon dioxide or produce less of it, captur-
L13 ing and storing carbon dioxide from fossil fuel use,
L14 and so on. Choices made about emissions reductions
L15 now and over the next few decades will have far-
L16 reaching consequences for climate-change impacts.
L17 The importance of mitigation is clear in compari-
L18 sons of impacts resulting from higher versus lower
L19 emissions scenarios considered in this report. Over
L20 the long term, lower emissions will lessen both the
L21 magnitude of climate-change impacts and the rate
L22 at which they appear. Smaller climate changes that
L23 come more slowly make the adaptation challenge
L24 more tractable.
L25

L26 However, no matter how aggressively heat-trapping
L27 emissions are reduced, some amount of climate
L28 change and resulting impacts will continue due to
L29 the effects of gases that have already been released.
L30 This is true for several reasons. First, some of these
L31 gases are very long-lived and the levels of atmo-
L32 spheric heat-trapping gases will remain elevated for
L33 hundreds of years or more. Second, the Earth's vast
L34 oceans have absorbed much of the heat added to the
L35 climate system due to the increase in heat-trapping
L36 gases, and will retain that heat for many decades.
L37 In addition, the factors that determine emissions,
L38 such as energy-supply systems, cannot be changed
L39 overnight. Consequently, there is also a need for
L40 adaptation.
L41

L42 Adaptation can include a wide range of activities.
L43 Examples include a farmer switching to growing
L44 a different crop variety better suited to warmer or
L45 drier conditions; a company relocating key busi-
L46 ness centers away from coastal areas vulnerable
L47 to sea-level rise and hurricanes; and a community
L48 altering its zoning and building codes to place fewer
L49 structures in harm's way and making buildings
L50

less vulnerable to damage from floods, fires, and
other extreme events. Some adaptation options that
are currently being pursued in various regions and
sectors to deal with climate change and/or other
environmental issues are identified in this report.
However, it is clear that there are limits to how
much adaptation can achieve.

Humans have adapted to changing conditions in the
past, but in the future, adaptations will be particu-
larly challenging because society won't be adapting
to a new steady state but rather to a moving target.
Climate will be continually changing, moving at a
relatively rapid rate, outside the range to which soci-
ety has adapted. The precise amounts and timing of
these changes will not be known with certainty.

In an increasingly interdependent world, U.S.
vulnerability to climate change is linked to the fates
of other nations. For example, conflicts or mass
migrations of people resulting from food scarcity
and other resource limits, health, or environmental
stresses in other parts of the world could threaten
national security. It is thus difficult to fully evaluate
the impacts of climate change on the United States
without considering the consequences of climate
change elsewhere. However, such analysis is beyond
the scope of this report.

Finally, this Assessment identifies a number of ar-
eas in which inadequate information or understand-
ing hampers our ability to estimate likely future
climate change and its impacts. For example, our
knowledge of changes in tornadoes, hail, and ice
storms is quite limited, making it difficult to know
if and how such events have changed as climate has
warmed, and how they might change in the future.
Research on ecological responses to climate change
is also limited, as is our understanding of social
responses. The section titled *An Agenda for Climate
Impacts Science* at the end of this report offers some
thoughts on the most important ways to improve our
knowledge. Results from such efforts would inform
future assessments that continue building our
understanding of humanity's impacts on climate,
and climate's impacts on us.

R1
R2
R3
R4
R5

R6

R7

R8

R9

R10

R11

R12

R13

R14

R15

R16

R17

R18

R19

R20

R21

R22

R23

R24

R25

R26

R27

R28

R29

R30

R31

R32

R33

R34

R35

R36

R37

R38

R39

R40

R41

R42

R43

R44

R45

R46

R47

R48

R49

R50

Key Findings

1. Global warming is unequivocal and primarily human-induced.

There is no question that global temperature has increased over the past 50 years. This observed increase is due primarily to human-induced emissions of heat-trapping gases. (p. 13)

2. Climate changes are underway in the United States and are projected to grow.

Climate-related changes are already observed in the United States and its coastal waters. These include increases in heavy downpours, rising temperature and sea level, rapidly retreating glaciers, thawing permafrost, lengthening growing seasons, lengthening ice-free seasons in the ocean and on lakes and rivers, earlier snowmelt, and alterations in river flows. These changes are projected to grow. (p. 27)

3. Widespread climate-related impacts are occurring now and are expected to increase.

Climate changes are already affecting water, energy, transportation, agriculture, ecosystems, and health. These impacts are different from region to region and will grow under projected climate change. (p. 41-106, 107-152)

4. Climate change will stress water resources.

Water is an issue in every region, but the nature of the potential impacts varies. Drought, related to reduced precipitation, increased evaporation, and increased water loss from plants, is an important issue in many regions, especially in the West. Floods and water quality problems are likely to be amplified by climate change in most regions. Declines in mountain snowpack are important in the West and Alaska where snowpack provides vital natural water storage. (p. 41, 129, 135, 139)

5. Crop and livestock production will be increasingly challenged.

Agriculture is considered one of the sectors most adaptable to changes in climate. However, increased heat, pests, water stress, diseases, and weather extremes will pose adaptation challenges for crop and livestock production. (p. 71)

6. Coastal areas are at increasing risk from sea-level rise and storm surge.

Sea-level rise and storm surge place many U.S. coastal areas at increasing risk of erosion and flooding, especially along the Atlantic and Gulf Coasts, Pacific Islands, and parts of Alaska. Energy and transportation infrastructure and other property in coastal areas are very likely to be adversely affected. (p. 111, 139, 145, 149)

7. Threats to human health will increase.

Health impacts of climate change are related to heat stress, water-borne diseases, poor air quality, extreme weather events, and diseases transmitted by insects and rodents. Robust public health infrastructure can reduce the potential for negative impacts. (p. 89)

8. Climate change will interact with many social and environmental stresses.

Climate change will combine with pollution, population growth, overuse of resources, urbanization, and other social, economic, and environmental stresses to create larger impacts than from any of these factors alone. (p. 99)

9. Tipping points have already been reached and have led to large changes.

Changes in climate have pushed ecosystems beyond tipping points. With further climate change, if more tipping points are crossed, additional important services that ecosystems provide to society will be diminished. (p. 82, 115, 155)

10. Future climate change and its impacts depend on choices made today.

The amount and rate of future climate change depend primarily on current and future human-caused emissions of heat-trapping gases and airborne particles. Responses involve reducing emissions to limit future warming, and adapting to the changes that are unavoidable. (p. 25, 29)



Global Climate Change

Key Messages:

- Human activities have led to large increases in heat-trapping gases over the past century.
- Over the last 50 years, global average temperature and sea level have increased, and precipitation patterns have changed.
- The global warming of the past 50 years is due primarily to human-induced increases in heat-trapping gases. Human “fingerprints” also have been identified in many other aspects of the climate system, including changes in ocean heat content, precipitation, atmospheric moisture, and arctic sea ice.
- Global temperatures are projected to continue to rise over this century; by how much and for how long depends on a number of factors, including the amount of heat-trapping gas emissions and how sensitive the climate is to those emissions.

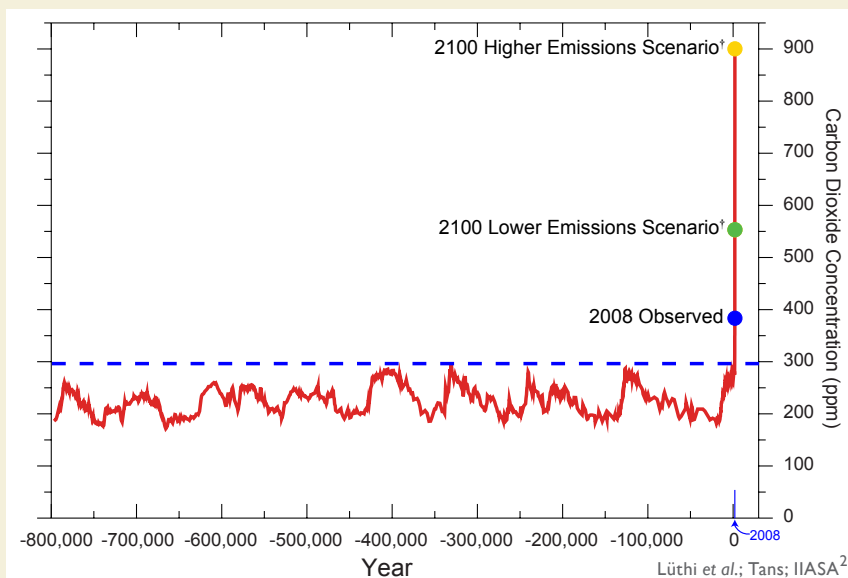
Key Sources



This introduction to global climate change explains very briefly what has been happening to the world’s climate and why, and what is projected to happen in the future. While this report focuses on climate change impacts in the United States, understanding these changes and their impacts requires an understanding of the global climate system.

Many changes have been observed in global climate over the past century. The nature and causes of these changes have been comprehensively chronicled in a variety of recent reports, such as those by the Intergovernmental Panel on Climate Change (IPCC) and the U.S. Climate Change Science Program (CCSP). This section does not intend to duplicate these comprehensive efforts, but rather to provide a brief synthesis, and to integrate more recent work with the assessments of the IPCC, CCSP, and others.

800,000 Year Record of Carbon Dioxide Concentration



Analysis of air bubbles trapped in an Antarctic ice core extending back 800,000 years documents the Earth’s changing carbon dioxide concentration. Over this long period, natural factors have caused the atmospheric carbon dioxide concentration to vary within a range of about 170 to 300 parts per million (ppm). Temperature-related data make clear that these variations have played a central role in determining the global climate. As a result of human activities, the present carbon dioxide concentration of about 385 ppm is about 30 percent above its highest level over at least the last 800,000 years. In the absence of strong control measures, emissions projected for this century would result in the carbon dioxide concentration increasing to a level that is roughly 2 to 3 times the highest level occurring over the glacial-interglacial era that spans the last 800,000 or more years.

Human activities have led to large increases in heat-trapping gases over the past century.

The Earth’s climate depends on the functioning of a large natural “greenhouse effect.” This effect is the result of heat-trapping gases (also known as greenhouse gases) like water vapor, carbon dioxide, ozone, methane, and nitrous oxide, which absorb heat radiated from the Earth’s surface and lower atmosphere and then radiate much of the energy back towards the surface. Without this natural greenhouse effect, the average surface temperature of the Earth would be about 60°F colder. However, human activities have been releasing additional heat-trapping gases, intensifying the natural greenhouse effect, thereby changing the Earth’s climate.

Climate is influenced by a variety of factors, both human-induced and natural. The increase in the carbon dioxide concentration has been the principal factor causing warming over the past 50 years. Its concentration has been building up in the Earth’s atmosphere since the beginning of the industrial era in the mid-1700s, primarily due to the burning of fossil fuels (coal, oil, and natural gas) and the clearing of forests. Human activities have also increased the emissions of other greenhouse gases, such as methane, nitrous oxide, and halocarbons.³

These emissions are thickening the blanket of heat-trapping gases in Earth’s atmosphere, causing surface temperatures to rise.

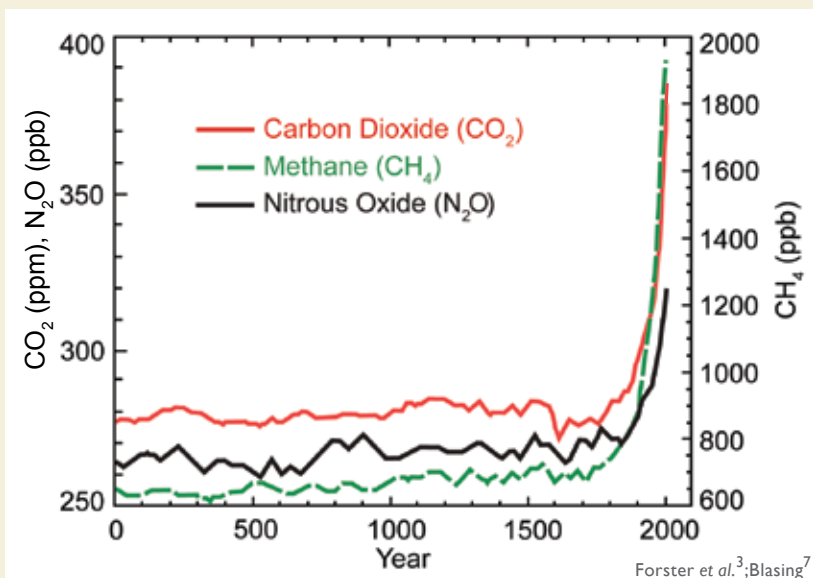
Heat-trapping gases

Carbon dioxide concentration has increased due to the use of fossil fuels in electricity generation, transportation, and industrial and household uses. It is also produced as a by-product during the manufacturing of cement. Deforestation provides a source of carbon dioxide and reduces its uptake by trees and other plants. Globally, over the past several decades, about 80 percent of human-induced carbon dioxide emissions came from the burning of fossil fuels, while about 20 percent resulted from deforestation and associated agricultural practices. The concentration of carbon dioxide in the atmosphere has increased by roughly 35 percent since the start of the industrial revolution.³

Methane concentration has increased mainly as a result of agriculture, raising livestock (which produce methane in their digestive tracts), mining, transportation, and use of certain fossil fuels, sewage, and decomposing garbage in landfills. About 70 percent of the emissions of atmospheric methane are now related to human activities.⁴

Nitrous oxide concentration is increasing as a result of fertilizer use and fossil fuel burning.

2,000 Years of Greenhouse Gas Concentrations



Increases in concentrations of these gases since 1750 are due to human activities in the industrial era. Concentration units are parts per million (ppm) or parts per billion (ppb), indicating the number of molecules of the greenhouse gas per million or billion molecules of air.

Halocarbon emissions come from the release of manufactured chemicals to the atmosphere. Examples include chlorofluorocarbons (CFCs), which were used extensively in refrigeration and for other industrial processes before their presence in the atmosphere was found to cause stratospheric ozone depletion. The abundance of these gases in the atmosphere is now decreasing as a result of international regulations designed to protect the ozone layer. Continued decreases in ozone-depleting halocarbon emissions are expected to reduce their relative influence on climate change in the future.^{3,5} Many halocarbon replacements, however, are potent greenhouse gases, and their concentrations are increasing.⁶

Ozone is a greenhouse gas, and is continually produced and destroyed in the atmosphere by chemical reactions. In the troposphere, the lowest 5 to 10 miles of the atmosphere near the surface, human activities have increased the ozone concentration through the release of gases such as carbon monoxide, hydrocarbons, and nitrogen oxides. These gases undergo chemical reactions to produce ozone in the presence of sunlight. In addition to trapping heat, excess ozone in the troposphere causes respiratory illnesses and other human health problems.

In the stratosphere, the layer above the troposphere, ozone exists naturally and protects life on Earth from exposure to excessive ultraviolet radiation from the Sun. As mentioned previously, halocarbons released by human activities destroy ozone in the stratosphere and have caused the ozone hole over Antarctica.⁸ Changes in the stratospheric ozone layer have contributed to changes in wind patterns and regional climates in Antarctica.⁹

Water vapor is the most important and abundant greenhouse gas in the atmosphere. Human activities produce only a very small increase in water vapor through irrigation and combustion processes.³ However, the surface warming caused by human-produced increases in other greenhouse gases leads to an increase in atmospheric water vapor, since a warmer climate increases evaporation and allows the atmosphere to hold more moisture. This creates an amplifying “feedback loop,” leading to more warming.

Other human influences

In addition to the global-scale climate effects of heat-trapping gases, human activities also produce additional local and regional effects. Some of these activities partially offset the warming caused by greenhouse gases, while others increase the warming. One such influence on climate is caused by tiny particles called “aerosols” (not to be confused with aerosol spray cans). For example, the burning of coal produces emissions of sulfur-containing compounds. These compounds form “sulfate aerosol” particles, which reflect some of the incoming sunlight away from the Earth, causing a cooling influence at the surface. Sulfate aerosols also tend to make clouds more efficient at reflecting sunlight, causing an additional indirect cooling effect.

Another type of aerosol, often referred to as soot or black carbon, absorbs incoming sunlight and traps heat in the atmosphere. Thus, depending on their type, aerosols can either mask or increase the warming caused by increased levels of greenhouse gases. On a globally averaged basis, the sum of these aerosol effects offsets some of the warming caused by heat-trapping gases.¹⁰

The effects of various greenhouse gases and aerosol particles on Earth’s climate depend in part on how long these gases and particles remain in the atmosphere. After emission, the atmospheric concentration of carbon dioxide remains elevated for thousands of years and that of methane for decades, while the elevated concentrations of aerosols only persist for days to weeks.^{11,12} The climate effects of reductions in emissions of carbon dioxide and other long-lived gases do not become apparent for at least several decades. In contrast, reductions in emissions of short-lived compounds can have a rapid, but complex effect since the geographic patterns of their climatic influence and the resulting surface temperature responses are quite different. One modeling study found that while the greatest emissions of short-lived pollutants in summertime by late this century are projected to come from Asia, the strongest climatic response is projected to be over the central United States.¹³

Human activities have also changed the land surface in ways that alter how much heat is reflected or absorbed by the surface. Such changes include the cutting and burning of forests, the replacement of other areas of natural vegetation with agriculture and cities, and large-scale irrigation. These transformations of the land surface can cause local (and even regional) warming or cooling. Globally, the net effect of these changes has probably been a slight cooling of the Earth’s surface over the past 100 years.^{14,15}

Natural influences

Two important natural factors also influence climate: the Sun and volcanic eruptions. Over the past three decades, human influences on climate have become increasingly obvious, and global temperatures have risen sharply. During the same period, the Sun’s energy output (as measured by satellites since 1979) has followed its historic 11-year cycle of

R1
R2
R3
R4
R5
R6
R7
R8
R9
R10
R11
R12
R13
R14
R15
R16
R17
R18
R19
R20
R21
R22
R23
R24
R25
R26
R27
R28
R29
R30
R31
R32
R33
R34
R35
R36
R37
R38
R39
R40
R41
R42
R43
R44
R45
R46
R47
R48
R49
R50



small ups and downs, but with no net increase (see figure page 20).¹⁶ The two major volcanic eruptions of the past 30 years have had short-term cooling effects on climate, lasting 2 to 3 years.¹⁷ Thus, these natural factors cannot explain the warming of recent decades; in fact, their net effect on climate has probably been a slight cooling influence over this period. Slow changes in Earth's orbit around the Sun and its tilt toward or away from the Sun are also a purely natural influence on climate, but are only important on timescales from thousands to many tens of thousands of years.

The climate changes that have occurred over the last century are not solely caused by the human and natural factors described above. In addition to these

influences, there are also fluctuations in climate that occur even in the absence of changes in human activities, the Sun, or volcanoes. One example is the El Niño phenomenon, which has important influences on many aspects of regional and global climate. Many other modes of variability have been identified by climate scientists and their effects on climate occur at the same time as the effects of human activities, the Sun, and volcanoes.

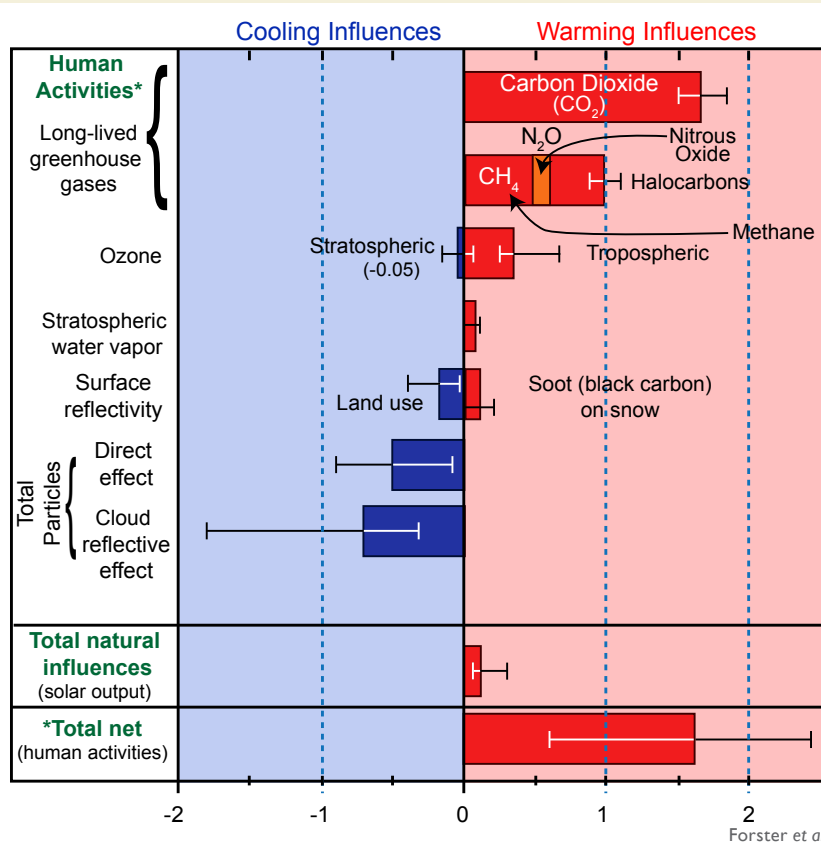
Carbon release and uptake

Once carbon dioxide is emitted to the atmosphere, some of it is absorbed by the oceans and taken up by vegetation, although this storage may be temporary. About 45 percent of the carbon dioxide emitted by human activities in the last 50 years is now

stored in the oceans and vegetation. The rest has remained in the air, increasing the atmospheric concentration.^{2,3,18} It is thus important to understand not only how much carbon dioxide is emitted, but also how much is taken up, over what time scales, and how these sources and sinks of carbon dioxide might change as climate continues to warm. For example, it is known from long records of Earth's climate history that under warmer conditions, carbon tends to be released, for example, from thawing permafrost, initiating a feedback loop in which more carbon release leads to more warming which leads to further release, and so on.^{19,20}

Global emissions of carbon dioxide have been accelerating. The growth rate increased from 1.3 percent per year in the 1990s to 3.3 percent per year between 2000 and 2006.²¹ The increasing emissions of carbon dioxide have clearly contributed to the increased concentration of carbon dioxide observed in the atmosphere, but are perhaps not the only factor. There is also evidence that a smaller fraction of the annual human-induced emissions is now being taken up than in the past, leading to a greater fraction remaining in the atmosphere and an accelerating rate of increase in the carbon dioxide concentration.²¹

Major Warming and Cooling Influences on Climate 1750-2005



The figure above shows the amount of warming influence (red bars) or cooling influence (blue bars) that different factors have had on Earth's climate over the industrial age (from about 1750 to the present). Results are in watts per square meter. The longer the bar, the greater the influence on climate. The top part of the box includes all the major human-induced factors, while the second part of the box includes the Sun, the only major natural factor with a long-term effect on climate. The cooling effect of individual volcanoes is also natural, but is relatively short-lived (2 to 3 years), thus their influence is not included in this figure. The bottom part of the box shows that the total net effect (warming influences minus cooling influences) of human activities is a strong warming influence. The thin lines on each bar provide an estimate of the range of uncertainty.

Ocean acidification

As the ocean absorbs carbon dioxide from the atmosphere, seawater is becoming less alkaline (its pH is decreasing) through a process generally referred to as ocean acidification. The pH of seawater has decreased significantly since 1750,^{22,23} and is projected to drop much more dramatically by the end of the century if carbon dioxide concentrations continue to increase.²⁴ Such ocean acidification is essentially irreversible over centuries. As discussed in the *Ecosystems* sector and *Coasts* region, ocean acidification affects the process of calcification by which living things create shells and skeletons, with substantial negative consequences for coral reefs, mollusks, and some plankton important to ocean food chains.²⁵

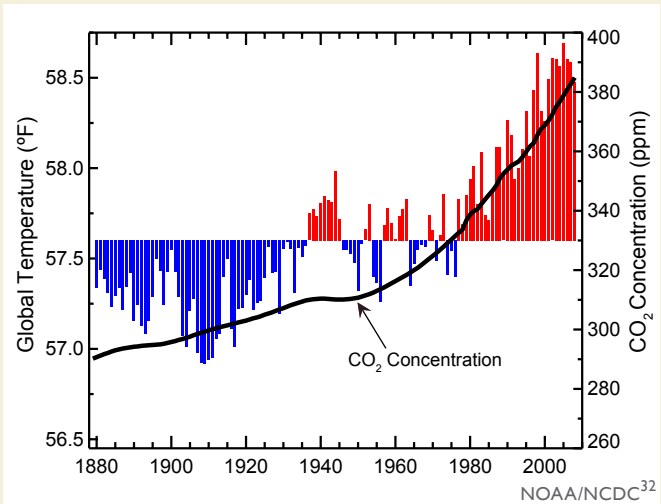
Over the last 100 years, global average temperature and sea level have increased, and precipitation patterns have changed.

Temperatures are rising

Global average surface air temperature has been increasing rapidly since 1970.²⁶ The estimated change in the average temperature of Earth's surface is based on measurements from thousands of weather stations, ships, and buoys around the world, as well as from satellites. These measurements are independently compiled, analyzed, and processed by different research groups. There are a number of important steps in the data processing. These include identifying and adjusting for the effects of changes in the instruments used to measure temperature, the measurement times and locations, the local environment around the measuring site, and such factors as satellite orbital drift. For instance, the growth of cities can cause localized "urban heat island" effects.

A number of research groups around the world have produced estimates of global-scale changes in surface temperature. The warming trend that is apparent in all of these temperature records is confirmed by other independent observations, such as the melting of Arctic sea ice, the retreat of mountain glaciers on every continent,²⁷ reductions in the extent of snow cover, earlier blooming of plants in spring, and increased melting of the Greenland

Global Temperature and Carbon Dioxide



Global annual average temperature (as measured over both land and oceans). Red bars indicate temperatures above and blue bars indicate temperatures below the average temperature for the period 1901-2000. The black line shows atmospheric carbon dioxide (CO₂) concentration in parts per million (ppm). While there is a clear long-term global warming trend, each individual year does not show a temperature increase relative to the previous year, and some years show greater changes than others³³. These year-to-year fluctuations in temperature are due to natural processes, such as the effects of El Niños, La Niñas, and the eruption of large volcanoes.

and Antarctic ice sheets.^{28,29} Because snow and ice reflect the Sun's heat, this melting causes more heat to be absorbed, which causes more melting, resulting in another feedback loop.²⁰

Additionally, temperature measurements above the surface have been made by weather balloons since the late 1940s, and from satellites since 1979. These measurements show warming of the troposphere, consistent with the surface warming.^{30,31} They also reveal cooling in the stratosphere.³⁰ This pattern of tropospheric warming and stratospheric cooling agrees with our understanding of how atmospheric temperature would be expected to change in response to increasing greenhouse gas concentrations and the observed depletion of stratospheric ozone.¹⁴

Precipitation patterns are changing

Precipitation is not distributed evenly over the globe. Its average distribution is governed primarily by atmospheric circulation patterns, the availability of moisture, and surface terrain effects. The first two of these factors are influenced by temperature. Thus, human-caused changes in temperature are expected to alter precipitation patterns.

Observations show that such shifts are occurring. Changes have been observed in the amount, intensity, frequency, and type of precipitation. Pronounced increases in precipitation over the past 100 years have been observed in eastern North America, southern South America, and northern Europe. Decreases have been seen in the Mediterranean, most of Africa, and southern Asia. Changes in the geographical distribution of droughts and flooding have been complex. In some regions, there have been increases in the occurrences of both droughts and floods.²⁸ As the world warms, northern regions and mountainous areas are experiencing more precipitation falling as rain rather than snow.³⁴ Widespread increases in heavy precipitation events have occurred, even in places where total amounts have decreased. These changes are associated with the fact that warmer air holds more water vapor evaporating from the world's oceans and land surface.³¹ This increase in atmospheric water vapor has been observed from satellites, and is primarily due to human influences.^{35,36}

Sea level is rising

After at least 2000 years of little change, sea level rose by roughly 8 inches over the past century. Satellite data available over the past 15 years show sea level rising at a rate roughly double the rate observed over the past century.³⁷

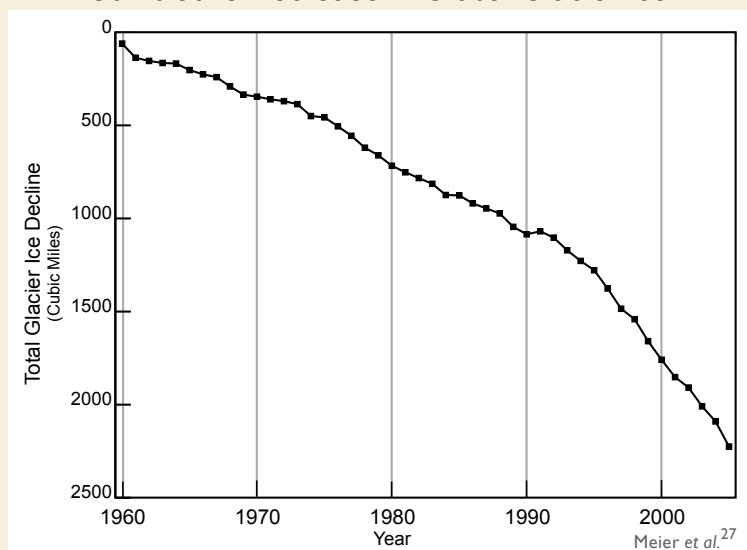
There are two principal ways in which global warming causes sea level to rise. First, ocean water expands as it warms, and therefore takes up more space. Warming has been observed in each of the world's major ocean basins, and has been directly linked to human influences.^{38,39}

Second, warming leads to the melting of glaciers and ice sheets, which raises sea level by adding water to the oceans. Glaciers have been retreating worldwide for at least the last century, and the rate of retreat has increased in the past decade.^{29,40} Only a few glaciers are actually advancing (in locations that were

well below freezing, and where increased precipitation has outpaced melting). The total volume of glaciers on Earth is declining sharply. The progressive disappearance of glaciers has implications not only for the rise in global sea level, but also for water supplies in certain densely-populated regions of Asia and South America.

The Earth has major ice sheets on Greenland and Antarctica. These ice sheets are currently losing ice volume by increased melting and calving of icebergs, contributing to sea-level rise. The Greenland Ice Sheet has also been experiencing record amounts of surface melting, and a large increase in the rate of mass loss in the past decade.⁴¹ If the entire Greenland Ice Sheet melted, it would raise sea level by about 20 feet. The Antarctic Ice Sheet consists of two portions, the West Antarctic Ice Sheet and the East Antarctic Ice Sheet. The West Antarctic Ice Sheet, the more vulnerable of the two, contains enough water to raise global sea levels by about 16 to 20 feet.²⁹ If the East Antarctic Ice Sheet melted entirely, it would raise global sea level by about 200 feet. Complete melting of these ice sheets over this century or the next is thought to be virtually impossible, although past climate records provide precedent for very significant decreases in ice volume, and therefore increases in sea level.^{42,43}

Cumulative Decrease in Global Glacier Ice



As temperatures have risen, glaciers around the world have shrunk. The graph shows the cumulative decline in glacier ice worldwide.



The global warming of the past 50 years is due primarily to human-induced increases in heat-trapping gases. Human “fingerprints” also have been identified in many other aspects of the climate system, including changes in ocean heat content, precipitation, atmospheric moisture, and arctic sea ice.

In 1996, the IPCC Second Assessment Report⁴⁴ cautiously concluded that “the balance of evidence suggests a discernible human influence on global climate.” Since then, a number of national and international assessments have come to much stronger conclusions about the reality of human effects on climate. Recent scientific assessments find that most of the warming of the Earth’s surface over the past 50 years has been caused by human activities.^{45,46}

This conclusion rests on multiple lines of evidence. Like the warming “signal” that has gradually emerged from the “noise” of natural climate variability, the scientific evidence for a human influence on global climate has accumulated over the past several decades, from many hundreds of studies. No single study is a “smoking gun.” Nor has any single study or combination of studies undermined the large body of evidence supporting the conclusion that human activity is the primary driver of recent warming.

The first line of evidence is our basic physical understanding of how greenhouse gases trap heat, how the climate system responds to increases in greenhouse gases, and how other human and natural factors influence climate. The second line of evidence is from indirect estimates of climate changes over the last 1,000 to 2,000 years. These records are obtained from living things and their remains (like tree rings and corals) and from physical quantities (like the ratio between lighter and heavier isotopes of oxygen in ice cores) which change in measurable ways as climate changes. The lesson from these data is that global surface temperatures over the last several decades are clearly unusual, in that they were higher than at any time during at least the past 400 years⁴⁷. For the Northern Hemisphere, the recent temperature rise is clearly unusual in at least the last 1,000 years.^{47,48}

The third line of evidence is based on the broad, qualitative consistency between observed changes in climate and the computer model simulations of how climate would be expected to change in response to human activities. For example, when climate models are run with historical increases in greenhouse gases, they show gradual warming of the Earth and ocean surface, increases in ocean heat content and the temperature of the lower atmosphere, a rise in global sea level, retreat of sea-ice and snow cover, cooling of the stratosphere, an increase in the amount of atmospheric water vapor, and changes in large-scale precipitation and pressure patterns. These and other aspects of modeled climate change are in agreement with observations.^{14,49}

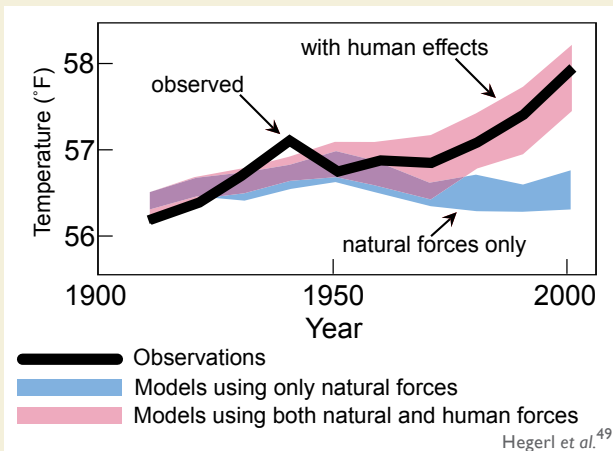
Finally, there is extensive statistical evidence from so-called “fingerprint” studies. Each factor that affects climate produces a unique pattern of climate response, much as each person has a unique fingerprint. Fingerprint studies exploit these unique signatures, and allow detailed comparisons of modeled and observed climate change patterns.⁴⁴ Scientists rely on such studies to attribute observed changes in climate to a particular cause or set of causes. In the real world, the climate changes that have occurred since the start of the Industrial Revolution are due to a complex mixture of human and natural causes. The importance of each individual influence in this mixture changes over time. Of course, there are not multiple Earths, which would allow an experimenter to change one factor at a time on each Earth, thus helping to isolate different fingerprints. Climate models can be used to perform the systematic experiments that are not possible in the real world: a single factor (like greenhouse gases) or a set of factors can be varied, and the response of the climate system to these individual or combined changes can thus be studied.⁵⁰

For example, when climate model simulations of the last century include all of the major influences on climate, both human-induced and natural, they can reproduce many important features of observed climate change patterns. When human influences are removed from the model experiments, results suggest that the surface of the Earth would actually have cooled slightly over the last 50 years. The clear message from fingerprint studies is that the

R1
R2
R3
R4
R5
R6
R7
R8
R9
R10
R11
R12
R13
R14
R15
R16
R17
R18
R19
R20
R21
R22
R23
R24
R25
R26
R27
R28
R29
R30
R31
R32
R33
R34
R35
R36
R37
R38
R39
R40
R41
R42
R43
R44
R45
R46
R47
R48
R49
R50



Separating Human and Natural Influences on Climate



The blue band shows how global average temperatures would have changed due to natural forces only, as simulated by climate models. The red band shows model projections of the effects of human and natural forces combined. The black line shows actual observed global average temperatures. As the blue band indicates, without human influences, temperature over the past century would actually have first warmed and then cooled slightly over recent decades.⁵⁸

observed warming over the last half-century cannot be explained by natural factors, and is instead caused primarily by human factors.^{14,50}

Another fingerprint of human effects on climate has been identified by looking at a slice through the layers of the atmosphere, and studying the pattern of temperature changes from the surface up through the stratosphere. In all climate models, increases in carbon dioxide cause warming at the surface and in the troposphere, but lead to cooling of the stratosphere. For straightforward physical reasons, models also calculate that the human-caused depletion of stratospheric ozone has had a strong cooling effect in the stratosphere. There is a good match between the model fingerprint in response to combined carbon dioxide and ozone changes and the observed pattern of tropospheric warming and stratospheric cooling (see figure on next page).¹⁴

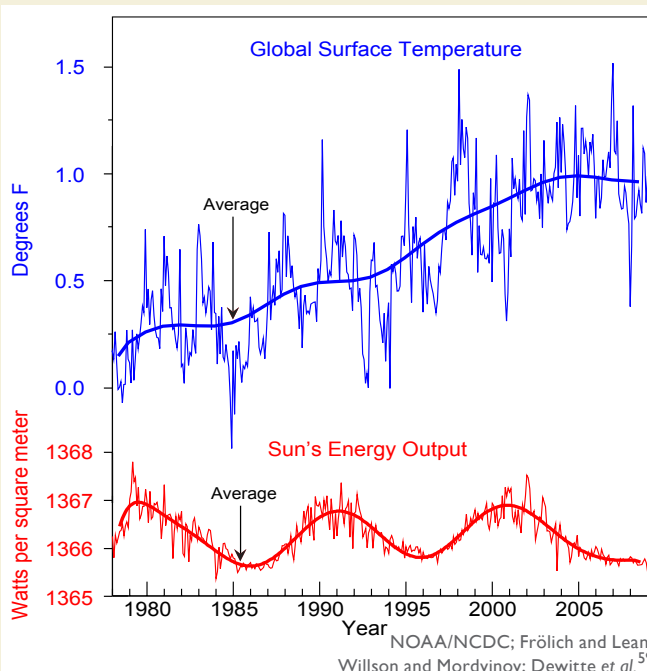
In contrast, if most of the observed temperature change had been due to an increase in solar output rather than an increase in greenhouse gases, Earth's atmosphere would have warmed throughout its full vertical extent, including the stratosphere.⁹ The observed pattern of atmo-

spheric temperature changes, with its pronounced cooling in the stratosphere, is therefore inconsistent with the hypothesis that changes in the Sun can explain the warming of recent decades. Moreover, direct satellite measurements of solar output show slight decreases during the recent period of warming.

The earliest fingerprint work⁵¹ focused on changes in surface and atmospheric temperature. Scientists then applied fingerprint methods to a whole range of climate variables,^{50,52} identifying human-caused climate signals in the heat content of the oceans,^{38,39} the height of the tropopause⁵³ (the boundary between the troposphere and stratosphere, which has shifted upward by hundreds of feet in recent decades), the geographical patterns of precipitation,⁵⁴ drought,⁵⁵ surface pressure,⁵⁶ and the runoff from major river basins.⁵⁷

Studies published after the appearance of the IPCC Fourth Assessment Report in 2007 have also found human fingerprints in the increased levels of atmospheric moisture^{35,36} (both close to the surface and over the full extent of the atmosphere), in the

Observed Changes in Surface Temperature and the Sun's Energy Output



The Sun's energy output has been measured by satellites since 1978. It has followed its natural 11-year cycle of small ups and downs, but with no net increase (bottom). Over the same period, global temperature has risen markedly (top).⁶⁰

L1
L2
L3
L4
L5
L6
L7
L8
L9
L10
L11
L12
L13
L14
L15
L16
L17
L18
L19
L20
L21
L22
L23
L24
L25
L26
L27
L28
L29
L30
L31
L32
L33
L34
L35
L36
L37
L38
L39
L40
L41
L42
L43
L44
L45
L46
L47
L48
L49
L50

R1
R2
R3
R4
R5
R6
R7
R8
R9
R10
R11
R12
R13
R14
R15
R16
R17
R18
R19
R20
R21
R22
R23
R24
R25
R26
R27
R28
R29
R30
R31
R32
R33
R34
R35
R36
R37
R38
R39
R40
R41
R42
R43
R44
R45
R46
R47
R48
R49
R50

decline of Arctic sea ice extent,⁶¹ and in the patterns of changes in Arctic and Antarctic surface temperatures.⁶²

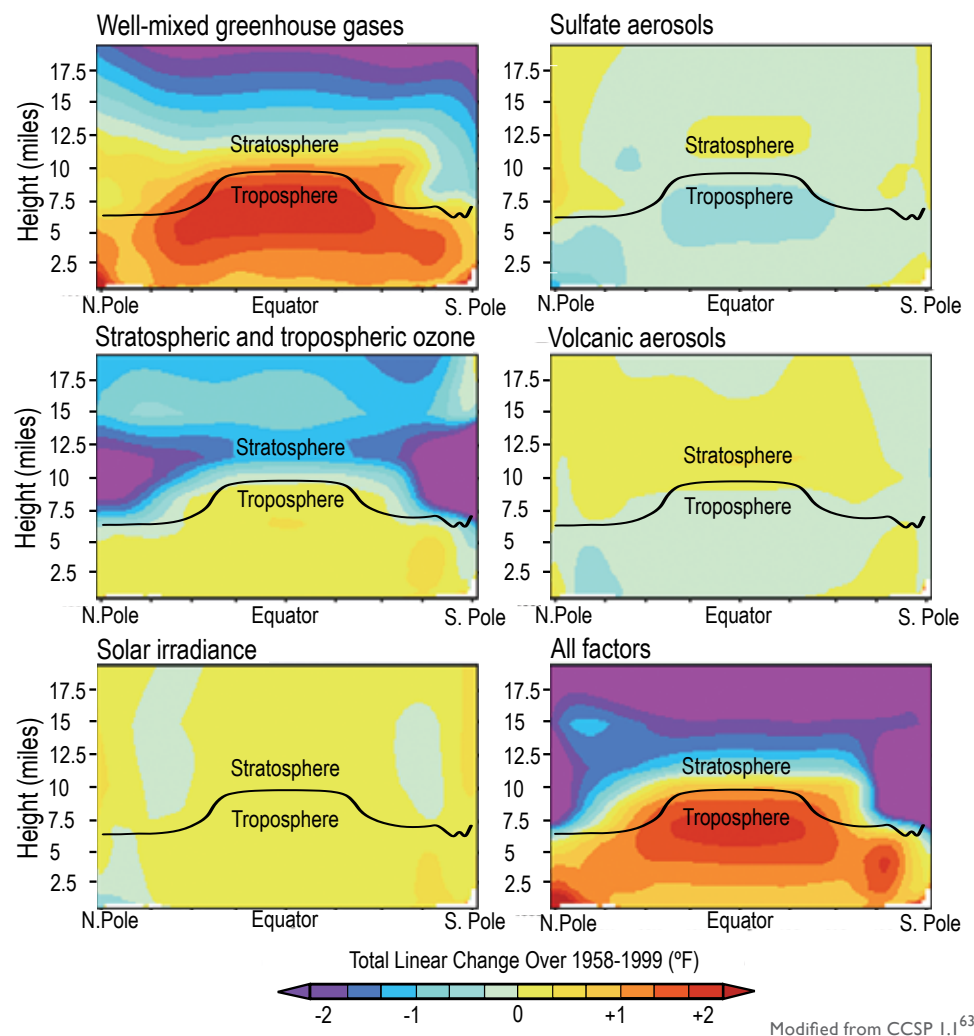
The message from this entire body of work is that the climate system is telling a consistent story of increasingly dominant human influence—the changes in temperature, ice extent, moisture, and circulation patterns fit together in a physically consistent way, like pieces in a complex puzzle.

Increasingly, this type of fingerprint work is shifting its emphasis. As noted, clear and compelling scientific evidence supports the case for a pronounced human influence on global climate. Much of the recent attention is now on climate changes at continental and regional scales,^{64,65} and on variables that can have large impacts on societies. For example, scientists have established causal links between human activities and the changes in snowpack, maximum and minimum temperature, and the seasonal timing of runoff over mountainous regions of the western United States.³⁴ Human activity is likely to have made a substantial contribution to ocean surface temperature changes in hurricane formation regions.⁶⁶⁻⁶⁸ Researchers are also looking beyond the physical climate system, and are beginning to tie changes in the distribution and seasonal behavior of plant and animal species to human-caused changes in temperature and precipitation.^{69,70}

For over a decade, one aspect of the climate change story seemed to show a significant difference between models and observations.¹⁴

In the tropics, all models predicted that with a rise in greenhouse gases, the troposphere would be expected to warm more rapidly than the surface. Observations from weather balloons, satellites, and surface thermometers seemed to show the opposite behavior (more rapid warming of the surface than the troposphere). This issue was a stumbling block in our understanding of the causes of climate change. It is now largely resolved.⁷¹ Research showed that there were large uncertainties in the satellite and weather balloon data. When uncertainties in models and observations are properly accounted for, newer observational data sets (with better treatment of known problems) are in agreement with climate model results.^{31,72-75}

Patterns of Temperature Change Produced by Various Atmospheric Factors, 1958-1999



Climate simulations of the vertical profile of temperature change due to various factors, and the effect due to all factors taken together. The panels above represent a cross-section of the atmosphere from the north pole to the south pole, and from the surface up into the stratosphere. The black lines show the location of the tropopause, the boundary between the lower atmosphere (troposphere) and the stratosphere.



L1 This does not mean, however, that all remain- R1
 L2 ing differences between models and observations R2
 L3 have been resolved. The observed changes in some R3
 L4 climate variables, such as Arctic sea ice,^{61,76} some R4
 L5 aspects of precipitation,^{54,77} and patterns of sur- R5
 L6 face pressure, appear to be proceeding much more R6
 L7 rapidly than models have projected. The reasons for R7
 L8 these differences are not well understood. Never- R8
 L9 theless, the bottom-line conclusion from climate R9
 L10 fingerprinting is that most of the observed changes R10
 L11 studied to date are consistent with each other, and R11
 L12 are also consistent with our scientific understand- R12
 L13 ing of how the climate system would be expected R13
 L14 to respond to the increase in heat-trapping gases R14
 L15 resulting from human activities.^{14,49} R15
 L16

L17 Scientists are sometimes asked whether extreme R17
 L18 weather events can be linked to human activities.²⁴ R18
 L19 Scientific research has concluded that human influ- R19
 L20 ences on climate are indeed changing the likelihood R20
 L21 of certain types of extreme events. For example, R21
 L22 an analysis of the European summer heat wave R22
 L23 of 2003 found that the risk of such a heat wave is R23
 L24 now roughly four times greater than it would have R24
 L25 been in the absence of human-induced climate R25
 L26 change.^{68,78} R26

L27 Like fingerprint work, such analyses of human- R27
 L28 caused changes in the risks of extreme events rely R28
 L29 on information from climate models, and on our R29
 L30 understanding of the physics of the climate system. R30
 L31 All of the models used in this work have imperfec- R31
 L32 tions in their representation of the complexities of R32
 L33 the “real world” climate system.^{79,80} These are due R33
 L34 to both limits in our understanding of the climate R34
 L35 system, and in our ability to represent its com- R35
 L36 plex behavior with available computer resources. R36
 L37 Despite this, models are extremely useful, for a R37
 L38 number of reasons. R38
 L39

L40 First, despite remaining imperfections, the current R40
 L41 generation of climate models accurately portrays R41
 L42 many important aspects of today’s weather pat- R42
 L43 terns and climate.^{79,80} Models are constantly being R43
 L44 improved, and are routinely tested against many R44
 L45 observations of Earth’s climate system. Second, R45
 L46 the fingerprint work shows that models capture R46
 L47 not only our present-day climate, but also key R47
 L48 features of the observed climate changes over R48
 L49 the past century.⁴⁷ Third, many of the large-scale R49
 L50

observed climate changes (such as the warming of R1
 the surface and troposphere, and the increase in the R2
 amount of moisture in the atmosphere) are driven R3
 by very basic physics, which is well-represented R4
 in models.³⁵ Fourth, climate models can be used to R5
 predict changes in climate that can be verified in R6
 the real world. Examples include the global cooling R7
 subsequent to the eruption of Mount Pinatubo and R8
 the stratospheric cooling with increasing carbon R9
 dioxide. Finally, models are the only tools that exist R10
 for trying to understand the climate changes likely R11
 to be experienced over the course of this century. R12
 No period in Earth’s geological history provides an R13
 exact analogue for the climatic conditions that will R14
 unfold in the coming decades. R15
 R16

Global temperatures are projected to continue to rise over this century; by how much and for how long depends on a number of factors, including the amount of heat-trapping gas emissions and how sensitive the climate is to those emissions.

R17 Some continued warming of the planet is projected R17
 R18 over the next few decades due to past emissions. R18
 R19 Choices made now will influence the amount of fu- R19
 R20 ture warming. Lower levels of heat-trapping emis- R20
 R21 sions will yield less future warming, while higher R21
 R22 levels will result in more warming, and more severe R22
 R23 impacts on society and the natural world. R23
 R24

Emissions scenarios

R25 The IPCC developed a set of scenarios in a Special R25
 R26 Report on Emissions Scenarios (SRES).⁸¹ These R26
 R27 have been extensively used to explore the potential R27
 R28 for future climate change. None of these scenarios R28
 R29 assumes explicit policies to limit climate change or R29
 R30 to stabilize atmospheric concentrations of heat- R30
 R31 trapping gases. Rather, emissions in these scenarios R31
 R32 vary based on different assumptions about changes R32
 R33 in population, rate of adoption of new technologies, R33
 R34 economic growth, and other factors. R34

R35 The IPCC emission scenarios also do not encom- R35
 R36 pass the full range of possible futures: emissions R36
 R37 can change less than those scenarios imply, or they R37
 R38 can change more. Current carbon dioxide emissions R38
 R39 are, in fact, above the highest emissions scenario R39
 R40
 R41
 R42
 R43
 R44
 R45
 R46
 R47
 R48
 R49
 R50

developed by the IPCC⁸² (see figure below). Whether this will continue is uncertain.

There are also lower possible emissions paths than those put forth by the IPCC. The Framework Convention on Climate Change, to which the United States and most other countries are signatories, calls for stabilizing concentrations of greenhouse gases in the atmosphere at a level that would avoid dangerous human interference with the climate system. What exactly constitutes such interference is subject to interpretation.

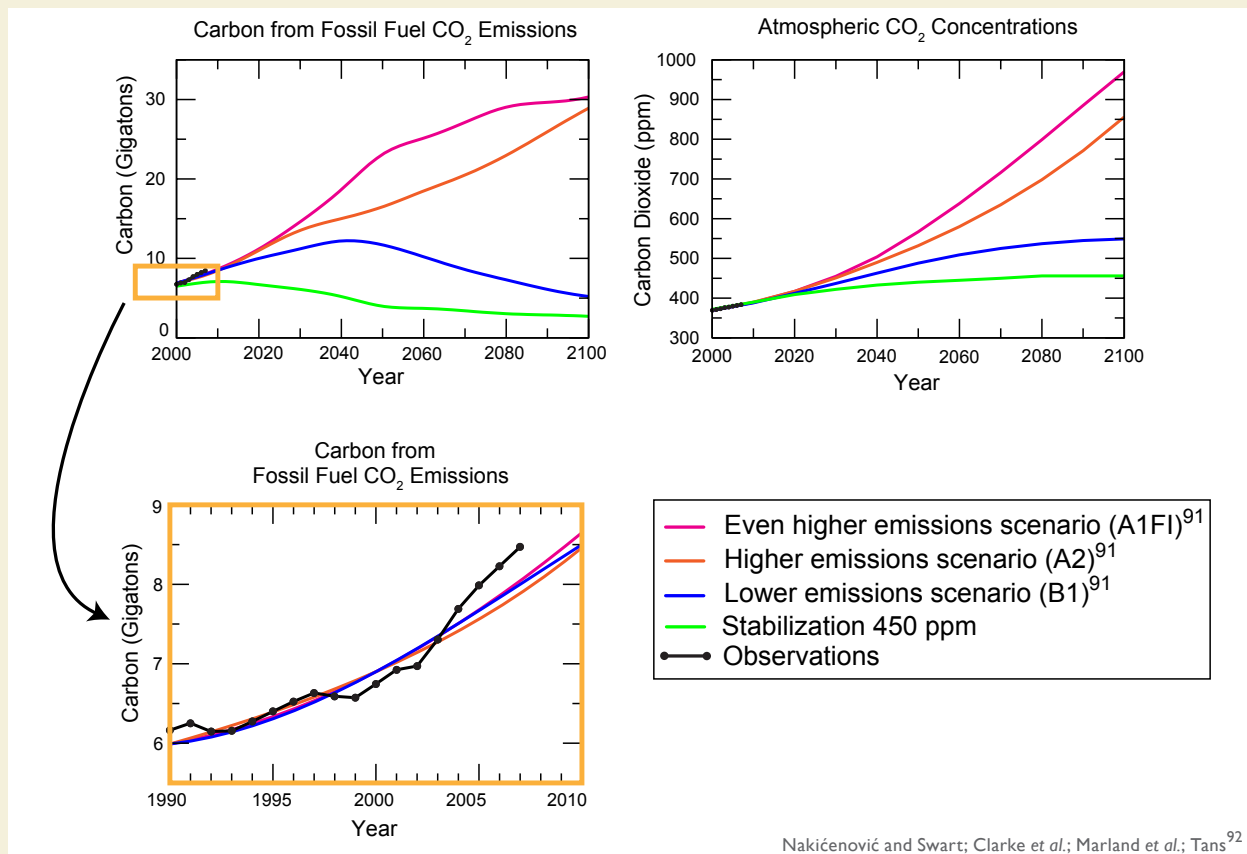
A variety of research studies suggest that a further 2°F increase (relative to the 1980-1999 period) would lead to severe, widespread, and irreversible impacts.⁸³⁻⁸⁵ To have a good chance (but not a guarantee) of avoiding temperatures above those levels, it has been estimated that atmospheric concentra-

tion of carbon dioxide would need to stabilize in the long term at around today's levels.⁸⁶⁻⁸⁹

Reducing emissions of carbon dioxide would reduce warming over this century and beyond. Sizable early reductions in emissions would significantly reduce the pace and the overall amount of climate change. Earlier reductions in emissions would have a greater effect in reducing climate change than comparable reductions made later. In addition, reducing emissions of some shorter-lived greenhouse gases, such as methane, and some types of particles, such as soot, would begin to reduce warming within weeks to decades.

The graphs below show emissions scenarios and resulting carbon dioxide concentrations for three IPCC scenarios^{90,91} and two stabilization scenarios.²⁵ The stabilization scenarios are aimed at stabi-

Scenarios of Future Carbon Dioxide Global Emissions and Concentrations



The graphs show recent and projected global emissions of carbon dioxide in gigatons of carbon, on the left, and atmospheric concentrations on the right under five emissions scenarios. The top three in the key are IPCC scenarios that assume no explicit climate policies (these are used in model projections that appear throughout this report). The bottom line is a “stabilization scenario,” designed to stabilize atmospheric carbon dioxide concentration at 450 parts per million. The inset expanded below these charts shows emissions for 1990-2007 under the three IPCC scenarios along with actual emissions (in black).



L1 lizing the atmospheric carbon dioxide concentra- R1
 L2 tion at roughly 450 and 550 parts per million (ppm); R2
 L3 this is 70 to 170 ppm above the 2008 concentration R3
 L4 of 385 ppm. Resulting temperature changes depend R4
 L5 on the level of carbon dioxide, how sensitive the R5
 L6 climate system is, and the amount of particles in R6
 L7 the atmosphere.⁸⁷ Only the 450 ppm stabilization R7
 L8 target has the potential to keep the global tempera- R8
 L9 ture rise at or below about 3.5°F from pre-industrial R9
 L10 and 2°F above the current average temperature, R10
 L11 a level beyond which many concerns have been R11
 L12 raised about dangerous human interference with the R12
 L13 climate system.^{88,89} Scenarios that stabilize carbon R13
 L14 dioxide below 450 ppm (not shown in the figure) R14
 L15 offer an increased chance of avoiding dangerous R15
 L16 climate change.^{88,89} R16

L17
 L18 Carbon dioxide is not the only greenhouse gas of R18
 L19 concern. Concentrations of other heat-trapping R19
 L20 gases like methane and nitrous oxide and particles R20
 L21 like soot will also have to be stabilized at low R21
 L22 enough levels to prevent global temperatures from R22
 L23 rising higher than the level mentioned above. When R23
 L24 these other gases are added, including the offset- R24
 L25 ting cooling effects of sulfate aerosol particles, R25
 L26 analyses suggest that stabilizing concentrations R26
 L27 around 400 parts per million of “equivalent carbon R27
 L28 dioxide” would yield about an 80 percent chance of R28
 L29 avoiding exceeding the 2°F above present tempera- R29
 L30 ture threshold. This would be true even if concen- R30
 L31 trations temporarily peaked as high as 475 parts R31
 L32 per million and then stabilized at 400 parts per R32
 L33 million roughly a century later.^{72,88,89,93-95} Reductions R33
 L34 in sulfate aerosol particles would necessitate lower R34
 L35 equivalent carbon dioxide targets. R35
 L36

L37 **Rising global temperature**

L38 All climate models project that human-caused R38
 L39 emissions of heat-trapping gases will cause further R39
 L40 warming in the future. Based on scenarios that R40
 L41 do not assume explicit climate policies to reduce R41
 L42 greenhouse gas emissions, global average tempera- R42
 L43 ture is projected to rise by 2 to 11.5°F by the end R43
 L44 of this century⁹⁰ (relative to the 1980-1999 time R44
 L45 period). Whether the actual warming in 2100 will R45
 L46 be closer to the low or the high end of this range R46
 L47 depends primarily on two factors: first, the fu- R47
 L48 ture level of emissions of heat-trapping gases, and R48
 L49 second, how sensitive climate is to past and future R49
 L50 emissions. The range of possible outcomes has R50

R1 been explored using a range of different emissions R1
 R2 scenarios, and a variety of climate models that en- R2
 R3 compass the known range of climate sensitivity. R3
 R4

R5 **Changing precipitation patterns**

R6 Projections of changes in precipitation largely R6
 R7 follow recently observed patterns of change, with R7
 R8 overall increases in the global average but substan- R8
 R9 tial shifts in where and how precipitation falls.⁹⁰ R9
 R10 Generally, higher latitudes are projected to receive R10
 R11 more precipitation, while the dry belt that lies just R11
 R12 outside the tropics expands further poleward,^{96,97} R12
 R13 and also receives less rain. Increases in tropical R13
 R14 precipitation are projected during rainy seasons R14
 R15 (such as monsoons), and especially over the tropical R15
 R16 Pacific. Certain regions, including the U.S. West R16
 R17 (especially the Southwest) and the Mediterranean, R17
 R18 are expected to become drier. The trend towards R18
 R19 more heavy downpours is expected to continue, R19
 R20 with precipitation becoming less frequent but more R20
 R21 intense.⁹⁰ More precipitation is expected to fall as R21
 R22 rain rather than snow. R22

R23
 R24 **Currently rare extreme events are becoming**
 R25 **more common**

R26 In a warmer future climate, models project there R26
 R27 will be an increased risk of more intense, more R27
 R28 frequent and longer-lasting heat waves.⁹⁰ The R28
 R29 European heat wave of 2003 is an example of the R29
 R30 type of extreme heat event that is likely to become R30
 R31 much more common.⁹⁰ If greenhouse gas emissions R31
 R32 continue to increase, by the 2040s more than half of R32
 R33 European summers will be hotter than the summer R33
 R34 of 2003, and by the end of this century, a summer R34
 R35 as hot as that of 2003 will be considered unusually R35
 R36 cool.⁷⁸ R36

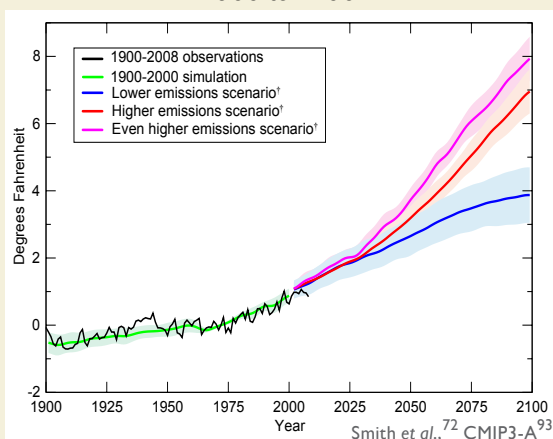
R37
 R38 Increased extremes of summer dryness and winter R38
 R39 wetness are projected for much of the globe, mean- R39
 R40 ing a generally greater risk of droughts and floods. R40
 R41 This has already been observed,⁵⁵ and is projected R41
 R42 to continue. In a warmer world, precipitation tends R42
 R43 to be concentrated into more intense events, with R43
 R44 longer periods of little precipitation in between.⁹⁰ R44
 R45

R46 Models project a general tendency for more intense R46
 R47 but fewer storms overall outside the tropics, with R47
 R48 more extreme wind events and higher ocean waves R48
 R49 in a number of regions in association with those R49
 R50

storms. Models also project a shift of storm tracks toward the poles in both hemispheres.⁹⁰

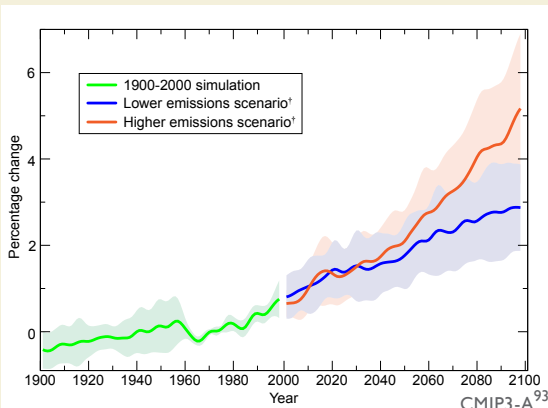
Changes in hurricanes are difficult to project because there are countervailing forces. Higher ocean temperatures lead to stronger storms with higher wind speeds and more rainfall.⁹⁸ But changes in wind speed and direction with height are also projected to increase in some regions, and this tends to work against storm formation and growth.⁹⁹⁻¹⁰¹ It currently appears that stronger, more rain-producing tropical storms and hurricanes are

Global Average Temperature 1900 to 2100



Observed and projected changes in the global average temperature under three IPCC no-policy emissions scenarios. The shaded areas show the likely ranges while the lines show the central projections from a set of climate models. A wider range of model types shows outcomes from 2 to 11.5°F.⁶⁸ Changes are relative to the 1960-1979 average.

Global Increase in Heavy Precipitation 1900 to 2100



Simulated and projected changes in the heaviest 5 percent of precipitation events. The shaded areas show the likely ranges while the lines show the central projections from a set of climate models. Changes are relative to the 1960-1979 average.

generally more likely, though more research is required on these issues.⁶⁸ A more detailed discussion of Atlantic hurricanes, which most impact the United States, can be found on page 34 in *National Climate Change*.

Sea level will continue to rise

Projecting future sea-level rise presents special challenges. Scientists have a well-developed understanding of the contributions of thermal expansion and melting glaciers to sea-level rise, so the models used to project sea-level rise include these processes. However, the contributions to past and future sea level rise from ice sheets are less well understood. Recent observations of the polar ice sheets show that a number of complex processes control the movement of ice to the sea, and thus affect the contributions of ice sheets to sea-level rise.²⁹ Some of these processes are already producing substantial loss of ice mass. Because these processes are not well understood it is difficult to predict their future contributions to sea-level rise.¹⁰²

Because of this uncertainty, the 2007 assessment by the IPCC could not quantify the contributions to sea-level rise due to changes in ice sheet dynamics, and thus projected a rise of the world's oceans from 8 inches to 2 feet by the end of this century.⁹⁰

More recent research has attempted to quantify the potential contribution to sea-level rise from the accelerated flow of ice sheets to the sea^{27,42} or to estimate future sea level based on its observed relationship to temperature.¹⁰³ The resulting estimates exceed those of the IPCC, and the average estimates under higher emissions scenarios are for sea-level rise between 3 and 4 feet. An important question that is often asked is, what is the upper bound of sea-level rise expected over this century? Few analyses have focused on this question. There is some evidence to suggest that it would be virtually impossible to have a rise of sea level higher than about 6.5 feet by the end of this century.⁴²

The changes in sea level experienced at any particular location along the coast depend not only on the increase in the global average sea level, but also on changes in regional currents and winds, proximity to the mass of melting ice sheets, and on the vertical movements of the land due to geological



L1 forces.¹⁰⁴ The consequences of sea-level rise at any
 L2 particular location depend on the amount of sea-
 L3 level rise relative to the adjoining land. Although
 L4 some parts of the U.S. coast are undergoing uplift
 L5 (rising), most shorelines are subsiding (sinking) to
 L6 various degrees—from a few inches to over 2 feet
 L7 per century.

L9 **Abrupt climate change**

L10 There is also the possibility of even larger changes
 L11 in climate than current scenarios and models
 L12 project. Not all changes in the climate are gradual.
 L13 The long record of climate found in ice cores, tree
 L14 rings, and other natural records show that Earth’s
 L15 climate patterns have undergone rapid shifts from
 L16 one stable state to another within as short a period
 L17 as a decade. The occurrence of abrupt changes in
 L18 climate becomes increasingly more likely as the
 L19 human disturbance of the climate system grows.⁹⁰
 L20 Such changes can occur so rapidly that they would
 L21 challenge the ability of human and natural systems
 L22 to adapt.¹⁰⁵ Examples of such changes are abrupt
 L23 shifts in drought frequency and duration. Ancient
 L24 climate records suggest that in the United States,
 L25 the Southwest may be at greatest risk for this kind
 L26 of change, but that other regions including the Mid-
 L27 west and Great Plains have also had these kinds of
 L28 abrupt shifts in the past and could experience them
 L29 again in the future.

L30
 L31 Rapid ice sheet collapse with related sea-level rise
 L32 is another type of abrupt change that is not well
 L33 understood or modeled that poses a risk for the fu-
 L34 ture. Recent observations show that melting on the
 L35 surface of an ice sheet produces water that flows
 L36 down through large cracks that create conduits
 L37 through the ice to the base of the ice sheet where it
 L38 lubricates ice previously frozen to the rock below.²⁹
 L39 Further, the interaction with warm ocean water,
 L40 where ice meets the sea, can lead to sudden losses
 L41 in ice mass and accompanying rapid global sea-
 L42 level rise. Observations indicate that ice loss has
 L43 increased dramatically over the last decade, though
 L44 scientists are not yet confident that they can project
 L45 how the ice sheets will respond in the future.

R1 There are also concerns regarding the potential for
 R2 abrupt release of methane from thawing of frozen
 R3 soils, from the sea floor, and from wetlands in the
 R4 tropics and the Arctic. While analyses suggest that
 R5 an abrupt release of methane is very unlikely to oc-
 R6 cur within 100 years, it is very likely that warming
 R7 will accelerate the pace of chronic methane emis-
 R8 sions from these sources, potentially increasing the
 R9 rate of global temperature increases.¹⁰⁶

R10
 R11 A third major area of concern regarding possible
 R12 abrupt change involves the operation of the ocean
 R13 currents that transport vast quantities of heat
 R14 around the globe. One branch of the ocean circula-
 R15 tion is in the North Atlantic. In this region, warm,
 R16 less-salty water flows from the tropics to the North
 R17 Atlantic in the upper layer of the ocean, while cold,
 R18 saltier water flows back from the North Atlantic
 R19 to the tropics in the ocean’s deep layers, creating a
 R20 “conveyor belt” for heat. Changes in this circulation
 R21 have profound impacts on the global climate sys-
 R22 tem, from changes in African and Indian monsoon
 R23 rainfall, to atmospheric circulation relevant to hur-
 R24 ricanes, to changes in climate over North America
 R25 and Western Europe.

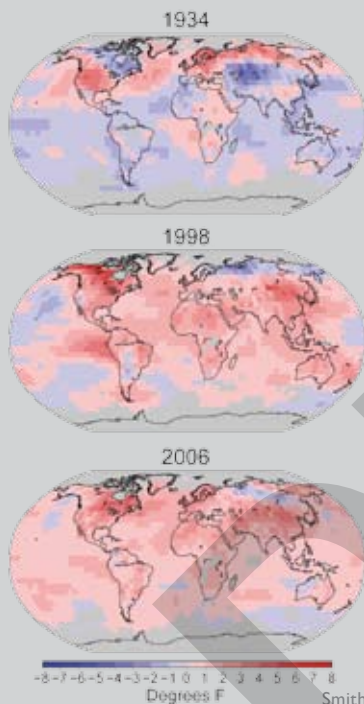
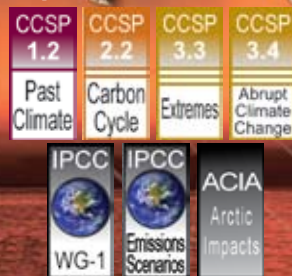
R26
 R27 Recent findings indicate that it is very likely that
 R28 the strength of this North Atlantic circulation will
 R29 decrease over the course of this century in response
 R30 to increasing greenhouse gases. The best estimate
 R31 is that the strength of this circulation will decrease
 R32 25 to 30 percent in this century, leading to a reduc-
 R33 tion in heat transfer to the North Atlantic. It is
 R34 considered very unlikely that this circulation would
 R35 collapse entirely during the next 100 years or so,
 R36 though it cannot be ruled out. While very unlikely,
 R37 the potential consequences of such an abrupt event
 R38 would be severe. Impacts would likely include
 R39 sea-level rise around the North Atlantic of up to 2.5
 R40 feet (in addition to the rise expected from thermal
 R41 expansion and melting glaciers and ice sheets),
 R42 changes in atmospheric circulation conditions that
 R43 influence hurricane activity, a southward shift of
 R44 tropical rainfall belts with resulting agricultural
 R45 impacts, and disruptions to marine ecosystems.⁷⁶

National Climate Change

Key Messages:

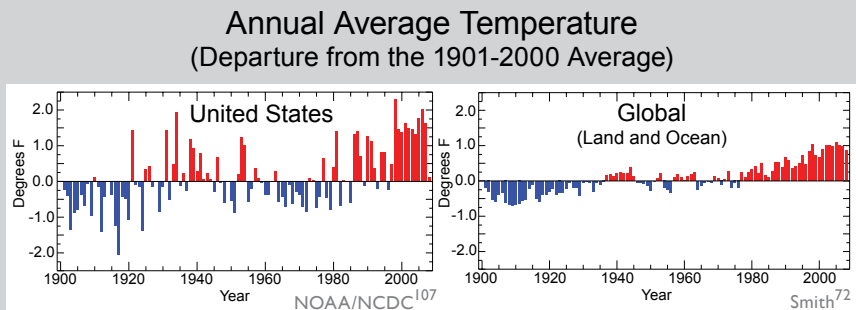
- U.S. average temperature has risen more than 2°F over the past 50 years and is projected to rise more in the future; how much more depends primarily on the amount of heat-trapping gases emitted globally and how sensitive the climate is to those emissions.
- Precipitation has increased an average of about 5 percent over the past 50 years. Projections of future precipitation generally indicate that northern areas will become wetter, and southern areas, particularly in the West, will become drier.
- The amount of rain falling in the heaviest downpours has increased approximately 20 percent on average in the past century, and this trend is very likely to continue, with the largest increases in the wettest places.
- Many types of extreme weather events, such as heat waves and regional droughts, have become more frequent and intense during the past 40 to 50 years.
- The destructive energy of Atlantic hurricanes has increased in recent decades. The intensity of these storms is likely to increase in this century.
- In the eastern Pacific, the strongest hurricanes have become stronger since the 1980s, even while the total number of storms has decreased.
- Sea level has risen along most of the U.S. coast over the last 50 years, and will rise more in the future.
- Cold-season storm tracks are shifting northward and the strongest storms are likely to become stronger and more frequent.
- Arctic sea ice is declining rapidly and this is very likely to continue.

Key Sources



The maps show annual temperature difference from the 1961 to 1990 average for the 3 years that were the hottest on record in the United States: 1998, 1934 and 2006 (in rank order). Red areas were warmer than average, blue were cooler than average. The 1930s were very warm in much of the United States, but they were not unusually warm globally. On the other hand, the warmth of 1998 and 2006, as for most years in recent decades, has been global in extent.

Like the rest of the world, the United States has been warming significantly over the past 50 years in response to the build up of heat-trapping gases in the atmosphere. When looking at national climate, however, it is important to recognize that climate responds to local, regional, and global factors. Therefore, national climate varies more than the average global climate. While various parts of the world have had particularly hot or cold periods earlier in the historical record, these periods have not been global in scale, whereas the warming of recent decades has been global in scale – hence the term *global* warming. It is also important to recognize that at both the global and national scales year-to-year fluctuations in natural weather and climate patterns can produce a period that does not follow the long-term trend. Thus, each year will not necessarily be warmer than every year before it, though the warming trend continues.



From 1901 to 2008, each year's temperature departure from the long-term average is one bar, with blue bars representing years cooler than the long-term average and red bars representing years warmer than that average. National temperatures vary much more than global temperatures, in part because of the moderating influence of the oceans.

U.S. average temperature has risen more than 2°F over the past 50 years and is projected to rise more in the future; how much more depends primarily on the amount of heat-trapping gases emitted globally and how sensitive the climate is to those emissions.

The series of maps and thermometers on these two pages shows the magnitude of the observed and projected changes in annual average temperature. The map for the period around 2000 shows that most areas of the United States have warmed 1 to 2°F compared to the 1960s and 1970s. Although not reflected in these maps of annual average temperature, this warming has generally resulted in longer warm seasons and shorter, less intense cold seasons.

The remaining maps show projected warming over the course of this century under a lower emissions scenario and a higher emissions scenario⁹¹ (see *Global Climate Change* section, page 23). Temperatures

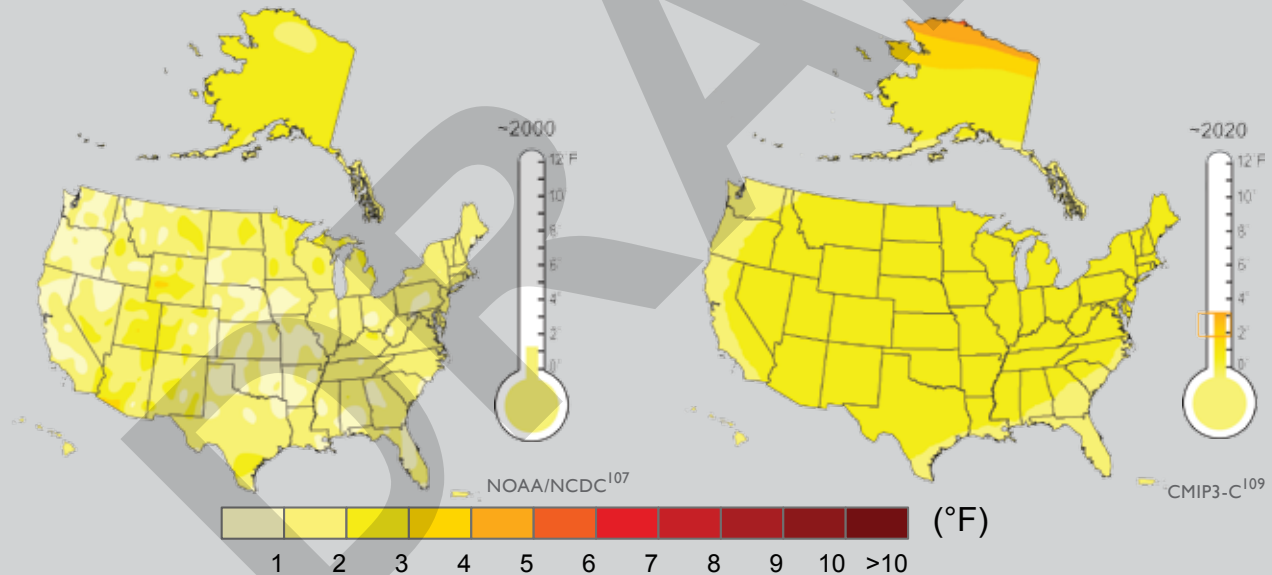
will continue to rise throughout the century under both emissions scenarios,⁹¹ although higher emissions result in more warming by the middle of the century and significantly more by the end of the century.

Temperature increases in the next couple of decades will be primarily determined by past emissions of heat-trapping gases. As a result, there is little difference in projected temperature between the higher and lower emissions scenarios⁹¹ in the near-term (around 2020), so only a single map is shown for this timeframe. Increases after the next couple of decades will be primarily determined by future emissions.⁹⁰ This is clearly evident in greater projected warming in the higher emissions scenario⁹¹ by the middle (around 2050) and end of this century (around 2090).

On a seasonal basis, most of the United States is projected to experience greater warming in summer than in winter, while Alaska experiences far more warming in winter than summer.¹⁰⁸

Present-Day (1993-2008) Average Change (°F) from 1961-1979 Baseline

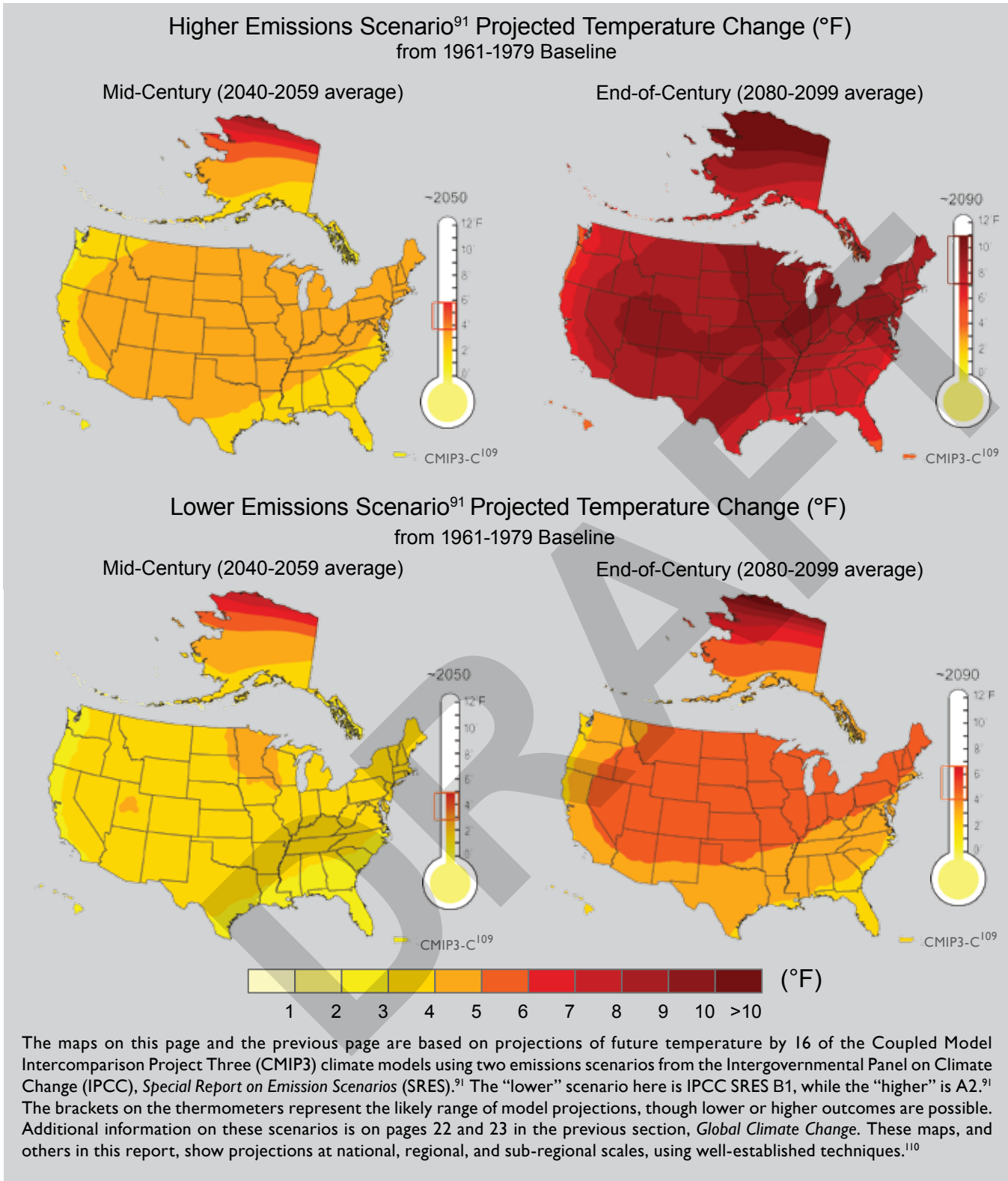
Near-Term (2010-2029) Projected Average Change (°F) from 1961-1979 Baseline



The maps and thermometers on this page and the next page show temperature differences (either measured or projected) from conditions as they existed during the period from 1961-1979. Comparisons to this period are made because the influence on climate from increasing greenhouse gas emissions has been greatest during the past five decades. The present-day map is based on the average observed temperatures from 1993-2008 minus the average from 1961-1979. Projected temperatures are based on results from 16 climate models for the periods 2010-2029, 2040-2059, and 2080-2099. The brackets on the thermometers represent the likely range of model projections, though lower or higher outcomes are possible. The mid-century and end-of-century maps show projections of model projections, though lower or higher outcomes are possible. The projection for the near-term is the average of the higher and lower emission scenarios⁹¹ because there is little difference in that timeframe.

L1 The average warming for the country as a whole is shown on the thermometers adjacent to each map. By the
 L2 end of the century, the average U.S. temperature is projected to increase by approximately 7 to 11°F under
 L3 the higher emissions scenario⁹¹ and by approximately 4 to 6.5°F under the lower emissions scenario.⁹¹ These
 L4 ranges are due to differences among climate model results for the same emissions scenarios.
 L5
 L6
 L7
 L8
 L9
 L10

R1
 R2
 R3
 R4
 R5
 R6
 R7
 R8
 R9
 R10
 R11
 R12
 R13
 R14
 R15
 R16
 R17
 R18
 R19
 R20
 R21
 R22
 R23
 R24
 R25
 R26
 R27
 R28
 R29
 R30
 R31
 R32
 R33
 R34
 R35
 R36
 R37
 R38
 R39
 R40
 R41
 R42
 R43
 R44
 R45
 R46
 R47
 R48
 R49
 R50



L46 The maps on this page and the previous page are based on projections of future temperature by 16 of the Coupled Model
 L47 Intercomparison Project Three (CMIP3) climate models using two emissions scenarios from the Intergovernmental Panel on Climate
 L48 Change (IPCC), *Special Report on Emission Scenarios* (SRES).⁹¹ The “lower” scenario here is IPCC SRES B1, while the “higher” is A2.⁹¹
 L49 The brackets on the thermometers represent the likely range of model projections, though lower or higher outcomes are possible.
 L50 Additional information on these scenarios is on pages 22 and 23 in the previous section, *Global Climate Change*. These maps, and
 others in this report, show projections at national, regional, and sub-regional scales, using well-established techniques.¹¹⁰

Precipitation has increased an average of about 5 percent over the past 50 years. Projections of future precipitation generally indicate that northern areas will become wetter, and southern areas, particularly in the West, will become drier.

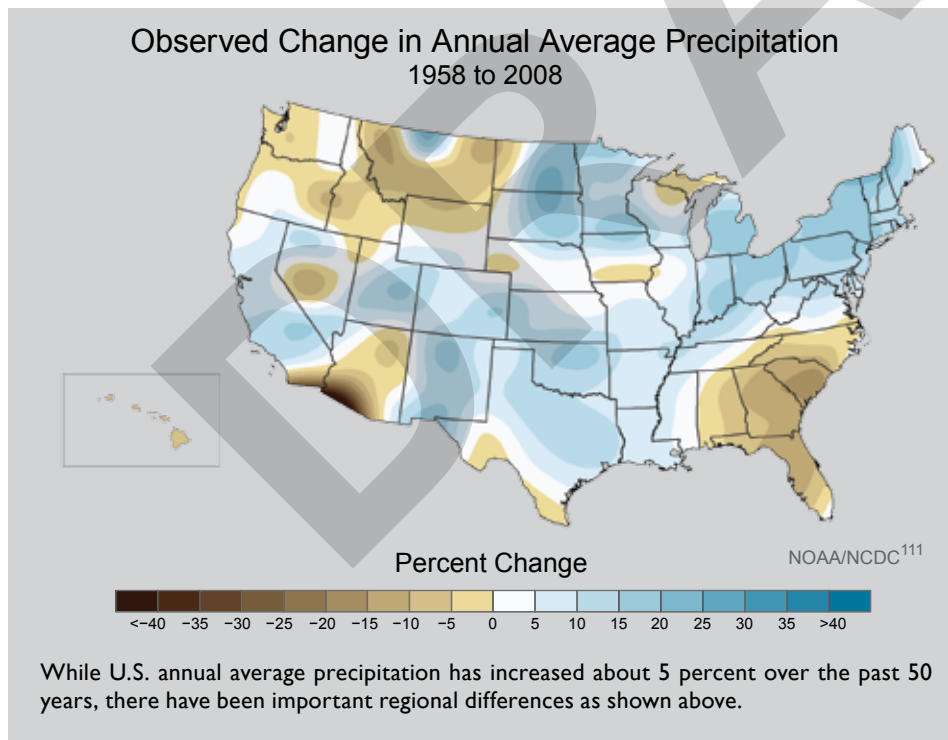
While precipitation over the United States as a whole has increased, there have been important regional and seasonal differences. Increasing trends throughout much of the year have been predominant in the Northeast and large parts of the Plains and Midwest. Decreases occurred in much of the Southeast in all but the fall season and in the Northwest in all seasons except spring. Precipitation also generally decreased during the summer and fall in the Southwest, while winter and spring, which are the wettest seasons in states such as California and Nevada, have had increases in precipitation.¹¹¹

Future changes in total precipitation due to human-induced warming are more difficult to project than changes in temperature. In some seasons, some areas will experience an increase in precipitation, other areas will experience a decrease, and others will see little discernible change. The difficulty arises in predicting the extent of those areas and the amount of change. Model projections of future pre-

cipitation generally indicate that northern areas will become wetter, and southern areas, particularly in the West, will become drier^{97,108}

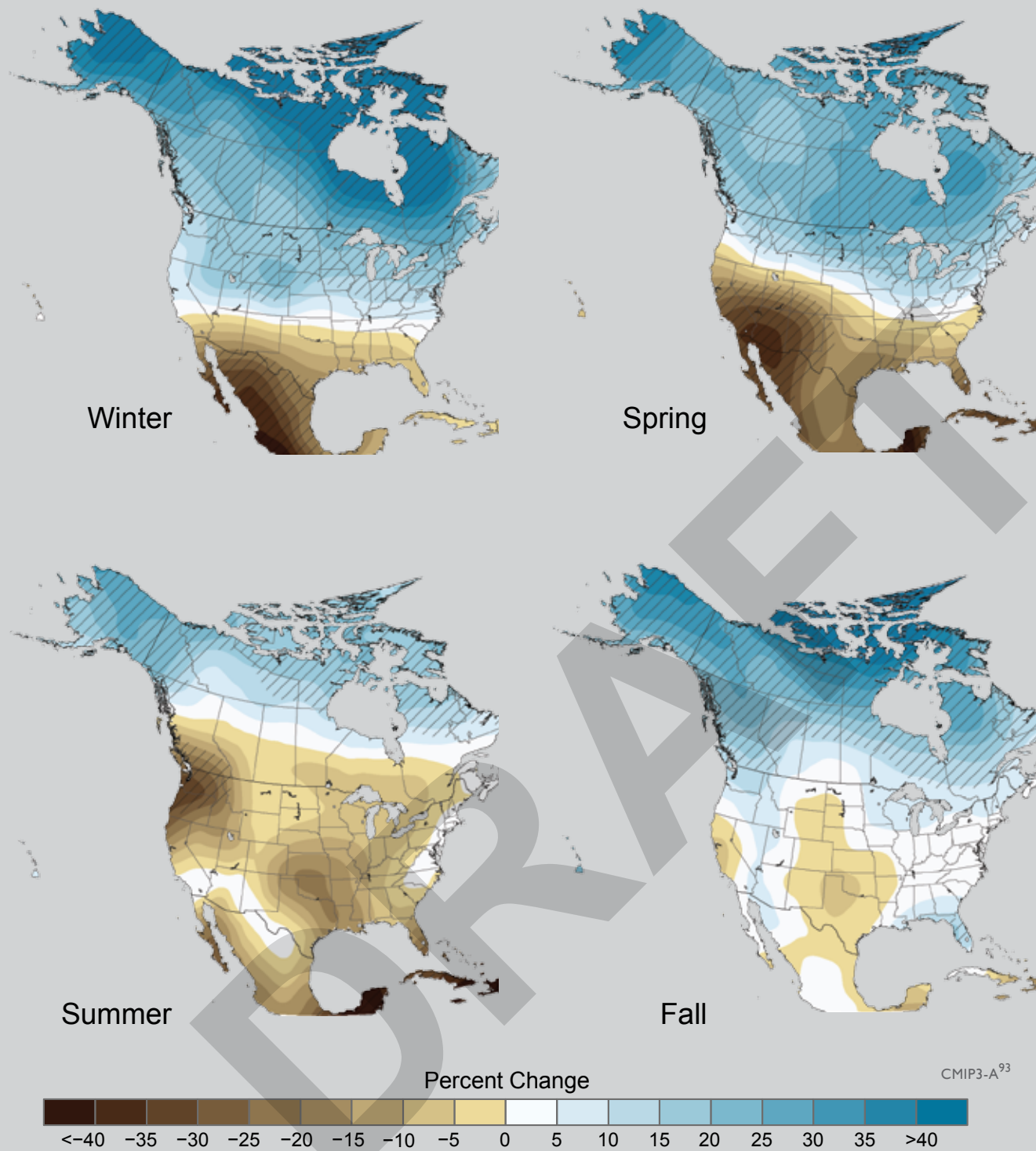
Confidence in projected changes is higher for winter and spring than for summer and fall. In winter and spring, northern areas are expected to receive significantly more precipitation than they do now, because the interaction of warm and moist air coming from the south with colder air from the north is projected to occur farther north than it did on average in the last century. The more northward incursions of warmer and moister air masses are expected to be particularly noticeable in northern regions that will change from very cold and dry atmospheric conditions to warmer but moister conditions.⁶⁸ Alaska, the Great Plains, the upper Midwest, and the Northeast are beginning to experience such changes for at least part of the year, with the likelihood of these changes increasing over time.

In some northern areas, warmer conditions will result in more precipitation falling as rain and less as snow. In addition, potential water resource benefits from increasing precipitation could be countered by the competing influences of increasing evaporation and runoff. In southern areas, significant reductions in precipitation are expected in winter and spring as the subtropical dry belt expands.¹⁰⁸ This is particularly pronounced in the Southwest, where it will have serious ramifications for water resources.

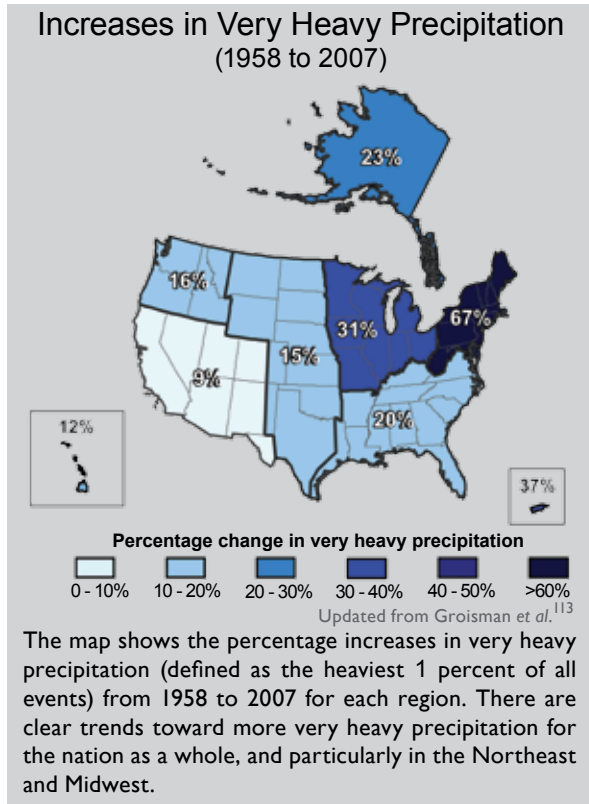


While U.S. annual average precipitation has increased about 5 percent over the past 50 years, there have been important regional differences as shown above.

Projected Change in North American Precipitation
by 2080-2099



The maps show projected future changes in precipitation relative to the recent past as simulated by 15 climate models. The simulations are for late this century, under a higher emissions scenario.⁹¹ For example, in the spring, climate models agree that northern areas are likely to get wetter, and southern areas drier. There is less confidence in exactly where the transition between wetter and drier areas will occur. Confidence in the projected changes is highest in the hatched areas.



The amount of rain falling in the heaviest downpours has increased approximately 20 percent on average in the past century, and this trend is very likely to continue, with the largest increases in the wettest places.

One of the clearest precipitation trends in the United States is the increasing frequency and intensity of heavy downpours. This increase was responsible for most of the observed increase in overall precipitation during the last 50 years. In fact, there has been little change or a decrease in the frequency of light and moderate precipitation during the past 30 years, while heavy precipitation has increased. In addition, while total average precipitation over the nation as a whole increased by about 7 percent over the past century, the amount of precipitation falling in the heaviest 1 percent of rain events increased nearly 20 percent.¹¹²

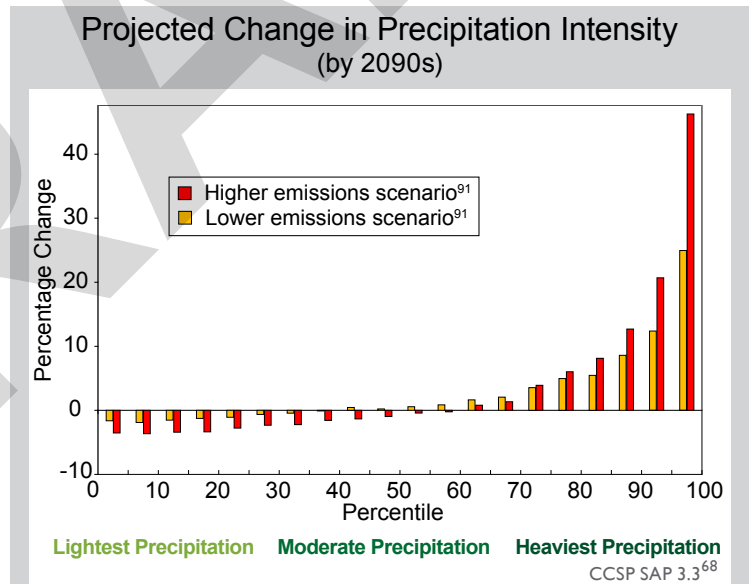
During the past 50 years, the greatest increases in heavy precipitation occurred in the Northeast and the Midwest. There have also been increases in heavy downpours in the other regions of the continental United States, as well as Alaska, Hawaii, and Puerto Rico.¹¹²

Climate models project continued increases in the heaviest downpours during this century, while the lightest precipitation is projected to decrease. Heavy downpours that are now 1-in-20-year occurrences are projected to occur about every 4 to 15 years by the end of this century, depending on location, and the intensity of heavy downpours also is expected to increase. The 1-in-20-year heavy downpour is expected to be between 10 and 25 percent heavier by the end of the century than it is now.¹¹²

Changes in extreme weather and climate events are among the most serious challenges to our nation in coping with a changing climate.

Many types of extreme weather events, such as heat waves and regional droughts, have become more frequent and intense during the past 40 to 50 years.

Many extremes and their associated impacts are now changing. For example, in recent decades most of North America has been experiencing more unusually hot days and nights, fewer unusually cold days and nights, and fewer frost days. Droughts are becoming more severe in some regions. The power and frequency of Atlantic hurricanes have increased substantially in recent decades, though the number of North American mainland landfalling hurricanes does



The figure shows projected changes from the 1990s average to the 2090s average in the intensity of precipitation in North America displayed in 5 percent increments from the lightest drizzles to the heaviest downpours. As shown here, the lightest precipitation is projected to decrease, while the heaviest will increase, continuing the observed trend. The higher emission scenario⁹¹ yields larger changes. Projections are based on the models used in the IPCC 2007 Fourth Assessment Report.

L1 not appear to have increased over the past
 L2 century. Outside the tropics, cold-season
 L3 storm tracks are shifting northward and
 L4 the strongest storms are becoming even
 L5 stronger. These trends in storms outside the
 L6 tropics are projected to continue throughout
 L7 this century.^{68,112,114}

L8 **Drought**

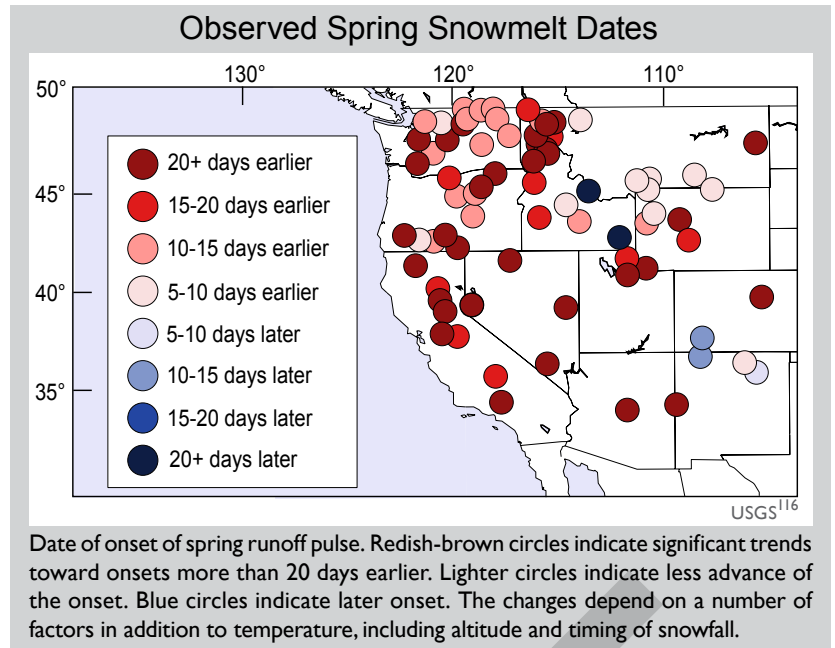
L9 Like precipitation, trends in drought have
 L10 strong regional variations. In much of the
 L11 Southeast and large parts of the West, the
 L12 frequency of drought has increased coincid-
 L13 ent with rising temperatures over the past 50
 L14 years. In other regions, such as the Midwest
 L15 and Great Plains, there has been a reduction
 L16 in drought frequency.
 L17

L18 Although there has been an overall increase
 L19 in precipitation and no clear trend in drought for
 L20 the nation as a whole, increasing temperatures
 L21 have made droughts more severe and widespread
 L22 than they would have otherwise been. Without the
 L23 observed increase in precipitation, higher tempera-
 L24 tures would have led to an increase in the area of
 L25 the contiguous United States in severe to extreme
 L26 drought, with some estimates of a 30 percent
 L27 increase.¹¹² In the future, droughts are likely to be-
 L28 come more frequent and severe in some regions.⁶⁸
 L29 The Southwest, in particular, is expected to experi-
 L30 ence increasing drought as changes in atmospheric
 L31 circulation patterns cause the dry zone just outside
 L32 the tropics to expand farther northward into the
 L33 United States.⁹⁷
 L34

L35 Rising temperatures have also led to earlier melt-
 L36 ing of the snowpack in the western United States.⁴⁰
 L37 Because snowpack runoff is critical to the water
 L38 resources in the western United States, changes in
 L39 the timing and amount of runoff can exacerbate
 L40 problems with already limited water supplies in the
 L41 region.
 L42

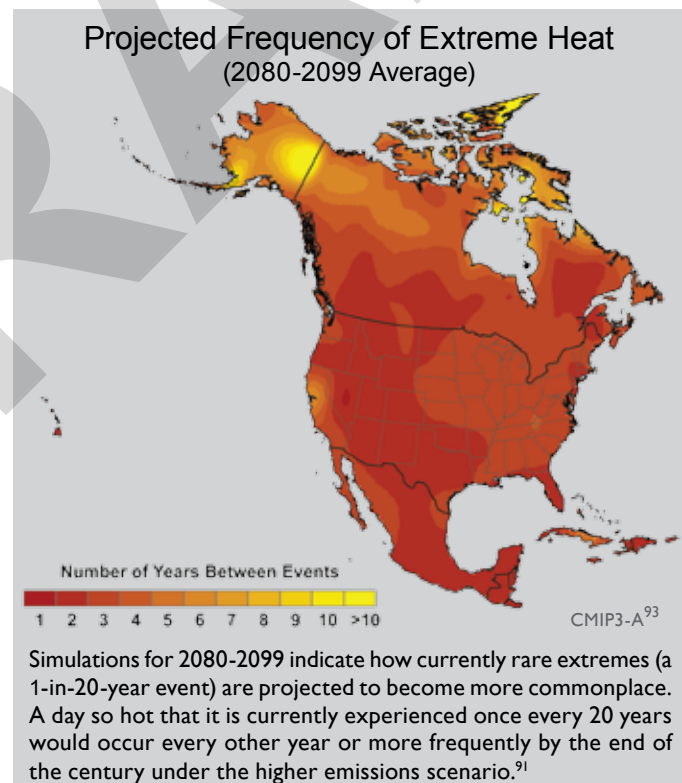
L43 **Heat waves**

L44 A heat wave is a period of several days to weeks
 L45 of abnormally hot weather, often with high humid-
 L46 ity. During the 1930s, there was a high frequency
 L47 of heat waves due to high daytime temperatures
 L48 resulting in large part from an extended multi-year
 L49 period of intense drought. By contrast, in the past 3
 L50



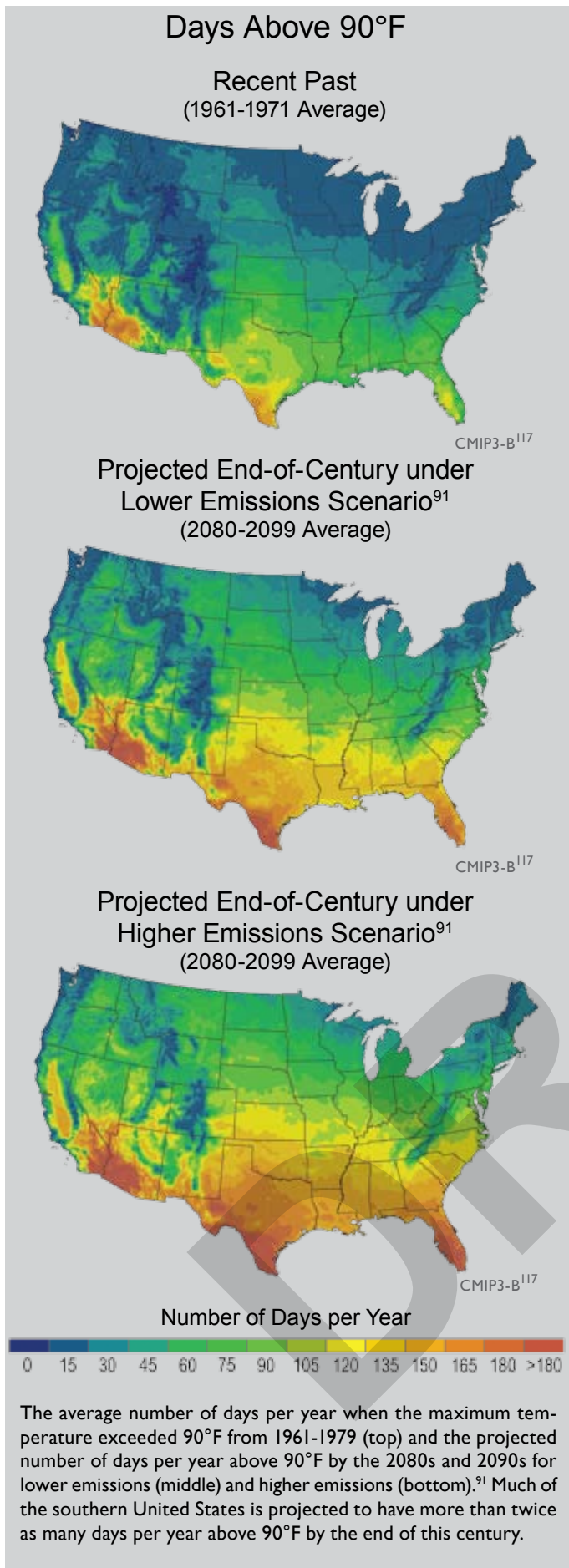
to 4 decades, there has been an increasing trend in high-humidity heat waves, which are characterized by persistence of extremely high nighttime temperatures.¹¹²

As average temperatures continue to rise throughout this century, the frequency of cold extremes will decrease and the frequency and intensity of high temperature extremes will increase.¹¹⁵ The number of days with high temperatures above 90°F



L1
L2
L3
L4
L5
L6
L7
L8
L9
L10
L11
L12
L13
L14
L15
L16
L17
L18
L19
L20
L21
L22
L23
L24
L25
L26
L27
L28
L29
L30
L31
L32
L33
L34
L35
L36
L37
L38
L39
L40
L41
L42
L43
L44
L45
L46
L47
L48
L49
L50

R1
R2
R3
R4
R5
R6
R7
R8
R9
R10
R11
R12
R13
R14
R15
R16
R17
R18
R19
R20
R21
R22
R23
R24
R25
R26
R27
R28
R29
R30
R31
R32
R33
R34
R35
R36
R37
R38
R39
R40
R41
R42
R43
R44
R45
R46
R47
R48
R49
R50



is projected to increase throughout the country as illustrated in the maps on the left. Parts of the South that currently have about 60 days per year with temperatures over 90°F are projected to experience 150 or more days a year above 90°F by the end of this century, under a higher emissions scenario.⁹¹ There is higher confidence in the regional patterns than in results for any specific location (see *An Agenda for Climate Impacts Science* section).

With rising high temperatures, extreme heat waves that are currently considered rare will occur more frequently in the future. Recent studies using an ensemble of models show that events that now occur once every 20 years are projected to occur about every other year in much of the country by the end of this century. In addition to occurring more frequently, at the end of this century these very hot days will be about 10°F hotter than they are today.⁶⁸

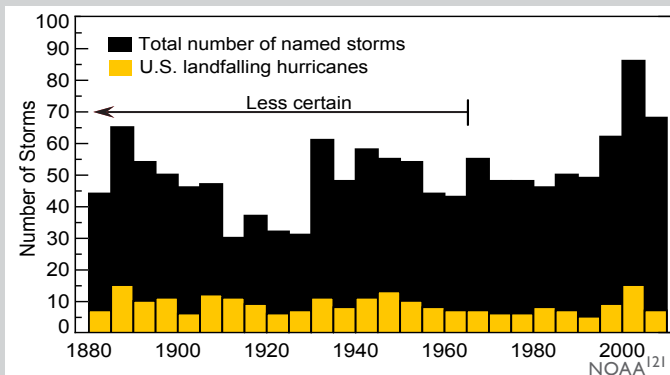
The destructive energy of Atlantic hurricanes has increased in recent decades. The intensity of these storms is likely to increase in this century.

Of all the world's tropical storm and hurricane basins, the North Atlantic has been the most thoroughly monitored and studied. The advent of routine aircraft monitoring in the 1940s and the use of satellite observations since the 1960s have greatly aided monitoring of tropical storms and hurricanes. In addition, observations of tropical storm and hurricane strength made from island and mainland weather stations and from ships at sea began in the 1800s and continue today. Because of new and evolving observing techniques and technologies, scientists pay careful attention to ensuring consistency in tropical storm and hurricane records from the earliest manual observations to today's automated measurements. This is accomplished through collection, analysis, and cross-referencing of data from numerous sources and, where necessary, the application of adjustment techniques to account for differences in observing and reporting methodologies through time. Nevertheless, data uncertainty is larger in the early part of the record. Confidence in the tropical storm and hurricane record increases after 1900 and is greatest during the satellite era, from 1965 to the present.¹¹²

The total number of hurricanes and strongest hurricanes (Category 4 and 5) observed from 1881 through 2008 shows multi-decade periods of above average activity in

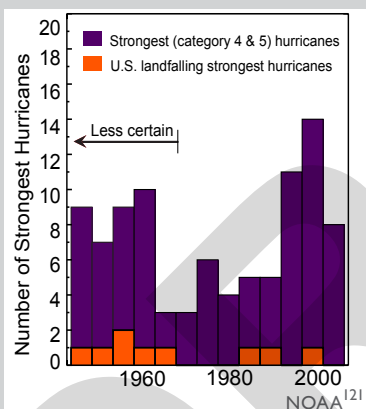
the 1800s, the mid-1900s, and since 1995. The power and frequency of Atlantic hurricanes have increased substantially in recent decades.¹¹² There has been little change in the total number of landfalling hurricanes, in part because a variety of factors affect whether a hurricane will make landfall. These include large-scale steering winds, atmospheric stability, wind shear, and ocean heat content. This highlights the importance of understanding the broader changes occurring throughout the Atlantic Basin beyond the storms making landfall along the U.S. coast.¹¹²

Atlantic Tropical Storms and Hurricanes



Top: Total numbers of North Atlantic named storms (tropical storms and hurricanes) (black) and total U.S. landfalling hurricanes (yellow) in five-year periods based on annual data from 1881 to 2008. The bar for the last 5-year period is based on the assumption that the level of activity from 2006 to 2008 persists through 2010. In the era before satellites, indicated by the arrow above, the total number of named storms is less certain and has been adjusted upward to account for missing storms. Adjustments are based on relationships established during the satellite era between the number of observed storms and the number that would have been missed if satellite data had not been available.

Atlantic Basin Strongest Hurricanes



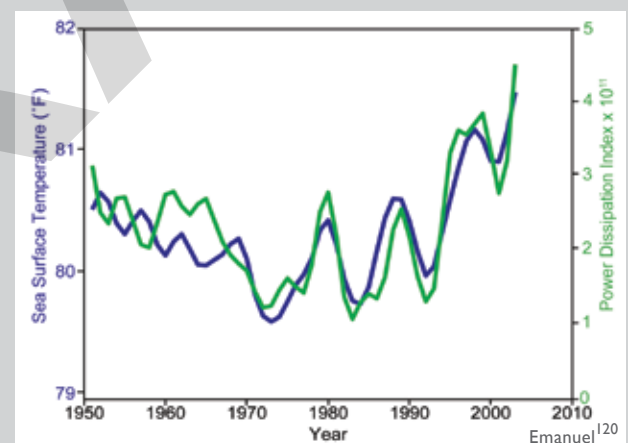
Bottom: Total number of strongest (Category 4 and 5) North Atlantic basin hurricanes (purple) and strongest U.S. landfalling hurricanes (orange) in 5-year periods based on annual data from 1946 to 2008. The bar for the last 5-year period is based on the assumption that the level of activity from 2006 to 2008 persists through 2010. From 1946 to the mid-1960s, as indicated by the arrow above, hurricane intensity was measured primarily by aircraft reconnaissance. Data prior to aircraft reconnaissance are not shown due to the greater uncertainty in estimates of a hurricane's maximum intensity. Satellites have increased the reliability of hurricane intensity estimates since the mid-1960s.

Tropical storms and hurricanes develop and gain strength over warm ocean waters. As oceans warm, they provide a source of energy for hurricane growth. During the past 30 years, annual sea surface temperatures in the main Atlantic hurricane development region increased nearly 2°F. This warming coincided with an increase in the destructive energy (as defined by the Power Dissipation Index, a combination of intensity, duration, and frequency) of Atlantic tropical storms and hurricanes. The strongest hurricanes (Category 4 and 5) have, in particular, increased in intensity.¹¹² The graph

below shows the strong correlation between hurricane power and sea surface temperature in the Atlantic and the overall increase in both during the past 30 years. Climate models project that hurricane intensity will continue to increase, though at a lesser rate than that observed in recent decades.¹⁰⁰

New evidence has emerged recently for other temperature related linkages that can help explain the increase in Atlantic hurricane activity. This includes the contrast in sea surface temperature between the main hurricane development region and the broader tropical ocean.^{99,118,119} Other causes beyond the rise in ocean temperature, such as atmospheric sta-

Observed Relationship Between Sea Surface Temperatures and Hurricane Power in the North Atlantic Ocean



Observed sea surface temperature (blue) and the Power Dissipation Index (green), which combines frequency, intensity and duration for North Atlantic hurricanes.¹²⁰ Hurricane rainfall and wind speeds are likely to increase in response to human-caused warming. Analyses of model simulations suggest that for each 1.8°F increase in tropical sea surface temperatures, rainfall rates will increase by 6 to 18 percent.⁶⁸

bility and circulation, can also influence hurricane power. For these and other reasons, a confident assessment requires further study.⁶⁸

Evidence of increasing hurricane strength in the Atlantic and other oceans with linkages to rising sea surface temperatures is also supported by satellite records dating back to 1981. An increase in the maximum wind speeds of the strongest hurricanes has been documented and linked to increasing sea surface temperatures.¹²²

Projections that sea surface temperatures in the main Atlantic hurricane development region will increase at even faster rates during the second half of this century under higher emissions scenarios⁹¹ highlight the need to better understand the relationship between increasing temperatures and hurricane intensity. As ocean temperatures continue to increase in the future, it is likely that hurricane rainfall and wind speeds will increase in response to human-caused warming.⁶⁸ Analyses of model simulations suggest that for each 1.8°F increase in tropical sea surface temperatures, core rainfall rates will increase by 6 to 18 percent and the surface wind speeds of the strongest hurricanes will increase by about 1 to 8 percent.¹¹⁴ Even without further coastal development, storm surge levels and

hurricane damages are likely to increase because of increasing hurricane intensity coupled with sea-level rise, the latter being a virtually certain outcome of the warming global climate.⁶⁸

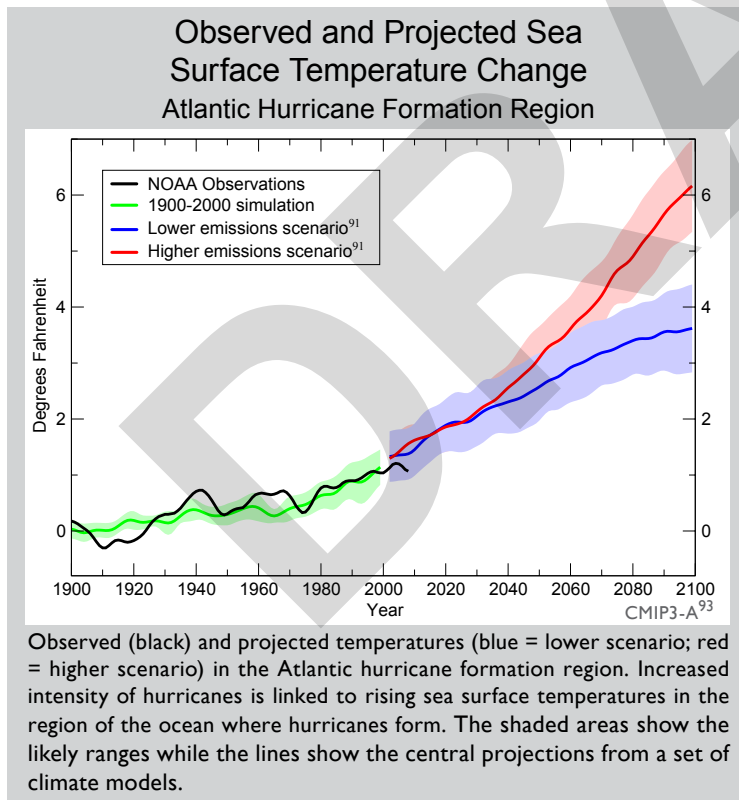
In the eastern Pacific, the strongest hurricanes have become stronger since the 1980s, even while the total number of storms has decreased.

Although on average more hurricanes form in the eastern Pacific than the Atlantic each year, cool ocean waters along the U.S. West Coast and atmospheric steering patterns help protect the contiguous U.S. from landfalls. Threats to the Hawaiian Islands are greater, but landfalling storms are rare in comparison to those of the U.S. East and Gulf Coasts. Nevertheless, changes in hurricane intensity and frequency could influence the impact of landfalling Pacific hurricanes in the future.

The total number of tropical storms and hurricanes in the eastern Pacific on seasonal to multi-decade time periods is generally opposite to that observed in the Atlantic. For example, during El Niño events it is common for hurricanes in the Atlantic to be suppressed while the eastern Pacific is more active.

This reflects the large-scale atmospheric circulation patterns that extend across both the Atlantic and the Pacific oceans.^{123,124}

Within the past three decades the total number of tropical storms and hurricanes and their destructive energy have decreased in the eastern Pacific.^{68,124} However, satellite observations have shown that like the Atlantic, the strongest hurricanes (the top 5 percent), have gotten stronger since the early 1980s.^{122,125} As ocean temperatures rise, the strongest hurricanes are likely to increase in both the eastern Pacific and the Atlantic.⁶⁸



Sea level has risen along most of the U.S. coast over the past 50 years, and will rise more in the future.

Recent global sea-level rise has been caused by the warming-induced expansion of the oceans, accelerated melting of most of the world’s glaciers and ice caps, and loss of ice on the Greenland and Antarctic ice sheets.³⁷ There is strong evidence that global sea level is currently rising at an increased rate.^{37,126} A warming global climate will cause further sea-level rise over this century and beyond.^{90,105}

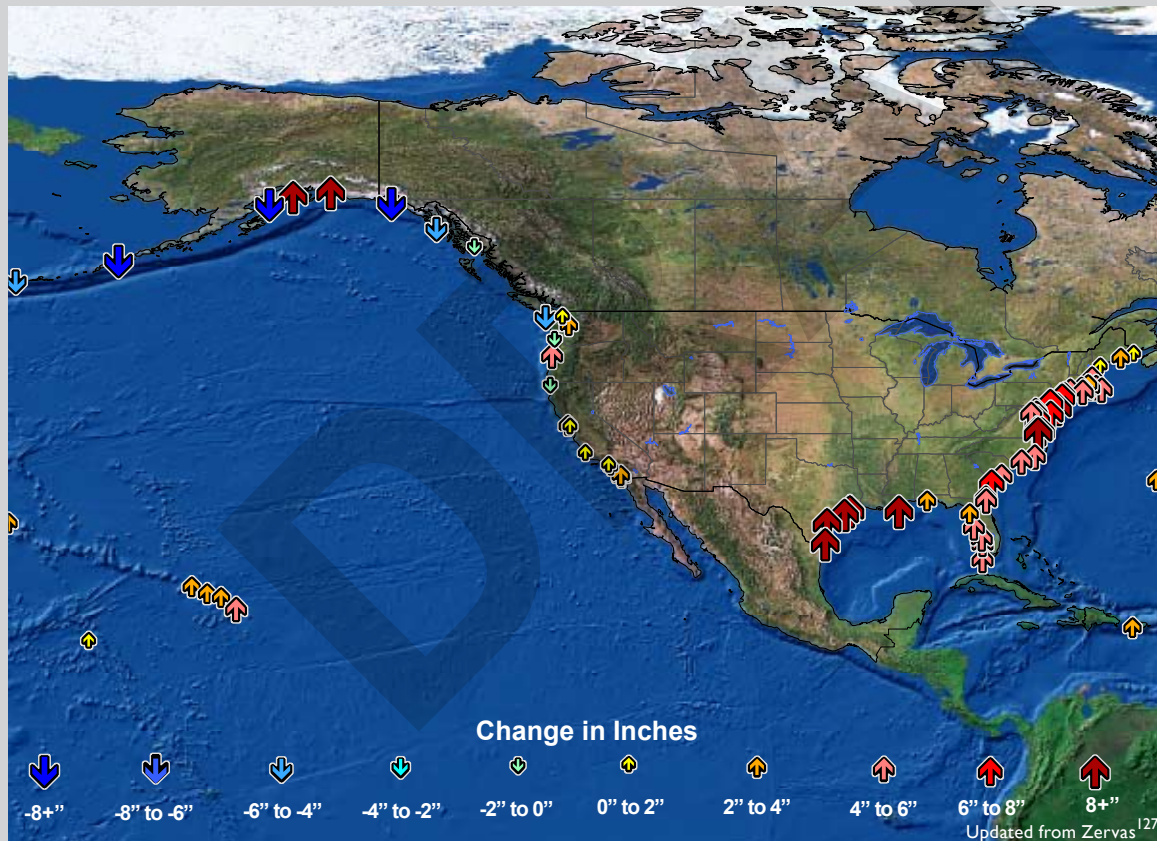
During the past 50 years, sea level has risen up to 8 inches or more along some coastal areas of the United States, and has fallen in other locations. The amount of relative sea-level rise experienced along different parts of the U.S. coast depends on the changes in elevation of the land that occur as a result of subsidence (sinking) or uplift (rising), as well as increases in global sea level due to warming. In addition, atmospheric and oceanic circulation, which will be affected by climate change, will influence regional sea level. Regional differences

in sea-level rise are also expected to be related to where the meltwater originates.¹⁰⁴

Human-induced sea-level rise is occurring globally. Large parts of the Atlantic Coast and Gulf of Mexico Coast have experienced significantly higher rates of relative sea-level rise than the global average during the last 50 years, with the local differences mainly due to land subsidence.¹²⁷ Portions of the Northwest and Alaska coast have, on the other hand, experienced slightly falling sea level as a result of long-term uplift as a consequence of glacier melting and other geological processes.

Regional variations in relative sea-level rise are expected in the future. For example, assuming historical geological forces continue, a 2-foot rise in global sea level (which is within the range of recent estimates) by the end of this century would result in a relative sea-level rise of 2.3 feet at New York City, 2.9 feet at Hampton Roads, Virginia, 3.5 feet at Galveston, Texas, and 1 foot at Neah Bay in Washington state.¹²⁸

Relative Sea-Level Changes on U.S. Coastlines, 1958 to 2008



Observed changes in relative sea level from 1958 to 2008 for locations on the U.S. coast. Some areas along the Atlantic and Gulf coasts saw increases greater than 8 inches over the past 50 years. Updated from Zervas¹²⁷

Cold-season storm tracks are shifting northward and the strongest storms are likely to become stronger and more frequent.

Large-scale storm systems are the dominant weather phenomenon during the cold season in the United States. Although the analysis of these storms is complicated by a relatively short length of most observational records and by the highly variable nature of strong storms, some clear patterns have emerged.¹¹²

Storm tracks have shifted northward over the last 50 years as evidenced by a decrease in the frequency of storms in mid-latitude areas of the Northern Hemisphere, while high-latitude activity has increased. There is also evidence of an increase in the intensity of extra-tropical storms in both the mid- and high-latitude areas of the Northern Hemisphere, but there is greater confidence in the increases occurring in high latitudes.¹¹² The northward shift is projected to continue, and strong cold season storms are likely to become stronger and more frequent, with greater wind speeds and more extreme wave heights.⁶⁸

Snowstorms

The northward shift in storm tracks is reflected in regional changes in the frequency of snowstorms. The South and lower Midwest saw reduced snowstorm frequency during the last century. In contrast, the Northeast and upper Midwest saw increases in snowstorms, although considerable decade-to-decade variations were present in all regions, influenced, for example, by the frequency of El Niño events.¹¹²

There is also evidence of an increase in lake-effect snowfall along and near the southern and eastern shores of the Great Lakes since 1950.⁹⁷ Lake-effect snow is produced by the strong flow of cold air across large areas of relatively warmer ice-free water. As the climate has warmed, ice coverage on the Great Lakes has fallen. The maximum seasonal coverage of Great Lakes ice decreased at a rate of 8.4 percent per decade from 1973 through 2008, amounting to a roughly 30 percent decrease in ice coverage (see *Midwest* region). This has created conditions conducive to greater evaporation of



Areas in New York state east of Lake Ontario received over 10 feet of lake-effect snow during a 10-day period in early February 2007.

moisture and thus heavier snowstorms. Among recent extreme lake-effect snow events was a February 2007 10-day storm total of over 10 feet of snow in western New York state. Climate models suggest that lake-effect snowfalls are likely to increase over the next few decades. In the longer term, lake-effect snows are likely to decrease as temperatures continue to rise, with the precipitation falling as rain.^{129,130}

Tornadoes and severe thunderstorms

Reports of severe weather including tornadoes and severe thunderstorms have increased during the past 50 years. However, the increase in the number of reports is widely believed to be due to improvements in monitoring technologies such as Doppler radars combined with changes in population and increasing public awareness. When adjusted to account for these factors, there is no clear trend in the frequency or strength of tornadoes since the 1950s for the United States as a whole.¹¹²

The distribution by intensity for the strongest 10 percent of hail and wind reports is little changed, providing no evidence of an observed increase in the severity of events.¹¹² Climate models project future increases in the frequency of environmental conditions favorable to severe thunderstorms.¹³¹ But the inability to adequately model the small-scale conditions involved in thunderstorm development remains a limiting factor in projecting the future character of severe thunderstorms and other small-scale weather phenomena.⁶⁸

Arctic sea ice is declining rapidly and this is very likely to continue.

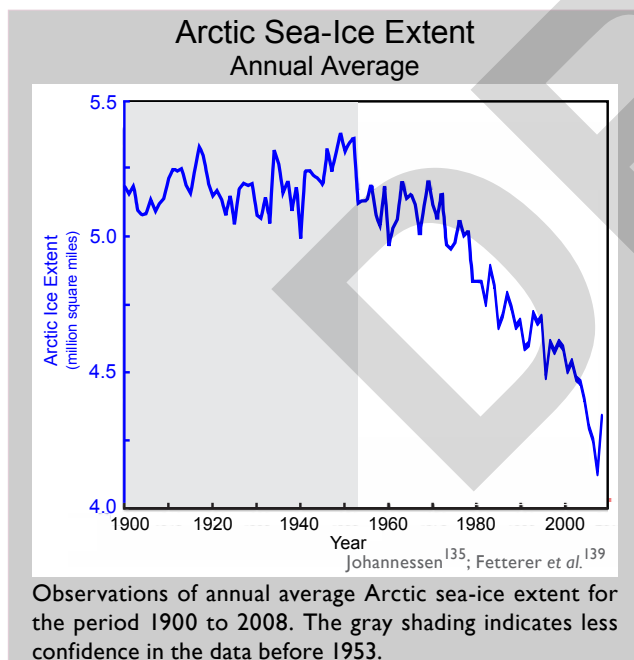
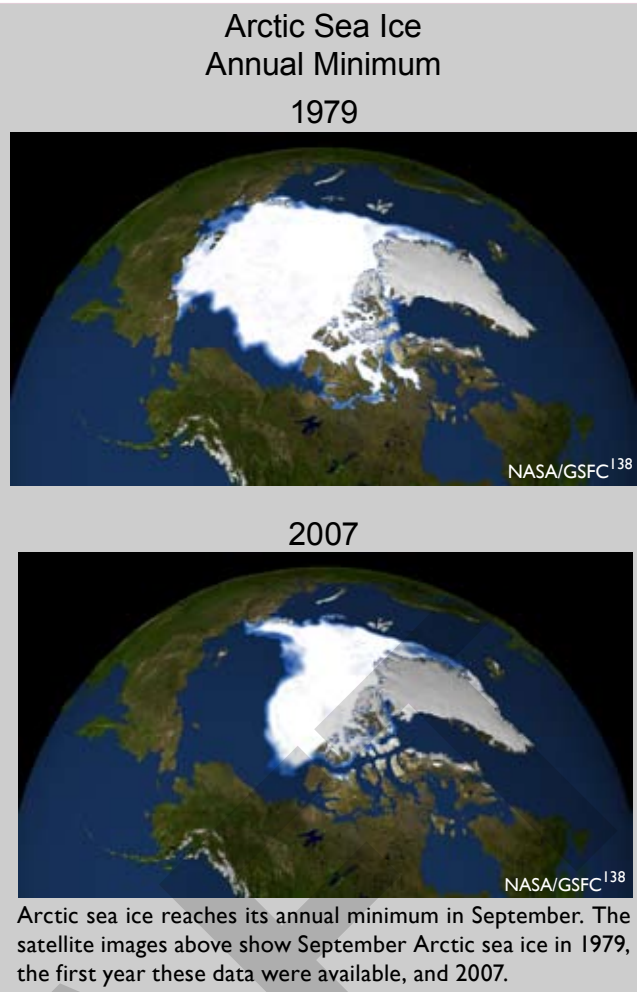
Sea ice is a very important part of the climate system. In addition to direct impacts on coastal areas of Alaska, it more broadly affects surface reflectivity, ocean currents, cloudiness, humidity, and the exchange of heat and moisture at the ocean's surface. Open ocean water is darker in color than sea ice, which causes it to absorb more of the Sun's heat, which increases the warming of the water even more.^{40,132}

The most complete record of sea ice is provided by satellite observations of sea-ice extent since the 1970s. Prior to that, aircraft, ship, and coastal observations in the Arctic make it possible to extend the record of Northern Hemisphere sea-ice extent back to at least 1900, although there is a lower level of confidence in the data prior to 1953.⁴⁰

Arctic sea-ice extent has fallen at a rate of 3 to 4 percent per decade over the last three decades. End-of-summer Arctic sea ice has fallen at an even faster rate of more than 11 percent per decade in that time. The observed decline in Arctic sea ice has been more rapid than projected by climate models.¹³³ Year-to-year changes in sea-ice extent and record low values are influenced by natural variations in atmospheric pressure and wind patterns.¹³⁴

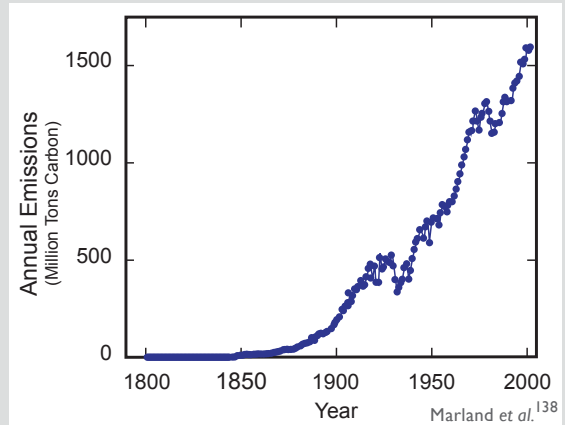
However, clear linkages between rising greenhouse gas concentrations and declines in Arctic sea ice have been identified in the climate record as far back as the early 1990s.⁶¹ The extreme loss in Arctic sea ice that occurred in 2007 would not have been possible without the long-term reductions that have coincided with a sustained increase in the atmospheric concentration of carbon dioxide and the rapid rise in global temperatures that have occurred since the mid-1970s.¹³⁵ Although the 2007 record low was not eclipsed in 2008, the 2008 sea-ice extent is well below the long-term average, reflecting a continuation of the long-term decline in Arctic sea ice. In addition, the total volume of Arctic sea ice in 2008 was likely a record low because the ice was unusually thin.¹³⁶

It is expected that declines in Arctic sea ice will continue in the coming decades with year-to-year fluctuations influenced by natural atmospheric variability. The overall rate of decline will be influenced mainly by the rate at which carbon dioxide and other greenhouse gas concentrations increase.¹³⁷



U.S. Emission and Absorption of Heat-Trapping Gases

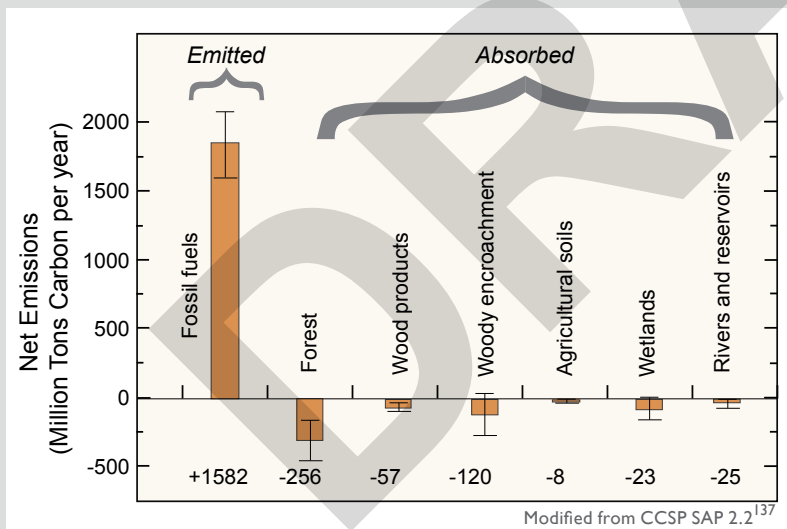
Since the industrial revolution, the United States has been the world's largest emitter of heat-trapping gases. With 4.5 percent of world's population, the United States is responsible for about 28 percent of the human-induced heat-trapping gases in the atmosphere today.¹³⁶ Although China has recently surpassed the United States in current total annual emissions, per capita emissions remain much higher in the United States. Carbon dioxide, the most important of the heat-trapping gases produced directly by human activities, is a cumulative problem because it has a long atmospheric lifetime. Roughly one-half of the carbon dioxide released from fossil fuel burning remains in the atmosphere after 100 years, and roughly one-fifth of it remains after 1,000 years.⁹⁰



U.S. annual emissions of CO₂ from fossil-fuel use.¹³⁸

U.S. carbon dioxide emissions grew dramatically over the past century. These emissions come almost entirely from burning fossil fuels. These sources of carbon dioxide are one side of the equation and on the other side are “sinks” that take up carbon dioxide. The growth of trees and other plants is an important natural carbon sink. In recent years, it is estimated that about 20 percent of U.S. carbon dioxide emissions have been offset by U.S. forest growth and other sinks (see figure below).¹³⁷ It is not known whether U.S. forests and other sinks will continue to take up roughly this amount of carbon dioxide in the future as climate change alters carbon release and uptake. For example, a warming-induced lengthening of the growing season would tend to increase carbon uptake. On the other hand, the increases in forest fires and in the decomposition rate of dead plant matter would decrease uptake, and might convert the carbon sink into a source.¹³⁷

The amount of carbon released and taken up by natural sources varies considerably from year to year depending on climatic and other conditions. For example, fires release carbon dioxide, so years with many



U.S. carbon dioxide emissions and uptake in millions of tons of carbon per year in 2003. The bar marked “Emitted” indicates the amount of carbon as carbon dioxide added to the atmosphere from U.S. emissions. The bars marked “Absorbed” indicate amounts of carbon as carbon dioxide removed from the atmosphere. The thin lines on each bar indicate estimates of uncertainty.

large fires result in more carbon release and less uptake as natural sinks (the vegetation) are lost. Similarly, the trees destroyed by intense storms or droughts release carbon dioxide as they decompose, and the loss results in reduced strength of natural sinks until regrowth is well underway. For example, Hurricane Katrina killed or severely damaged over 320 million large trees. As these trees decompose over the next few years, they will release an amount of carbon dioxide equivalent to that taken up by all U.S. forests in a year.¹¹² The net change in carbon storage in the long run will depend on how much is taken up by the regrowth as well as how much was released by the original disturbance.

Water Resources

Key Messages:

- Climate change has already altered, and will continue to alter, the water cycle, affecting where, when, and how much water is available for all uses.
- Floods and droughts are likely to become more common and more intense as regional and seasonal precipitation patterns change, and rainfall becomes more concentrated into heavy events (with longer, hotter dry periods in between).
- Precipitation and runoff are likely to increase in the Northeast and Midwest in winter and spring, and decrease in the West, especially the Southwest, in spring and summer.
- In areas where snowpack dominates, the timing of runoff will continue to shift to earlier in the spring and flows will be lower in late summer.
- Surface water quality and groundwater quantity will be affected by a changing climate.
- Climate change will place additional burdens on already stressed water systems.
- The past century is no longer a reasonable guide to the future for water management.

Key Sources

CCSP 3.3	CCSP 3.4	CCSP 4.3	CCSP 4.5	CCSP 4.6	CCSP 4.7
Extremes	Abrupt Climate Change	Impacts	Energy	Health	Transportation
CCSP 5.1	CCSP 5.3	IPCC WG-1	IPCC WG-2	IPCC Water	
Data Uses & Limitations	Decision Support				

Changes in the water cycle, which are consistent with the warming observed over the past several decades, include:

- changes in precipitation patterns and intensity
- changes in the incidence of drought
- widespread melting of snow and ice
- increasing atmospheric water vapor
- increasing evaporation
- increasing water temperatures
- reductions in lake and river ice
- changes in soil moisture and runoff

For the future, marked regional differences are projected, with increases in annual precipitation, runoff, and soil moisture in much of the Midwest and Northeast, and declines in much of the West and Southwest.



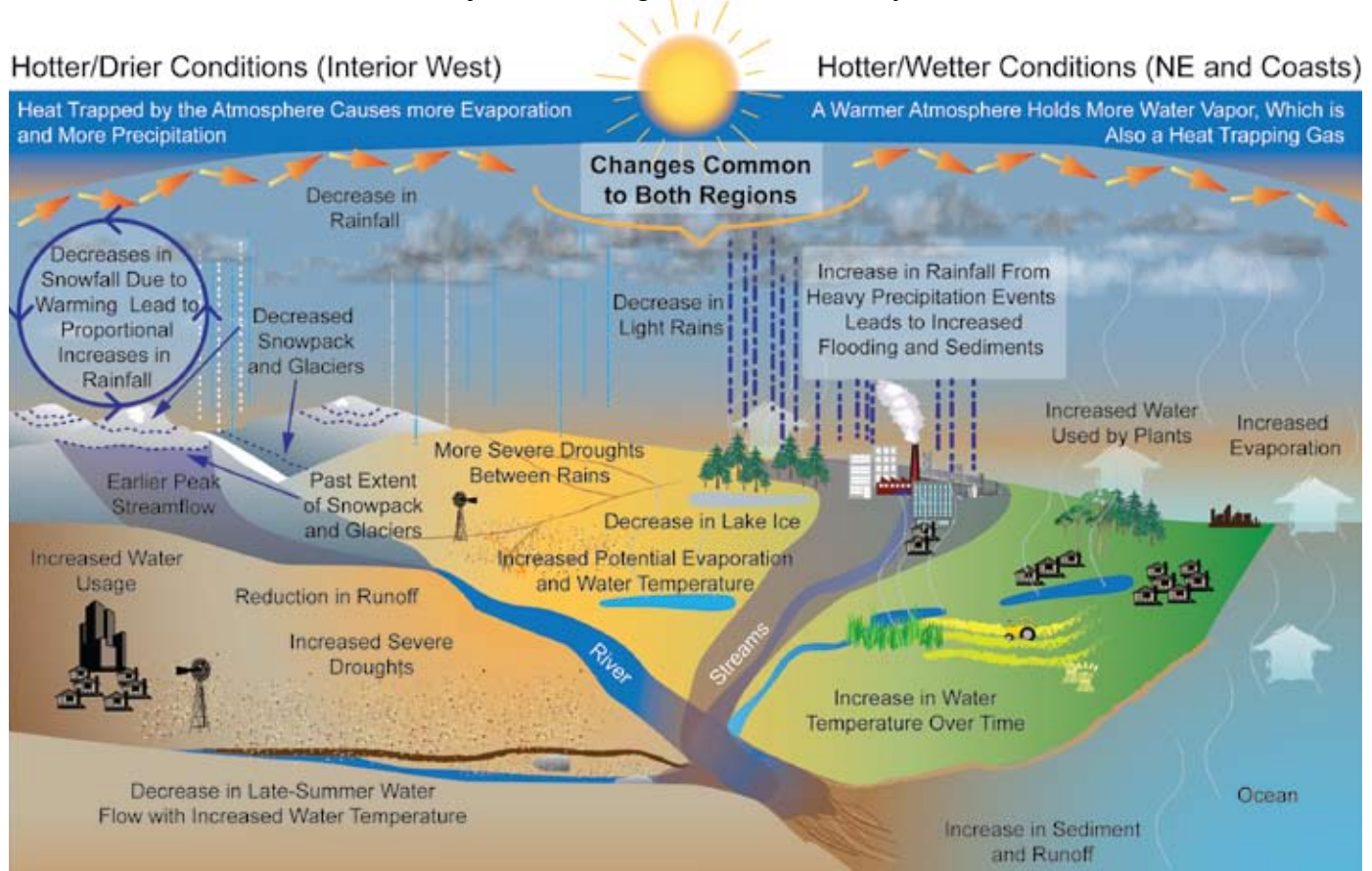
Skagit River and surrounding mountains in the Northwest

Climate change impacts include too little water in some places, too much water in other places, and degraded water quality. Some locations will be subject to all of these conditions during different times of the year. Water cycle changes are expected to continue and will adversely affect energy production and use, human health, transportation, agriculture, and ecosystems.¹⁴²

Climate change has already altered, and will continue to alter, the water cycle, affecting where, when, and how much water is available for all uses.

Substantial changes to the water cycle are expected as the planet warms because the movement of water in the atmosphere and oceans is one of the primary mechanisms for the redistribution of heat around the world. Evidence is mounting that human-induced climate change is already altering many of the existing patterns of precipitation in the United States, including when, where, how much, and what kind of precipitation falls.^{68,142} A warmer climate increases evaporation of water from land and sea, and allows more moisture to be held in the atmosphere. For every 1°F rise in temperature, the water holding capacity of the atmosphere increases by about 4 percent.⁴⁹

Projected Changes in the Water Cycle



The water cycle exhibits many changes as the earth warms. Wet and dry areas respond differently.

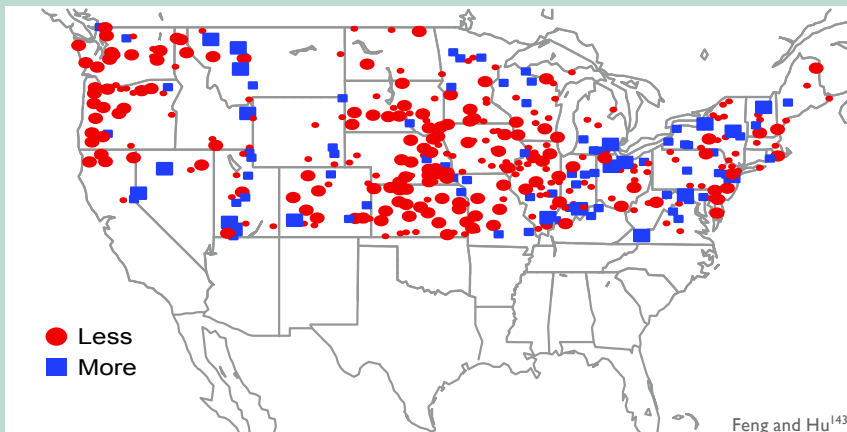
NOAA/NCDC

In addition, changes in atmospheric circulation will tend to move storm tracks northward with the result that dry areas will become drier and wet areas wetter. Hence, the arid Southwest will experience longer and more severe droughts from the combination of increased evaporation and reductions in precipitation.¹⁰⁸

The additional atmospheric moisture contributes to more overall precipitation in some areas, especially in much of the Northeast and Alaska. Over the past 50 years, precipitation and streamflow have increased in much of the Northeast and Midwest, with a reduction in drought duration and severity. Much of the Southeast and West has had reductions in precipitation and increases in drought severity and duration, especially in the Southwest.

In most areas of the country, the fraction of precipitation falling as rain versus snow has increased during the last 50 years. Despite this general shift from snow to rain, snowfalls along

Changes in Snowfall Contributions to Wintertime Precipitation 1949 to 2005

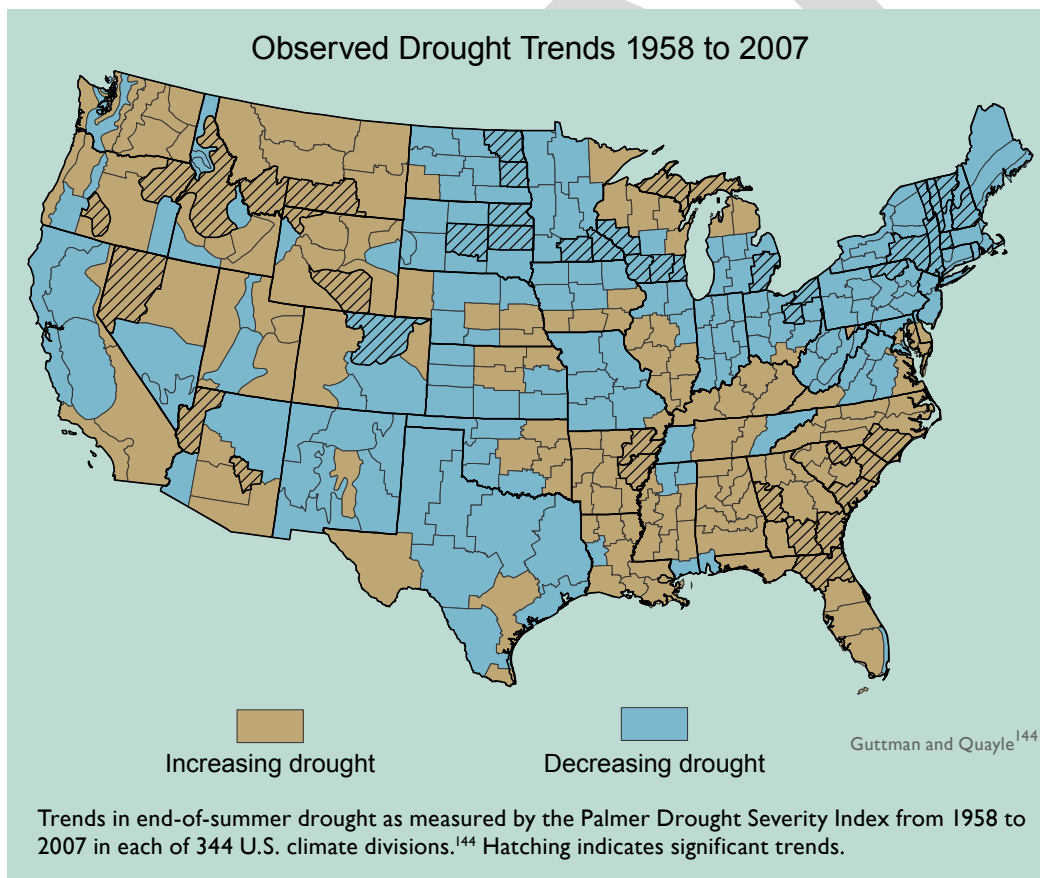


Trends in winter snow-to-total precipitation ratio from 1949 to 2005. Red circles indicate less snow, while blue squares indicate more snow. Large circles and squares indicate the most significant trends.¹⁴³ Areas south of 37°N latitude were excluded from the analysis because most of that area receives little snowfall. White areas above that line have inadequate data for this analysis.

Observed Water-Related Changes During the Last Century¹⁴²

Observed Change	Direction of Change	Region Affected
One to four week earlier peak streamflow due to earlier warming-driven snowmelt	Earlier	West and Northeast
Proportion of precipitation falling as snow	Decreasing	West and Northeast
Duration and extent of snow cover	Decreasing	Most of the United States
Mountain snow water equivalent	Decreasing	West
Annual precipitation	Increasing	Most of the United States
Annual precipitation	Decreasing	Southwest
Frequency of heavy precipitation events	Increasing	Most of the United States
Runoff and streamflow	Decreasing	Colorado and Columbia River Basins
Streamflow	Increasing	Most of East
Amount of ice in mountain glaciers	Decreasing	U.S. western mountains, Alaska
Water temperature of lakes and streams	Increasing	Most of the United States
Ice cover on lakes and rivers	Decreasing	Great Lakes and Northeast
Periods of drought	Increasing	Parts of West and East
Salinization of surface waters	Increasing	Florida, Louisiana
Widespread thawing of permafrost	Increasing	Alaska

Observed Drought Trends 1958 to 2007



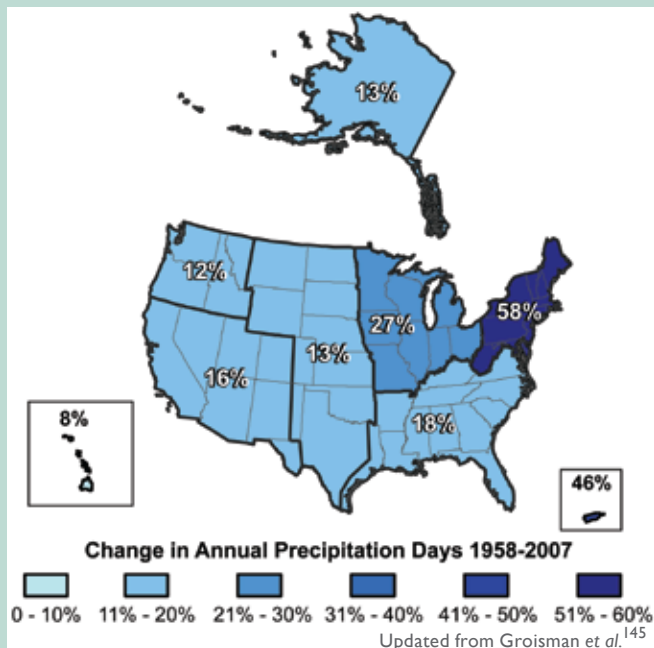
L1 the downwind coasts of the Great Lakes have
 L2 increased. Factors contributing to this increase
 L3 include reduced ice cover due to warming,
 L4 which lengthens the period of open water.
 L5 In addition, cold air moving over relatively
 L6 warm, open lake water induces strong evapo-
 L7 ration, often causing heavy lake-effect snow.
 L8 Heavy snowfall and snowstorm frequency
 L9 have increased in many northern parts of the
 L10 United States. In the South however, where
 L11 temperatures are already marginal for heavy
 L12 snowfall, climate warming has led to a reduc-
 L13 tion in heavy snowfall and snowstorm fre-
 L14 quency. These trends suggest a northward shift
 L15 in snowstorm occurrence.⁶⁸

Floods and droughts are likely to become more common and more intense as regional and seasonal precipitation patterns change, and rainfall becomes more concentrated into heavy events (with longer, hotter dry periods in between).

L16 While it sounds counterintuitive, a warmer
 L17 world produces both wetter and drier conditions.
 L18 Even though total global precipitation increases,
 L19 the regional and seasonal distribution of precipi-
 L20 tation changes, and more precipitation comes in
 L21 heavier rains (which can cause flooding) rather
 L22 than light events. In the past century, averaged over
 L23 the United States, total precipitation has increased
 L24 by about 7 percent, while the heaviest 1 percent of
 L25 rain events increased by nearly 20 percent.⁶⁸ This
 L26 has been especially noteworthy in the East, where
 L27 the annual number of days with very heavy precipi-
 L28 tation has increased most in the past 50 years, as
 L29 shown in the adjacent figure. Flooding often occurs
 L30 when heavy precipitation persists for weeks to
 L31 months in large river basins. Such extended periods
 L32 of heavy precipitation have also been increasing
 L33 over the past century, most notably in the past two
 L34 to three decades in the United States.¹¹²

L45 Observations also show that over the past several
 L46 decades, extended dry periods have become more
 L47 frequent in parts of the United States, especially
 L48 the Southwest and the eastern United States.^{146,147}
 L49 Longer periods between rainfalls, combined with
 L50

Increases in Very Heavy Precipitation Days (1958 to 2007)



R1 The map shows the percentage increases in the average number
 R2 of days with very heavy precipitation (defined as the heaviest 1
 R3 percent of all events) from 1958 to 2007 for each region. There
 R4 are clear trends toward more very heavy precipitation days for
 R5 the nation as a whole, and particularly in the Northeast and
 R6 Midwest.

R7 higher air temperatures, dry out soils and vegeta-
 R8 tion, causing drought.

R9 For the future, precipitation intensity is projected
 R10 to increase everywhere, with the largest increases
 R11 occurring in areas in which average precipitation
 R12 increases the most. For example, the Midwest and
 R13 Northeast, where total precipitation is expected to
 R14 increase the most, will also experience the largest
 R15 increases in heavy precipitation events. The num-
 R16 ber of dry days between precipitation events is also
 R17 projected to increase, especially in the more arid
 R18 areas. Mid-continental areas and the Southwest
 R19 are particularly threatened by future drought. The
 R20 magnitude of the projected changes in extremes is
 R21 expected to be greater than changes in averages,
 R22 and hence detectable sooner.^{49,68,90,142,148,}

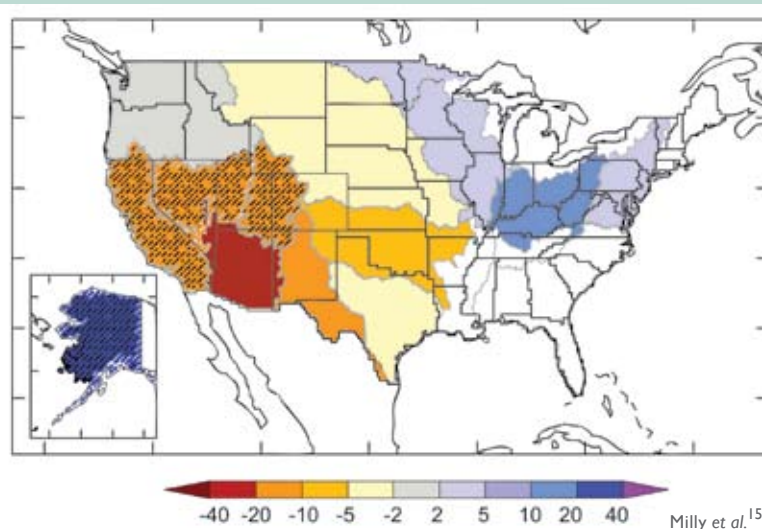
Precipitation and runoff are likely to increase in the Northeast and the Midwest in winter and spring, and decrease in the West, especially the Southwest, in the spring and summer.

Runoff, which accumulates as streamflow, is the amount of precipitation that is not evaporated, stored as snowpack or soil moisture, or filtered down to groundwater. The proportion of precipitation that runs off is determined by a variety of factors, including temperature, wind speed, humidity, solar intensity at the ground, vegetation, and soil moisture. While runoff generally tracks precipitation, increases and decreases in precipitation do not necessarily lead to equal increases and decreases in runoff. For example, droughts cause soil moisture reductions that can reduce expected runoff until soil moisture is replenished. Conversely, water-saturated soils can generate floods with only moderate additional precipitation. During the last century, consistent increases in precipitation have been found in the Midwest and Northeast along with increased runoff.^{149,150} Climate models consistently project that the East will experience increased runoff, while there will be substantial declines in the interior West, especially the Southwest. Projections for runoff in California and other parts of the West also show reductions, although less than in the interior West. In short, wet areas are projected to get wetter and dry areas drier. Climate models also consistently project heat-related summer soil moisture reductions in the middle of the continent.^{115,142,146,149}

In areas where snowpack dominates, the timing of runoff will continue to shift to earlier in the spring and flows will be lower in late summer.

Large portions of the West and some areas in the Northeast rely on snowpack as a natural reservoir to hold winter precipitation until it later runs off as streamflow in spring, summer, and fall. Over the last 50

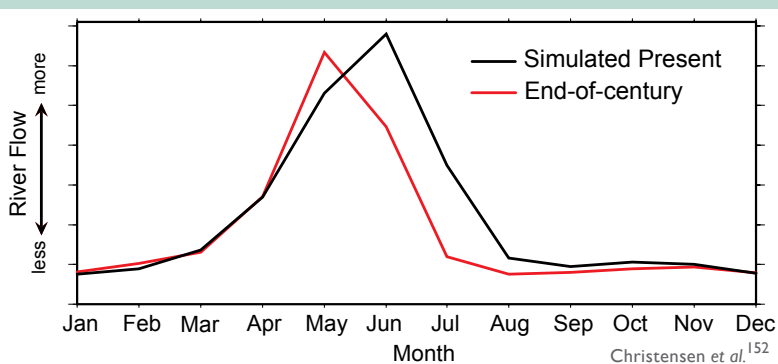
Projected Changes in Annual Runoff



Projected changes in median runoff for 2041-2060, relative to a 1901-1970 baseline, are mapped by water-resource region. Colors indicate percentage changes in runoff. Hatched areas indicate greater confidence due to strong agreement among model projections. U.S. white areas indicate divergence among model projections. Results are based on emissions in between the lower and higher emissions scenarios.⁹¹

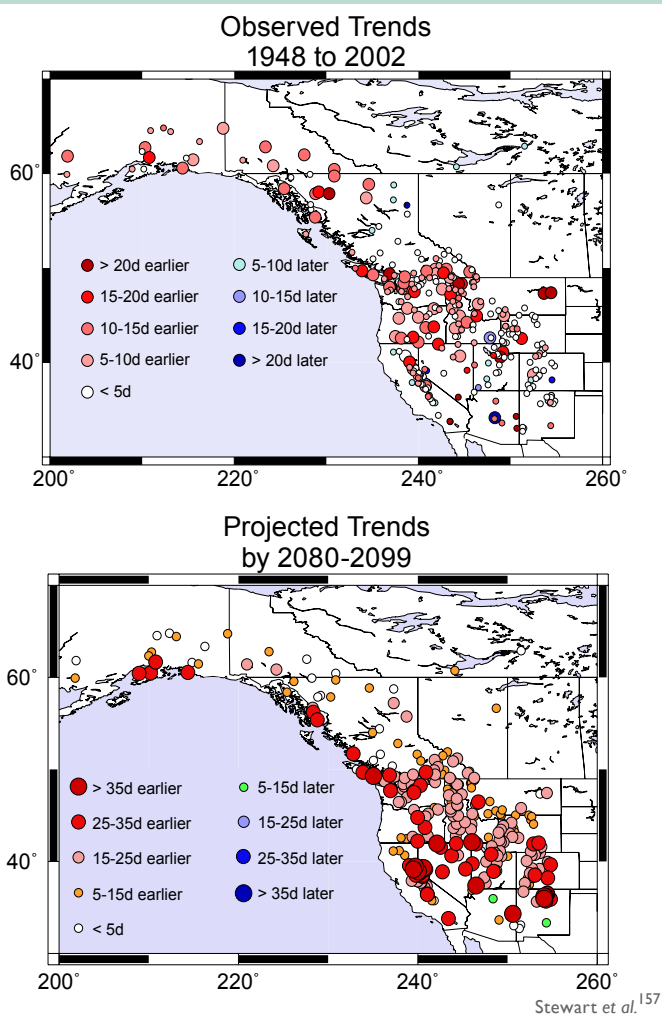
years, there have been widespread temperature-related reductions in snowpack in the West, with the largest reductions occurring in lower elevation mountains in the Northwest and California where snowfall occurs at temperatures close to the freezing point.^{142,153} The Northeast has also experienced snowpack reductions during a similar period. Observations indicate a transition to more rain and less snow in both the West and Northeast in the last 50 years.^{143,154-156} Runoff in snowmelt-dominated areas is occurring up to 20 days earlier in the West, and up to 14 days earlier in the Northeast.^{157,158} Future projections for most snowmelt-dominated basins in the West consistently indicate earlier spring

Projected Changes in Annual Runoff Pattern



General schematic of changes in the annual pattern of runoff for snowmelt-dominated streams. Compared to the historical pattern, runoff peak is projected to shift to earlier in the spring and late summer flows are expected to be lower. The above example is for the Green River, which is part of the Colorado River watershed.¹⁵²

Trends in Peak Streamflow Timing



Top map shows changes in runoff timing in snowmelt-driven streams from 1948 to 2002 with red circles indicating earlier runoff, and blue circles indicating later runoff. Bottom map shows projected changes in snowmelt-driven streams by 2080-2099, compared to 1951-1980, under a higher emissions scenario.⁹¹

runoff, in some cases up to 60 days earlier.^{157,159} For the Northeast, projections indicate spring runoff will advance by up to 14 days.¹⁵⁰ Earlier runoff produces lower late-summer streamflows, which stress human and environmental systems through less water availability and higher water temperatures.¹⁴⁵ Scientific analyses to determine the causes of recent changes in snowpack, runoff timing, and increased winter temperatures have attributed these changes to human-caused climate change.^{34,160,161}

Surface water quality and groundwater quantity will be affected by a changing climate.

Changes in water quality

Increased air temperatures lead to higher water temperatures, which have already been detected in many streams, especially during low-flow periods. In lakes and reservoirs, higher water temperatures lead to longer periods of summer stratification (when surface and bottom waters do not mix). Dissolved oxygen is reduced in lakes, reservoirs, and rivers at higher temperatures. Oxygen is an essential resource for many living things, and its availability is reduced at higher temperatures both because the amount that can be dissolved in water is lower and because respiration rates of living things are higher. Low oxygen stresses aquatic animals such as cold-water fish and the insects and crustaceans on which they feed.¹⁴² Lower oxygen levels also decrease the self-purification capabilities of rivers.

The negative effects of water pollution, including sediments, nitrogen from agriculture, disease pathogens, pesticides, herbicides, salt, and thermal pollution, will be amplified by observed and projected increases in precipitation intensity and longer periods when streamflows are low.¹⁴⁶ The U.S. Environmental Protection Agency expects the number of waterways considered “impaired” by water pollution to increase.¹⁶² Heavy downpours lead to increased sediment in runoff and outbreaks of water-borne diseases.^{163,164} Increases in pollution carried to lakes, estuaries, and the coastal ocean, especially when coupled with increased temperature, can result in blooms of harmful algae and bacteria. However, regions that experience increased streamflow will have the benefit of pollution being more diluted.

Water-quality changes during the last century were probably due to causes other than climate change, primarily changes in pollutants.¹⁴⁹

Changes in groundwater

Many parts of the United States are heavily dependent on groundwater for drinking, residential, and agricultural water supplies.¹⁶⁴ How climate change will affect groundwater is not well known,



L1
L2
L3
L4
L5
L6
L7
L8
L9
L10
L11
L12
L13
L14
L15
L16
L17
L18
L19
L20
L21
L22
L23
L24
L25
L26
L27
L28
L29
L30
L31
L32
L33
L34
L35
L36
L37
L38
L39
L40
L41
L42
L43
L44
L45
L46
L47
L48
L49
L50

R1
R2
R3
R4
R5
R6
R7
R8
R9
R10
R11
R12
R13
R14
R15
R16
R17
R18
R19
R20
R21
R22
R23
R24
R25
R26
R27
R28
R29
R30
R31
R32
R33
R34
R35
R36
R37
R38
R39
R40
R41
R42
R43
R44
R45
R46
R47
R48
R49
R50



Heavy rain can cause sediments to become suspended in water, reducing its quality, as seen in the brown swath above in New York City's Ashokan reservoir following Hurricane Floyd in September 1999.

but increased water demands by society in regions that already rely on groundwater will clearly stress this resource, which is often drawn down faster than it can be recharged.¹⁶⁴ In many locations, groundwater is closely connected to surface water and thus trends in surface-water supplies over time affect groundwater. Changes in the water cycle that reduce precipitation or increase evaporation and runoff would reduce the amount of water available for recharge. Changes in vegetation and soils that occur as temperature changes or due to fire or pest outbreaks are also likely to affect recharge by altering evaporation and infiltration rates. More frequent and larger floods are likely to increase groundwater recharge in semi-arid and arid areas,

where most recharge occurs through dry streambeds after heavy rainfalls and floods.¹⁴²

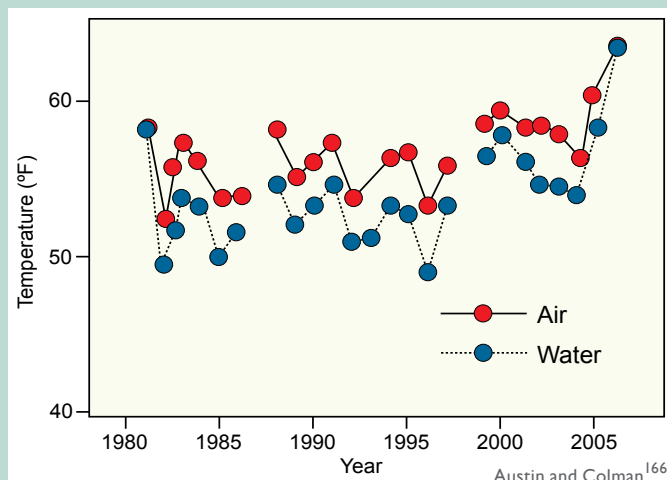
Sea-level rise is expected to increase saltwater intrusion into coastal freshwater aquifers, making them unusable without desalination.¹⁴⁶ Increased evaporation or reduced recharge into coastal aquifers will exacerbate saltwater intrusion. Shallow groundwater aquifers that exchange water with streams are likely to be the most sensitive part of the groundwater system to climate change. Small reductions in groundwater levels can lead to large reductions in streamflow and increases in groundwater levels can increase streamflow.¹⁶⁵ Further, the interface between streams and groundwater is an important site for pollution removal by microorganisms. Their activity will change in response to increased temperature and increased or decreased streamflow as climate changes, and this will affect water quality. Like water quality, research on the impacts of climate change on groundwater has been minimal.¹⁴⁹

Climate change will place additional burdens on already stressed water systems.

In many places, the nation's water systems are already taxed due to aging infrastructure, population increases, and conflicts between water for farming, municipalities, hydropower, recreation, and ecosystems.¹⁶⁷⁻¹⁶⁹ Climate change will add another factor to existing water management challenges, thus increasing vulnerability.¹⁷⁰ The U.S. Bureau of Reclamation has identified many areas in the West that are already at risk for serious conflict over water in the absence of climate change¹⁷¹ (see figure next page).

Adapting to gradual changes, such as changes in average amounts of precipitation, is less difficult than adapting to changes in extremes. Where extreme events, such as droughts or floods, become more intense or more frequent with climate change, the economic and social costs of these events will increase.¹⁷² Water systems have lifetimes of many years and are designed with spare

Lake Superior Summer Air and Water Temperatures 1979 to 2006



The recent large jump in summer water temperature is related to the recent large reduction in ice cover (see Midwest region).

capacity. These systems are thus able to cope with small changes in average conditions.¹⁷² Water resource planning today considers a broad range of stresses and hence adaptation to climate change will be one factor among many in deciding what actions will be taken to minimize vulnerability.¹⁷²⁻¹⁷⁴

Rapid regional population growth

The U.S. population is estimated to have grown to more than 300 million people, nearly a 7 percent increase since the 2000 Census. Current Census Bureau projections are for this growth rate to continue, with the national population projected to reach 350 million by 2025 and 420 million by 2050. The highest rates of population growth to 2025 are projected to occur in areas such as the Southwest that are at risk for reductions in water supplies due to climate change.¹⁶⁷

Aging water infrastructure

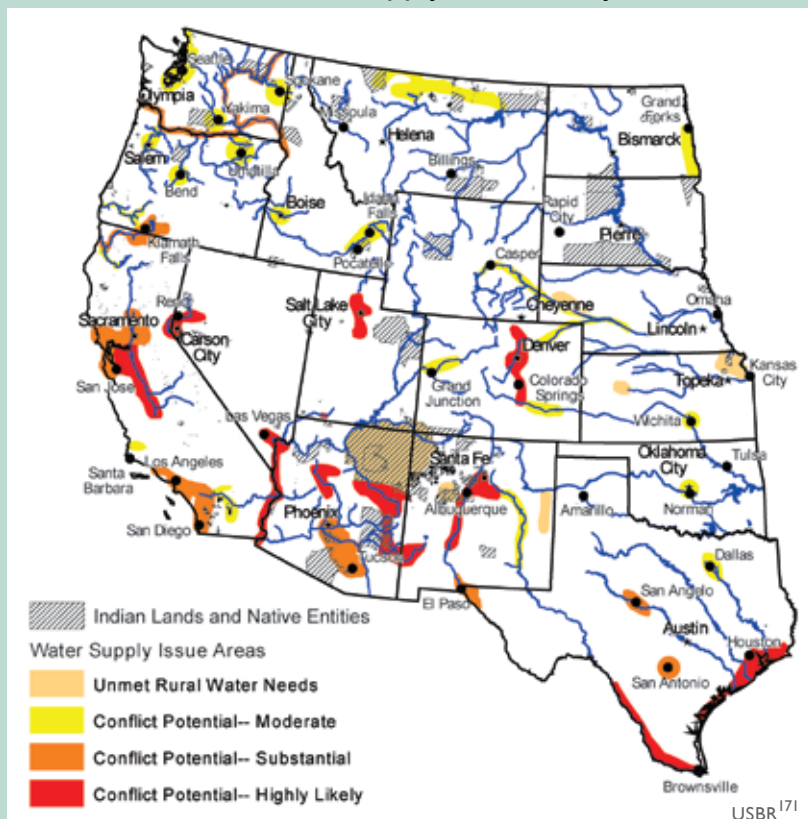
The nation’s drinking water and wastewater infrastructure is aging. In older cities, some buried water mains are over 100 years old and breaks of these lines are a significant problem. Sewer overflows resulting in the discharge of untreated wastewater also occur frequently. The Environmental Protection Agency has identified a large potential funding shortfall for drinking water and wastewater infrastructure.¹⁶⁸ Heavy downpours will exacerbate existing problems in many cities, especially where stormwater catchments and sewers are combined.



Damage to the city water system in Asheville, North Carolina, due to heavy rain in 2004.

Drinking water and sewer infrastructure is very expensive to install and maintain. Climate change will present a new set of challenges for designing upgrades to the nation’s water delivery and sewage removal infrastructure.¹⁶⁸

Potential Water Supply Conflicts by 2025



The map shows regions in the West where water supply conflicts are likely to occur by 2025 based on a combination of factors including population trends and potential endangered species’ needs for water. The red zones are where the conflicts are most likely to occur. This analysis does not factor in the effects of climate change, which is expected to exacerbate many of these already-identified issues.¹⁷¹

Existing water disputes across the country

Many locations in the United States are already undergoing water stress. The Great Lakes states are establishing an interstate compact to protect against reductions in lake levels and potential water exports. Georgia, Alabama, and Florida are in a dispute over water for drinking, recreation, farming, environmental purposes, and hydropower in the Apalachicola–Chattahoochee–Flint River system.^{175,176}

The State Water Project in California is facing a variety of problems in the Sacramento Delta, including endangered species, saltwater intrusion, and potential loss of islands due to flood- or earthquake-caused levee failures.¹⁷⁷⁻¹⁸² A dispute over endangered fish in the Rio Grande has been ongoing for many years.¹⁸³ The Klamath River in Oregon

L1 and California has been the location of a multi-
 L2 year disagreement over native fish, hydropower,
 L3 and farming.^{184,185} The Colorado River has been the
 L4 site of numerous interstate quarrels over the last
 L5 century.^{186,187} Large, unquantified Native American
 L6 water rights challenge existing uses in the West.¹⁸⁸
 L7 By changing the existing patterns of precipitation
 L8 and runoff, climate change will add another stress
 L9 to existing problems.

L11 **Changing water demands**

L12 Water demands are expected to change with in-
 L13 creased temperatures. Evaporation is projected to
 L14 increase over most of the United States as tempera-
 L15 tures rise. Higher temperatures and longer dry peri-
 L16 ods are expected to lead to increased water demand
 L17 for irrigation. This may be partially offset by more
 L18 efficient use of water by plants due to rising atmo-
 L19 spheric carbon dioxide. Higher temperatures are
 L20 projected to increase cooling water withdrawals by
 L21 electrical generating stations. In addition, greater
 L22 cooling requirements in summer will increase elec-
 L23 tricity use, which in turn will require more cooling
 L24 water for power plants. Industrial and municipal
 L25 demands are expected to increase slightly.¹⁴⁶

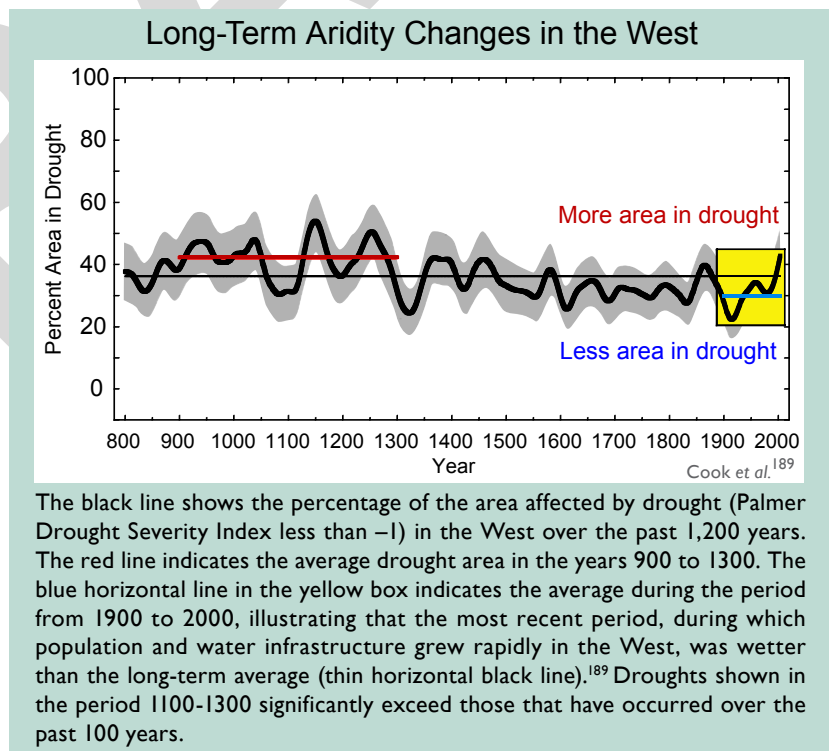
L28 **The past century is no longer a L29 reasonable guide to the future for water L30 management.**

L32 Water planning and management have been
 L33 based on historical fluctuations in records
 L34 of stream flows, lake levels, precipitation,
 L35 temperature, and water demands. All aspects
 L36 of water management including reservoir
 L37 sizing, reservoir flood operations, maximum
 L38 urban stormwater runoff amounts, and
 L39 projected water demands have been based on
 L40 these records. Water managers have proven
 L41 adept at balancing supplies and demand
 L42 through the significant climate variability of
 L43 the past century.¹⁴² Because climate change
 L44 will significantly modify many aspects of the
 L45 water cycle, the assumption of an unchanging
 L46 climate is no longer appropriate for many
 L47 aspects of water planning. Past assumptions
 L48 derived from the historic record about supply
 L49 and demand will need to be revisited for
 L50 existing and proposed water projects.^{142,151,174}

Drought studies that consider the past 1,200 years
 indicate that in the West, the last century was sig-
 nificantly wetter than most other centuries. Multi-
 decade “megadroughts” in the years 900 to 1300
 were substantially worse than the worst droughts of
 the last century, including the Dust Bowl era. The
 causes of these events are only partially known;
 if they were to reoccur, they would clearly stress
 water management even in the absence of climate
 change (see figure).^{97,149,189}

The intersection of substantial changes in the water
 cycle with multiple stresses such as population
 growth and competition for water supplies means
 that water planning will be doubly challenging.
 The ability to modify operational rules and water
 allocations is likely to be critical for the protection
 of infrastructure, for public safety, to ensure reli-
 ability of water delivery, and to protect the environ-
 ment. There are, however, many institutional and
 legal barriers to such changes in both the short and
 long term.¹⁹⁰ Four examples:

- The allocation of the water in many interstate rivers is governed by compacts, international treaties, federal laws, court decrees, and other agreements that are difficult to modify.



- Reservoir operations are governed by “rule curves” that require a certain amount of space to be saved in a reservoir at certain times of year to capture a potential flood. Developed by the Army Corps of Engineers based on historical flood data, many of these rule curves have never been modified, and modifications might require Environmental Impact Statements.¹⁵¹
- In most parts of the West, water is allocated based on a “first in time means first in right” system, and because agriculture was developed before cities were established, large volumes of water typically are allocated to agriculture. Transferring agricultural rights to municipalities, even for short periods during drought, can involve substantial expense and time and can be socially divisive.

Highlights of Water-Related Impacts by Sector

Sector	Examples of Impacts
Human Health	Heavy downpours increase incidence of waterborne disease and floods, resulting in hazards to human life and health. ¹⁶³
Energy Production and Use	Hydropower production is reduced due to low flows in some regions. Power generation is reduced in fossil fuel and nuclear plants due to increased water temperatures and reduced cooling water availability. ¹⁹¹
Transportation	Floods and droughts disrupt transportation. Heavy downpours affect harbor infrastructure and inland waterways. Declining Great Lakes levels reduce freight capacity. ¹⁹²
Agriculture and Forests	Intense precipitation can delay spring planting and damage crops. Earlier spring snowmelt leads to increased number of forest fires. ¹⁹³
Ecosystems	Coldwater fish threatened by rising water temperatures. Some warmwater fish will expand ranges. ⁷⁰

- Conserving water does not necessarily lead to a right to that saved water, thus creating a disincentive for conservation.

Total U.S. water diversions peaked in the 1980s, which implies that expanding supplies in many areas to meet new needs will not be a viable option, especially in arid areas likely to experience less precipitation. However, over the last 30 years, per capita water use has decreased significantly (due, for example, to more efficient technologies such as drip irrigation) and it is anticipated that per capita use will continue to decrease, thus easing stress.¹⁴⁹

Adaptation: New York City Begins Planning for Climate Change

The New York City Department of Environmental Protection (DEP), the agency in charge of providing the city’s drinking water and wastewater treatment, is beginning to alter its planning to take into account the effects of climate change – sea-level rise, higher temperatures, increases in extreme events, and changing precipitation patterns – on the city’s water systems. In partnership with Columbia University, DEP is evaluating climate change projections, impacts, indicators, and adaptation and mitigation strategies.

City planners have begun to address these issues by defining risks using probabilistic climate scenarios and considering potential adaptations that relate to operations/management, infrastructure, and policy. For example, DEP is examining the feasibility of relocating critical control systems to higher floors in low-lying buildings or to higher ground, building flood walls, and modifying design criteria to reflect changing hydrologic processes.

Important near-term goals of the overall effort include updating the existing 100-year flood elevations using climate model projections and identifying additional monitoring stations needed to track changes. DEP will also establish a system for reporting the impacts of extreme weather events on the City’s watershed and infrastructure. In the immediate future, DEP will evaluate flood protection measures for three existing water pollution control plants that are scheduled for renovation.¹⁹⁴

Spotlight on the Colorado River



The Colorado River system supplies water to over 30 million people in the Southwest including Los Angeles, Phoenix, Las Vegas, and Denver. Reservoirs in the system, including the giant lakes Mead and Powell, were nearly full in 1999, with almost four times the annual flow of the river stored. By 2007, the system had lost approximately half of that storage after enduring the worst drought in 100 years of record keeping.²⁹ Runoff was reduced due to low winter precipitation, and warm, dry, and windy springs that substantially reduced snowpack.

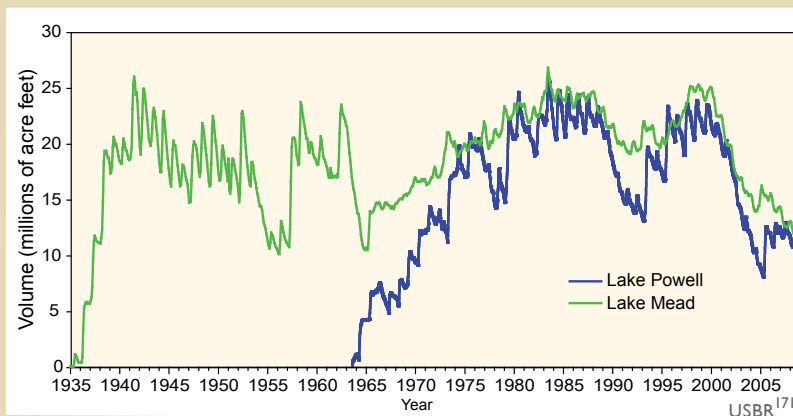


Matching photographs taken 18 months apart during the most serious period of recent drought show a significant decrease in Lake Powell.

Numerous studies over the last 30 years have indicated that the river is likely to experience reductions in runoff due to climate change. In addition, diversions from the river to meet the needs of cities and agriculture are approaching its average flow. Under current conditions, even without climate change, large year-to-year fluctuations in reservoir storage are possible.¹⁵² If reductions in flow projected to accompany global climate change occur, water managers will be challenged to satisfy all existing demands, let alone the increasing demands of a rapidly growing population.^{167,195}

Efforts are underway to address these challenges. In 2005, the Department of Interior's Bureau of Reclamation began a process to formalize operating rules for lakes Mead and Powell during times of low flows and to apportion limited water among the states.¹⁹⁶

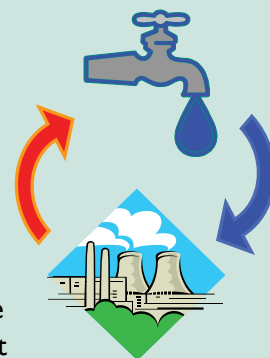
Change in Water Volume of Lakes Mead and Powell



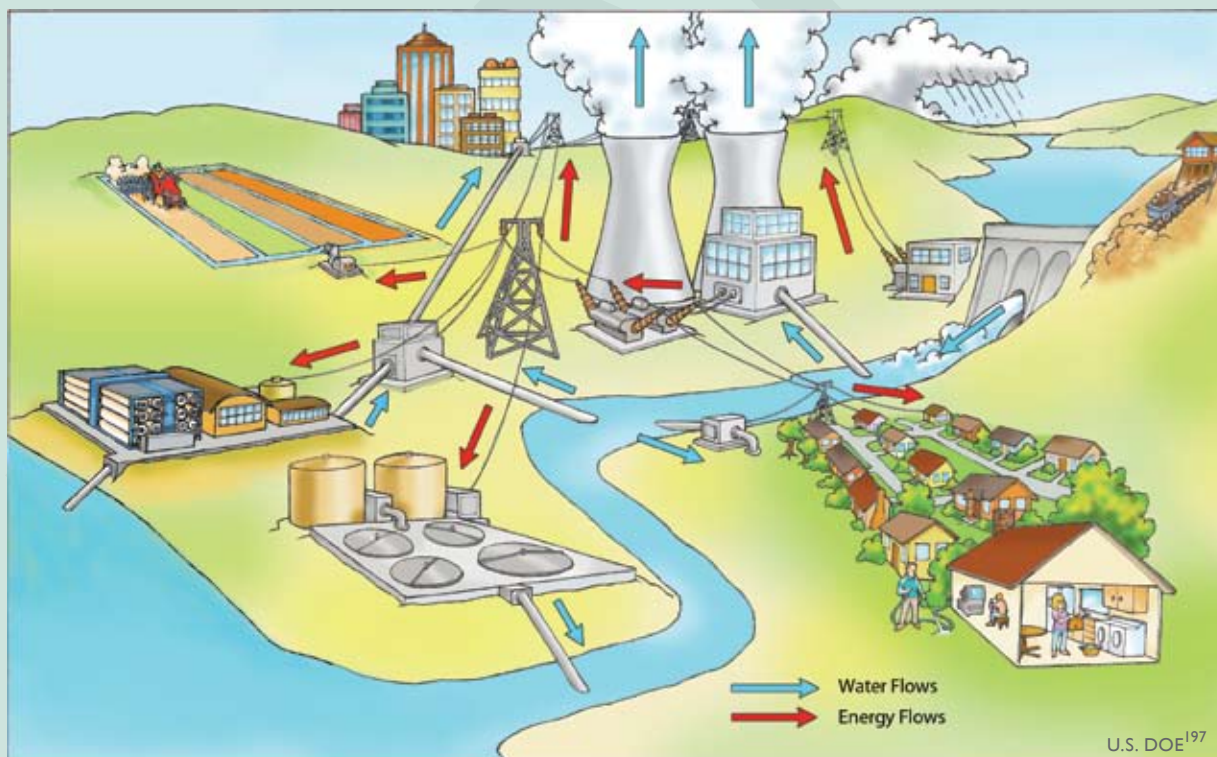
Lake Mead (green) was first filled in 1935, and Lake Powell (blue) was first filled in 1963. In 1999, the lakes were nearly full, but by 2007, the lakes had lost nearly half of their storage water after the worst drought in 100 years.

Water and Energy Connections

Water and energy are tightly interconnected; water systems use large amounts of energy, and energy systems use large amounts of water. Both are expected to be under increasing pressure in the future and both will be affected by a changing climate. In the energy sector, water is used directly for hydropower, and cooling water is critical for nearly all other forms of electrical power generation. Withdrawals of freshwater used to cool power plants that use heat to generate electricity are very large, nearly equaling the water withdrawn for irrigation. Water consumption by power plants is about 20 percent of all non-agricultural uses, or half that of all domestic use.¹⁹⁷



In the water sector, two very unusual attributes of water, significant weight due to its relatively high density, and high heat capacity, make water use energy intensive. Large amounts of energy are needed for pumping, heating, and treating drinking and wastewater. Water supply and treatment consumes roughly 4 percent of the nation's power supply, and electricity accounts for about 75 percent of the cost of municipal water processing and transport. In California, 30 percent of all non-power plant natural gas is used for water-related activities.^{198,199} The energy required to provide water depends on its source (groundwater, surface water, desalinated water, treated wastewater, or recycled water), the distance the water is conveyed, the amount of water moved, and the local topography. Surface water often requires more treatment than groundwater. Desalination requires large amounts of energy to produce freshwater. Treated wastewater and recycled water (used primarily for agriculture and industry) require energy for treatment, but little energy for supply and conveyance. Conserving water has the dual benefit of conserving energy and potentially reducing greenhouse gas emissions if fossil fuels are the predominant source of that energy.



Water and energy are intimately connected. Water is used by the power generation sector for cooling, and energy is used by the water sector for pumping, drinking water treatment, and wastewater treatment. Without energy, there would be limited water distribution, and without water, there would be limited energy production.

Energy Supply and Use

Key Sources



Key Messages:

- Warming will be accompanied by decreases in demand for heating energy and increases in demand for cooling energy. The latter will result in significant increases in electricity use and peak demand in most regions.
- Energy production is likely to be constrained by rising temperatures and limited water supplies in many regions.
- Energy production and delivery systems are exposed to sea-level rise and extreme weather events in vulnerable regions.
- Climate change is likely to affect some renewable energy sources across the nation, such as hydropower production in regions subject to changing patterns of precipitation or snowmelt.

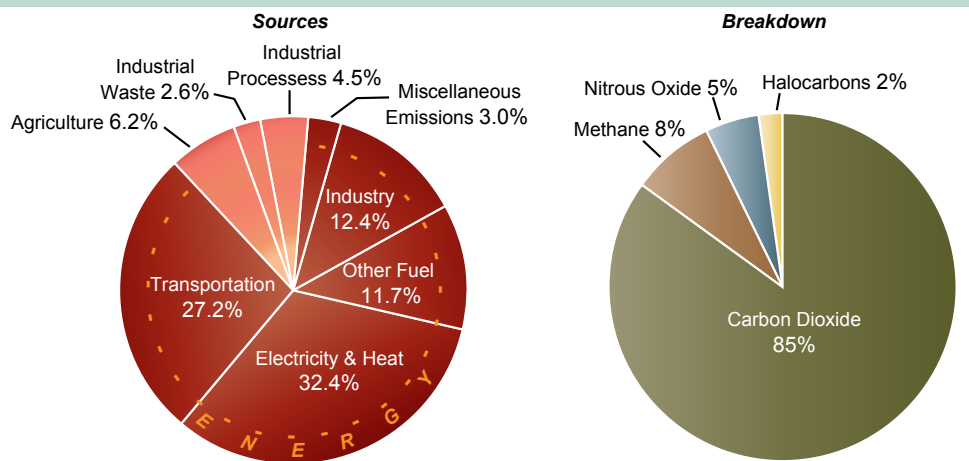
Energy is at the heart of the global warming challenge.³ It is humanity's production and use of energy that is the primary cause of global warming, and in turn, climate change will eventually affect our production and use of energy. The vast majority of U.S. greenhouse gas emissions, about 87 percent, come from energy production and use.²⁰⁰

At the same time, other U.S. trends are increasing energy use: population shifts to the South and Southwest where air conditioning use is high, an increase in the square footage built per person, increased electrification of the residential and commercial sectors, and increased market penetration of air conditioning.²⁰¹

Many of the effects of climate change on energy production and use in the United States are not well studied. Some of the effects of climate change, however, have clear implications for

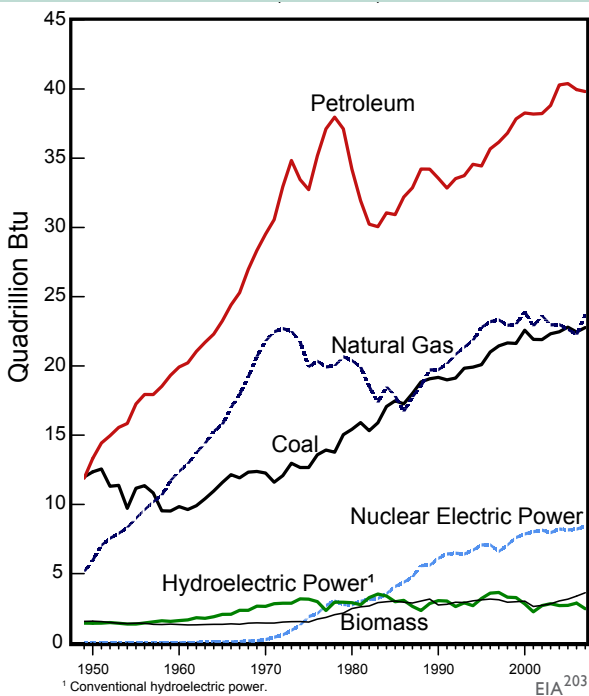
energy production and use. For instance, rising temperatures are expected to increase energy requirements for cooling and reduce energy requirements for heating.^{201,164} Changes in precipitation have the potential to affect prospects for hydropower, positively or negatively.²⁰¹ Increases in hurricane intensity are likely to cause further disruptions to oil and gas operations in the Gulf, like those experienced in 2005 with Hurricane Katrina and in 2008 with Hurricane Ike.²⁰¹ Concerns about climate

Sources of U.S. Greenhouse Emissions



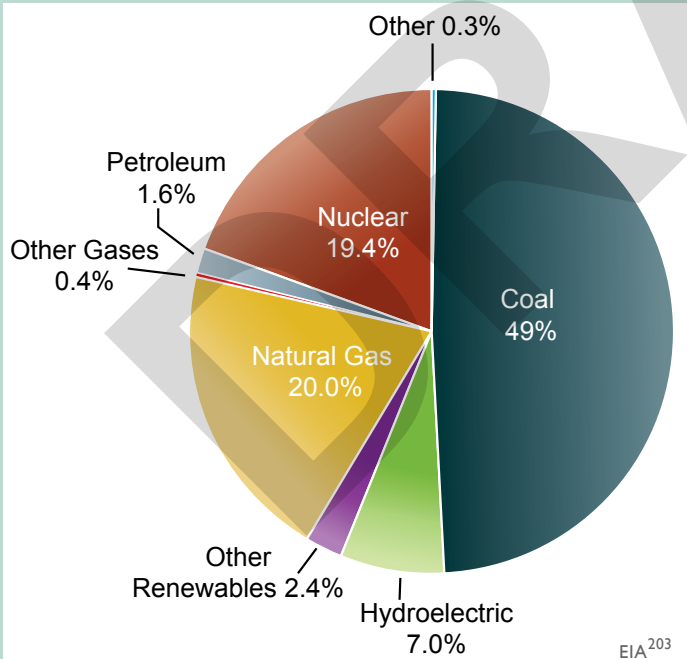
Adapted from U.S. EPA²⁰²
About 87 percent of U.S. greenhouse gas emissions come from energy production and use, as shown in the left pie chart. The right pie chart breaks down these emissions by greenhouse gas.

Primary Energy Consumption by Major Source (1949 to 2007)



U.S. energy supply is dominated by fossil fuels. Petroleum, the top source of energy shown above, is primarily used for transportation (70 percent of oil use). Natural gas is used in roughly equal parts to generate electricity, power industrial processes, and heat water and buildings. Coal is primarily used to generate electricity (91 percent of coal use). Nuclear power is used entirely for electricity generation.

U.S. Electricity Sources



Coal, natural gas, and nuclear power plants together account for 90 percent of current U.S. electricity production.

change impacts will almost certainly alter perceptions and valuations of energy technology alternatives. These effects are very likely to be relevant for energy policies, decisions, and institutions in the United States, affecting courses of action and appropriate strategies for risk management.²⁰¹

The overall scale of the national energy economy is very large, and the energy industry has both the financial and the managerial resources to be adaptive. Impacts due to climate change are likely to be most apparent at sub-national scales, such as regional effects of extreme weather events and reduced water availability, and effects of increased cooling demands on especially vulnerable places and populations.²⁰⁴

Warming will be accompanied by decreases in demand for heating energy and increases in demand for cooling energy. The latter will result in significant increases in electricity use and peak demand in most regions.

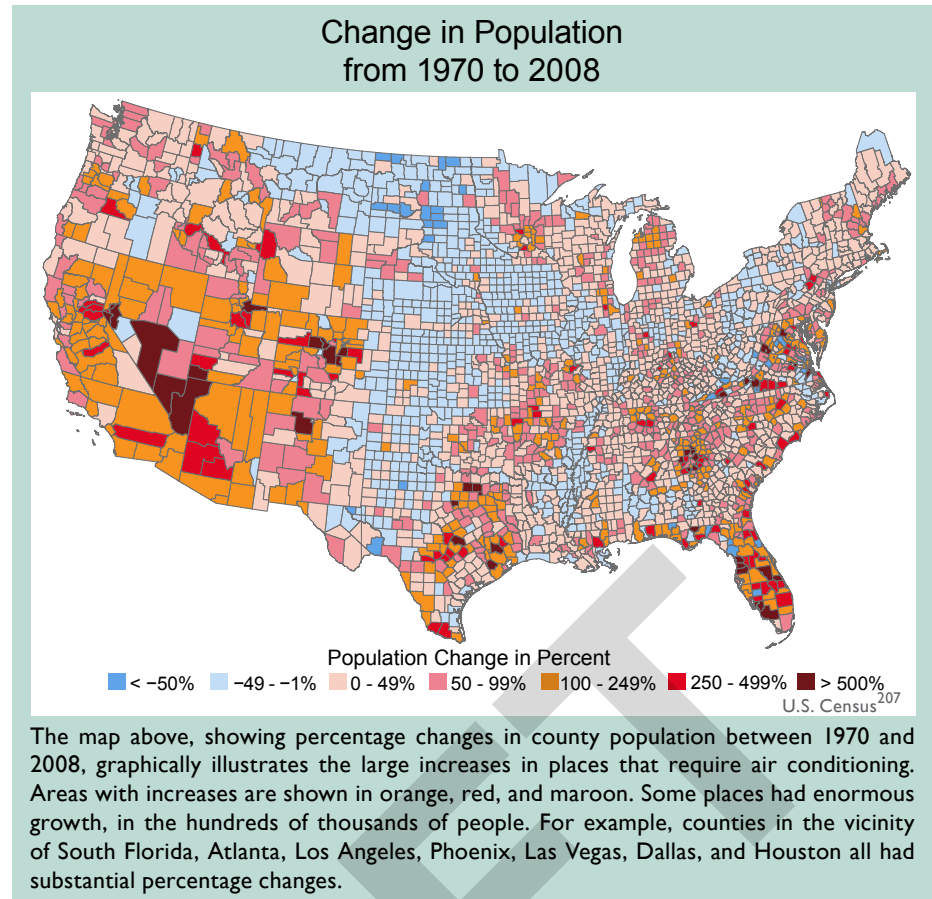
Research on the effects of climate change on energy production and use has largely been limited to impacts on energy use in buildings. These studies have considered effects of global warming on energy requirements for heating and cooling in buildings in the United States.²⁰⁵ They find that the demand for cooling energy increases from 5 to 20 percent per 1.8°F of warming, and the demand for heating energy drops by 3 to 15 percent per 1.8°F of warming.²⁰⁵ These ranges reflect different assumptions about factors such as the rate of market penetration of improved building equipment technologies.²⁰⁵

Studies project that temperature increases due to global warming are very likely to increase peak demand for electricity in most regions of the country.²⁰⁵ An increase in peak demand can lead to a disproportionate increase in energy infrastructure investment.²⁰⁵

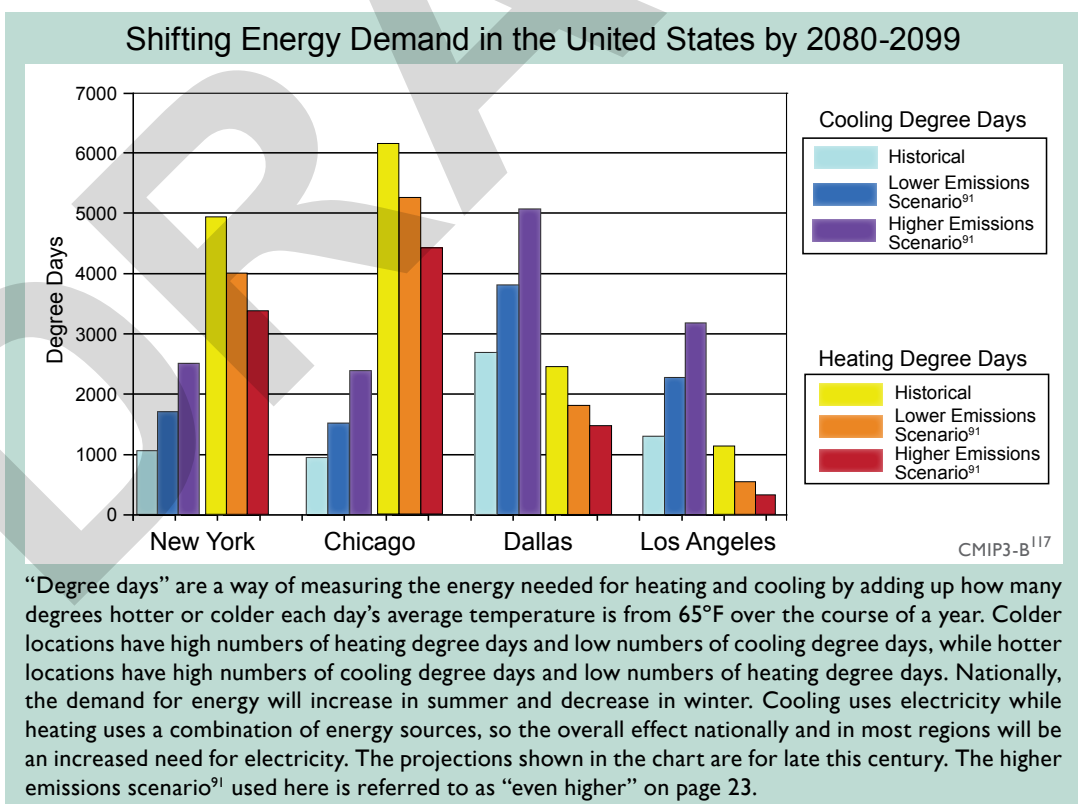
Since nearly all of the cooling of buildings is provided by electricity use, whereas the vast majority of the heating of buildings is provided by natural

R1
R2
R3
R4
R5
R6
R7
R8
R9
R10
R11
R12
R13
R14
R15
R16
R17
R18
R19
R20
R21
R22
R23
R24
R25
R26
R27
R28
R29
R30
R31
R32
R33
R34
R35
R36
R37
R38
R39
R40
R41
R42
R43
R44
R45
R46
R47
R48
R49
R50

L1 gas and fuel oil,^{201,206} the projected
 L2 changes imply increased demands
 L3 for electricity. This is especially the
 L4 case where climate change would
 L5 result in significant increases in the
 L6 heat index in summer, and where
 L7 relatively little space cooling has
 L8 been needed in the past, but demands
 L9 are likely to increase in the future.²⁰⁵
 L10 The increase in energy demand is
 L11 likely to be accelerated by population
 L12 movements to the South and South-
 L13 west, which are regions of especially
 L14 high per capita electricity use, due to
 L15 demands for cooling in commercial
 L16 buildings and households.²⁰⁵ Because
 L17 nearly half of the nation's electric-
 L18 ity is currently generated from coal,
 L19 these factors have the potential to
 L20 increase total national carbon dioxide
 L21 emissions in the absence of improved
 L22 energy efficiency, development of
 L23 non-carbon energy sources, and/or
 L24 carbon capture and storage.²⁰⁵



L25
 L26 Other effects of climate change on energy consumption are less clear, because little research has been
 L27 done.²⁰⁵ For instance, in addition to cooling, air conditioners also remove moisture from the air; thus the
 L28 increase in humidity projected to accompany global warming is likely to increase electricity consumption
 L29 by air conditioners even
 L30 further.²⁰⁵ As other ex-
 L31 amples, warming would
 L32 increase the use of air
 L33 conditioners in highway
 L34 vehicles, and water scar-
 L35 city in some regions has
 L36 the potential to increase
 L37 energy demands for
 L38 water pumping. Improv-
 L39 ing the information
 L40 available about these
 L41 other kinds of effects is
 L42 a priority.



Energy production is likely to be constrained by rising temperatures and limited water supplies in many regions.

In some regions, reductions in water supply due to decreases in precipitation and/or water from melting snowpack are likely to be significant, increasing the competition for water among various sectors including energy production (see *Water Resources* sector).^{191,208}

The production of energy from fossil fuels (coal, oil, and natural gas) is inextricably linked to the availability of adequate and sustainable supplies of water.^{191,208} While providing the United States with the majority of its annual energy needs, fossil fuels also place a high demand on the nation’s water resources in terms of both quantity and quality impacts.^{191,208} Generation of electricity in thermal power plants (coal, nuclear, gas, or oil) is water intensive. Power plants rank only slightly behind irrigation in terms of freshwater withdrawals in the United States.¹⁹¹

There is a high likelihood that water shortages will limit power plant electricity production in many regions, projecting future water constraints on electricity production in thermal power plants for Arizona, Utah, Texas, Louisiana, Georgia, Alabama, Florida, California, Oregon, and Washington state by 2025.¹⁹¹ Additional parts of the United States could face similar constraints as a result of drought, growing populations, and increasing demand for water for various uses, at least seasonally.²⁰⁹ Situations where the development of new power plants is being slowed down or halted due to inadequate cooling water are becoming more frequent throughout the nation.¹⁹¹

The issue of competition among various water uses is dealt with in more detail in the *Water Resources* sector. In connection with these issues and other regional water scarcity impacts, energy is likely to be needed to move and manage water. This is one of many examples of interactions among the impacts of climate change on various sectors that, in this case, affects energy requirements.



Nuclear, coal, and natural gas power plants require large amounts of water for cooling.¹⁹¹

In addition to the problem of water availability, there are issues related to an increase in water temperature. Use of warmer water reduces the efficiency of thermal power plant cooling technologies. And, warmer water discharged from power plants can alter species composition in aquatic ecosystems.²¹⁰ Large coal and nuclear plants have been limited in their operations by reduced river levels caused by higher temperatures and thermal limits on water discharge.¹⁹¹

The efficiency of thermal power plants, fossil or nuclear, is sensitive to ambient air and water temperatures; higher temperatures reduce power outputs by affecting the efficiency of cooling.¹⁹¹ Although this effect is not large in percentage terms, even a relatively small change could have significant implications for total national electric power supply.¹⁹¹ For example, an average reduction of 1 percent in electricity generated by thermal power plants nationwide would mean a loss of 25 billion kilowatt-hours per year,²¹¹ about the amount of electricity consumed by 2 million Americans, a loss that would need to be supplied in some other way or offset through measures that improve energy efficiency.

Energy production and delivery systems are exposed to sea-level rise and extreme weather events in vulnerable regions.

Sea-level rise

A significant fraction of America's energy infrastructure is located near the coasts, from power plants, to oil refineries, to facilities that receive oil and gas deliveries.¹⁹¹ Rising sea levels are likely to lead to direct losses, such as equipment damage from flooding or erosion, and indirect effects, such as the costs of raising vulnerable assets to higher levels or building new facilities farther inland, increasing transportation costs.¹⁹¹ The U.S. East Coast and Gulf Coast have been identified as particularly vulnerable to sea-level rise because the land is relatively flat and also sinking in many places.¹⁹¹

Extreme events

Observed and projected increases in a variety of extreme events will have significant impacts on the energy sector. As witnessed in 2005, hurricanes can have a debilitating impact on energy infrastructure. Direct losses to the energy industry in 2005 are estimated at \$15 billion,¹⁹¹ with millions more in restoration and recovery costs. As one example, the Yscloskey Gas Processing Plant (located on

the Louisiana coast) was forced to close for six months following Hurricane Katrina, resulting in lost revenues to the plant's owners and employees, and higher prices to consumers, as gas had to be procured from alternative sources.¹⁹¹

The impacts of an increase in severe weather are not limited to hurricane-prone areas. For example, rail transportation lines, which carry approximately two-thirds of the coal to the nation's power plants,²¹² often follow riverbeds, especially in the Appalachian region.¹⁹¹ More intense rainstorms, which have been observed and projected,^{68,112} can lead to rivers flooding, which can "wash out" or degrade nearby railbeds and roadbeds.¹⁹¹ This is also a problem in the Midwest, which experienced major flooding of the Mississippi River in 1993 and 2008.²¹³

Development of new energy facilities could be restricted by siting concerns related to sea-level rise, exposure to extreme events, and increased capital costs resulting from a need to provide greater protection from extreme events.¹⁹¹

The electricity grid is also vulnerable to climate change effects, from temperature changes to severe weather events.¹⁹¹ The most familiar example is ef-

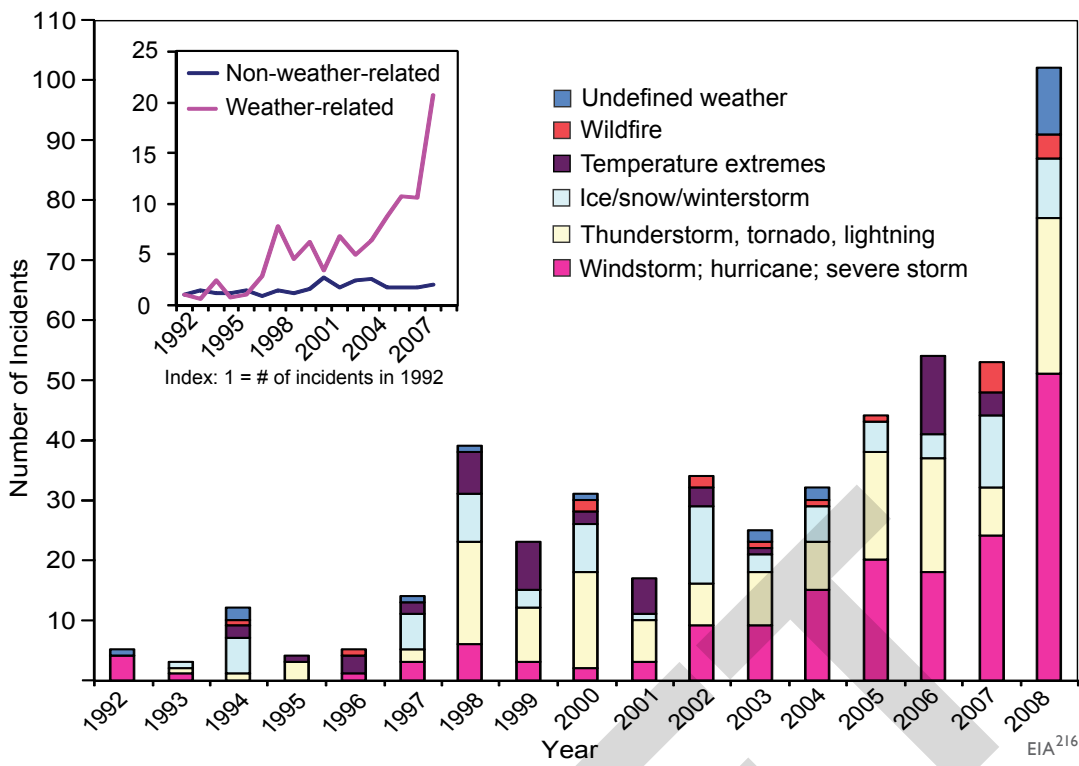
Regional Spotlight: Gulf Coast Oil and Gas



The Gulf Coast is home to the U.S. oil and gas industries, representing nearly 30 percent of the nation's crude oil production and approximately 20 percent of its natural gas production. One-third of the national refining and processing capacity lies on coastal plains adjacent to the Gulf of Mexico. Several thousand offshore drilling platforms, dozens of refineries, and thousands of miles of pipelines are vulnerable to damage and disruption due to sea-level rise and the high winds and storm surge associated with hurricanes and other tropical storms. For example, hurricanes Katrina and Rita halted all oil and gas production from the Gulf, disrupted nearly 20 percent of the nation's refinery capacity, and closed many oil and gas pipelines.²¹⁴ Relative sea-level rise in parts of the Gulf Coast region (Louisiana and East Texas) is projected to be as high as 2 to 4 feet by 2050 to 2100, due to the combination of global sea-level rise caused by warming oceans and melting ice and local land sinking.²¹⁵ Combined with onshore and offshore storm activity, this would represent an increased threat to this regional energy infrastructure. Some adaptations to these risks are beginning to emerge (see Adaptation box, page 58).

Offshore oil production is particularly susceptible to extreme weather events. Hurricane Ivan in 2004 destroyed seven platforms in the Gulf of Mexico, significantly damaged 24 platforms, and damaged 102 pipelines. Hurricanes Katrina and Rita in 2005 destroyed more than 100 platforms and damaged 558 pipelines. For example, Chevron's \$250 million "Typhoon" platform was damaged beyond repair. Plans are being made to sink its remains to the seafloor.

Significant Weather-Related U.S. Electric Grid Disturbances



The number of incidents caused by extreme weather has increased tenfold since 1992. The portion of all events that are caused by weather-related phenomena has more than tripled from about 20 percent in the early 1990s to about 65 percent in recent years. The weather-related events are more severe, with an average of about 180,000 customers affected per event compared to about 100,000 for non-weather-related events (and 50,000 excluding the massive blackout of August 2003).²⁰¹ The data shown include disturbances that occurred on the nation's large-scale "bulk" electric transmission systems. Most outages occur in local distribution networks and are not included in the graph. Although the figure does not demonstrate a cause-effect relationship between climate change and grid disruption, it does suggest that weather and climate extremes often have important effects on grid disruptions. We do know that more frequent weather and climate extremes are likely in the future,⁶⁸ which poses unknown new risks for the electric grid.

Adaptation: Addressing Oil Infrastructure Vulnerabilities in the Gulf Coast

Port Fourchon, Louisiana, supports 75 percent of deepwater oil and gas production in the Gulf of Mexico, and its role in supporting oil production in the region is increasing. The Louisiana Offshore Oil Port, located about 20 miles offshore, links daily imports of 1 million barrels of oil and production of 300,000 barrels in the Gulf of Mexico to 50 percent of national refining capacity. One road, Louisiana Highway 1, connects Port Fourchon with the nation. It transports machinery, supplies, and workers and is the evacuation route for onshore and offshore workers. Responding to threats of storm surge and flooding, related in part to concerns about climate change, Louisiana is currently upgrading Highway 1, including elevating it above the 500-year flood level and building a higher bridge over Bayou LaFourche and the Boudreaux Canal.²¹⁷



Regional Spotlight: Florida's Energy Infrastructure



Florida's energy infrastructure is particularly vulnerable to sea-level rise and storm impacts. Most of the petroleum products consumed in Florida are delivered by barge to three ports, two on the east coast and one on the west coast. The interdependencies of natural gas distribution, transportation fuel distribution and delivery, and electrical generation and distribution were found to be major issues in Florida's recovery from recent major hurricanes.¹⁹¹



ffects of severe weather events on power lines, such as from ice storms, thunderstorms, and hurricanes. In the summer heat wave of 2006, for example, electric power transformers failed in several areas (including St. Louis, Missouri, and Queens, New York) due to high temperatures, causing interruptions of electric power supply. It is not yet possible to project effects of climate change on the grid, because so many of the effects would be more localized than current climate change models can depict; but, weather-related grid disturbances are recognized as a challenge for strategic planning and risk management.

Climate change is likely to affect some renewable energy sources across the nation, such as hydropower production in regions subject to changing patterns of precipitation or snowmelt.

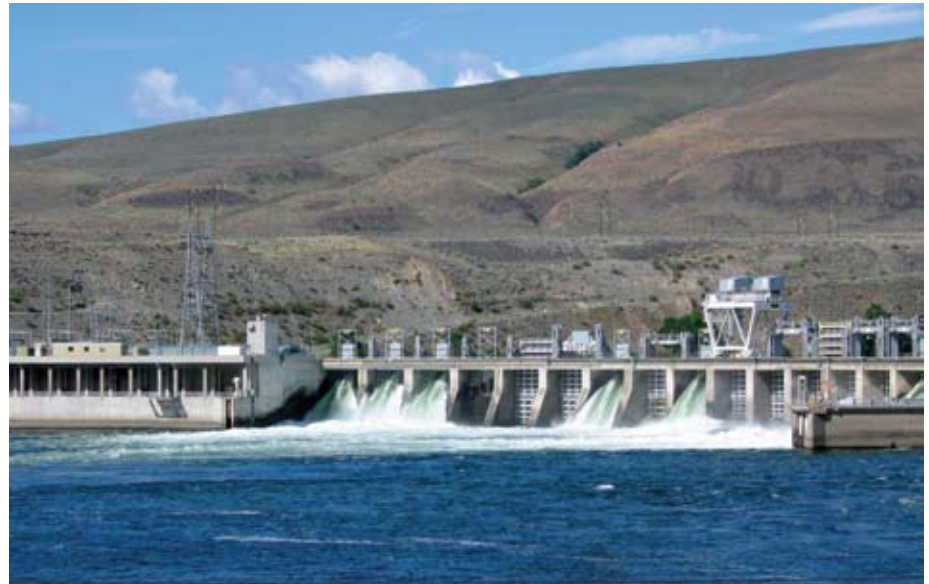
Renewable sources currently account for about 9 percent of electricity production in the United States.²⁰³ Hydroelectric power is by far the largest renewable contributor to electricity generation,¹⁹¹ accounting for about 7 percent of total U.S. electricity.²¹⁸ Like many things discussed in this report, renewable energy resources have strong interrelationships with climate change; using renewable energy can reduce the magnitude of climate change, while climate change can affect the prospects for using some renewable energy sources.

Hydropower is a major source of electricity in some regions of the United States, notably in the

Northwest.¹⁹¹ It is likely to be significantly affected by climate change in regions subject to reduced precipitation and/or water from melting snowpack. Significant changes are already being detected in the timing and amount of streamflows in many western rivers,¹⁶⁴ consistent with the predicted effects of global warming. More precipitation coming as rain rather than snow, reduced snowpack, earlier peak runoff, and related effects are beginning to affect hydropower availability.¹⁶⁴ Hydroelectric generation is very sensitive to changes in precipitation and river discharge. For example, every 1 percent decrease in precipitation results in a 2 to 3 percent drop in streamflow;²¹⁹ every 1 percent decrease in streamflow in the Colorado River Basin results in a 3 percent drop in power generation.¹⁹¹ Such magnifying sensitivities occur because water flows through multiple power plants in a river basin.¹⁹¹ Climate impacts on hydropower occur when either the total amount or the timing of runoff is altered, such as when natural water storage in snowpack and glaciers is reduced under hotter conditions. Glaciers, snowpack, and their associated runoff are already declining in the West, and larger declines are projected.¹⁶⁴

Hydropower operations are also affected by changes to air temperatures, humidity, or wind patterns due to climate change.¹⁹¹ These variables cause changes in water quantity and quality, including water temperature. Warmer air and water generally increases the evaporation of water from the surface of reservoirs, reducing the amount of water available for power production and other uses. Huge reservoirs with large surface areas, located in arid,

L1 sunny parts of the country,
 L2 such as Lake Mead (located
 L3 on Arizona-Nevada border
 L4 on the Colorado River),
 L5 are particularly susceptible
 L6 to increased evaporation
 L7 due to warming, meaning
 L8 less water will be avail-
 L9 able for all uses, including
 L10 hydropower.¹⁹¹ And, where
 L11 hydropower dams flow into
 L12 waterways that support
 L13 trout, salmon or other cold-
 L14 water fisheries, warming
 L15 of reservoir releases might
 L16 have detrimental conse-
 L17 quences that require changes
 L18 in operations that reduce power production.¹⁹¹ Such
 L19 impacts will increasingly translate into competition
 L20 for water resources.



Hydroelectric dam in the Northwest

R1
R2
R3
R4
R5
R6
R7
R8
R9
R10
R11
R12
R13
R14
R15
R16
R17
R18
R19
R20
R21
R22
R23
R24
R25
R26
R27
R28
R29
R30
R31
R32
R33
R34
R35
R36
R37
R38
R39
R40
R41
R42
R43
R44
R45
R46
R47
R48
R49
R50

fuels). The limited research to date on these impor-
 tant issues does not support firm conclusions about
 where such impacts would occur and how signifi-
 cant they would be.²⁰⁵ This is an area that calls
 for much more study (see *An Agenda for Climate
 Impacts Science* section, Recommendation 2).

L22 Climate change will affect other renewable en-
 L23 ergy sources as well, including potential effects of
 L24 changing cloud cover on solar energy resources,
 L25 effects of climate on winds, and effects of tempera-
 L26 ture and water availability on biomass production
 L27 (particularly related to water requirements for bio-

**Regional Spotlight: Energy
 Impacts of Alaska's Rapid
 Warming**



Significant impacts of warming on the energy sector can
 already be observed in Alaska, where temperatures have risen
 about twice as much as the rest of the nation. In Alaska, frozen
 ground and ice roads are an important means of winter travel,
 and warming has resulted in a much shorter cold season. Impacts
 on the oil and natural gas industries on Alaska's North Slope have
 been one of the results. For example, the season during which oil
 and gas exploration and extraction equipment can be operated on the
 tundra has been shortened due to warming. In addition, the thawing of
 permafrost, on which buildings, pipelines, airfields, and coastal installations
 supporting oil and gas development are located, adversely affects these structures
 and increases the cost of maintaining them.¹⁹¹

Different energy impacts are expected in the marine environment as
 sea ice continues to retreat and thin. These trends are expected to
 improve shipping accessibility, including oil and gas transport by sea,
 around the margins of the Arctic Basin – at least in the summer. The
 improved accessibility, however, will not be uniform throughout the
 different regions. Offshore oil exploration and extraction might benefit
 from less extensive and thinner sea ice, although equipment will have to
 be designed to withstand increased wave forces and ice movement.^{191,220}





Transportation

Key Messages:

- Sea-level rise and storm surge will increase the risk of major coastal impacts, including both temporary and permanent flooding of airports, roads, rail lines, and tunnels.
- Flooding from increasingly intense downpours will increase the risk of disruptions and delays in air, rail, and road transportation, and damage from mudslides in some areas.
- The increase in extreme heat will limit some transportation operations and cause pavement and track damage. Decreased extreme cold will provide some benefits.
- Increased intensity of strong hurricanes would lead to more evacuations, infrastructure damage and failure, and transportation interruptions.
- Arctic warming will continue to reduce sea ice, lengthening the ocean transport season, but also resulting in greater coastal erosion due to waves. Permafrost thaw in Alaska will damage infrastructure. The ice-road season will become shorter.

R1
R2
R3
R4
R5
R6
R7
R8
R9
R10
R11
R12
R13
R14
R15
R16
R17
R18
R19
R20
R21



R22
R23
R24
R25
R26
R27
R28
R29
R30
R31
R32
R33
R34
R35
R36
R37
R38
R39
R40
R41
R42
R43
R44
R45
R46
R47
R48
R49
R50

The U.S. transport sector is a significant source of greenhouse gases, accounting for 27 percent of U.S. emissions.²²¹ While it is widely recognized that emissions from transportation have a major impact on climate, climate change will also have a major impact on transportation.

Climate change impacts pose significant challenges to our nation’s multi-modal transportation system and cause disruptions in other sectors across the economy. For example, major flooding in the Midwest in 1993 and 2008 restricted regional travel of all types, and disrupted freight and rail shipments across the country, such as those bringing coal to power plants and chlorine to water treatment systems. The U.S. transportation network is vital to the nation’s economy, safety, and quality of life.

Extreme events present major challenges for transportation, and such events are becoming more frequent and intense. Historical weather patterns are no longer a reliable predictor of the future.²²² Transportation planners have not typically accounted for climate change in their long-term planning and project development. The longevity of transportation infrastructure, the long-term nature of climate change, and the potential impacts identified by recent studies warrant serious attention to climate change in planning new or rehabilitated transportation systems.²²³

The strategic examination of national, regional, state, and local networks is an important step toward understanding the risks posed by climate change. A range of adaptation responses can be employed to reduce risks through redesign or relocation of infrastructure, increased redundancy of critical services, and operational improvements. Adapting to climate change is an evolutionary process. Through adoption of longer planning horizons, risk management, and adaptive responses, vulnerable transportation infrastructure can be made more resilient.²¹⁵



Buildings and debris float up against a railroad bridge on the Cedar River during record flooding in June 2008, in Cedar Rapids, Iowa.

Sea-level rise and storm surge will increase the risk of major coastal impacts, including both temporary and permanent flooding of airports, roads, rail lines, and tunnels.

Sea-level rise

Transportation infrastructure in U.S. coastal areas is increasingly vulnerable to sea-level rise. Given the high population density near the coasts, the potential exposure of transportation infrastructure to flooding is immense. Population swells in these areas during the summer months because beaches are very important tourist destinations.²²²

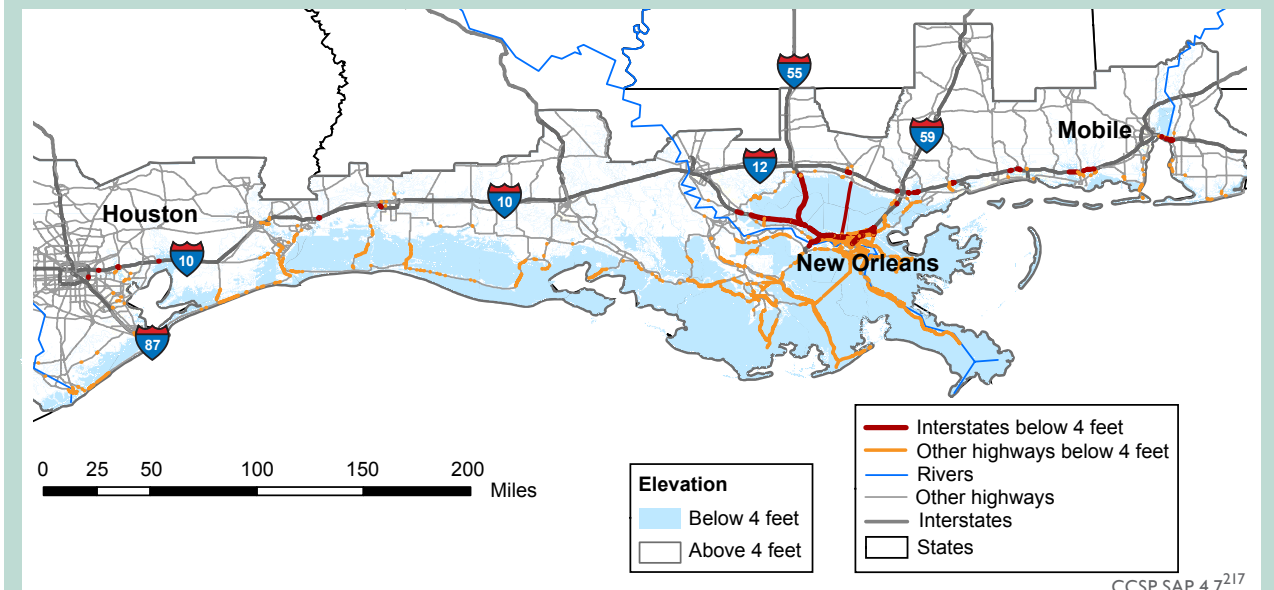
In the Gulf Coast area alone, an estimated 2,400 miles of major roadway and 246 miles of freight rail lines are at risk of permanent flooding within 50 to 100 years as global warming and land subsidence (sinking) combine to produce an anticipated relative sea-level rise in the range of 4 feet.²¹⁷ Since the Gulf Coast region's transportation network is interdependent and relies on minor roads and other low-lying infrastructure, the risks of service disruptions due to sea-level rise are likely to be even greater.²¹⁷

Coastal areas are also major centers of economic activity. Six of the nation's top 10 freight gateways (measured by the value of shipments) will be threatened by sea-level rise.²²² Seven of the 10 largest ports (by tons of traffic) are located on the Gulf Coast.²²² The region is also home to the U.S. oil and gas industry, with its offshore drilling platforms, refineries, and pipelines. Roughly two-thirds of all U.S. oil imports are transported through this region²²⁴ (see *Energy* sector). Sea-level rise would potentially affect commercial transportation activity valued in the hundreds of billions of dollars annually through inundation of area roads, railroads, airports, seaports, and pipelines.²¹⁷

Storm surge

More intense storms, especially when coupled with sea-level rise, will result in far-reaching and damaging storm surge. An estimated 60,000 miles of coastal highway are already exposed to periodic flooding from coastal storms and high waves.²²² Some of these highways currently serve as evacuation routes during hurricanes and other coastal storms, and these routes could become seriously compromised in the future.

Gulf Coast Area Roads at Risk from Sea-Level Rise



Within 50 to 100 years, 2,400 miles of major roadway are projected to be inundated by sea-level rise in the Gulf Coast region. The map shows roadways at risk in the event of a sea-level rise of about 4 feet, within the range of projections for this region in this century under medium- and high-emissions scenarios.⁹¹ In total, 24 percent of interstate highway miles and 28 percent of secondary road miles in the Gulf Coast region are at elevations below 4 feet.²¹⁷

Regional Spotlight: Gulf Coast



Sea-level rise, combined with high rates of subsidence in some areas, will make much of the existing infrastructure more prone to frequent or permanent inundation; 27

percent of the major roads, 9 percent of the rail lines, and 72 percent of the ports in the area shown on the map on the previous page are built on land at or below 4 feet in elevation, a level within the range of projections for relative sea-level rise in this region in this century. Increased storm intensity might lead to increased service disruption and infrastructure damage. More than half of the area’s major highways (64 percent of interstates, 57 percent of arterials), almost half of the rail miles, 29 airports, and virtually all of the ports, are below 23 feet in elevation and subject to flooding and damage due to hurricane storm surge. These factors merit consideration in today’s transportation decisions and planning processes.²¹⁷

Coastal areas are projected to experience continued development pressures as both retirement and tourist destinations. Many of the most populous counties of the Gulf Coast, which already experience the effects of tropical storms, are expected to grow rapidly in the coming decades.²²² This growth will generate demand for more transportation infrastructure and services, challenging transportation planners to meet the demand, address current and future flooding, and plan for future conditions.²²³

Land

More frequent inundation and interruptions in travel on coastal and low-lying roadways and rail lines due to storm surge are projected, potentially requiring changes to minimize disruptions. More frequent evacuations due to severe storm surges are also likely. Across the United States, many coastal cities have subways, tunnels, parking lots, and other transportation infrastructure below

ground. Underground tunnels and other low-lying infrastructure will experience more frequent and severe flooding. Higher sea levels and storm surges will also erode road base and undermine bridge supports. The loss of coastal wetlands and barrier islands will lead to further coastal erosion due to the loss of natural protection from wave action.

Water

Impacts on harbor infrastructure from wave damage and storm surges are projected to increase. Changes will be required in harbor and port facilities to accommodate higher tides and storm surges. There will be reduced clearance under some waterway bridges for boat traffic. Changes in the navigability of channels are expected; some will become more accessible (and extend farther inland) because of deeper waters, while others will be restricted because of changes in sedimentation rates and sandbar locations. In some areas, waterway systems will become part of open water as barrier islands disappear. Some channels are likely to have to be dredged more frequently as has been done across large open-water bodies in Texas.²²²

Regional Spotlight: New York Metropolitan Area



With the potential for significant sea-level rise estimated under business-as-usual emissions, the combined effects of sea-level rise and storm surge are projected to dramatically increase the frequency of flooding. What is currently called a 100-year storm is projected to occur as often as every 5 years. Portions of lower Manhattan and coastal areas of Brooklyn, Queens, Staten Island, and Nassau County, would experience a marked increase in flooding frequency. Much of the critical transportation infrastructure, including tunnels, subways, and airports, lies well within the range of projected storm surge and would be flooded during such events.^{222,225}

L1
L2
L3
L4
L5
L6
L7
L8
L9
L10
L11
L12
L13
L14
L15
L16
L17
L18
L19
L20
L21
L22
L23
L24
L25
L26
L27
L28
L29
L30
L31
L32
L33
L34
L35
L36
L37
L38
L39
L40
L41
L42
L43
L44
L45
L46
L47
L48
L49
L50

R1
R2
R3
R4
R5
R6
R7
R8
R9
R10
R11
R12
R13
R14
R15
R16
R17
R18
R19
R20
R21
R22
R23
R24
R25
R26
R27
R28
R29
R30
R31
R32
R33
R34
R35
R36
R37
R38
R39
R40
R41
R42
R43
R44
R45
R46
R47
R48
R49
R50

L1 **Air**
 L2 Airports in coastal cities are often located adjacent
 L3 to rivers, estuaries, or open ocean. Airport runways
 L4 in coastal areas face inundation unless effective
 L5 protective measures are taken. There is the po-
 L6 tential for closure or restrictions for several of the
 L7 nation’s busiest airports that lie in coastal zones,
 L8 affecting service to the highest density populations
 L9 in the United States.

L12 **Flooding from increasingly intense
 L13 downpours will increase the risk of
 L14 disruptions and delays in air, rail, and
 L15 road transportation, and damage from
 L16 mudslides in some areas.**

L18 Heavy downpours have already increased substan-
 L19 tially in the United States; the heaviest 1 percent
 L20 of precipitation events increased by 20 percent,
 L21 while total precipitation increased by only 7 percent
 L22 over the past century.¹¹² Such intense precipitation
 L23 is likely to increase the frequency and severity
 L24 of events such as the Great Flood of 1993, which
 L25 caused catastrophic flooding along 500 miles of
 L26 the Mississippi and Missouri river system, paralyz-
 L27 ing surface transportation systems, including rail,
 L28 truck, and marine traffic. Major east-west traffic
 L29 was halted for roughly six weeks in an area stretch-
 L30 ing from St. Louis, Missouri, west to Kansas City,
 L31 Missouri and north to Chicago, Illinois, affecting
 L32 one-quarter of all U.S. freight that either originated
 L33 or terminated in the flood-affected region.²²²

L35 The June 2008 Midwest flood was the second
 L36 record-breaking flood in the past 15 years. Dozens
 L37 of levees were breached or overtopped in Iowa,
 L38 Illinois, and Missouri, flooding huge areas, includ-
 L39 ing nine square miles in and around Cedar Rapids,
 L40 Iowa. Numerous highway and rail bridges were
 L41 impassable due to flooding of approaches and
 L42 transport was shut down along many stretches of
 L43 highway, rail lines, and normally navigable water-
 L44 ways.

L46 Planners have generally relied on weather extremes
 L47 of the past as a guide to the future, planning, for
 L48 example, for a “100-year flood,” which is now
 L49 likely to come more frequently as a result of
 L50 climate change. Historical analysis of weather data

R1 has thus become less reliable as a forecasting tool.
 R2 The accelerating changes in climate make it more
 R3 difficult to predict the frequency and intensity of
 R4 weather events that can affect transportation.²²²

R7 **Land**

R7 The increase in heavy precipitation will inevita-
 R8 bly cause increases in weather-related accidents,
 R9 delays, and traffic disruptions in a network already
 R10 challenged by increasing congestion.²¹⁵ There will
 R11 be increased flooding of evacuation routes, and
 R12 construction activities will be disrupted. Changes
 R13 in rain, snowfall, and seasonal flooding will impact
 R14 safety and maintenance operations on the nation’s
 R15 roads and railways. For example, if more precipita-
 R16 tion falls as rain rather than snow in winter and
 R17 spring, there will be an increased risk of landslides,
 R18 slope failures, and floods from the runoff, causing
 R19 road closures as well as the need for road repair and
 R20 reconstruction²²² (see *Water Resources* sector).

R22 Increased flooding of roadways, rail lines and
 R23 underground tunnels is expected. Drainage systems
 R24 will be overloaded more frequently and severely,
 R25 causing backups and street flooding. Areas where
 R26 flooding is already common will face more fre-
 R27 quent and severe problems. For example, Louisiana
 R28 Highway 1, a critical link in the transport of oil
 R29 from the Gulf of Mexico, has recently experienced
 R30 increased flooding, prompting authorities to elevate
 R31 the road (see Adaptation Box page 58).²¹⁷ Increases
 R32 in road washouts, damage to railbed support struc-
 R33 tures, and landslides and mudslides that damage
 R34 roads and other infrastructure are expected. If soil
 R35 moisture levels become too high, the structural
 R36 integrity of roads, bridges, and tunnels, which in
 R37 some cases are already under age-related stress and
 R38 in need of repair, could be compromised. Stand-
 R39 ing water will have adverse impacts on road base.
 R40 For example, damage due to long term submersion
 R41 of roadways in Louisiana was estimated to be \$50
 R42 million for just 200 miles of state-owned highway.
 R43 The Louisiana Department of Transportation and
 R44 Development noted that a total of 1,800 miles of
 R45 roads were under water for long periods, requiring
 R46 costly repairs.²¹⁷ Pipelines are likely to be damaged
 R47 because intense precipitation can cause the ground
 R48 to sink underneath the pipeline; in shallow river-
 R49 beds, pipelines are more exposed to the elements
 R50



Adaptation: Climate Proofing a Road

Completion of a road around the 42-square mile island of Kosrae in the U.S.-affiliated Federated States of Micronesia provides a good example of adaptation to climate change. A road around the island’s perimeter existed, except for a 10-mile gap. Filling this gap would provide all-weather land access to a remote village and allow easier access to the island’s interior.



In planning this new section of road, authorities decided to “climate-proof” it against projected increases in heavy downpours and sea-level rise. This led to the section of road being placed higher above sea level and with an improved drainage system to handle the projected heavier rainfall. While there were additional capital costs for incorporating this drainage system, the accumulated costs, including repairs and maintenance, would be lower after about 15 years, equating to a good rate of return on investment. Adding this improved drainage system to roads that are already built is more expensive than on new construction, but still has been found to be cost effective.²²⁶

and can be subject to scouring and shifting due to heavy precipitation.²¹⁷

Water

Facilities on land at ports and harbors will be vulnerable to short term flooding from heavy downpours, interrupting shipping service. Changes in silt and debris buildup resulting from extreme precipitation events will affect channel depth, increasing dredging costs. The need to expand stormwater treatment facilities, which can be a significant expense for container and other terminals with large impermeable surfaces, will increase.

Air

Increased delays due to heavy downpours are likely to affect operations, causing increasing flight delays and cancellations.²²² Stormwater runoff that exceeds the capacity of collection and drainage systems will cause flooding, delays, and airport closings. Heavy downpours will affect the structural integrity of airport facilities, such as through flood damage to runways and other infrastructure. All of these impacts have implications for emergency evacuation planning, facility maintenance, and safety.²²²

The increase in extreme heat will limit some transportation operations and cause pavement and track damage. Decreased extreme cold will provide some benefits.

Land

Longer periods of extreme heat in summer might damage roads in several ways, including softening of asphalt that leads to rutting from heavy traffic.¹⁶⁴ Sustained air temperature over 90°F is a significant threshold for such problems. Extreme heat can cause deformities in rail tracks, at minimum resulting in speed restrictions and, at worst, causing derailments. Air temperatures above 100°F can lead to equipment failure. Extreme heat also causes thermal expansion of bridge joints, adversely affecting bridge operations and increasing maintenance costs. Vehicle overheating and tire deterioration are additional concerns.²²² Higher temperatures also will increase refrigeration needs for goods during transport, particularly in the South, raising transportation costs.²¹⁷

Increases in very hot days and heat waves are expected to limit construction activities due to health and safety concerns for highway workers. Guidance from the U.S. Occupational Safety and Health

Regional Spotlight: the Midwest



An example of intense precipitation affecting transportation infrastructure was the record-breaking 24-hour rainstorm in July 1996, which resulted in flash flooding in Chicago and its suburbs, with major impacts. Extensive travel delays occurred on metropolitan highways and railroads, and streets and bridges were damaged. Commuters were unable to reach Chicago for up to three days, and more than 300 freight trains were delayed or rerouted.²²²

The June 2008 Midwest floods caused I-80 in eastern Iowa to be closed for more than five days, disrupting major east-west shipping routes for trucks and the east-west rail lines through Iowa. These floods exemplify the kind of extreme precipitation events and their direct impacts on transportation that are likely to become more frequent in a warming world. These extremes create new and more difficult problems that must be addressed in the design, construction, rehabilitation, and operation of the nation's transportation infrastructure.

Administration states that concern for heat stress for moderate to heavy work begins at about 80°F as measured by an index that combines temperature, wind, humidity, and direct sunlight. For dry climates, such as Phoenix and Denver, National Weather Service Heat Indices above 90°F might allow work to proceed, while higher humidity areas such as New Orleans or Miami should consider 80 to 85°F as an initial level for work restrictions.²²⁷ These trends and associated impacts will be exacerbated in many places by urban heat island effects (see *Human Health and Society* sectors).

Wildfires are projected to increase, especially in the Southwest (see *Southwest* region), threatening communities and infrastructure directly and bringing about road and rail closures in affected areas.

In many northern states, warmer winters will bring about reductions in snow and ice removal costs, lessen adverse environmental impacts from the use of salt and chemicals on roads and bridges, extend the construction season, and improve the mobility and safety of passenger and freight travel through reduced winter hazards. On the other hand, more freeze-thaw conditions are projected to occur in northern states, creating frost heaves and potholes on road and bridge surfaces and resulting in load restrictions on certain roads to minimize the damage. With the expected earlier onset of seasonal warming, the period of springtime load restrictions might be reduced in some areas, but it is likely to expand in others with shorter winters but longer thaw seasons. Longer construction seasons will be a benefit in colder locations.²²²

Water

Warming is projected to mean a longer shipping season but lower water levels for the Great Lakes and St. Lawrence Seaway. Higher temperatures, reduced lake ice, and increased evaporation are expected to combine to produce lower water levels as climate warming proceeds (see *Midwest* region). With lower lake levels, ships will be unable to carry as much cargo and hence shipping costs will increase. A recent study, for example, found that the projected reduction in Great Lakes water levels would result in an estimated 13 to 29 percent increase in shipping costs for Canadian commercial navigation by 2050, all else remaining equal.²²²

If low water levels become more common because of drier conditions due to climate change, this could create problems for river traffic, reminiscent of the stranding of more than 4,000 barges on the Mississippi River during the drought in 1988. Freight movements in the region could be seriously impaired, and extensive dredging could be required to keep shipping channels open. On the other hand, a longer shipping season afforded by a warmer climate could offset some of the resulting adverse economic effects.

Navigable Inland Waterways



CCSP SAP 4.7²¹⁷

Inland waterways are an important part of the transportation network in various parts of the United States. For example, these waterways provide 20 states with access to the Gulf of Mexico.²¹⁷ As conditions become drier, these main transportation pathways are likely to be adversely affected by the resulting lower water levels, creating problems for river traffic. Names of navigable rivers are shown above.

In cold areas, the projected decrease in very cold days will mean less ice accumulation on vessels, decks, riggings, and docks; less ice fog; and fewer ice jams in ports.²²²

Air

Rising temperatures will affect airport ground facilities, runways in particular, in much the same way they affect roads. Airports in some areas are likely to benefit from reduction in the cost of snow and ice removal and the impacts of salt and chemical use, though some locations have seen increases in snowfall. Airlines could benefit from reduced need to de-ice planes.

More heat extremes will create added operational difficulties, for example, causing greater energy consumption by planes on the ground. Extreme heat also affects aircraft lift; because hotter air is less dense, it reduces the lift produced by the wing and the thrust produced by the engine – problems exacerbated at high altitudes and high temperatures. As a result, planes need to take off faster, and if runways are not sufficiently long for aircraft to build up enough speed to generate lift, aircraft weight must be reduced. Thus, increases in extreme heat will result in payload restrictions, could cause flight cancellations and service disruptions

at affected airports, and could require some airports to lengthen runways. Recent hot summers have seen flights cancelled due to heat, especially in high altitude locations. Economic losses are expected at affected airports. A recent illustrative analysis projects a 17 percent reduction in freight carrying capacity for a single Boeing 747 at the Denver airport by 2030 and a 9 percent reduction at the Phoenix airport due to increased temperature and water vapor.²²²

Drought

Rising air temperatures increase evaporation, contributing to dry conditions, especially when accompanied by decreasing precipitation. Even where total annual precipitation does not decrease, precipitation is projected to become less frequent in many parts

of the country.⁶⁸ Drought is expected to be an increasing problem in some regions; this, in turn, has impacts on transportation. For example, increased susceptibility to wildfires during droughts could threaten roads and other transportation infrastructure directly, or cause road closures due to fire threat or reduced visibility such as has occurred in Florida and California in recent years. There is also increased susceptibility to mudslides in areas deforested by wildfires. Airports could suffer from decreased visibility due to wildfires. River transport is seriously affected by drought, with reductions in the routes available, shipping season, and cargo carrying capacity.

Increased intensity of strong hurricanes would lead to more evacuations, infrastructure damage and failure, and transportation interruptions.

More intense hurricanes in some regions are a projected effect of climate change. Three aspects of tropical storms are relevant to transportation: precipitation, winds, and wind-induced storm surge. Stronger hurricanes have longer periods of intense precipitation, higher wind speeds (damage increases exponentially with wind speed²²⁸),

R1
R2
R3
R4
R5
R6
R7
R8
R9
R10
R11
R12
R13
R14
R15
R16
R17
R18
R19
R20
R21
R22
R23
R24
R25
R26
R27
R28
R29
R30
R31
R32
R33
R34
R35
R36
R37
R38
R39
R40
R41
R42
R43
R44
R45
R46
R47
R48
R49
R50



L1 and higher storm surge and waves. Transportation
 L2 planners, designers, and operators might need to
 L3 adopt probabilistic approaches to developing trans-
 L4 portation projects rather than relying on standards
 L5 and the deterministic approaches of the past. The
 L6 uncertainty associated with projecting impacts over
 L7 a 50- to 100-year time period makes risk manage-
 L8 ment a reasonable approach for realistically incor-
 L9 porating climate change into decision-making and
 L10 investment.²¹⁵

L11 **Land**

L12 There will be a greater probability of infrastruc-
 L13 ture failures such as highway and rail bridge decks
 L14 being displaced and railroad tracks being washed
 L15 away. Storms leave debris on roads and rail lines,
 L16 which can damage the infrastructure and interrupt
 L17 travel and shipments of goods. In Louisiana, the
 L18 Department of Transportation and Development

R1 spent \$74 million for debris removal alone in the
 R2 wake of hurricanes Katrina and Rita. The Missis-
 R3 sippi Department of Transportation expected to
 R4 spend in excess of \$1 billion to replace the Biloxi
 R5 and Bay St. Louis bridges, repair other portions of
 R6 roadway, and remove debris. As of June 2007, more
 R7 than \$672 million had been spent.
 R8

R9 There will be more frequent and potentially more
 R10 extensive emergency evacuations. Damage to signs,
 R11 lighting fixtures, and supports will increase. The
 R12 lifetime of highways that have been exposed to
 R13 flooding is expected to decrease. Road and rail
 R14 infrastructure for passenger and freight services are
 R15 likely to face increased flooding by strong hurri-
 R16 canes. In the Gulf Coast, more than one-third of the
 R17 rail miles are likely to flood when subjected to a
 R18 storm surge of 18 feet.²¹⁷
 R19
 R20
 R21
 R22

L23 **Spotlight on Hurricane Katrina**



L24
 L25
 L26
 L27
 L28
 L29
 L30
 L31
 L32
 L33
 L34
 L35
 L36
 L37
 L38
 L39
 L40
 L41
 L42
 L43
 L44
 L45
 L46
 L47
 L48
 L49
 L50

R23 Hurricane Katrina was one of the most
 R24 destructive and expensive natural disasters in
 R25 U.S. history, claiming more than 1,800 lives and
 R26 causing an estimated \$134 billion in damage.^{217,229} It
 R27 also seriously disrupted transportation systems as key
 R28 highway and railroad bridges were heavily damaged or de-
 R29 stroyed, necessitating rerouting of traffic and placing increased
 R30 strain on other routes, particularly other rail lines. Replacement of
 R31 major infrastructure took from months to years. The CSX Gulf Coast line
 R32 was re-opened after five months and \$250 million in reconstruction costs, while the
 R33 Biloxi-Ocean Springs Bridge took more than two years to reopen. Barge shipping was halted, as
 R34 was grain export out of the Port of New Orleans, the nation's largest site of grain exports. The extensive
 R35 oil and gas pipeline network was shut down by the loss of electrical power, producing shortages of natural
 R36 gas and petroleum products. Total recovery costs for the roads, bridges, and utilities as well as debris
 R37 removal have been estimated at \$15 billion to \$18 billion.²¹⁷
 R38

L39 Redundancies in the transportation system, as well as the storm
 L40 timing and track, helped keep the storm from having major or
 L41 long-lasting impacts on national-level freight flows. For example,
 L42 truck traffic was diverted from the collapsed bridge that carries
 L43 highway I-10 over Lake Pontchartrain to highway I-12, which
 L44 parallels I-10 well north of the Gulf Coast. The primary north-
 L45 south highways that connect the Gulf Coast with major inland
 L46 transportation hubs were not damaged and were open for nearly
 L47 full commercial freight movement within days. The railroads were
 L48 able to route some traffic not bound directly for New Orleans through Memphis and other Midwest rail
 L49 hubs. While a disaster of historic proportions, the effects of Hurricane Katrina could have been even
 L50 worse if not for the redundancy and resilience of the transportation network in the area.



Hurricane Katrina damage to bridge

L1 **Water**

L2 All aspects of shipping are disrupted by major
 L3 storms. For example, freight shipments need to
 L4 be diverted from the storm region. Activities at
 L5 offshore drilling sites and coastal pumping faci-
 L6 lities are generally suspended and extensive damage
 L7 to these facilities can occur, as was amply demon-
 L8 strated during the 2005 hurricane season. Refiner-
 L9 ies and pipelines are also vulnerable to damage
 L10 and disruption due to the high winds and storm
 L11 surge associated with hurricanes and other tropical
 L12 storms (see *Energy* sector). Barges that are unable
 L13 to get to safe harbors can be destroyed or severely
 L14 damaged. Waves and storm surge will damage
 L15 harbor infrastructure such as cranes, docks, and
 L16 other terminal facilities. There are implications for
 L17 emergency evacuation planning, facility mainte-
 L18 nance, and safety management.

L19 **Air**

L20 More frequent interruptions in air service and
 L21 airport closures can be expected. Airport facili-
 L22 ties including terminals, navigational equipment,
 L23 perimeter fencing, and signs are likely to sustain
 L24 increased wind damage. Airports are frequently
 L25 located in low-lying areas and can be expected to
 L26 flood with more intense storms. As a response to
 L27 this vulnerability, some airports, such as LaGuard-
 L28 ia in New York City, are already protected by
 L29 levees. Eight airports in the Gulf Coast region of
 L30 Louisiana and Texas are located in historical 100-
 L31 year flood plains; the 100-year flood events will be
 L32 more frequent in the future creating the likelihood
 L33 of serious costs and disruption.²¹⁷

L34 **Arctic warming will continue to
 L35 reduce sea ice, lengthening the ocean
 L36 transport season, but also resulting in
 L37 greater coastal erosion due to waves.
 L38 Permafrost thaw in Alaska will damage
 L39 infrastructure. The ice-road season will
 L40 become shorter.**

L41 **Special issues in Alaska**

L42 Warming has been most rapid in high northern
 L43 regions. As a result, Alaska is warming at twice the
 L44 rate of the rest of the nation, bringing both major
 L45 opportunities and major challenges. Alaska's trans-
 L46 portation infrastructure differs sharply from that of

the lower 48 states. Although Alaska is twice the
 size of Texas, its population and road mileage are
 more like Vermont's. Only 30 percent of Alaska's
 roads are paved. Air travel is much more common
 than in other states. Alaska has 84 commercial air-
 ports and more than 3,000 airstrips, many of which
 are the only means of transport for rural communi-
 ties. Unlike other states, over much of Alaska, the
 land is generally more accessible in winter, when
 the ground is frozen and ice roads and bridges
 formed by frozen rivers are available.

Sea ice decline

The striking thinning and downward trend in the
 extent of Arctic sea ice is regarded as a consider-
 able opportunity for shippers. Continued reduction
 in sea ice should result in opening of additional
 ice-free ports, improved access to ports and natu-
 ral resources in remote areas, and longer shipping
 seasons, but it is likely to increase erosion rates on
 land as well, raising costs for maintaining ports and
 other transportation infrastructure.^{132,220}

Later this century and beyond, shippers are looking
 forward to new Arctic shipping routes, including
 the fabled Northwest Passage, which could provide
 significant costs savings in shipping times and
 distances. However, the next few decades are likely
 to be very unpredictable for shipping through these
 new routes. The past three decades have seen very
 high year-to-year variability of sea-ice extent in the
 Canadian Arctic, despite the overall decrease in
 September sea-ice extent. The loss of sea ice from
 the shipping channels of the Canadian Archipelago
 might actually allow more frequent intrusions of
 icebergs, which would continue to impede shipping
 through the Northwest Passage.

Lack of sea ice, especially on the northern shores of
 Alaska, creates conditions whereby storms pro-
 duce waves that cause serious coastal erosion.^{137,219}
 Already a number of small towns, roads, and
 airports are threatened by retreating coastlines,
 necessitating the planned relocation of these
 communities.^{132,220}

Thawing ground

The challenges warming presents for transportation
 on land are considerable.¹⁶⁴ For highways, thawing
 of permafrost causes settling of the roadbed and

R1
 R2
 R3
 R4
 R5
 R6
 R7
 R8
 R9
 R10
 R11
 R12
 R13
 R14
 R15
 R16
 R17
 R18
 R19
 R20
 R21
 R22
 R23
 R24
 R25
 R26
 R27
 R28
 R29
 R30
 R31
 R32
 R33
 R34
 R35
 R36
 R37
 R38
 R39
 R40
 R41
 R42
 R43
 R44
 R45
 R46
 R47
 R48
 R49
 R50



Arctic Sea Ice Decline



The pink line shows the average September sea ice extent from 1979 through the present. The white area shows September 2007 sea ice extent. In 2008, the extent was slightly larger than 2007, but the ice was thinner, resulting in a lower total volume of sea ice. In addition, recent years have had less ice that persisted over numerous years and more first-year ice, which melts more quickly.¹³⁹

frost heaves that adversely affect the integrity of the road structure and its load-carrying capacity. The majority of Alaska’s highways are located in areas where permafrost is discontinuous, and dealing with thaw settlement problems already claims a significant portion of highway maintenance dollars.

Bridges and large culverts are particularly sensitive to movement caused by thawing permafrost and are often much more difficult than roads to repair and modify for changing site conditions. Thus, designing these facilities to take climate change into account is even more critical than is the case for roads.

Another impact of climate change on bridges is increased scouring. Hotter, drier summers in Alaska have led to increased glacial melting and longer pe-

riods of high streamflows, causing both increased sediment in rivers and scouring of bridge supporting piers and abutments. Temporary ice roads and bridges are commonly used in many parts of Alaska to access northern communities and provide support for the mining and oil and gas industries. Rising temperatures have already shortened the season during which these critical facilities can be used. Like the highway system, the Alaska Railroad crosses permafrost terrain, and frost heave and settlement from thawing affect some portions of the track, increasing maintenance costs.^{28,132,220}

A significant number of Alaska’s airstrips in the southwest, northwest, and interior of the state are built on permafrost. These airstrips will require major repairs or relocation if their foundations are compromised by thawing.

The cost of maintaining Alaska’s public infrastructure is projected to increase 10 to 20 percent by 2030 due to warming, costing the state an additional \$4 billion to \$6 billion, with roads and airports accounting for about half of this cost.²³⁰ Private infrastructure impacts have not been evaluated.²¹⁷

The Trans-Alaska Pipeline System, which stretches from Prudhoe Bay in the north to the ice-free port of Valdez in the south, crosses a wide range of permafrost types and varying temperature conditions. More than half of the 800-mile pipeline is elevated on vertical supports over potentially unstable permafrost. Because the system was designed in the early 1970s on the basis of permafrost and climate conditions of the 1950 to 1970 period, it requires continuous monitoring and some supports have had to be replaced.

Travel over the tundra for oil and gas exploration and extraction is limited to the period when the ground is sufficiently frozen to avoid damage to the fragile tundra. In recent decades, the number of days that exploration and extraction equipment could be used has dropped from 200 days to 100 days per year due to warming. With warming, the number of exploration days is expected to decline even more.

L1
L2
L3
L4
L5
L6
L7
L8
L9
L10
L11
L12
L13
L14
L15
L16
L17
L18
L19
L20
L21
L22
L23
L24
L25
L26
L27
L28
L29
L30
L31
L32
L33
L34
L35
L36
L37
L38
L39
L40
L41
L42
L43
L44
L45
L46
L47
L48
L49
L50

R1
R2
R3
R4
R5
R6
R7
R8
R9
R10
R11
R12
R13
R14
R15
R16
R17
R18
R19
R20
R21
R22
R23
R24
R25
R26
R27
R28
R29
R30
R31
R32
R33
R34
R35
R36
R37
R38
R39
R40
R41
R42
R43
R44
R45
R46
R47
R48
R49
R50

Agriculture

Key Messages:

- Many crops show positive responses to elevated carbon dioxide and lower levels of warming, but higher levels of warming often negatively affect growth and yields.
- Extreme events such as heavy downpours and droughts are likely to reduce crop yields because excesses or deficits of water have negative impacts on plant growth.
- Weeds, diseases, and insect pests benefit from warming, and weeds also benefit from a higher carbon dioxide concentration, increasing stress on crop plants and requiring more attention to pest and weed control.
- Forage quality in pastures and rangelands generally declines with increasing carbon dioxide concentration because of the effects on plant nitrogen and protein content, reducing the land's ability to supply adequate livestock feed.
- Increased heat, disease, and weather extremes are likely to reduce livestock productivity.

Key Sources

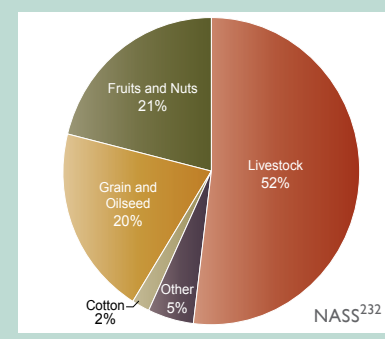


Agriculture in the United States is extremely diverse in the range of crops and animals grown and produces over \$200 billion a year in food commodities, with livestock accounting for more than half. Climate change will increase productivity in certain crops and regions and reduce productivity in others (see for example *Midwest* and *Great Plains* regions).¹⁹³

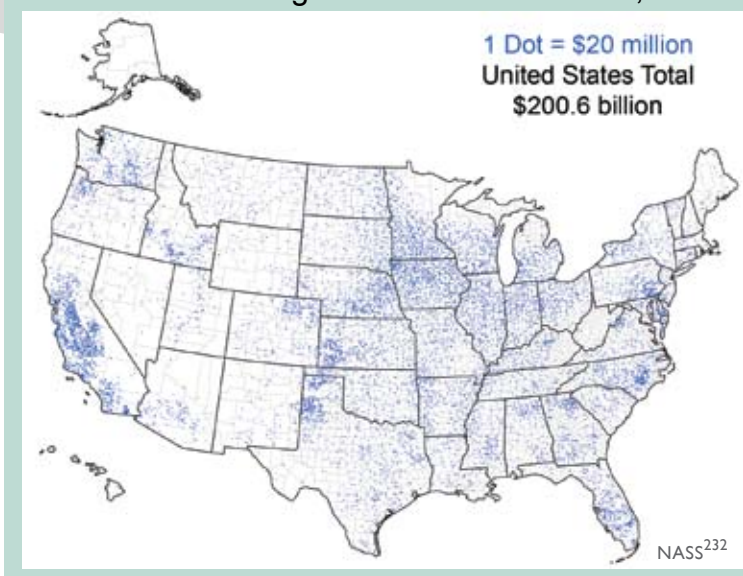
While climate change clearly affects agriculture, climate is also affected by agriculture, which contributes 13.5 percent of all human-induced greenhouse gas emissions globally. In the United States, agriculture represents 8.6 percent of the nation's total greenhouse gas emissions, including 80 percent of its nitrous oxide emissions and 31 percent of its methane emissions.²³¹

Increased agricultural productivity will be required in the future to supply the needs of an increasing population. Agricultural productivity is dependent upon the climatic and land resources. Climate change can have both beneficial and detrimental impacts on plants. Throughout history agricultural enterprises have coped with changes in climate through changes in management and in crop or animal selection. However, the projected climate changes are likely to challenge the United States capacity to as efficiently produce food, feed, fuel, and livestock products.

Relative Contributions to Agricultural Products, 2002



Market Value of Agricultural Products Sold, 2002



Many crops show positive responses to elevated carbon dioxide and lower levels of warming, but higher levels of warming often negatively affect growth and yields.

Crop responses in a changing climate reflect the interplay among three factors: rising temperatures, changing water resources, and increasing carbon dioxide concentrations. Warming generally causes plants that are below their optimum temperature to grow faster, with obvious benefits. For some plants, such as cereal crops, however, faster growth means there is less time for the grain itself to grow and mature, reducing yields.¹⁹³ For some annual crops, this can be compensated for by adjusting the planting date to avoid late season heat stress.¹⁶⁴

The grain-filling period (the time when the seed grows and matures) of wheat and other small grains shortens dramatically with rising temperatures. Analysis of crop responses suggests that even moderate increases in temperature will decrease yields of corn, wheat, sorghum, bean, rice, cotton, and peanut crops.

Some crops are particularly sensitive to high nighttime temperatures, which have been rising even faster than daytime temperatures.⁶⁸ Nighttime temperatures are expected to continue to rise in the future. These changes in temperature are especially critical to the reproductive phase of growth because warm nights increase the respiration rate and reduce the amount of carbon that is captured during the day by photosynthesis to be retained in

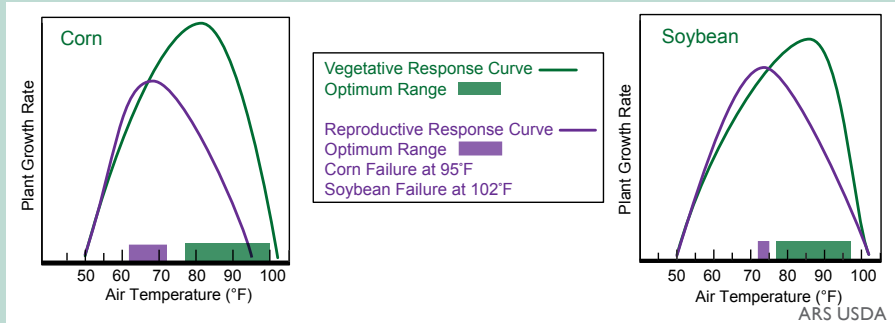
the fruit or grain. Further, as temperatures continue to rise and drought periods increase, crops will be more frequently exposed to temperature thresholds at which pollination and grain-set processes begin to fail and quality of vegetable crops decreases. Grain, soybean, and canola crops have relatively low optimal temperatures, and thus will have reduced yields and will increasingly begin to experience failure as warming proceeds.¹⁹³ Common snap beans show substantial yield reduction when nighttime temperatures exceed 80°F.

Higher temperatures will mean a longer growing season for crops that do well in the heat, such as melon, okra, and sweet potato, but a shorter growing season for crops more suited to cooler conditions, such as potato, lettuce, broccoli, and spinach.¹⁹³ Higher temperatures also cause plants to use more water to keep cool. This is one example of how the interplay between rising temperatures and water availability is critical to how plants respond to climate change. But fruits, vegetables, and grains can suffer even under well-watered conditions if temperatures exceed the maximum level for pollen viability in a particular plant; if temperatures exceed the threshold for that plant, it won't produce seed and so it won't reproduce.¹⁹³

Temperature increases will cause the optimum latitude for crops to move northward, while decreases in temperature will cause shifts toward the equator. Where plants can be efficiently grown depends upon climate conditions, of which temperature is one of the major factors.

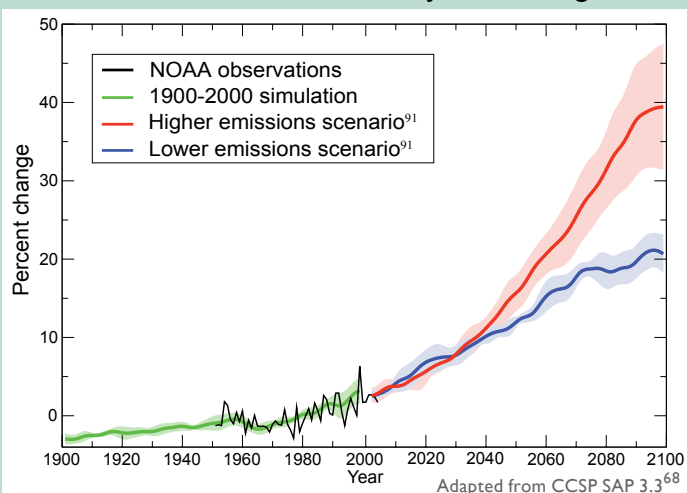
Plants need adequate water to maintain their temperature within an optimal range. Without water for cooling, plants will suffer heat stress. In many regions, irrigation water is used to maintain adequate temperature conditions for the growth of cool season plants (such as many vegetables), even in warm environments. With increasing demand and competition for freshwater supplies, the water needed for these crops might be increasingly limited. If water supply variability increases, it will affect plant growth

Corn and Soybean Temperature Response



For each plant variety, there is an optimal temperature for vegetative growth, with growth dropping off as temperatures increase or decrease. Similarly, there is a range of temperatures at which a plant will produce seed. Outside of this range, the plant will not reproduce. As the graphs show, corn will fail to reproduce at temperatures above 95°F and soybean above 102°F.

Increase in Percent of Very Warm Nights



The graph shows the observed and projected change in percent of very warm nights from the 1950-1990 average in the United States. Under the lower emissions scenario,⁹¹ the percentage of very warm nights is projected to increase about 20 percent by 2100. Under the higher emissions scenario,⁹¹ it is projected to increase by about 40 percent.⁶⁸ The shaded areas show the likely ranges while the lines show the central projections from a set of climate models. The projections appear smooth because they show the calculated average of many models.

and cause reduced yields. The amount and timing of precipitation during the growing season are also critical, and will be affected by climate change. Changes in season length are also important and affect crops differently.¹⁹³

Higher carbon dioxide levels generally cause plants to grow larger. For some crops, this is not necessarily a benefit because they are often less nutritious, with reduced nitrogen and protein content. Carbon dioxide also makes some plants more water-use efficient, meaning they produce more plant material, such as grain, on less water.¹⁹³ This is a benefit in water-limited areas and in seasons with less than normal rainfall amounts.

In some cases, adapting to climate change could be as simple as changing planting dates, which can be an effective no- or low-cost option for taking advantage of a longer growing season or avoiding crop exposure to adverse climatic conditions such as high temperature stress or low rainfall periods. Effectiveness will depend on the region, crop, and the rate and amount of warming. It is unlikely to be effective if a farmer goes to market when the supply-demand balance drives prices down. Predicting the optimum planting date for maximum profits will be very challenging in a future with increased

uncertainty regarding climate effects on not only local productivity, but also on supply from competing regions.

Another adaptation strategy involves changing to crop varieties with improved tolerance to heat or drought, or those that are adapted to take advantage of a longer growing season. This is less likely to be cost-effective for perennial crops, for which changing varieties is extremely expensive and new plantings take several years to reach maximum productivity. Even for annual crops, changing varieties is not always a low-cost option. Seed for new stress-tolerant varieties can be expensive, and new varieties often require investments in new planting equipment or require adjustments in a wide range of farming practices. In some cases, it is difficult to breed for genetic tolerance to elevated temperature or to identify an alternative variety that is

adapted to the new climate and to local soils, practices, and market demands.

Fruits that require long winter chilling periods will experience declines. Many varieties of fruits (such as popular varieties of apples and berries) require between 400 and 1,800 cumulative hours below 45°F each winter to produce abundant yields the following summer and fall. By late this century, under higher emissions scenarios,⁹¹ winter temperatures in many important fruit-producing regions such as the Northeast will be too consistently warm to meet these requirements. Cranberries have a particularly high chilling requirement, and there are no known low-chill varieties. Massachusetts and New Jersey supply nearly half the nation's cranberry crop. By the middle of this century, under higher emissions scenarios,⁹¹ it is unlikely that these areas will support cranberry production due to a lack of the winter chilling they need.^{233,234} Such impacts will vary by region. For example, though there will still be risks of early-season frosts and damaging winter thaws, warming is expected to improve the climate for fruit production in the Great Lakes region.¹⁶⁴

A seemingly paradoxical impact of warming is that it appears to be increasing the risk of plant frost

Effects of Increased Air Pollution on Crop Yields

Ground-level ozone (a component of smog) is an air pollutant that is formed when nitrogen oxides emitted from fossil fuel burning interact with other compounds, such as unburned gasoline vapors, in the atmosphere,²³⁷ in the presence of sunlight. Higher air temperatures result in greater concentrations of ozone. Ozone levels at the land surface have risen in rural areas of the United States over the past 50 years, and they are forecast to continue increasing with warming, especially under higher emissions scenarios.⁹¹ Plants are sensitive to ozone, and crop yields are reduced as ozone levels increase. Some crops that are particularly sensitive to ozone pollution include soybeans, wheat, oats, green beans, peppers, and some types of cotton.¹⁹³

damage. Mild winters and warm, early springs, which are beginning to occur more frequently as climate warms, induce premature plant development and blooming, resulting in exposure of vulnerable young plants and plant tissues to subsequent late-season frosts. For example, the 2007 spring freeze in the eastern United States caused widespread devastation of crops and natural vegetation because the frost occurred during the flowering period of many trees and during early grain development on wheat plants.²³⁵ Another example is occurring in the Rocky Mountains where in addition to the process described above, reduced snow cover leaves young plants unprotected from spring frosts, with some plant species already beginning to suffer as a result²³⁶ (see *Ecosystems* sector).

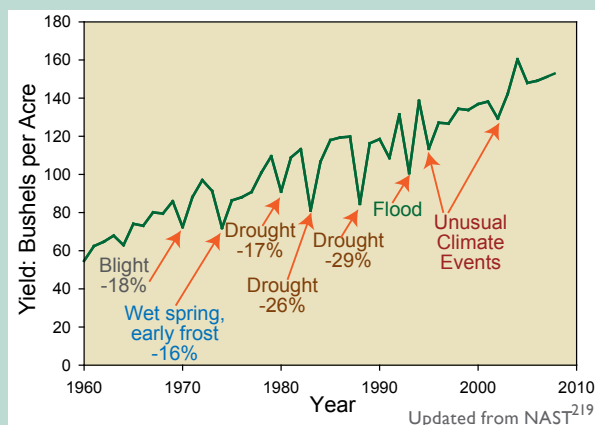
Extreme events such as heavy downpours and droughts are likely to reduce crop yields because excesses or deficits of water have negative impacts on plant growth.

One of the most pronounced effects of climate change is the increase in heavy downpours. Precipitation has become less frequent but more intense, and this pattern is projected to continue across the United States.¹¹² One consequence of excessive rainfall is delayed spring planting, which jeopardizes profits for farmers paid a premium for early season production of high-value crops such as melon, sweet corn, and tomatoes. Field flooding during the growing season causes crop losses due to low oxygen levels in the soil, increased susceptibility to root diseases, and increased soil compaction due to the use of heavy farm equipment on wet soils. In spring 2008, heavy rains caused the Mississippi River to rise to about 7 feet above flood

stage, inundating hundreds of thousands of acres of cropland. The flood hit just as farmers were preparing to harvest wheat and plant corn, soybeans, and cotton. Preliminary estimates of agricultural losses are around \$8 billion.²¹³ Some farmers were put out of business and others will be recovering for years to come. The flooding caused severe erosion in some areas and also caused an increase in runoff and leaching of agricultural chemicals into surface water and groundwater.²³³

Another impact of heavy downpours is that wet conditions at harvest time result in reduced quality of many crops. Storms with heavy rainfall often are accompanied by wind gusts, and both strong winds and rain can flatten crops, causing significant damage. Vegetable and fruit crops are sensitive to even short-term, minor stresses, and as such are par-

U.S. Corn Yields 1960 to 2008



While technological improvements have resulted in a general increase in corn yields, extreme weather events have caused dramatic reductions in yields in particular years. Increased variation in yield is likely to occur as temperatures increase and rainfall becomes more variable during the growing season. Without dramatic technological breakthroughs, yields are unlikely to continue their historical upward trend as temperatures rise above the optimum level for vegetative and reproductive growth.

L1 ticularly vulnerable to weather extremes.¹⁹³ More
L2 rainfall concentrated into heavy downpours also in-
L3 creases the likelihood of water deficiencies at other
L4 times because of reductions in rainfall frequency.
L5

L6 Drought frequency and severity are projected to in-
L7 crease in the future over much of the United States,
L8 particularly under higher emissions scenarios.^{90,91}
L9 Increased drought will be occurring at a time when
L10 crop water requirements also are increasing due to
L11 rising temperatures. Water deficits are detrimental
L12 for all crops.²³³
L13

L14 Temperature extremes also will pose problems.
L15 Even crop species that are well-adapted to warmth,
L16 such as tomatoes, can have reduced yield and/
L17 or quality when daytime maximum temperatures
L18 exceed 90°F for even short periods during critical
L19 reproductive stages.¹¹² For many high-value crops,
L20 just hours or days of moderate heat stress at critical
L21 growth stages can reduce grower profits by nega-
L22 tively affecting visual or flavor quality, even when
L23 total yield is not reduced.²³⁸
L24

L25
L26 **Weeds, diseases, and insect pests**
L27 **benefit from warming, and weeds also**
L28 **benefit from a higher carbon dioxide**
L29 **concentration, increasing stress on crop**
L30 **plants and requiring more attention to**
L31 **pest and weed control.**
L32

L33 Weeds benefit more than cash crops from
L34 higher temperatures and carbon dioxide
L35 levels.¹⁹³ One concern with continued
L36 warming is the northward expansion of
L37 invasive weeds. Southern farmers currently
L38 lose more of their crops to weeds than do
L39 northern farmers. For example, southern
L40 farmers lose 64 percent of the soybean
L41 crop to weeds, while northern farmers lose
L42 22 percent.²³⁹ Some extremely aggressive
L43 weeds plaguing the South (such as kudzu)
L44 have historically been confined to areas
L45 where winter temperatures do not drop
L46 below specific thresholds. As temperatures
L47 continue to rise, these weeds will expand
L48 their ranges northward into important
L49 agricultural areas.²⁴⁰ Kudzu currently has
L50

invaded 2.5 million acres of the Southeast and is a
carrier of the fungal disease soybean rust, which
represents a major and expanding threat to U.S.
soybean production.²³⁴

Controlling weeds currently costs the United States
more than \$11 billion a year, with the majority
spent on herbicides;²⁴¹ so both herbicide use and
costs are likely to increase as temperatures and
carbon dioxide levels rise. At the same time, the
most widely used herbicide in the United States,
glyphosate (RoundUp®), loses its efficacy on weeds
grown at carbon dioxide levels that are projected
to occur in the coming decades. Higher concentra-
tions of the chemical and more frequent spraying
thus will be needed, increasing economic and envi-
ronmental costs associated with chemical use.²³³

Many insect pests and crop diseases thrive due
to warming, increasing losses and necessitating
greater pesticide use. Warming aids insects and
diseases in several ways. Rising temperatures
allow both insects and pathogens to expand their
ranges northward. In addition, rapidly rising winter
temperatures allow more insects to survive over the
winter, whereas cold winters once controlled their
populations. Some of these insects, in addition to
directly damaging crops, also carry diseases that
harm crops. Crop diseases in general are likely to
increase as earlier springs and warmer winters al-

Increasing CO₂ Reduces Herbicide Effectiveness



Current CO₂ (380 ppm)

Potential Future CO₂ (680 ppm)

The left photo shows weeds in a plot grown at a carbon dioxide (CO₂) concentration of about 380 parts per million (ppm), which approximates the current level. The right photo shows a plot in which the CO₂ level has been raised to about 680 ppm.²³³

L1
L2
L3
L4
L5
L6
L7
L8
L9
L10
L11
L12
L13
L14
L15
L16
L17
L18
L19
L20
L21
L22
L23
L24
L25
L26
L27
L28
L29
L30
L31
L32
L33
L34
L35
L36
L37
L38
L39
L40
L41
L42
L43
L44
L45
L46
L47
L48
L49
L50

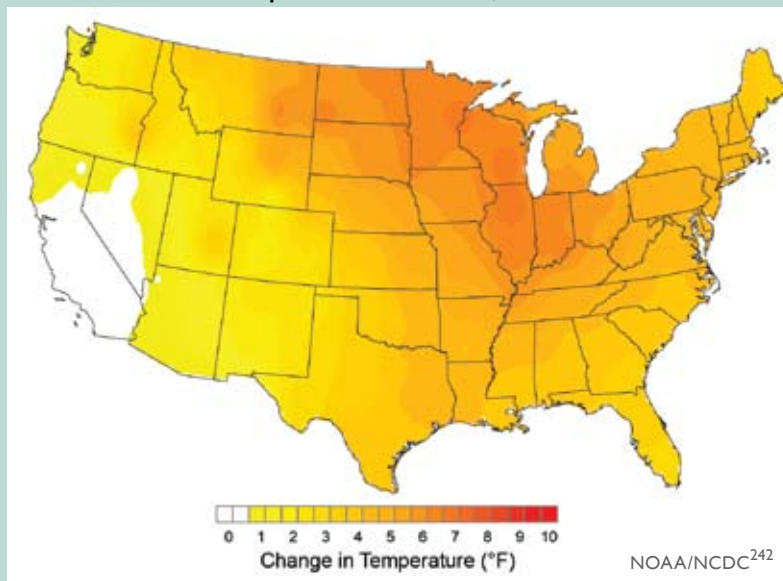
low proliferation and higher survival rates of disease pathogens and parasites.^{193,234} The longer growing season will allow some insects to produce more generations in a single season, greatly increasing their populations. Finally, plants grown in higher carbon dioxide conditions tend to be less nutritious, so insects must eat more to meet their protein requirements, causing greater destruction to crops.¹⁹³

Due to the increased presence of pests, spraying is already much more common in warmer areas than in cooler areas. For example, Florida sweet corn growers spray their fields 15 to 32 times a year to fight pests such as corn borer and corn earworm, while New York farmers average zero to five times. In addition, higher temperatures are known to reduce the effectiveness of certain classes of pesticides (pyrethroids and spinosad).

A particularly unpleasant example of how carbon dioxide tends to favor undesirable plants is found in the response of poison ivy to rising carbon dioxide concentrations. Poison ivy thrives in air with extra carbon dioxide in it, growing bigger and producing a more toxic form of the oil, urushiol, which causes painful skin reactions in 80 percent of people. Contact with poison ivy is one of the most widely reported ailments at poison centers in the United States, causing more than 350,000 cases of contact dermatitis each year. The growth stimulation of poison ivy due to increasing carbon dioxide concentration exceeds that of most other woody species. Given continued increases in carbon dioxide emissions, poison ivy is expected to become more abundant and more toxic in the future, with implications for forests and human health.²³⁴

Higher temperatures, longer growing seasons, and increased drought will lead to increased agricultural water use in some areas. Obtaining the maximum “carbon dioxide fertilization” benefit often requires more efficient use of water and fertilizers that better synchronize plant demand with supply.

Winter Temperature Trends, 1975 to 2007



Temperatures are rising faster in winter than in any other season, especially in many key agricultural regions. This allows many insect pests and crop diseases to expand and thrive, creating increasing challenges for agriculture. As indicated by the map, the Midwest and northern Great Plains have experienced increases of more than 7°F in average winter temperatures over the past 30 years.

R1
R2
R3
R4
R5
R6
R7
R8
R9
R10
R11
R12
R13
R14
R15
R16
R17
R18
R19
R20
R21
R22
R23
R24
R25
R26
R27
R28
R29
R30
R31
R32
R33
R34
R35
R36
R37
R38
R39
R40
R41
R42
R43
R44
R45
R46
R47
R48
R49
R50

Farmers are likely to respond to more aggressive and invasive weeds, insects, and pathogens with increased use of herbicides, insecticides, and fungicides. Where increases in water and chemical inputs become necessary, this will increase costs for the farmer, as well as having society-wide impacts by depleting water supply, increasing reactive nitrogen and pesticide loads to the environment, and increasing risks to food safety and human exposure to pesticides.

Forage quality in pastures and rangelands generally declines with increasing carbon dioxide concentration because of the effects on plant nitrogen and protein content, reducing the land’s ability to supply adequate livestock feed.

Beef cattle production takes place in every state in the United States, with the greatest number raised in regions that have an abundance of native or planted pastures for grazing. Generally, eastern pasturelands are planted and managed, whereas western rangelands are native pastures, which are not seeded and receive much less rainfall. There are transformations now underway in many semi-arid rangelands as a result of increasing atmospheric

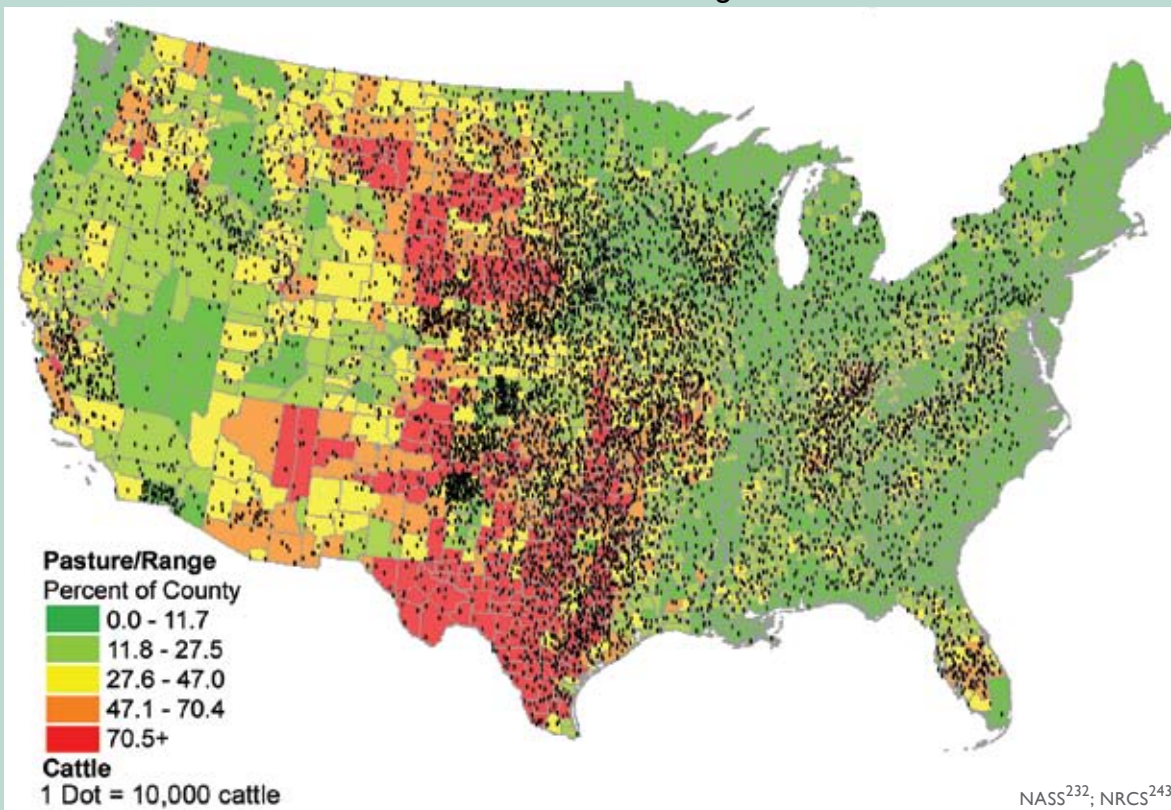
L1 carbon dioxide concentration and the associated
 L2 climate change. These transformations include
 L3 which species of grasses dominate, as well as the
 L4 forage quality of the dominant grasses. Increases in
 L5 carbon dioxide are generally reducing the quality
 L6 of the forage, so that more acreage is needed to
 L7 provide animals with the same nutritional value,
 L8 resulting in an overall decline in livestock pro-
 L9 ductivity. In addition, woody shrubs and invasive
 L10 cheatgrass are encroaching into grasslands, further
 L11 reducing their forage value.¹⁹³ The combination of
 L12 these factors leads to an overall decline in livestock
 L13 productivity.

L15 The rising atmospheric carbon dioxide concentra-
 L16 tion affects forage quality because plant nitrogen
 L17 and protein concentrations often decline with
 L18 higher concentrations of carbon dioxide.¹⁹³ This
 L19 reduction in protein reduces forage quality and
 L20 counters the positive effects of carbon dioxide-
 L21 enrichment on plant production and carbohydrates.
 L22 Rising carbon dioxide concentration might reduce
 L23 the digestibility of forages that are already of poor

quality. Reductions in forage quality could have
 pronounced detrimental effects on animal growth,
 reproduction, and survival, and could render live-
 stock production unsustainable unless animal diets
 are supplemented with protein, adding more costs
 to the production. On shortgrass prairie, for exam-
 ple, carbon dioxide enrichment reduced the protein
 concentration of autumn forage below critical
 maintenance levels for livestock in 3 out of 4 years
 and reduced the digestibility of forage by 14 per-
 cent in mid-summer and by 10 percent in autumn.
 Significantly, the grass type that thrived the most
 under excess carbon dioxide conditions also had the
 lowest protein concentration.¹⁹³

At the scale of a region, the composition of forage
 plant species is determined mostly by climate and
 soils. The primary factor controlling the distribu-
 tion and abundance of plants is water: both the
 amount of water plants use and water availability
 over time and space. The ability to anticipate veg-
 etation changes at local scales and over shorter pe-
 riods is limited because at these scales the response

Distribution of Beef Cattle and Pasture/Rangeland in Continental U.S.



The colors show the percent of the county that is cattle pasture or rangeland, with red indicating the highest percentage. Each dot represents 10,000 cattle. Livestock production occurs in every state. Increasing concentration of carbon dioxide reduces the quality of forage, necessitating more acreage and resulting in a decline in livestock production.

L1 of vegetation to global-scale changes depends on
 L2 a variety of local processes including the rate of
 L3 disturbances such as fire and grazing, and the rate
 L4 at which plant species can move across sometimes-
 L5 fragmented landscapes. Nevertheless, some general
 L6 patterns of vegetation change are beginning to
 L7 emerge. For example, experiments indicate that a
 L8 higher carbon dioxide concentration favors weeds
 L9 and invasive plants over native species because
 L10 invasives have traits (such as rapid growth rate or
 L11 prolific seed production) that allow a larger growth
 L12 response to carbon dioxide. In addition, the effect
 L13 of a higher carbon dioxide concentration on plant
 L14 species composition appears to be greatest where
 L15 the land has been disturbed (such as by fire or graz-
 L16 ing) and nutrient and light availability are high.¹⁹³

L17
 L18 Increases in temperature lengthen the growing sea-
 L19 son, and thus are likely to extend forage production
 L20 into the late fall and early spring. However, overall
 L21 productivity remains dependent on precipitation
 L22 during the growing season.¹⁹³

L23
 L24
 L25
 L26
 L27 **Increased heat, disease, and weather
 extremes are likely to reduce livestock
 productivity.**

L28
 L29 Like human beings, cows, pigs, and poultry are
 L30 warm-blooded animals that are sensitive to heat. In
 L31 terms of production efficiency, studies show that
 L32 the negative effects of hotter summers will out-
 L33 weigh the positive effects of warmer winters. The
 L34 more the U.S. climate warms, the more production
 L35 will fall. For example, an analysis projected that a
 L36 warming in the range of 9 to 11°F (as in the higher
 L37 emissions scenarios⁹¹) would cause a 10 percent
 L38 decline in livestock yields in cow/calf and dairy
 L39 operations in Appalachia, the Southeast (including
 L40 the Mississippi Delta), and southern Plains regions,
 L41 while a warming of 2.7°F would cause less than a 1
 L42 percent decline.

L43
 L44 Temperature and humidity interact to cause stress
 L45 in animals, just as in humans; the higher the heat
 L46 and humidity, the greater the stress and discomfort,
 L47 and the larger the reduction in the animals' ability
 L48 to produce milk, gain weight, and reproduce. Milk
 L49 production declines in dairy operations, the number
 L50 of days it takes for cows to reach their target weight

grows longer in meat operations, conception rate in
 cattle falls, and swine growth rates decline due to
 heat. As a result, swine, beef, and milk production
 are all projected to decline in a warmer world.¹⁹³

R1
 R2
 R3
 R4
 R5
 R6 The projected increases in air temperatures will
 R7 negatively affect confined animal operations (dairy,
 R8 beef, and swine) located in the central United
 R9 States, increasing production costs as a result of
 R10 reductions in performance associated with lower
 R11 feed intake and increased requirements for energy
 R12 to maintain healthy livestock. These costs do not
 R13 account for the increased death of livestock as-
 R14 sociated with extreme weather events such as heat
 R15 waves. Nighttime recovery is an essential element
 R16 of survival when livestock are stressed by extreme
 R17 heat. A feature of recent heat waves is the lack of
 R18 nighttime relief. Large numbers of deaths have oc-
 R19 curred in recent heat waves, with individual states
 R20 reporting losses of 5,000 head of cattle in a single
 R21 heat wave in one summer.¹⁹³

R22
 R23 Warming also affects parasites and disease patho-
 R24 gens. The earlier arrival of spring and warmer win-
 R25 ters allow greater proliferation and survival of para-
 R26 sites and disease pathogens. In addition, changes in
 R27 rainfall distributions are likely to lead to changes in
 R28 diseases sensitive to moisture. Heat stress reduces
 R29 animals' ability to cope with other stresses, such
 R30 as diseases and parasites. Furthermore, changes
 R31 in rainfall distributions could lead to changes in
 R32 diseases sensitive to relative humidity.

R33
 R34 Maintaining livestock production would require
 R35 modifying facilities to reduce heat stress on ani-
 R36 mals, using the best understanding of both the
 R37 chronic and acute stresses that livestock will
 R38 encounter to determine the optimal modification
 R39 strategy.

R40
 R41 Changing livestock species as an adaptation strat-
 R42 egy is a much more extreme, high-risk, and, in
 R43 most cases, high-cost option than changing crop
 R44 varieties. Accurate predictions of climate trends
 R45 and development of the infrastructure and market
 R46 for the new livestock products are essential to mak-
 R47 ing this an effective response.

Ecosystems

Key Messages:

- Ecosystem processes, such as those that control growth and decomposition, have been affected by climate change.
- Large-scale shifts have occurred in the ranges of species and the timing of the seasons and animal migration, and are very likely to continue.
- Fires, insect pests, disease pathogens, and invasive weed species have increased, and these trends are likely to continue.
- Deserts and drylands are likely to become hotter and drier, feeding a self-reinforcing cycle of invasive plants, fire, and erosion.
- Coastal and near-coastal ecosystems are already under multiple stresses. Climate change and ocean acidification will exacerbate these stresses.
- Arctic sea-ice ecosystems are already being adversely affected by the loss of summer sea ice and further changes are expected.
- The habitats of some mountain species and coldwater fish, such as salmon and trout, are very likely to contract in response to warming.
- Some of the benefits ecosystems provide to society will be threatened by climate change, while others will be enhanced.

Key Sources



The natural functioning of the environment provides both goods – such as food and other products that are bought and sold – and services, which our society depends upon. For example, ecosystems store large amounts of carbon in plants and soils; they regulate water flow and water quality; and they stabilize local climates. These services are not assigned a financial value, but society nonetheless depends on them. Ecosystem processes are the underpinning of these services: photosynthesis, the process by which plants capture carbon dioxide from the atmosphere and create new growth; the plant and soil processes that recycle nutrients from decomposing matter and maintain soil fertility; and the processes by which plants draw water from soils and return water to the atmosphere. These ecosystem processes are affected by climate and by the concentration of carbon dioxide in the atmosphere.⁷⁰

The diversity of living things (biodiversity) in ecosystems is itself an important resource that maintains the ability of these systems to provide the services upon which society depends. Many factors affect biodiversity including: climatic conditions; the influences of competitors, predators, parasites, and diseases; disturbances such as fire; and other physical factors. Human-induced climate change,

in conjunction with other stresses, is exerting major influences on natural environments and biodiversity, and these influences are generally expected to grow with increased warming.⁷⁰

Ecosystem processes, such as those that control growth and decomposition, have been affected by climate change.

Climate has a strong influence on the processes that control growth and development in ecosystems. Temperature increases generally speed up plant growth, rates of decomposition, and how rapidly the cycling of nutrients occurs, though other factors, such as whether sufficient water is available, also influence these rates. The growing season is lengthening as higher temperatures occur earlier in the spring. Forest growth has risen over the past several decades as a consequence of a number of factors – young forests reaching maturity, an increased concentration of carbon dioxide in the atmosphere, a longer growing season, and increased deposition of nitrogen from the atmosphere. Based on the current understanding of these processes, the individual effects are difficult to disentangle.²⁴⁴

Butterfly Range Shifts Northward



As climate warms, many species in the United States are shifting their ranges northward and to higher elevations. The map shows the response of Edith's checkerspot butterfly populations to a warming climate over the past 136 years in the American West. Over 70 percent of the southernmost populations (shown in yellow) have gone extinct. The northernmost populations and those above 8,000 feet elevation in the cooler climate of California's Sierra Nevada (shown in green) are still thriving. These differences in numbers of population extinctions across the geographic range of the butterfly have resulted in the average location shifting northward and to higher elevations over the past century, illustrating how climate change is altering the ranges of many species. Because their change in range is slow, most species are not expected to be able to keep up with the rapid climate change projected in the coming decades.²⁴⁵

A higher atmospheric carbon dioxide concentration causes trees and other plants to capture more carbon from the atmosphere, but experiments show that trees put much of this extra carbon into producing fine roots and twigs, rather than new wood. The effect of carbon dioxide in increasing growth thus seems to be relatively modest, and generally is seen most strongly in young forests on fertile soils where there is also sufficient water to sustain this growth. In the future, as atmospheric carbon dioxide continues to rise, and as climate continues to change, forest growth in some regions is projected to increase, especially in relatively young forests on fertile soils.²⁴⁴

Forest productivity is thus projected to increase in much of the East, while it is projected to decrease in much of the West where water is scarce and projected to become more so. Wherever droughts increase, forest productivity will decrease and tree death will increase. In addition to occurring in much of the West, these conditions are projected to occur in parts of Alaska and in the eastern part of the Southeast.²⁴⁴

Large-scale shifts have occurred in the ranges of species and the timing of the seasons and animal migration, and are very likely to continue.

Climate change already is having impacts on animal and plant species throughout the United States. Some of the most obvious changes are related to the timing of the seasons: when plants bud in spring, when birds and other animals migrate, and so on. In the United States, spring now arrives an average of 10 days to two weeks earlier than it did 20 years ago. The growing season is lengthening over much of the continental United States. Many migratory bird species are arriving earlier. For example, a study of northeastern birds that migrate long distances found that birds wintering in the southern United States now arrive back in the Northeast an average of 13 days earlier than they did during the first half of the last century. Birds wintering in South America arrive back in the Northeast an average of four days earlier.⁷⁰

Another major change is in the geographic distribution of species. The ranges of many species in the United States have shifted northward and upward in elevation. For example, the ranges of many butterfly species have expanded northward, contracted at the southern edge, and shifted to higher elevations as warming has continued. A study of Edith's checkerspot butterfly showed that 40 percent of the populations below 2,400 feet have gone extinct, despite the availability of suitable habitat and food supply. The checkerspot's most southern populations also have gone extinct, while new populations have been established north of the previous northern boundary for the species.⁷⁰

For butterflies, birds, and other species, one of the concerns with such changes in geographic range and timing of migration is the potential for mismatches between species and the resources they need to survive. The rapidly changing landscape, such as new highways and expanding urban areas, can create barriers that limit habitat and increase species loss. Failure of synchronicity between butterflies and the resources they depend

upon has led to local population extinctions of the checkerspot butterfly during extreme drought and low-snowpack years in California.⁷⁰

Tree species shifts

Forest tree species also are expected to shift their ranges northward and upslope in response to climate change, although specific quantitative predictions are very difficult to make because of the complexity of human land use and many other factors. This would result in major changes in the character of U.S. forests and the types of forests that will be most prevalent in different regions. In the United States, some common forest types are projected to expand, such as oak-hickory; others are projected to contract, such as maple-beech-birch. Still others, such as spruce-fir, are likely to disappear from the United States altogether.²⁴⁴

In Alaska, vegetation changes are already underway due to warming. The tree line is shifting northward into tundra, encroaching on the habitat for many migratory birds and land animals such as caribou that depend on the open tundra landscape.²⁴⁶

Marine species shifts and effects on fisheries

The distribution of marine fish and plankton are predominantly determined by climate, so it is not surprising that marine species in U.S. waters are moving northward and that the timing of plankton blooms is shifting. Extensive shifts in the ranges and distributions of both warm-water and coldwater species of fish have been documented.⁷⁰ For example, in the waters around Alaska, climate change already is causing significant alterations in marine ecosystems with important implications for fisheries and the people who depend on them (see *Alaska* region).

In the Pacific, climate change is expected to cause an eastward shift in the location of tuna stocks.²⁴⁷ It is clear that such shifts are related to climate, including natural modes of climate variability such as the cycles of El Niño and La Niña.

However, it is unclear how these modes of ocean variability will change as global climate continues to change, and there-

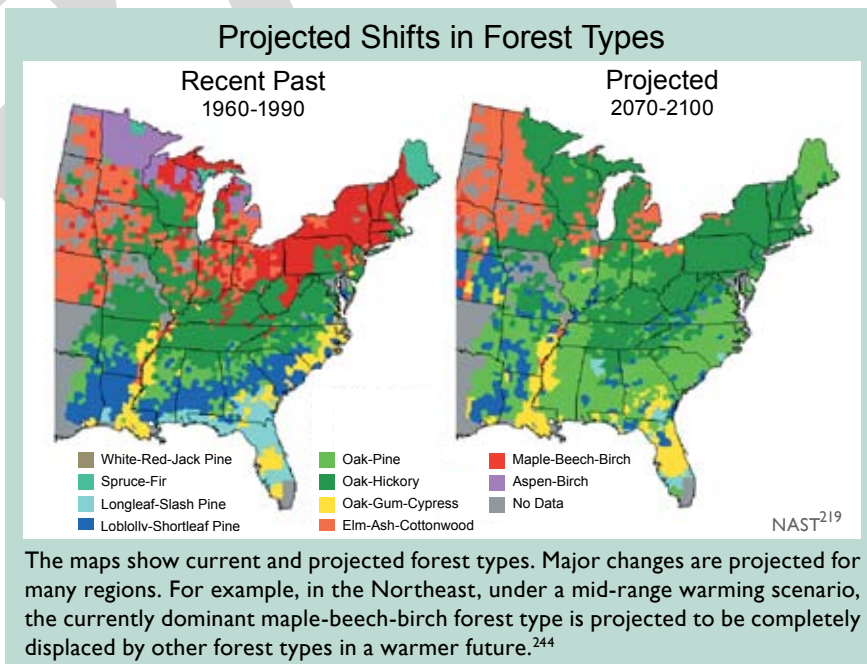
fore it is very difficult to predict quantitatively how marine fish and plankton species' distributions might shift as a function of climate change.⁷⁰

Breaking up of existing ecosystems

As warming drives changes in timing and geographic range for various species, it is important to note that entire communities of species do not shift intact. Rather, the range and timing of each species shifts in response to its sensitivity to climate change, its mobility, its lifespan, and the availability of the resources it needs (such as soil, moisture, food, and shelter). The ranges of animals can generally shift much faster than those of plants, and large migratory animals can move faster than small ones. In addition, migratory pathways must be available, such as northward flowing rivers which serve as conduits for fish. Some migratory pathways might be blocked by development and habitat fragmentation. All of these variations result in the breakup of existing ecosystems and formation of new ones, with unknown consequences.²²⁰

Extinctions and climate change

Interactions among impacts of climate change and other stressors can increase the risk of species extinction. Extinction rates of plants and animals have already risen considerably, with the vast majority of these extinctions attributed to loss of habitat or over-exploitation.²⁴⁸ Climate change has been identified as a serious risk factor for the fu-



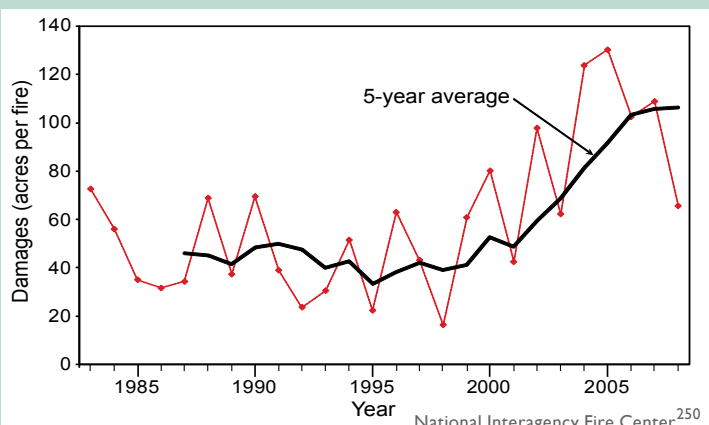
ture, however, since it is one of the environmental stresses on species and ecosystems that is continuing to increase.²⁴⁸ The Intergovernmental Panel on Climate Change has estimated that if a warming of 3.5 to 5.5°F occurs, 20 to 30 percent of species that have been studied would be in climate zones that are far outside of their current ranges, and would therefore likely be at risk of extinction.²⁴⁹ One reason this percentage is so high is that climate change would be superimposed on other stresses including habitat loss and continued overharvesting of some species, resulting in considerable stress on populations and species.

Fires, insect pests, disease pathogens, and invasive weed species have increased, and these trends are likely to continue.

Forest fires

In the western United States, both the frequency of large wildfires and the length of the fire season have increased substantially in recent decades, due to earlier spring snowmelt and higher spring and summer temperatures. These changes in climate have reduced the availability of moisture, drying out the vegetation that provides the fuel for fires. Alaska also has experienced large increases in fire, with the area burned more than doubling in recent decades. As in the western United States, higher air temperature is a key factor. In Alaska, for example, June air temperatures alone explained approximately 38 percent of the increase in the area burned annually from 1950 to 2003.²⁴⁴

Size of U.S. Wildfires, 1983 to 2008



Data on wildland fires in the United States show that the number of acres burned per fire has increased since the 1980s.

Insect pests

Insect pests are economically important stresses on forest ecosystems in the United States. Coupled with pathogens, they cost \$1.5 billion in damage per year. Forest insect pests are sensitive to climatic variations in many stages of their lives. Changes in climate have contributed significantly to several major insect pest outbreaks in the United States and Canada over the past several decades. The mountain pine beetle has infested lodgepole pine in British Columbia. Over 33 million acres of forest have been affected, by far the largest such outbreak in recorded history. Another 1.5 million acres have been infested by pine beetle in Colorado. Spruce beetle has affected more than 2.5 million acres in Alaska (see *Alaska* region) and western Canada. The combination of drought and high temperatures also has led to serious insect infestations and death of piñon pine in the Southwest, and to various insect pest attacks throughout the forests of the eastern United States.²⁴⁴

Rising temperatures increase insect outbreaks in a number of ways. First, warmer winters allow larger populations of insects to survive the cold season that normally limits their numbers. Second, the longer warm season allows them to develop faster, sometimes completing two life cycles instead of one in a single growing season. Third, warmer conditions help expand their ranges northward. And fourth, drought stress reduces trees' ability to resist insect attack (for example, by pushing back against boring insects with the pressure of their sap). Spruce beetle, pine beetle, spruce budworm, and woolly adelgid (which attacks eastern hemlocks) are just some of the insects that are proliferating in the United States, causing devastation in many forests. These outbreaks are projected to increase with ongoing warming. Trees killed by insects also provide more dry fuel for wildfires.^{70,244,251}

Disease pathogens and their carriers

One consequence of a longer, warmer growing season and less extreme cold in winter is that opportunities are created for many insect pests and disease pathogens to flourish. Accumulating evidence links the spread of disease pathogens to a warming climate. For example, a recent study showed that widespread amphibian extinctions in the mountains of Costa Rica are linked to changes in climatic

L1 conditions which are thought to have enabled the
L2 proliferation of an amphibian disease.^{70,252}
L3

L4 Diseases that affect wildlife and the living things
L5 that carry these diseases have been expanding their
L6 geographic ranges as climate heats up. Depending
L7 on their specific adaptations to current climate,
L8 many parasites, and the insects, spiders, and
L9 scorpions that carry and transmit diseases, die
L10 or fail to develop below threshold temperatures.
L11 Therefore, as temperatures rise, more of these
L12 disease-carrying creatures survive. For some
L13 species, rates of reproduction, population growth,
L14 and biting, tend to increase with increasing
L15 temperatures, up to a limit. Some parasites'
L16 development rates and infectivity periods also
L17 increase with temperature.⁷⁰ An analysis of diseases
L18 among marine species found that diseases were
L19 increasing for mammals, corals, turtles, and
L20 mollusks, while no trends were detected for sharks,
L21 rays, crabs, and shrimp.⁷⁰
L22

L23 **Invasive plants**

L24 Problems involving invasive plant species arise
L25 from a mix of human-induced changes, including
L26 disturbance of the land surface (such as through
L27 over-grazing or clearing natural vegetation for
L28 development), deliberate or accidental transport of
L29 non-native species, the increase in available nitrogen
L30 through over-fertilization of crops, and the rising
L31 carbon dioxide concentration and the resulting
L32 climate change.²⁴⁴ Human-induced climate change
L33 is not generally the initiating factor, nor the
L34 most important one, but it is becoming a more
L35 important part of the mix.
L36

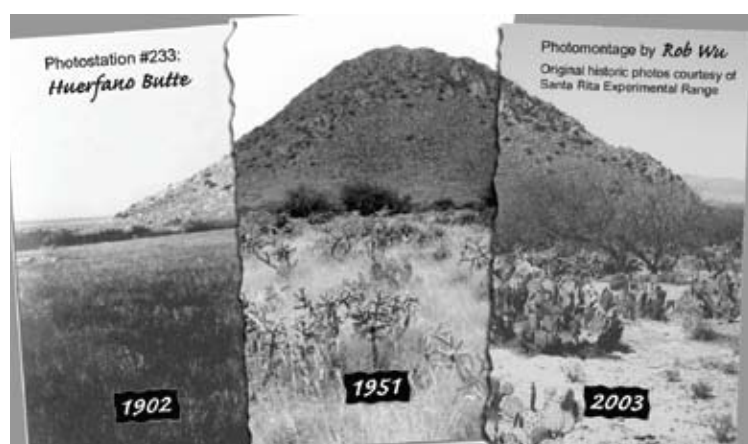
L37 The increasing carbon dioxide concentration
L38 stimulates the growth of most plant species,
L39 and some invasive plants respond with greater
L40 growth rates than native plants. Beyond this,
L41 invasive plants appear to better tolerate a
L42 wider range of environmental conditions and
L43 might be more successful in a warming world
L44 because they can migrate and establish themselves
L45 in new sites more rapidly than native
L46 plants.⁷⁰ They are also not usually dependent
L47 on external pollinators or seed dispersers to
L48 reproduce. For all of these reasons, invasive
L49 plant species present a growing problem that is
L50 extremely difficult to control once unleashed.⁷⁰

R1 **Deserts and drylands are likely to R2 become hotter and drier, feeding a self- R3 reinforcing cycle of invasive plants, fire, R4 and erosion.**

R5
R6 The arid region of the Southwest is projected to be-
R7 come drier in this century. There is emerging evi-
R8 dence that these changes are already underway.³⁴
R9 Deserts in the United States also are projected to
R10 expand to the north, east, and upward in elevation
R11 in response to projected warming and associated
R12 changes in climate.
R13

R14 Increased drying in the region contributes to a va-
R15 riety of changes that exacerbate a cycle of desertifi-
R16 cation. Increased drought conditions cause peren-
R17 nial plants to die due to water stress and increased
R18 susceptibility to plant diseases. At the same time,
R19 non-native grasses have invaded the region. As
R20 these grasses increase in abundance, they pro-
R21 vide more fuel for fires, causing fire frequency to
R22 increase in a self-reinforcing manner that leads to
R23 further losses of vegetation. When it does rain, the
R24 rain tends to come in heavy downpours, and since
R25 there is less vegetation to protect the soil, water
R26 erosion increases. Higher air temperatures and de-
R27 creased soil moisture reduce soil stability, further
R28 exacerbating erosion. And with a growing popula-
R29 tion needing water for urban uses, hydroelectric
R30 generation, and agriculture, there is increasing
R31 pressure on mountain water sources that would oth-
R32 erwise flow to desert river areas.^{70,149}
R33
R34
R35
R36
R37
R38
R39
R40
R41
R42
R43
R44
R45
R46
R47
R48
R49
R50

R36 **Desertification of Arid Grassland R37 near Tucson, Arizona, 1902 to 2003**



R36 The photo series shows the progression from arid grassland to desert
R37 (desertification) over a 100-year period. The change is the result of grazing
R38 management and reduced rainfall in the Southwest.^{251,253,254}
R39
R40
R41
R42
R43
R44
R45
R46
R47
R48
R49
R50

L1 The response of arid lands to climate change also
 L2 depends on how other factors interact with climate
 L3 at local scales. Large-scale, unregulated livestock
 L4 grazing in the Southwest during the late 1800s and
 L5 early 1900s is widely regarded as having contrib-
 L6 uted to widespread desertification. Grazing peaked
 L7 around 1920 on public lands in the West. By the
 L8 1970s, grazing had been reduced by about 70
 L9 percent, but the arid lands have been very slow to
 L10 recover from its impacts. Warmer and drier climate
 L11 conditions are expected to slow recovery even
 L12 more. In addition, the land resource in the South-
 L13 west is currently managed more for providing water
 L14 for people than for protecting the productivity of the
 L15 landscape. As a result, the land resource is likely to
 L16 be further degraded and its recovery hampered.²⁴⁴

L19 **Coastal and near-coastal ecosystems are**
 L20 **already under multiple stresses. Climate**
 L21 **change and ocean acidification will**
 L22 **exacerbate these stresses.**

L24 Coastal and near-shore marine ecosystems are vul-
 L25 nerable to a host of climate change-related effects,
 L26 including increasing air and water temperatures,
 L27 ocean acidification, changes in runoff from the
 L28 land, sea-level rise, and altered currents. Some of
 L29 these changes have already led to coral bleaching,
 L30 shifts in species ranges, increased storm intensity in
 L31 some regions, dramatic reductions in sea ice extent
 L32 and thickness along the Alaskan coast,¹³⁷ and other
 L33 significant changes to the nation’s coastlines and
 L34 marine ecosystems.⁷⁰

L36 The interface between land and sea is important,
 L37 as many species depend on it at some point in their
 L38 life cycle, including many endangered species. In
 L39 addition, coastal areas buffer inland areas from
 L40 the effects of wave action and storms.²⁴⁸ Coastal
 L41 wetlands, intertidal areas, and other near-shore
 L42 ecosystems are subject to a variety of environmen-
 L43 tal stresses.^{255,256} Sea-level rise, increased coastal
 L44 storm intensity, and rising temperatures contrib-
 L45 ute to increased vulnerability of coastal wetland
 L46 ecosystems. It has been estimated that 3 feet of
 L47 sea-level rise (within the range of projections for
 L48 this century) would inundate 65 percent of the
 L49 coastal marshlands and swamps in the contiguous
 L50 United States.²⁵⁷ The combination of sea-level rise,

local land sinking, and related factors already have
 resulted in substantially higher relative sea-level
 rise along the Gulf of Mexico and the Southeast
 Atlantic coast, more so than farther north on the
 Atlantic Coast or on the Pacific Coast.^{43,255} In
 Louisiana alone, over one-third of the coastal plain
 that existed a century ago has since been lost,²⁵⁵
 which is mostly due to local land sinking.⁷⁰ Barrier
 islands also are being lost at an increasing rate²⁵⁸
 (see *Southeast* region), and they are particularly im-
 portant in protecting the coastline in some regions
 vulnerable to sea-level rise and storm surge.

Coral reefs

Coral reefs are very diverse ecosystems that sup-
 port many other species by providing food and
 habitat. In addition to their ecological value, coral
 reefs provide billions of dollars in services includ-
 ing tourism, fish breeding habitat, and protection
 of coastlines. In addition to climate change-related
 stresses, corals in many places face a host of other
 challenges associated with human activities such as
 poorly regulated tourism, destructive fishing, and
 pollution.⁷⁰

Corals are marine animals that host symbiotic algae
 that help nourish the animals and give the corals
 their color. When corals are stressed by increases
 in water temperatures or ultraviolet light, they lose
 their algae and turn white, a process called coral
 bleaching. If the stress persists, the corals die.
 Intensities and frequencies of bleaching events,
 clearly driven by warming in surface water, have
 increased substan-tially over the past 30 years,
 leading to the death or severe damage of about one-
 third of the world’s corals.⁷⁰

The United States has extensive coral reef eco-
 systems in the Caribbean, Atlantic, and Pacific
 oceans. In 2005, the Caribbean Basin experienced
 unprecedented water temperatures that resulted
 in dramatic coral bleaching with some sites in the
 U.S. Virgin Islands seeing 90 percent of the coral
 bleached. Some corals began to recover when water
 temperatures decreased, but later that year disease
 appeared, striking the previously bleached and
 weakened coral. To date, 50 percent of the corals
 in Virgin Islands National Park have died from the
 bleaching and disease events. In the Florida Keys,
 summer bleaching in 2005 was also followed by
 disease in September.⁷⁰



L1 But rising temperature is not the only stress coral
L2 reefs face. As the carbon dioxide concentration in
L3 the air increases, more carbon dioxide is absorbed
L4 into the world's oceans, leading to their acidifica-
L5 tion. This makes less calcium carbonate available
L6 for corals and other sea life to build their skeletons
L7 and shells.²⁵⁹ If carbon dioxide concentrations
L8 continue to rise and the resulting acidification pro-
L9 ceeds, eventually, corals and other ocean life that
L10 rely on calcium carbonate will not be able to build
L11 these skeletons and shells at all. The implications
L12 of such extreme changes in ocean ecosystems are
L13 not clear, but there is now evidence that in some
L14 ocean areas, such as along the Northwest coast,
L15 acidification is already occurring^{70,260} (see *Coasts*
L16 region for a more in-depth discussion about ocean
L17 acidification).

L19
L20 **Arctic sea ice ecosystems are already**
L21 **being adversely affected by the loss of**
L22 **summer sea ice and further changes are**
L23 **expected.**
L24

L25 Perhaps most vulnerable of all to the impacts of
L26 warming are Arctic ecosystems that rely on sea ice,
L27 which is vanishing rapidly and is projected to dis-
L28 appear entirely in summertime within this century.
L29 Algae that bloom on the underside of the sea ice
L30 form the base of a food web linking microscopic
L31 animals and fish to seals, whales, polar bears, and
L32 people. As the sea ice disappears, so too do these
L33 algae. The ice also provides a vital platform for
L34 ice-dependent seals (such as the ringed seal) to give
L35 birth, nurse their pups, and rest. Polar bears use the
L36 ice as a platform from which to hunt their prey. The
L37 walrus rests on the ice near the continental shelf
L38 between its dives to eat clams and other shellfish.
L39 As the ice edge retreats away from the shelves to
L40 deeper areas, there will be no clams nearby.^{70,132}

L41
L42 The Bering Sea, off the west coast of Alaska,
L43 produces our nation's largest commercial fish
L44 harvests as well as providing food for many Native
L45 Alaskan peoples. Ultimately, the fish populations
L46 (and animals including seabirds, seals, walruses,
L47 and whales) depend on plankton blooms regulated
L48 by the extent and location of the ice edge in spring.
L49 As the sea ice continues to decline, the location,
L50 timing, and species composition of the blooms is

changing. The spring melt of sea ice in the Ber-
ing Sea has long provided material that feeds the
clams, shrimp, and other life forms on the ocean
floor that, in turn, provide food for the walruses,
gray whales, bearded seals, eider ducks, and many
fish. The earlier ice melt resulting from warming,
however, leads to later phytoplankton blooms that
are largely consumed by microscopic animals near
the sea surface, vastly decreasing the amount of
food reaching the living things on the ocean floor.
This will radically change the makeup of the fish
and other creatures, with significant repercussions
for both subsistence and commercial fishing.⁷⁰

Ringed seals give birth in snow caves on the sea
ice, which protect their pups from extreme cold
and predators. Warming leads to earlier snow melt,
which causes the snow caves to collapse before the
pups are weaned. The small, exposed pups might
die of hypothermia or be vulnerable to predation by
arctic foxes, polar bears, gulls, and ravens. Gulls
and ravens are arriving in the Arctic earlier as
springs become warmer, increasing the birds' op-
portunity to prey on the seal pups.⁷⁰

Polar bears are the top predators of the sea ice
ecosystem. Because they prey primarily on ice-
associated seals, they are especially vulnerable to
the disappearance of sea ice. The bears' ability to
catch seals depends on the presence of sea ice. In
that habitat, polar bears take advantage of the fact
that seals must surface to breathe in limited open-
ings in the ice cover. In the open ocean, bears lack
a hunting platform, seals are not restricted in where
they can surface, and successful hunting is very
rare. On shore, polar bears feed little, if at all.



About two-thirds of the world's polar bears are projected to
be gone by the middle of this century. It is projected that there
will be no wild polar bears in Alaska in 75 years.⁷⁰

R1
R2
R3
R4
R5
R6
R7
R8
R9
R10
R11
R12
R13
R14
R15
R16
R17
R18
R19
R20
R21
R22
R23
R24
R25
R26
R27
R28
R29
R30
R31
R32
R33
R34
R35
R36
R37
R38
R39
R40
R41
R42
R43
R44
R45
R46
R47
R48
R49
R50

L1
L2
L3
L4
L5
L6
L7
L8
L9
L10
L11
L12
L13
L14
L15
L16
L17
L18
L19
L20
L21
L22
L23
L24
L25
L26
L27
L28
L29
L30
L31
L32
L33
L34
L35
L36
L37
L38
L39
L40
L41
L42
L43
L44
L45
L46
L47
L48
L49
L50

In addition, the rapid rate of warming in Alaska and the rest of the Arctic in recent decades is sharply reducing the snow cover in which polar bears build dens and the sea ice they use as foraging habitat. Female polar bears build snow dens in which they hibernate for four to five months each year and in which they give birth to their cubs. Born weighing only about 1 pound, the tiny cubs depend on the snow den for warmth.

About two-thirds of the world’s polar bears are projected to be gone by the middle of this century. It is projected that there will be no wild polar bears left in Alaska in 75 years.⁷⁰

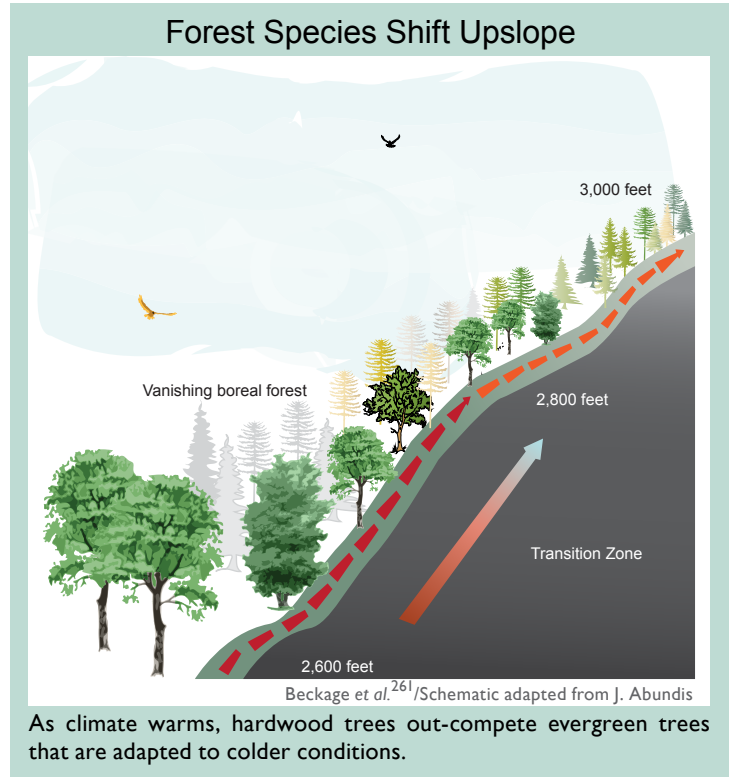
Continued warming will inevitably entail major changes in the sea-ice ecosystem, to the point that its viability is in jeopardy. Some species will become extinct, while others might adapt to new habitats. The chances of species surviving the current changes might depend critically on the rate of change. The current rates of change in the sea-ice ecosystem are very rapid relative to the life spans of animals including seals, walruses, and polar bears, and as such, are a major threat to their survival.⁷⁰



The pika, pictured above, is a small mammal whose habitat is limited to cold areas near the tops of mountains. As climate warms, little suitable habitat is left. Of 25 pika populations studied in the Great Basin between the Rocky Mountains and the Sierra Nevada, more than one-third have gone extinct in recent decades.^{262,263}

The habitats of some mountain species and coldwater fish, such as salmon and trout, are very likely to contract in response to warming.

Animal and plant species that live in the mountains are among those particularly sensitive to rapid climate change. They include animal species such as the grizzly bear, bighorn sheep, pika, mountain goat, and wolverine. Major changes already have been observed in the pika as previously reported populations have disappeared entirely as climate has warmed over recent decades.⁷⁰ One reason mountain species are so vulner-



able is that their suitable habitats are being compressed as climatic zones shift upward in elevation. Some species try to shift uphill with the changing climate, but there might be other constraints related to food, other species present, and other variables. In addition, as species move up the mountains, those near the top simply run out of habitat.⁷⁰

Fewer wildflowers are projected to grace the slopes of the Rocky Mountains as global warming causes earlier spring snowmelt. Larkspur, aspen fleabane, and aspen sunflower grow at an altitude of about 9,500 feet where the winter snows are deep. Once the snow melts, the flowers form buds and prepare to bloom. But warmer springs mean that the snow melts earlier, leaving the buds exposed to frost. (The percentage of buds that were frosted has doubled over the past decade.) Frost does not kill the plants, but it does make them unable to seed and reproduce, meaning there will be no next generation. Insects and other animal species depend on the flowers for food, and other species depend on those species, so the loss is likely to propagate through the food chain.²³⁶

Shifts in tree species on mountains in New England, where temperatures have risen 2 to 4°F in the last 40 years, offer another example. Some

R1
R2
R3
R4
R5
R6
R7
R8
R9
R10
R11
R12
R13
R14
R15
R16
R17
R18
R19
R20
R21
R22
R23
R24
R25
R26
R27
R28
R29
R30
R31
R32
R33
R34
R35
R36
R37
R38
R39
R40
R41
R42
R43
R44
R45
R46
R47
R48
R49
R50

L1 mountain tree species have shifted uphill by 350
 L2 feet in the last 40 years. Tree communities were
 L3 relatively unchanged at low and high elevations, but
 L4 in the transition zone in between (at about 2,600
 L5 feet elevation) the changes have been dramatic.
 L6 Cold-loving tree species declined from 43 to 18
 L7 percent, while warmer-loving trees increase from
 L8 57 to 82 percent. Overall, the transition zone has
 L9 shifted about 350 feet uphill in just a few decades,
 L10 a surprisingly rapid rate since these are trees that
 L11 live for hundreds of years. One possibility is that as
 L12 trees were damaged or killed by air pollution, it left
 L13 an opportunity for the warming-induced transition
 L14 to occur more quickly. These results indicate that
 L15 the composition of high-elevation forests is chang-
 L16 ing rapidly.²⁶¹

L18 **Coldwater fish**

L19 Salmon and other coldwater fish species in the
 L20 United States are at particular risk from warm-
 L21 ing. Salmon are under threat from a variety of
 L22 human activities, but global warming is a grow-
 L23 ing source of stress. Rising temperatures impact
 L24 salmon in several important ways. As precipitation
 L25 increasingly falls as rain rather than snow, it feeds
 L26 floods that wash away salmon eggs incubating in
 L27 the streambed. Warmer water leads eggs to hatch
 L28 earlier in the year, so the young are smaller and
 L29 more vulnerable to predators. Warmer conditions
 L30 increase the fish's metabolism, taking energy away
 L31 from growth and forcing the fish to find more food,
 L32 but earlier hatching of eggs could put them out of
 L33 sync with the insects they eat. Earlier melting of
 L34 snow leaves rivers and streams warmer and shall-
 L35 lower in summer and fall. Diseases and parasites
 L36 tend to flourish in warmer water. Studies suggest
 L37 that up to 40 percent of Northwest salmon popula-
 L38 tions might be lost by 2050.²⁶⁴

L40 Large declines in trout populations also are pro-
 L41 jected to occur around the United States. Over half
 L42 of the wild trout populations are likely to disappear
 L43 from the southern Appalachian Mountains because
 L44 of the effects of warming stream temperatures.
 L45 Losses of western trout populations might exceed
 L46 60 percent in certain regions. About 90 percent of
 L47 bull trout, which live in western rivers in some of
 L48 the country's most wild places, are projected to be
 L49 lost due to warming. Pennsylvania is predicted to
 L50 lose 50 percent of its trout habitat in the coming

decades. Projected losses of trout habitat for some
 warmer states, such as North Carolina and Virgin-
 ia, are up to 90 percent.²⁶⁵

Some of the benefits ecosystems provide to society will be threatened by climate change, while others will be enhanced.

Human well-being depends on the Earth's ecosys-
 tems and the services that they provide to sustain
 and fulfill human life.²⁶⁶ These services are impor-
 tant to human well-being because they contribute
 to basic material needs, physical and psychological
 health, security, and economic activity. A recent
 assessment reported that of 24 vital ecosystem ser-
 vices, 15 were being degraded by human activity.²⁴⁸
 Climate change is one of several human-induced
 stresses that threaten to intensify and extend these
 adverse impacts to biodiversity, ecosystems, and
 the services they provide. Two of many possible
 examples follow.

Forests and carbon storage

Forests provide many services important to the
 well-being of Americans: air and water quality
 maintenance, water flow regulation, and watershed
 protection; wildlife habitat and biodiversity conser-
 vation; recreational opportunities and aesthetic and
 spiritual fulfillment; raw materials for wood and
 paper products; and climate regulation and carbon
 storage. A changing climate will alter forests and
 the services they provide. Most of these changes
 are likely to be detrimental.

In the United States, forest growth and long-lived
 forest products currently offset about 20 percent of
 U.S. fossil fuel carbon emissions.^{140,258} This carbon
 "sink" is an enormous service provided by forests
 and its persistence or growth will be important to
 limiting the atmospheric carbon dioxide concentra-
 tion. The scale of the challenge of increasing this
 sink is very large. To offset an additional 10 percent
 of the U.S. emissions through tree planting would
 require converting one-third of current croplands to
 forests.²⁴⁴

Recreational opportunities

Tourism is one of the largest economic sectors in the world, and it is also one of the fastest growing;²⁶⁷ the jobs created by recreational tourism provide economic benefits not only to individuals but also to communities. Slightly more than 90 percent of the U.S. population participates in some form of outdoor recreation, representing nearly 270 million participants,²⁶⁸ and several billion days spent each year in a wide variety of outdoor recreation activities.

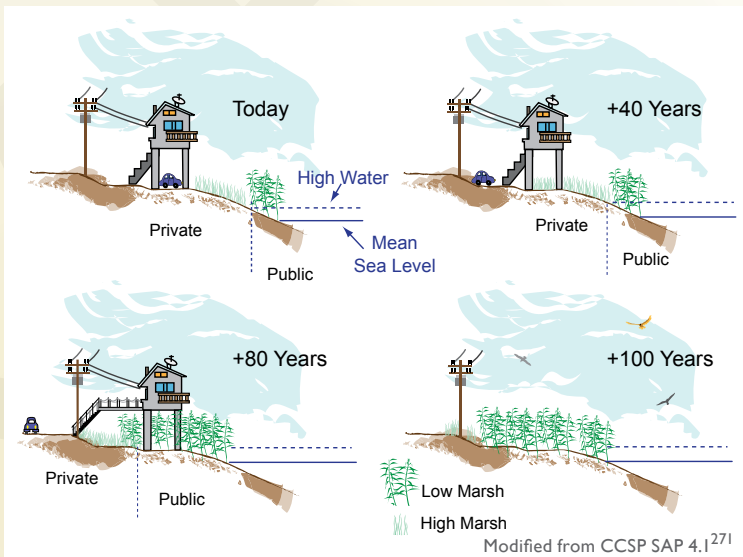
Since much recreation and tourism occurs outside, increased temperature and precipitation have a direct effect on the enjoyment of these activities, and on the desired number of visitor days and associated level of visitor spending as well as tourism employment. Weather conditions are an important factor influencing tourism visits. In addition, outdoor recreation and tourism often depends on the availability and quality of natural resources,²⁶⁹ such

as beaches, forests, wetlands, snow, and wildlife, all of which will be affected by climate change. Thus, climate change can have direct effects on the natural resources that people enjoy. The length of the season for, and desirability of, several of the most popular activities – walking; visiting a beach, lakeshore, or river; sightseeing; swimming; and picnicking²⁶⁸ – are likely to be enhanced by small near-term increases in temperature. Other activities are likely to be harmed by even small increases in global warming, such as snow- and ice-dependent activities including skiing, snowmobiling, and ice fishing.

The net economic effect of near-term climate change on recreational activities is likely to be positive. In the longer term, however, as climate change effects on ecosystems and seasonality become more pronounced, the net economic effect on tourism and recreation is not known with certainty.¹⁷²

Adaptation: Preserving Coastal Wetlands

Coastal wetlands are rich ecosystems that protect the shore from damage during storm surges and provide society with other services. One strategy designed to preserve coastal wetlands as sea level rises is the “rolling easement.” Rolling easements allow some development near the shore, but prohibit construction of seawalls or other armoring to protect buildings; they recognize nature’s right-of-way to advance inland as sea level rises. Massachusetts and Rhode Island prohibit shoreline armoring along the shores of some estuaries so that ecosystems can migrate inland, and several states limit armoring along ocean shores.²⁷⁰



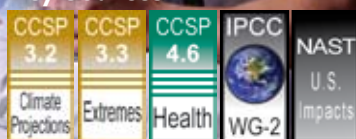
In the case shown here, the coastal marsh would reach the footprint of the house 40 years in the future. Because the house is on pilings, it could still be occupied if it is connected to a community sewage treatment system; a septic system would probably fail due to proximity to the water table. After 80 years, the marsh would have taken over the yard, and the footprint of the house would extend onto public property. The house could still be occupied but reinvestment in the property would be unlikely. After 100 years, this house would be removed, although some other houses in the area could still be occupied. Eventually, the entire area would return to nature. A home with a rolling easement would depreciate in value rather than appreciate like other coastal real estate. But if the loss were expected to occur 100 years from now, it would only reduce the current property value by 1 to 5 percent, for which the owner could be compensated.²⁷¹

Human Health

Key Messages:

- Significant increases in the risk of illness and death related to extreme heat and heat waves are very likely.
- Climate change is likely to contribute to poor air quality, adversely affecting health.
- Extreme weather events cause physical and mental health problems. Some of these events are projected to increase.
- Some diseases transmitted by food, water, and insects are likely to increase.
- Rising temperature and carbon dioxide concentration increase pollen production and prolong the pollen season in a number of plants with highly allergenic pollen, presenting a health risk.
- Certain groups, including children, the elderly, and the poor, are most vulnerable to a range of climate-related health effects.

Key Sources

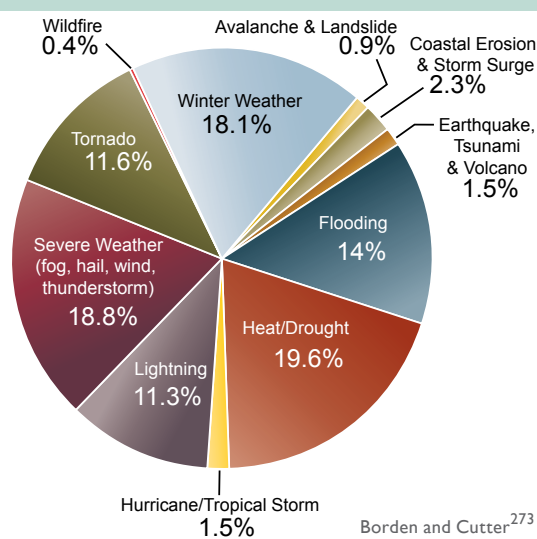


Climate change poses unique challenges to human health. Unlike health threats caused by a particular toxin or disease pathogen, there are many ways that climate change can lead to potentially harmful health effects. There are direct health impacts from heat waves and severe storms, ailments caused or exacerbated by air pollution and airborne allergens, and many climate-sensitive infectious diseases.¹⁶³

Realistically assessing the potential health effects of climate change must include consideration of the capacity to manage new and changing climatic conditions.¹⁶³ Whether or not increased health risks due to climate change are realized will depend largely on societal responses and underlying vulnerability. The probability of exacerbated health risks due to climate change points to a need to maintain a strong public health infrastructure to help limit future impacts.¹⁶³

Increased risks associated with diseases originating outside the United States must also be considered because we live in an increasingly globalized world. Many poor nations are expected to suffer even greater health consequences from climate change.²⁷² With global trade and travel, disease flare-ups in any part of the world can potentially reach the United States. In addition, weather and climate extremes such as severe storms and drought can undermine public health infrastructure, further stress environmental resources, destabilize economies, and potentially create security risks both within the United States and internationally.²¹⁹

Hazard-Related Deaths in the U.S.



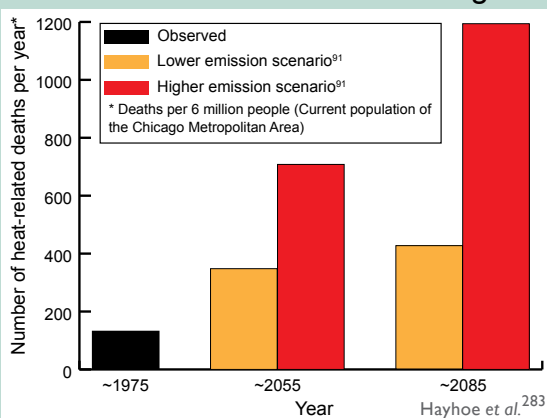
The pie chart shows the distribution of deaths for 11 hazard categories as a percent of the total 19,958 deaths due to these hazards from 1970 to 2004. Heat/drought ranks highest, followed by severe weather, which includes events with multiple causes such as lightning, wind, and rain.²⁷³ This analysis ended prior to the 2005 hurricane season which resulted in approximately 2,000 deaths.²²⁹

Significant increases in the risk of illness and death related to extreme heat and heat waves are very likely.

Temperatures are rising and the probability of severe heat waves is increasing. Analyses suggest that currently rare extreme heat waves will become much more common in the future.⁶⁸ At the same time, the U.S. population is aging, and older people are more vulnerable to hot weather and heat waves. The percentage of the U.S. population over age 65 is currently 12 percent and is projected to be 21 percent by 2050 (over 86 million people).^{163,274} Diabetics are also at greater risk of heat-related death, and the prevalence of obesity and diabetes is increasing. Heat-related illnesses range from heat exhaustion to kidney stones.^{275,276}

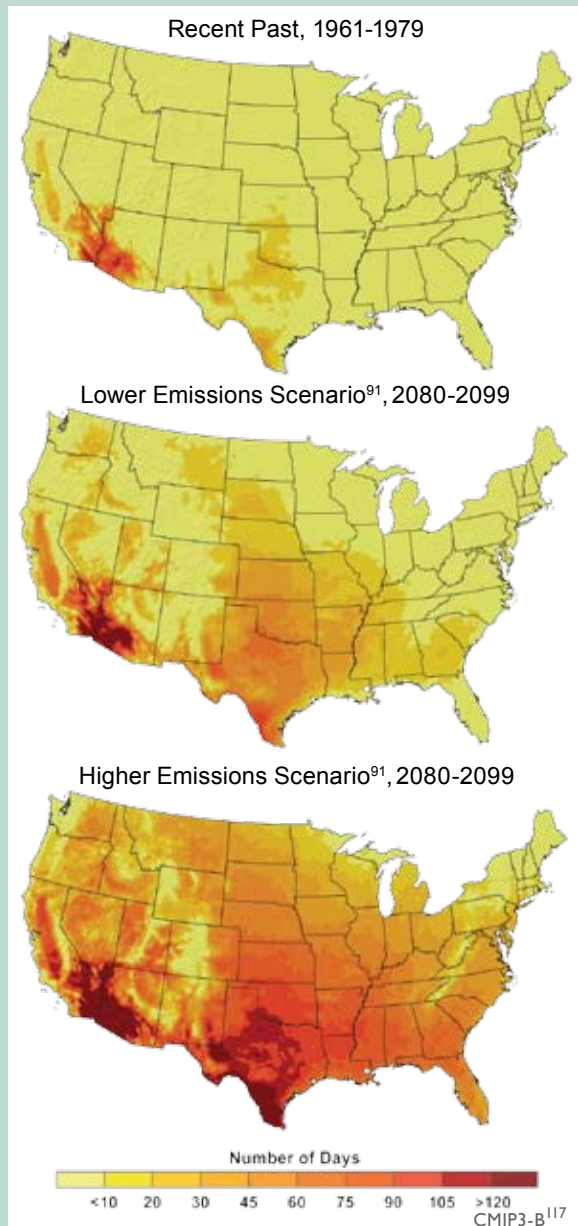
Heat is already the leading cause of weather-related deaths in the United States. More than 3,400 deaths between 1999 and 2003 were reported as resulting from exposure to excessive heat.²⁷⁷ An analysis of nine U.S. cities shows that deaths due to heat increase with rising temperature and humidity.²⁷⁸ From the 1970s to the 1990s, however, heat-related deaths declined.²⁷⁹ This likely resulted from a rapid increase in the use of air conditioning. In 1978, 44 percent of households were without air conditioning, whereas in 2005, only 16 percent of the U.S. population lived without it (and only 3 percent did

Projected Increase in Heat-Related Deaths in Chicago



Increases in heat-related deaths are projected in cities around the nation, especially under higher emissions scenarios.⁹¹ This analysis included some adaptation measures. The graph shows the projected number of deaths per year, averaged over a three-decade period around 1975, 2055, and 2085 for the City of Chicago under lower and higher emissions.⁹¹

Number of Days Over 100°F



The number of days in which the temperature exceeds 100°F by late this century, compared to the 1960s and 1970s, is projected to increase strongly across the United States. For example, an area in Texas that currently experiences about 10 days per year over 100°F is expected to experience more than 100 days per year in which the temperature exceeds 100°F by the end of the century under the higher emissions scenario.⁹¹

not have it in the South).^{280,281} With air conditioning reaching near saturation, a recent study found that the general decline in heat-related deaths seem to have leveled off since the mid-1990s.²⁸²

As human-induced warming is projected to raise average temperatures by about 6 to 11°F in this century under a higher emissions scenario,⁹¹ heat

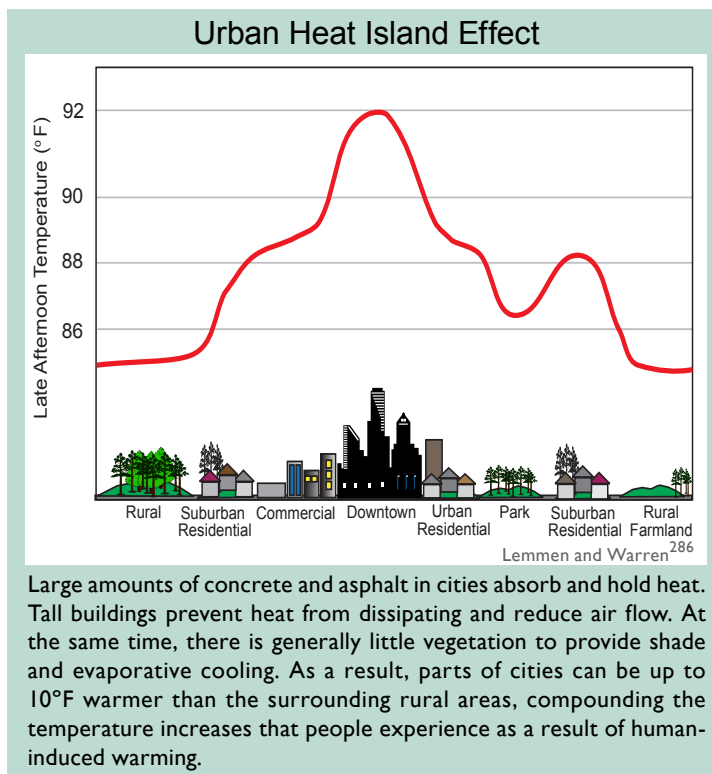
waves are expected to continue to increase in frequency, severity, and duration.^{68,112} For example, by the end of this century, the number of heat-wave days in Los Angeles is projected to double,²⁸⁴ and the number in Chicago to quadruple,²⁸⁵ if emissions are not reduced.

Projections for Chicago suggest that the average number of deaths due to heat waves would more than double by 2050 under a lower emissions scenario⁹¹ and quadruple under a high emissions scenario⁹¹ (see figure).²⁸³

A study of climate change impacts in California projects that, by the 2090s, annual heat-related deaths in Los Angeles would increase by two to three times under a lower emissions scenario and by five to seven times under a higher emissions scenario, compared to a 1990s baseline of about 165 deaths.

These estimates assume that people will have become somewhat more accustomed to higher temperatures. Without such acclimatization, these estimates are projected to be about 20 to 25 percent higher.²⁸⁴

The full effect of global warming on heat-related illness and death involves a number of factors including actual changes in temperature (averages, highs, and lows); and human population characteristics, such as age, wealth, and fitness. In addition, adaptation at the scale of a city includes options such as heat wave early warning systems, urban design to reduce heat loads, and enhanced services during heat waves.¹⁶³



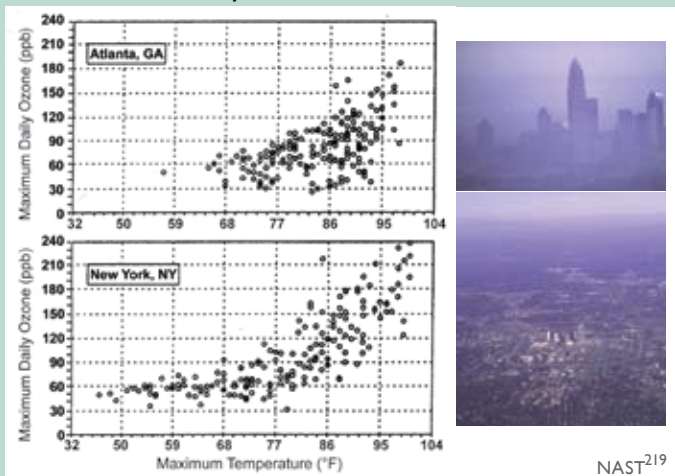
Reduced extreme cold

In a warmer world, the number of deaths caused by extremely low temperatures would be expected to drop, although in general, it is uncertain how climate change will affect net mortality.¹⁶³ Nevertheless, a recent study that analyzed daily mortality and weather data with regard to 6,513,330 deaths in 50 U.S. cities between 1989 and 2000 shows a marked difference between deaths resulting from hot and cold temperatures. The researchers found that, on average, cold snaps increased death rates by 1.6 percent, while heat waves triggered a 5.7 percent increase in death rates.²⁸⁹ The study concluded

Adaptation: Reducing Deaths During Heat Waves

In the mid-1990s, Philadelphia became the first U.S. city to implement a system for reducing the risk of death during heat waves. The city focuses its efforts on the elderly, homeless, and poor. During a heat wave, a heat alert is issued and news organizations are provided with tips on how vulnerable people can protect themselves. The health department and thousands of block captains use a buddy system to check on elderly residents in their homes; electric utilities voluntarily refrain from shutting off services for non-payment; and public cooling places extend their hours. The city operates a “Heatline” where nurses are standing by to assist callers experiencing health problems; if callers are deemed “at risk,” mobile units are dispatched to the residence. The city also has implemented a “Cool Homes Program” for elderly low-income residents, which provides measures such as roof coatings and roof insulation that save energy and lower indoor temperatures. Philadelphia’s system is estimated to have saved 117 lives over its first 3 years of operation.^{287,288}

Temperature and Ozone



The graphs illustrate the observed association between ground-level ozone (a component of smog) concentration and temperature in Atlanta and New York City (May to October 1988 to 1990) in parts per billion (ppb).²¹⁹ The projected higher temperatures across the United States in this century are likely to increase the occurrence of high ozone concentrations, although this will also depend on emissions of ozone precursors and meteorological factors. Ground-level ozone can exacerbate respiratory diseases and cause short-term reductions in lung function.

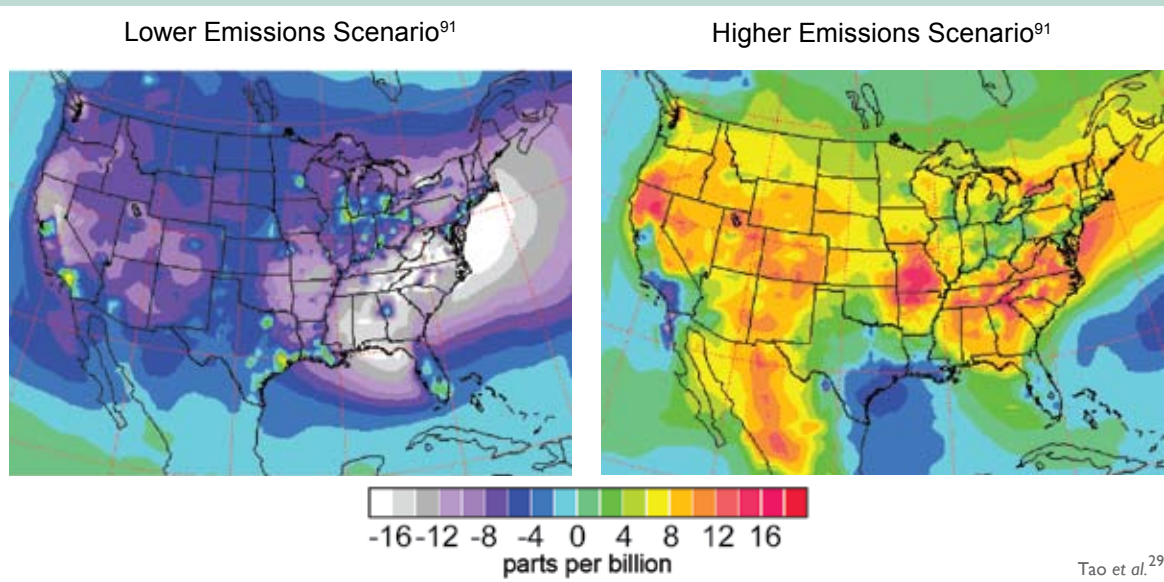
that the reduction in deaths as a result of relatively milder winters attributable to global warming will not make up for the more severe health effects of summertime heat extremes.²⁸⁹

It has been suggested that because death rates are higher in winter than in summer, warming might decrease deaths overall, but this ignores the fact that influenza and pneumonia cause many winter deaths, and it is unclear how these highly seasonal diseases are affected by temperature.¹⁶³

Climate change is likely to contribute to poor air quality, adversely affecting health.

Poor air quality, especially in cities, is a serious concern across the United States. Half of all Americans, 158 million people, live in counties where air pollution exceeds national health standards.²⁹⁰ While the Clean Air Act has improved air quality, higher temperatures and associated stagnant air masses are expected to make it more challenging to meet air quality standards, particularly for ground-level ozone (a component of smog).¹³ It

Projected Change in Ground-Level Ozone, 2090s



The maps show projected changes in ground-level ozone (a component of smog) for the 2090s, averaged over the summer months (June through August), relative to 1996-2000, under lower and higher emissions scenarios, which include both greenhouse gases and emissions that lead to ozone formation (some of which decrease under the lower emissions scenario).⁹¹ By themselves, higher temperatures and other projected climate changes would increase ozone levels under both scenarios. However, the maps indicate that future projections of ozone depend heavily on emissions, with the higher emissions scenario⁹¹ increasing ozone by large amounts, while the lower emissions scenario⁹¹ results in an overall decrease in ground-level ozone by the end of the century.²⁹¹

has been firmly established that breathing ozone results in short-term decreases in lung function and damages the cells lining the lungs. It also increases the incidence of asthma-related hospital visits and premature deaths.²⁷² Vulnerability to ozone effects is greater for those who spend time outdoors, especially with physical exertion, because this results

in a higher cumulative dose to their lungs. As a result, children, outdoor workers, and athletes are at higher risk for these ailments.¹⁶³

Ground-level ozone concentrations are affected by many factors including weather conditions, emissions of gases from vehicles and industries that lead

Spotlight on Air Quality in California



Californians currently experience the worst air quality in the nation. More than 90 percent of the population lives in areas that violate air quality standards for ground-level ozone or small particles. These pollutants cause an estimated 8,800 deaths and over a billion dollars in health care costs every year in California.²⁹² Higher temperatures are projected to increase the frequency, intensity, and duration of conditions conducive to air pollution formation, potentially increasing the number of days conducive to air pollution by 75 to 85 percent in Los Angeles and the San Joaquin Valley, toward the end of this century, under a higher emissions scenario, and by 25 to 35 percent under a lower emissions scenario.^{91,293} Air quality could be further compromised by wildfires, which are already increasing as a result of warming.²⁹⁴

Adaptation: Improving Urban Air Quality

Because poor air quality is related to temperature, it is projected to become worse with human-induced climate change. Many areas in the country already have plans in place for responding to air quality problems. For example, the Air Quality Alert program in Rhode Island encourages residents to reduce air pollutant emissions by limiting car travel and the use of small engines, lawn mowers, and charcoal lighter fluids on days when ground-level ozone is high. Television weather reports include alerts when ground-level ozone is high, warning especially susceptible people to limit their time outdoors. To help cut down on the use of cars, all regular bus routes are free on Air Quality Alert days.²⁹⁵

Pennsylvania offers the following suggestions for high ozone days:

- Refuel vehicles after dark. Avoid spilling gasoline and stop fueling when the pump shuts off automatically.
- Conserve energy. Do not overcool homes. Turn off lights and appliances that are not in use. Wash clothes and dishes only in full loads.
- Limit daytime driving. Consider carpooling or taking public transportation. Properly maintain vehicles, which also helps to save fuel.
- Limit outdoor activities, such as mowing the lawn or playing sports, to the evening hours.
- Avoid burning leaves, trash, and other materials.

Traffic restrictions imposed during the 1996 summer Olympics in Atlanta quantified the direct respiratory health benefits of reducing the number of cars and the amount of their tailpipe emissions from an urban environment. Peak morning traffic decreased by 23 percent, and peak ozone levels dropped by 28 percent. As a result, childhood asthma-related emergency room visits fell by 42 percent.²⁹⁶

L1 to ozone formation (especially nitrogen oxides and
 L2 volatile organic compounds), natural emissions of
 L3 volatile organic compounds from plants, and pol-
 L4 lution blown in from other places.^{290,297} A warmer
 L5 climate is projected to increase the natural emis-
 L6 sions of volatile organic compounds, accelerate
 L7 ozone formation, and increase the frequency and
 L8 duration of stagnant air masses that allow pollution
 L9 to accumulate, which will exacerbate health symp-
 L10 toms.²⁹⁸ Increased temperatures and water vapor
 L11 due to human-induced carbon dioxide emissions
 L12 have been found to increase ozone more in areas
 L13 with already elevated concentrations, meaning that
 L14 global warming tends to exacerbate ozone pollu-
 L15 tion most in already polluted areas. Under constant
 L16 pollutant emissions, by the middle of this century,
 L17 Red Ozone Alert Days (when the air is unhealthy
 L18 for everyone) in the 50 largest cities in the east-
 L19 ern United States are projected to increase by 68
 L20 percent due to warming alone.²⁹⁸ Such conditions
 L21 would challenge the ability of communities to meet
 L22 health-based air quality standards such as those in
 L23 the Clean Air Act.

L24
 L25 Health risks from heat waves and risks from exac-
 L26 erbated air pollution are not necessarily indepen-
 L27 dent. The formation of ground-level ozone occurs
 L28 under hot and stagnant conditions – essentially
 L29 the same weather conditions accompanying heat
 L30 waves. Such interactions among risk factors are
 L31 likely to increase as climate change continues.

L32
 L33
 L34 **Extreme weather events cause physical
 L35 and mental health problems. Some of
 L36 these events are projected to increase.**
 L37

L38 Injury, illness, emotional trauma, and death are
 L39 known to result from extreme weather events.⁶⁸
 L40 The number and intensity of some of these events
 L41 are already increasing and are projected to increase
 L42 further in the future.^{68,112} Human health impacts in
 L43 the United States are generally expected to be less
 L44 severe than in poorer countries where the emer-
 L45 gency preparedness and public health infrastruc-
 L46 ture is less developed. For example, early warning
 L47 and evacuation systems and effective sanitation
 L48 lessen the health impacts of extreme events.⁶⁸ This
 L49 assumes that medical and emergency relief systems
 L50 in the United States will function well and that

timely and effective adaptation measures will be R1
 developed and deployed. There have already been R2
 serious failures of these systems in the aftermath of R3
 hurricanes Katrina and Rita, so coping with future R4
 impacts will require significant improvements. R5
 R6

R7 **Extreme storms**

R8 Over 2,000 Americans were killed in the 2005 R8
 hurricane season, more than double the average R9
 number of lives lost to hurricanes in the United R10
 States over the previous 65 years.¹⁶³ But the human R11
 health impacts of extreme storms go beyond direct R12
 injury and death to indirect effects such as carbon R13
 monoxide poisoning from portable electric genera- R14
 tors in use following hurricanes, an increase in R15
 stomach and intestinal illness among evacuees, and R16
 mental health impacts such as depression and post- R17
 traumatic stress disorder.¹⁶³ Failure to fully account R18
 for both direct and indirect health impacts might R19
 result in inadequate preparation for and response to R20
 future extreme weather events.¹⁶³ R21
 R22

R23 **Floods**

R24 Heavy downpours have increased in recent decades R24
 and are projected to increase further as the world R25
 continues to warm.^{68,112} In the United States, the R26
 amount of precipitation falling in the heaviest 1 R27
 percent of rain events increased by 20 percent in R28
 the past century, while total precipitation increased R29
 by 7 percent. Over the last century, there was a R30
 50 percent increase in the frequency of days with R31
 precipitation over 4 inches in the upper Midwest.¹¹² R32
 Other regions, notably the South, also have seen R33
 strong increases in heavy downpours, with most of R34
 these coming in the warm season and almost all of R35
 the increase coming in the last few decades. R36
 R37

R38 Heavy rains can lead to flooding, which can cause R38
 health impacts including direct injuries as well as R39
 increased incidence of waterborne diseases due to R40
 pathogens such as *Cryptosporidium* and *Giardia*.¹⁶³ R41
 Downpours can trigger sewage overflows, contami- R42
 nating drinking water and endangering beachgoers. R43
 The consequences will be particularly severe in the R44
 roughly 770 U.S. cities and towns, including New R45
 York, Chicago, Washington DC, Milwaukee, and R46
 Philadelphia, that have “combined sewer systems;” R47
 an older design that carries storm water and sew- R48
 age in the same pipes.²⁹⁹ During heavy rains, these R49
 systems often cannot handle the volume, and raw R50



L1 sewage spills into lakes or waterways, including
 L2 drinking-water supplies and places where people
 L3 swim.²⁵³
 L4

L5 In 1994, the EPA established a policy that mandates
 L6 that communities substantially reduce or eliminate
 L7 their combined sewer overflow, but this mandate
 L8 remains unfulfilled.³⁰⁰ In 2004, the EPA estimated
 L9 it would cost \$203 billion to address these and other
 L10 needs of publicly-owned wastewater treatment
 L11 systems.³⁰¹
 L12

L13 Using 2.5 inches of precipitation in one day as the
 L14 threshold for initiating a combined sewer overflow
 L15 event, the frequency of these events in Chicago is
 L16 expected to rise by 50 percent to 120 percent by the
 L17 end of this century,³⁰² posing further risks to drink-
 L18 ing and recreational water quality.
 L19
 L20
 L21

Wildfires

Wildfires in the United States are already increas-
 ing due to warming. In the West, there has been a
 nearly fourfold increase in large wildfires in recent
 decades, with greater fire frequency, longer fire du-
 rations, and longer wildfire seasons. This increase
 is strongly associated with increased spring and
 summer temperatures and earlier spring snowmelt,
 which have caused drying of soils and vegeta-
 tion.^{163,294} In addition to direct injuries and deaths
 due to burns, wildfires can cause eye and respira-
 tory illnesses due to fire-related air pollution.¹⁶³

Some diseases transmitted by food, water, and insects are likely to increase.

A number of important disease-causing agents
 (pathogens) commonly transmitted by food, water,
 or animals are susceptible to changes in replication,

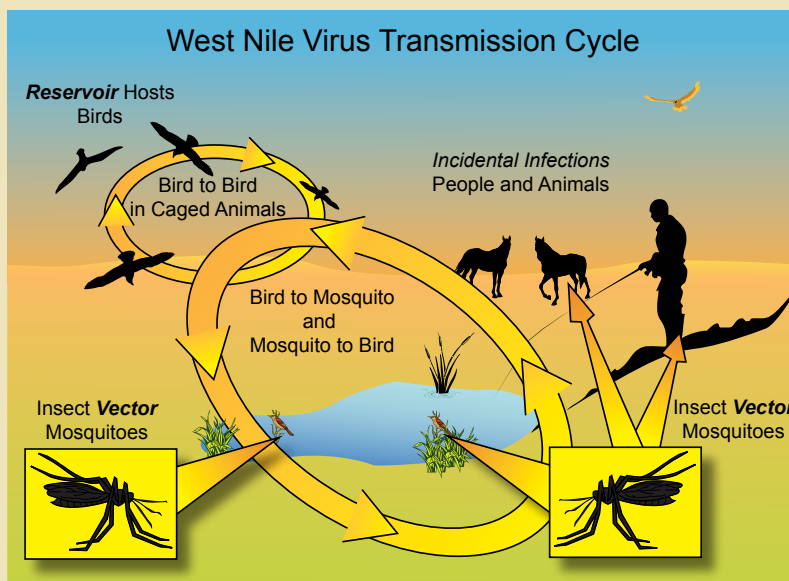
Spotlight on West Nile Virus



The first outbreak of West Nile virus in the United States occurred in the summer of 1999, likely a result of international air transport. Within five years, the disease had spread across the continental United States, transmitted by mosquitoes that acquire the virus from infected birds. While bird migrations were the primary mode of disease spread, during the epidemic summers of 2002 to 2004, epicenters of West Nile virus were linked to locations with either drought or above-average temperatures.

Since 1999, West Nile virus has caused over 28,000 reported cases, and over 1,100 Americans have died from it.³⁰³ During 2002, a more virulent strain of West Nile virus emerged in the United States. Recent analyses indicate that this mutated strain responds strongly to higher temperatures, suggesting that greater risks from the disease may result from increases in the frequency of heatwaves,³⁰⁴ though the risk will also depend on the effectiveness of mosquito control programs.

While West Nile virus causes mild flu-like symptoms in most people, about one in 150 infected people develop serious illness, including the brain inflammation diseases encephalitis and meningitis.



R1
 R2
 R3
 R4
 R5
 R6
 R7
 R8
 R9
 R10
 R11
 R12
 R13
 R14
 R15
 R16
 R17
 R18
 R19
 R20
 R21
 R22
 R23
 R24
 R25
 R26
 R27
 R28
 R29
 R30
 R31
 R32
 R33
 R34
 R35
 R36
 R37
 R38
 R39
 R40
 R41
 R42
 R43
 R44
 R45
 R46
 R47
 R48
 R49
 R50

survival, persistence, habitat range, and transmission as a result of changing climatic conditions such as increasing temperature, precipitation, and extreme weather events.¹⁶³

- Cases of food poisoning due to *Salmonella* and other bacteria peak within one to six weeks of the highest reported ambient temperatures.¹⁶³
- Cases of waterborne *Cryptosporidium* and *Giardia* increase following heavy downpours. These parasites can be transmitted in drinking water and through recreational water use.¹⁶³
- Climate change affects the life cycle and distribution of the mosquitoes, ticks, and rodents that carry West Nile virus, equine encephalitis, Lyme disease, and *Hantavirus*. However, moderating factors such as housing quality, land use patterns, pest control programs, and a robust public health infrastructure are likely to prevent the large-scale spread of these diseases in the United States.^{163,305}
- Heavy rain and flooding can contaminate certain food crops with feces from nearby livestock or wild animals, increasing the likelihood of food-borne disease associated with fresh produce.¹⁶³
- *Vibrio* sp. (shellfish poisoning) accounts for 20 percent of the illnesses and 95 percent of the deaths associated with eating infected shellfish, although the overall incidence of illness from *Vibrio* infection remains low. There is a close association between temperature, *Vibrio* sp. abundance, and clinical illness. The U.S. infection rate increased 41 percent from 1996 to 2006,¹⁶³ concurrent with rising temperatures.
- As temperatures rise, tick populations that carry Rocky Mountain spotted fever are projected to shift from south to north.³⁰⁶
- The introduction of disease-causing agents from other regions of the world is an additional threat.¹⁶³

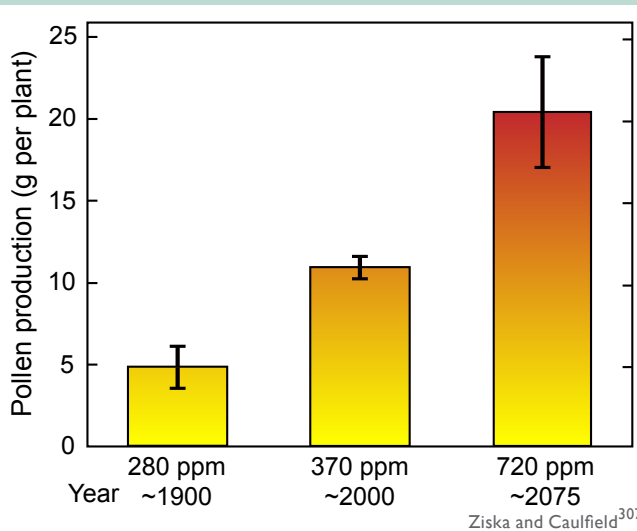
While the United States has programs such as the Safe Drinking Water Act that help protect against some of these problems, climate change will present new challenges.

Rising temperature and carbon dioxide concentration increase pollen production and prolong the pollen season in a number of plants with highly allergenic pollen, presenting a health risk.

Rising carbon dioxide levels have been observed to increase the growth and toxicity of some plants that cause health problems. Climate change has caused an earlier onset of the spring pollen season in the United States.²⁷² It is reasonable to conclude that allergies caused by pollen have also experienced associated changes in seasonality.²⁷² Several laboratory studies suggest that increasing carbon dioxide concentrations and temperatures increase ragweed pollen production and prolong the ragweed pollen season.^{163,272}

Poison ivy growth and toxicity is also greatly increased by carbon dioxide, with plants growing larger and more allergenic. These increases exceed those of most beneficial plants. For example, poison ivy vines grow twice as much per year in air with a doubled preindustrial carbon dioxide concentration as they do in unaltered air; this is nearly five times the increase reported for tree species in other

Pollen Counts Rise with Increasing Carbon Dioxide



Pollen production from ragweed grown in chambers at the carbon dioxide concentration of a century ago (about 280 parts per million [ppm]) was about 5 grams per plant; at today's approximate carbon dioxide level, it was about 10 grams; and at a level projected to occur about 2075 under the higher emissions scenario,⁹¹ it was about 20 grams.³⁰⁷



Poison ivy



analyses.³⁰⁸ Recent and projected increases in carbon dioxide also have been shown to stimulate the growth of stinging nettle and leafy spurge, two weeds that cause rashes when they come into contact with human skin.^{309,310}

Certain groups, including children, the elderly, and the poor, are most vulnerable to a range of climate-related health effects.

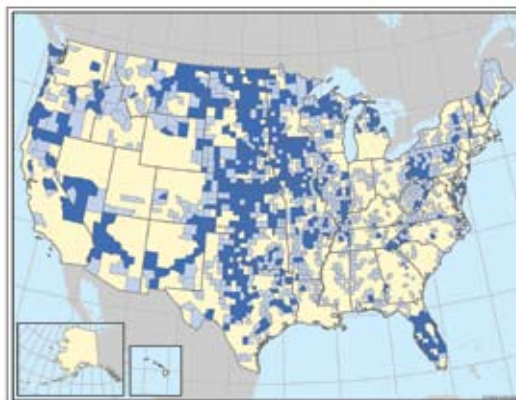
Infants and children, pregnant women, the elderly, people with chronic medical conditions, outdoor workers, and people living in poverty are especially at risk from increasing heat stress, air pollution, extreme weather events, and diseases carried by food, water, and insects.¹⁶³

Geographic Vulnerability of U.S. Residents to Selected Climate-Related Health Impacts

a) Location of Hurricane Landfalls, 1995 to 2000



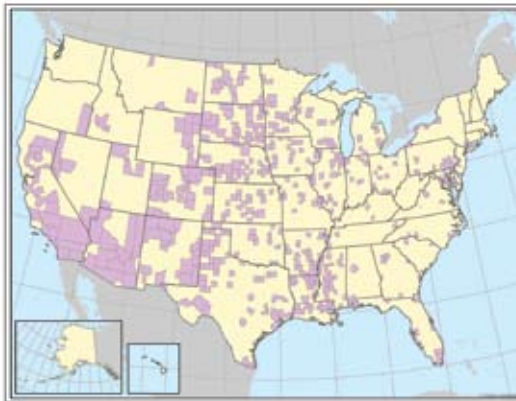
c) Percentage of Population Aged 65 or older



b) Locations of Extreme Heat Events, 1995 to 2000



d) West Nile Virus Cases, 2004



CCSP SAP 4.6¹⁶³

Maps indicating U.S. counties, or in some cases states, with existing vulnerability to climate-sensitive health outcomes: a) location of hurricane landfalls; b) extreme heat events (defined by the Centers for Disease Control as temperatures 10 or more degrees F above the average high temperature for the region and lasting for several weeks); c) percentage of population over age 65 (dark blue indicates that percentage is over 17.6 percent, light blue 14.4 to 16.5 percent); d) locations of West Nile virus cases reported in 2004. These examples demonstrate both the diversity of climate-sensitive health outcomes and the geographic variability of where they occur. Events over short time spans, in particular West Nile virus cases, are not necessarily predictive of future vulnerability.

L1 Children’s small ratio of body mass to surface
 L2 area and other factors make them vulnerable to
 L3 heat-related illness and death. Their increased
 L4 breathing rate relative to body size, additional time
 L5 spent outdoors, and developing respiratory tracts,
 L6 heighten their sensitivity to air pollution. In addition,
 L7 children’s immature immune systems increase
 L8 their risk of serious consequences from water- and
 L9 food-borne diseases, while developmental factors
 L10 make them more vulnerable to complications from
 L11 severe infections such as *E. coli* or *Salmonella*.¹⁶³

L12
 L13 The greatest health burdens related to climate
 L14 change are likely to fall on the poor, especially
 L15 those lacking adequate shelter and access to other
 L16 resources such as air conditioning.¹⁶³

L17
 L18 Elderly people are more likely to have debilitating
 L19 chronic diseases or limited mobility. The elderly
 L20 are also generally more sensitive to extreme heat
 L21 for several reasons. They have a reduced ability to
 L22 regulate their own body temperature or sense when
 L23 they are too hot. They are at greater risk of heart
 L24 failure, which is further exacerbated when cardiac
 L25 demand increases in order to cool the body during
 L26 a heat wave. Also, people taking medications, such
 L27 as diuretics for high blood pressure, have a higher
 L28 risk of dehydration.¹⁶³

L29
 L30 The multiple health risks associated with diabetes
 L31 will increase the vulnerability of the U.S. popula-
 L32 tion to increasing temperatures. The
 L33 number of Americans with diabetes has
 L34 grown to about 24 million people, or
 L35 roughly 8 percent of the U.S. population.
 L36 Almost 25 percent of the population 60
 L37 years and older had diabetes in 2007.³¹¹
 L38 Fluid imbalance and dehydration create
 L39 higher risks for diabetics during heat
 L40 waves. People with diabetes-related
 L41 heart disease are at especially increased
 L42 risk of dying in heat waves.

R1 High obesity rates in the United States are a con-
 R2 tributing factor in currently high levels of diabe-
 R3 tes. Similarly, a factor in rising obesity rates is a
 R4 sedentary lifestyle and automobile dependence;
 R5 60 percent of Americans do not meet minimum
 R6 daily exercise requirements. Making cities more
 R7 walkable and bikeable would thus have multiple
 R8 benefits: improved personal fitness and weight loss;
 R9 reduced local air pollution and associated respirato-
 R10 ry illness; and reduced greenhouse gas emissions.³¹²

R11
 R12 The United States has considerable capacity to
 R13 adapt to climate change, but during recent extreme
 R14 weather and climate events, actual practices have
 R15 not always protected people and property. Vulner-
 R16 ability to extreme events is highly variable, with
 R17 disadvantaged groups and communities (such as the
 R18 poor, infirm, and elderly) experiencing consider-
 R19 able damage and disruptions to their lives. Adapta-
 R20 tion tends to be reactive, unevenly distributed, and
 R21 focused on coping rather than preventing problems.
 R22 Future reduction in vulnerability will require
 R23 consideration of how best to incorporate planned
 R24 adaptation into long-term municipal and public ser-
 R25 vice planning, including energy, water and health
 R26 services, in the face of changing climate-related
 R27 risks combined with ongoing changes in population
 R28 and development patterns.^{163,164}



Society

Key Messages:

- Population shifts and development choices are making more Americans vulnerable to the expected impacts of climate change.
- Vulnerability is greater for those who have few resources and few choices.
- City residents and city infrastructure have unique vulnerabilities to climate change.
- Climate change affects communities through changes in climate-sensitive resources that occur both locally and at great distances.
- Insurance is one of the industries particularly vulnerable to increasing extreme weather events such as severe storms, but it can also help society manage the risks.
- The United States is connected to a world that is unevenly vulnerable to climate change and thus will be affected by impacts in other parts of the globe.

Key Sources



Climate change will affect society through impacts on the necessities and comforts of life: water, energy, housing, transportation, food, natural ecosystems, and health. This section focuses on some characteristics of society that make it vulnerable to the potential impacts of climate change and how the risks and costs may be distributed. Many impacts of climate change on society, for example, sea-level rise and increased water scarcity, are covered in other sections of this report. This section is not a comprehensive analysis of societal vulnerabilities, but rather highlights key examples.

Because societies and their built environments have developed under a climate that fluctuates within a relatively confined range of conditions, most impacts of a rapidly changing climate will present challenges. Society is especially vulnerable to extremes, such as heat waves and floods, many of which are increasing as climate changes.³¹³ And while there are likely to be some benefits and opportunities in the early stages of warming, as climate continues to change, negative impacts are projected to dominate.¹⁶⁴

Climate change will affect different segments of society differently because of their varying exposures and adaptive capacities. The impacts of climate change also do not affect society in isola-

tion. Rather, impacts can be exacerbated when climate change occurs in combination with the effects of an aging and growing population, pollution and poverty, and natural environmental fluctuations.^{164,172,274} Unequal adaptive capacity in the world as a whole also will pose challenges to the United States. Poorer countries are projected to be disproportionately affected by the impacts of climate change and the United States is strongly connected to the world beyond its borders through markets, trade, investments, shared resources, migrating species, health, travel and tourism, environmental refugees (those fleeing deteriorating environmental conditions), and security.



Cedar Rapids, Iowa, June 12, 2008

Population shifts and development choices are making more Americans vulnerable to the expected impacts of climate change.

Climate is one of the key factors in Americans’ choices of where to live. As the U.S. population grows, ages, and becomes further concentrated in cities and coastal areas, society is faced with additional challenges. Climate change is likely to exacerbate these challenges as changes in temperature, precipitation, sea levels, and extreme weather events increasingly affect homes, communities, water supplies, land resources, transportation, urban infrastructure, and regional characteristics that people have come to value and depend on.

Population growth in the United States over the past century has been most rapid in the South, near the coasts, and in large urban areas (see figure on page 55 in the *Energy* sector). The four most populous states in 2000 – California, Texas, Florida, and New York – accounted for 38 percent of the total growth in U.S. population during that time, and share significant vulnerability to coastal storms, severe drought, sea-level rise, air pollution, and urban heat island effects.³¹³ But migration patterns are now shifting: the population of the Mountain West (Montana, Idaho, Wyoming, Nevada, Utah, Colorado, Arizona, and New Mexico) is projected to increase by 65 percent from 2000 to 2030, representing one-third of all U.S. population growth.^{274,314} Southern coastal areas on both the Atlantic and the Gulf of Mexico are projected to continue to see population growth.³¹³

Overlaying projections of future climate change and its impacts on expected changes in U.S. population and development patterns reveals a critical insight: more Americans will be living in the areas that are most vulnerable to the effects of climate change.²⁷⁴

America’s coastlines have seen pronounced population growth in regions most at risk of hurricane activity, sea-level rise, and storm surge; putting more people and property in harm’s way as the probability of harm increases.²⁷⁴ On the Atlantic and Gulf coasts where hurricane activity is prevalent, the coastal land in many areas is sinking while sea level is rising; human activities are exacerbating the

loss of coastal wetlands that once helped buffer the coastline from erosion due to storms. The devastation caused by recent hurricanes highlights the vulnerability of these areas.²²⁴

The most rapidly growing area of the country is the Mountain West, a region projected to face more frequent and severe wildfires and have less water available, particularly during the high-demand period of summer. Continued population growth in these arid and semi-arid regions would stress water supplies. Because of high demand for irrigating agriculture, overuse of rivers and streams is common in the arid West, particularly along the Front Range of the Rocky Mountains in Colorado, in Southern California, and in the Central Valley of California. Rapid population and economic growth in these arid and semi-arid regions has dramatically increased vulnerability to water shortages (see *Water Resources* sector and *Southwest* region).²⁷⁴

Many questions are raised by ongoing development patterns in the face of climate change. Will growth continue as projected in vulnerable areas, despite the risks? Will there be a retreat from the coastline as it becomes more difficult to insure vulnerable properties? Will there be pressure for the government to insure properties that private insurers have rejected? How can the vulnerability of new development be minimized? How can we ensure that communities adopt measures to manage the significant changes that are projected in sea level, temperature, rainfall, and extreme weather events?

Development choices are based on people’s needs and desires for places to live, economies that provide employment, ecosystems that provide services, and community-based social activities. Thus, the future vulnerability of society will be influenced by how and where people choose to live. Some choices, such as expanded development in coastal regions, can increase vulnerabilities to climate-related events, even without any change in climate.

Vulnerability is greater for those who have few resources and few choices.

Vulnerabilities to climate change depend not only on where people are but also on who they are.



L1 In general, groups that are especially vulnerable
 L2 include the very young, the very old, the sick, and
 L3 the poor. These groups represent a more significant
 L4 portion of the total population in some regions and
 L5 localities than others. For example, the elderly more
 L6 often cite a warm climate as motivating their choice
 L7 of where to live and thus make up a larger share of
 L8 the population in warmer areas.³⁰⁵

L10 People with few resources often live in marginal
 L11 locations, such as in river floodplains or low-
 L12 lying coastal areas, which increases their risk. For
 L13 example, the experience with Hurricane Katrina
 L14 showed that the poor and elderly were the most
 L15 vulnerable because of where they lived and their
 L16 limited ability to get out of harm's way. Thus, those
 L17 who have the least often proportionately lose the
 L18 most. And it is clear that people with access to
 L19 financial resources, including insurance, have a
 L20 greater capacity to adapt to, recover, or escape from
 L21 adverse impacts of climate change than those who
 L22 do not have such access. The fate of the poor can be
 L23 permanent dislocation, leading to the loss of social
 L24 relationships and community support networks
 L25 provided by schools, churches, and neighborhoods.

L27 Native American communities have unique vulner-
 L28 abilities. Native Americans who choose to live on
 L29 established reservations are restricted to reservation
 L30 boundaries and therefore have limited relocation
 L31 options. In Alaska, over 100 villages on the coast
 L32 and in low-lying areas along rivers are subject to
 L33 increased flooding and erosion due to warming.³¹⁵
 L34 Warming also reduces the availability and acces-
 L35 sibility of many traditional food sources for Native
 L36 Alaskans, such as seals that live on ice and caribou
 L37 whose migration patterns depend on being able to
 L38 cross frozen rivers and wetlands. These vulnerable
 L39 people face losing their livelihoods, their com-
 L40 munities, and in some cases, their culture, which
 L41 depends on traditional ways of collecting and
 L42 sharing food.^{132,220} Native cultures in the Southwest
 L43 are particularly vulnerable to impacts of climate
 L44 change on water quality and availability.

L46 In the future (as in the past), the impacts of climate
 L47 change are likely to fall disproportionately on the
 L48 disadvantaged.³¹³ For example, the sensitivity of
 L49 California's population to increased air and water
 L50 pollution, heat waves, and other weather-related



Chalmette, Louisiana after Hurricane Katrina



Katrina flood waters

problems shows significant racial and socioeco-
 nomic differences, particularly for those who
 live and work without air conditioning.³¹⁶ Stud-
 ies specifically examining the impacts of climate
 change on the African American community in
 the United States have concluded that they are both
 economically and physically more vulnerable to
 climate-related disasters, illness, and price shocks.
 Economic impacts of climate change such as higher
 prices for food, water, and energy are also expected
 to impose new economic burdens on low-income
 households.³¹⁷ However, these same studies have
 concluded that investments in clean energy and im-
 proved air quality would significantly benefit these
 vulnerable populations.³¹⁸

City residents and city infrastructure have unique vulnerabilities to climate change.

Over 80 percent of the U.S. population resides in
 urban areas, which are among the most rapidly
 changing environments on Earth. In recent de-
 cades, cities have become increasingly spread out,
 complex, and interconnected with regional and
 national economies and infrastructure.³¹⁹ Cities
 also experience a host of social problems, includ-
 ing neighborhood degradation, traffic congestion,
 crime, unemployment, poverty, and inequities in
 health and well-being.³²⁰ Climate-related changes
 such as increased heat, water shortages, and
 extreme weather events will add further stress to
 existing problems. The impacts of climate change
 on cities are compounded by aging infrastructure,
 buildings, and populations; as well as increased air

R1
 R2
 R3
 R4
 R5
 R6
 R7
 R8
 R9
 R10
 R11
 R12
 R13
 R14
 R15
 R16
 R17
 R18
 R19
 R20
 R21
 R22
 R23
 R24
 R25
 R26
 R27
 R28
 R29
 R30
 R31
 R32
 R33
 R34
 R35
 R36
 R37
 R38
 R39
 R40
 R41
 R42
 R43
 R44
 R45
 R46
 R47
 R48
 R49
 R50

L1 pollution and population growth. Further, infra-
 L2 structure designed to handle past variations in
 L3 climate can instill a false confidence in its ability to
 L4 handle future changes. However, urban areas also
 L5 present opportunities for adaptation through tech-
 L6 nology, infrastructure, planning, and design.³¹³
 L7

L8 As cities grow, they alter local climates through the
 L9 urban heat island effect. This effect occurs because
 L10 cities absorb, produce, and retain more heat than
 L11 the surrounding countryside. The urban heat island
 L12 effect has raised average urban air temperatures
 L13 by 2 to 5°F more than surrounding areas over the
 L14 past 100 years, and by up to 20°F more at night.³²¹
 L15 Such temperature increases, on top of the general
 L16 increase caused by human-induced warming, affect
 L17 urban dwellers in many ways, influencing health,
 L18 comfort, energy costs, air quality, water quality
 L19 and availability, and violent crime (which increases
 L20 at high temperatures) (see *Human Health, Energy,*
 L21 *and Water Resources* sectors).^{172,313,322,323}
 L22
 L23

R1 More frequent heavy downpours and floods in
 R2 urban areas will cause greater property damage, a
 R3 heavier burden on emergency management, in-
 R4 creased clean-up and rebuilding costs, and a grow-
 R5 ing financial toll on businesses and homeowners.
 R6 The Midwest floods of 2008 provide a recent vivid
 R7 example of such tolls. Heavy downpours and urban
 R8 floods can also overwhelm combined sewer and
 R9 storm-water systems and release pollutants to wa-
 R10 terways.³¹³ Unfortunately, for many cities, current
 R11 planning and existing infrastructure are designed
 R12 for the historical one-in-100 year event, whereas
 R13 cities are likely to experience this same flood level
 R14 much more frequently as a result of the climate
 R15 change projected over this century.^{146,164,324}
 R16

R17 Cities are also likely to be affected by climate
 R18 change in unforeseen ways, necessitating diversion
 R19 of city funds for emergency responses to extreme
 R20 weather.³¹³ There is the potential for increased sum-
 R21 mer electricity blackouts owing to greater demand
 R22 for air conditioning.³²⁵ Unreliable electric power,
 R23

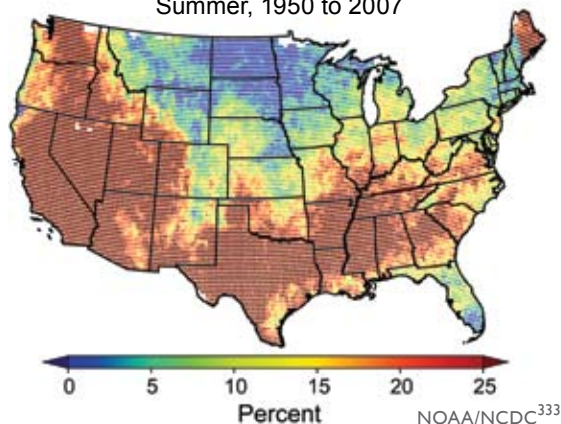
L24 **Heat, Drought, and Stagnant Air Degrade Air Quality and Quality of Life**

L25 Heat waves and poor air quality already threaten the lives of thousands of people each year. Experience
 L26 and research have shown that these events are interrelated as the atmospheric conditions that produce
 L27 heat waves are often accompanied by stagnant air and poor air quality.³²⁶ The simultaneous occurrence of
 L28 heat waves, drought, and stagnant air negatively affects quality of life, especially in cities.
 L29
 L30

L31 One such event occurred in the United States during the summer of 1988, causing 5,000 to 10,000 deaths
 L32 and economic losses of more than \$70 billion (in 2002 dollars).^{229,327} Half of the nation was affected by
 L33 drought, and 5,994 all-time daily high temperature records were set around the country in July alone
 L34 (more than three times the most recent 10-year
 L35 average).^{328,329} Poor air quality resulting from the lack
 L36 of rainfall, high temperatures, and stagnant conditions
 L37 led to an unprecedented number of unhealthy air
 L38 quality days throughout large parts of the country.^{327,329}
 L39 Continued climate change is projected to increase the
 L40 likelihood of such episodes.^{68,330}
 L41
 L42

L43 Interactions such as those between heat wave and
 L44 drought will affect adaptation planning. For example,
 L45 electricity use increases during heat waves due to
 L46 increased air conditioning demand.^{330,331} During
 L47 droughts, cooling water availability is at its lowest.
 L48 Thus, during a simultaneous heat wave and drought,
 L49 electricity demand for cooling will be high when power
 L50 plant cooling water availability is at its lowest.³⁴⁰

R24 **Stagnation When Heat Waves Exist**
 R25 Summer, 1950 to 2007



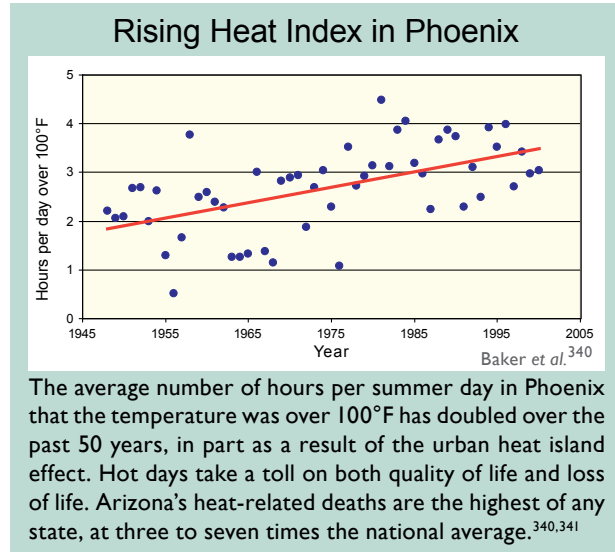
R26 The map shows the frequency of occurrence of stagnant
 R27 air conditions when heat wave conditions were also
 R28 present. Since 1950, across the Southeast, southern Great
 R29 Plains, and most of the West, the air was stagnant more
 R30 than 25 percent of the time during heat waves.
 R31
 R32
 R33
 R34
 R35
 R36
 R37
 R38
 R39
 R40
 R41
 R42
 R43
 R44
 R45
 R46
 R47
 R48
 R49
 R50

L1 which affected minority neighborhoods during
 L2 New York City’s 1999 heat wave, can pose health
 L3 risks and environmental justice issues because of
 L4 their disproportionate effect on minority popula-
 L5 tions.³¹⁸ In southern California’s cities, additional
 L6 summer electricity demand will intensify conflicts
 L7 between hydropower and flood-control objec-
 L8 tives.¹⁶⁴ Increased costs of repairs and maintenance
 L9 are projected for transportation systems, including
 L10 roads, railways, and airports, as they are negatively
 L11 affected by heavy downpours and extreme heat¹⁹⁰
 L12 (see *Transportation* sector). Coping with increased
 L13 flooding will require replacement or improvements
 L14 in storm drains, flood channels, levees, and dams.
 L15

L16 In addition, coastal cities are also vulnerable to
 L17 sea-level rise, storm surge, and increased hurricane
 L18 intensity. Cities such as New Orleans, Miami, and
 L19 New York are particularly at risk, and would have
 L20 difficulty coping with the sea-level rise projected
 L21 by the end of the century under a higher emissions
 L22 scenario.^{91,164} Remnants of hurricanes moving in-
 L23 land also threaten cities of the Appalachian Moun-
 L24 tains, which are vulnerable if hurricane frequency
 L25 or intensity increases. Since most large U.S. cities
 L26 are on coasts, rivers, or both, climate change will
 L27 lead to increased potential flood damage. The larg-
 L28 est impacts are expected when sea-level rise, heavy
 L29 runoff, high tides, and storms coincide.³¹³ Analyses
 L30 of New York and Boston indicate that the potential
 L31 impacts of climate change are likely to be negative,
 L32 but that vulnerability can be reduced by behavioral
 L33 and policy changes.^{313,334-336}
 L34

L35 Urban areas concentrate the human activities that
 L36 are largely responsible for heat-trapping emissions.
 L37 The demands of urban residents are also associated
 L38 with a much larger footprint on areas far removed
 L39 from these population centers.³³⁷ Cities thus have a
 L40 large role to play in reducing heat-trapping emis-
 L41 sions, and many are pursuing such actions. For
 L42 example, over 700 cities have committed to the U.S.
 L43 Mayors’ Climate Protection Agreement to advance
 L44 emissions reduction goals.
 L45

L46 Cities also have considerable potential to adapt to
 L47 climate change through technological, institutional,
 L48 structural, and behavioral changes. For example, a
 L49 number of cities have warning programs in place
 L50 to reduce heat-related illness and death (see *Huma-*



Health sector). Relocating development away from low-lying areas, building new infrastructure with future sea-level rise in mind, and promoting water conservation are examples of structural and institutional strategies. Choosing road materials that can handle higher temperatures is an adaptation option that relies on new technology (see *Transportation* sector). Cities can reduce heat loads by increasing reflective surfaces and green spaces. Some actions have multiple benefits. For example, increased planting of trees and other vegetation in cities has been shown to be associated with a reduction in crime,³³⁸ in addition to reducing local temperatures, and thus energy demand for air conditioning.

Human well-being is influenced by economic conditions, natural resources and amenities, public health and safety, infrastructure, government, and social and cultural resources. Climate change will influence all of these, but an understanding of the many interacting impacts, as well as the ways society can adapt to them, remains in its infancy.^{305,339}

Climate change affects communities through changes in climate-sensitive resources that occur both locally and at great distances.

Human communities are intimately connected to resources beyond their geographical boundaries. Thus, communities will be vulnerable to the potential impacts of climate change on sometimes-distant resources. For example, communities that have developed near areas of agricultural production, such

R1
R2
R3
R4
R5
R6
R7
R8
R9
R10
R11
R12
R13
R14
R15
R16
R17
R18
R19
R20
R21
R22
R23
R24
R25
R26
R27
R28
R29
R30
R31
R32
R33
R34
R35
R36
R37
R38
R39
R40
R41
R42
R43
R44
R45
R46
R47
R48
R49
R50

as the Midwest corn belt or the wine-producing regions of California and the Northwest, depend on the continued productivity of those regions, which would be compromised by increased temperature or severe weather.³¹³ Some agricultural production that is linked to cold climates is likely to disappear entirely: recent warming has altered the required temperature patterns for maple syrup production, shifting production northward from New England into Canada. Similarly, cranberries require a long winter chill period, which is shrinking as climate warms²³⁴ (see *Northeast* region). Most cities depend on water supplies from distant watersheds, and those depending on diminishing supplies (such as the Sierra Nevada snowpack) are vulnerable. Northwest communities also depend upon forest resources for their economic base, and many island, coastal, and “sunbelt” communities depend on tourism.

Recreation and tourism play important roles in the economy and quality of life of many Americans. In some regions tourism and recreation are major job creators, bringing billions of dollars to regional economies. Across the nation, fishing, hunting, skiing, snowmobiling, diving, beach-going, and other outdoor activities make important economic contributions and are a part of family traditions that have value that goes beyond financial returns. A changing climate will mean reduced opportunities for some activities and locations and expanded opportunities for others.^{305,342} Hunting and fishing will change as animals’ habitats shift and as relationships among species in natural communities are disrupted by their different responses to rapid

climate change. Water-dependent recreation in areas projected to get drier, such as the Southwest, and beach recreation in areas that are expected to see rising sea levels, will suffer. Some regions will see an expansion of the season for warm weather recreation such as hiking and bicycle riding.

Insurance is one of the industries particularly vulnerable to increasing extreme weather events such as severe storms, but it can also help society manage the risks.

Insurance – the world’s largest industry – is one of the primary mechanisms through which the costs of climate change are distributed across society.

Most of the climate change impacts described in this report have economic consequences. A significant portion of these flow through public and private insurance markets, which essentially aggregate and distribute society’s risk. Insurance thus provides a window into the myriad ways in which the costs of climate change will manifest, and serves as a form of economic adaptation and a messenger of these impacts through the terms and price signals it sends its customers.³⁴⁴

In an average year, about 90 percent of insured catastrophe losses worldwide are weather-related. In the United States, about half of all these losses are insured, which amounted to \$320 billion between 1980 and 2005 (inflation-adjusted to 2005 dollars). While major events such as hurricanes grab head-

lines, the aggregate effect of smaller events accounts for at least 60 percent of total insured losses on average.³⁴⁴ Many of the smallest scale property losses and weather-related life/health losses are unquantified.³⁴⁵

Escalating exposures to catastrophic weather events, coupled with private insurers’ withdrawal from various markets, are placing the federal government at increased

Examples of Impacts On Recreation

Recreational Activity	Scenario of Potential Impact of Climate Change	Economic Impact
Skiing, Northeast	20 percent reduction in ski season length	\$800 million loss per year, potential resort closures ²³⁴
Snowmobiling, Northeast	Reduction of season length under higher emissions scenario ⁹¹	Complete loss of opportunities in New York and Pennsylvania within a few decades, 80 percent reduction in season length for region by end of century ^{234,342} .
Beaches, North Carolina	14 of 17 beaches permanently underwater by 2080	Lost opportunities for beach and fishing trips = \$3.9 billion over 75 years ³⁴³

L1 financial risk as insurer of last resort. The National Flood Insur-
 L2 ance Program would have gone bankrupt after the storms of
 L3 2005 had they not been given the ability to borrow about \$20 bil-
 L4 lion from the U.S. Treasury.¹⁷² For public and private insurance
 L5 programs alike, rising losses require a combination of risk-based
 L6 premiums and improved loss prevention.

L7
 L8 While economic and demographic factors have no doubt contrib-
 L9 uted to observed increases in losses,³⁴⁶ these factors do not fully
 L10 explain the upward trend in costs or numbers of events.^{344,347}
 L11 For example, during the time period covered in the figure to
 L12 the right, population increased by a factor of 1.3 while losses
 L13 increased by a factor of 15 to 20 in inflation-corrected dollars.
 L14 Analyses asserting little or no role of climate change in increas-
 L15 ing the risk of losses tend to focus on a highly limited set of haz-
 L16 ards and locations. They also often fail to account for the vaga-
 L17 raries of natural cycles and inflation adjustments, or to normalize
 L18 for countervailing factors such as improved pre- and post-event
 L19 loss prevention (such as dikes, building codes, and early warning
 L20 systems).^{348,349}

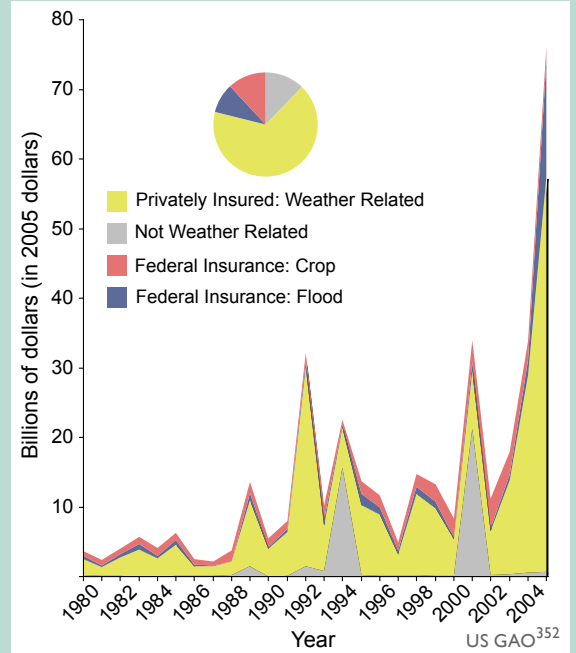
L21
 L22 What is known with far greater certainty is that future increases
 L23 in losses will be attributable to climate change as it increases the
 L24 frequency and intensity of many types of extreme weather, such
 L25 as severe thunderstorms and heat waves.^{131,350}

L26
 L27 Insurance is emblematic of the increasing globalization of cli-
 L28 mate risks. Because large U.S.-based companies operate around
 L29 the world, their customers and assets are exposed to climate
 L30 impacts wherever they occur. Most of the growth in the insurance
 L31 industry is in emerging markets, which will structurally increase
 L32 U.S. insurers' exposure to climate risk because those regions are
 L33 more vulnerable and are experiencing particularly high rates of
 L34 population growth and development.³⁵¹

L35
 L36 The movement of populations into harm's way creates a rising
 L37 baseline of insured losses upon which the consequences of cli-
 L38 mate change will be superimposed. These observations reinforce
 L39 a recurring theme in this report: the past can no longer be used as
 L40 the basis for planning for the future.

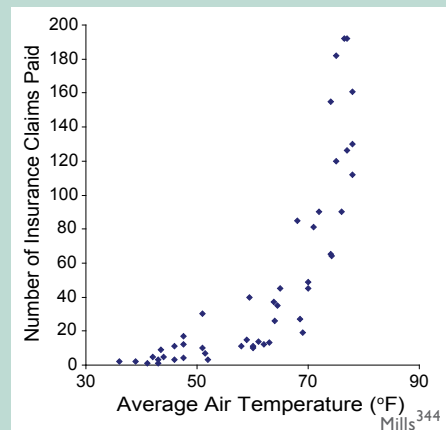
L41
 L42 It is a challenge to design insurance systems that properly price
 L43 risks, reward loss prevention, and do not foster risk taking (for
 L44 example by repeatedly rebuilding flooded homes). Market failures
 L45 of this sort compound society's vulnerability to climate change.
 L46 Rising losses²⁵³ are already affecting the availability and afford-
 L47 ability of insurance. Several million customers in the United
 L48 States, no longer able to purchase private insurance coverage, are
 L49 taking refuge in state-mandated insurance pools, or going with-
 L50 out insurance altogether. Offsetting rising insurance costs is one

Insured Losses from Catastrophes, 1980 to 2005



Weather-related insurance losses in the United States are increasing. Typical weather-related losses today are similar to those that resulted from the 9/11 attack (shown in gray at 2001 in the graph). About half of all economic losses are insured, so actual losses are roughly twice those shown on the graph. Data on smaller-scale losses (many of which are weather-related) are significant but are not included in this graph as they are not comprehensively reported by the U.S. insurance industry.

Lightning-Related Insurance Claims



There is a strong observed correlation between higher temperatures and the frequency of lightning-induced insured losses in the United States. Each marker represents aggregate monthly U.S. lightning-related insurance claims paid by a large national insurer over a five-year period, 1991-1995. All else being equal, these claims are expected to increase with temperature.^{344,353,354}

L1 benefit of mitigation and adaptation investments to
 L2 reduce the impacts of climate change.
 L3 Virtually all segments of the insurance industry
 L4 are vulnerable to the impacts of climate change.
 L5 Examples include damage to property, crops, for-
 L6 est products, livestock, and transportation infra-
 L7 structure; business and supply-chain interruptions
 L8 caused by weather extremes, water shortages, and
 L9 electricity outages; legal consequences;³⁵⁵ and
 L10 compromised health or loss of life. Increasing risks
 L11 to insurers and their customers are driven by many
 L12 factors including reduced periods of time between
 L13 loss events, increasing variability, shifting types
 L14 and location of events, and widespread simultane-
 L15 ous losses.

L16
 L17 In light of these challenges, insurers are emerging
 L18 as partners in climate science and the formulation
 L19 of public policy and adaptation strategies.³⁵⁶ Some
 L20 have promoted adaptation by providing premium
 L21 incentives for customers who fortify their proper-
 L22 ties, engaging in the process of determining build-
 L23 ing codes and land-use plans, and participating in
 L24 the development and financing of new technologies
 L25 and practices. For example, FEMA’s Community
 L26 Rating System is a point system that rewards com-
 L27 munities that undertake floodplain management
 L28 activities to reduce flood risk beyond the minimum
 L29 requirement set by the National Flood Insurance
 L30 Program. Everyone in these communities is reward-
 L31 ed with lower flood insurance premiums (–5 to –45
 L32 percent).³⁵⁷ Others have recognized that mitigation
 L33 and adaptation can work hand in hand in a coord-
 L34 inated climate risk-management strategy and are
 L35 offering “green” insurance products designed to
 L36 capture these dual benefits.^{90,351}

The United States is connected to a world that is unevenly vulnerable to climate change and thus will be affected by impacts in other parts of the globe.

L44 American society will not experience the poten-
 L45 tial impacts of climate change in isolation. In an
 L46 increasingly connected world, impacts elsewhere
 L47 will have political, social, economic, and environ-
 L48 mental ramifications for the United States. As in
 L49 the United States, vulnerability to the potential
 L50 impacts of climate change worldwide varies by

location, population characteristics, and economic R1
 status. The rising concentration of people in cit- R2
 ies is occurring globally, but is most prevalent R3
 in lower-income countries. Many large cities are R4
 located in vulnerable areas such as floodplains and R5
 coasts. In most of these cities, the poor often live in R6
 the most marginal of these environments, in areas R7
 that are susceptible to extreme events, and their R8
 ability to adapt is limited by their lack of financial R9
 resources.¹⁷² R10

R11
 R12 In addition, over half of the world’s population –
 R13 including most of the world’s major cities – depends
 R14 on glacier melt or snowmelt to supply water for
 R15 drinking and municipal uses. Today, some locations
 R16 are experiencing abundant water supplies and even
 R17 frequent floods due to increases in glacier melt
 R18 rates due to increased temperatures worldwide.
 R19 Soon, however, this trend is projected to reverse as
 R20 even greater temperature increases reduce glacier
 R21 mass and cause more winter precipitation to fall as
 R22 rain and less as snow.⁹⁰ R22

R23
 R24 As conditions worsen elsewhere, the number of
 R25 people wanting to immigrate to the United States
 R26 will increase. The direct cause of increased migra-
 R27 tion, such as extreme climatic events, will be diffi-
 R28 cult to separate from other forces that drive people
 R29 to migrate. Climate change also has the potential to
 R30 alter trade relationships by changing the compara-
 R31 tive trade advantages of regions or nations. As with
 R32 migration, shifts in trade can have multiple causes. R32

R33
 R34 Accelerating emissions in economies that are
 R35 rapidly expanding, such as China and India, pose
 R36 future threats to the climate system and already are
 R37 associated with air pollution episodes that reach the
 R38 United States. R38

R39
 R40 Meeting the challenge of improving conditions for
 R41 the world’s poor has economic implications for the
 R42 United States, as does intervention and resolution
 R43 of intra- and intergroup conflicts. Where climate
 R44 change exacerbates such challenges, for example by
 R45 limiting access to scarce resources or increasing in-
 R46 cidence of damaging weather events, consequences
 R47 are likely for the U.S. economy and security.³⁵⁸ R47



Northeast

The Northeast has significant geographic and climatic diversity within its relatively small area. The character and economy of the Northeast have been shaped by many aspects of its climate including its snowy winters, colorful autumns, and variety of extreme events such as nor'easters, ice storms, and heat waves. This familiar climate has already begun changing in noticeable ways.

Since 1970, the annual average temperature in the Northeast has increased by 2°F, with winter temperatures rising twice this much.¹⁵⁰ This warming has resulted in many other climate-related changes, including:

- More frequent days with temperatures above 90°F,
- A longer growing season,
- Increased heavy precipitation,
- Less winter precipitation falling as snow and more as rain,
- Reduced snowpack,
- Earlier breakup of winter ice on lakes and rivers,
- Earlier spring snowmelt resulting in earlier peak river flows, and
- Rising sea surface temperatures and sea level.

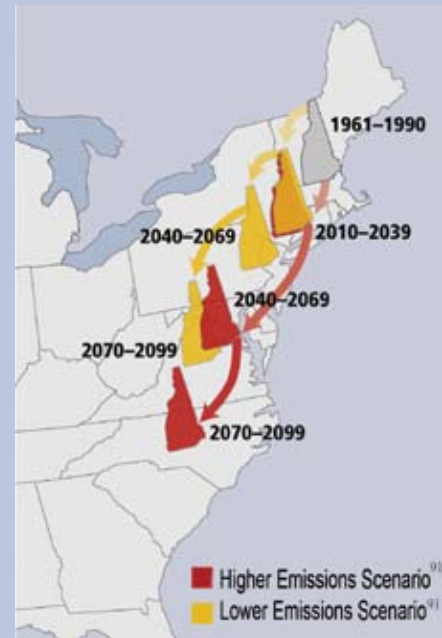
Each of these observed changes is consistent with the changes expected in this region from global warming. The Northeast is projected to face continued warming and more extensive climate-related changes, some of which could dramatically alter the region's economy, landscape, character, and quality of life.

Over the next several decades, temperatures are projected to rise an additional 2.5 to 4°F in winter and 1.5 to 3.5°F in summer.

By mid-century and beyond, however, today's emissions choices would generate starkly different climate futures; the lower the emissions, the smaller the climatic changes and resulting impacts.^{150,359} By late this century, under a higher-emissions scenario⁹¹:

- Winters in the Northeast are projected to be much shorter with fewer cold days and more precipitation.
- The length of the winter snow season would be cut in half across northern New York, Vermont, New Hampshire, and Maine, and reduced to a week or two in southern parts of the region.
- Cities that today experience few days above 100°F each summer would average 20 such days per summer, while certain cities, such as Hartford and Philadelphia, would average nearly 30 days over 100°F.
- Short-term (one- to three-month) droughts are projected to occur as frequently as once each summer in the Catskill and Adirondack mountains, and across the New England states.
- Hot summer conditions would arrive three weeks earlier and last three weeks longer into the fall.
- Sea level in this region is projected to rise about 2 feet, with the potential for a much larger rise, for reasons discussed in the *Global* and *National Climate Change* sections (see pages 25 and 37).

Climate on the Move: Changing Summers in New Hampshire



Hayhoe et al.³⁵⁹/Fig. from Frumhoff et al.²³⁴

Yellow arrows track what summers are projected to feel like under a lower emissions scenario⁹¹, while red arrows track projections for a higher emissions scenario⁹¹ (referred to as “even higher” on page 23). For example, under the higher emission scenario,⁹¹ by late this century residents of New Hampshire would experience a summer climate more like what occurs today in North Carolina.³⁵⁹



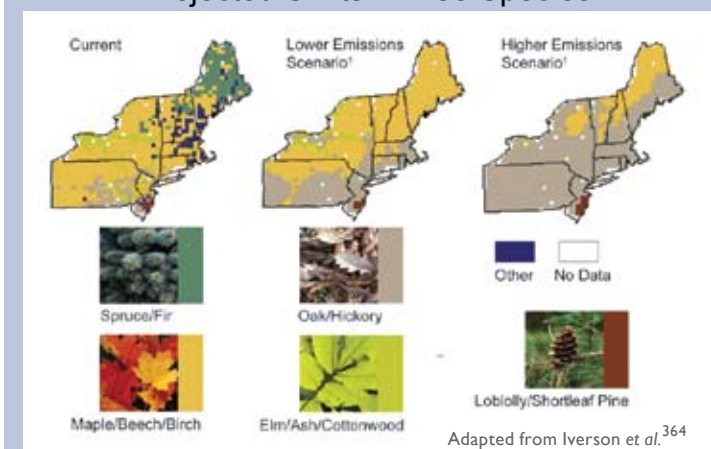
Extreme heat and declining air quality are likely to pose increasing problems for human health, especially in urban areas.

Heat waves, which are currently rare in the region, are projected to become much more commonplace in a warmer future, with major implications for human health (see *Human Health* sector).^{163,360}

In addition to the physiological stresses associated with hotter days and nights,³⁶⁰ for cities that now experience ozone pollution problems, the number of days that fail to meet federal air quality standards is projected to increase with rising temperatures if there are no further additional controls on ozone-causing pollutants^{163,361} (see *Human Health* sector). Sharp reductions in emissions will be needed to keep ozone within existing standards.

Projected changes in the summer heat provide a clear sense of how different the climate of the Northeast is projected to be under lower versus higher emissions scenarios. Changes of this kind will require greater use of air conditioning (see *Energy* sector).

Projected Shifts in Tree Species



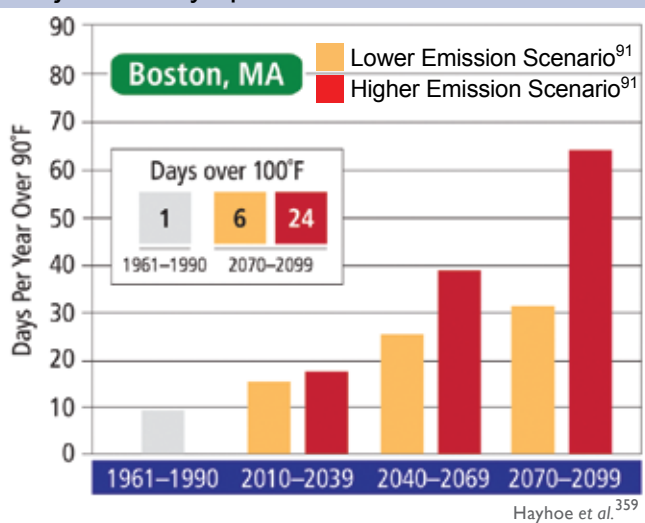
Much of the Northeast's forest is composed of the hardwoods maple, beech, and birch, while mountain areas and more northern parts of the region are dominated by spruce/fir forests. As climate changes over this century, suitable habitat for spruce and fir is expected to contract dramatically. Suitable maple/beech/birch habitat is projected to shift significantly northward under a higher emissions scenario (referred to as "even higher" on page 23),⁹¹ but to shift far less under a lower emissions scenario.^{91,363}

Agricultural production, including dairy, fruit, and maple syrup, are likely to be adversely affected as favorable climates shift.

Large portions of the Northeast are likely to become unsuitable for growing popular varieties of apples, blueberries, and cranberries under a higher emissions scenario.^{91,362,363} Climate conditions suitable for maple/beech/birch forests are projected to shift dramatically northward, eventually leaving only a small portion of the Northeast with a maple sugar business.³⁶⁴

The dairy industry is the most important agricultural sector in this region, with annual production worth \$3.6 billion.³⁶⁵ Heat stress in dairy cows depresses both milk production and birth rates for periods of weeks to months.^{193,366} By late this century, all but the northern parts of Maine, New Hampshire, New York, and Vermont are projected to suffer declines in July milk production under the higher emissions scenario.⁹¹ In parts of Connecticut, Massachusetts, New Jersey, New York, and Pennsylvania, a large decline in milk production, up to 20 percent or greater, is projected. Under the lower emissions scenario,⁹¹ however, reductions in milk production of up to 10 percent remain confined primarily to the southern parts of the region. This analysis used average monthly temperature

Projected Days per Year over 90°F in Boston



The graph shows model projections of the number of summer days with temperatures over 90°F in Boston, Massachusetts, under lower and higher (referred to as "even higher" on page 23) emissions scenarios.⁹¹ The inset shows projected days over 100°F.³⁵⁹

L1 and humidity data that do not
 L2 capture daily variations in heat
 L3 stress and projected increases
 L4 in extreme heat. Nor did the
 L5 analysis directly consider farmer
 L6 responses, such as installation
 L7 of potentially costly cooling sys-
 L8 tems. On balance, these projec-
 L9 tions are likely to underestimate
 L10 impacts on the dairy industry.¹⁵⁰

L11
 L12

L13 **Severe flooding due to**
 L14 **sea-level rise and heavy**
 L15 **downpours is likely to**
 L16 **occur more frequently.**

L17

L18 The densely populated coasts
 L19 of the Northeast face substan-
 L20 tial increases in the extent
 L21 and frequency of storm surge,
 L22 coastal flooding, erosion,
 L23 property damage, and loss of
 L24 wetlands.^{367,369} New York state
 L25 alone has more than \$2.3 trillion
 L26 in insured coastal property.³⁶⁸ Much of this coastline is exceptionally vulnerable to sea-level rise and related
 L27 impacts. Some major insurers have withdrawn coverage from thousands of homeowners in coastal areas of
 L28 the Northeast, including New York City.

L29

L30 Rising sea level is projected to increase the frequency and severity of damaging storm surges and flooding.
 L31 Under a higher emissions scenario,⁹¹ what is now considered a once-in-a-century coastal flood in New York
 L32 City is projected to occur at least twice as often by mid-century, and 10 times as often (or once per decade

L33

L34

L35 **Adaptation: Raising a Sewage Treatment Plant in Boston**

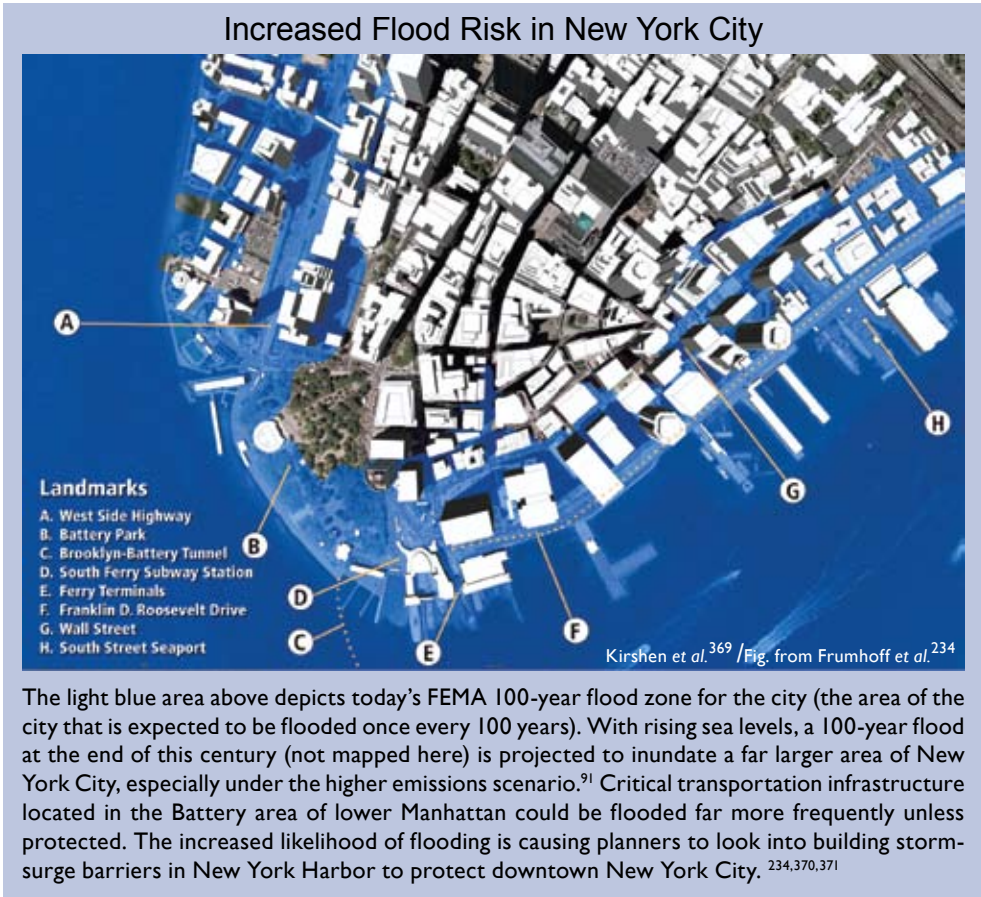
L36

L37

L38 Boston's Deer Island sewage treatment plant was designed and
 L39 built taking future sea-level rise into consideration. Because
 L40 the level of the plant relative to the level of the ocean at the
 L41 outfall is critical to the amount of rainwater and sewage that
 L42 can be treated, the plant was built 1.9 feet higher than it would
 L43 otherwise have been to accommodate the amount of sea-level
 L44 rise projected to occur by 2050, the planned life of the facility.

L45

L46 The planners recognized that the future would be different
 L47 from the past and they decided to plan for the future based
 L48 on the best available information. They assessed what could be easily and inexpensively changed
 L49 at a later date versus those things that would be more difficult and expensive to change later. For
 L50 example, increasing the plant's height would be less costly to incorporate in the original design, while
 protective barriers could be added at a later date, as needed, at a relatively small cost.



The light blue area above depicts today's FEMA 100-year flood zone for the city (the area of the city that is expected to be flooded once every 100 years). With rising sea levels, a 100-year flood at the end of this century (not mapped here) is projected to inundate a far larger area of New York City, especially under the higher emissions scenario.⁹¹ Critical transportation infrastructure located in the Battery area of lower Manhattan could be flooded far more frequently unless protected. The increased likelihood of flooding is causing planners to look into building storm-surge barriers in New York Harbor to protect downtown New York City.^{234,370,371}

R1
 R2
 R3
 R4
 R5
 R6
 R7
 R8
 R9
 R10
 R11
 R12
 R13
 R14
 R15
 R16
 R17
 R18
 R19
 R20
 R21
 R22
 R23
 R24
 R25
 R26
 R27
 R28
 R29
 R30
 R31
 R32
 R33
 R34
 R35
 R36
 R37
 R38
 R39
 R40
 R41
 R42
 R43
 R44
 R45
 R46
 R47
 R48
 R49
 R50

on average) by late this century. With a lower emissions scenario,⁹¹ today's 100-year flood is projected to occur once every 22 years on average by late this century.³⁶⁹

The projected reduction in snow cover will adversely affect winter recreation and the industries that rely upon it.

Winter snow and ice sports, which contribute some \$7.6 billion annually to the regional economy, will be particularly affected by warming.³⁴² Of this total, alpine skiing and other snow sports (not including snowmobiling) account for \$4.6 billion annually. Snowmobiling, which now rivals skiing as the largest winter recreation industry in the nation, accounts for the remaining \$3 billion.³⁷² Other winter traditions, ranging from skating and ice fishing on frozen ponds and lakes, to cross-country (Nordic) skiing, snowshoeing, and dog sledding, are integral to the character of the Northeast, and for many residents and visitors, its desirable quality of life.

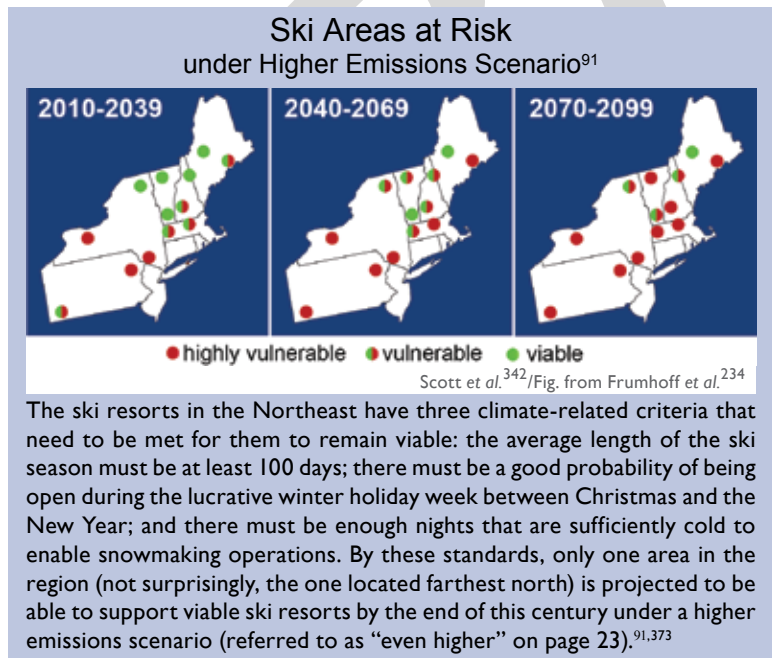
Warmer winters will shorten the average ski and snowboard seasons, increase artificial snowmaking requirements, and drive up operating costs. While snowmaking can enhance the prospects for ski resort success, it requires a great deal of water and energy, as well as very cold nights, which are becoming less frequent. Without the opportunity

to benefit from snowmaking, the prospects for the snowmobiling industry are even worse. Most of the region is likely to have a marginal or non-existent snowmobile season by mid-century.

The center of lobster fisheries is projected to continue its northward shift and the cod fishery on Georges Bank is likely to be diminished.

Lobster catch has increased dramatically in the Northeast as a whole over the past three decades, though not uniformly.^{374,375} Catches in the southern part of the region peaked in the mid-1990s, and have since declined sharply, beginning with a 1997 die-off in Rhode Island and Buzzards Bay (Massachusetts) associated with the onset of a temperature-sensitive bacterial shell disease, and accelerated by a 1999 lobster die-off in Long Island Sound. Currently, the southern extent of the commercial lobster harvest appears to be limited by this temperature-sensitive shell disease, and these effects are expected to increase with rising near-shore water temperatures. Analyses also suggest that lobster survival and settlement in northern regions of the Gulf of Maine could be increased by warming water, a longer growing season, more rapid growth, an earlier hatching season, an increase in nursery grounds suitable for larvae, and faster development of plankton.³⁷⁶

Cod populations throughout the North Atlantic are adapted to a wide range of seasonal ocean temperatures, including average annual temperatures near the seafloor ranging from 36 to 54°F. A maximum ocean temperature of 54°F represents the threshold of thermally suitable habitat for cod and the practical limit of cod distribution.³⁷⁷ Temperature also influences both the location and timing of spawning, which in turn affects the subsequent growth and survival of young cod. Studies indicate that increases in average annual bottom temperatures above 47°F will lead to a decline in growth and survival.^{378,379} Climate change will thus introduce an additional stress to an already-stressed fishery.



L1
L2
L3
L4
L5
L6
L7
L8
L9
L10
L11
L12
L13
L14
L15
L16
L17
L18
L19
L20
L21
L22
L23
L24
L25
L26
L27
L28
L29
L30
L31
L32
L33
L34
L35
L36
L37
L38
L39
L40
L41
L42
L43
L44
L45
L46
L47
L48
L49
L50

R1
R2
R3
R4
R5
R6
R7
R8
R9
R10
R11
R12
R13
R14
R15
R16
R17
R18
R19
R20
R21
R22
R23
R24
R25
R26
R27
R28
R29
R30
R31
R32
R33
R34
R35
R36
R37
R38
R39
R40
R41
R42
R43
R44
R45
R46
R47
R48
R49
R50



Southeast

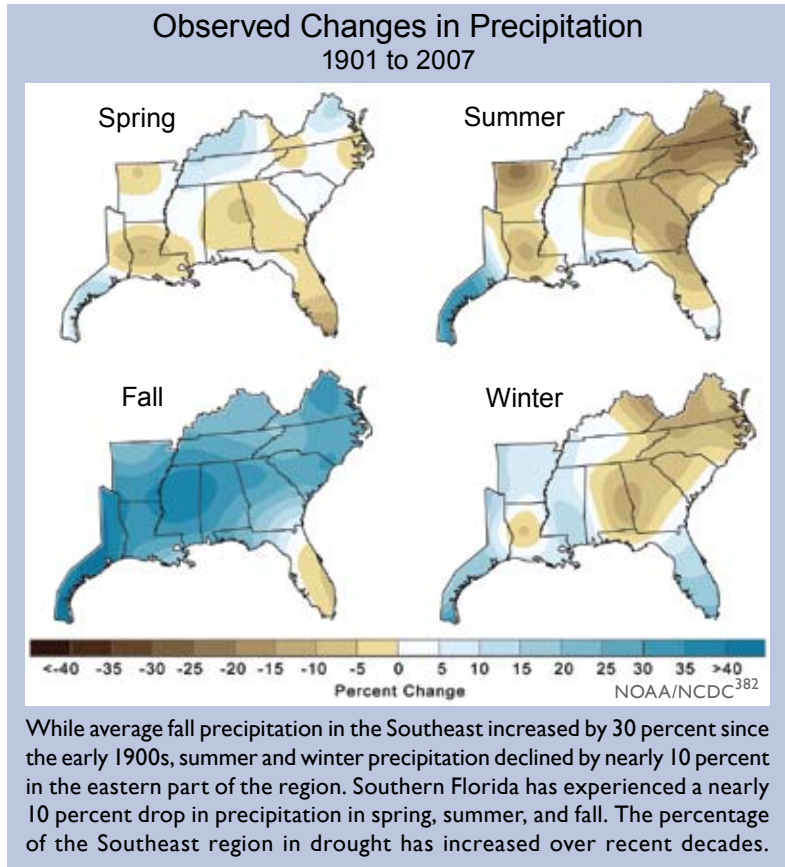
L1
L2
L3
L4
L5
L6
L7
L8
L9
L10
L11
L12
L13
L14
L15
L16
L17
L18
L19
L20
L21
L22
L23
L24
L25
L26
L27
L28
L29
L30
L31
L32
L33
L34
L35
L36
L37
L38
L39
L40
L41
L42
L43
L44
L45
L46
L47
L48
L49
L50

R1
R2
R3
R4
R5
R6
R7
R8
R9
R10
R11
R12
R13
R14
R15
R16
R17
R18
R19
R20
R21
R22
R23
R24
R25
R26
R27
R28
R29
R30
R31
R32
R33
R34
R35
R36
R37
R38
R39
R40
R41
R42
R43
R44
R45
R46
R47
R48
R49
R50

The climate of the Southeast is uniquely warm and wet, with mild winters and high humidity, compared with the rest of the continental United States. The average annual temperature of the Southeast did not change significantly over the past century as a whole. Since 1970, however, annual average temperature has risen about 2°F, with the greatest seasonal increase in temperature occurring during the winter months. The number of freezing days in the Southeast has declined by four to seven days per year for most of the region since the mid-1970s.

Average autumn precipitation has increased by 30 percent for the region since 1901. The decline in fall precipitation in South Florida contrasts strongly with the regional average. There has been an increase in heavy downpours in many parts of the region,^{380,381} while the percentage of the region experiencing moderate to severe drought increased over the past three decades. The area of moderate to severe spring and summer drought has increased by 12 percent and 14 percent, respectively, since the mid-1970s. Even in the fall months, when precipitation tended to increase in most of the region, the extent of drought increased by 9 percent.

Climate models project continued warming in all seasons across the Southeast and an increase in the rate of warming through the end of this century. The projected rates of warming are more than double those experienced in the Southeast since 1975, with the greatest temperature increases projected to occur in the summer months. The number of very hot days is projected to rise at a greater rate than the average temperature. Under a lower



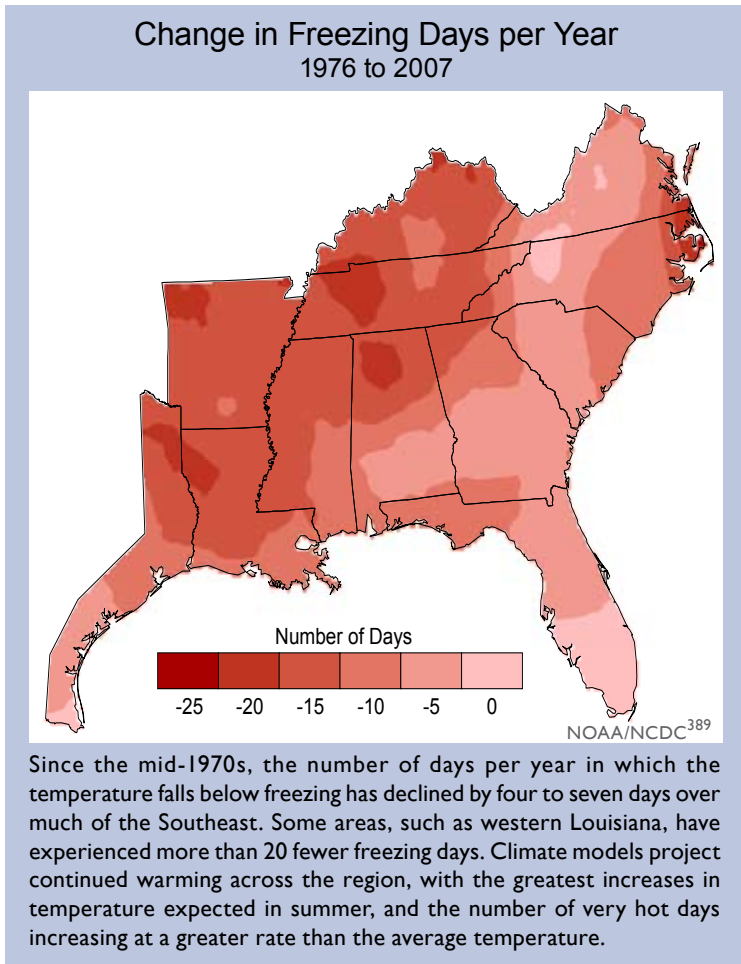
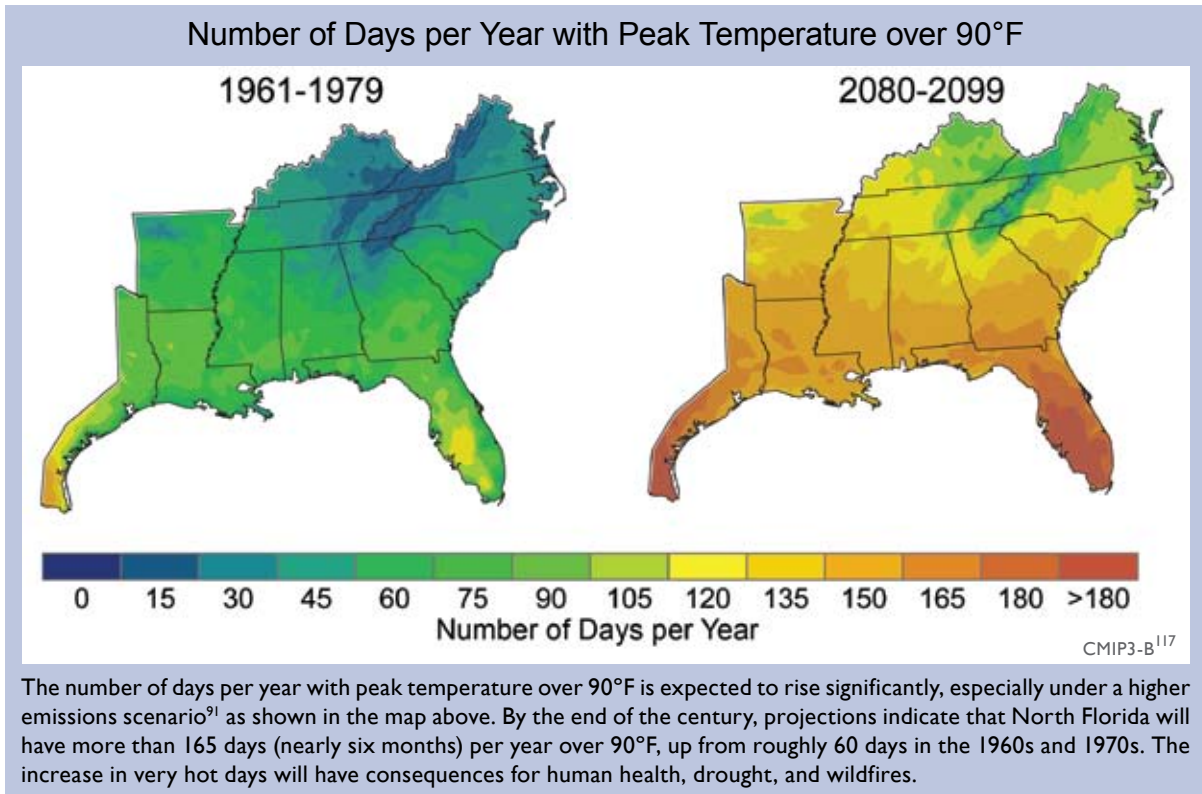
emissions scenario,⁹¹ average temperatures in the region are projected to rise by about 4.5°F by the 2080s, while a higher emissions scenario⁹¹ yields about 9°F of average warming (with about a 10.5°F increase in summer, and a much higher heat index). Spring and summer rainfall is projected to decline in South Florida during this century. Except for indications that the amount of rainfall from individual hurricanes will increase,⁶⁸ climate models

Average Change in Temperature and Precipitation in the Southeast					
	Temperature Change in °F			Precipitation change in %	
	1901-2008	1970-2008		1901-2008	1970-2008
Annual	0.3	1.6	Annual	6.0	-7.7
Winter	0.2	2.7	Winter	1.2	-9.6
Spring	0.4	1.2	Spring	1.7	-29.2
Summer	0.4	1.6	Summer	-4.0	3.6
Fall	0.2	1.1	Fall	27.4	0.1

This is a summary of observed climatic changes in the Southeast for two different periods.³⁸³ Most of the warming over the past century has occurred in the last several decades.

L1
L2
L3
L4
L5
L6
L7
L8
L9
L10
L11
L12
L13
L14
L15
L16
L17
L18
L19
L20
L21
L22
L23
L24
L25
L26
L27
L28
L29
L30
L31
L32
L33
L34
L35
L36
L37
L38
L39
L40
L41
L42
L43
L44
L45
L46
L47
L48
L49
L50

R1
R2
R3
R4
R5
R6
R7
R8
R9
R10
R11
R12
R13
R14
R15
R16
R17
R18
R19
R20
R21
R22
R23
R24
R25
R26
R27
R28
R29
R30
R31
R32
R33
R34
R35
R36
R37
R38
R39
R40
R41
R42
R43
R44
R45
R46
R47
R48
R49
R50



provide divergent results for future precipitation for the remainder of the Southeast. Models suggest that Gulf Coast states will tend to have less rainfall in winter and spring, compared with the more northern states in the region (see map on page 31 in the *National Climate Change* section). Because higher temperatures lead to more evaporation of moisture from soils and water loss from plants, the frequency, duration, and intensity of droughts are likely to continue to increase.

The destructive potential of Atlantic hurricanes has increased since 1970, correlated with an increase in sea surface temperature. A similar relationship with the frequency of land falling hurricanes has not been established^{98,384-387} (see *National Climate Change* section for a discussion of past trends and future projections). An increase in average summer wave heights along the U.S. Atlantic coastline since 1975 has been attributed to a progressive increase in hurricane power.^{112,388} The intensity of hurricanes is likely to increase during this century with higher peak wind speeds, rainfall intensity, and storm surge height and strength.^{90,112} Even with no increase in hurricane intensity, coastal inundation and shoreline retreat would increase as sea-level rise accelerates, which is one of the most certain and most costly consequences of a warming climate.¹⁶⁴

L1 **Projected increases in air and water**
 L2 **temperatures will cause heat-related**
 L3 **stresses for people, plants, and animals.**
 L4

L5 The warming projected for the Southeast during
 L6 the next 50 to 100 years will create heat-related
 L7 stress for people, agricultural crops, livestock,
 L8 trees, transportation and other infrastructure, fish,
 L9 and wildlife. The average temperature change is
 L10 not as important for all of these sectors and natu-
 L11 ral systems as the projected increase in maximum
 L12 and minimum temperatures. Examples of potential
 L13 impacts include:

- L14
- L15 • Increased illness and death due to greater
 L16 summer heat stress, unless effective adaptation
 L17 measures are implemented.¹⁶⁴
- L18 • Decline in forest growth and agricultural crop
 L19 production due to the combined effects of ther-
 L20 mal stress and declining soil moisture.³⁹⁰
- L21 • Increased buckling of pavement and
 L22 railways.^{217,222}
- L23 • Decline in dissolved oxygen in stream, lakes,
 L24 and shallow aquatic habitats leading to fish
 L25 kills and loss of aquatic species diversity.
- L26 • Decline in production of cattle and other
 L27 rangeland livestock.³⁹¹ Significant impacts on
 L28 beef cattle occur at continuous temperatures
 L29 in the 90 to 100°F range, increasing in danger
 L30 as the humidity level increases (see *Agricul-*
 L31 *ture* sector).³⁹¹ Poultry and swine are primarily
 L32 raised in indoor operations, so warming would
 L33 increase energy requirements.¹⁹³

L34

L35 A reduction in very cold days is likely to reduce
 L36 the loss of human life due to cold-related stress,
 L37 while heat stress and related deaths in the sum-
 L38 mer months are likely to increase. The reduction
 L39 in cold-related deaths is not expected to offset the
 L40 increase in heat-related deaths (see *Human Health*
 L41 sector). Other effects of the projected increases in
 L42 temperature include more frequent outbreaks of
 L43 shellfish-borne diseases in coastal waters, altered
 L44 distribution of native plants and animals, local
 L45 loss of many threatened and endangered species,
 L46 displacement of native species by invasive species,
 L47 and more frequent and intense wildfires.

R1 **Decreased water availability is very**
 R2 **likely to impact the region’s economy as**
 R3 **well as its natural systems.**
 R4

R5 Decreased water availability due to increased
 R6 temperature and longer periods of time between
 R7 rainfall events, coupled with an increase in societal
 R8 demand is very likely to affect many sectors of the
 R9 Southeast’s economy. The amount and timing of
 R10 water available to natural systems also is affected
 R11 by climate change, as well as by human response
 R12 strategies such as increasing storage capacity
 R13 (dams)¹⁴² and increasing acreage of irrigated crop-
 R14 land.³⁹² The 2007 water shortage in the Atlanta re-
 R15 gion created serious conflicts between three states,
 R16 the U.S. Army Corps of Engineers (which operates
 R17 the dam at Lake Lanier), and the U.S. Fish and
 R18 Wildlife Service, which is charged with protecting
 R19 endangered species. As humans seek to adapt to
 R20 climate change by manipulating water resources,
 R21 streamflow and biological diversity are likely to be
 R22 reduced.¹⁴² During droughts, recharge of ground-
 R23 water will decline as the temperature and spacing
 R24 between rainfall events increases. Responding by
 R25 increasing groundwater pumping will further stress
 R26 or deplete aquifers and place increasing strain on
 R27 surface water resources. Increasing evaporation
 R28 and plant water loss rates alter the balance of runoff
 R29 and groundwater recharge, which is likely to lead
 R30 to saltwater intrusion into shallow aquifers in many
 R31 parts of the Southeast.¹⁴²



R32

R33

R34

R35

R36

R37

R38

R39

R40

R41

R42

R43

R44

R45

R46

R47

R48

R49

R50

In Atlanta and Athens, Georgia, 2007 was the second driest year on record. Among the numerous effects of the rainfall shortage were restrictions on water use in some cities and low water levels in area lakes. In the photo, a dock lies on dry land near Aqualand Marina on Lake Lanier (located northeast of Atlanta) in December 2007.

Land Lost During 2005 Hurricanes



In 2005, 217 square miles of land and wetlands were lost to open water during hurricanes Rita and Katrina. The photos and maps show the Chandeleur Islands, east of New Orleans, before and after the 2005 hurricanes; 85 percent of the islands' above-water land mass was eliminated.

Sea-level rise and the likely increase in hurricane intensity and associated storm surge will be among the most serious consequences of climate change.

An increase in average sea level of up to 2 feet and the likelihood of increased hurricane intensity and associated storm surge are likely to be among the most costly consequences of climate change for this region. As sea level rises, coastal shorelines will retreat. Wetlands will be inundated and eroded away, and low-lying areas including cities will be inundated more frequently – some permanently – by the advancing sea. Current buildings and infrastructure were not designed to withstand the intensity of the projected storm surge, which would cause catastrophic damage. As temperature increases and rainfall patterns change, soil moisture and runoff to the coast are likely to be more

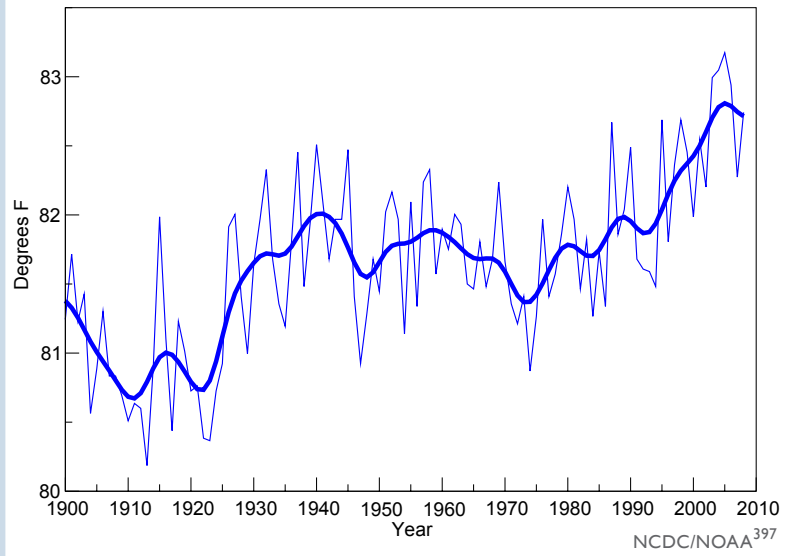
variable. The salinity of estuaries, coastal wetlands, and tidal rivers is likely to increase in the southeastern coastal zone, thereby altering coastal ecosystems and displacing them farther inland if no barriers exist. More frequent storm surge flooding and permanent inundation of coastal ecosystems and communities is likely in some low-lying areas, particularly along the central Gulf Coast where the land surface is sinking.^{393,394} Rapid acceleration in the rate of increase in sea-level rise could threaten a large portion of the Southeast coastal zone (see *Global Climate Change* section). The likelihood of a catastrophic increase in the rate of sea-level rise is dependent upon ice sheet response to warming, which is the subject of much scientific uncertainty.⁹⁰ Such rapid rise in sea level is likely to result in the crossing of thresholds, resulting in the destruction of barrier islands and wetlands.^{258,390}

L1
L2
L3
L4
L5
L6
L7
L8
L9
L10
L11
L12
L13
L14
L15
L16
L17
L18
L19
L20
L21
L22
L23
L24
L25
L26
L27
L28
L29
L30
L31
L32
L33
L34
L35
L36
L37
L38
L39
L40
L41
L42
L43
L44
L45
L46
L47
L48
L49
L50

R1
R2
R3
R4
R5
R6
R7
R8
R9
R10
R11
R12
R13
R14
R15
R16
R17
R18
R19
R20
R21
R22
R23
R24
R25
R26
R27
R28
R29
R30
R31
R32
R33
R34
R35
R36
R37
R38
R39
R40
R41
R42
R43
R44
R45
R46
R47
R48
R49
R50

L1 Compared to the present coastal situation, for
 L2 which vulnerability is quite high, an increase
 L3 in hurricane intensity will further affect low-
 L4 lying coastal ecosystems and coastal com-
 L5 munities along the Gulf and South Atlantic
 L6 coastal margin. An increase in intensity is
 L7 very likely to increase inland and coastal
 L8 flooding, coastal erosion rates, wind damage
 L9 to coastal forests, and wetland loss. Major
 L10 hurricanes also pose a severe risk to people,
 L11 personal property, and public infrastructure
 L12 in the Southeast, and this risk is likely to
 L13 be exacerbated.^{393,394} Hurricanes have their
 L14 greatest impact at the coastal margin where
 L15 they make landfall, causing storm surge,
 L16 severe beach erosion, inland flooding, and
 L17 wind-related casualties for both cultural and
 L18 natural resources. Some of these impacts
 L19 extend farther inland, affecting larger areas.
 L20 Recent examples of societal vulnerability
 L21 to severe hurricanes include Katrina and
 L22 Rita in 2005, which were responsible for
 L23 the loss of more than 1,800 lives and the net
 L24 loss of 217 square miles of low-lying coastal
 L25 marshes and barrier islands in southern
 L26 Louisiana.^{390,396}

Sea Surface Temperature
 Atlantic Hurricane Main Development Region
 August through October, 1900 to 2008



Ocean surface temperature during the peak hurricane season, August through October, in the main development region for Atlantic hurricanes.³⁹⁷ Higher sea surface temperatures in this region of the ocean have been associated with more intense hurricanes. As ocean temperatures continue to increase in the future, it is likely that hurricane rainfall and wind speeds will increase in response to human-caused warming (see *National Climate Change* section).⁶⁸

Ecological thresholds are likely to be crossed throughout the region, causing major disruptions to ecosystems and to the benefits they provide to people.

Ecological systems provide numerous important services that have high economic and cultural value in the Southeast. Ecological effects cascade among both living and physical systems, as illustrated in the following examples of ecological disturbances that result in abrupt responses, as opposed to gradual and proportional responses to warming:

- The sudden loss of coastal landforms that serve as a storm-surge barrier for natural resources and as a homeland for coastal communities (such as in a major hurricane).^{255,390}
- An increase in sea level can have no apparent effect until an elevation is reached that allows widespread, rapid saltwater intrusion into coastal forests and freshwater aquifers.³⁹⁸
- Lower soil moisture and higher temperatures leading to intense wildfires or pest outbreaks (such as the southern pine beetle) in southeastern forests;³⁹⁹ intense droughts leading to the drying of lakes, ponds, and wetlands; and the local or global extinction of riparian and aquatic species.¹⁴²



Flooding damage in Louisiana due to Hurricane Katrina

R1
R2
R3
R4
R5
R6
R7
R8
R9
R10
R11
R12
R13
R14
R15
R16
R17
R18
R19
R20
R21
R22
R23
R24
R25
R26
R27
R28
R29
R30
R31
R32
R33
R34
R35
R36
R37
R38
R39
R40
R41
R42
R43
R44
R45
R46
R47
R48
R49
R50

- A precipitous decline of wetland-dependent coastal fish and shellfish populations due to the rapid loss of coastal marsh.⁴⁰⁰

Quality of life will be affected by increasing heat stress, water scarcity, severe weather events, and reduced availability of insurance for at-risk properties.

Over the past century, the southeastern “sunbelt” has attracted people, industry, and investment. The population of Florida more than doubled during the

past three decades, and growth rates in most other southeastern states were in the range of 45 to 75 percent. Future population growth and the quality of life for existing residents is likely to be affected by the many challenges associated with climate change, such as reduced insurance availability, increased insurance cost, and increases in water scarcity, sea-level rise, extreme weather events, and heat stress.

Adaptation: Reducing Exposure to Flooding and Storm Surge

Three different types of adaptation to sea-level rise are available for low-lying coastal areas.^{173,270} One is to move buildings and infrastructure further inland to get out of the way of the rising sea. Another is to accommodate rising water through changes in building design and construction, such as elevating buildings on stilts. Flood insurance programs even require this in some areas with high probabilities of floods. The third adaptation option is to try to protect existing development by building levees and river flood control structures. This option is being pursued in some highly vulnerable areas of the Gulf and South Atlantic coasts. Flood control structures can be designed to be effective in the face of higher sea level and storm surge. Some hurricane levees and floodwalls were not just replaced after Hurricane Katrina, they were redesigned to withstand higher storm surge and wave action.⁴⁰¹

The costs and environmental impacts of building such structures can be significant. Furthermore, building levees can actually increase future risks. This is sometimes referred to as the levee effect or the safe-development paradox. Levees that provide protection from, for example, the storm surge from a Category 3 hurricane, increase real and perceived safety and thereby lead to increased development. This increased development means there will be greater damage if and when the storm surge from a Category 5 hurricane tops the levee than there would have been if no levee had been constructed.²⁵³

In addition to levees, enhancement of key highways used as hurricane evacuation routes and improved hurricane evacuation planning is a common adaptation underway in all Gulf Coast states.²¹⁷ Other protection options that are being practiced along low-lying coasts include the enhancement and protection of natural features such as forested wetlands, saltmarshes, and barrier islands.³⁹⁰



Recent upgrades that raised the height of this earthen levee increased protection against storm surge in the New Orleans area.

L1
L2
L3
L4
L5
L6
L7
L8
L9
L10
L11
L12
L13
L14
L15
L16
L17
L18
L19
L20
L21
L22
L23
L24
L25
L26
L27
L28
L29
L30
L31
L32
L33
L34
L35
L36
L37
L38
L39
L40
L41
L42
L43
L44
L45
L46
L47
L48
L49
L50

R1
R2
R3
R4
R5
R6
R7
R8
R9
R10
R11
R12
R13
R14
R15
R16
R17
R18
R19
R20
R21
R22
R23
R24
R25
R26
R27
R28
R29
R30
R31
R32
R33
R34
R35
R36
R37
R38
R39
R40
R41
R42
R43
R44
R45
R46
R47
R48
R49
R50





Midwest

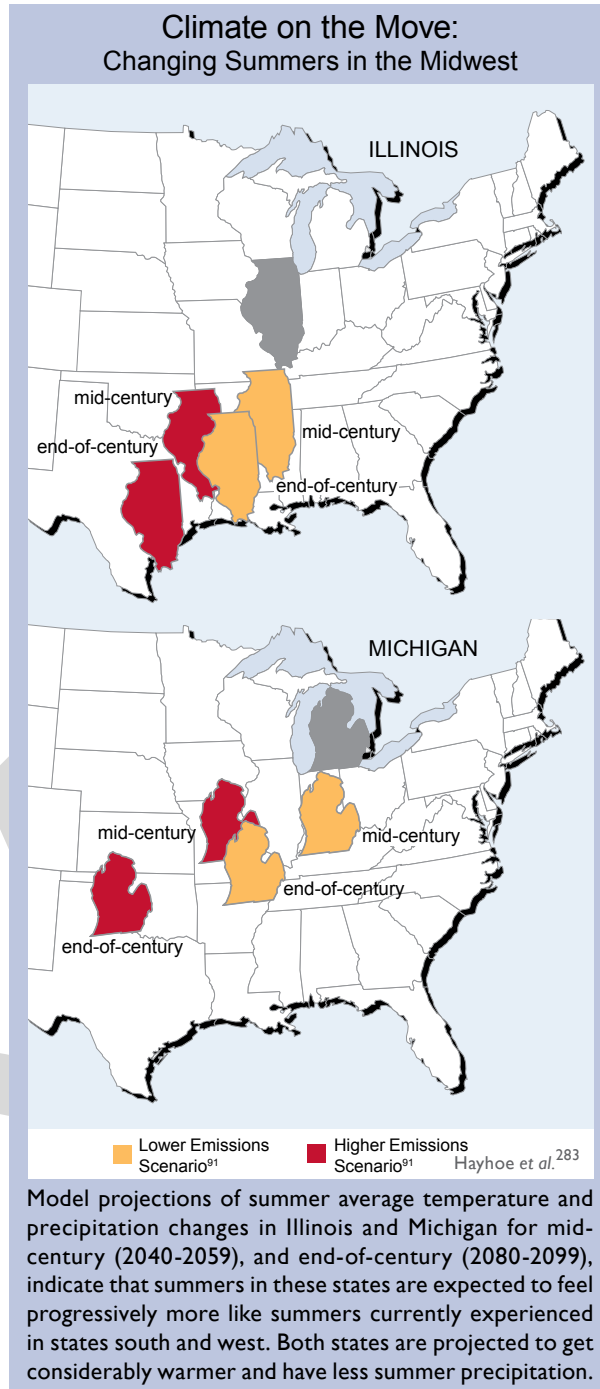
L1
L2
L3
L4
L5
L6
L7
L8
L9
L10
L11
L12
L13
L14
L15
L16
L17
L18
L19
L20
L21
L22
L23
L24
L25
L26
L27
L28
L29
L30
L31
L32
L33
L34
L35
L36
L37
L38
L39
L40
L41
L42
L43
L44
L45
L46
L47
L48
L49
L50

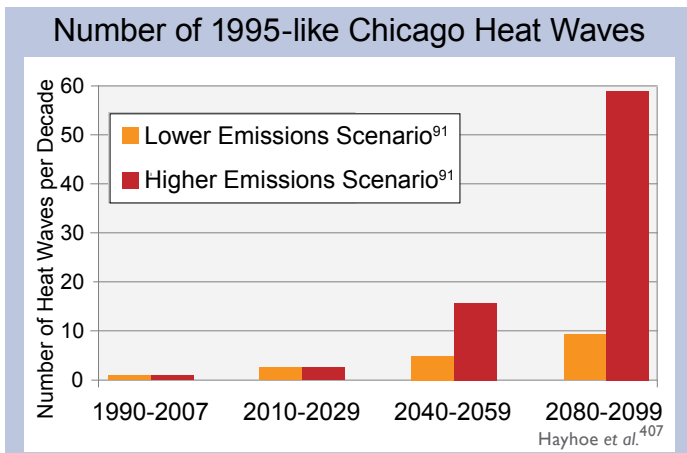
R1
R2
R3
R4
R5
R6
R7
R8
R9
R10
R11
R12
R13
R14
R15
R16
R17
R18
R19
R20
R21
R22
R23
R24
R25
R26
R27
R28
R29
R30
R31
R32
R33
R34
R35
R36
R37
R38
R39
R40
R41
R42
R43
R44
R45
R46
R47
R48
R49
R50

The Midwest’s climate is shaped by the presence of the Great Lakes and the region’s location in the middle of the North American continent. This location, far from the temperature-moderating effects of the oceans, contributes to large seasonal swings in air temperature from hot, humid summers to cold winters. In recent decades, a noticeable increase in average temperatures in the Midwest has been observed, despite the strong year-to-year variations. The largest increase has been measured in winter, extending the length of the frost-free or growing season by more than one week, mainly due to earlier dates for the last spring frost. Heavy downpours are now twice as frequent as they were a century ago. Both summer and winter precipitation have been above average for the last three decades, the wettest period in a century. The Midwest has experienced two record-breaking floods in the past 15 years. There has also been a decrease in lake ice, including on the Great Lakes. Since the 1980s, large heat waves have been more frequent in the Midwest than anytime in the last century, other than the Dust Bowl years of the 1930s.^{112,283,402-404}

During the summer, public health and quality of life, especially in cities, will be negatively affected by increasing heat waves, reduced air quality, and insect and waterborne diseases. In the winter, warming will have mixed impacts.

Heat waves that are more frequent, more severe, and longer-lasting are projected. The frequency of hot days and the length of the heat-wave season both will be more than twice as great under the higher emissions scenario⁹¹ compared to the lower emissions scenario.^{91,283, 402,403,405} Events such as the Chicago heat wave of 1995, which resulted in over 700 deaths, will become more common. Under the lower emissions scenario,⁹¹ such a heat wave is projected to occur every other year in Chicago by the end of the century, while under the higher emissions scenario,⁹¹ there would be about three such heat waves per year. Even more severe heat waves, such as the one that claimed tens of thousands of lives in Europe in 2003, are projected to become more frequent in a warmer world, occurring as often as every other year in the Midwest by the end of this century under the higher emissions scenario.^{91,283,403,406} Some health impacts can be reduced by better preparation for such events.²⁸⁸





By the end of the century, heat waves like the one that occurred in Chicago in 1995 are projected to occur every other year under the lower emissions scenario.⁹¹ Under the higher emissions scenario,⁹¹ such events are projected to occur more than three times every year. In this analysis, heat waves were defined as at least one week of daily maximum temperatures greater than 90°F and nighttime minimum temperatures greater than 70°F, with at least two consecutive days with daily temperatures greater than 100°F and nighttime temperatures greater than 80°F.

During heat waves, high electricity demand combines with climate-related limitations on energy production capabilities (see *Energy Production and Use* sector), increasing the likelihood of electricity shortages and resulting in brownouts or even blackouts. This combination can leave people without air conditioning and ventilation when they need it most, as occurred during the 1995 Chicago/Milwaukee heat wave. In general, electricity demand for air conditioning is projected to significantly in-

crease in summer. Improved energy planning could reduce electricity disruptions.

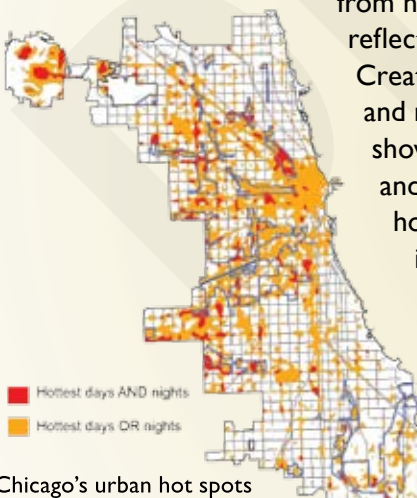
The urban heat island effect can further add to the local daytime and nighttime temperatures (see *Human Health* sector). Heat waves take a greater toll in illness and death when there is little relief from the heat at night.

Another health-related issue arises from the fact that climate change can affect air quality. A warmer climate generally means more ground-level ozone (a component of smog), which can cause respiratory problems, especially for those who are young, old, or have asthma or allergies. Unless the emissions of pollutants that lead to ozone formation are reduced significantly, there will be more ground-level ozone as a result of the projected climate changes in the Midwest due to increased air temperatures, more stagnant air, and increased emissions from vegetation.^{283,291,402,403,408-410}

Insects such as ticks and mosquitoes that carry diseases will survive winters more easily and produce larger populations in a warmer Midwest.^{283,402,403} One potential risk is an increasing incidence of diseases such as West Nile virus. Waterborne diseases will present an increasing risk to public health because so many pathogens thrive in warmer conditions.¹⁶³

Adaptation: Chicago Tries to Cool the Urban Heat Island

Efforts to reduce urban heat island effects become even more important in a warming climate. The City of Chicago has produced a map of urban hotspots to use as a planning tool to target areas that could most benefit from heat-island reduction initiatives such as reflective or green roofing, and tree planting. Created using satellite images of daytime and nighttime temperatures, the map shows the hottest 10 percent of both day and night temperatures in red, and the hottest 10 percent of either day or night in orange.



“Green roofs” are cooler than the surrounding conventional roofs.

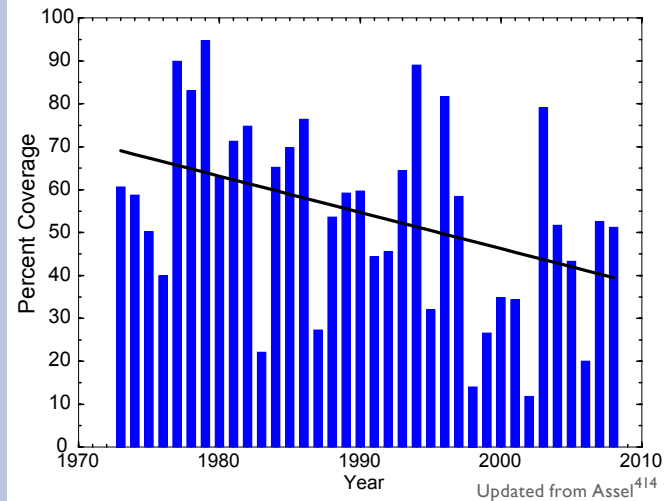
The City is working to reduce urban-heat buildup and the need for air conditioning by using reflective roofing materials. This thermal image shows that the radiating temperature of the City Hall’s “green roof”—covered with soil and vegetation—is up to 77°F cooler than the nearby conventional roofs.⁴¹¹

L1 In winter, oil and gas demand for heating will
 L2 decline. Warming will also decrease the number
 L3 of days with snow on the ground, which is
 L4 expected to improve traffic safety.²²² On the
 L5 other hand, warming will decrease outdoor
 L6 winter recreational opportunities.

L7
 L8
 L9 **Significant reductions in Great Lakes water levels, which are projected under higher emissions scenarios⁹¹, lead to impacts on shipping, infrastructure, beaches, and ecosystems.**
 L10
 L11
 L12
 L13
 L14
 L15

L16 The Great Lakes are a natural resource of tremendous significance, containing 20 percent of the planet's fresh surface water and serving as the dominant feature of the industrial heartland of the nation. Higher temperatures will mean more evaporation and hence a likely reduction in the Great Lakes water levels. Reduced lake ice increases evaporation in winter, contributing to the decline. Under a lower emissions scenario,⁹¹ water levels in the Great Lakes are projected to fall no

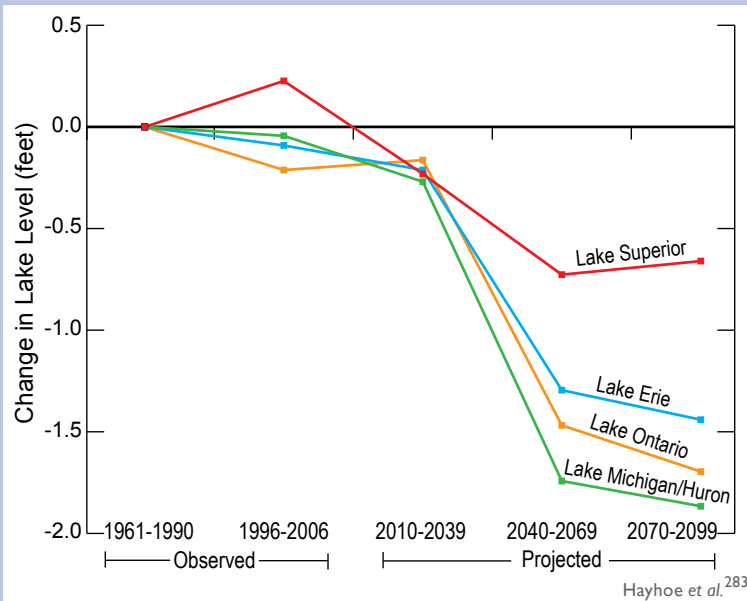
Observed Changes in Great Lakes Ice Cover
 Seasonal Maximum Coverage, 1973 to 2008



Reductions in winter ice cover lead to more evaporation, causing lake levels to drop even farther. While the graph indicates large year-to-year variations, there is a clear decrease in the extent of Great Lakes ice coverage, as shown by the black trend line.

more than 1 foot by the end of the century, but under a higher emissions scenario,⁹¹ they are projected to fall between 1 and 2 feet.²⁸³ The greater the temperature rise, the higher the likelihood of a larger decrease in lake levels.⁴¹² Even a decrease of 1 foot, combined with normal fluctuations, can result in significant lengthening of the distance to the lakeshore in many places. There are also potential impacts on beaches, coastal ecosystems, dredging requirements, infrastructure, and shipping. For example, lower lake levels reduce “draft,” or the distance between the waterline and the bottom of a ship, which lessens a ship’s ability to carry freight. Large vessels, sized for passage through the St. Lawrence Seaway, lose up to 240 tons of capacity for each inch of draft lost.^{283,402,403,413} These impacts will have costs, including increased shipping, repair and maintenance costs, and lost recreation and tourism dollars.

Projected Changes in Great Lakes Levels under Higher Emissions Scenario⁹¹

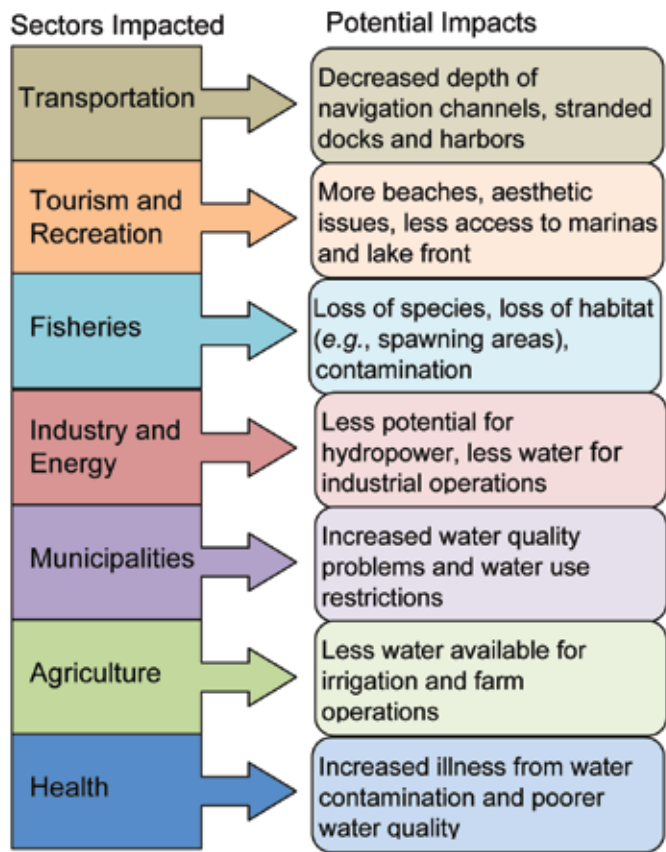


Average Great Lakes levels depend on the balance between precipitation (and corresponding runoff) in the Great Lakes Basin on one hand and evaporation and outflow on the other. As a result, lower emissions scenarios⁹¹ with less warming show less reduction in lake levels than higher emissions scenarios.⁹¹ Projected changes in lake levels are based on simulations by the NOAA Great Lakes model for projected climate changes under a higher emissions scenario.⁹¹

R1
 R2
 R3
 R4
 R5
 R6
 R7
 R8
 R9
 R10
 R11
 R12
 R13
 R14
 R15
 R16
 R17
 R18
 R19
 R20
 R21
 R22
 R23
 R24
 R25
 R26
 R27
 R28
 R29
 R30
 R31
 R32
 R33
 R34
 R35
 R36
 R37
 R38
 R39
 R40
 R41
 R42
 R43
 R44
 R45
 R46
 R47
 R48
 R49
 R50



Lower Water Levels in the Great Lakes



Adapted from Field et al.¹⁶⁴

Reduced water levels in the Great Lakes will have interconnected impacts across many sectors, creating mismatches between water supply and demand, and necessitating trade-offs. Regions outside the Midwest will also be affected. For example, a reduction in hydropower potential would affect the Northeast, and a reduction in irrigation water would affect regions that depend on agricultural produce from the Midwest.

The likely increase in precipitation in winter and spring, more heavy downpours, and greater evaporation in summer would lead to more periods of both floods and water deficits.

Precipitation is projected to increase in winter and spring, and to become more intense throughout the year. This pattern is expected to lead to more frequent flooding, increasing infrastructure damage, and impacts on human health. Such heavy downpours can overload drainage systems and water treatment facilities, increasing the risk of water-borne diseases. Such an incident occurred in Milwaukee in 1993 when the water supply was contaminated with the parasite *Cryptosporidium*, causing 403,000 reported cases of gastrointestinal illness and 54 deaths.

In Chicago, rainfall of more than 2.5 inches per day is an approximate threshold beyond which combined water and sewer systems overflow into Lake Michigan (such events occurred 2.5 times per decade from 1961 to 1990). This generally results in beach closures to reduce the risk of disease transmission. Rainfall above this threshold is projected to occur twice as often by the end of this century under the lower emissions scenario⁹¹ and three times as often under the higher emissions scenario.^{91,283,403} Similar increases are expected across the Midwest.



The Great Flood of 1993 caused flooding along 500 miles of the Mississippi and Missouri river systems. The photo shows its effects on U.S. Highway 54, just north of Jefferson City, Missouri.

More intense rainfall can lead to floods that cause significant impacts regionally and even nationally. For example, the Great Flood of 1993 caused catastrophic flooding along 500 miles of the Mississippi and Missouri river systems, affecting one-quarter of all U.S. freight (see *Transportation* sector).^{222,415-417} Another example was a record-breaking 24-hour rainstorm in July 1996, which resulted in flash flooding in Chicago and its suburbs, causing extensive damage and disruptions, with some commuters not being able to reach Chicago for three days (see *Transportation* sector).²²² There was also a record-breaking storm in August 2007. Increases in such events

L1
L2
L3
L4
L5
L6
L7
L8
L9
L10
L11
L12
L13
L14
L15
L16
L17
L18
L19
L20
L21
L22
L23
L24
L25
L26
L27
L28
L29
L30
L31
L32
L33
L34
L35
L36
L37
L38
L39
L40
L41
L42
L43
L44
L45
L46
L47
L48
L49
L50

R1
R2
R3
R4
R5
R6
R7
R8
R9
R10
R11
R12
R13
R14
R15
R16
R17
R18
R19
R20
R21
R22
R23
R24
R25
R26
R27
R28
R29
R30
R31
R32
R33
R34
R35
R36
R37
R38
R39
R40
R41
R42
R43
R44
R45
R46
R47
R48
R49
R50

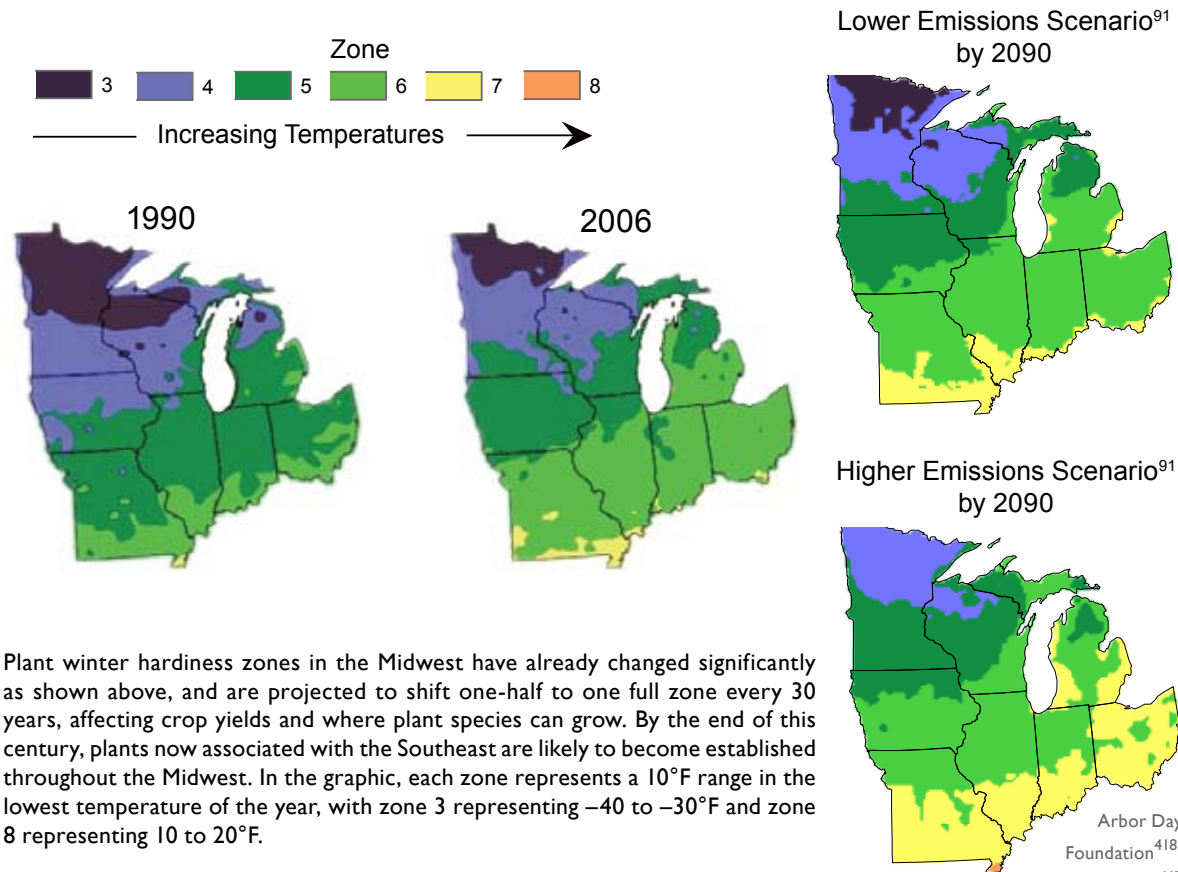
L1 are likely to cause greater property damage, higher
 L2 insurance rates, a heavier burden on emergency
 L3 management, increased clean-up and rebuilding
 L4 costs, and a growing financial toll on businesses,
 L5 homeowners, and insurers.
 L6

L7 In the summer, with increasing evaporation rates
 L8 and longer periods between rainfalls, the likelihood
 L9 of drought will increase and water levels in rivers,
 L10 streams, and wetlands are likely to decline. Lower
 L11 water levels also could create problems for river
 L12 traffic, reminiscent of the stranding of more than
 L13 4,000 barges on the Mississippi River during the
 L14 1988 drought. Reduced summer water levels are
 L15 also likely to reduce the recharge of groundwater,
 L16 cause small streams to dry up (reducing native fish
 L17 populations), and reduce the area of wetlands in the
 L18 Midwest.
 L19
 L20
 L21
 L22
 L23

While the longer growing season provides the potential for increased crop yields, increases in heat waves, floods, droughts, insects, and weeds will present increasing challenges to managing crops, livestock, and forests.

The projected increase in winter and spring precipitation and flooding is likely to delay planting and crop establishment. Longer growing seasons and increased carbon dioxide have positive effects on some crop yields, but this is likely to be counterbalanced in part by the negative effects of additional disease-causing pathogens, insect pests, and weeds (including invasive weeds).¹⁹³ Livestock production is expected to become more costly as higher temperatures stress livestock, decreasing productivity and increasing costs associated with the needed ventilation and cooling equipment.¹⁹³

Observed and Projected Changes in Plant Hardiness Zones



R1
 R2
 R3
 R4
 R5
 R6
 R7
 R8
 R9
 R10
 R11
 R12
 R13
 R14
 R15
 R16
 R17
 R18
 R19
 R20
 R21
 R22
 R23
 R24
 R25
 R26
 R27
 R28
 R29
 R30
 R31
 R32
 R33
 R34
 R35
 R36
 R37
 R38
 R39
 R40
 R41
 R42
 R43
 R44
 R45
 R46
 R47
 R48
 R49
 R50

L1 Plant winter hardiness zones (each zone represents
 L2 a 10°F change in minimum temperature) in the
 L3 Midwest are likely to shift one-half to one full zone
 L4 about every 30 years. By the end of the century,
 L5 plants now associated with the Southeast are likely
 L6 to become established throughout the Midwest.
 L7 Impacts on forests are likely to be mixed, with the
 L8 positive effects of higher carbon dioxide and nitro-
 L9 gen levels acting as fertilizers potentially negated
 L10 by the negative effects of decreasing air quality.²⁴⁴
 L11 In addition, more frequent droughts, and hence fire
 L12 hazards, and an increase in destructive insect pests,
 L13 such as gypsy moths, hinder plant growth. Insects,
 L14 historically controlled by cold winters, more easily
 L15 survive milder winters and produce larger popula-
 L16 tions in a warmer climate (see *Agriculture* sector).

L19 **Native species are very likely to face**
 L20 **increasing threats from rapidly changing**
 L21 **climate conditions, pests, diseases, and**
 L22 **invasive species moving in from warmer**
 L23 **regions.**

L25 As air temperatures increase, so will water temper-
 L26 atures. This will lead to an earlier and longer period
 L27 in summer during which mixing of the relatively
 L28 warm surface lake water with the colder water
 L29 below is reduced. This stratification effectively cuts
 L30 off oxygen from bottom layers, increasing the risk
 L31 of oxygen-poor or oxygen-free “dead zones” that
 L32 kill fish and other living things. Warmer water and
 L33 low-oxygen conditions in the bottom layer of lakes
 L34 also mobilize mercury and other contaminants
 L35 in lake sediments. These increasing quantities of
 L36 contaminants will be taken up in the aquatic food
 L37 chain, adding to the existing health hazard for spe-
 L38 cies that eat fish from the lakes, including people.



Populations of coldwater fish, such as brook trout,
 lake trout, and whitefish, are expected to decline
 dramatically, while populations of coolwater fish
 such as muskie, and warmwater species such as
 small-mouth bass and bluegill, will take their place.
 Aquatic ecosystem disruptions are likely to be
 compounded by invasions by non-native species,
 which tend to thrive under a wide range of environ-
 mental conditions. Native species, adapted to a nar-
 rower range of conditions, are expected to decline.

All major groups of animals, including birds,
 mammals, amphibians, reptiles, and insects, will
 be affected by impacts on local populations, and
 by competition from other species moving into the
 Midwest region.⁷⁰ The potential for animals to shift
 their range to keep pace with the changing climate
 will be inhibited by major urban areas and the pres-
 ence of the Great Lakes.





Great Plains

L1
L2
L3
L4
L5
L6
L7
L8
L9
L10
L11
L12
L13
L14
L15
L16
L17
L18
L19
L20
L21
L22
L23
L24
L25
L26
L27
L28
L29
L30
L31
L32
L33
L34
L35
L36
L37
L38
L39
L40
L41
L42
L43
L44
L45
L46
L47
L48
L49
L50

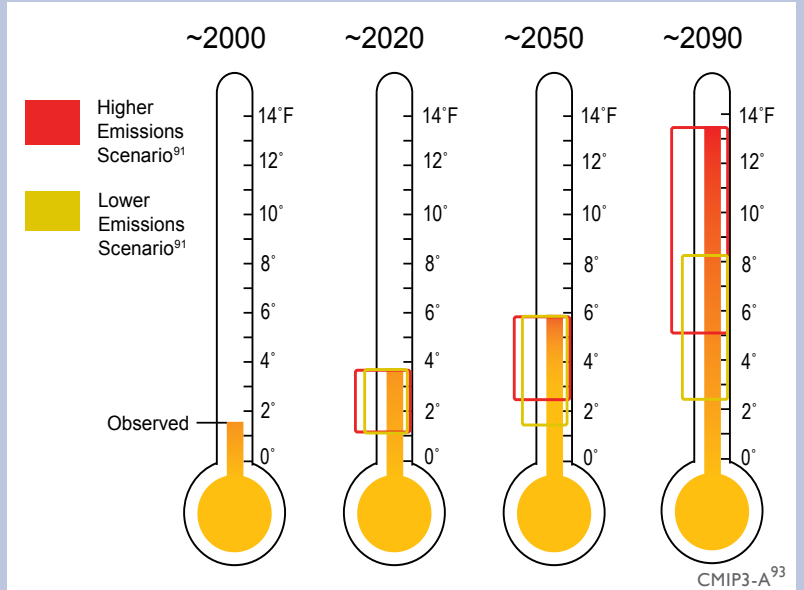
R1
R2
R3
R4
R5
R6
R7
R8
R9
R10
R11
R12
R13
R14
R15
R16
R17
R18
R19
R20
R21
R22
R23
R24
R25
R26
R27
R28
R29
R30
R31
R32
R33
R34
R35
R36
R37
R38
R39
R40
R41
R42
R43
R44
R45
R46
R47
R48
R49
R50

The Great Plains is characterized by strong seasonal climate variations. Over thousands of years, records preserved in tree rings, sediments, and sand deposits provide evidence of recurring periods of extended drought (such as the Dust Bowl of the 1930s) alternating with wetter conditions.^{97,419}

Today, semi-arid conditions in the western Great Plains gradually transition to a moister climate in the eastern parts of the region. To the north, winter days in North Dakota average 25°F, while it is not unusual to have a West Texas winter day over 75°F. In West Texas, there are between 70 and 100 days per year over 90°F, whereas North Dakota has only 10 to 20 such days on average.

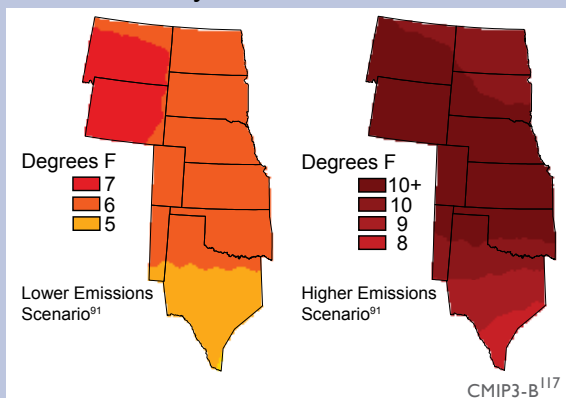
Significant trends in regional climate are apparent over the last few decades. Average temperatures have increased throughout the region, with the largest changes occurring in winter months and over the northern states. Relatively cold days are becoming less frequent and relatively hot days more frequent.⁴²⁰ Precipitation also has increased over most of the area.^{149,421}

Observed and Projected Temperature Rise



The average temperature in the Great Plains already has increased roughly 1.5°F relative to a 1960s and 1970s baseline. By the end of the century, temperatures are projected to continue to increase by 2.5°F to more than 13°F compared with the 1960 to 1979 baseline, depending on future emissions of heat-trapping gases. The brackets on the thermometers represent the likely range of model projections, though lower or higher outcomes are possible.

Summer Temperature Change by 2080-2099



Temperatures in the Great Plains are projected to increase significantly by the end of this century, with the northern part of the region experiencing the greatest projected increase in temperature.

Temperatures are projected to continue to increase over this century, with larger changes expected under scenarios of higher heat-trapping emissions as compared to lower heat-trapping emissions. Summer changes are projected to be larger than those in winter in the southern and central Great Plains.¹⁰⁸ Precipitation is also projected to change, particularly in winter and spring. Conditions are anticipated to become wetter in the north and drier in the south.

Projected changes in long-term climate and more frequent extreme events such as heat waves, droughts, and heavy rainfall will affect many aspects of life in the Great Plains. These include the region's already threatened water resources, essential agricultural and ranching activities, unique natural and protected areas, and the health and prosperity of its inhabitants.

Projected increases in temperature, evaporation, and drought frequency add to concerns about the region’s declining water resources.

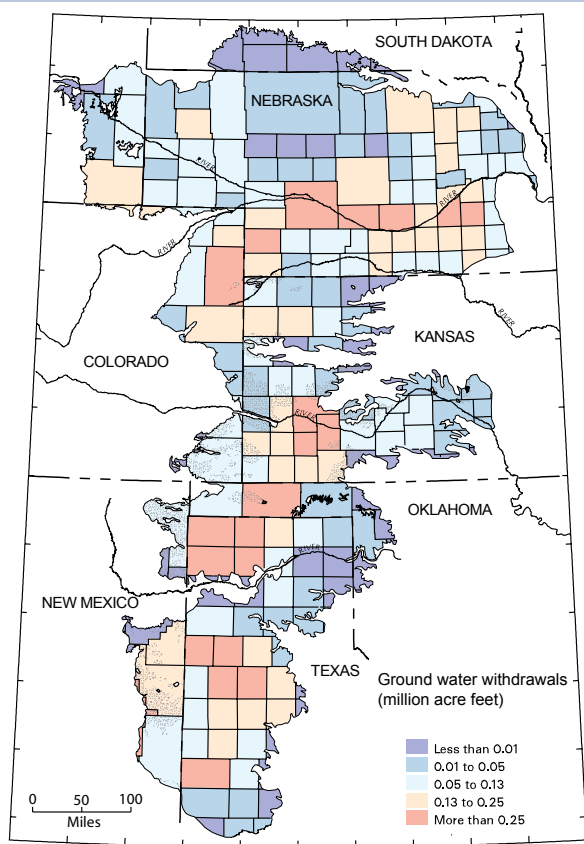
Water is the most important factor affecting activities on the Great Plains. Most of the water used in the Great Plains comes from the High Plains aquifer (sometimes referred to by the name of its largest formation, the Ogallala aquifer), which stretches from South Dakota to Texas. The aquifer holds both current recharge from precipitation and so-called “ancient” water, water trapped by silt and soil washed down from the Rocky Mountains during the last ice age.

As population increased in the Great Plains and irrigation became widespread, annual withdrawals began to outpace natural recharge.⁴²² Today, an

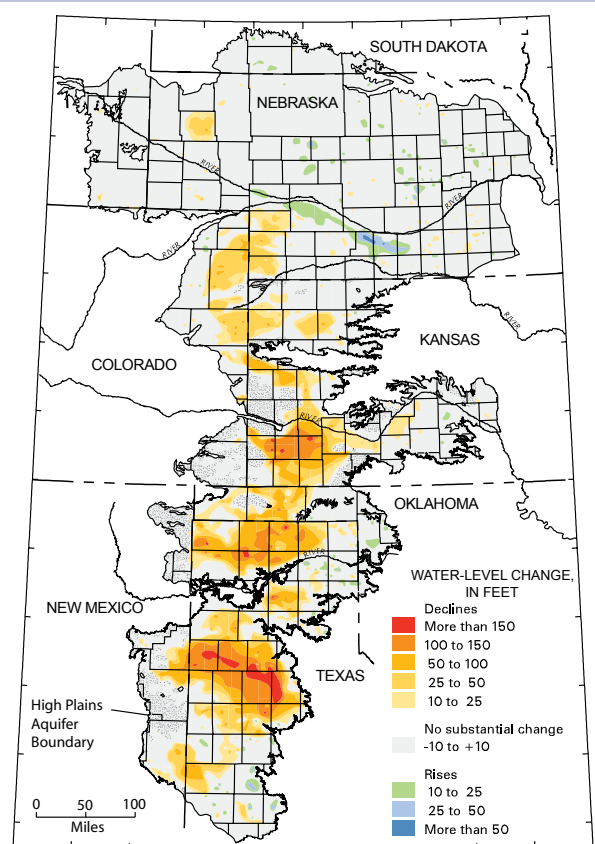
average of 19 billion gallons of groundwater are pumped from the aquifer each day. This water irrigates 13 million acres of land and provides drinking water to over 80 percent of the region’s population.⁴²³ Since 1950, aquifer water levels have dropped an average of 13 feet, equivalent to a 9 percent decrease in aquifer storage. In heavily irrigated parts of Texas, Oklahoma, and Kansas, reductions are much larger, from 100 feet to over 250 feet.

Projections of increasing temperatures, faster evaporation rates, and more sustained droughts brought on by climate change will only add more stress to overtaxed water sources.^{149,254,424,425} Current water use on the Great Plains is unsustainable, as the High Plains aquifer continues to be tapped faster than the rate of recharge.

Groundwater Withdrawals for Irrigation 1950 to 2005



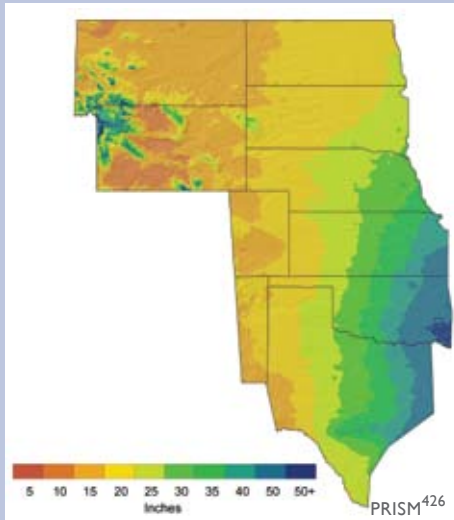
Water Level Changes in the High Plains Aquifer 1950 to 2005



McGuire⁴²²

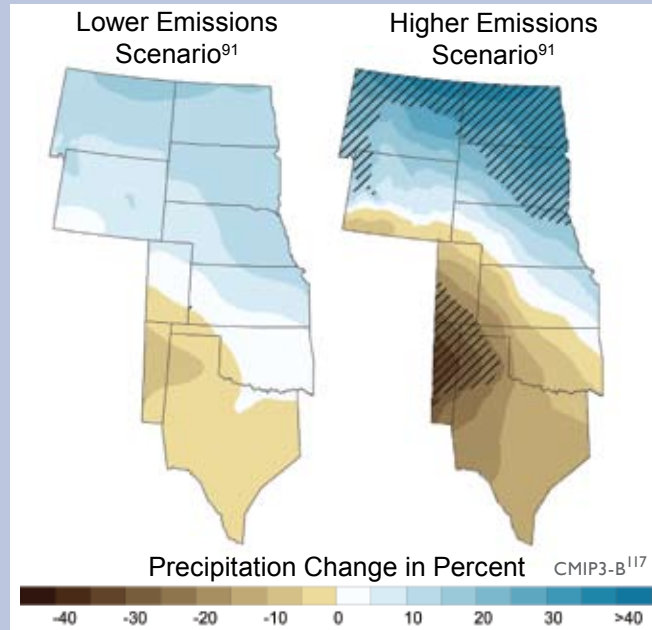
Irrigation is one of the main factors stressing water resources in the Great Plains. In parts of the region, more than 81 trillion gallons of water (pink areas on the irrigation map) were withdrawn for irrigation in Texas, Oklahoma, and Kansas from 1950 to 2005. During the same time period, water levels in parts of the High Plains aquifer in those states decreased by more than 150 feet (red areas on the water level change map).

Average Annual Observed Precipitation 1971-2000



The Great Plains currently experiences a sharp precipitation gradient from east to west, from more than 50 inches of precipitation per year in eastern Oklahoma and Texas to less than 10 inches in some of the western parts of the region.

Projected Spring Precipitation Change by 2080s-2090s



Northern areas of the Great Plains are projected to experience a wetter climate by the end of this century, while southern areas are projected to experience a drier climate. The change in precipitation is compared with a 1960-1979 baseline. Confidence in the projected changes is highest in the hatched areas.

The Dust Bowl: Combined Effects of Land Use and Climate

Over the past century, large-scale conversion of grasslands to crops and rangeland has altered the natural environment of the Great Plains.¹⁴⁹ Irrigated fields have increased evaporation rates, reducing summer temperatures and increasing local precipitation.^{427,428}

The Dust Bowl of the 1930s epitomizes what can happen as a result of interactions between climate and human activity. In the 1920s, increasing demand for food encouraged poor agricultural practices. Small-scale producers ploughed under native grasses to plant wheat, removing the protective cover the land required to retain its moisture.



Dust bowl of 1935 in Stratford, Texas

Variations in ocean temperature contributed to a slight increase in air temperatures, just enough to disrupt the winds that typically draw moisture from the south into the Great Plains. As the intensively tilled soils dried up, topsoil from an estimated 100 million acres of the Great Plains blew across the continent.

The Dust Bowl dramatically demonstrated the potentially devastating effects of poor land-use practices combined with climate variability and change.⁴²⁹

A similar trend is apparent today. Water is being pumped from the Ogallala aquifer faster than it can recharge. In many areas, playa lakes are poorly managed (see page 127). Existing stresses on water resources in the Great Plains due to unsustainable water usage are likely to be exacerbated by future changes in temperature and precipitation, this time largely due to human-induced climate change.

Agriculture, ranching, and natural lands, already under pressure due to an increasingly limited water supply, are very likely to also be stressed by rising temperatures.

Agricultural, range, and croplands cover more than 70 percent of the Great Plains, producing wheat, hay, corn, barley, cattle, and cotton. Agriculture is fundamentally sensitive to climate. Heat and water stress from droughts and heat waves can decrease yields and wither crops.^{430,431} The influence of long-term trends in temperature and precipitation can be just as great.⁴³¹

As temperatures increase over the coming century, optimal zones for growing particular crops will shift. Pests that were historically unable to survive in the Great Plains’ cooler areas are expected to spread northward. Milder winters and earlier springs also will encourage greater numbers and earlier emergence of insects.¹⁴⁹ Rising carbon dioxide levels in the atmosphere can increase crop growth, but also make some types of weeds grow even faster.⁴³²

Projected increases in precipitation are unlikely to be sufficient to offset decreasing soil moisture and water availability in the Great Plains due to rising temperatures and aquifer depletion. In some areas, there is not expected to be enough water for agriculture to sustain even current usage.

With limited water supply comes an increased vulnerability of agriculture to climate change. Further stresses on water supply for agriculture and ranching are likely as the region’s cities continue to grow, increasing competition between urban and rural users.⁴³³ The largest impacts are expected in heavily irrigated areas in the southern Great Plains, already plagued by unsustainable water use and greater frequency of extreme heat.¹⁴⁹

Successful adaptation will require diversification of crops and livestock, as well as transitions from irrigated to rain-fed agriculture.⁴³⁴⁻⁴³⁶ Producers who can adapt to changing climate conditions are likely to see their businesses survive; some might even thrive. Others, without resources or ability to adapt effectively, will lose out.

Climate change is likely to affect native plant and animal species by altering key habitats such as the wetland ecosystems known as prairie potholes or playa lakes.

Ten percent of the Great Plains is protected lands, home to unique ecosystems and wildlife. The region is a haven for hunters and anglers, with its ample supplies of wild game such as moose, elk, and deer; birds such as goose, quail, and duck; and fish such as walleye and bass.

Climate-driven changes are likely to combine with human stresses to further increase the vulnerability of natural ecosystems to pests, invasive species, and loss of native species. Changes in temperature and precipitation affect the composition and diversity of native animals and plants through altering their breeding patterns, water and food supply, and habitat availability.¹⁴⁹ In a changing climate, populations of some pests such as red fire ants and rodents, better adapted to a warmer climate, are projected to increase.^{437,438} Grassland and plains birds, already besieged by habitat fragmentation, could experience significant shifts and reductions in their range.⁴³⁹

Urban sprawl, agriculture, and ranching practices already threaten the Great Plains’ distinctive wetlands. Many of these are home to endangered and iconic species. In particular, prairie wetland ecosystems provide crucial habitat for migratory waterfowl and shorebirds.



Mallard ducks are one of the many species that inhabit the playa lakes, also known as prairie potholes.

Ongoing shifts in the region’s population from rural areas to urban centers will interact with a changing climate, resulting in a variety of consequences.

Inhabitants of the Great Plains include a rising number of urban dwellers, a long tradition of rural communities, and extensive Native American

L1
L2
L3
L4
L5
L6
L7
L8
L9
L10
L11
L12
L13
L14
L15
L16
L17
L18
L19
L20
L21
L22
L23
L24
L25
L26
L27
L28
L29
L30
L31
L32
L33
L34
L35
L36
L37
L38
L39
L40
L41
L42
L43
L44
L45
L46
L47
L48
L49
L50

R1
R2
R3
R4
R5
R6
R7
R8
R9
R10
R11
R12
R13
R14
R15
R16
R17
R18
R19
R20
R21
R22
R23
R24
R25
R26
R27
R28
R29
R30
R31
R32
R33
R34
R35
R36
R37
R38
R39
R40
R41
R42
R43
R44
R45
R46
R47
R48
R49
R50

Playa Lakes and Prairie Potholes

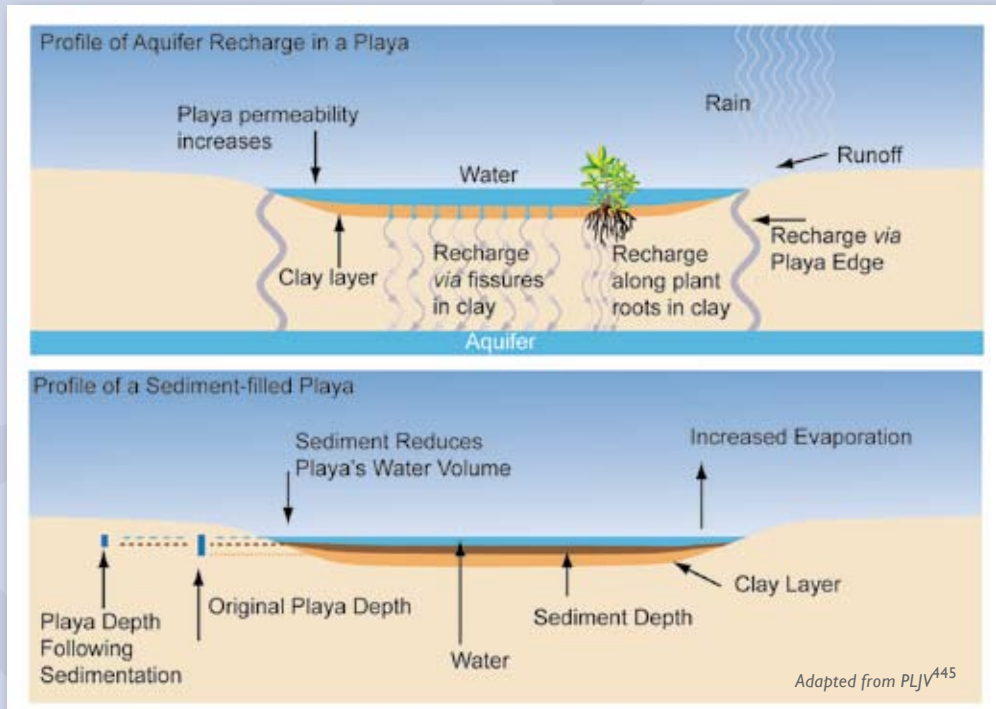
Shallow ephemeral lakes dot the Great Plains, anomalies of water in the arid landscape. In the north they are known as prairie potholes; in the south, playa lakes. Playa lakes create unique microclimates that support diverse wildlife and plant communities. A playa can lie with little or no water for long periods, or have several wet/dry cycles each year. When it rains, what appeared to be only a few clumps of short, dry grasses just a few days earlier suddenly teems with frogs, toads, clam shrimp, and aquatic plants.



Playa lakes in west Texas fill up after a heavy spring rain.

The playas provide a perfect home for migrating birds to feed, mate, and raise their young. Millions of shorebirds and waterfowl, including Canada geese, mallard ducks, and Sandhill cranes, depend on the playas for their breeding grounds. From the prairie potholes of North Dakota to the playa lakes of West Texas, the abundance and diversity of native bird species directly depends on these lakes.^{440,441}

Despite their small size, playa lakes and prairie potholes also play a critical role in supplying water to the Great Plains. The contribution of the playa lakes to this sensitively balanced ecosystem needs to be monitored and maintained in order to avoid unforeseen impacts on our natural resources. Before cultivation, water from these lakes was the primary source of the recharge to the High Plains aquifer.⁴⁴²



But many playas are disappearing and others are threatened by growing urban populations, extensive agriculture, and other filling and tilling practices.⁴⁴³ In recent years, agricultural demands have drawn down the playas to irrigate crops. Agricultural waste and fertilizer residues drain into playas, decreasing the quality of the water, or clogging them so the water cannot trickle down to refill the aquifer. Climate change is expected to add to these stresses, with increasing temperatures and changing rainfall patterns altering rates of evaporation, recharge, and runoff to the playa lake systems.⁴⁴⁴



L1 populations. Although farming and ranching remain
 L2 primary uses of the land – taking up much of the region’s
 L3 geographical area – growing cities provide housing and
 L4 jobs for more than two-thirds of the population. For
 L5 everyone on the Great Plains, though, a changing climate
 L6 and a limited water supply are likely to challenge their
 L7 ability to thrive, leading to conflicting interests in the
 L8 allocation of increasingly scarce water resources.^{313,433}

L9
 L10 **Native American communities**

L11 The Great Plains region is home to 65 Native American
 L12 tribes. Native populations on rural tribal lands have
 L13 limited capacities to respond to climate change.³¹³
 L14 Many reservations already face severe problems with
 L15 both water quantity and quality – problems likely to be
 L16 exacerbated by climate change and other human-induced
 L17 stresses.

L18
 L19 **Rural communities**

L20 As young adults migrate out of small, rural communities,
 L21 the towns are increasingly populated by a vulnerable
 L22 demographic of very old and very young, placing them
 L23 more at risk for health issues than urban communities.
 L24 Combined effects of changing demographics and climate
 L25 are likely to make it more difficult to supply adequate
 L26 and efficient public health services and educational
 L27 opportunities to rural areas. Climate-driven shifts in
 L28 optimal crop types and increased risk of drought, pests,
 L29 and extreme events will add more economic stress and
 L30 tension to traditional communities.^{430,433}

L31
 L32 **Urban populations**

L33 Although the Great Plains is not yet known for large
 L34 cities, many mid-sized towns throughout the region
 L35 are growing rapidly. One in four of the most rapidly
 L36

growing cities in the nation is located in the Great
 Plains⁴⁴⁶ (see *Society* sector). Most of these growing
 centers can be found in the southern parts of the
 region, where water resources are already seriously
 constrained. Urban populations, particularly the young,
 elderly, and economically disadvantaged, also might be
 disproportionately affected by heat.⁴⁴⁷

New opportunities

There is growing recognition that the enormous wind
 power potential of the Great Plains could provide new
 avenues for future employment and land use. Texas
 already produces the most wind power of any state.
 Wind energy production also is prominent in Oklahoma.
 North and South Dakota have rich wind potential.¹⁹¹

As climate change creates new environmental condi-
 tions, effective adaptation strategies become increasingly
 essential to ecological and socioeconomic survival. A
 great deal of the Great Plains’ adaptation potential might
 be realized through agriculture. For example, plant spe-
 cies that mature earlier and are more resistant to disease
 and pests are more likely to thrive under warmer condi-
 tions.

Other emerging adaptation strategies include dynamic
 cropping systems and increased crop diversity. In partic-
 ular, mixed cropping-livestock systems maximize avail-
 able resources while minimizing the need for external
 inputs such as irrigation that draws down precious water
 supplies.⁴³⁶ In many parts of the region, diverse cropping
 systems and improved water use efficiency will be key
 to sustaining crop and rangeland systems.⁴⁴⁸ Reduced
 water supplies might cause some farmers to alter the
 intensive cropping systems currently in use.^{193,219}

Adaptation: Agricultural Practices to Reduce Water Loss and Soil Erosion

Conservation of water is critical to efficient crop production in areas where water can be scarce. Following the Dust Bowl in the 1930s, Great Plains farmers implemented a number of improved farming practices to increase the effectiveness of rainfall capture and retention in the soil and protect the soil against water and wind erosion. Examples include rotating crops, retaining crop residues, increasing vegetative cover, and altering plowing techniques.



With observed and projected increases in summer temperatures and in the frequency and intensity of heavy downpours, it will become even more important to protect against increasing loss of water and soil. Across the upper Great Plains, where strong storms are projected to occur more frequently, producers are being encouraged to increase the amount of crop residue left on the soil or to plant cover crops in the fall to protect the soil in the spring before crops are planted.

Across the southern Great Plains, some farmers are returning to dryland farming rather than relying on irrigation for their crops. Preserving crop residue helps the soil absorb more moisture from rain and eases the burden on already-stressed groundwater. These efforts have been promoted by the U.S. Department of Agriculture through research and extension efforts such as Kansas State University’s Center for Sustainable Agriculture and Alternative Crops.



Southwest

L1
L2
L3
L4
L5
L6
L7
L8
L9
L10
L11
L12
L13
L14
L15
L16
L17
L18
L19
L20
L21
L22
L23
L24
L25
L26
L27
L28
L29
L30
L31
L32
L33
L34
L35
L36
L37
L38
L39
L40
L41
L42
L43
L44
L45
L46
L47
L48
L49
L50

R1
R2
R3
R4
R5
R6
R7
R8
R9
R10
R11
R12
R13
R14
R15
R16
R17
R18
R19
R20
R21
R22
R23
R24
R25
R26
R27
R28
R29
R30
R31
R32
R33
R34
R35
R36
R37
R38
R39
R40
R41
R42
R43
R44
R45
R46
R47
R48
R49
R50

The Southwest region stretches from the southern Rocky Mountains to the Pacific Coast. Elevations range from the lowest in the country to among the highest, with climates ranging from the driest to some of the wettest. Past climate records based on changes in Colorado River flows indicate that drought is a frequent feature of the Southwest, with some of the longest documented “megadroughts” on Earth. Since the 1940s, the region has experienced its most rapid population and urban growth. During this time, there were both unusually wet periods (including much of 1980s and 1990s) and dry periods (including much of 1950s and 1960s).⁴⁴⁹ The prospect of future droughts becoming more severe as a result of global warming is a significant concern, especially because the Southwest continues to lead the nation in population growth.

Human-induced climate change appears to be well underway in the Southwest. Recent warming is among the most rapid in the nation, significantly more than the global average in some areas. This is driving declines in spring snowpack and Colo-

rado River flow.^{34,160,161} Projections suggest continued strong warming, with much larger increases under higher emissions scenarios⁹¹ compared to lower scenarios. Projected summertime temperature increases are greater than the annual-average increases in some parts of the region, and are likely to be exacerbated locally by expanding urban heat island effects.⁴⁵⁰ Further water cycle changes are projected, which, combined with increasing temperatures, signal a serious water supply challenge in the decades and centuries ahead.^{34,159}

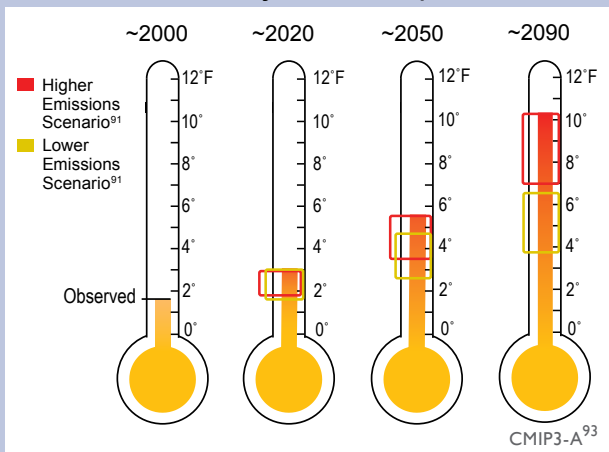
Water supplies will become increasingly scarce, calling for tradeoffs among competing uses, and potentially leading to conflict.

Water is, quite literally, the lifeblood of the Southwest. The largest use of water in the region is associated with agriculture, including some of the nation’s most important crop-producing areas in California. Water is also an important source of hydroelectric power, and water is required for the large population growth in the region, particularly that of major cities such as Phoenix and Las Vegas. Water also plays a critical role in supporting healthy ecosystems across the region, both on land and in rivers and lakes.

Water supplies in some areas of the Southwest are already becoming limited, and this trend toward scarcity is likely to be a harbinger of future water shortages.^{34,451} Groundwater pumping is lowering water tables, while rising temperatures reduce river flows in vital rivers including the Colorado.³⁴ Limitations imposed on water supply by projected temperature increases are likely to be made worse by substantial reductions in rain and snowfall in the spring months, when precipitation is most needed to fill reservoirs to meet summer demand.¹⁵¹

A warmer and drier future means extra care will be needed in planning the allocation of water for

Observed and Projected Temperature Rise



The average temperature in the Southwest has already increased roughly 1.5°F compared to a 1960-1979 baseline period. By the end of the century, average annual temperature is projected to rise approximately 4°F to 10°F above the historical baseline, averaged over the Southwest region. The brackets on the thermometers represent the likely range of model projections, though lower or higher outcomes are possible.

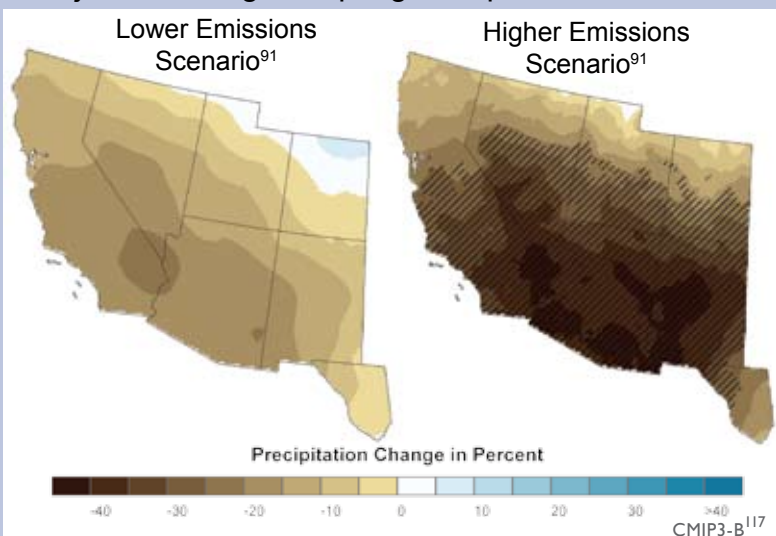
the coming decades. The Colorado Compact, negotiated in the 1920s, allocated the Colorado River’s water among the seven basin states. It was based, however, on unrealistic assumptions about how much water was available because the observations of runoff during the early 1900s turned out to be part of the greatest and

longest high-flow period of the last five centuries.⁴⁵² Today, even in normal decades, the Colorado River does not have enough water to meet the agreed-upon allocations. During droughts and under projected future conditions, the situation looks even bleaker.

Under exceptional circumstances, water designated for agriculture could provide a backup supply for urban water needs. Similarly, non-renewable groundwater could be tapped during especially dry periods. Both of these options, however, come at the cost of either current or future agricultural production.

Water is already a subject of contention in the Southwest, and climate change – coupled with rapid population growth – promises to increase the likelihood of water-related

Projected Change in Spring Precipitation, 2080-2099

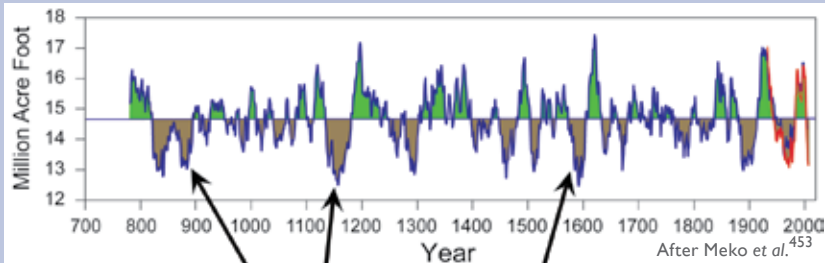


Percentage change in March-April-May precipitation for 2080-2099 compared to 1961-1979 for a lower emissions scenario⁹¹ (left) and a higher emissions scenario⁹¹ (right). Confidence in the projected changes is highest in the hatched areas.

Future of Drought in the Southwest

Droughts are a long-standing feature of the Southwest’s climate. The droughts of the last 110 years pale in comparison to some of the decades-long “megadroughts” that the region has experienced over the last 2000 years.⁴¹⁹ During the closing decades of the 1500s, for example, major droughts gripped parts of the Southwest.¹⁸⁹ These droughts sharply reduced the flow of the Colorado River^{452,453} and the all-important Sierra Nevada headwaters for California,⁴⁵⁴ and dried out the region as a whole. As of 2009, much of the Southwest remains in a drought that began around 1999. This event is the most severe western drought of the last 110 years, and is being exacerbated by record warming.⁴⁵⁵

Over this century, projections point to an increasing probability of drought for the region.^{90,115} Many aspects of these projections, including a northward shift in winter and spring storm tracks, are consistent with observed trends over recent decades.^{96,456,457} Thus, the most likely future for the Southwest is a substantially drier one (although there is presently no consensus on how the region’s summer monsoon [rainy season] might change in the future). Combined with the historical record of



Colorado River flow has been reconstructed back over 1200 years based primarily on tree-ring data. These data reveal that some droughts in the past have been more severe and longer lasting than any experienced in the last 100 years. The red line indicates actual measurements of river flow during the last 100 years. Models indicate that, in the future, droughts will continue to occur, but will become hotter, and thus more severe, over time.⁹⁰

severe droughts and the current uncertainty regarding the exact causes and drivers of these past events, the Southwest must be prepared for droughts that could potentially result from multiple causes. The combined effects of natural climate variability and human-induced climate change could turn out to be a devastating “one-two punch” for the region.

L1 conflict. Projected temperature increases, com- R1
 L2 bined with river-flow reductions, will increase the R2
 L3 risk of water conflicts between sectors, states, and R3
 L4 even nations. In recent years, negotiations regard- R4
 L5 ing existing water supplies have taken place among R5
 L6 the seven states sharing the Colorado River and R6
 L7 the two states (New Mexico and Texas) sharing the R7
 L8 Rio Grande. Mexico and the United States already R8
 L9 disagree on meeting their treaty allocations of Rio R9
 L10 Grande and Colorado River water. R10

L11
 L12 In addition, many water settlements between the R11
 L13 U.S. Government and Native American tribes have R12
 L14 yet to be fully worked out. The Southwest is home R13
 L15 to dozens of Native communities whose status as R14
 L16 sovereign nations means they hold rights to the R15
 L17 water for use on their land. However, the amount R16
 L18 of water actually available to each nation is deter- R17
 L19 mined through negotiations and litigation. Increas- R18
 L20 ing water demand in the Southwest is driving R19
 L21 current negotiations and litigation of tribal water R20
 L22 rights. While several nations have legally settled R21
 L23 their water rights, many other tribal negotiations R22
 L24 are either currently underway or pending. Compet- R23
 L25 ing demands from treaty rights, rapid development, R24
 L26 and changes in agriculture in the region, exacer- R25
 L27 bated by years of drought and climate change, have R26
 L28 the potential to spark significant conflict over an R27
 L29 already over-allocated and dwindling resource. R28

L30
 L31
 L32 **Increasing temperature, drought, wildfire, and invasive species will**
 L33 **accelerate transformation of the**
 L34 **landscape.**
 L35
 L36

L37 Climate change already appears to be influenc- R29
 L38 ing both natural and managed ecosystems of the R30
 L39 Southwest.^{455,458} Future landscape impacts are likely R31
 L40 to be substantial, threatening biodiversity, pro- R32
 L41 tected areas, and ranching and agricultural lands. R33
 L42 These changes are often driven by multiple factors, R34
 L43 including changes in temperature and drought pat- R35
 L44 terns, wildfire, invasive species, and pests. R36

L45
 L46 Conditions observed in recent years can serve as R37
 L47 indicators for future change. For example, tempera- R38
 L48 ture increases have made the current drought in R39
 L49 the region more severe than the natural droughts of R40
 L50 the last several centuries. As a result, about 4,600

square miles of piñon-juniper woodland in the Four R1
 Corners region of the Southwest have experienced R2
 substantial die-off of piñon pine trees.⁴⁵⁵ Record R3
 wildfires are also being driven by rising tempera- R4
 tures and related reductions in spring snowpack R5
 and soil moisture.⁴⁵⁸ R6

R7
 R8 How climate change will affect fire in the South- R8
 west varies according to location. In general, total R9
 area burned is projected to increase.⁴⁵⁹ How this R10
 plays out at individual locations, however, depends R11
 on regional changes in temperature and precipita- R12
 tion, as well as on whether fire in the area is cur- R13
 rently limited by fuel availability or by rainfall.⁴⁶⁰ R14
 For example, fires in wetter, forested areas are R15
 expected to increase in frequency, while areas R16
 where fire is limited by the availability of fine fuels R17
 experience decreases.⁴⁶⁰ Climate changes could R18
 also create subtle shifts in fire behavior, allowing R19
 more “runaway fires” – fires that are thought to R20
 have been brought under control, but then rekind- R21
 le.⁴⁶¹ The magnitude of fire damages, in terms of R22
 economic impacts as well as direct endangerment, R23
 also increases as urban development increasingly R24
 impinges on forested areas.^{460,462} R25

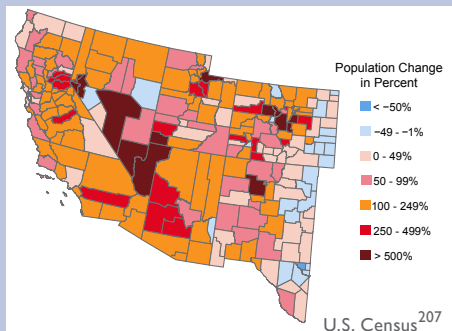
R26
 R27 Climate-fire dynamics will also be affected by R27
 changes in the distribution of ecosystems across the R28
 Southwest. Increasing temperatures and shifting R29
 precipitation patterns will drive declines in high- R30
 elevation ecosystems such as alpine forests and R31
 tundra.^{459,463} Under higher emissions scenarios,⁹¹ R32
 high-elevation forests in California, for example, R33
 are projected to decline by 60 to 90 percent be- R34
 fore the end of the century.^{284,459} At the same time, R35
 grasslands are projected to expand, another factor R36
 likely to increase fire risk. R37

R38
 R39 As temperatures rise, some iconic landscapes of R39
 the Southwest will be greatly altered as species R40
 shift their ranges northward and upward to cooler R41
 climates, and fires attack unaccustomed ecosys- R42
 tems which lack natural defenses. The Sonoran R43
 Desert, for example, famous for the saguaro cactus, R44
 would look very different if more woody species R45
 spread northward from Mexico into areas currently R46
 dominated by succulents (such as cacti) or native R47
 grasses.⁴⁶⁴ The desert is already being invaded R48
 by red brome and buffle grasses that do well in R49
 high temperatures and are native to Africa and the R50

R1
 R2
 R3
 R4
 R5
 R6
 R7
 R8
 R9
 R10
 R11
 R12
 R13
 R14
 R15
 R16
 R17
 R18
 R19
 R20
 R21
 R22
 R23
 R24
 R25
 R26
 R27
 R28
 R29
 R30
 R31
 R32
 R33
 R34
 R35
 R36
 R37
 R38
 R39
 R40
 R41
 R42
 R43
 R44
 R45
 R46
 R47
 R48
 R49
 R50



Change in Population from 1970 to 2008



The map above of percentage changes in county population between 1970 and 2008 shows that the Southwest has experienced very rapid growth in recent decades (indicated in orange, red, and maroon).

L1
L2
L3
L4
L5
L6
L7
L8
L9
L10
L11
L12
L13
L14
L15
L16
L17
L18
L19
L20
L21
L22
L23
L24
L25
L26
L27
L28
L29
L30
L31
L32
L33
L34
L35
L36
L37
L38
L39
L40
L41
L42
L43
L44
L45
L46
L47
L48
L49
L50

Mediterranean. Not only do these noxious weeds out-compete some native species in the Sonoran Desert, they also fuel hot, cactus-killing fires. With these invasive plant species and climate change, the Saguaro and Joshua Tree national parks could end up with

far fewer of their namesake plants.⁴⁶⁵ In California, two-thirds of the more than 5,500 native plant species are projected to experience range reductions up to 80 percent before the end of this century under projected warming.⁴⁶⁶ In their search for optimal conditions, some species will move uphill, others northward, breaking up present-day ecosystems; those species moving southward to higher elevations might cut off future migration options as temperatures continue to increase.

The potential for successful plant and animal adaptation to coming change is further hampered by existing regional threats such as human-caused

fragmentation of the landscape, invasive species, river-flow reductions, and pollution. Given the mountainous nature of the Southwest, and the associated impediments to species shifting their ranges, climate change likely places other species at risk. Some areas have already been identified as possible refuges, where species at risk could continue to live if these areas were preserved for this purpose.⁴⁶⁶ Other rapidly changing landscapes will require major adjustments, not only from plant and animal species, but also the region’s ranchers, foresters, and other inhabitants.

Increased frequency and altered timing of flooding will increase risks to people, ecosystems, and infrastructure.

Paradoxically, a warmer atmosphere and an intensified water cycle are likely to mean not only a greater likelihood of drought for the Southwest, but also an increased risk of flooding. Winter precipitation in Arizona, for example, is already becoming more variable, with a trend toward both more frequent extremely dry and extremely wet winters.⁴⁷² Some water systems rely on smaller reservoirs being filled up each year. More frequent dry winters suggest an increased risk of these systems running short of water. However, a greater potential for flooding also means reservoirs cannot

R1
R2
R3
R4
R5
R6
R7
R8
R9
R10
R11
R12
R13
R14
R15
R16
R17
R18
R19
R20
R21
R22
R23
R24
R25
R26
R27
R28
R29
R30
R31
R32
R33
R34
R35
R36
R37
R38
R39
R40
R41
R42
R43
R44
R45
R46
R47
R48
R49
R50

A Biodiversity Hotspot

The Southwest is home to two of the world’s 34 designated “biodiversity hotspots.” These at-risk regions have two special qualities: they hold unusually large numbers of plant and animal species that are endemic (found nowhere else), and they have already lost over 70 percent of their native vegetation.^{467,468} About half the world’s species of plants and land animals occur only in these 34 locations, though they cover just 2.3 percent of the Earth’s land surface.

One of these biodiversity hotspots is the Madrean Pine-Oak Woodlands. Once covering 178 square miles, only isolated patches remain in the United States, mainly on mountaintops in southern Arizona, New Mexico, and West Texas. The greatest diversity of pine species in the world grows in this area: 44 of the 110 varieties,⁴⁶⁹ as well as more than 150 species of oak.⁴⁷⁰ Some 5,300 to 6,700 flowering plant species inhabit the ecosystem, and over 500 bird species, 23 of which are endemic. More hummingbirds are found here than anywhere else in the United States. There are 384 species of reptiles, 37 of which are endemic, and 328 species of mammals, six of which are endemic. There are 84 fish species, 18 of which are endemic. Some 200 species of butterfly thrive here, of which 45 are endemic, including the Monarch that migrates 2,500 miles north to Canada each year.⁴⁷¹ Ecotourism has become the economic driver in many parts of this region, but logging, land clearing for agriculture, urban development, and now climate change threaten the region’s viability.

L1 be filled to capacity as safely in years where that
 L2 is possible. Flooding also causes reservoirs to fill
 L3 with sediment at a faster rate, thus reducing their
 L4 water-storage capacities.
 L5

L6 On a global scale, precipitation patterns are already
 L7 observed to be shifting, with more rain falling in
 L8 heavy downpours that can lead to flooding.^{90,473}
 L9 Rapid landscape transformation due to vegetation
 L10 die-off and wildfire as well as loss of wetlands
 L11 along rivers is also likely to reduce flood-buffering
 L12 capacity. Moreover, increased flood risk in the
 L13 Southwest is likely to result from a combination of
 L14 decreased snow cover on the lower slopes of high
 L15 mountains, and an increased fraction of winter pre-
 L16 cipitation falling as rain and therefore running off
 L17 more rapidly.¹⁵⁴ The increase in rain on snow events
 L18 will also result in rapid runoff and flooding.⁴⁷⁴
 L19

L20 The most obvious impact of more frequent flooding
 L21 is a greater risk to human beings and their infra-
 L22 structure. This applies to locations along major riv-
 L23 ers, but also to much broader and highly vulnerable
 L24 areas such as the Sacramento–San Joaquin River
 L25 Delta system. Stretching from the San Francisco
 L26 Bay nearly to the state capital of Sacramento, the
 L27 Sacramento–San Joaquin River Delta and Suisun
 L28 Marsh makes up the largest estuary on the West
 L29 Coast of North America. With its rich soils and
 L30 rapid subsidence rates – in some locations as high
 L31 as 2 or more feet per decade – the entire Delta
 L32 region is now below mean water level, protected by
 L33 more than a thousand miles of levees and dams.⁴⁷⁵
 L34 Projected changes in the timing and amount of river
 L35 flow, particularly in winter and spring, is estimated
 L36 to more than double the risk of Delta flooding
 L37 events by mid-century, and result in an eight-fold
 L38 increase before the end of the century.⁴⁷⁶ Taking
 L39 into account the additional risk of a major seismic
 L40 event and increases in sea level due to climate
 L41 change over this century, the California Bay–Delta
 L42 Authority has concluded that the Delta and Suisun
 L43 Marsh are not sustainable under current practices;
 L44 efforts are underway to identify and implement ad-
 L45 aptation strategies aimed at reducing these risks.⁴⁷⁶
 L46
 L47
 L48
 L49
 L50

Unique tourism and recreation opportunities are likely to suffer.

Tourism and recreation are important aspects of the region’s economy. Increasing temperatures will affect important winter activities such as downhill and cross-country skiing, snowshoeing, and snowmobiling that require snow on the ground. Projections indicate later snow and less snow coverage in ski resort areas, particularly those at lower elevations and in the southern part of the region.²⁸⁴ Decreases from 40 to almost 90 percent are likely in end-of-season snowpack under a higher emissions scenario⁹¹ in counties with major ski resorts from New Mexico to California.⁴⁷⁷ In addition to shorter seasons, earlier wet snow avalanches – more than six weeks earlier by the end of this century under a higher emissions scenario⁹¹ – could force ski areas to shut down affected runs before the season would otherwise end.⁴⁷⁸ Resorts require a certain number of days just to break even; cutting the season short by even a few weeks, particularly if those occur during the lucrative holiday season, could easily render a resort unprofitable.

Even in non-winter months, ecosystem degradation will affect the quality of the experience for hikers, bikers, birders, and others who enjoy the Southwest’s natural beauty. Water sports that depend on the flows of rivers and sufficient water in lakes and reservoirs are already being affected, and much larger changes are expected.

Cities and agriculture face increasing risks from a changing climate.

Resource use in the Southwest is involved in a constant three-way tug-of-war among preserving natural ecosystems, supplying the needs of rapidly expanding urban areas, and protecting the lucrative agricultural sector, which particularly in California, is largely based on highly temperature- and water-sensitive specialty crops. Urban areas are also sensitive to temperature-related impacts on air quality, electricity demand, and the health of their inhabitants.

R1
 R2
 R3
 R4
 R5
 R6
 R7
 R8
 R9
 R10
 R11
 R12
 R13
 R14
 R15
 R16
 R17
 R18
 R19
 R20
 R21
 R22
 R23
 R24
 R25
 R26
 R27
 R28
 R29
 R30
 R31
 R32
 R33
 R34
 R35
 R36
 R37
 R38
 R39
 R40
 R41
 R42
 R43
 R44
 R45
 R46
 R47
 R48
 R49
 R50



L1 The magnitude of projected temperature increases
 L2 for the Southwest, particularly when combined
 L3 with urban heat island effects for major cities such
 L4 as Phoenix, Albuquerque, Las Vegas, and many
 L5 California cities, represent significant stresses
 L6 to health, electricity, and water supply in a re-
 L7 gion that already experiences very high summer
 L8 temperatures.^{284,325,450}

L10 If present-day levels of ozone-producing emis-
 L11 sions are maintained, rising temperatures also
 L12 imply declining air quality in urban areas such as
 L13 those in California which already experience some
 L14 of the worst air quality in the nation (see *Society*
 L15 *sector*).⁴⁷⁹ Continued rapid population growth is
 L16 expected to exacerbate these concerns.

L18 With more intense, longer-lasting heat wave events
 L19 projected to occur over the coming century, de-
 L20 mands for air conditioning are expected to deplete
 L21 electricity supplies, increasing risks of brownouts
 L22 and blackouts.³²⁵ Electricity supplies will also be
 L23 affected by changes in the timing of river flows and
 L24 where hydroelectric systems have limited storage
 L25 capacity and reservoirs.^{480,481}

L27 Much of the region's agriculture will experi-
 L28 ence detrimental impacts in a warmer future,
 L29 particularly specialty crops in California such as
 L30 apricots, almonds, artichokes, figs, kiwis, olives,

and walnuts.^{482,483} These and other specialty crops R1
 require a minimum number of hours at a chill- R2
 ing temperature threshold in the winter to become R3
 dormant and set fruit for the following year.⁴⁸² R4
 Accumulated winter chilling hours have already R5
 decreased across central California and its coastal R6
 valleys. This trend is projected to continue to the R7
 point where chilling thresholds for many key crops R8
 would no longer be met. A steady reduction in win- R9
 ter chilling could have serious economic impacts on R10
 fruit and nut production in the region. California's R11
 losses due to future climate change are estimated R12
 between zero and 40 percent for wine and table R13
 grapes, almonds, oranges, walnuts, and avocados, R14
 varying significantly by location.⁴⁸³ R15
 R16

Adaptation strategies for agriculture in Califor- R17
 nia include more efficient irrigation and shifts R18
 in cropping patterns, which have the potential to R19
 help compensate for climate-driven increases in R20
 water demand for agriculture due to rising tem- R21
 peratures.⁴⁸⁴ The ability to use groundwater and/or R22
 water designated for agriculture as backup sup- R23
 plies for urban uses in times of severe drought is R24
 expected to become more important in the future as R25
 climate change dries out the Southwest; however, R26
 these supplies are at risk of being depleted as urban R27
 populations swell. R28
 R29
 R30

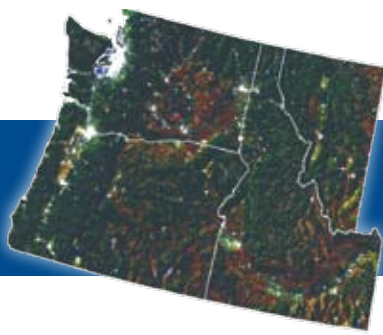
Adaptation: Strategies for Fire

Living with present-day levels of fire risk, along with projected increases in risk, involves actions by residents along the urban-forest interface as well as fire and land management officials. Some basic strategies for reducing damage to structures due to fires are being encouraged by groups like National Firewise Communities, an interagency program that encourages wildfire preparedness measures such as creating defensible space around residential structures by thinning trees and brush, choosing fire-resistant plants, selecting ignition-resistant building materials and design features, positioning structures away from slopes, and working with firefighters to develop emergency plans.

Additional strategies for responding to the increased risk of fire as climate continues to change could include adding fire-fighting resources⁴⁶¹ and improving evacuation procedures and communications infrastructure. Also important would be regularly updated insights into what the latest climate science implies for changes in types, locations, timing, and potential severity of fire risks over seasons to decades and beyond; implications for related political, legal, economic, and social institutions; and improving predictions for regeneration of burnt-over areas and the implications for subsequent fire risks. Reconsideration of policies that encourage growth of residential developments in or near forests is another potential avenue for adaptive strategies.⁴⁶²



Northwest



The Northwest's rapidly growing population, as well as its forests, mountains, rivers, and coastlines, are already experiencing human-induced climate change and its impacts.³⁴ Regionally-averaged temperature rose about 1.5°F over the past century⁴⁸⁵ (with some areas experiencing increases up to 4°F) and is projected to increase another 3 to 10°F during this century.⁴⁸⁶ Higher emissions scenarios[†] would result in warming in the upper end of the projected range. Increases in winter precipitation and decreases in summer precipitation are projected by many climate models,⁴⁸⁷ though these projections are less certain than those for temperature. Impacts related to changes in snowpack, streamflows, sea level, forests, and other important aspects of life in the Northwest are already underway, with more severe impacts expected over coming decades in response to continued and more rapid warming.

Declining springtime snowpack leads to reduced summer streamflows, straining water supplies.

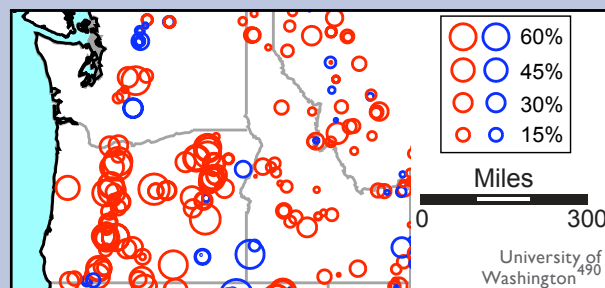
The Northwest is highly dependent on temperature-sensitive springtime snowpack to meet growing, and often competing, water demands such as municipal and industrial uses, agricultural irrigation, hydropower production, navigation, recreation, and in-stream flows that protect aquatic ecosystems including threatened and endangered species. Higher cool season (October through March) temperatures cause more precipitation to fall as rain rather than snow and contribute to earlier snowmelt. April 1 snowpack, a key indicator of natural water storage available for the warm season, has already declined substantially throughout the region. The average decline in the Cascade Mountains, for example, was about 25 percent over the past 40 to 70 years, with most of this due to the 2.5°F increase in cool season temperatures over that period.^{108,488} Further declines in Northwest snowpack are projected to result from additional

warming over this century, varying with latitude, elevation, and proximity to the coast. April 1 snowpack is projected to decline as much as 40 percent in the Cascades by the 2040s.⁴⁸⁹ Throughout the region, earlier snowmelt will cause a reduction in the amount of water available during the warm season.⁶⁸

In areas where it snows, a warmer climate means major changes in the timing of runoff: streamflow increases in winter and early spring, and then decreases in late spring, summer, and fall. This shift in streamflow timing has already been observed over the past 50 years,²⁵³ with the peak of spring runoff shifting from a few days earlier in some places to as much as 25 to 30 days earlier in others.¹⁵⁷

This trend is projected to continue, with runoff shifting 20 to 40 days earlier within this century.¹⁵⁷ Reductions in summer water availability will vary with the temperatures experienced in different parts of the region. In relatively warm areas on the western slopes of the Cascade Mountains, for example, reductions in warm season (April through September) runoff of 30 percent or more are projected by mid-century, whereas colder areas in the Rocky Mountains are expected to see reductions on the order of 10 percent. Areas dominated by rain rather than snow are not expected to see major shifts in the timing of runoff.⁴⁹²

Trends in April 1 Snow Water Equivalent 1950 to 2002



April 1 snowpack (a key indicator of natural water storage available for the warm season) has declined throughout the Northwest. In the Cascade Mountains, April 1 snowpack declined by an average of 25 percent, with some areas experiencing up to 60 percent declines. On the map, decreasing trends are in red and increasing trends are in blue.⁴⁹¹

L1 Extreme high and low streamflows also are ex- R1
 L2 pected to change with warming. Increasing winter R2
 L3 rainfall (as opposed to snowfall) is expected to lead R3
 L4 to more winter flooding in relatively warm water- R4
 L5 sheds on the west side of the Cascades. The already R5
 L6 low flows of late summer are projected to decrease R6
 L7 further due to both earlier snowmelt and increased R7
 L8 evaporation and water loss from vegetation. Pro- R8
 L9 jected decreases in summer precipitation would R9
 L10 exacerbate these effects. Some sensitive watersheds R10
 L11 are projected to experience both increased flood R11
 L12 risk in winter and increased drought risk in sum- R12
 L13 mer due to warming. R13

L14
 L15 The region’s water supply infrastructure was built R15
 L16 based on the assumption that most of the water R16
 L17 needed for summer uses would be stored naturally R17
 L18 in snowpack. For example, the storage capacity in R18
 L19 Columbia Basin reservoirs is only 30 percent of the R19
 L20 annual runoff, and many small urban water sup- R20
 L21 ply systems on the west side of the Cascades store R21
 L22 less than 10 percent of their annual flow.⁴⁹³ Besides R22
 L23 providing water supply and managing flows for R23
 L24 hydropower, the region’s reservoirs are operated for R24
 L25 flood-protection purposes and, as such, might have R25
 L26 to release (rather than store) large amounts of run- R26
 L27 off during the winter and early spring to maintain R27
 L28 enough space for flood protection. Earlier flows R28
 L29 would thus place more of the year’s runoff into the R29
 L30 category of hazard rather than resource. An ad- R30
 L31 vance in the timing of snowmelt runoff would also R31

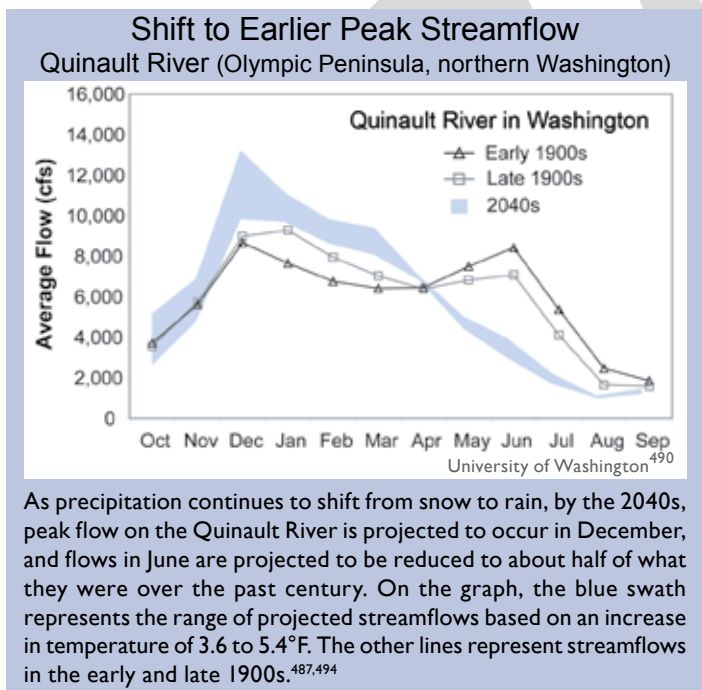
R1 increase the length of the summer dry period, with R1
 R2 important consequences for water supply, ecosys- R2
 R3 tems, and wildfire management.¹⁵⁷ R3
 R4

R5 One of the largest demands on water resources in R5
 R6 the region is hydroelectric power production. About R6
 R7 70 percent of the Northwest’s electricity is provided R7
 R8 by hydropower, a far greater percentage than in R8
 R9 any other region. Warmer summers will increase R9
 R10 electricity demands for air conditioning and refrig- R10
 R11 eration at the same time of year that lower stream- R11
 R12 flows will lead to reduced hydropower generation. R12
 R13 At the same time, water is needed for irrigated R13
 R14 agriculture, protecting fish species, reservoir and R14
 R15 river recreation, and urban uses. Conflicts between R15
 R16 all of these water uses are expected to increase, R16
 R17 forcing complex trade-offs between competing R17
 R18 objectives.^{487,494} R18
 R19
 R20

R21 **Increased insect outbreaks, wildfires,**
 R22 **and changing species composition**
 R23 **in forests will pose challenges for**
 R24 **ecosystems, and the forest products**
 R25 **industry.** R25

R26
 R27 Higher summer temperatures and earlier spring R27
 R28 snowmelt are expected to increase the risk of forest R28
 R29 fires in the Northwest by increasing summer mois- R29
 R30 ture deficits; this pattern has already been observed R30
 R31 in recent decades. Drought stress and higher tem- R31
 R32 peratures will decrease tree growth in most low- R32
 R33 and mid-elevation forests and also will increase the R33
 R34 frequency and intensity of mountain pine beetle R34
 R35 and other insect attacks,²⁴⁴ further increasing fire R35
 R36 risk and reducing timber production, an important R36
 R37 part of the regional economy. The mountain pine R37
 R38 beetle outbreak in British Columbia has destroyed R38
 R39 33 million acres of trees so far, about 40 percent of R39
 R40 the marketable pine trees in the province. By 2018, R40
 R41 it is projected that the infestation will have run R41
 R42 its course and over 78 percent of the mature pines R42
 R43 will have been killed; this will affect more than R43
 R44 one-third of the total area of British Columbia’s R44
 R45 forests⁴⁹⁵ (see *Ecosystems* sector). Idaho’s Sawtooth R45
 R46 Mountains are also now threatened by pine beetle R46
 R47 infestation. R47
 R48

R49 In the short term, high elevation forests on the west R49
 R50 side of the Cascade Mountains are expected to R50

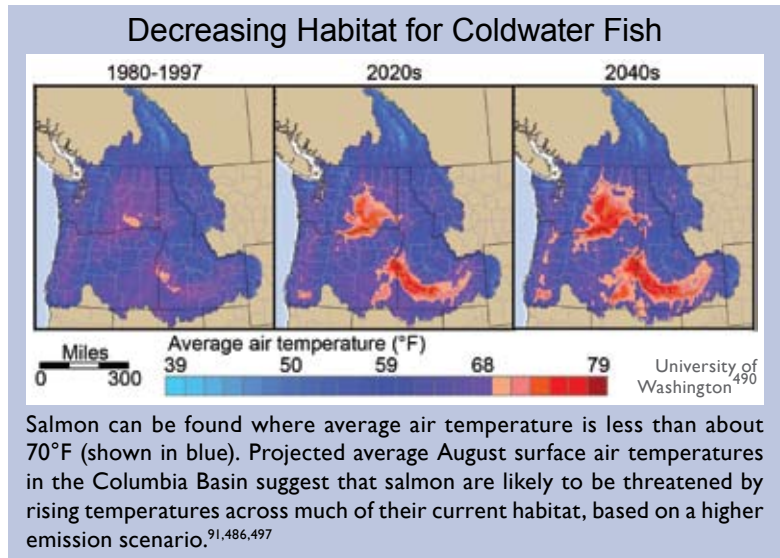


L1 see increased growth. In the longer term, forest
 L2 growth is expected to decrease as summertime
 L3 soil moisture deficits limit forest productivity,
 L4 with low-elevation forests experiencing these
 L5 changes first. The extent and species composi-
 L6 tion of forests also are expected to change as tree
 L7 species respond to climatic changes. There is also
 L8 the potential for extinction of local populations
 L9 and loss of biological diversity if environmental
 L10 changes outpace species' ability to shift their
 L11 ranges and form successful new ecosystems.

L13 Agriculture, especially production of tree fruit
 L14 such as apples, is also an important part of the
 L15 regional economy. Decreasing irrigation sup-
 L16 plies, increasing pests and disease, and increased
 L17 competition from weeds are likely to have negative
 L18 effects on agricultural production.

L21 **Salmon and other coldwater species**
 L22 **will experience additional stresses as a**
 L23 **result of rising water temperatures and**
 L24 **declining summer streamflows.**

L26 Northwest salmon populations are at historically
 L27 low levels due to stresses imposed by a variety of
 L28 human activities including dam building, logging,
 L29 pollution, and over-fishing. Climate change affects
 L30 salmon throughout their life stages and poses an
 L31 additional stress. As more winter precipitation falls
 L32 as rain rather than snow, higher winter stream-
 L33 flows scour streambeds, damaging spawning nests
 L34 and washing away incubating eggs. Earlier peak
 L35 streamflows flush young salmon from rivers to
 L36 estuaries before they are physically mature enough
 L37 for the transition, increasing a variety of stresses
 L38 including the risk of being eaten by predators.
 L39 Lower summer streamflows and warmer water
 L40 temperatures create less favorable summer stream
 L41 conditions for salmon and other coldwater fish
 L42 species in many parts of the Northwest. In addition,
 L43 diseases and parasites that infect salmon tend to
 L44 flourish in warmer water. Climate change also im-
 L45 pacts the ocean environment, where salmon spend
 L46 several years of their lives. Historically, warm
 L47 periods in the coastal ocean have coincided with
 L48 relatively low abundances of salmon, while cooler
 L49 ocean periods have coincided with relatively high
 L50 salmon numbers.



R17 Most wild Pacific salmon populations are extinct
 R18 or imperiled in 56 percent of their historical range
 R19 in the Northwest and California,⁴⁹⁶ and populations
 R20 are down more than 90 percent in the Columbia
 R21 River system. Many species are listed as either
 R22 threatened or endangered under the Federal En-
 R23 dangered Species Act. Studies suggest that about
 R24 one-third of the current habitat for the Northwest's
 R25 salmon and other coldwater fish will no longer be
 R26 suitable for them by the end of this century as key
 R27 temperature thresholds are exceeded. Because cli-
 R28 mate change impacts on their habitat are projected
 R29 to be negative, climate change is expected to ham-
 R30 per efforts to restore depleted salmon populations.

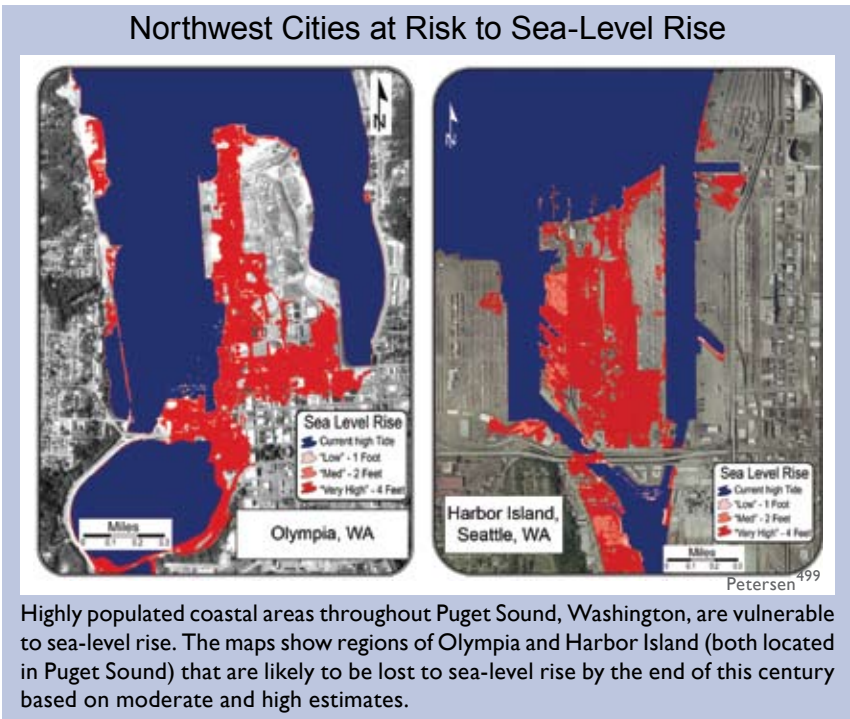
R33 **Sea-level rise along vulnerable coastlines**
 R34 **will result in increased erosion and the**
 R35 **loss of land.**

R37 Climate change is projected to exacerbate many
 R38 of the stresses and hazards currently facing the
 R39 coastal zone. Sea-level rise will increase erosion of
 R40 the Northwest coast and cause the loss of beaches
 R41 and significant coastal land areas. Among the most
 R42 vulnerable parts of the coast are the heavily popu-
 R43 lated south Puget Sound region, which includes
 R44 the cities of Olympia, Tacoma, and Seattle, Wash-
 R45 ington. Some climate models project changes in
 R46 atmospheric pressure patterns that suggest a more
 R47 southwesterly direction of future winter winds.
 R48 Combined with higher sea levels, this would accel-
 R49 erate coastal erosion all along the Pacific Coast.
 R50 Sea-level rise in the Northwest (as elsewhere) is

L1
L2
L3
L4
L5
L6
L7
L8
L9
L10
L11
L12
L13
L14
L15
L16
L17
L18
L19
L20
L21
L22
L23
L24
L25
L26
L27
L28
L29
L30
L31
L32
L33
L34
L35
L36
L37
L38
L39
L40
L41
L42
L43
L44
L45
L46
L47
L48
L49
L50

determined by global rates of sea-level rise, changes in coastal elevation associated with local vertical movement of the land, and atmospheric circulation patterns that influence wind-driven “pile-up” of water along the coast. A mid-range estimate of relative sea-level rise for the Puget Sound basin is about 13 inches by 2100. However, higher levels of up to 50 inches by 2100 in more rapidly subsiding portions of the basin are also possible given the large uncertainties about accelerating rates of ice melt from Greenland and Antarctica in recent years.⁴⁹⁸

An additional concern is landslides on coastal bluffs. The projected heavier winter rainfall suggests an increase in saturated soils and, therefore, an increased number of landslides. Increased frequency and/or severity of landslides is expected to be especially problematic in areas where there has been intensive development on unstable slopes. Within Puget Sound, the cycle of beach erosion and bluff landslides will be exacerbated by sea-level rise, increasing beach erosion, and decreasing slope stability.



R1
R2
R3
R4
R5
R6
R7
R8
R9
R10
R11
R12
R13
R14
R15
R16
R17
R18
R19
R20
R21
R22
R23
R24
R25
R26
R27
R28
R29
R30
R31
R32
R33
R34
R35
R36
R37
R38
R39
R40
R41
R42
R43
R44
R45
R46
R47
R48
R49
R50

Adaptation: Improved Planning to Cope with Future Changes

States, counties, and cities in the Northwest are beginning to develop strategies to adapt to climate change. In 2007, Washington state convened stakeholders to develop adaptation strategies for water, agriculture, forests, coasts, infrastructure, and human health. Recommendations included improved drought planning, improved monitoring of diseases and pests, incorporating sea-level rise in coastal planning, and public education. An implementation strategy is under development.

In response to concerns about increasing flood risk, King County, Washington, approved plans in 2007 to fund repairs to the county’s aging levee system. The county also will replace more than 57 “short-span” bridges with wider span structures that allow more debris and floodwater to pass underneath rather than backing up and causing the river to flood. The county has begun incorporating porous concrete and rain gardens into road projects to manage the effects of stormwater runoff during heavy rains, which are increasing as climate changes. King County also has published an adaptation guidebook that is becoming a model that other local governments can refer to in order to organize adaptation actions within their municipal planning processes.⁵⁰⁰

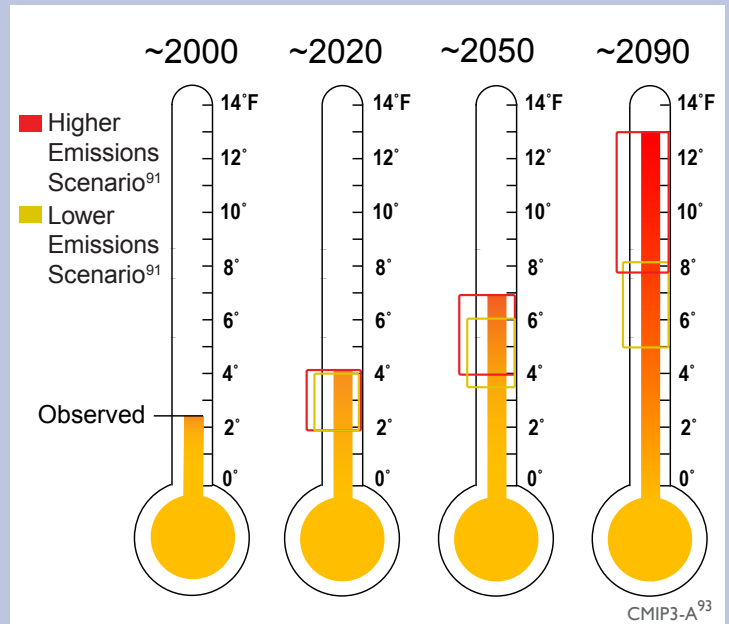
Concern about sea-level rise in Olympia, Washington, contributed to the city’s decision to relocate its primary drinking water source from a low-lying surface water source to wells on higher ground. The city adjusted its plans for construction of a new City Hall to locate the building in an area less vulnerable to sea-level rise than the original proposed location. The building’s foundation also was raised by 1 foot.

Alaska

Over the past 50 years, Alaska has warmed at more than twice the rate of the rest of the United States. Its annual average temperature has increased 3.4°F, while winters have warmed even more, by 6.3°F.⁵⁰¹ As a result, climate change impacts are much more pronounced than in other regions of the United States. The higher temperatures are already contributing to earlier spring snowmelt, reduced sea ice, widespread glacier retreat, and permafrost warming.^{220,501} These observed changes are consistent with climate model projections of greater warming over Alaska, especially in winter, as compared to the rest of the country.

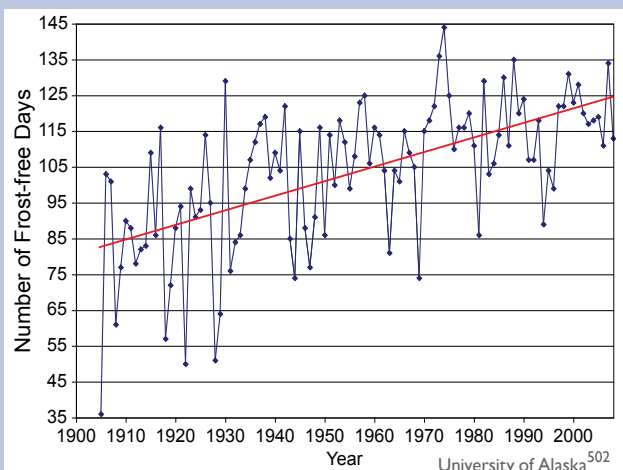
Climate models also project increases in precipitation over Alaska. Simultaneous increases in evaporation due to higher air temperatures, however, are expected to lead to drier conditions overall, with reduced soil moisture.⁹⁰ In the future, therefore, model projections suggest a longer summer growing season combined with an increased likelihood of summer drought and wildfires.

Observed and Projected Temperature Rise



Alaska's annual average temperature has increased 3.4°F over the past 50 years. The observed increase shown above compares the average temperature of 1993-2007 with a 1960s-1970s baseline, an increase of over 2°F. The brackets on the thermometers represent the likely range of model projections, though lower or higher outcomes are possible. By the end of this century, the average temperature is projected to rise by 5 to 13°F above the 1960s-1970s baseline.

Fairbanks Frost-Free Season, 1904 to 2008



Over the past 100 years, the length of the frost-free season in Fairbanks, Alaska, has increased by 50 percent. The trend toward a longer frost-free season is projected to produce benefits in some sectors and detriments in others.

Average annual temperatures in Alaska are projected to rise about 3.5 to 7°F by the middle of this century. How much temperatures rise later in the century depends strongly on global emissions choices, with increases of 5 to 8°F projected with lower emissions,⁹¹ and increases of 8 to 13°F with higher emissions.⁹¹ Higher temperatures are expected to continue to reduce Arctic sea ice coverage. Reduced sea ice provides opportunities for increased shipping and resource extraction. At the same time, however, it increases coastal erosion, raises the risk of accidents as offshore commercial activity increases, and is expected to drive major shifts of marine species such as pollock and other commercial fish stocks.

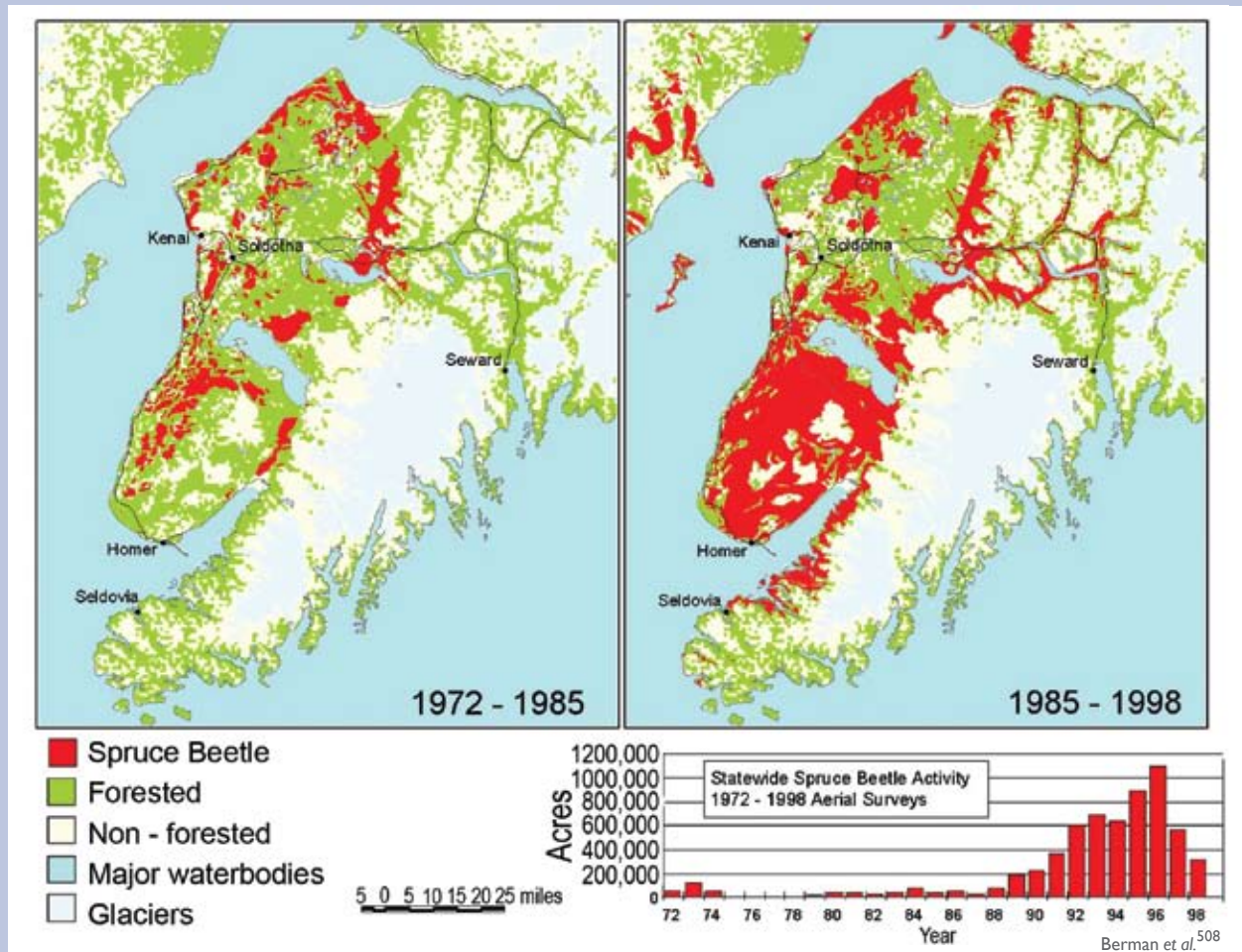
Longer summers and higher temperatures are causing drier conditions, even in the absence of strong trends in precipitation.

Between 1970 and 2000, the snow-free season increased by approximately 10 days across Alaska, primarily due to earlier snowmelt in the spring.^{503,504} A longer growing season has potential economic benefits, providing a longer period of outdoor and commercial activity such as tourism. However, there are also downsides. For example, white spruce forests in Alaska’s interior are experiencing declining growth due to drought stress⁵⁰⁵ and continued warming could lead to widespread death of trees.⁵⁰⁶ The decreased soil moisture in Alaska also suggests that agriculture in Alaska might not benefit from the longer growing season.

Insect outbreaks and wildfires are increasing with warming.

Climate plays a key role in determining the extent and severity of insect outbreaks and wildfires.^{506,507} During the 1990s, for example, south-central Alaska experienced the largest outbreak of spruce beetles in the world.^{244,506} This outbreak occurred because rising temperatures allowed the spruce beetle to survive over the winter and to complete its life cycle in just one year instead of the normal two years. Healthy trees ordinarily defend themselves by pushing back against burrowing beetles with their pitch. From 1989 to 1997, however, the region experienced an extended drought, leaving the trees too stressed to fight off the infestation.

Alaska Spruce Beetle Infestation
Kenai Peninsula, 1972 to 1998



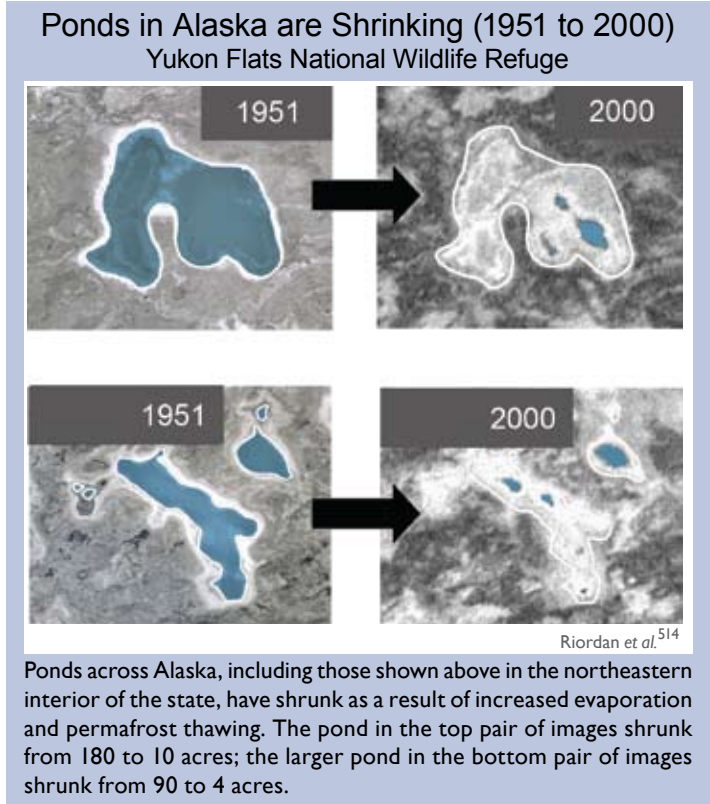
Warming in Alaska has caused insect outbreaks to increase. Red areas indicate spruce beetle infestations on the Kenai Peninsula. Over 5 million acres of Alaska spruce forests were destroyed.

L1 Prior to 1990, the spruce budworm was not able to
 L2 reproduce in interior Alaska.⁵⁰⁶ Hotter, drier sum-
 L3 mers, however, now mean that the forests there are
 L4 threatened by an outbreak of spruce budworms.⁵⁰⁹
 L5 This trend is expected to increase in the future
 L6 if summers in Alaska become hotter and drier.⁵⁰⁶
 L7 Large areas of dead trees, such as those left behind
 L8 by pest infestations, are highly flammable and thus
 L9 much more vulnerable to wildfire than living trees.

L10
 L11 The area burned in North America’s northern forest
 L12 that spans Alaska and Canada tripled from the
 L13 1960s to the 1990s. Two of the three most exten-
 L14 sive wildfire seasons in Alaska’s 56-year record
 L15 occurred in 2004 and 2005, and half of the most
 L16 severe fire years on record have occurred since
 L17 1990.⁵¹⁰ Under changing climate conditions, the av-
 L18 erage area burned per year in Alaska is projected to
 L19 double by the middle of this century.⁵⁰⁷ By the end
 L20 of this century, area burned by fire is projected to
 L21 triple under a moderate greenhouse gas emissions
 L22 scenario and to quadruple under a higher emissions
 L23 scenario.⁹¹ Such increases in area burned would
 L24 result in numerous impacts, including hazardous
 L25 air quality conditions such as those suffered by
 L26 residents of Fairbanks during the summers of 2004
 L27 and 2005, as well as increased risks to rural Native
 L28 Alaskan communities because of reduced avail-
 L29 ability of the fish and game that make up their diet.
 L30 This would cause them to adopt a more “Western”
 L31 diet,⁵¹¹ known to be associated with increased risk
 L32 of cancers, diabetes, and cardiovascular disease.⁵¹²

L33
 L34
 L35 **Lakes are declining in area.**

L36
 L37 Across the southern two-thirds of Alaska, the
 L38 area of closed-basin lakes (lakes without stream
 L39 inputs and outputs) has decreased over the past 50
 L40 years. This is likely due to the greater evapora-
 L41 tion and thawing of permafrost that result from
 L42 warming.^{513,514} A continued decline in the area of
 L43 surface water would present challenges for the
 L44 management of natural resources and ecosystems
 L45 on National Wildlife Refuges in Alaska. These
 L46 refuges, which cover over 77 million acres (21 per-
 L47 cent of Alaska) and comprise 81 percent of the U.S.
 L48 National Wildlife Refuge System, provide a breed-
 L49 ing habitat for millions of waterfowl and shorebirds
 L50 that winter in the lower 48 states. Wetlands are



Ponds across Alaska, including those shown above in the northeastern interior of the state, have shrunk as a result of increased evaporation and permafrost thawing. The pond in the top pair of images shrunk from 180 to 10 acres; the larger pond in the bottom pair of images shrunk from 90 to 4 acres.

also important to Native peoples who hunt and fish for their food in interior Alaska. Many villages are located adjacent to wetlands that support an abundance of wildlife resources. The sustainability of these traditional lifestyles is thus threatened by a loss of wetlands.

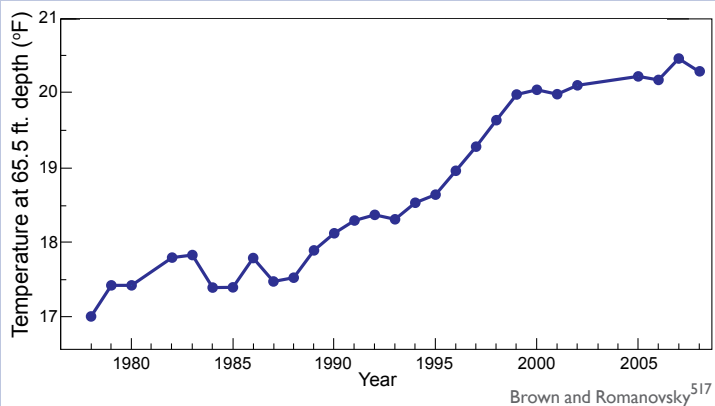
Thawing permafrost damages roads, runways, water and sewer systems, and other infrastructure.

Permafrost temperatures have increased throughout Alaska since the late 1970s.¹⁴⁹ The largest increases have been measured in the northern part of the state.⁵¹⁵ While permafrost in interior Alaska so far has experienced less warming than permafrost in northern Alaska, it is more vulnerable to thawing during this century because it is generally just below the freezing point, while permafrost in northern Alaska is colder.

Land subsidence (sinking) associated with the thawing of permafrost presents substantial challenges to engineers attempting to preserve infrastructure in Alaska.⁵¹⁶ Public infrastructure at risk for damage includes roads, runways, and water and sewer systems. It is estimated that thawing

R1
 R2
 R3
 R4
 R5
 R6
 R7
 R8
 R9
 R10
 R11
 R12
 R13
 R14
 R15
 R16
 R17
 R18
 R19
 R20
 R21
 R22
 R23
 R24
 R25
 R26
 R27
 R28
 R29
 R30
 R31
 R32
 R33
 R34
 R35
 R36
 R37
 R38
 R39
 R40
 R41
 R42
 R43
 R44
 R45
 R46
 R47
 R48
 R49
 R50

Permafrost Temperature, 1978 to 2008
Deadhorse, northern Alaska

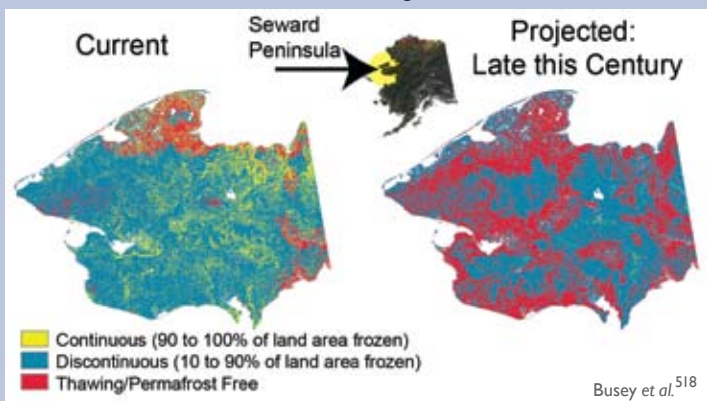


Permafrost temperatures have risen throughout Alaska, with the largest increases in the northern part of the state.

permafrost would add between \$3.6 billion and \$6.1 billion (10 to 20 percent) to future costs for publicly owned infrastructure by 2030 and between \$5.6 billion and \$7.6 billion (10 to 12 percent) by 2080.²³⁰ Analyses of the additional costs of permafrost thawing to private property have not yet been conducted.

Thawing ground also has implications for oil and gas drilling. As one example, the number of days per year in which travel on the tundra is allowed under Alaska Department of Natural Resources standards has dropped from more than 200 to about 100 days in the past 30 years. This results in a 50 percent reduction in days that oil and gas exploration and extraction equipment can be used.^{220,246}

Changing Permafrost Distribution
Moderate Warming Scenario



The maps show projected thawing on the Seward Peninsula by the end of this century under a moderate warming scenario approximately halfway between the lower and higher emissions scenarios⁹¹ described on page 23.

Coastal storms increase risks to villages and fishing fleets.

Alaska has more coastline than the 49 other states combined. Frequent storms in the Gulf of Alaska and the Bering, Chukchi, and Beaufort Seas already affect the coasts during much of the year. Alaska's coastlines, many of which are low in elevation, are increasingly threatened by a combination of the loss of their protective sea ice buffer, increasing storm activity, and thawing coastal permafrost.

Increasing storm activity in autumn in recent years⁵²⁰ has delayed or prevented barge operations that supply coastal communities with fuel. Commercial fishing fleets and other marine traffic are also strongly affected by Bering Sea storms. High-

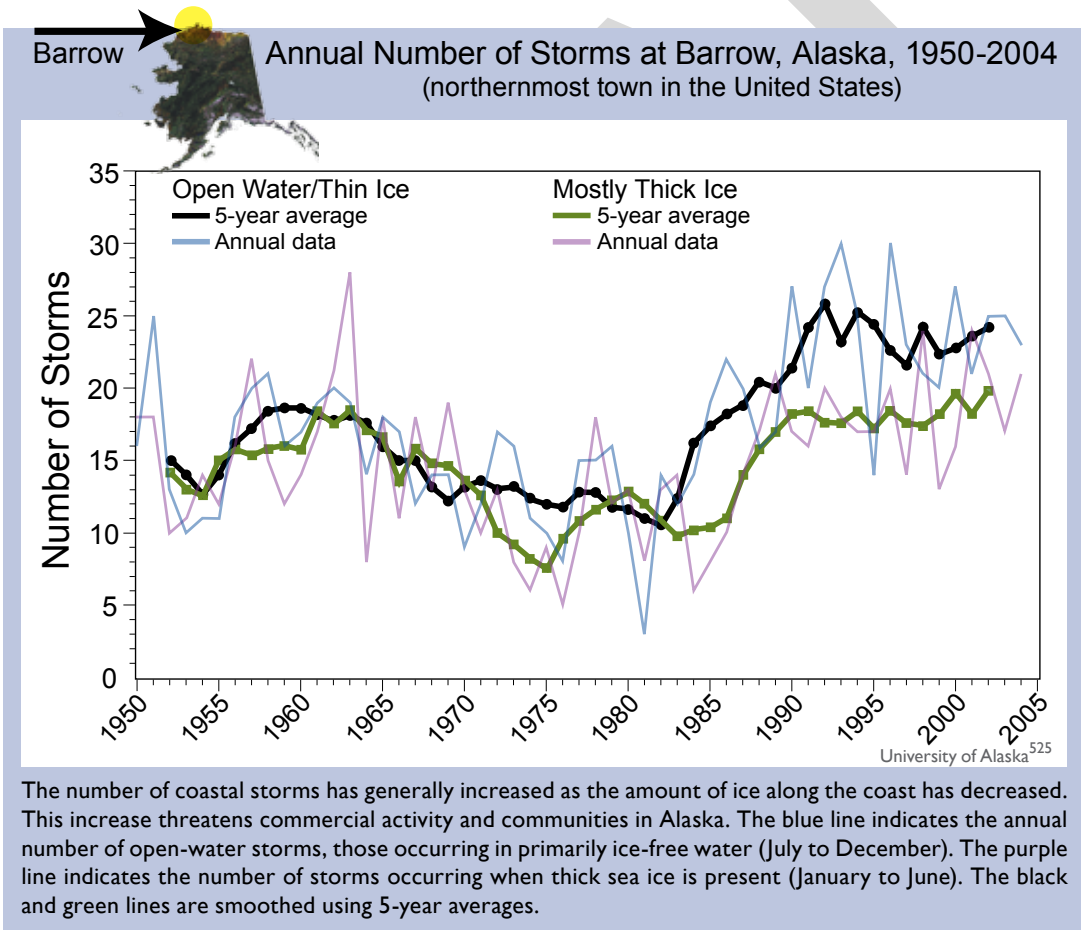
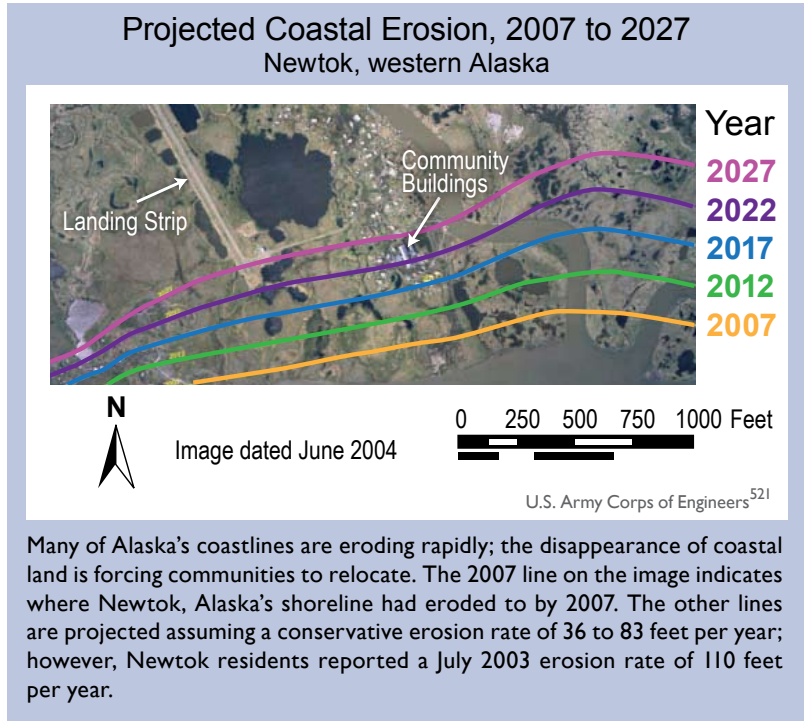
Adaptation: Keeping Soil Around the Pipeline Cool

When permafrost thaws, it can cause the soil to sink or settle, damaging structures built upon or within that soil. A warming climate and burial of supports for the Trans-Alaska Pipeline System both contribute to thawing of the permafrost around the pipeline. In locations on the pipeline route where soils were ice-rich, a unique above-ground system was developed to keep the ground cool. Thermal siphons were designed to disperse heat to the air that would otherwise be transferred to the soil, and these siphons were placed on the pilings that support the pipeline. While this unique technology added significant expense to the pipeline construction, it helps to greatly increase the useful lifetime of this structure.⁵¹⁹



L1 wind events have become more frequent along
 L2 the western and northern coasts. The same
 L3 regions are experiencing increasingly long
 L4 sea-ice-free seasons and hence longer periods
 L5 during which coastal areas are especially vul-
 L6 nerable to wind and wave damage. Downtown
 L7 streets in Nome, Alaska, have flooded in recent
 L8 years. Coastal erosion is causing the shorelines
 L9 of some areas to retreat at average rates of tens
 L10 of feet per year. The ground beneath several
 L11 native communities is literally crumbling into
 L12 the sea, forcing residents to confront difficult
 L13 and expensive choices between relocation and
 L14 engineering strategies that require continuing
 L15 investments despite their uncertain effective-
 L16 ness (see *Society* sector). The rate of erosion
 L17 along Alaska's northeastern coastline has
 L18 doubled over the past 50 years.⁵²²

L19
 L20 Over the coming century, an increase of sea
 L21 surface temperatures and a reduction of ice
 L22 cover are likely to lead to northward shifts in the Pacific storm track and increased impacts on coastal Alaska.^{523,524}
 L23 Climate models project the Bering Sea to experience the largest decreases in atmospheric pressure in the Northern
 L24 Hemisphere, suggesting an increase in storm activity in the region.⁹⁰ In addition, the longer ice-free season is likely to
 L25 make more heat and moisture available for storms in the Arctic Ocean, increasing their frequency and/or intensity.



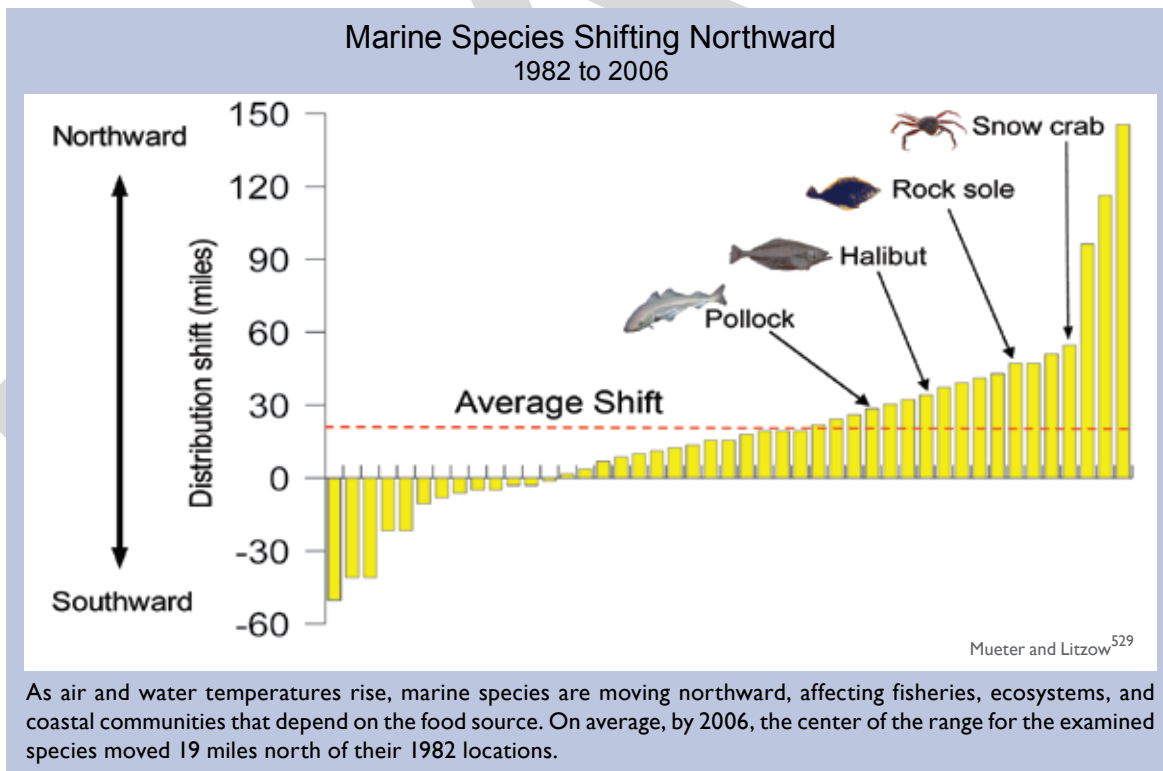
Displacement of marine species will affect key fisheries.

Alaska leads the United States in the value of its commercial fishing catch. Most of the nation’s salmon, crab, halibut, and herring come from Alaska. In addition, many Native communities depend on local harvests of fish, walrus, seals, whales, seabirds, and other marine species for their food supply. Climate change causes significant alterations in marine ecosystems with important implications for fisheries. Ocean acidification associated with a rising carbon dioxide concentration represents an additional threat to cold-water marine ecosystems^{23,526} (see *Ecosystems* sector and *Coasts* region).

One of the most productive areas for Alaska fisheries is the northern Bering Sea off Alaska’s west coast. The world’s largest single fishery is the Bering Sea pollock fishery, which has undergone major declines in recent years. Over much of the past decade, as air and water temperatures rose, sea ice in this region declined sharply. Populations of fish, seabirds, seals, walrus, and other species depend on plankton blooms that are regulated by the extent and location of the ice edge in spring. As the sea ice

retreats, the location, timing, and species composition of the blooms changes, reducing the amount of food reaching the living things on the ocean floor. This radically changes the species composition and populations of fish and other marine life forms, with significant repercussions for fisheries⁵²⁷ (see *Ecosystems* sector).

Over the course of this century, changes already observed on the shallow shelf of the northern Bering Sea are likely to affect a much broader portion of the Pacific-influenced sector of the Arctic Ocean. As such changes occur, the most productive commercial fisheries are likely to become more distant from existing fishing ports and processing infrastructure, requiring either relocation or greater investment in transportation time and fuel costs. These changes also will affect the ability of native peoples to successfully hunt and fish for the food they need to survive. Coastal communities already are noticing a displacement of walrus and seal populations. Bottom-feeding walrus populations are threatened when their sea ice platform retreats from the shallow coastal feeding grounds on which they depend.⁵²⁸



Islands



Climate change presents the Pacific and Caribbean islands with unique challenges. The U.S. affiliated Pacific Islands are home to approximately 1.7 million people in the Hawaiian Islands; Palau; the Samoan Islands of Tutuila, Manua, Rose, and Swains; and islands in the Micronesian archipelago, the Carolines, Marshalls, and Marianas.⁵³⁰ These include volcanic, continental, and limestone islands, atolls, and islands of mixed geologies.⁵³⁰ The degree to which climate change and variability will impact each of the roughly 30,000 islands in the Pacific depends upon a variety of factors, including the island's geology, area, height above sea level, extent of reef formation, and the size of its freshwater aquifer.⁵³¹

In addition to Puerto Rico and the U.S. Virgin Islands, there are 40 island nations in the Caribbean that are home to approximately 38 million people.⁵³² Population growth, often concentrated in coastal areas, escalates the vulnerability of both Pacific and Caribbean island communities to the effects of climate change, as do weakened traditional support systems. Tourism and fisheries, both of which are climate-sensitive, play a large economic role in these communities.⁵³⁰

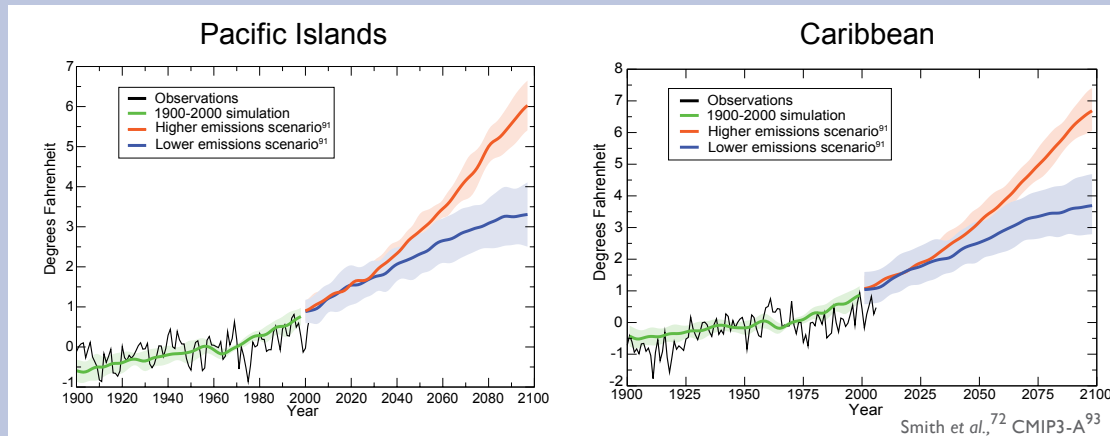
Small islands are considered among the most vulnerable to climate change because extreme events have major impacts on them. Changes in weather patterns and the frequency and intensity of extreme events, sea-level rise, coastal erosion, coral reef bleaching, ocean acidification, and contamination of freshwater resources by salt water are among the impacts small islands face.⁵³³

Islands have experienced rising temperatures and sea levels in recent decades. Projections for the rest of this century suggest:

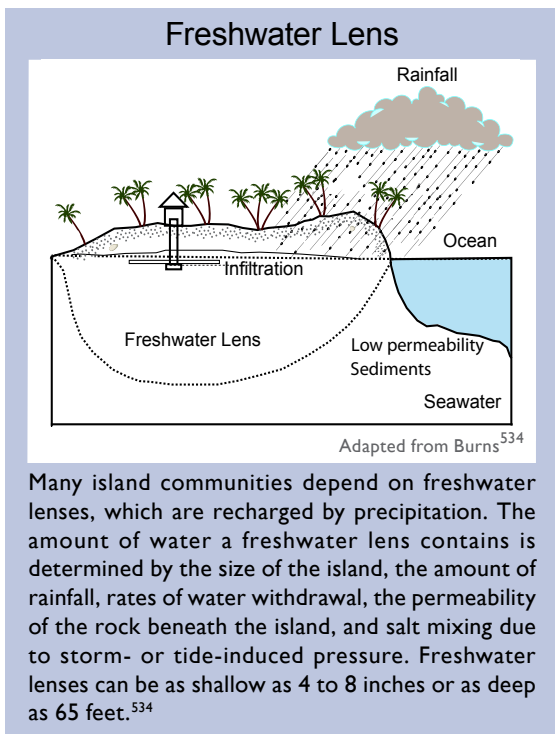
- Increases in air and ocean surface temperatures in both the Pacific and Caribbean,⁹⁰
- An overall decrease in rainfall in the Caribbean; and
- An increased frequency of heavy downpours and increased rainfall during summer months (rather than the normal rainy season in winter months) for the Pacific (although the range of projections regarding rainfall in the Pacific is still quite large).

The number of heavy rain events is very likely to increase.⁹⁰ Hurricane (typhoon) wind speeds and rainfall rates are likely to increase with continued

Air Temperature Change, Observed and Projected, 1900 to 2100
relative to 1960-1979 average



Air temperatures have increased over the last 100 years in both the Pacific Island and Caribbean regions. Larger increases are projected in the future, with higher emissions scenarios⁹¹ producing considerably greater increases. The shaded areas show the likely ranges while the lines show the central projections from a set of climate models.



Many island communities depend on freshwater lenses, which are recharged by precipitation. The amount of water a freshwater lens contains is determined by the size of the island, the amount of rainfall, rates of water withdrawal, the permeability of the rock beneath the island, and salt mixing due to storm- or tide-induced pressure. Freshwater lenses can be as shallow as 4 to 8 inches or as deep as 65 feet.⁵³⁴

warming.⁶⁸ Islands and other low-lying coastal areas will be at increased risk from coastal inundation due to sea-level rise and storm surge, with major implications for coastal communities, infrastructure, natural habitats, and resources.

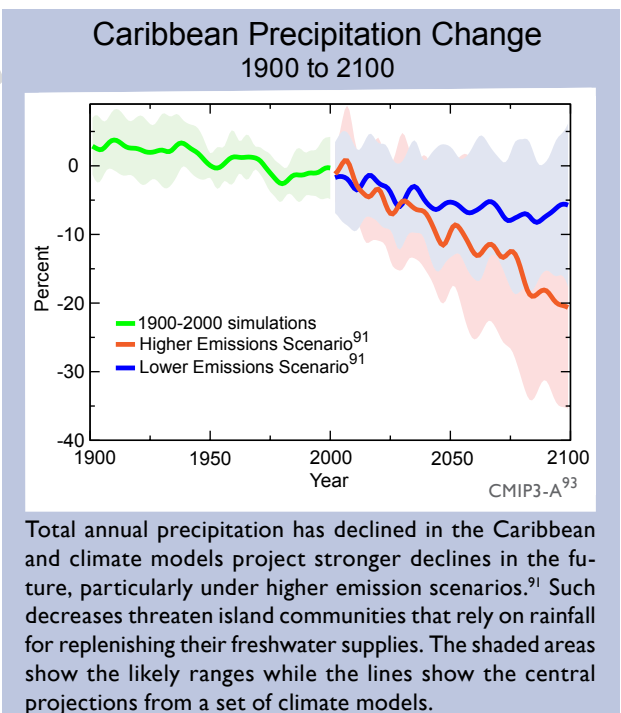
The availability of freshwater is likely to be reduced, with significant implications for island communities, economies, and resources.

Most island communities in the Pacific and the Caribbean have limited sources of the freshwater needed to support unique ecosystems and biodiversity, public health, agriculture, and tourism. Conventional freshwater resources include rainwater collection, groundwater, and surface water.⁵³⁴ For drinking and bathing, smaller Pacific islands primarily rely on individual rainwater catchment systems, while groundwater from the freshwater lens is used for irrigation. The size of freshwater lenses in atolls is influenced by factors such as rates of recharge (through precipitation), rates of use, and extent of tidal inundation.⁵³¹ Since rainfall triggers the formation of the freshwater lens, changes in precipitation, such as the significant decreases projected for the Caribbean, can significantly affect the availability of water. Because tropical storms replenish water supplies, potential changes in these storms are a great concern.

While it might be seen initially as a benefit, increased rainfall in the Pacific Islands during the summer months is likely to result in increased flooding, which would reduce drinking water quality and crop yields.⁵³⁴ In addition, many islands have weak distribution systems and old infrastructure, which result in significant water leakage, decreasing their ability to use freshwater efficiently. Water pollution (such as from agriculture or sewage), exacerbated by storms and floods, can contaminate the freshwater supply, impacting public health. Sea-level rise also impacts island water supplies by causing salt water to contaminate the freshwater lens and by causing an increased frequency of flooding due to storm high tides.⁵³¹ Finally, a rapidly rising population is straining the limited water resources, as would an increased incidence and/or intensity of storms⁵³⁴ or periods of prolonged drought.

Island communities, infrastructure, and ecosystems are vulnerable to coastal inundation due to sea-level rise and coastal storms.

Sea-level rise will have enormous effects on many island nations. Flooding will become more frequent due to higher storm tides, and coastal land will be permanently lost as the sea inundates low-



L1 lying areas and the shorelines erode. Loss of land
 L2 will reduce freshwater supplies⁵³¹ and affect living
 L3 things in coastal ecosystems. For example, the
 L4 Northwestern Hawaiian Islands, which are low-
 L5 lying and therefore at great risk from increasing sea
 L6 level, have a high concentration of endangered and
 L7 threatened species, some of which exist nowhere
 L8 else.⁵³⁵ The loss of nesting and nursing habitat is
 L9 expected to threaten the survival of already vulner-
 L10 able species.⁵³⁵

L12 In addition to gradual sea-level rise, extreme high
 L13 water level events can result from a combination
 L14 of coastal processes.²⁷¹ For example, the harbor
 L15 in Honolulu, Hawaii, experienced the highest
 L16 daily average sea level ever recorded in Septem-
 L17 ber 2003. This resulted from the combination of
 L18 long-term sea-level rise, normal seasonal heating
 L19 (which causes the volume of water to expand and
 L20 thus the level of the sea to rise), seasonal high tide,
 L21 and an ocean circulation event which temporarily
 L22 raised local sea level.⁵³⁶ The interval between such
 L23 extreme events has decreased from more than 20
 L24 years to approximately 5 years as average sea level
 L25 has risen.⁵³⁶

L27 Hurricanes, typhoons, and other storm events, with
 L28 their intense precipitation and storm surge, cause

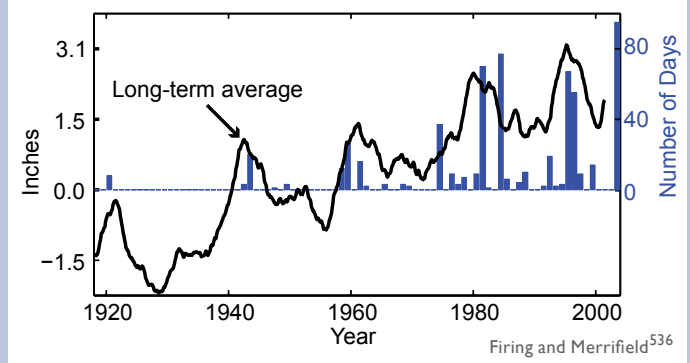
L30 **Adaptation: Securing Water Resources**

L32 In the islands, “water is gold.” Effective adaptation to climate-related
 L33 changes in the availability of freshwater is thus a high priority. While
 L34 island communities cannot completely counter the threats to water sup-
 L35 plies posed by global warming, effective adaptation approaches can help
 L36 reduce the damage.

L38 When existing resources fall short, managers look to unconventional
 L39 resources, such as desalinating seawater, importing water by ship, and
 L40 using treated wastewater for non-drinking uses. Desalination costs are
 L41 declining, though concerns remain about the impact on marine life, the
 L42 disposal of concentrated brines that might contain chemical waste, and the large energy use (and as-
 L43 sociated carbon footprint) of the process.¹⁴⁶ With limited natural resources, the key to successful water
 L44 resource management in the islands will continue to be “conserve, recover, and reuse.”⁵³⁰

L46 Pacific Island communities are also making use of the latest science. This effort started during the 1997 to
 L47 1998 El Niño, when managers began using seasonal forecasts to prepare for droughts by increasing public
 L48 awareness and encouraging water conservation. In addition, resource managers can improve infrastruc-
 L49 ture, such as by fixing water distribution systems to minimize leakage and by increasing freshwater stor-
 L50 age capacity.⁵³⁰

Extreme Sea-Level Days: Honolulu, Hawaii



Sea-level rise will result in permanent land loss and reductions in freshwater supplies, as well as threaten coastal ecosystems. “Extreme” sea-level days (with a daily average of more than 6 inches above the long-term average⁹⁰) can result from the combined effects of gradual sea-level rise due to warming and other phenomena, including seasonal heating and high tides.

major impacts to Pacific and Caribbean island communi-
 ties, including loss of life, damage to infrastructure and
 property, and contamination of freshwater supplies.⁵³⁷ As
 the climate continues to warm, the peak wind intensities
 and near-storm precipitation from future tropical cyclones
 are likely to increase,⁹⁰ which, combined with sea-level
 rise, is expected to cause higher storm surge levels. If
 such events occur frequently, communities would face
 challenges in recovering between events, resulting in
 long-term deterioration of infrastructure, freshwater and
 agricultural resources, and other impacts.²⁴⁷



A billboard on Pohnpei, in the Federated States of Micronesia, encourages water conservation in preparation for the 1997 to 1998 El Niño.

R1
 R2
 R3
 R4
 R5
 R6
 R7
 R8
 R9
 R10
 R11
 R12
 R13
 R14
 R15
 R16
 R17
 R18
 R19
 R20
 R21
 R22
 R23
 R24
 R25
 R26
 R27
 R28
 R29
 R30
 R31
 R32
 R33
 R34
 R35
 R36
 R37
 R38
 R39
 R40
 R41
 R42
 R43
 R44
 R45
 R46
 R47
 R48
 R49
 R50



Coastal houses and an airport in the U.S.-affiliated Federated States of Micronesia rely on mangroves' protection from erosion and damage due to rising sea level, waves, storm surges, and wind.

Critical infrastructure, including homes, airports, and roads, tends to be located along the coast. Flooding related to sea-level rise and hurricanes and typhoons negatively impacts port facilities and harbors, and causes closures of roads, airports, and bridges.⁵³⁸ Long-term infrastructure damage

would affect social services such as disaster risk management, health care, education, management of freshwater resources, and economic activity in sectors such as tourism and agriculture.

Climate changes affecting coastal and marine ecosystems will have major implications for tourism and fisheries.

Marine and coastal ecosystems of the islands are particularly vulnerable to the impacts of climate change. Sea-level rise, increasing water temperatures, rising storm intensity, coastal inundation, and flooding from extreme events, beach erosion, ocean acidification, increased incidences of coral disease, and increased invasions by non-native species are among the threats that endanger the ecosystems that provide safety, sustenance, economic viability, and cultural and traditional values to island communities.⁵³⁹

Tourism is a vital part of the economy for many islands. In 1999, the Caribbean had tourism-based gross earnings of \$17 billion, providing 900,000 jobs and making the Caribbean one of the most tourism dependent regions in the world.⁵³² In the South Pacific, tourism can contribute as much as 47 percent of gross domestic product.⁵⁴⁰ In Hawaii, tourism generated \$12.4 billion for the state in 2006, with over 7 million visitors.⁵⁴¹

Sea-level rise can erode beaches, and along with increasing water temperatures, can destroy or degrade natural resources such as mangroves and coral reef ecosystems that attract tourists.²⁴⁷ Extreme weather events can affect transportation systems and interrupt communications. The availability of

freshwater is critical to sustaining tourism, but is subject to the climate-related impacts described on the previous page. Public health concerns about diseases such as dengue would also negatively affect tourism.

Coral reefs sustain fisheries and tourism, have biodiversity value, scientific and educational value, and form natural protection against wave erosion.⁵⁴² For Hawaii alone, net benefits of reefs to the economy are estimated at \$360 million annually, and the overall asset value is conservatively estimated to be nearly \$10 billion.⁵⁴² In the Caribbean, coral reefs provide annual net benefits from fisheries, tourism, and shoreline protection services of between \$3.1 billion and \$4.6 billion. The loss of income by 2015 from degraded reefs is conservatively estimated at several hundred million dollars annually.^{532,543}

Coral reef ecosystems are particularly susceptible to the impacts of climate change, as even small increases in water temperature can cause coral bleaching,⁵⁴⁴ damaging and killing corals. Ocean acidification due to a rising carbon dioxide concentration poses an additional threat (see *Ecosystems* sector and *Coasts* region). Coral reef ecosystems are also especially vulnerable to invasive species.⁵⁴⁵ These impacts, combined with changes in the occurrence and intensity of El Niño events, rising sea level, and increasing storm damage,²⁴⁷ will have major negative effects on coral reef ecosystems.

Fisheries feed local people and island economies. Almost all communities within the Pacific Islands derive over 25 percent of their animal protein from fish, with some deriving up to 69 percent.⁵⁴⁶ For island fisheries sustained by healthy coral reef and marine ecosystems, climate change impacts exacerbate stresses such as overfishing,²⁴⁷ affecting both fisheries and tourism that depend on abundant and diverse reef fish. The loss of live corals results in local extinctions and a reduced number of reef fish species.⁵⁴⁷

Nearly 70 percent of the world's annual tuna harvest, approximately 3.2 million tons, comes from the Pacific Ocean.⁵⁴⁸ Climate change is projected to cause a decline in tuna stocks and an eastward shift in their location, affecting the catch of certain countries.²⁴⁷

R1
R2
R3
R4
R5
R6
R7
R8
R9
R10
R11
R12
R13
R14
R15
R16
R17
R18
R19
R20
R21
R22
R23
R24
R25
R26
R27
R28
R29
R30
R31
R32
R33
R34
R35
R36
R37
R38
R39
R40
R41
R42
R43
R44
R45
R46
R47
R48
R49
R50



Coasts

Approximately one-third of all Americans live in counties immediately bordering the nation's ocean coasts.^{549,550} In addition to accommodating major cities, the coasts and the exclusive economic zone extending 200 miles offshore provide enjoyment, recreation, seafood, transportation of goods, and energy. Coastal and ocean activities contribute more than \$1 trillion to the nation's gross domestic product and the ecosystems hold rich biodiversity and provide invaluable services.⁵⁵¹ However, intense human uses have taken a toll on coastal environments and their resources. Up to 38 percent of all fish stocks have been diminished by over-fishing, large "dead zones" depleted of oxygen have developed as a result of pollution by excess nitrogen runoff, toxic blooms of algae are increasingly frequent, coral reefs are badly damaged or becoming overgrown with algae, and about half of the nation's coastal wetlands have been lost – and most of this loss has occurred during the past 50 years.

Global climate change imposes additional stresses on coastal environments. Rising sea level is already eroding shorelines, drowning wetlands, and threatening the built environment.^{43,224} The destructive potential of Atlantic tropical storms and hurricanes has increased since 1970 in association with increasing Atlantic sea surface temperatures, and it is likely that hurricane rainfall and wind speeds will increase in response to global warming.¹¹² Coastal water temperatures have risen by about 2°F in several regions, and

the geographic distributions of marine species have shifted.^{37,68,347} Precipitation increases on land have increased river runoff, polluting coastal waters with more nitrogen and phosphorous, sediments, and other contaminants. Furthermore, increasing acidification resulting from the uptake of carbon dioxide by ocean waters threatens corals, shellfish, and other living things that form their shells and skeletons from calcium carbonate²³ (see *Ecosystems* sector). All of these forces converge and interact at the coasts, making these areas particularly sensitive to the impacts of climate change.

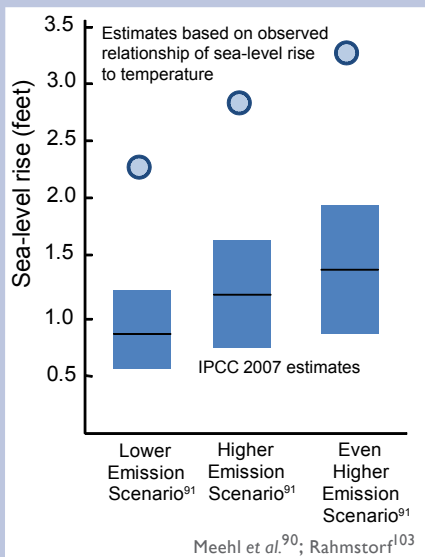
Significant sea-level rise and storm surge will adversely affect coastal cities and ecosystems around the nation; low-lying and subsiding areas are most vulnerable.

The rise in sea level relative to the land surface in any given location is a function of both the amount of global average sea-level rise and the degree to which the land is rising or falling. During the past century in the United States, relative sea level changes ranged from falling several inches to rising as much as 2 feet.²²⁵ High rates of relative sea-level rise, coupled with cutting off the supply of sediments from the Mississippi River and other human alterations, have resulted in the loss of 1,900 square miles of Louisiana's coastal wetlands during the past century, weakening their capacity

Multiple Stresses Confront Coastal Regions

Various forces of climate change at the coasts pose a complex array of management challenges and adaptation requirements. For example, relative sea level is expected to rise at least 2 feet in Chesapeake Bay (located between Maryland and Virginia) where the land is subsiding, threatening portions of cities, inhabited islands, most tidal wetlands, and other low-lying regions. Climate change also will affect the volume of the bay, its salinity distribution and circulation, as will changes in precipitation and freshwater runoff. These changes, in turn, will affect summertime oxygen depletion and efforts to reduce the agricultural nitrogen runoff that causes it. Meanwhile the warming of the bay's waters will make survival there difficult for northern species such as eelgrass and soft clams, while allowing southern species and invaders riding in ships' ballast water to move in and change the mix of species that are caught and must be managed. Additionally, more acidic waters resulting from rising carbon dioxide levels will make it difficult for oysters to build their shells and will complicate the recovery of this key species.⁵⁵³

Projected Sea-Level Rise



Estimates of sea-level rise by the end of the century for three emissions scenarios.⁹¹ Intergovernmental Panel on Climate Change 2007 projections (range shown as bars) exclude changes in ice sheet flow.⁹⁰ Light blue circles represent estimates derived using the observed relationship of sea-level rise to temperature.¹⁰³ Areas where coastal land is sinking, for example by as much as 1.5 feet in this century along portions of the Gulf Coast, would experience that much additional sea-level rise relative to the land.¹²⁸

to absorb the storm surge of hurricanes such as Katrina.⁵⁵² Shoreline retreat is occurring along most of the nation’s exposed shores.

The amount of sea-level rise likely to be experienced during this century depends mainly on the expansion of the ocean volume due to warming and the response of glaciers and polar ice sheets. Complex processes control the discharges from polar ice sheets and some are already producing substantial additions of water to the ocean.⁵⁵⁴ Because these processes are not well understood it is difficult to predict their future contributions to sea-level rise.^{90,555}

As discussed in the *Global Climate Change* section, recent estimates of global sea-level rise substantially exceed the IPCC estimates, suggesting sea-level rise between 3 and 4 feet in this century. Even a 2-foot rise in relative sea level over a century would result in the loss of a large portion of the nation’s remaining coastal wetlands, as they are not able to build new soil at a fast enough rate.¹⁶⁴ Accelerated sea-level rise would also affect sea-grasses, coral reefs, and other important habitats, fragment barrier islands, and place into jeopardy existing homes, businesses, and infrastructure, including roads, ports, and water and sewage systems. Portions of major cities, including Boston and New York, would be subject to inundation by ocean water during storm surges or even during regular high tides.²³⁴

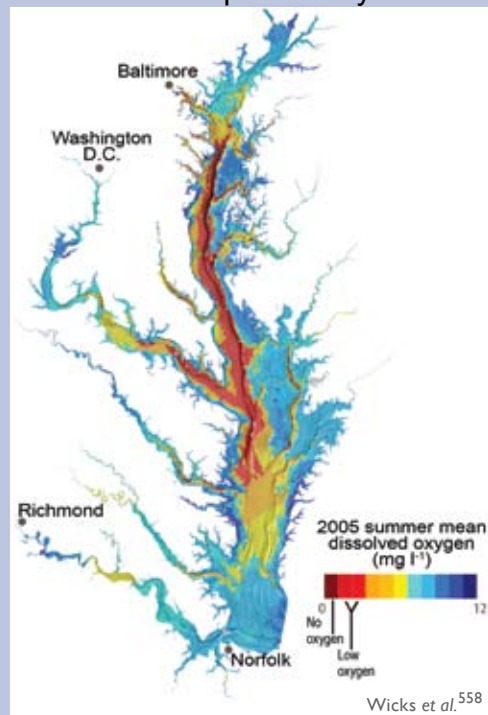


A “ghost swamp” in south Louisiana shows the effects of saltwater intrusion.

More spring runoff and warmer coastal waters will increase the seasonal reduction in oxygen resulting from excess nitrogen from agriculture.

Coastal dead zones in places such as the northern Gulf of Mexico⁵⁵⁶ and the Chesapeake Bay⁵⁵⁷ are likely to increase in size and intensity as warming increases unless efforts to control runoff of agricultural fertilizers are redoubled. Greater spring runoff into East Coast estuaries and the Gulf of Mexico would flush more nitrogen into coastal waters stimulating harmful blooms of algae and the excess production of microscopic plants that settle near the seafloor and deplete oxygen supplies as they decompose. In addition, greater runoff reduces salinity, which when coupled with warmer surface water increases the difference in density between surface and bottom waters, thus preventing the replacement of oxygen in the deeper waters. As dissolved oxygen levels decline below a certain level, living things cannot survive. They leave the area if they can, and die if they cannot.

Dead Zones in the Chesapeake Bay



Climate change is likely to expand and intensify “dead zones,” areas where bottom water is depleted of dissolved oxygen because of nitrogen pollution, threatening living things.

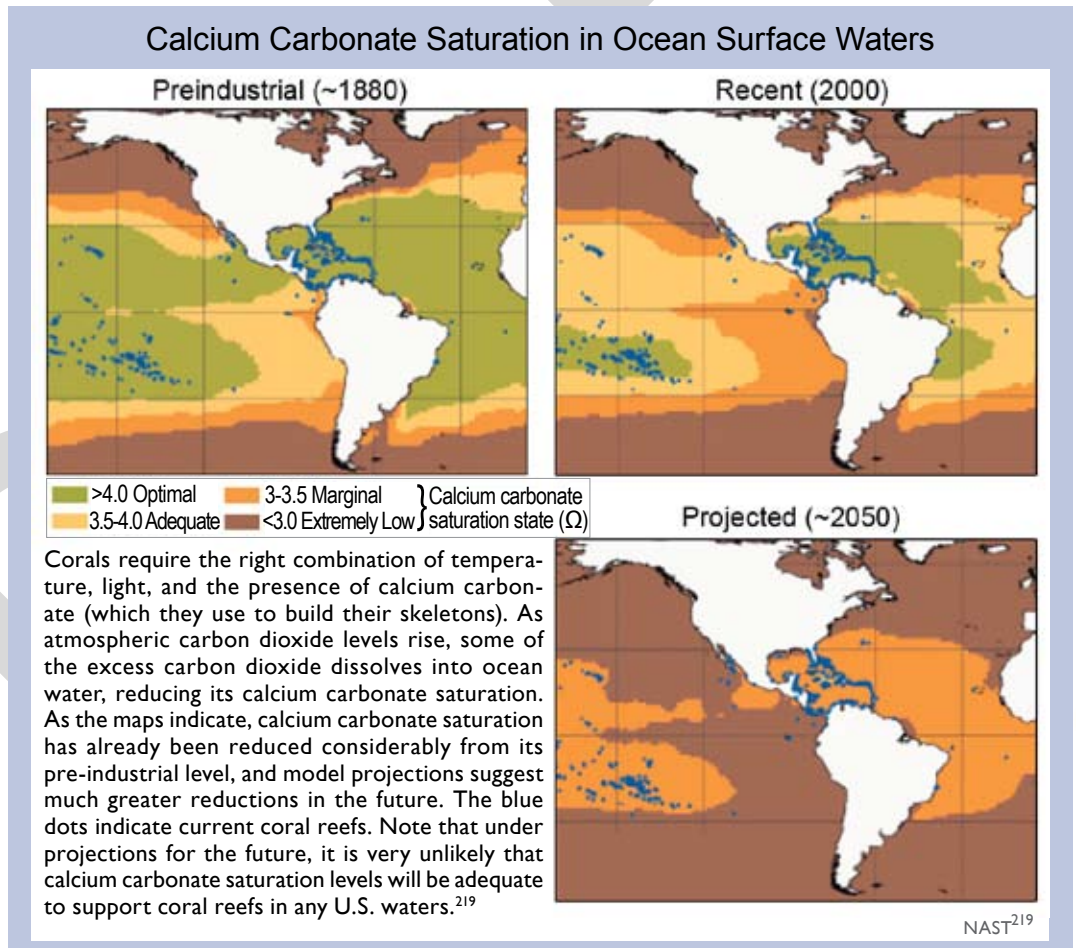
Coastal waters are very likely to continue to warm by as much 4 to 8°F in this century, both in summer and winter.²³⁴ As with animals and plants on land, this will result in a northward shift in the geographic distribution of marine life along the coasts; this is already being observed.^{70,347} Species that cannot tolerate the higher temperatures will move northward while species from farther south move in. Warming also opens the door to invasion by species that humans are intentionally or unintentionally transporting around the world, for example in the ballast water carried by ships. Species that were previously unable to establish populations because of cold winters are likely to find the warmer conditions more welcoming and gain a foothold, particularly as native species are under stress from climate change and other human activities. Non-native clams and small crustaceans already have had major effects on the San Francisco Bay ecosystem and the health of its fishery resources.⁵⁵⁹

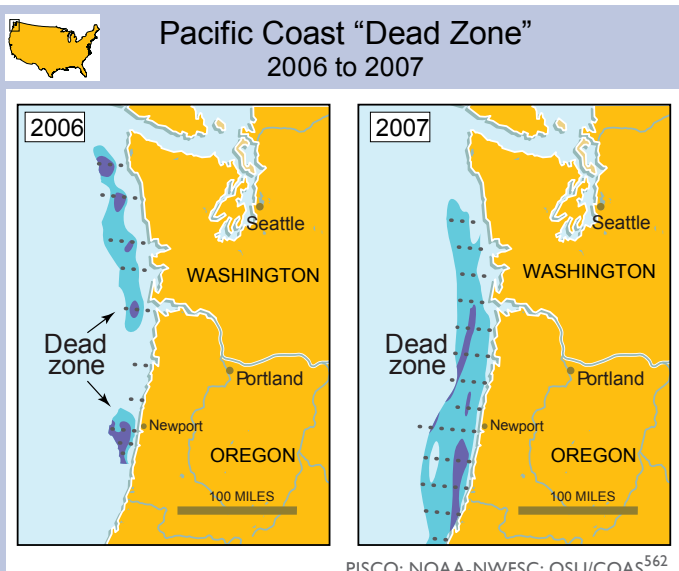
Higher water temperatures and ocean acidification due to increasing atmospheric carbon dioxide will present major additional stresses to coral reefs, resulting in significant die-offs and limited recovery.

In addition to carbon dioxide’s heat-trapping effect, the increase in its concentration in the atmosphere is gradually acidifying the ocean. About one-third of the carbon dioxide emitted by human activities has been absorbed by the ocean, resulting in a decrease in the ocean’s pH. Since the beginning of the industrial era, ocean pH has declined demonstrably and is projected to decline much more by 2100 if current emissions trends continue. Further declines in pH

are very likely to continue to affect the ability of living things to create and maintain shells or skeletons of calcium carbonate. This is because at a lower pH less of the dissolved carbon is available as carbonate ions. The living things affected include important plankton species in the open ocean, mollusks and other shellfish, and reef-building corals.^{70,260}

Ocean acidification (see *Global Climate Change*) reduces the ability of living things to create calcium carbonate-containing shells or skeletons. The living things affected include important plankton species in the open ocean, mollusks and other shellfish, and corals.^{22,23,70} The effects on reef-building corals are likely to be particularly severe during this century. Coral calcification rates are likely to decline by more than 30 percent under a doubling of atmospheric carbon dioxide concentrations, with erosion outpacing reef formation at even lower concentrations.²² In addition, the reduction in pH also affects photosynthesis, growth, and reproduction. The upwelling of deeper ocean water, deficient in carbonate, and thus potentially detrimental to the food chains supporting juvenile salmon has recently been observed along the U.S. West Coast.²⁶⁰





Climate change affects coastal currents that moderate ocean temperatures and the productivity of ecosystems. As such, it is believed to be a factor in the low-oxygen “dead zone” that has appeared along the coast of Washington and Oregon in recent years.⁵⁶¹ In the maps above, light blue indicates low-oxygen areas and purple shows areas that are the most severely oxygen depleted.

Acidification imposes yet another stress on reef-building corals, which are also subject to bleaching – the expulsion of the microscopic plants that live inside the corals and are essential to their survival – as a result of heat stress⁷⁰ (see *Ecosystems* sector and *Islands* region). As a result of these and other stresses, the corals that form the reefs in the Florida Keys, Puerto Rico, Hawaii, and the Pacific Islands are projected to be lost if carbon dioxide concentrations continue to rise at their current rate.⁵⁶⁰

Changing ocean currents will affect coastal ecosystems.

Because it affects the distribution of heat in the atmosphere and the oceans, climate change will affect the currents that move along the nation’s coasts, such as the California Current that bathes the West Coast from British Columbia to Baja California.⁷⁰ This southward flowing current produces upwelling of deeper ocean water along the coast that is vital to moderation

of temperatures and the high productivity of Pacific Coast ecosystems. Such coastal currents are subject to periodic variations caused by the El Niño-Southern Oscillation and the Pacific Decadal Oscillation, which have substantial effects on the success of salmon and other fishery resources. Climate change is expected to affect such coastal currents, and possibly the larger scale natural oscillations as well, though these effects are not well understood yet. The recent emergence of oxygen-depletion events on the continental shelf off Oregon and Washington (a dead zone not directly caused by agricultural runoff and waste discharges such as those in the Gulf of Mexico or Chesapeake Bay) is one example.⁵⁶¹

Adaptation: Coping with Sea-Level Rise

Adaptation to sea-level rise is already taking place in three main categories: (1) protection by building hard structures such as levees and seawalls (although hard structures can, in some cases, actually increase risks and worsen beach erosion and wetland retreat), (2) accommodation by elevating or redesigning structures, enhancing wetlands, or adding sand from elsewhere to beaches (the latter is not a permanent solution, and can encourage development in vulnerable locations), and (3) planned retreat from the coastline as sea level rises.²⁷⁰



Several states have laws or regulations that require setbacks for construction based on the planned life of the development and observed erosion rates. Michigan, North Carolina, Rhode Island, and South Carolina are using such a moving baseline to guide planning. Maine’s Coastal Sand Dune Rules prohibit buildings of a certain size that are unlikely to remain stable with a sea-level rise of 2 feet. The Massachusetts Coastal Hazards Commission is preparing a 20-year infrastructure and protection plan to improve hazards management and the Maryland Commission on Climate Change has recently made comprehensive recommendations to reduce the state’s vulnerability to sea-level rise and coastal storms by addressing building codes, public infrastructure, zoning, and emergency preparedness. Governments and private interests are beginning to take sea-level rise into account in planning levees and bridges, and in the siting and design of facilities such as sewage treatment plants (see *Northeast* region).

L1
L2
L3
L4
L5
L6
L7
L8
L9
L10
L11
L12
L13
L14
L15
L16
L17
L18
L19
L20
L21
L22
L23
L24
L25
L26
L27
L28
L29
L30
L31
L32
L33
L34
L35
L36
L37
L38
L39
L40
L41
L42
L43
L44
L45
L46
L47
L48
L49
L50

R1
R2
R3
R4
R5
R6
R7
R8
R9
R10
R11
R12
R13
R14
R15
R16
R17
R18
R19
R20
R21
R22
R23
R24
R25
R26
R27
R28
R29
R30
R31
R32
R33
R34
R35
R36
R37
R38
R39
R40
R41
R42
R43
R44
R45
R46
R47
R48
R49
R50



An Agenda for Climate Impacts Science

Both mitigation and adaptation decisions are becoming increasingly necessary. Advancing our knowledge in the many aspects of science that affect the climate system has already contributed greatly to decision making on climate change issues. Further advances in climate science including better understanding and projections regarding rainfall, storm tracks, storm intensity, heat waves, and sea-level rise will improve decision making capabilities.

The focus below, however, is on advancing our knowledge specifically on climate change impacts and those aspects of climate change responsible for these impacts in order to continue to guide decision making.

Recommendation 1: Expand our understanding of climate change impacts.

There is a clear need to increase understanding of how ecosystems, social and economic systems, human health, and infrastructure will be affected by climate change in the context of other stresses. New understanding will come from a mix of activities including sustained and systematic observations, field and laboratory experiments, model development, and integrated impact assessments. These will incorporate shared learning among researchers, practitioners (such as engineers and water managers), and local stakeholders.

Ecosystems

Ecosystem changes, in response to changes in climate and other environmental conditions, have already been documented. These include changes in the chemistry of the atmosphere and precipitation, vegetation patterns, growing season length, plant productivity, animal species distributions, and the frequency and severity of pest outbreaks and fires. In the marine environment,

changes include the health of corals and other living things due to temperature stress and ocean acidification. These observations not only document climate-change impacts, but also provide critical input to understanding how and why these changes occur, and how changes in ecosystems in turn affect climate. In this way, records of observed changes can improve projections of future impacts related to various climate-change scenarios.

In addition to observations, large-scale, whole-ecosystem experiments are essential for improving projections of impacts. Ecosystem-level experiments that vary multiple factors, such as temperature, moisture, ground-level ozone, and atmospheric carbon dioxide, would provide process-level understanding of the ways ecosystems could respond to climate change in the context of other environmental stresses. Such experiments are particularly important for ecosystems with the greatest potential to experience massive change due to the crossing of thresholds or tipping points.

Insights regarding ecosystem responses to climate change gained from both observations and experiments are the essential building blocks of ecosystem simulation models. These models, when rigorously developed and tested, provide powerful tools for exploring the ecosystem consequences of alternative future climates. The incorporation of ecosystem models into an integrated assessment framework that includes socioeconomic, atmospheric and ocean chemistry, and atmospheric-ocean general circulation models should be a major goal of impacts research. This knowledge can provide a base for research studies into ways to manage critical ecosystems in an environment that is continually changing.

Economic systems, human health, and the built environment

As natural systems experience variations due to a changing climate, social and economic systems will

L1 be affected. Food production, water resources, forests,
 L2 parks, and other managed systems provide life support
 L3 for society. Their sustainability will depend on how well
 L4 they can adapt to a future climate that is different from
 L5 historical experience.

L6
 L7 At the same time, climate change is exposing human
 L8 health and the built environment to increasing risks.
 L9 Among the likely impacts are an expansion of the
 L10 ranges of insects and other animals that carry diseases
 L11 and a greater incidence of health-threatening air pollu-
 L12 tion events compounded by unusually hot weather as-
 L13 sociated with climate change. In coastal areas, sea-level
 L14 rise and storm surge threaten infrastructure including
 L15 homes, roads, ports, and oil and gas drilling and distri-
 L16 bution facilities. In other parts of the country, floods,
 L17 droughts, and other weather and climate extremes pose
 L18 increasing threats.

L19
 L20 Careful observations along with climate and Earth
 L21 system models run with a range of emissions scenarios
 L22 can help society evaluate these risks and plan actions to
 L23 minimize them. Work in this area would include assess-
 L24 ments of the performance of delivery systems, such as
 L25 those for regional water and electricity supply, so that
 L26 climate change impacts and costs can be evaluated in
 L27 terms of changes in risk to system performance. It will
 L28 be particularly important to understand when the effects
 L29 on these systems are extremely large and/or rapid,
 L30 similar to tipping points and thresholds in ecosystems.

L31
 L32 In addition, the climate change experienced outside the
 L33 United States will have implications for our nation. A
 L34 better understanding of these international linkages,
 L35 including those related to trade, security, and large-scale
 L36 movements of people in response to climate change, is
 L37 required.

L38
 L39
 L40 **Recommendation 2:**
 L41 **Refine ability to project climate change,**
 L42 **including extreme events, at local scales.**
 L43

L44 One of the main messages to emerge from the past
 L45 decade of synthesis and assessments is that while
 L46 climate change is a global issue, it has a great deal of
 L47 regional variability. There is an indisputable need to
 L48 improve understanding of climate system effects at
 L49 these smaller scales, because these are often the scales
 L50 of decision-making in society. Understanding impacts at



R1 local scales will also help to target finite resources for
 R2 adaptation measures. Although much progress has been
 R3 made in understanding important aspects of this vari-
 R4 ability, uncertainties remain. Further work is needed on
 R5 how to quantify cumulative uncertainties across spatial
 R6 scales and the uncertainties associated with complex,
 R7 intertwined natural and social systems.

R8
 R9 Because region-specific climate changes will occur in
 R10 the context of other environmental and social changes
 R11 that are also region-specific, it is important to continue
 R12 to refine our understanding of regional details, espe-
 R13 cially those related to precipitation and soil moisture.
 R14 This requires further testing of models against observa-
 R15 tions using established metrics designed to evaluate and
 R16 improve the realism of regional model simulations.

R17
 R18 Success will also require continued development of
 R19 improved, higher resolution global climate models and
 R20 increased computational capacity, as well as extensive
 R21 climate model experiments. Improved downscaling
 R22 methods will increase the value of geographically
 R23 specific climate projections for decision makers in
 R24 government, business, and the general population.

R25
 R26 Extreme weather and climate events are a key com-
 R27 ponent of regional climate. Additional attention needs
 R28 to be focused on improved observations (made on the
 R29 relevant time and space scales to capture high-impact
 R30 extreme events) and associated research and analysis of
 R31 the potential for future changes in extremes. Impacts
 R32 analyses indicate that extreme weather and climate
 R33 events often play a major role in determining climate-
 R34 change consequences.

R35
 R36
 R37 **Recommendation 3:**
 R38 **Expand capacity to provide decision makers**
 R39 **and the public with relevant information on**
 R40 **climate change and its impacts.**
 R41

R42 The United States has tremendous potential to create
 R43 more comprehensive measurement, archive, and data-
 R44 access systems and to convey needed information that
 R45 could provide great benefit to society. There are several
 R46 aspects to fulfilling this goal: defining what is most
 R47 relevant, gathering the needed information, expanding
 R48 the capacity to deliver information, and improving the
 R49 tools for decision makers to use this information to the
 R50 best advantage. All of these aspects should involve an

L1 interactive and iterative process of continual learning
 L2 between those who provide information and those who
 L3 use it. Through such a process, monitoring systems,
 L4 distribution networks, and tools for using information
 L5 can all be refined to meet user needs.
 L6

L7 For example, tools used by researchers that could also
 L8 be useful to decision makers include those that analyze
 L9 and display the probability of occurrence of a range of
 L10 outcomes to help in assessing risks.
 L11

L12 Improved climate monitoring can be efficiently
 L13 achieved by following the Climate Monitoring Prin-
 L14 ciples recommended by the National Academy of
 L15 Sciences and the Climate Change Science Strategic Plan
 L16 in addition to integrating current efforts of governments
 L17 at all levels. Such a strategy complements a long-term
 L18 commitment to the measurement of the set of essential
 L19 climate variables identified by both the Climate Change
 L20 Science Program and the Global Climate Observing
 L21 System. Attention must be placed on the variety of time
 L22 and space scales critical for decision making.
 L23

L24 Improved impacts monitoring would include informa-
 L25 tion on the physical and economic effects of extreme
 L26 events (such as floods and droughts), available, for
 L27 example, from emergency preparedness and resource
 L28 management authorities. This would require regular
 L29 archiving of information about impacts.
 L30

L31 Improved access to data and information archives could
 L32 substantially enhance society's ability to respond to
 L33 climate change. While many data related to climate
 L34 impacts are already freely and readily available to a
 L35 broad range of users, other data, such as damage costs,
 L36 are not, and efforts should be made to make them
 L37 available. Easily accessible information should include
 L38 a set of agreed-upon baseline indicators and measures
 L39 of environmental conditions that can be used to track
 L40 the effects of changes in climate. Services that provide
 L41 reliable, well-documented, and easily used climate
 L42 information are an essential part of this much-needed
 L43 capacity, and where those services are not yet available,
 L44 they should be made available to support uses identified
 L45 as highest priority.
 L46
 L47
 L48
 L49
 L50

Recommendation 4: Improve understanding of thresholds likely to lead to abrupt changes in climate or ecosystems.

R1
 R2
 R3
 R4
 R5
 R6
 R7
 R8
 R9
 R10
 R11
 R12
 R13
 R14
 R15
 R16
 R17
 R18
 R19
 R20
 R21
 R22
 R23
 R24
 R25
 R26
 R27
 R28
 R29
 R30
 R31
 R32
 R33
 R34
 R35
 R36
 R37
 R38
 R39
 R40
 R41
 R42
 R43
 R44
 R45
 R46
 R47
 R48
 R49
 R50

Paleoclimatic data show that climate can and has
 changed quite abruptly when certain thresholds are
 crossed. Similarly, there is evidence that ecological
 and human systems can undergo abrupt change when
 tipping points are reached.

Within the climate system there are a number of key
 risks to society for which understanding is still quite
 limited. Additional research is needed in some key
 areas, for example, identifying thresholds that lead to
 rapid changes in ice sheet dynamics. Sea-level rise is
 a major concern and improved understanding of the
 sensitivity of the major ice sheets to sustained warming
 requires improved observing capability, analysis, and
 modeling of the ice sheets and their interactions with
 nearby oceans. Estimates of sea-level rise in previous
 assessments, such as the recent Intergovernmental Panel
 on Climate Change 2007 report, did not attempt to fully
 quantify the magnitude and rate of future sea-level rise
 due to inadequate scientific understanding of potential
 instabilities of the Greenland and Antarctic ice sheets.

Tipping points in biological systems include the tem-
 perature thresholds above which insects survive winter,
 and can complete two life cycles instead of one in a
 single growing season, contributing to infestations that
 kill large numbers of trees. The devastation caused by
 bark beetles in Canada, and increasingly in the U.S.
 West, provides an example of how crossing such a
 threshold can set off massive destruction in an ecosys-
 tem with far-reaching consequences. Similarly, there is
 increasing concern about the acidification of the world's
 oceans due to rising atmospheric carbon dioxide levels.
 There are ocean acidity thresholds beyond which corals
 and other living things, including some that form the
 base of important marine food chains, will no longer
 be able to form the shells and other body structures
 they need to survive. Improving understanding of such
 thresholds is an important goal for future research.

**Recommendation 5:
Improve understanding of the most effective ways to reduce the rate and magnitude of climate change, as well as unintended consequences of such activities.**

This report underscores the importance of reducing the concentrations of heat-trapping gases in the atmosphere. Impacts of climate change during this century and beyond are projected to be far larger and more rapid in scenarios in which greenhouse gas concentrations continue to grow rapidly compared to scenarios in which concentrations grow more slowly. Additional research will help identify the desired mix of mitigation options necessary to control the rate and magnitude of climate change.

In addition to their intended reduction of atmospheric concentrations of greenhouse gases, mitigation options also have the potential for unintended consequences, which should also be examined in future research. For example, widespread adoption of biofuels could lead to increases in water and fertilizer use and create competition among land uses for food production, biofuels production, and natural ecosystems that provide many benefits to society. Improved understanding of such unintended consequences, and identification of those options that carry the largest negative impacts, can help decision makers make more informed choices regarding the possible tradeoffs inherent in various mitigation strategies.

**Recommendation 6:
Enhance understanding of how society can adapt to climate change.**

There is currently limited knowledge about the ability of communities, regions, and sectors to adapt to future climate change. It is essential to improve understanding of how to enhance society's capacity to adapt to a changing climate in the context of other environmental stresses. Priority should be placed on interdisciplinary research on adaptation that takes into account the interconnectedness of the Earth system and the complex nature of the social, political, and economic environment in which adaptation decisions must be made.

The potential exists to provide insights into the possible effectiveness and limits of adaptation options that might be considered in the future. To realize this potential, new research will be required that documents past responses to climate variability and other environmental changes, analyzes the underlying reasons for them, and explains how individual and institutional decisions were made. However, human-induced climate change is projected to be larger and more rapid than any experienced by modern society so there are limits to what can be learned from the past.

A major difficulty in the analysis of adaptation strategies in this report has been the lack of information about the potential costs of adaptation measures, their effectiveness under various scenarios of climate change, the time horizons required for their implementation, and unintended consequences. These types of information should be systematically gathered and shared with decision makers as they consider a range of adaptation options. It is also clear that there is a substantial gap between the available information about climate change and the development of new guidelines for infrastructure such as housing, transportation, water systems, commercial buildings, and energy systems. There are also social and institutional obstacles to appropriate action, even in the face of adequate knowledge. These obstacles need to be better understood so that they can be reduced or eliminated.

Finally, it is important to carry out regular assessments of adaptation measures that address combined scenarios of future climate change, population growth, and economic development paths. This is an important opportunity for shared learning in which researchers, practitioners, and stakeholders collaborate using observations, models, and dialogue to explore adaptation as part of long-term sustainable development planning.





Concluding Thoughts

Responding to changing conditions

Human-induced climate change is happening now, and impacts are already apparent. Greater impacts are projected, particularly if heat-trapping gas emissions continue unabated. Previous assessments have established these facts, and this report confirms, solidifies, and extends these conclusions for the United States. It reports the latest understanding of how climate change is already affecting important sectors and regions. In particular, it reports that some climate change impacts appear to be increasing faster than previous assessments had suggested. This report represents a significant update to previous work, as it draws from the U.S. Climate Change Science Program's Synthesis and Assessment Products and other recent studies that examine how climate change and its effects are projected to continue to increase over this century and beyond.

Climate choices

Choices about emissions now and in the coming years will have far-reaching consequences for climate change impacts. A consistent finding of this assessment is that the rate and magnitude of future climate change and resulting impacts depend critically on the level of global atmospheric heat-trapping gas concentrations as well as the types and concentrations of atmospheric particles (aerosols). Lower emissions of heat-trapping gases will delay the appearance of climate change impacts and lessen their magnitude. Unless the rate of emissions is substantially reduced, impacts are expected to become increasingly severe for more people and places.

Similarly, there are choices to be made about adaptation strategies that can help to reduce or avoid some of the undesirable impacts of climate change. There is much to learn about the effectiveness of the various types of adaptation responses and how they will interact with each other and with mitigation actions.

Responses to the climate change challenge will almost certainly evolve over time as society learns by doing. Determining and refining societal responses will be an iterative process involving scientists, policymakers, and public and private decision makers at all levels. Implementing these response strategies will require careful planning and continual feedback on the impacts of mitigation and adaptation policies for government, industry, and society.

The value of assessments

Science has revolutionized our ability to observe and model the Earth's climate and living systems, to understand how they are changing, and to project future changes in ways that were not possible in prior generations. These advances have enabled the assessment of climate change, impacts, vulnerabilities, and response strategies. Assessments serve a very important function in providing the scientific underpinnings of informed policy. They can identify advances in the underlying science, provide critical analysis of issues, and highlight key findings and key unknowns that can guide decision making. Regular assessments also serve as progress reports to evaluate and improve policy making and other types of decision making related to climate change.

L1
 L2 Impacts and adaptation research includes complex
 L3 human dimensions, such as economics, manage-
 L4 ment, governance, behavior, and equity. Compre-
 L5 hensive assessments provide an opportunity to
 L6 evaluate the social implications of climate change
 L7 within the context of larger questions of how com-
 L8 munities and the nation as a whole create sustain-
 L9 able development paths.

R1
 R2
 R3
 R4
 R5
 R6
 R7
 R8
 R9

L10
 L11

R10
 R11

A vision for future U.S. assessments

R12

L12
 L13
 L14 Over the past decade, U.S. federal agencies have
 L15 undertaken two coordinated, national-scale efforts
 L16 to evaluate the impacts of global climate change
 L17 on this country. Each effort produced a report to
 L18 the nation – *Climate Change Impacts on the United*
 L19 *States*, published in 2000, and this report, *Global*
 L20 *Climate Change Impacts in the United States*,
 L21 published in 2009. A unique feature of the first
 L22 report was that in addition to reporting the current
 L23 state of the science, it created a national discourse
 L24 on climate change that involved hundreds of sci-
 L25 entists and thousands of stakeholders including
 L26 farmers, ranchers, resource managers, city planners,
 L27 business people, and local and regional government
 L28 officials. A notable feature of the second report is
 L29 the incorporation of information from the 21 topic-
 L30 specific Synthesis and Assessment Products, many
 L31 motivated by stakeholder interactions.

R13
 R14
 R15
 R16
 R17
 R18
 R19
 R20
 R21
 R22
 R23
 R24
 R25
 R26
 R27
 R28
 R29
 R30
 R31

L32
 L33 A vision for future climate change assessments
 L34 includes both sustained, extensive stakeholder
 L35 involvement, and targeted, scientifically rigorous
 L36 reports that address concerns in a timely fashion.
 L37 The value of stakeholder involvement includes
 L38 helping scientists understand what information
 L39 society wants and needs. In addition, the problem-
 L40 solving abilities of stakeholders will be essential to
 L41 designing, initiating, and evaluating mitigation and
 L42 adaptation strategies and their interactions. The best
 L43 decisions about these strategies will come when
 L44 there is widespread understanding of the complex
 L45 issue of climate change – the science and its many
 L46 implications for our nation.

R32
 R33
 R34
 R35
 R36
 R37
 R38
 R39
 R40
 R41
 R42
 R43
 R44
 R45
 R46

L47
 L48
 L49
 L50

R47
 R48
 R49
 R50



Federal Advisory Committee Authors

David M. Anderson is the Director for the World Data Center for Paleoclimatology, Chief of the Paleoclimatology Branch of NOAA's National Climatic Data Center, and an Associate Professor at the University of Colorado.

Donald F. Boesch is President of the University of Maryland Center for Environmental Science. His area of expertise is biological oceanography.

Virginia R. Burkett is the Chief Scientist for Global Change Research at the U.S. Geological Survey. Her areas of expertise are coastal ecology, wetland management, and forestry.

Lynne M. Carter is the Director of the Adaptation Network, a non-profit organization, and a project of the Earth Island Institute. Through assessment and action, she works to build resilience in communities and ecosystems in the face of a changing climate.

Stewart J. Cohen is senior researcher with the Adaptation and Impacts Research Division of Environment Canada, and an Adjunct Professor with the Department of Forest Resources Management of the University of British Columbia.

Nancy B. Grimm is a Professor of Life Sciences at Arizona State University. She studies how human-environment interactions and climate variability influence biogeochemical processes in both riverine and urban ecosystems.

Jerry L. Hatfield is the Laboratory Director of the USDA-ARS National Soil Tilth Laboratory in Ames, Iowa. His expertise is in the quantifications of spatial and temporal interactions across the soil-plant-atmosphere continuum.

Katharine Hayhoe is a Research Associate Professor in the Department of Geosciences at Texas Tech University and Principal Scientist and CEO of ATMOS Research & Consulting. Her research examines the potential impacts of human activities on the global environment.

Anthony C. Janetos is the Director of the Joint Global Change Research Institute, a joint venture between the Pacific Northwest National Laboratory and the University of Maryland. His area of expertise is biology.

Thomas R. Karl, (Co-Chair), is the Director of NOAA's National Climatic Data Center. His areas of expertise include monitoring for climate change and changes in extreme climate and weather events. He is also president of the American Meteorological Society.

Jack A. Kaye currently serves as Associate Director for Research of the Earth Science Division within NASA's Science Mission Directorate. He is responsible for NASA's research and data analysis programs in Earth System Science.

Jay H. Lawrimore is Chief of the Climate Analysis Branch at NOAA's National Climatic Data Center. He has led a team of scientists that monitors the Earth's climate on an operational basis.

James J. McCarthy is Alexander Agassiz Professor of Biological Oceanography at Harvard University. His areas of expertise are biology and oceanography. He is also President of the American Association for the Advancement of Science.

A. David McGuire is a Professor of Ecology in the U.S. Geological Survey's Alaska Cooperative Fish and Wildlife Research Unit located at the University of Alaska Fairbanks. His areas of expertise are ecosystem ecology and terrestrial feedbacks to the climate system.

Jerry M. Melillo, (Co-Chair), is the Director of The Ecosystems Center at the Marine Biological Laboratory in Woods Hole. He specializes in understanding the impacts of human activities on the biogeochemistry of ecological systems.

Edward L. Miles is the Virginia and Prentice Bloedel Professor of Marine Studies and Public Affairs at the University of Washington. His fields of specialization are international science and technology policy, marine policy and ocean management, and the impacts of climate variability and change.

Evan Mills is currently a Staff Scientist at the U.S. Department of Energy's Lawrence Berkeley National Laboratory. His areas of expertise are energy systems and risk management in the context of climate change.

Jonathan T. Overpeck is a climate system scientist at the University of Arizona, where he is also the Director of the Institute of the Environment, as well as a Professor of Geosciences and a Professor of Atmospheric Sciences.

Federal Advisory Committee Authors

Jonathan A. Patz is a Professor & Director of Global Environmental Health at the University of Wisconsin in Madison. He has earned medical board certification in both Occupational/Environmental Medicine and Family Medicine.

Thomas C. Peterson, (Co-Chair), is a physical scientist at NOAA's National Climatic Data Center in Asheville, North Carolina. His areas of expertise include data fidelity, international data exchange and global climate analysis using both *in situ* and satellite data.

Roger S. Pulwarty is a physical scientist and the Director of the National Integrated Drought Information System Program at NOAA in Boulder, Colorado. His interests are in climate risk assessment and adaptation.

Benjamin D. Santer is an atmospheric scientist at Lawrence Livermore National Laboratory. His research focuses on climate model evaluation, the use of statistical methods in climate science, and identification of "fingerprints" in observed climate records.

Michael J. Savonis has 25 years of experience in transportation policy, with extensive expertise in air quality and emerging environmental issues. He currently serves as a Senior Policy Advisor at the Federal Highway Administration.

H. Gerry Schwartz Jr., is an internationally known expert in environmental and civil engineering. He is past-president of both the Water Environment Federation and the American Society of Civil Engineers, a member of the National Academy of Engineering, and a private consultant.

Eileen L. Shea serves as Director of the NOAA Integrated Data and Environmental Applications Center and Chief of the Climate Monitoring and Services Division, National Climatic Data Center, NOAA/NESDIS. Her educational experience focused on marine science, environmental law, and resource management.

John M.R. Stone is an Adjunct Research Professor in the Department of Geography and Environmental Studies at Carleton University. He has spent the last 20 years managing climate research in Canada and helping to influence the dialogue between science and policy.

Bradley H. Udall is the Director of the University of Colorado Western Water Assessment. He was formerly a consulting engineer at Hydrosphere Resource Consultants. His expertise includes water and policy issues of the American West and especially the Colorado River. He is an affiliate of NOAA's Earth System Research Laboratory.

John E. Walsh is a President's Professor of Global Change at the University of Alaska, Fairbanks and Professor Emeritus of Atmospheric Sciences at the University of Illinois. His research interests include the climate of the Arctic, extreme weather events as they relate to climate, and climate-cryosphere interactions.

Michael F. Wehner is a member of the Scientific Computing Group at the Lawrence Berkeley National Laboratory in Berkeley, California. He has been active in both the design of global climate models and in the analysis of their output.

Thomas J. Wilbanks is a Corporate Research Fellow at the Oak Ridge National Laboratory and leads the Laboratory's Global Change and Developing Country Programs. He conducts research on such issues as sustainable development and responses to concerns about climate change.



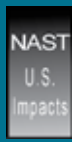

Donald J. Wuebbles is the Harry E. Preble Professor of Atmospheric Sciences at the University of Illinois. His research emphasizes the study of chemical and physical processes of the atmosphere towards improved understanding of the Earth's climate and atmospheric composition.

PRIMARY SOURCES OF INFORMATION

<p>CCSP Goal 1: Improve knowledge of the Earth's past and present climate and environment, including its natural variability, and improve understanding of the causes of observed variability and change.</p>	
<p>CCSP 1.1 Temperature Trends</p>	<p><i>Temperature Trends in the Lower Atmosphere: Steps for Understanding and Reconciling Differences</i></p> <p>Thomas R. Karl, NOAA; Susan J. Hassol, STG Inc.; Christopher D. Miller, NOAA; William L. Murray, STG Inc.</p>
<p>CCSP 1.2 Past Climate</p>	<p><i>Past Climate Variability and Change in the Arctic and at High Latitudes</i></p> <p>Richard B. Alley, Pennsylvania State Univ.; Julie Brigham-Grette, Univ. of Massachusetts; Gifford H. Miller, Univ. of Colorado; Leonid Polyak, Ohio State Univ.; James W.C. White, Univ. of Colorado; Joan J. Fitzpatrick, USGS</p>
<p>CCSP 1.3 Re-Analysis</p>	<p><i>Re-Analyses of Historical Climate Data for Key Atmospheric Features: Implications for Attribution of Causes of Observed Change</i></p> <p>Randall M. Dole, Martin P. Hoerling, Siegfried Schubert, NOAA</p>
<p>CCSP Goal 2: Improve quantification of the forces bringing about changes in the Earth's climate and related systems.</p>	
<p>CCSP 2.1 GHG Emissions</p>	<p>Part A: <i>Scenarios of Greenhouse Gas Emissions and Atmospheric Concentrations</i> Part B: <i>Global-Change Scenarios: Their Development and Use</i></p> <p>Leon E. Clarke, James A. Edmonds, Hugh M. Pitcher, Pacific Northwest National Lab.; Henry D. Jacoby and John M. Reilly, MIT; Richard G. Richels, Electric Power Research Institute; Edward A. Parson, Univ. of Michigan; Virginia R. Burkett, USGS; Karen Fisher-Vanden, Dartmouth College; David W. Keith, Univ. of Calgary; Linda O. Mearns, NCAR; Cynthia E. Rosenzweig, NASA; Mort D. Webster, MIT; John C. Houghton DOE/Office of Biological and Environmental Research</p>
<p>CCSP 2.2 Carbon Cycle</p>	<p><i>The First State of the Carbon Cycle Report (SOCCR)</i> <i>North American Carbon Budget and Implications for the Global Carbon Cycle</i></p> <p>Anthony W. King, ORNL; Lisa Dilling, Univ. of Colorado/NCAR; Gregory P. Zimmerman, ORNL; David Fairman, Consensus Building Institute Inc.; Richard A. Houghton, Woods Hole Research Center; Gregg Marland, ORNL; Adam Z. Rose, Pennsylvania State Univ. and Univ. Southern California; Thomas J. Wilbanks, ORNL</p>
<p>CCSP 2.3 Aerosol Impacts</p>	<p><i>Atmospheric Aerosol Properties and Climate Impacts</i></p> <p>Mian Chin, NASA; Ralph A. Kahn, NASA; Stephen E. Schwartz, DOE/BNL; Lorraine A. Remer, NASA/GSFC; Hogbin Yu, NASA/GSFC/UMBC; David Rind, NASA/GISS; Graham Feingold, NOAA/ESRL; Patricia K. Quinn, NOAA/PMEL; David G. Streets, DOE/ANL; Philip DeCola, NASA HQ; Rangasayi Halthore, NASA HQ/NRL</p>
<p>CCSP 2.4 Ozone Trends</p>	<p><i>Trends in Emissions of Ozone-Depleting Substances, Ozone Layer Recovery, & Implications for Ultraviolet Radiation Exposure</i></p> <p>A.R. Ravishankara, NOAA; Michael J. Kurylo, NASA; Christine Ennis, NOAA/ESRL</p>

CCSP Goal 3: Reduce uncertainty in projections of how the Earth's climate and related systems may change in the future.	
CCSP 3.1 Climate Models	<i>Climate Models: An Assessment of Strengths and Limitations</i> David C. Bader and Curt Covey, Lawrence Livermore National Lab.; William J. Gutowski Jr., Iowa State Univ.; Isaac M. Held, NOAA/GFDL; Kenneth E. Kunkel, Illinois State Water Survey; Ronald L. Miller, NASA/GISS; Robin T. Tokmakian, Naval Postgraduate School; Minghua H. Zhang, State Univ. of New York Stony Brook; Anjuli S. Bamzai, DOE
CCSP 3.2 Climate Projections	<i>Climate Projections Based on Emissions Scenarios for Long-Lived and Short-Lived Radiatively Active Gases and Aerosols</i> Hiram Levy II, NOAA/GFDL; Drew Shindell, NASA/GISS; Alice Gilliland, NOAA /ARL; M. Daniel Schwarzkopf, NOAA/GFDL; Larry W. Horowitz, NOAA/GFDL; Anne M. Waple, STG Inc.
CCSP 3.3 Extremes	<i>Weather and Climate Extremes in a Changing Climate: Regions of Focus: North America, Hawaii, Caribbean, and U.S. Pacific Islands</i> Thomas R. Karl, NOAA; Gerald A. Meehl, NCAR; Christopher D. Miller, NOAA; Susan J. Hassol, STG Inc.; Anne M. Waple, STG Inc.; William L. Murray, STG Inc.
CCSP 3.4 Abrupt Climate Change	<i>Abrupt Climate Change</i> John P. McGeehin, USGS; John A. Barron, USGS; David M. Anderson, NOAA; David J. Verardo, NSF; Peter U. Clark, Oregon State Univ.; Andrew J. Weaver, Univ. of Victoria; Konrad Steffen, Univ. of Colorado; Edward R. Cook, Columbia Univ.; Thomas L. Delworth, NOAA; Edward Brook, Oregon State Univ.
CCSP Goal 4: Understand the sensitivity and adaptability of different natural and managed ecosystems and human systems to climate and related global changes.	
CCSP 4.1 Sea Level Rise	<i>Coastal Sensitivity to Sea-Level Rise: A Focus on the Mid-Atlantic Region</i> James G. Titus, US EPA; K. Eric Anderson, USGS; Donald R. Cahoon, USGS; Dean B. Gesch, USGS; Stephen K. Gill, NOAA; Benjamin T. Gutierrez, USGS; E. Robert Thieler, USGS; S. Jeffress Williams, USGS
CCSP 4.2 Ecosystem Thresholds	<i>Thresholds of Climate Change in Ecosystems</i> Daniel B. Fagre, USGS; Colleen W. Charles, USGS
CCSP 4.3 Impacts	<i>The Effects of Climate Change on Agriculture, Land Resources, Water Resources and Biodiversity in the United States</i> Peter Backlund, NCAR; Anthony Janetos, PNNL/Univ. of Maryland; David Schimel, National Ecological Observatory Network; Margaret Walsh, USDA
CCSP 4.4 Ecosystem Adaptation	<i>Preliminary Review of Adaptation Options for Climate-Sensitive Ecosystems and Resources</i> Jill S. Baron, USGS and Colorado State Univ.; Linda A. Joyce, USDA Forest Service; Brad Griffith, USGS; Peter Kareiva, The Nature Conservancy; Brian D. Keller, NOAA; Margaret Palmer, Univ. of Maryland; Charles Peterson, Univ. of North Carolina; J. Michael Scott, USGS and Univ. of Idaho; Susan Herrod Julius, US EPA; Jordan M. West, US EPA
CCSP 4.5 Energy	<i>Effects of Climate Change on Energy Production and Use in the United States</i> Thomas J. Wilbanks, ORNL; Vatsal Bhatt, Brookhaven National Lab.; Daniel E. Bilello, Stanley R. Bull, National Renewable Energy Lab.; James Ekmann, National Energy Technology Lab.; William C. Horak, Brookhaven National Lab.; Y. Joe Huang, Mark D. Levine, Lawrence Berkeley National Lab.; Michael J. Sale, ORNL; David K. Schmalzer, Argonne National Lab.; Michael J. Scott, Pacific Northwest National Lab.

	<p><i>Analyses of the Effects of Global Change on Human Health and Welfare and Human Systems</i></p>
	<p>Janet L. Gamble, US EPA; Kristie L. Ebi, ESS LLC.; Anne E. Grambsch, US EPA; Frances G. Sussman, Environmental Economics Consulting; Thomas J. Wilbanks, ORNL</p>
<p>CCSP Goal 5: Explore the uses and identify the limits of evolving knowledge to manage risks and opportunities related to climate variability and change.</p>	
	<p><i>Uses and Limitations of Observations, Data, Forecasts, and Other Projections in Decision Support for Selected Sectors and Regions</i></p>
	<p>John Haynes, NASA; Fred Vukovich, SAIC; Molly K. Macauley, RFF; Daewon W. Byun, Univ. of Houston; David Renne, NREL; Gregory Glass, Johns Hopkins School of Public Health; Holly Hartmann, Univ. of Arizona</p>
	<p><i>Best Practice Approaches for Characterizing, Communicating and Incorporating Scientific Uncertainty in Climate Decision Making</i></p>
<p>Other Assessments Referenced</p>	
	<p><i>Decision Support Experiments and Evaluations using Seasonal-to-Interannual Forecasts and Observational Data: A Focus on Water Resources</i></p>
	<p>Working Group I - <i>Climate Change 2007: The Physical Science Basis</i></p>
	<p>Susan Solomon, Dahe Qin, Martin Manning, Zhenlin Chen, Melinda Marquis, Kristen B. Averyt, Melina M.B. Tignor, Henry LeRoy Miller, Jr.</p>
	<p>Working Group II - <i>Climate Change 2007: Impacts, Adaptation and Vulnerability</i></p>
<p>Working Group III - <i>Climate Change 2007: Mitigation of Climate Change</i></p>	
<p><i>Special Report on Emissions Scenarios</i></p>	
<p>Bert Metz, Ogunlade R. Davidson, Peter R. Bosch, Rutu Dave, Leo A. Meyer</p>	
<p>Nebojsa Nakicenovic, Robert Swart</p>	

	<p><i>Climate Change and Water</i></p>
	<p>Bryson Bates, Zbigniew W. Kundzewicz, Shaohong Wu, Jean P. Palutikof</p>
	<p><i>Potential Impacts of Climate Change on U.S. Transportation</i></p>
	<p>Henry G. Schwartz, Jr., Alan C. Clark, G. Edward Dickey, George C. Eads, Robert E. Gallamore, Genevieve Giuliano, William J. Gutowski, Jr., Randell H. Iwasaki, Klaus H. Jacob, Thomas R. Karl, Robert J. Lempert, Luisa M. Paiewonsky, S. George H. Philander, Christopher R. Zeppie</p>
	<p><i>Climate Change Impacts on the United States: The Potential Consequences of Climate Variability and Change</i></p>
	<p>Jerry M. Melillo, Anthony C. Janetos, Thomas R. Karl, Eric J. Barron, Virginia Rose Burkett, Thomas F. Cecich, Robert W. Corell, Katharine L. Jacobs, Linda A. Joyce, Barbara Miller, M. Granger Morgan, Edward A. Parson, Richard G. Richels, David S. Schimel</p>
	<p><i>Impacts of a Warming Arctic, Arctic Climate Impact Assessment</i></p>
	<p>Robert W. Corell, Susan J. Hassol, Pål Prestrud, Patricia A. Anderson, Snorri Baldursson, Elizabeth Bush, Terry V. Callaghan, Paul Grabhorn, Gordon McBean, Michael MacCracken, Lars-Otto Reiersen, Jan Idar Solbakken, Gunter Weller</p>

ACRONYMS

ARS: Agricultural Research Service
 CCSP: Climate Change Science Program
 CIESIN: Center for International Earth Science Information Network
 CIRES: Cooperative Institute for Research in Environmental Sciences
 CMIP: Coupled Model Intercomparison Project
 DOE: Department of Energy
 EIA: Energy Information Administration
 GAO: General Accounting Office
 IARC: International Arctic Research Center
 IPCC: Intergovernmental Panel on Climate Change
 NASA: National Aeronautics and Space Administration
 NASS: National Agricultural Statistics Service
 NAST: National Assessment Synthesis Team
 NCDC: National Climatic Data Center
 NESDIS: National Environmental Satellite, Data, and Information Service

NOAA: National Oceanic and Atmospheric Administration
 NRCS: Natural Resources Conservation Service
 NSIDC: National Snow and Ice Data Center
 NWS: National Weather Service
 NWFS: Northwest Fisheries Science Center
 PISCO: Partnership for Interdisciplinary Studies of Coastal Oceans
 PLJV: Playa Lakes Joint Venture
 SAP: Synthesis and Assessment Product
 SRH: Southern Regional Headquarters
 USACE: United States Army Corps of Engineers
 USBR: United States Bureau of Reclamation
 USDA: United States Department of Agriculture
 USDOE: United States Department of Energy
 USEPA: United States Environmental Protection Agency
 USFS: United States Forest Service
 USGAO: United States Government Accountability Office
 USGS: United States Geological Survey

- ¹ CCSP, 2009: *Best Practice Approaches for Characterizing, Communicating, and Incorporating Scientific Uncertainty in Decisionmaking*. [Morgan, G., H. Dowlatabadi, M. Henrion, D. Keith, R. Lempert, S. McBrid, M. Small, and T. Willbanks (eds.)]. Synthesis and Assessment Product 5.2. National Oceanic and Atmospheric Administration, Washington DC.
- ² Historical data: Lüthi, D., M. Le Floch, B. Bereiter, T. Blunier, J.-M. Barnola, U. Siegenthaler, D. Raynaud, J. Jouzel, H. Fischer, K. Kawamura, and T.F. Stocker, 2008: High-resolution carbon dioxide concentration record 650,000-800,000 years before present. *Nature*, **453(7193)**, 379-382.
1959-2008 data: Tans, P., 2008: *Trends in Atmospheric Carbon Dioxide: Mauna Loa*. NOAA Earth System Research Laboratory (ESRL). [Web site] <<http://www.esrl.noaa.gov/gmd/ccgg/trends/>> Data available at <ftp://ftp.cmdl.noaa.gov/ccg/co2/trends/co2_annmean_mlo.txt>
2100 projected data: International Institute for Applied System Analysis (IIASA) GGI Scenario Database, 2008. <<http://www.iiasa.ac.at/Research/GGI/DB/>>
- ³ Forster, P., V. Ramaswamy, P. Artaxo, T. Berntsen, R. Betts, D.W. Fahey, J. Haywood, J. Lean, D.C. Lowe, G. Myhre, J. Nganga, R. Prinn, G. Raga, M. Schulz, and R. Van Dorland, 2007: Changes in atmospheric constituents and in radiative forcing. In: *Climate Change 2007: The Physical Basis*. Contribution of Working Group I to the Fourth Assessment Report of the Intergovernmental Panel on Climate Change [Solomon, S., D. Qin, M. Manning, Z. Chen, M. Marquis, K.B. Averyt, M. Tignor, and H.L. Miller (eds.)]. Cambridge University Press, Cambridge, UK, and New York, pp. 129-234.
- ⁴ Denman, K.L., G. Brasseur, A. Chidthaisong, P. Ciaus, P.M. Cox, R.E. Dickinson, D. Hauglustaine, C. Heinze, E. Holland, D. Jacob, U. Lohmann, S. Ramachandran, P.L. da Silva Dias, S.C. Wofsy, and X. Zhang, 2007: Couplings between changes in the climate system and biogeochemistry. In: *Climate Change 2007: The Physical Science Basis*. Contribution of Working Group I to the Fourth Assessment Report of the Intergovernmental Panel on Climate Change [Solomon, S., D. Qin, M. Manning, Z. Chen, M. Marquis, K.B. Averyt, M. Tignor and H.L. Miller (eds.)]. Cambridge University Press, Cambridge, UK, and New York, pp. 499-587.
- ⁵ Ko, M., J.S. Daniel, J.R. Herman, P.A. Newman, and V. Ramaswamy, 2008: The future and recovery. In: *Trends in Emissions of Ozone-Depleting Substances, Ozone Layer Recovery, and Implications for Ultraviolet Radiation Exposure*. [Ravishankara, A.R., M.J. Kurylo, and C.A. Ennis (eds.)]. Synthesis and Assessment Product 2.4. NOAA's National Climatic Data Center, Asheville, NC, pp. 133-154.
- ⁶ Ravishankara, A.R., M.J. Kurylo, and A.-M. Schmoltner, 2008: Introduction. In: *Trends in Emissions of Ozone-Depleting Substances, Ozone Layer Recovery, and Implications for Ultraviolet Radiation Exposure*. [Ravishankara, A.R., M.J. Kurylo, and C.A. Ennis (eds.)]. Synthesis and Assessment Product 2.4. NOAA's National Climatic Data Center, Asheville, NC, pp. 23-28.
- ⁷ Blasing, T.J., 2008: *Recent Greenhouse Gas Concentrations*. Carbon Dioxide Information Analysis Center, Oak Ridge National Laboratory. <http://cdiac.ornl.gov/pns/current_ghg.html>
- ⁸ Fahey, D.W. (lead author), 2007: *Twenty Questions and Answers about the Ozone Layer: 2006 Update*. World Meteorological Organization, Geneva, Switzerland, 50 pp. <<http://www.esrl.noaa.gov/csd/assessments/2006/twentyquestions.html>>
- ⁹ Thompson, D.W.J. and S. Solomon, 2002: Interpretation of recent Southern Hemisphere climate change. *Science*, **296(5569)**, 895-899.
- ¹⁰ Kahn, R.A., H. Yu, S.E. Schwartz, M. Chin, G. Feingold, L.A. Remer, D. Rind, R. Halthore, and P. DeCola, 2009: Introduction. In: *Atmospheric Aerosol Properties and Climate Impacts*. [Chin, M., R.A. Kahn, and S.E. Schwartz (eds.)]. Synthesis and Assessment Product 2.3. National Aeronautics and Space Administration, Washington, DC, pp. 9-20.
- ¹¹ Solomon, S., G.-K. Plattner, R. Knutti, and P. Friedlingstein, 2009: Irreversible climate change because of carbon dioxide emissions. *Proceedings of the National Academy of Sciences*, **106(6)**, 1704-1709.
- ¹² Archer, D., 2005: Fate of fossil fuel CO₂ in geologic time. *Journal of Geophysical Research*, **110**, C09S05, doi:10.1029/2004JC002625.
- ¹³ Shindell, D.T., H. Levy II, A. Gilliland, M.D. Schwarzkopf, and L.W. Horowitz, 2008: Climate change from short-lived emissions due to human activities. In: *Climate Projections Based on Emissions Scenarios for Long-Lived and Short-Lived Radiatively Active Gases and Aerosols*. [Levy II, H., D.T. Shindell, A. Gilliland, M.D. Schwarzkopf, and L.W. Horowitz, (eds.)]. Synthesis and Assessment Product 3.2. U.S. Climate Change Science Program, Washington, DC, pp. 27-60.
- ¹⁴ Santer, B.D., J.E. Penner, and P.W. Thorne, 2006: How well can the observed vertical temperature changes be reconciled with our understanding of the causes of these changes? In: *Temperature Trends in the Lower Atmosphere: Steps for Understanding and Reconciling Differences* [Karl, T.R., S.J. Hassol, C.D. Miller, and W.L. Murray (eds.)]. Synthesis and Assessment Product 1.1 U.S. Climate Change Science Program, Washington, DC, pp. 89-118.
- ¹⁵ Hansen, J., M. Sato, R. Ruedy, L. Nazarenko, A. Lacis, G.A. Schmidt, G. Russell, I. Aleinov, M. Bauer, S. Bauer, N. Bell, B. Cairns, V. Canuto, M. Chandler, Y. Cheng, A. Del Genio, G. Faluveg E. Fleming, A. Friend, T. Hall, C. Jackman, M. Kelley, N. Kiang, D. Koch, J. Lean, J. Lerner, K. Lo, S. Menon, R. Miller, P. Minnis, T. Novakov, V. Oinas, Ja. Perlwitz, Ju. Perlwitz, D. Rind, A. Romanou, D. Shindell, P. Stone, S. Sun, N. Tausnev, D. Thresher, B. Wielicki, T. Wong, M. Yao, and S. Zhang 2005: Efficacy of climate forcings. *Journal of Geophysical Research*, **110**, D18104, doi:10.1029/2005JD005776.
- ¹⁶ National Research Council, 2005: *Radiative Forcing of Climate Change: Expanding the Concept and Addressing Uncertainties*. National Academies Press, Washington DC, 207 pp.
- ¹⁷ Hansen, J., M. Sato, R. Ruedy, A. Lacis, and V. Oinas, 2000: Global warming in the twenty-first century: an alternative scenario. *Proceedings of the National Academy of Sciences*, **97(18)**, 9875-9880.
- ¹⁸ Field, C.B., J. Sarmiento, and B. Hales, 2007: The carbon cycle of North America in a global context. In: *The First State of the Carbon Cycle Report (SOCCR): The North American Carbon Budget and Implications for the Global Carbon Cycle* [King, A.W., L. Dilling, G.P. Zimmerman, D.M. Fairman, R.A. Houghton, G. Marland, A.Z. Rose, and T.J.

- Wilbanks (eds.]. Synthesis and Assessment Product 2.2. NOAA's National Climatic Data Center, Asheville, NC, pp. 21-28.
- ¹⁹ Tarnocai, C., C.-L. Ping, and J. Kimble, 2007: Carbon cycles in the permafrost region of North America. In: *The First State of the Carbon Cycle Report (SOCCR): The North American Carbon Budget and Implications for the Global Carbon Cycle* [King, A.W., L. Dilling, G.P. Zimmerman, D.M. Fairman, R.A. Houghton, G. Marland, A.Z. Rose, and T.J. Wilbanks (eds.]. Synthesis and Assessment Product 2.2. NOAA's National Climatic Data Center, Asheville, NC, pp. 127-138.
- ²⁰ Jansen, E., J. Overpeck, K.R. Briffa, J.-C. Duplessy, F. Joos, V. Masson-Delmotte, D. Olago, B. Otto-Bliesner, W.R. Peltier, S. Rahmstorf, R. Ramesh, D. Raynaud, D. Rind, O. Solomina, R. Villalba and D. Zhang, 2007: Palaeoclimate. In: *Climate Change 2007: The Physical Science Basis*. Contribution of Working Group I to the Fourth Assessment Report of the Intergovernmental Panel on Climate Change [Solomon, S., D. Qin, M. Manning, Z. Chen, M. Marquis, K.B. Averyt, M. Tignor, and H.L. Miller (eds.]. Cambridge University Press, Cambridge, UK, and New York, pp. 433-497.
- ²¹ Canadell, J.G., C. Le Quééré, M.R. Raupach, C.B. Field, E.T. Buitenhuis, P. Ciais, T.J. Conway, N.P. Gillett, R.A. Houghton, and G. Marland, 2007: Contributions to accelerating atmospheric CO₂ growth from economic activity, carbon intensity, and efficiency of natural sinks. *Proceedings of the National Academy of Sciences*, **104(47)**, 18866-18870.
- ²² Royal Society, 2005: *Ocean Acidification Due to Increasing Atmospheric Carbon Dioxide*. Policy Document 12/05. Royal Society, London, 60 pp.
- ²³ Orr, J.C., V.J. Fabry, O. Aumont, L. Bopp, S.C. Doney, R.A. Feely, A. Gnanadesikan, N. Gruber, A. Ishida, F. Joos, R.M. Key, K. Lindsay, E. Maier-Reimer, R. Matear, P. Monfray, A. Mouchet, R.G. Najjar, G.-K. Plattner, K.B. Rodgers, C.L. Sabine, J.L. Sarmiento, R. Schlitzer, R.D. Slater, I.J. Totterdell, M.-F. Weirig, Y. Yamanaka, and A. Yool, 2005: Anthropogenic ocean acidification over the twenty-first century and its impact on calcifying organisms. *Nature*, **437(7059)**, 681-686.
- ²⁴ Allen, M.R., 2003: Liability for climate change. *Nature*, **421(6926)**, 891-892.
- ²⁵ Clarke, L., J. Edmonds, H. Jacoby, H. Pitcher, J. Reilly, and R. Richels, 2007: *Scenarios of Greenhouse Gas Emissions and Atmospheric Concentrations*. Sub-report 2.1A of Synthesis and Assessment Product 2.1. Department of Energy, Office of Biological & Environmental Research, Washington, DC, 154 pp.
- ²⁶ The spatial average of annual-average surface air temperatures around the globe is commonly referred to as the global average surface air temperature.
- ²⁷ Meier, M.F., M.B. Dyurgerov, U.K. Rick, S. O'Neel, W.T. Pfeffer, R.S. Anderson, S.P. Anderson, and A.F. Glazovsky, 2007: Glaciers dominate eustatic sea-level rise in the 21st century. *Science*, **317(5841)**, 1064-1067.
- ²⁸ Trenberth, K.E., P.D. Jones, P. Ambenje, R. Bojariu, D. Easterling, A. Klein Tank, D. Parker, F. Rahimzadeh, J.A. Renwick, M. Rusticucci, B. Soden, and P. Zhai, 2007: Observations: surface and atmospheric climate change. In: *Climate Change 2007: The Physical Basis*. Contribution of Working Group I to the Fourth Assessment Report of the Intergovernmental Panel on Climate Change [Solomon, S., D. Qin, M. Manning, Z. Chen, M. Marquis, K.B. Averyt, M. Tignor, and H.L. Miller (eds.]. Cambridge University Press, Cambridge, UK, and New York, pp. 235-335.
- ²⁹ Steffen, K., P.U. Clark, J.G. Cogley, D. Holland, S. Marshall, E. Rignot, and R. Thomas, 2008: Rapid changes in glaciers and ice sheets and their impacts on sea level. In: *Abrupt Climate Change*. Synthesis and Assessment Product 3.4. U.S. Geological Survey, Reston, VA, pp. 60-142.
- ³⁰ Lanzante, J.R., T.C. Peterson, F.J. Wentz, and K.Y. Vinnikov, 2006: What do observations indicate about the change of temperatures in the atmosphere and at the surface since the advent of measuring temperatures vertically? In: *Temperature Trends in the Lower Atmosphere: Steps for Understanding and Reconciling Differences* [Karl, T.R., S.J. Hassol, C.D. Miller, and W.L. Murray (eds.]. Synthesis and Assessment Product 1.1 U.S. Climate Change Science Program, Washington, DC, pp. 47-70.
- ³¹ Santer, B.D., P.W. Thorne, L. Haimberger, K.E. Taylor, T.M.L. Wigley, J.R. Lanzante, S. Solomon, M. Free, P.J. Gleckler, P.D. Jones, T.R. Karl, S.A. Klein, C. Mears, D. Nychka, G.A. Schmidt, S.C. Sherwood, and F.J. Wentz, 2008: Consistency of modelled and observed temperature trends in the tropical troposphere. *International Journal of Climatology*, **28(13)**, 1703-1722.
- ³² Uncertainties in the data are an order of magnitude smaller than the trend according to Karl, T.R., J.R. Christy, R.A. Clarke, G.V. Gruza, J. Jouzel, M.E. Mann, J. Oerlemans, M.J. Salinger, and S.-W. Wang, 2001: Observed climate variability and change. In: *Climate Change 2001: The Scientific Basis*. Contribution of Working Group I to the Third Assessment Report of the Intergovernmental Panel on Climate Change [Houghton, J.T., Y. Ding, D.J. Griggs, M. Noguer, P.J. van der Linden, X. Dai, K. Maskell, and C.A. Johnson (eds.]. Cambridge University Press, Cambridge, UK, and New York, pp. 99-181.
- Temperature data:*
Smith, T.M. and R.W. Reynolds, 2004: Improved extended reconstruction of SST (1854–1997). *Journal of Climate*, **17(12)**, 2466-2477.
Jones, P.D., M. New, D.E. Parker, S. Martin, and I.G. Rigor, 1999: Surface air temperature and its changes over the past 150 years. *Reviews of Geophysics*, **37(2)**, 173-199.
- Carbon dioxide data:*
Data from 1974 to present: Tans, P., 2008: *Trends in Atmospheric Carbon Dioxide: Mauna Loa*. NOAA Earth System Research Laboratory (ESRL). [Web site] <<http://www.esrl.noaa.gov/gmd/ccgg/trends/>> Data available at <ftp://ftp.cmdl.noaa.gov/ccg/co2/trends/co2_annmean_mlo.txt>
1958-1974 data are from the Scripps Institution of Oceanography (Keeling) Mauna Loa Observatory record. <<http://scrippsco2.ucsd.edu/>>
Pre-1958 values are annual points taken from a smooth fit to the Law Dome data: Etheridge, D.M., L.P. Steele, R.L. Langenfelds, R.J. Francey, J.-M. Barnola, and V.I. Morgan, 1996: Natural and anthropogenic changes in atmospheric CO₂ over the last 1000 years from air in Antarctic ice and firn. *Journal of Geophysical Research*, **101(D2)**, 4115-4128.
- ³³ Easterling, D. and M. Wehner, 2009: Is the climate warming or cooling? *Geophysical Research Letters*, in press, 30 March 2009, doi:10.1029/2009GL037810.
- ³⁴ Barnett, T.P., D.W. Pierce, H.G. Hidalgo, C. Bonfils, B.D. Santer, T. Das, G. Bala, A.W. Wood, T. Nozawa, A.A. Mirin, D.R. Cayan, and M.D. Dettinger, 2008: Human-induced changes in the hydrology of the western United States. *Science*, **319(5866)**, 1080-1083.
- ³⁵ Willett, K.M., N.P. Gillett, P.D. Jones, and P.W. Thorne, 2007: Attribution of observed surface humidity changes to human influence. *Nature*, **449(7163)**, 710-712.
- ³⁶ Santer, B.D., C. Mears, F.J. Wentz, K.E. Taylor, P.J. Gleckler, T.M.L. Wigley, T.P. Barnett, J.S. Boyle, W. Brüggemann, N.P. Gillett, S.A. Klein, G.A. Meehl, T. Nozawa, D.W. Pierce, P.A. Stott, W.M. Washington, and M.F. Wehner, 2007: Identification of human-induced changes in atmospheric moisture content. *Proceedings of the National Academy of Sciences*, **104(39)**, 15248-15253.

- ³⁷ Bindoff, N.L., J. Willebrand, V. Artale, A. Cazenave, J. Gregory, S. Gulev, K. Hanawa, C. Le Quéré, S. Levitus, Y. Nojiri, C.K. Shum, L.D. Talley, and A. Unnikrishnan, 2007: Observations: oceanic climate change and sea level. In: *Climate Change 2007: The Physical Science Basis*. Contribution of Working Group I to the Fourth Assessment Report of the Intergovernmental Panel on Climate Change [Solomon, S., D. Qin, M. Manning, Z. Chen, M. Marquis, K.B. Averyt, M. Tignor, and H.L. Miller (eds.)]. Cambridge University Press, Cambridge, UK, and New York, pp. 385-432.
- ³⁸ Barnett, T.P., D.W. Pierce, K.M. AchutaRao, P.J. Gleckler, B.D. Santer, J.M. Gregory, and W.M. Washington, 2005: Penetration of human-induced warming into the world's oceans. *Science*, **309(5732)**, 284-287.
- ³⁹ Pierce, D.W., T.P. Barnett, K.M. AchutaRao, P.J. Gleckler, J.M. Gregory, and W.M. Washington, 2006: Anthropogenic warming of the oceans: observations and model results. *Journal of Climate*, **19(10)**, 1873-1900.
- ⁴⁰ Lemke, P., J. Ren, R.B. Alley, I. Allison, J. Carrasco, G. Flato, Y. Fujii, G. Kaser, P. Mote, R.H. Thomas, and T. Zhang, 2007: Observations: changes in snow, ice and frozen ground. In: *Climate Change 2007: The Physical Basis*. Contribution of Working Group I to the Fourth Assessment Report of the Intergovernmental Panel on Climate Change [Solomon, S., D. Qin, M. Manning, Z. Chen, M. Marquis, K.B. Averyt, M. Tignor, and H.L. Miller (eds.)]. Cambridge University Press, Cambridge, UK, and New York, pp. 337-383.
- ⁴¹ Luthcke, S.B., H.J. Zwally, W. Abdalati, D.D. Rowlands, R.D. Ray, R.S. Nerem, F.G. Lemoine, J.J. McCarthy, and D.S. Chinn, 2006: Recent Greenland ice mass loss by drainage system from satellite gravity observations. *Science*, **314(5803)**, 1286-1289.
- ⁴² Pfeffer, W.T., J.T. Harper, and S. O'Neel, 2008: Kinematic constraints on glacier contributions to 21st-century sea-level rise. *Science*, **321(5894)**, 1340-1343.
- ⁴³ Williams, S.J., B.T. Gutierrez, J.G. Titus, S.K. Gill, D.R. Cahoon, E.R. Thieler, K.E. Anderson, D. FitzGerald, V. Burkett, and J. Samenow, 2009: Sea-level rise and its effects on the coast. In: *Coastal Elevations and Sensitivity to Sea-level Rise: A Focus on the Mid-Atlantic Region* [J.G. Titus (coordinating lead author), K.E. Anderson, D.R. Cahoon, D.B. Gesch, S.K. Gill, B.T. Gutierrez, E.R. Thieler, and S.J. Williams (lead authors)]. Synthesis and Assessment Product 4.1. U.S. Environmental Protection Agency, Washington, DC, pp. 11-24.
- ⁴⁴ IPCC (Intergovernmental Panel on Climate Change), 1996: Summary for policy makers. In: *Climate Change 1995: The Science of Climate Change*. Contribution of Working Group I to the Second Assessment Report of the Intergovernmental Panel on Climate Change [Houghton, J.T., L.G. Meiro Filho, B.A. Callander, N. Harris, A. Kattenberg, and K. Maskell (eds.)]. Cambridge University Press, Cambridge, UK, and New York, pp. 1-7.
- ⁴⁵ IPCC (Intergovernmental Panel on Climate Change), 2007: Summary for policymakers. In: *Climate Change 2007: The Physical Science Basis*. Contribution of Working Group I to the Fourth Assessment Report of the Intergovernmental Panel on Climate Change [Solomon, S., D. Qin, M. Manning, Z. Chen, M. Marquis, K.B. Averyt, M. Tignor, and H.L. Miller (eds.)]. Cambridge University Press, Cambridge, UK, and New York, pp. 1-18.
- ⁴⁶ Wigley, T.M.L., V. Ramaswamy, J.R. Christy, J.R. Lanzante, C.A. Mears, B.D. Santer, and C.K. Folland, 2006: Executive summary. In: *Temperature Trends in the Lower Atmosphere: Steps for Understanding and Reconciling Differences* [Karl, T.R., S.J. Hassol, C.D. Miller, and W.L. Murray (eds.)]. Synthesis and Assessment Product 1.1 U.S. Climate Change Science Program, Washington, DC, pp. 1-15.
- ⁴⁷ National Research Council, 2006: *Surface Temperature Reconstructions for the Last 2,000 Years*. National Academies Press, Washington DC, 196 pp.
- ⁴⁸ Mann, M.E., Z. Zhang, M.K. Hughes, R.S. Bradley, S.K. Miller, S. Rutherford, and F. Ni, 2008: Proxy-based reconstructions of hemispheric and global surface temperature variations over the past two millennia. *Proceedings of the National Academy of Sciences*, **105(36)**, 13252-13257.
- ⁴⁹ Hegerl, G.C., F.W. Zwiers, P. Braconnot, N.P. Gillett, Y. Luo, J.A. Marengo Orsini, N. Nicholls, J.E. Penner, and P.A. Stott, 2007: Understanding and attributing climate change. In: *Climate Change 2007: The Physical Science Basis*. Contribution of Working Group I to the Fourth Assessment Report of the Intergovernmental Panel on Climate Change [Solomon, S., D. Qin, M. Manning, Z. Chen, M. Marquis, K.B. Averyt, M. Tignor, and H.L. Miller (eds.)]. Cambridge University Press, Cambridge, UK, and New York, pp. 663-745.
- ⁵⁰ LeTreut, H., R. Somerville, U. Cubasch, Y. Ding, C. Mauritzen, A. Mokssit, T. Peterson, and M. Prather, 2007: Historical overview of climate changes science. In: *Climate Change 2007: The Physical Basis*. Contribution of Working Group I to the Fourth Assessment Report of the Intergovernmental Panel on Climate Change [Solomon, S., D. Qin, M. Manning, Z. Chen, M. Marquis, K.B. Averyt, M. Tignor, and H.L. Miller (eds.)]. Cambridge University Press, Cambridge, UK, and New York, pp. 93-127.
- ⁵¹ Santer, B.D., T.M.L. Wigley, T.P. Barnett, and E. Anyamba, 1996: Detection of climate change, and attribution of causes. In: *Climate Change 1995: The Science of Climate Change*. Contribution of Working Group I to the Second Assessment Report of the Intergovernmental Panel on Climate Change [Houghton, J.T., L.G. Meiro Filho, B.A. Callander, N. Harris, A. Kattenberg, and K. Maskell (eds.)]. Cambridge University Press, Cambridge, UK, and New York, pp. 407-443.
- ⁵² Mitchell, J.F.B., D.J. Karoly, G.C. Hegerl, F.W. Zwiers, M.R. Allen, and J. Marengo, 2001: Detection of climate change and attribution of causes. In: *Climate Change 2001: The Scientific Basis*. Contribution of Working Group I to the Third Assessment Report of the Intergovernmental Panel on Climate Change [Houghton, J.T., Y. Ding, D.J. Griggs, M. Noguer, P.J. van der Linden, X. Dai, K. Maskell, and C.A. Johnson (eds.)]. Cambridge University Press, Cambridge, UK, and New York, pp. 695-738.
- ⁵³ Santer, B.D., M.F. Wehner, T.M.L. Wigley, R. Sausen, G.A. Meehl, K.E. Taylor, C. Ammann, J. Arblaster, W.M. Washington, J.S. Boyle, and W. Brüggemann, 2003: Contributions of anthropogenic and natural forcing to recent tropopause height changes. *Science*, **301(5632)**, 479-483.
- ⁵⁴ Zhang, X., F.W. Zwiers, G.C. Hegerl, F.H. Lambert, N.P. Gillett, S. Solomon, P.A. Stott and T. Nozawa, 2007: Detection of human influence on twentieth-century precipitation trends. *Nature*, **448(7152)**, 461-465.
- ⁵⁵ Burke, E.J., S.J. Brown, and N. Christidis, 2006: Modeling the recent evolution of global drought and projections for the twenty-first century with the Hadley Centre climate model. *Journal of Hydrometeorology*, **7(5)**, 1113-1125.
- ⁵⁶ Gillett, N.P., F.W. Zwiers, A.J. Weaver, and P.A. Stott, 2003: Detection of human influence on sea level pressure. *Nature*, **422(6929)**, 292-294.
- ⁵⁶ Gedney, N., P.M. Cox, R.A. Betts, O. Boucher, C. Huntingford, and P.A. Stott, 2006: Detection of a direct carbon dioxide effect in continental river runoff records. *Nature*, **439(7078)**, 835-838.
- ⁵⁷ Dole, R. and M. Hoerling, 2008: Introduction. In: *Reanalysis of Historical Climate Data for Key Atmospheric Features: Implications for Attribution of Causes of Observed Change*. [Dole, R., M. Hoerling, and S. Schubert (eds.)]. Synthesis and Assessment Prod-

- uct 1.3. NOAA's National Climatic Data Center, Asheville, NC, pp. 5-10.
- ⁵⁹ The temperature data for the globe is the standard NOAA/NCDC temperature product.
The solar data are a composite of 3 different data sets:
Fröhlich, C. and J. Lean, 2004: Solar radiative output and its variability: evidence and mechanisms. *Astronomy and Astrophysics Review*, **12(4)**, 273-320.
Willson, R.C. and A.V. Mordvinov, 2003: Secular total solar irradiance trend during solar cycles 21023. *Geophysical Research Letters*, **30(5)**, 1199, doi:10.1029/2002GL016038.
Dewitte, S., D. Crommelynck, S. Mekaoui, and A. Jouko, 2004: Measurement and uncertainty of the long-term total solar irradiance trend. *Solar Physics*, **224(1-2)**, 209-216.
- ⁶⁰ Lean, J.L. and D.H. Rind, 2008: How natural and anthropogenic influences alter global and regional surface temperatures: 1889 to 2006. *Geophysical Research Letters*, **35**, L18701, doi:10.1029/2008GL034864.
- ⁶¹ Min, S.-K., X. Zhang, F.W. Zwiers, and T. Agnew, 2008: Human influence on Arctic sea ice detectable from early 1990s onwards. *Geophysical Research Letters*, **35**, L21701, doi:10.1029/2008GL035725.
- ⁶² Gillett, N.P., D.A. Stone, P.A. Stott, T. Nozawa, A.Y. Karpechko, G.C. Hegerl, M.F. Wehner, and P.D. Jones, 2008: Attribution of polar warming to human influence. *Nature Geoscience*, **1(11)**, 750-754.
- ⁶³ Ramaswamy, V., J.W. Hurrell, G.A. Meehl, A. Phillips, B.D. Santer, M.D. Schwarzkopf, D.J. Seidel, S.C. Sherwood, and P.W. Thorne, 2006: Why do temperatures vary vertically (from the surface to the stratosphere) and what do we understand about why they might vary and change over time? In: *Temperature Trends in the Lower Atmosphere: Steps for Understanding and Reconciling Differences* [Karl, T.R., S.J. Hassol, C.D. Miller, and W.L. Murray (eds.)]. Synthesis and Assessment Product 1.1. U.S. Climate Change Science Program, Washington, DC, pp. 15-28.
- ⁶⁴ Stott, P.A., 2003: Attribution of regional-scale temperature changes to anthropogenic and natural causes. *Geophysical Research Letters*, **30(14)**, 1724, doi:10.1029/2003GL017324.
- ⁶⁵ Zwiers, F.W. and X. Zhang, 2003: Towards regional-scale climate change detection. *Journal of Climate*, **16(5)**, 793-797.
- ⁶⁶ Santer, B.D., T.M.L. Wigley, P.J. Glecker, C. Bonfils, M.F. Wehner, K. AchutaRao, T.P. Barnett, J.S. Boyle, W. Brüggemann, M. Fiorino, N.P. Gillett, J.E. Hansen, P.D. Jones, S.A. Klein, G.A. Meehl, S.C.B. Raper, R.W. Reynolds, K.E. Taylor, and W.M. Washington, 2006: Forced and unforced ocean temperature changes in Atlantic and Pacific tropical cyclogenesis regions. *Proceedings of the National Academy of Sciences*, **103(38)**, 13905-13910.
- ⁶⁷ Gillett, N.P., P.A. Stott, and B.D. Santer, 2008: Attribution of cyclogenesis region sea surface temperature change to anthropogenic influence. *Geophysical Research Letters*, **35**, L09707, doi:10.1029/2008GL033670.
- ⁶⁸ Gutowski, W.J., G.C. Hegerl, G.J. Holland, T.R. Knutson, L.O. Mearns, R.J. Stouffer, P.J. Webster, M.F. Wehner, and F.W. Zwiers, 2008: Causes of observed changes in extremes and projections of future changes. In: *Weather and Climate Extremes in a Changing Climate: Regions of Focus: North America, Hawaii, Caribbean, and U.S. Pacific Islands* [Karl, T.R., G.A. Meehl, C.D. Miller, S.J. Hassol, A.M. Waple, and W.L. Murray (eds.)]. Synthesis and Assessment Product 3.3. U.S. Climate Change Science Program, Washington, DC, pp. 81-116.
- ⁶⁹ Root, T.L., D.P. MacMynowski, M.D. Mastrandrea, and S.H. Schneider, 2005: Human-modified temperatures induce species changes: joint attribution. *Proceedings of the National Academy of Sciences*, **102(21)**, 7465-7469.
- ⁷⁰ Janetos, A., L. Hansen, D. Inouye, B.P. Kelly, L. Meyerson, B. Peterson, and R. Shaw, 2008: Biodiversity. In: *The Effects of Climate Change on Agriculture, Land Resources, Water Resources, and Biodiversity in the United States* [Backlund, P., A. Janetos, D. Schimel, J. Hatfield, K. Boote, P. Fay, L. Hahn, C. Izaurralde, B.A. Kimball, T. Mader, J. Morgan, D. Ort, W. Polley, A. Thomson, D. Wolfe, M.G. Ryan, S.R. Archer, R. Birdsey, C. Dahm, L. Heath, J. Hicke, D. Hollinger, T. Huxman, G. Okin, R. Oren, J. Randerson, W. Schlesinger, D. Lettenmaier, D. Major, L. Poff, S. Running, L. Hansen, D. Inouye, B.P. Kelly, L. Meyerson, B. Peterson, and R. Shaw (eds.)]. Synthesis and Assessment Product 4.3. U.S. Department of Agriculture, Washington, DC, pp. 151-181.
- ⁷¹ CCSP, 2006: *Temperature Trends in the Lower Atmosphere: Steps for Understanding and Reconciling Differences* [Karl, T.R., S.J. Hassol, C.D. Miller, and W.L. Murray (eds.)]. Synthesis and Assessment Product 1.1. U.S. Climate Change Science Program, Washington, DC, 164 pp.
- ⁷² Smith, T.M., R.W. Reynolds, T.C. Peterson, and J. Lawrimore, 2008: Improvements to NOAA's historical merged land-ocean surface temperature analysis (1880-2006). *Journal of Climate*, **21(10)**, 2283-2296.
- ⁷³ Haimberger, L., C. Tavolato, and S. Sperka, 2008: Toward elimination of the warm bias in historic radiosonde temperature records – some new results from a comprehensive intercomparison of upper air data. *Journal of Climate*, **21(18)**, 4587-4606.
- ⁷⁴ Sherwood, S.C., C.L. Meyer, R.J. Allen, and H.A. Titchner, 2008: Robust tropospheric warming revealed by iteratively homogenized radiosonde data. *Journal of Climate*, **21(20)**, 5336-5352.
- ⁷⁵ Titchner, H.A., P.W. Thorne, M.P. McCarthy, S.F.B. Tett, L. Haimberger, and D.E. Parker, 2008: Critically reassessing tropospheric temperature trends from radiosondes using realistic validation experiments. *Journal of Climate*, **22(3)**, 465-485.
- ⁷⁶ Delworth, T.L., P.U. Clark, M. Holland, W.E. Johns, T. Kuhlbrodt, J. Lynch-Stieglitz, C. Morrill, R. Seager, A.J. Weaver, and R. Zhang, 2008: The potential for abrupt change in the Atlantic Meridional Overturning Circulation. In: *Abrupt Climate Change*. Synthesis and Assessment Product 3.4. U.S. Geological Survey, Reston, VA, pp. 258-359.
- ⁷⁷ Wentz, F.J., L. Ricciardulli, K. Hilburn, and C. Mears, 2007: How much more rain will global warming bring? *Science*, **317(5835)**, 233-235.
- ⁷⁸ Stott, P.A., D.A. Stone, and M.R. Allen, 2004: Human contribution to the European heatwave of 2003. *Nature*, **432(7017)**, 610-614.
- ⁷⁹ CCSP, 2008: Introduction. In: *Climate Models: An Assessment of Strengths and Limitations* [Bader, D.C., C. Covey, W.J. Gutowski Jr., I.M. Held, K.E. Kunkel, R.L. Miller, R.T. Tokmakian, and M.H. Zhang (authors)]. Synthesis and Assessment Product 3.1 Department of Energy, Office of Biological and Environmental Research, Washington, DC, pp. 7-12.
- ⁸⁰ Randall, D.A., R.A. Wood, S. Bony, R. Coman, T. Fichetef, J. Fyfe, V. Kattsov, A. Pitman, J. Shukla, J. Srinivasan, R.J. Stouffer, A. Sumi, and K.E. Taylor, 2007: Climate models and their evaluation. In: *Climate Change 2007: The Physical Basis*. Contribution of Working Group I to the Fourth Assessment Report of the Intergovernmental Panel on Climate Change [Solomon, S., D. Qin, M. Manning, Z. Chen, M. Marquis, K.B. Averyt, M. Tignor, and H.L. Miller (eds.)]. Cambridge University Press, Cambridge, UK, and New York, pp. 589-662.
- ⁸¹ Nakićenović, N. and R. Swart (eds.), 2000: *Special Report on Emissions Scenarios*. A special report of Working Group III of the Intergovernmental Panel on Climate Change. Cambridge University Press, Cambridge, UK, and New York, 599 pp. <<http://www.ipcc.ch/ipccreports/sres/emission/index.htm>>

- ⁸² Raupach, M.R., G. Marland, P. Ciais, C. Le Quéré, J.G. Canadell, G. Klepper, and C.B. Field, 2007: Global and regional drivers of accelerating CO₂ emissions. *Proceedings of the National Academy of Sciences*, **104(24)**, 10288-10293.
- ⁸³ O'Neill, B.C. and M. Oppenheimer, 2004: Climate change impacts are sensitive to the concentration stabilization path. *Proceedings of the National Academy of Sciences*, **101(47)**, 16411-16416.
- ⁸⁴ Schneider, S.H. and M.D. Mastrandrea, 2005: Probabilistic assessment of "dangerous" climate change and emissions pathways. *Proceedings of the National Academy of Sciences*, **102(44)**, 15728-15735.
- ⁸⁵ Lenton, T.M., H. Held, E. Kriegler, J.W. Hall, W. Lucht, S. Rahmstorf, and H.J. Schellnhuber, 2008: Tipping elements in the Earth's climate system. *Proceedings of the National Academy of Sciences*, **105(6)**, 1786-1793.
- ⁸⁶ Hansen, J., M. Sato, R. Ruedy, P. Kharecha, A. Lacis, R. Miller, L. Nazarenko, K. Lo, G.A. Schmidt, G. Russell, I. Aleinov, S. Bauer, E. Baum, B. Cairns, V. Canuto, M. Chandler, Y. Cheng, A. Cohen, A. Del Genio, G. Faluvegi, E. Fleming, A. Friend, T. Hall, C. Jackman, J. Jonas, M. Kelley, N.Y. Kiang, D. Koch, G. Labow, J. Lerner, S. Menon, T. Novakov, V. Oinas, Ja. Perlwitz, Ju. Perlwitz, D. Rind, A. Romanou, R. Schmunk, D. Shindell, P. Stone, S. Sun, D. Streets, N. Tausnev, D. Thresher, N. Unger, M. Yao, and S. Zhang, 2007: Dangerous human-made interference with climate: a GISS modelE study. *Atmospheric Chemistry and Physics*, **7(9)**, 2287-2312.
- ⁸⁷ Ramanathan, V. and Y. Feng, 2008: On avoiding dangerous anthropogenic interference with the climate system: formidable challenges ahead. *Proceedings of the National Academy of Sciences*, **105(38)**, 14245-14250.
- ⁸⁸ Meinshausen, M., 2006: What does a 2°C target mean for greenhouse gas concentrations? - A brief analysis based on multi-gas emission pathways and several climate sensitivity uncertainty estimates. In: *Avoiding Dangerous Climate Change* [Schellnhuber, J.S., W. Cramer, N. Nakićenović, T.M.L. Wigley, and G. Yohe (eds.)]. Cambridge University Press, Cambridge, UK, and New York, pp. 265-280.
- ⁸⁹ Meinshausen, M., B. Hare, T.M.L. Wigley, D. van Vuuren, M.G.J. den Elzen, and R. Swart, 2006: Multi-gas emission pathways to meet climate targets. *Climatic Change*, **75(1)**, 151-194.
- ⁹⁰ Meehl, G.A., T.F. Stocker, W.D. Collins, P. Friedlingstein, A.T. Gaye, J.M. Gregory, A. Kitoh, R. Knutti, J.M. Murphy, A. Noda, S.C.B. Raper, I.G. Watterson, A.J. Weaver, and Z.-C. Zhao, 2007: Global climate projections. In: *Climate Change 2007: The Physical Basis*. Contribution of Working Group I to the Fourth Assessment Report of the Intergovernmental Panel on Climate Change [Solomon, S., D. Qin, M. Manning, Z. Chen, M. Marquis, K.B. Averyt, M. Tignor, and H.L. Miller (eds.)]. Cambridge University Press, Cambridge, UK, and New York, pp. 747-845.
- ⁹¹ Refer to the description of the emissions scenarios in the *Global Climate Change* section on pages 22-25. "Lower emissions scenario" refers to IPCC SRES B1, "higher emissions scenario" refers to A2 and "even higher emissions scenario" refers to A1FI.
- ⁹² IPCC Emissions Scenarios (Even Higher, Higher Emission Scenario, Lower Emission Scenario): Nakićenović, N. and R. Swart (eds.), 2000: Appendix VII: Data tables. In: *Special Report on Emissions Scenarios*. A special report of Working Group III of the Intergovernmental Panel on Climate Change. Cambridge University Press, Cambridge, UK, and New York. <http://www.grida.no/publications/other/ipcc_sr/?src=/climate/ipcc/emission/> Emission trajectories are spline fits as per Raupach, M.R., G. Marland, P. Ciais, C. Le Quéré, J.G. Canadell, G. Klepper, and C.B. Field, 2007: Global and regional drivers of accelerating CO₂ emissions. *Proceedings of the National Academy of Sciences*, **104(24)**, 10288-10293.
- Stabilization scenarios (550 ppm, 450 ppm): CCSP 2.1a Scenario Information 070707 data file. From: Clarke, L., J. Edmonds, H. Jacoby, H. Pitcher, J. Reilly, and R. Richels, 2007: *Scenarios of Greenhouse Gas Emissions and Atmospheric Concentrations*. Sub-report 2.1A of Synthesis and Assessment Product 2.1. Department of Energy, Office of Biological & Environmental Research, Washington, DC. The emissions and concentrations shown were from MINICAM 1 and 2. See CCSP 2.1A Executive summary for more information. Spread sheet available at <<http://www.climate-science.gov/Library/sap/sap2-1/finalreport/default.htm>>
- Observations of CO₂ emissions (Fossil Fuel CO₂ Emissions graphic) are updates to: Marland, G., B. Andres, T. Boden, 2008: *Global CO₂ Emissions from Fossil-Fuel Burning, Cement Manufacture, and Gas Flaring: 1751-2005*. Carbon Dioxide Information Analysis Center, Oak Ridge National Laboratory, Oak Ridge, TN. <http://cdiac.ornl.gov/ftp/ndp030/global.1751_2005.ems>
- Observations of CO₂ concentrations (Atmospheric CO₂ Concentrations graphic): Tans, P., 2008: *Trends in Atmospheric Carbon Dioxide: Mauna Loa*. NOAA Earth System Research Laboratory (ESRL). [Web site] <<http://www.esrl.noaa.gov/gmd/ccgg/trends/>> Data available at <ftp://ftp.cmdl.noaa.gov/ccg/co2/trends/co2_annmean_mlo.txt>
- ⁹³ CMIP3-A: This analysis uses 15 models simulations from the WCRP CMIP3 that were available at resolutions finer than 4 degrees (CCSM3.0, CSIRO, UKMO-HadCM3, IPSL, ECHAM5/MPI, CGCM3.1(T47), GFDL2.0, UKMO-HadGEM1, MIROC3.2(medres), MRI-CGCM2.3.2a, CNRM, GFDL2.1, INM-CM3, ECHO-G, PCM). See Wehner, M., 2005: Changes in daily precipitation and surface air temperature extremes in the IPCC AR4 models. *US CLIVAR Variations*, **3(3)**, 5-9.
- Hatching indicates at least two out of three models agree on the sign of the projected change in precipitation.
- We acknowledge the modeling groups, the Program for Climate Model Diagnosis and Intercomparison (PCMDI) and the WCRP's Working Group on Coupled Modelling (WGCM) for their roles in making available the WCRP CMIP3 multi-model dataset, <<http://www.pcmdi.llnl.gov/projects/cmip/index.php>>. Support of this dataset is provided by the Office of Science, U.S. Department of Energy. For an overview and documentation of the CMIP3 modeling activity, see Meehl, G.A., C. Covey, T. Delworth, M. Latif, B. McAvaney, J.F.B. Mitchell, R.J. Stouffer, and K.E. Taylor, 2007: The WCRP CMIP3 multi-model dataset: a new era in climate change research. *Bulletin of the American Meteorological Society*, **88(9)**, 1383-1394.
- ⁹⁴ Hare, B. and M. Meinshausen, 2006: How much warming are we committed to and how much can be avoided? *Climatic Change*, **75(1)**, 111-149.
- ⁹⁵ den Elzen, M.G.J. and M. Meinshausen, 2006: Multi-gas emission pathways for meeting the EU 2°C climate target. In: *Avoiding Dangerous Climate Change* [Schellnhuber, J.S., W. Cramer, N. Nakićenović, T.M.L. Wigley and G. Yohe (eds.)]. Cambridge University Press, Cambridge, UK, and New York, pp. 299-310.
- ⁹⁶ Seidel, D.J., Q. Fu, W.J. Randel, and T.J. Reichler, 2008: Widening of the tropical belt in a changing climate. *Nature Geoscience*, **1(1)**, 21-24.
- ⁹⁷ Cook, E.R., P.J. Bartlein, N. Diffenbaugh, R. Seager, B.N. Shuman, R.S. Webb, J.W. Williams, and C. Woodhouse, 2008: Hydrological variability and change. In: *Abrupt Climate Change*. Synthesis and Assessment Product 3.4. U.S. Geological Survey, Reston, VA, pp. 143-257.
- ⁹⁸ Emanuel, K., 2005: Increasing destructiveness of tropical cyclones over the past 30 years. *Nature*, **436(7051)**, 686-688.

- ⁹⁹ Vecchi, G.A., K.L. Swanson, and B.J. Soden, 2008: Whither hurricane activity? *Science*, **322(5902)**, 687-689.
- ¹⁰⁰ Emanuel, K., R. Sundararajan, and J. Williams, 2008: Hurricanes and global warming: Results from downscaling IPCC AR4 simulations. *Bulletin of the American Meteorological Society*, **89(3)**, 347-367.
- ¹⁰¹ Vecchi, G.A. and B.J. Soden, 2007: Effect of remote sea surface temperature change on tropical cyclone potential intensity. *Nature*, **450(7172)**, 1077-1070.
- ¹⁰² Alley, R.B., P.U. Clark, P. Huybrechts, and I. Joughin, 2005: Ice-sheet and sea-level changes ice-sheet and sea-level changes. *Science*, **310(5747)**, 456-460.
- ¹⁰³ Rahmstorf, S., 2007: A semi-empirical approach to projecting future sea-level rise. *Science*, **315(5810)**, 368-370.
- ¹⁰⁴ Mitrovica, J.X., N. Gomez, and P.U. Clark, 2009: The sea-level fingerprint of West Antarctic collapse. *Science*, **323(5915)**, 753.
- ¹⁰⁵ Clark, P.U., A.J. Weaver, E. Brook, E.R. Cook, T.L. Delworth, and K. Steffen, 2008: Introduction: Abrupt changes in the Earth's climate system. In: *Abrupt Climate Change*. Synthesis and Assessment Product 3.4. U.S. Geological Survey, Reston, VA, pp. 19-59.
- ¹⁰⁶ Brook, E., D. Archer, E. Dlugokencky, S. Frohking, and D. Lawrence, 2008: Potential for abrupt changes in atmospheric methane. In: *Abrupt Climate Change*. Synthesis and Assessment Product 3.4. U.S. Geological Survey, Reston, VA, pp. 360-452.
- ¹⁰⁷ Temperatures for the contiguous U.S. are based on data from the US Historical Climatology Network Version 2 (Menne *et al.* 2008). Temperatures for Alaska, Hawaii, and Puerto Rico are based on data from the Cooperative Observers Network adjusted to remove non-climatic influences such as changes in instruments and observer practices and changes in the station environment (Menne and Williams, 2008).
- US time series on page 27 is calculated with data for the contiguous US, Alaska, and Hawaii. US map on page 28 lower left includes observed temperature change in Puerto Rico. Winter temperature trend map in the agriculture section, page 76, is for the contiguous US only.
- References for this endnote:**
- Menne, M.J., C.N. Williams, and R.S. Vose, 2008: The United States Historical Climatology Network Monthly Temperature Data - Version 2. *Bulletin of the American Meteorological Society*, Early online release, 25 February 2009, doi:10.1175/2008BAMS2613.1
- Menne, M.J. and C.N. Williams Jr., 2008: Homogenization of temperature series via pairwise comparisons. *Journal of Climate*, **22(7)**, 1700-1717.
- ¹⁰⁸ Christensen, J.H., B. Hewitson, A. Busuioc, A. Chen, X. Gao, I. Held, R. Jones, R.K. Kolli, W.-T. Kwon, R. Laprise, V. Magaña Rueda, L. Mearns, C.G. Menéndez, J. Räisänen, A. Rinke, A. Sarr, and P. Whetton, 2007: Regional climate projections. In: *Climate Change 2007: The Physical Basis*. Contribution of Working Group I to the Fourth Assessment Report of the Intergovernmental Panel on Climate Change [Solomon, S., D. Qin, M. Manning, Z. Chen, M. Marquis, K.B. Averyt, M. Tignor, and H.L. Miller (eds.)]. Cambridge University Press, Cambridge, UK, and New York, pp. 847-940.
- ¹⁰⁹ CMIP3-C: Analysis for the CONUS was based on methods described in: Hayhoe, K., D. Cayan, C.B. Field, P.C. Frumhoff, E.P. Maurer, N.L. Miller, S.C. Moser, S.H. Schneider, K.N. Cahill, E.E. Cleland, L. Dale, R. Drapek, R.M. Hanemann, L.S. Kalkstein, J. Lenihan, C.K. Lunch, R.P. Neilson, S.C. Sheridan, and J.H. Verville, 2004: Emission pathways, climate change, and impacts on California. *Proceedings of the National Academy of Sciences*, **101(34)**, 12422-12427; and Hayhoe, K., C. Wake, B. Anderson, X.-Z. Liang, E. Maurer, J. Zhu, J. Bradbury, A. DeGaetano, A.M. Stoner, and D. Wuebbles, 2008: Regional climate change projec-

tions for the Northeast USA. *Mitigation and Adaptation Strategies for Global Change*, **13(5-6)**, 425-436. This analysis uses 16 models simulations from the WCRP CMIP3. Where models had multiple runs, only the first run available from each model was used. See <http://gdo-dcp.ucllnl.org/downscaled_cmip3_projections/dcpInterface.html> for more information.

The Alaskan projections are based on 14 models that best captured the present climate of Alaska; see Walsh, J.E., W.L. Chaman, V. Romanovsky, J.H. Christensen, and M. Stendel, 2008: Global climate model performance over Alaska and Greenland. *Journal of Climate*, **21(23)**, 6156-6174.

Caribbean and Pacific islands analyses use 15 models simulations from the WCRP CMIP3 that were available at resolutions finer than 4 degrees (CCSM3.0, CSIRO, UKMO-HadCM3, IPSL, ECHAM5/MPI, CGCM3.1(T47), GFDL2.0, UKMO-HadGEM1, MIROC3.2(medres), MRI-CGCM2.3.2a, CNRM, GFDL2.1, INM-CM3, ECHO-G, PCM). See Wehner, M., 2005: Changes in daily precipitation and surface air temperature extremes in the IPCC AR4 models. *US CLIVAR Variations*, **3(3)**, 5-9.

We acknowledge the modeling groups, the Program for Climate Model Diagnosis and Intercomparison (PCMDI) and the WCRP's Working Group on Coupled Modelling (WGCM) for their roles in making available the WCRP CMIP3 multi-model dataset, <<http://www.pcmdi.llnl.gov/projects/cmip/index.php>>. Support of this dataset is provided by the Office of Science, U.S. Department of Energy. For an overview and documentation of the CMIP3 modeling activity, see Meehl, G.A., C. Covey, T. Delworth, M. Latif, B. McAvaney, J.F.B. Mitchell, R.J. Stouffer, and K.E. Taylor, 2007: The WCRP CMIP3 multi-model dataset: a new era in climate change research. *Bulletin of the American Meteorological Society*, **88(9)**, 1383-1394.

- ¹¹⁰ Detailed local-scale projections about temperature and precipitation changes displayed in this report were generated using well-documented "statistical downscaling" techniques [Wood *et al.*, 2002] for the contiguous U.S. and Alaska. These techniques use statistical relationships between surface observations and climate simulations of the past to develop modifications for the global model results. These modifications are then applied to the climate projections for the future scenarios. The approach is also used to drive daily simulations by a well-established hydrological modeling framework for the contiguous U.S. [Liang *et al.*, 1994]. This method, which modifies global climate model simulations to better account for landscape variations and other features affecting climate at the regional to local scale, has been previously applied to generate high-resolution regional climate projections for the Northeast, Midwest, Northwest, and Southwest [Wood *et al.*, 2004; Hayhoe *et al.*, 2004; Hayhoe *et al.*, 2008; Cayan *et al.*, 2008; Cherkauer *et al.*, 2009]. Comparison of these methods with dynamically downscaled projections generated using regional climate model simulations provide strong justification for the use of such techniques [Wood *et al.*, 2004; Hayhoe *et al.*, 2008].

References for this endnote:

Cayan, D., E. Maurer, M. Dettinger, M. Tyree, and K. Hayhoe, 2008: Climate change scenarios for the California region. *Climatic Change*, **87(Supplement 1)**, S21-S42.

Cherkauer, K. and T. Sinha, 2009: Hydrologic impacts of projected future climate change in the Lake Michigan region. *Journal of Great Lakes Research*, in press.

Hayhoe, K., D. Cayan, C.B. Field, P.C. Frumhoff, E.P. Maurer, N.L. Miller, S.C. Moser, S.H. Schneider, K.N. Cahill, E.E. Cleland, L. Dale, R. Drapek, R.M. Hanemann, L.S. Kalkstein, J. Lenihan, C.K. Lunch, R.P. Neilson, S.C. Sheridan, and J.H. Verville, 2004: Emission pathways, climate change, and impacts on Cali-

- for California. *Proceedings of the National Academy of Sciences*, **101(34)**, 12422-12427.
- Hayhoe, K., C. Wake, B. Anderson, X.-Z. Liang, E. Maurer, J. Zhu, J. Bradbury, A. DeGaetano, A.M. Stoner, and D. Wuebbles, 2008: Regional climate change projections for the Northeast USA. *Mitigation and Adaptation Strategies for Global Change*, **13(5-6)**, 425-436.
- Liang, X., D. Lettenmaier, E. Wood, and S. Burges, 1994: A simple hydrologically-based model of land surface water and energy fluxes for general circulation models. *Journal of Geophysical Research*, **99(D7)**, 14415-14428.
- Maurer, E.P., A.W. Wood, J.C. Adam, D.P. Lettenmaier, and B. Nijssen, 2002: A long-term hydrologically-based data set of land surface fluxes and states for the conterminous United States. *Journal of Climate*, **15(22)**, 3237-3251.
- Wood, A.W., L.R. Leung, V. Sridhar, and D.P. Lettenmaier, 2004: Hydrologic implications of dynamical and statistical approaches to downscaling climate model outputs. *Climatic Change*, **62(1-3)**, 189-216.
- Wood, A.W., E.P. Maurer, A. Kumar, and D.P. Lettenmaier, 2002: Long range experimental hydrologic forecasting for the eastern U.S. *Journal of Geophysical Research*, **107(D20)**, 4429, doi:10.1029/2001JD000659.
- ¹¹¹ NOAA's National Climatic Data Center, 2008: *The USHCN Version 2 Serial Monthly Dataset*. [Web site] <<http://www.ncdc.noaa.gov/oa/climate/research/ushcn/>>
- ¹¹² Kunkel, K.E., P.D. Bromirski, H.E. Brooks, T. Cavazos, A.V. Douglas, D.R. Easterling, K.A. Emanuel, P.Ya. Groisman, G.J. Holland, T.R. Knutson, J.P. Kossin, P.D. Komar, D.H. Levinson, and R.L. Smith, 2008: Observed changes in weather and climate extremes. In: *Weather and Climate Extremes in a Changing Climate: Regions of Focus: North America, Hawaii, Caribbean, and U.S. Pacific Islands* [Karl, T.R., G.A. Meehl, C.D. Miller, S.J. Hassol, A.M. Waple, and W.L. Murray (eds.)]. Synthesis and Assessment Product 3.3. U.S. Climate Change Science Program, Washington, DC, pp. 35-80.
- ¹¹³ Groisman, P.Ya., R.W. Knight, T.R. Karl, D.R. Easterling, B. Sun, and J.H. Lawrimore, 2004: Contemporary changes of the hydrological cycle over the contiguous United States, trends derived from *in situ* observations. *Journal of Hydrometeorology*, **5(1)**, 64-85. The climate regions are different than those used in this article but the methodology is identical.
- ¹¹⁴ Karl, T.R., G.A. Meehl, T.C. Peterson, K.E. Kunkel, W.J. Gutowski Jr., and D.R. Easterling, 2008: Executive summary. In: *Weather and Climate Extremes in a Changing Climate. Regions of Focus: North America, Hawaii, Caribbean, and U.S. Pacific Islands* [Karl, T.R., G.A. Meehl, C.D. Miller, S.J. Hassol, A.M. Waple, and W.L. Murray (eds.)]. Synthesis and Assessment Product 3.3. U.S. Climate Change Science Program, Washington, DC, pp. 1-9.
- ¹¹⁵ Seager, R., M. Ting, I. Held, Y. Kushnir, J. Lu, G. Vecchi, H.-P. Huang, N. Harnik, A. Leetmaa, N.-C. Lau, C. Li, J. Velez, and N. Naik, 2007: Model projections of an imminent transition to a more arid climate in southwestern North America. *Science*, **316(5828)**, 1181-1184.
- ¹¹⁶ USGS, 2005: *Changes in Streamflow Timing in the Western United States in Recent Decades*. USGS fact sheet 2005-3018. U.S. Geological Survey, National Streamflow Information Program, La Jolla, CA, 4 pp. <<http://pubs.usgs.gov/fs/2005/3018/>>
- ¹¹⁷ CMIP3-B: Analysis for the CONUS was based on methods described in: Hayhoe, K., D. Cayan, C.B. Field, P.C. Frumhoff, E.P. Maurer, N.L. Miller, S.C. Moser, S.H. Schneider, K.N. Cahill, E.E. Cleland, L. Dale, R. Drapek, R.M. Hanemann, L.S. Kalkstein, J. Lenihan, C.K. Lunch, R.P. Neilson, S.C. Sheridan, and J.H. Verville, 2004: Emission pathways, climate change, and impacts on California. *Proceedings of the National Academy of Sciences*, **101(34)**, 12422-12427; and Hayhoe, K., C. Wake, B. Anderson, X.-Z. Liang, E. Maurer, J. Zhu, J. Bradbury, A. DeGaetano, A.M. Stoner, and D. Wuebbles, 2008: Regional climate change projections for the Northeast USA. *Mitigation and Adaptation Strategies for Global Change*, **13(5-6)**, 425-436. This analysis uses 16 models simulations from the WCRP CMIP3. Where models had multiple runs, only the first run available from each model was used. See <http://gdo-dcp.ucllnl.org/downscaled_cmip3_projections/dcpInterface.html> for more information.
- We acknowledge the modeling groups, the Program for Climate Model Diagnosis and Intercomparison (PCMDI) and the WCRP's Working Group on Coupled Modelling (WGCM) for their roles in making available the WCRP CMIP3 multi-model dataset, <<http://www.pcmdi.llnl.gov/projects/cmip/index.php>>. Support of this dataset is provided by the Office of Science, U.S. Department of Energy. For an overview and documentation of the CMIP3 modeling activity, see Meehl, G.A., C. Covey, T. Delworth, M. Latif, B. McAvaney, J.F.B. Mitchell, R.J. Stouffer, and K.E. Taylor, 2007: The WCRP CMIP3 multi-model dataset: a new era in climate change research. *Bulletin of the American Meteorological Society*, **88(9)**, 1383-1394.
- ¹¹⁸ Swanson, K.L., 2008: Nonlocality of Atlantic tropical cyclone intensities. *Geochemistry, Geophysics, Geosystems*, **9**, Q04V01, doi:10.1029/2007GC001844.
- ¹¹⁹ Knutson, T.R., J.J. Sirutis, S.T. Garner, G.A. Vecchi, and I. Held, 2008: Simulated reduction in Atlantic hurricane frequency under twenty-first-century warming conditions. *Nature Geoscience*, **1(6)**, 359-364.
- ¹²⁰ Emanuel, K., 2007: Environmental factors affecting tropical cyclone power dissipation. *Journal of Climate*, **20(22)**, 5497-5509.
- ¹²¹ Number of strongest hurricanes, number of landfalling strongest hurricanes, and number of landfalling hurricanes are based on data obtained from NOAA's Oceanographic and Meteorological Laboratory: <<http://www.aoml.noaa.gov/hrd/hurdat/ushurrlst18512005-gt.txt>> with updates. The total number of named storms are adjusted to account for missing tropical storms and hurricanes in the pre-satellite era using the method of Vecchi and Knutson (2008). Basin and landfalling totals are displayed in 5-year increments (pentads) from 1881 through 2010. The final 5-year period was standardized to a comparable 5-year period assuming the level of activity from 2006 to 2008 persists through 2010.
- Vecchi, G.A. and T.R. Knutson, 2008: On estimates of historical North Atlantic tropical cyclone activity. *Journal of Climate*, **21(14)**, 3580-3600.
- ¹²² Elsner, J.B., J.P. Kossin, and T.H. Jagger, 2008: The increasing intensity of the strongest tropical cyclones. *Nature*, **455(7209)**, 92-95.
- ¹²³ Bell, G.D. and M. Chelliah, 2006: Leading tropical modes associated with interannual and multidecadal fluctuations in North Atlantic hurricane activity. *Journal of Climate*, **19(4)**, 590-612.
- ¹²⁴ Levinson, D.H. and J. Lawrimore (eds.), 2008: State of the climate in 2007. *Bulletin of the American Meteorological Society*, **89(7, Supplement)**, S1-S179.
- ¹²⁵ Kossin, J.P., K.R. Knapp, D.J. Vimont, R.J. Murnane, and B.A. Harper, 2007: A globally consistent reanalysis of hurricane variability and trends. *Geophysical Research Letters*, **34**, L04815, doi:10.1029/2006GL028836.
- ¹²⁶ Rahmstorf, S., A. Cazenave, J.A. Church, J.E. Hansen, R.F. Keeling, D.E. Parker, and R.C.J. Somerville, 2007: Recent climate observations compared to projections. *Science*, **316(5825)**, 709.
- ¹²⁷ Zervas, C., 2001: *Sea Level Variations of the United States 1985-1999*. NOAA technical report NOS CO-OPS 36. National Oceanic and Atmospheric Administration, Silver Spring, MD.

- anic and Atmospheric Administration, Silver Spring, MD, 66 pp. <<http://tidesandcurrents.noaa.gov/publications/techrpt36doc.pdf>>
- ¹²⁸ Sea-level rise numbers are calculated based on an extrapolation of NOAA tide gauge stations with records exceeding 50 years, as reported in Zervas, C., 2001: *Sea Level Variations of the United States 1985-1999*. NOAA technical report NOS CO-OPS 36. National Oceanic and Atmospheric Administration, Silver Spring, MD, 66 pp. <<http://tidesandcurrents.noaa.gov/publications/techrpt36doc.pdf>>
- ¹²⁹ Kunkel, K.E., N.E. Westcott, and D.A.R. Knistovich, 2002: Assessment of potential effects of climate changes on heavy lake-effect snowstorms near Lake Erie. *Journal of Great Lakes Research*, **28(4)**, 521-536.
- ¹³⁰ Burnett, A.W., M.E. Kirby, H.T. Mullins, and W.P. Patterson, 2003: Increasing Great Lake-effect snowfall during the twentieth century: a regional response to global warming? *Journal of Climate*, **16(21)**, 3535-3542.
- ¹³¹ Trapp, R.J., N.S. Diffenbaugh, H.E. Brooks, M.E. Baldwin, E.D. Robinson, and J.S. Pal, 2007: Changes in severe thunderstorm environment frequency during the 21st century caused by anthropogenically enhanced global radiative forcing. *Proceedings of the National Academy of Sciences*, **104(50)**, 19719-19723.
- ¹³² ACIA, 2005: *Arctic Climate Impact Assessment*. Cambridge University Press, Cambridge, UK, and New York, 1042 pp. <<http://www.acia.uaf.edu/pages/scientific.html>>
- ¹³³ Stroeve, J., M.M. Holland, W. Meier, T. Scambos, and M. Serreze, 2007: Arctic sea ice decline: faster than forecast. *Geophysical Research Letters*, **34**, L09501, doi:10.1029/2007GL029703.
- ¹³⁴ L'Heureux, M.L., A. Kumar, G.D. Bell, M.S. Halpert, and R.W. Higgins, 2008: Role of the Pacific-North American (PNA) pattern in the 2007 Arctic sea ice decline. *Geophysical Research Letters*, **35**, L20701, doi:10.1029/2008GL035205.
- ¹³⁵ Johannessen, O.M., 2008: Decreasing Arctic sea ice mirrors increasing CO₂ on decadal time scale. *Atmospheric and Oceanic Science Letters*, **1(1)**, 51-56.
- ¹³⁶ National Snow and Ice Data Center, 2008: *Arctic Sea Ice Down to Second-Lowest Extent; Likely Record-Low Volume*. Press release October 2, 2008. <http://nsidc.org/news/press/20081002_seaice_pressrelease.html>
- ¹³⁷ Polyak, L., J. Andrews, J. Brigham-Grette, D. Darby, A. Dyke, S. Funder, M. Holland, A. Jennings, J. Savelle, M. Serreze, and E. Wolff, 2009: History of sea ice in the Arctic. In: *Past Climate Variability and Change in the Arctic and at High Latitude*. Synthesis and Assessment Product 1.2. U.S. Geological Survey, Reston, VA, pp. 358-420.
- ¹³⁸ Images from *Sea Ice Yearly Minimum 1979-2007*. [Web site] NASA/Goddard Space Flight Center Scientific Visualization Studio. Thanks to Rob Gerston (GSFC) for providing the data. <<http://svs.gsfc.nasa.gov/goto?3464>>
- ¹³⁹ Fetterer, F., K. Knowles, W. Meier, and M. Savoie, 2002: updated 2008: *Sea Ice Index*. [Web site] National Snow and Ice Data Center, Boulder, CO. <http://nsidc.org/data/seaice_index/>
- ¹⁴⁰ Pacala, S., R. Birdsey, S. Bridgman, R.T. Conant, K. Davis, B. Hales, R. Houghton, J.C. Jenkins, M. Johnston, G. Marland, and K. Paustian, 2007: The North American carbon budget past and present. In: *The First State of the Carbon Cycle Report (SOCCR): The North American Carbon Budget and Implications for the Global Carbon Cycle* [King, A.W., L. Dilling, G.P. Zimmerman, D.F. Fairman, R.A. Houghton, G. Marland, A.Z. Rose, and T.J. Wilbanks (eds.)]. Synthesis and Assessment Product 2.2. National Oceanic and Atmospheric Administration, National Climatic Data Center, Asheville, NC, pp. 29-36.
- ¹⁴¹ Marland, G., R.J. Andres, T.J. Blasing, T.A. Boden, C.T. Broniak, J.S. Gregg, L.M. Losey, and K. Treanton, 2007: Energy, industry, and waste management activities: an introduction to CO₂ emissions from fossil fuels. In: *The First State of the Carbon Cycle Report (SOCCR): The North American Carbon Budget and Implications for the Global Carbon Cycle* [King, A.W., L. Dilling, G.P. Zimmerman, D.M. Fairman, R.A. Houghton, G. Marland, A.Z. Rose, and T.J. Wilbanks (eds.)]. Synthesis and Assessment Product 2.2. National Oceanic and Atmospheric Administration, National Climatic Data Center, Asheville, NC, pp. 57-64.
- ¹⁴² Bates, B.C., Z.W. Kundzewicz, S. Wu, and J.P. Palutikof (eds.), 2008: *Climate Change and Water*. Technical paper of the Intergovernmental Panel on Climate Change. IPCC Secretariat, Geneva, Switzerland, 210 pp.
- ¹⁴³ Feng, S. and Q. Hu, 2007: Changes in winter snowfall/precipitation ratio in the contiguous United States. *Journal of Geophysical Research*, **112**, D15109, doi:10.1029/2007JD008397.
- ¹⁴⁴ Guttman, N.B. and R.G. Quayle, 1996: A historical perspective of U.S. climate divisions. *Bulletin of the American Meteorological Society*, **77(2)**, 293-303. Operational practices described in this paper continue.
- ¹⁴⁵ Groisman, P.Ya., R.W. Knight, D.R. Easterling, T.R. Karl, G.C. Hegerl, and V.N. Razuvaev, 2005: Trends in intense precipitation in the climate record. *Journal of Climate*, **18(9)**, 1326-1350.
- ¹⁴⁶ Kundzewicz, Z.W., L.J. Mata, N.W. Arnell, P. Döll, P. Kabat, B. Jiménez, K.A. Miller, T. Oki, Z. Sen, and I.A. Shiklomanov, 2007: Freshwater resources and their management. In: *Climate Change 2007: Impacts, Adaptation and Vulnerability*. Contribution of Working Group II to the Fourth Assessment Report of the Intergovernmental Panel on Climate Change [Parry, M.L., O.F. Canziani, J.P. Palutikof, P.J. van der Linden, and C.E. Hanson (eds.)]. Cambridge University Press, Cambridge, UK, and New York, pp. 173-210.
- ¹⁴⁷ Groisman, P.Ya. and R.W. Knight, 2008: Prolonged dry episodes over the conterminous United States: new tendencies emerging during the last 40 years. *Journal of Climate*, **21(9)**, 1850-1862.
- ¹⁴⁸ Tebaldi, C., K. Hayhoe, J.M. Arblaster, and G.A. Meehl, 2006: Going to the extremes: an intercomparison of model-simulated historical and future changes in extreme events. *Climatic Change*, **79(3-4)**, 185-211.
- ¹⁴⁹ Lettenmaier, D., D. Major, L. Poff, and S. Running, 2008: Water resources. In: *The Effects of Climate Change on Agriculture, Land Resources, Water Resources, and Biodiversity in the United States* [Backlund, P., A. Janetos, D. Schimel, J. Hatfield, K. Boote, P. Fay, L. Hahn, C. Izaurre, B.A. Kimball, T. Mader, J. Morgan, D. Ort, W. Polley, A. Thomson, D. Wolfe, M.G. Ryan, S.R. Archer, R. Birdsey, C. Dahm, L. Heath, J. Hicke, D. Hollinger, T. Huxman, G. Okin, R. Oren, J. Randerson, W. Schlesinger, D. Lettenmaier, D. Major, L. Poff, S. Running, L. Hansen, D. Inouye, B.P. Kelly, L. Meyerson, B. Peterson, and R. Shaw (eds.)]. Synthesis and Assessment Product 4.3. U.S. Department of Agriculture, Washington, DC, pp. 121-150.
- ¹⁵⁰ Hayhoe, K., C.P. Wake, T.G. Huntington, L. Luo, M.D. Schwartz, J. Sheffield, E. Wood, B. Anderson, J. Bradbury, A. DeGaetano, T.J. Troy, and D. Wolfe, 2007: Past and future changes in climate and hydrological indicators in the U.S. Northeast. *Climate Dynamics*, **28(4)**, 381-407.
- ¹⁵¹ Milly, P.C.D., J. Betancourt, M. Falkenmark, R.M. Hirsch, Z.W. Kundzewicz, D.P. Lettenmaier, and R.J. Stouffer, 2008: Stationarity is dead: Whither water management? *Science*, **319(5863)**, 573-574.
- ¹⁵² Christensen, N.S., A.W. Wood, N. Voisin, D.P. Lettenmaier, and R.N. Palmer, 2004: The effects of climate change on the hydrology and water resources of the Colorado River basin. *Climatic Change*, **62(1-3)**, 337-363.

- ¹⁵³ Mote, P., A. Hamlet, and E. Salathé, 2008: Has spring snowpack declined in the Washington Cascades? *Hydrology and Earth System Sciences*, **12(1)**, 193-206.
- ¹⁵⁴ Knowles, N., M.D. Dettinger, and D.R. Cayan, 2006: Trends in snowfall versus rainfall in the western United States. *Journal of Climate*, **19(18)**, 4545-4559.
- ¹⁵⁵ Huntington T.G., G.A. Hodgkins, B.D. Keim, and R.W. Dudley, 2004: Changes in the proportion of precipitation occurring as snow in New England (1949 to 2000). *Journal of Climate*, **17(13)**, 2626-2636.
- ¹⁵⁶ Burakowski, E.A., C.P. Wake, B. Braswell, and D.P. Brown, 2008: Trends in wintertime climate in the northeastern United States: 1965–2005. *Journal of Geophysical Research*, **113**, D20114, doi:10.1029/2008JD009870/
- ¹⁵⁷ Stewart, I.T., D.R. Cayan, and M.D. Dettinger, 2004: Changes in snowmelt runoff timing in western North America under a ‘business as usual’ climate change scenario. *Climatic Change*, **62(1-3)**, 217-232.
- ¹⁵⁸ Stewart, I.T., D.R. Cayan, and M.D. Dettinger, 2005: Changes toward earlier streamflow timing across western North America. *Journal of Climate*, **18(8)**, 1136-1155.
- ¹⁵⁹ Rauscher, S.A., J.S. Pal, N.S. Diffenbaugh, and M.M. Benedetti, 2008: Future changes in snowmelt-driven runoff timing over the western United States. *Geophysical Research Letters*, **35**, L16703, doi:10.1029/2008GL034424.
- ¹⁶⁰ Pierce, D.W., T.P. Barnett, H.G. Hidalgo, T. Das, C. Bonfils, B.D. Santer, G. Bala, M.D. Dettinger, D.R. Cayan, A. Mirin, A.W. Wood, and T. Nozawa, 2008: Attribution of declining western U.S. snowpack to human effects. *Journal of Climate*, **21(23)**, 6425-6444.
- ¹⁶¹ Bonfils, C., B.D. Santer, D.W. Pierce, H.G. Hidalgo, G. Bala, T. Das, T.P. Barnett, D.R. Cayan, C. Doutriaux, A.W. Wood, A. Mirin, and T. Nozawa, 2008: Detection and attribution of temperature changes in the mountainous western United States. *Journal of Climate*, **21(23)**, 6404-6424.
- ¹⁶² U.S. Environmental Protection Agency, 2008: *National Water Program Strategy: Response to Climate Change*. U.S. Environmental Protection Agency, Washington, DC, 97 pp. <<http://www.epa.gov/water/climatechange/>>
- ¹⁶³ Ebi, K.L., J. Balbus, P.L. Kinney, E. Lipp, D. Mills, M.S. O’Neill, and M. Wilson, 2008: Effects of global change on human health. In: *Analyses of the Effects of Global Change on Human Health and Welfare and Human Systems* [Gamble, J.L. (ed.), K.L. Ebi, F.G. Sussman, and T.J. Wilbanks (authors)]. Synthesis and Assessment Product 4.6. U.S. Environmental Protection Agency, Washington, DC, pp. 2-1 to 2-78.
- ¹⁶⁴ Field, C.B., L.D. Mortsch, M. Brklacich, D.L. Forbes, P. Kovacs, J.A. Patz, S.W. Running, and M.J. Scott, 2007: North America. In: *Climate Change 2007: Impacts, Adaptation and Vulnerability*. Contribution of Working Group II to the Fourth Assessment Report of the Intergovernmental Panel on Climate Change [Parry, M.L., O.F. Canziani, J.P. Palutikof, P.J. van der Linden, and C.E. Hanson (eds.)]. Cambridge University Press, Cambridge, UK, and New York, pp. 617-652.
- ¹⁶⁵ Winter, J., J.W. Harvey, O.L. Franke, and W.M. Alley, 1998: *Ground Water and Surface Water: A Single Resource*. USGS circular 1139. U.S. Geological Survey, Denver, CO, 79 pp. <<http://pubs.usgs.gov/circ/circ1139/>>
- ¹⁶⁶ Austin, J.A. and S.M. Colman, 2007: Lake Superior summer water temperatures are increasing more rapidly than regional air temperatures: a positive ice-albedo feedback. *Geophysical Research Letters*, **34**, L06604, doi:10.1029/2006GL029021.
- ¹⁶⁶ U.S. General Accounting Office, 2003: *Freshwater Supply: States’ Views of How Federal Agencies Could Help Them Meet the Challenges of Expected Shortages*. GAO-03-514. General Accounting Office, Washington, DC, 110 pp. <<http://www.gao.gov/new.items/d03514.pdf>>
- ¹⁶⁸ U.S. Environmental Protection Agency, 2002: *The Clean Water and Drinking Water Infrastructure Gap Analysis*. EPA-816-R-02-020. U.S. Environmental Protection Agency, Washington, DC, 50 pp. <<http://www.epa.gov/safewater/gapreport.pdf>>
- ¹⁶⁹ National Research Council, 2004: *Confronting the Nation’s Water Problems: The Role of Research*. National Academies Press, Washington, DC, 310 pp.
- ¹⁷⁰ Yohe, G.W., R.D. Lasco, Q.K. Ahmad, N.W. Arnell, S.J. Cohen, C. Hope, A.C. Janetos, and R.T. Perez, 2007: Perspectives on climate change and sustainability. In: *Climate Change 2007: Impacts, Adaptation and Vulnerability*. Contribution of Working Group II to the Fourth Assessment Report of the Intergovernmental Panel on Climate Change [Parry, M.L., O.F. Canziani, J.P. Palutikof, P.J. van der Linden, and C.E. Hanson (eds.)]. Cambridge University Press, Cambridge, UK, and New York, pp. 811-841.
- ¹⁷¹ U.S. Bureau of Reclamation, 2005: *Water 2025: Preventing Crises and Conflict in the West*. U.S. Bureau of Reclamation, Washington, DC, 32 pp. Updated from USBR <<http://www.usbr.gov/uc/crsp/GetSiteInfo>>
- ¹⁷² Wilbanks, T.J., P. Romero Lankao, M. Bao, F. Berkhout, S. Cairncross, J.-P. Ceron, M. Kapshe, R. Muir-Wood, and R. Zapata-Marti, 2007: Industry, settlement and society. In: *Climate Change 2007: Impacts, Adaptation and Vulnerability*. Contribution of Working Group II to the Fourth Assessment Report of the Intergovernmental Panel on Climate Change [Parry, M.L., O.F. Canziani, J.P. Palutikof, P.J. van der Linden, and C.E. Hanson (eds.)]. Cambridge University Press, Cambridge, UK, and New York, pp. 357-390.
- ¹⁷³ Adger, W.N., S. Agrawala, M.M.Q. Mirza, C. Conde, K. O’Brien, J. Pulhin, R. Pulwarty, B. Smit, and K. Takahashi, 2007: Assessment of adaptation practices, options, constraints and capacity. In: *Climate Change 2007: Impacts, Adaptation and Vulnerability*. Contribution of Working Group II to the Fourth Assessment Report of the Intergovernmental Panel on Climate Change [Parry, M.L., O.F. Canziani, J.P. Palutikof, P.J. van der Linden, and C.E. Hanson (eds.)]. Cambridge University Press, Cambridge, UK, and New York, pp. 717-743.
- ¹⁷⁴ Hartmann, H.C., 2008: Decision support for water resources management. In: *Uses and Limitations of Observations, Data, Forecasts and Other Projections in Decision Support for Selected Sectors and Regions*. Synthesis and Assessment Product 5.1. U.S. Climate Change Science Program, Washington, DC, pp. 45-55.
- ¹⁷⁵ Ruhl, J.B., 2005: Water wars, eastern style: divvying up the Apalachicola-Chattahoochee-Flint River Basin. *Journal of Contemporary Water Research & Education*, **131**, 47-54.
- ¹⁷⁶ Leitman, S., 2008: Lessons learned from transboundary management efforts in the Apalachicola-Chattahoochee-Flint Basin, USA. In: *Transboundary Water Resources: A Foundation for Regional Stability in Central Asia* [Moerlins, J.E., M.K. Khankhasayev, S.F. Leitman, and E.J. Makhmudov (eds.)]. Springer, Dordrecht and London, pp. 195-208.
- ¹⁷⁷ Gobalet, K., P. Schulz, T. Wake, and N. Siefkin, 2004: Archaeological perspectives on Native American fisheries of California, with emphasis on steelhead and salmon. *Transactions of the American Fisheries Society*, **133(4)**, 801-833.
- ¹⁷⁸ Hammersmark, C., W. Fleenor, and S. Schladow, 2005: Simulation of flood impact and habitat extent for a tidal freshwater marsh restoration. *Ecological Engineering*, **25(2)**, 137-152.
- ¹⁷⁹ Kondolf, G., P. Angermeier, K. Cummins, T. Dunne, M. Healey, W. Kimmerer, P. Moyle, D. Murphy, D. Patten, S. Railsback, D. Reed, R. Spies, and R. Twiss, 2008: Projecting cumulative benefits of multiple river restoration projects: An example from the

- Sacramento-San Joaquin River system in California. *Environmental Management*, **42**(6), 933-945.
- ¹⁸⁰ McKee, L., N. Ganju, and D. Schoellhamer, 2006: Estimates of suspended sediment entering San Francisco Bay from the Sacramento and San Joaquin Delta, San Francisco Bay, California. *Journal of Hydrology*, **323**(1-4), 335-352.
- ¹⁸¹ Trenham, P., H. Shaffer, and P. Moyle, 1998: Biochemical identification and assessment of population subdivision in morphologically similar native and invading smelt species (*Hypomesus*) in the Sacramento San Joaquin estuary, California. *Transactions of the American Fisheries Society*, **127**(3), 417-424.
- ¹⁸² Vengosh, A., J. Gill, M. Davisson, and G. Hudson, 2002: A multi-isotope (B, Sr, O, H, and C) and age dating (³H-³He and ¹⁴C) study of groundwater from Salinas Valley, California: hydrochemistry, dynamics, and contamination processes. *Water Resources Research*, **38**(1), 1008, doi:10.1029/2001WR000517.
- ¹⁸³ Haggerty, G.M., D. Tave, R. Schmidt-Petersen, and J. Stomp, 2008: Raising endangered fish in New Mexico. *Southwest Hydrology*, **7**(4), 20-21.
- ¹⁸⁴ National Research Council, 2004: *Endangered and Threatened Fishes in the Klamath River Basin: Causes of Decline and Strategies for Recovery*. National Academies Press, Washington, DC, 398 pp.
- ¹⁸⁵ National Research Council, 2008: *Hydrology, Ecology, and Fishes of the Klamath River Basin*. National Academies Press, Washington, DC, 272 pp.
- ¹⁸⁶ National Research Council, 2007: *Colorado River Basin Water Management: Evaluating and Adjusting to Hydroclimatic Variability*. National Academies Press, Washington, DC, 218 pp.
- ¹⁸⁷ U.S. Bureau of Reclamation, 2007: *Colorado River Interim Guidelines for Lower Basin Shortages and Coordinated Operations for Lake Powell and Lake Mead: Final Environmental Impact Statement*. [U.S. Bureau of Reclamation, Boulder City, NV], 4 volumes. <<http://www.usbr.gov/lc/region/programs/strategies/FEIS/index.html>>
- ¹⁸⁸ University of Arizona, Undated: *Native American Water Rights in Arizona*. [Web site] <<http://www.library.arizona.edu/about/libraries/govdocs/waterdoc.html>>
- ¹⁸⁹ Cook, E.R., C.A. Woodhouse, C.M. Eakin, D.M. Meko, and D.W. Stahle, 2004: Long-term aridity changes in the western United States. *Science*, **306**(5698), 1015-1018.
- ¹⁹⁰ Ingram, H., D. Feldman, N. Mantua, K.L. Jacobs, D. Fort, N. Beller-Simms, and A.M. Waple, 2008: The changing context. In: *Decision-Support Experiments and Evaluations Using Seasonal-to-Interannual Forecasts and Observational Data: A Focus on Water Resources* [Beller-Simms, N., H. Ingram, D. Feldman, N. Mantua, K.L. Jacobs, and A.M. Waple (eds.)]. Synthesis and Assessment Product 5.3. NOAA's National Climatic Data Center, Asheville, NC, pp. 7-28.
- ¹⁹¹ Bull, S.R., D.E. Bilello, J. Ekmann, M.J. Sale, and D.K. Schmalzer, 2007: Effects of climate change on energy production and distribution in the United States. In: *Effects of Climate Change on Energy Production and Use in the United States* [Wilbanks, T.J., V. Bhatt, D.E. Bilello, S.R. Bull, J. Ekmann, W.C. Horak, Y.J. Huang, M.D. Levine, M.J. Sale, D.K. Schmalzer, and M.J. Scott (eds.)]. Synthesis and Assessment Product 4.5. U.S. Climate Change Science Program, Washington, DC, pp. 45-80.
- ¹⁹² Hyman, R.C., J.R. Potter, M.J. Savonis, V.R. Burkett, and J.E. Tump, 2008: Why study climate change impacts on transportation? In: *Impacts of Climate Change and Variability on Transportation Systems and Infrastructure: Gulf Coast Study, Phase I* [Savonis, M.J., V.R. Burkett, and J.R. Potter (eds.)]. Synthesis and Assessment Product 4.7. U.S. Department of Transportation, Washington, DC, pp. 1-1 to 1F-2 [48 pp.]
- ¹⁹³ Hatfield, J., K. Boote, P. Fay, L. Hahn, C. Izaurralde, B.A. Kimball, T. Mader, J. Morgan, D. Ort, W. Polley, A. Thomson, and D. Wolfe, 2008: Agriculture. In: *The Effects of Climate Change on Agriculture, Land Resources, Water Resources, and Biodiversity in the United States* [Backlund, P., A. Janetos, D. Schimel, J. Hatfield, K. Boote, P. Fay, L. Hahn, C. Izaurralde, B.A. Kimball, T. Mader, J. Morgan, D. Ort, W. Polley, A. Thomson, D. Wolfe, M.G. Ryan, S.R. Archer, R. Birdsey, C. Dahm, L. Heath, J. Hicke, D. Hollinger, T. Huxman, G. Okin, R. Oren, J. Randerson, W. Schlesinger, D. Lettenmaier, D. Major, L. Poff, S. Running, L. Hansen, D. Inouye, B.P. Kelly, L. Meyerson, B. Peterson, and R. Shaw (eds.)]. Synthesis and Assessment Product 4.3. U.S. Department of Agriculture, Washington, DC, pp. 21-74.
- ¹⁹⁴ Feldman, D.L., K.L. Jacobs, G. Garfin, A. Georgakakos, B. Morehouse, P. Restrepo, R. Webb, B. Yarnal, D. Basketfield, H.C. Hartmann, J. Kochendorfer, C. Rosenzweig, M. Sale, B. Udall, and C. Woodhouse, 2008: Making decision-support information useful, useable, and responsive to decision-maker needs In: *Decision-Support Experiments and Evaluations Using Seasonal-to-Interannual Forecasts and Observational Data: A Focus on Water Resources* [Beller-Simms, N., H. Ingram, D. Feldman, N. Mantua, K.L. Jacobs, and A.M. Waple (eds.)]. Synthesis and Assessment Product 5.3. NOAA's National Climatic Data Center, Asheville, NC, pp. 101-140.
- ¹⁹⁵ McCabe, G.J. and D.M. Wolock, 2007: Warming may create substantial water supply shortages in the Colorado River basin. *Geophysical Research Letters*, **34**, L22708, doi:10.1029/2007GL031764.
- ¹⁹⁶ Brekke, L., B. Harding, T. Piechota, B. Udall, C. Woodhouse, and D. Yates (eds.), 2007: Appendix U: Climate Technical Work Group Report: Review of science and methods for incorporating climate change information into Bureau of Reclamation's Colorado River Basin planning studies. In: *Colorado River Interim Guidelines for Lower Basin Shortages and Coordinated Operations for Lake Powell and Lake Mead: Final Environmental Impact Statement*. Bureau of Reclamation, Boulder City, NV, 110 pp. <<http://www.usbr.gov/lc/region/programs/strategies/FEIS/index.html>>
- ¹⁹⁷ U.S. Department of Energy, 2006: *Energy Demands on Water Resources. Report to Congress on the Interdependency of Energy and Water*. Sandia National Laboratories, Albuquerque, NM, 80 pp. <<http://www.sandia.gov/energy-water/docs/121-RptToCongress-EWwEIAcomments-FINAL.pdf>>
- ¹⁹⁸ California Energy Commission, 2005: *California's Water -- Energy Relationship*. CEC-700-2005-011-SF. California Energy Commission, [Sacramento], 174 pp. <<http://www.energy.ca.gov/2005publications/CEC-700-2005-011/CEC-700-2005-011-SF.PDF>>
- ¹⁹⁹ California Energy Commission, 2006: *Refining Estimates of Water-related Energy Use in California*. CEC-500-2006-118. California Energy Commission, [Sacramento]. <<http://www.energy.ca.gov/2006publications/CEC-500-2006-118/CEC-500-2006-118.PDF>>
- ²⁰⁰ U.S. Energy Information Administration, 2008: *Energy in Brief: What are Greenhouse Gases and How Much are Emitted by the United States?* [Web site] Energy Information Administration, Washington, DC. <http://tonto.eia.doe.gov/energy_in_brief/greenhouse_gas.cfm>
- ²⁰¹ Wilbanks, T.J., et al., 2007: Executive summary. In: *Effects of Climate Change on Energy Production and Use in the United States* [Wilbanks, T.J., V. Bhatt, D.E. Bilello, S.R. Bull, J. Ekmann, W.C. Horak, Y.J. Huang, M.D. Levine, M.J. Sale, D.K. Schmalzer, and M.J. Scott (eds.)]. Synthesis and Assessment Product 4.5. U.S. Climate Change Science Program, Washington, DC, pp. x-xii.
- ²⁰² From *Inventory of U.S. Greenhouse Gas Emissions and Sinks: 1990-2003*, U.S. EPA. Allocations from "Electricity & Heat" and

- “Industry” to end uses are WRI estimates based on energy use data from the International Energy Agency (IEA, 2005). All data is for 2003. All calculations are based on CO₂ equivalents, using 100-year global warming potentials from the IPCC (1996).
- ²⁰³ U.S. Energy Information Administration, 2008: *Annual Energy Review 2007*. U.S. Department of Energy, Washington, DC, 400 pp. <<http://www.eia.doe.gov/emeu/aer/pdf/aer.pdf>>
- ²⁰⁴ Bhatt, V., J. Eckmann, W.C. Horak, and T.J. Wilbanks, 2007: Possible indirect effects on energy production and distribution in the United States. In: *Effects of Climate Change on Energy Production and Use in the United States* [Wilbanks, T.J., V. Bhatt, D.E. Bilello, S.R. Bull, J. Eckmann, W.C. Horak, Y.J. Huang, M.D. Levine, M.J. Sale, D.K. Schmalzer, and M.J. Scott (eds.)]. Synthesis and Assessment Product 4.5. U.S. Climate Change Science Program, Washington, DC, pp. 81-97.
- ²⁰⁵ Scott, M.J. and Y.J. Huang, 2007: Effects of climate change on energy use in the United States. In: *Effects of Climate Change on Energy Production and Use in the United States* [Wilbanks, T.J., V. Bhatt, D.E. Bilello, S.R. Bull, J. Eckmann, W.C. Horak, Y.J. Huang, M.D. Levine, M.J. Sale, D.K. Schmalzer, and M.J. Scott (eds.)]. Synthesis and Assessment Product 4.5. U.S. Climate Change Science Program, Washington, DC, pp. 8-44.
- ²⁰⁶ U.S. Department of Energy, 2008: *Buildings Energy Data Book*. U.S. Department of Energy, Office of Energy Efficiency and Renewable Energy, [Washington, DC]. <<http://buildingsdatabook.eere.energy.gov/>>
- ²⁰⁷ U.S. Census Bureau, 2002: *Population of States and Counties of the United States: 1790-2000*. U.S. Census Bureau, Washington, DC, 226 pp. <<http://www.census.gov/population/www/censusdata/his-cendata.html>> See <<http://www.census.gov/popest/counties/>> for 2008 estimates.
- ²⁰⁸ Feldman, D.L., K.L. Jacobs, G. Garfin, A. Georgakakos, B. Morehouse, R. Webb, B. Yarnal, J. Kochendorfer, C. Rosenzweig, M. Sale, B. Udall, and C. Woodhouse, 2008: Decision-support experiments within the water resource management sector. In: *Decision-Support Experiments and Evaluations Using Seasonal-to-Interannual Forecasts and Observational Data: A Focus on Water Resources* [Beller-Simms, N., H. Ingram, D. Feldman, N. Mantua, K.L. Jacobs, and A.M. Waple (eds.)]. Synthesis and Assessment Product 5.3. NOAA's National Climatic Data Center, Asheville, NC, pp. 65-100.
- ²⁰⁹ Hightower, M. and S.A. Pierce, 2008: The energy challenge. *Nature*, **452(7185)**, 285-286.
- ²¹⁰ Wilbanks, T.J., et al., 2007: Conclusions and research priorities. In: *Effects of Climate Change on Energy Production and Use in the United States* [Wilbanks, T.J., V. Bhatt, D.E. Bilello, S.R. Bull, J. Eckmann, W.C. Horak, Y.J. Huang, M.D. Levine, M.J. Sale, D.K. Schmalzer, and M.J. Scott (eds.)]. Synthesis and Assessment Product 4.5. U.S. Climate Change Science Program, Washington, DC, pp. 98-108.
- ²¹¹ Maulbetsch, J.S. and M.N. DiFilippo, 2006: *Cost and Value of Water Use at Combined Cycle Power Plants*. CEC-500-2006-034. California Energy Commission, PIER Energy-Related Environmental Research. <<http://www.energy.ca.gov/2006publications/CEC-500-2006-034/CEC-500-2006-034.PDF>>
- ²¹² U.S. Energy Information Administration, 2002: *National Trends in Coal Transportation*. [Web site] Energy Information Administration, Washington, DC. <<http://www.eia.doe.gov/cneaf/coal/ctrdb/natlrends.html>>
- ²¹³ NOAA's National Climatic Data Center, 2008: *Climate of 2008: Midwestern U.S. Flood Overview*. [Web site] <<http://www.ncdc.noaa.gov/oa/climate/research/2008/flood08.html>>
- ²¹⁴ CBO Testimony, 2005: *Macroeconomic and Budgetary Effects of Hurricanes Katrina and Rita*. Statement of Douglas Holtz-Easkin, Director, before the Committee on the Budget, U.S. House of Representatives. Congressional Budget Office, Washington, DC, 21 pp. <<http://www.cbo.gov/doc.cfm?index=6684>>
- ²¹⁵ Potter, J.R., V.R. Burkett, and M.J. Savonis, 2008: Executive summary. In: *Impacts of Climate Change and Variability on Transportation Systems and Infrastructure: Gulf Coast Study, Phase I* [Savonis, M. J., V.R. Burkett, and J.R. Potter (eds.)]. Synthesis and Assessment Product 4.7. U.S. Department of Transportation, Washington, DC, pp. ES-1 to ES-10.
- ²¹⁶ U.S. Energy Information Administration, U.S. Department of Energy [data]; assembled by E. Mills, Lawrence Berkeley National Laboratory. Data available at <http://www.eia.doe.gov/cneaf/electricity/page/disturb_events.html>
- ²¹⁷ Kafalenos, R.S., K.J. Leonard, D.M. Beagan, V.R. Burkett, B.D. Keim, A. Meyers, D.T. Hunt, R.C. Hyman, M.K. Maynard, B. Fritsche, R.H. Henk, E.J. Seymour, L.E. Olson, J.R. Potter, and M.J. Savonis, 2008: What are the implications of climate change and variability for Gulf coast transportation? In: *Impacts of Climate Change and Variability on Transportation Systems and Infrastructure: Gulf Coast Study, Phase I* [Savonis, M.J., V.R. Burkett, and J.R. Potter (eds.)]. Synthesis and Assessment Product 4.7. U.S. Department of Transportation, Washington, DC, pp. 4-1 to 4F-27 [104 pp.]
- ²¹⁸ U.S. Energy Information Administration, 2008: *Monthly Energy Review*. U.S. Department of Energy, Energy Information Administration, Washington, DC. <<http://www.eia.doe.gov/emeu/mer/contents.html>>
- ²¹⁹ National Assessment Synthesis Team (NAST), 2001: *Climate Change Impacts on the United States: The Potential Consequences of Climate Variability and Change*. Cambridge University Press, Cambridge, UK, and New York, 612 pp. <<http://www.usgcrp.gov/usgcrp/Library/nationalassessment/>>
- ²²⁰ ACIA, 2004: *Impacts of a Warming Arctic: Arctic Climate Impact Assessment*. Cambridge University Press, Cambridge, UK, and New York, 139 pp. <<http://www.acia.uaf.edu>>
- ²²¹ U.S. Environmental Protection Agency, 2008: *Inventory of U.S. Greenhouse Gas Emissions and Sinks: 1990-2006*. USEPA 430-R-08-005. U.S. Environmental Protection Agency, Washington, DC, 473 pp. <<http://www.epa.gov/climatechange/emissions/usinventoryreport.html>>
- ²²² National Research Council, 2008: *Potential Impacts of Climate Change on U.S. Transportation*. Transportation Research Board special report 290. Transportation Research Board, Washington, DC, 280 pp. <<http://onlinepubs.trb.org/onlinepubs/sr/sr290.pdf>>
- ²²³ Schimel, D., A. Janetos, P. Backlund, J. Hatfield, M.G. Ryan, S.R. Archer, and D. Lettenmaier, 2008: Synthesis. In: *The Effects of Climate Change on Agriculture, Land Resources, Water Resources, and Biodiversity in the United States* [Backlund, P., A. Janetos, D. Schimel, J. Hatfield, K. Boote, P. Fay, L. Hahn, C. Izaurralde, B.A. Kimball, T. Mader, J. Morgan, D. Ort, W. Polley, A. Thomson, D. Wolfe, M.G. Ryan, S.R. Archer, R. Birdsey, C. Dahm, L. Heath, J. Hicke, D. Hollinger, T. Huxman, G. Okin, R. Oren, J. Randerson, W. Schlesinger, D. Lettenmaier, D. Major, L. Poff, S. Running, L. Hansen, D. Inouye, B.P. Kelly, L. Meyerson, B. Peterson, and R. Shaw (eds.)]. Synthesis and Assessment Product 4.3. U.S. Department of Agriculture, Washington, DC, pp. 183-193.
- ²²⁴ Burkett, V.R., R.C. Hyman, R. Hagelman, S.B. Hartley, M. Sheppard, T.W. Doyle, D.M. Beagan, A. Meyers, D.T. Hunt, M.K. Maynard, R.H. Henk, E.J. Seymour, L.E. Olson, J.R. Potter, and N.N. Srinivasan, 2008: Why study the Gulf Coast? In: *Impacts of Climate Change and Variability on Transportation Systems and Infrastructure: Gulf Coast Study, Phase I* [Savonis, M.J., V.R. Burkett, and J.R. Potter (eds.)]. Synthesis and Assessment Product

- 4.7. U.S. Department of Transportation, Washington, DC, pp. 2-1 to 2F-26. [66 pp.]
- ²²⁵ Peterson, T.C., M. McGuirk, T.G. Houston, A.H. Horvitz, and M.F. Wehner, 2008: *Climate Variability and Change with Implications for Transportation*. National Research Council, Washington, DC, 90 pp. <<http://onlinepubs.trb.org/onlinepubs/sr/sr290Many.pdf>>
- ²²⁶ Hay, J.E., R. Warrick, C. Cheatham, T. Manarangi-Trott, J. Konno, and P. Hartley, 2005: *Climate Proofing: A Risk-based Approach to Adaptation*. Asian Development Bank, Manila, The Philippines, 191 pp. <<http://www.adb.org/Documents/Reports/Climate-Proofing/default.asp>>
- ²²⁷ OSHA, 2008: Heat stress. In: *OSHA Technical Manual, Section III: Chapter 4. Occupational Safety & Health Administration*, Washington, DC. <http://www.osha.gov/dts/osta/otm/otm_iii/otm_iii_4.html>
- ²²⁸ Landsea, C.W., 1993: A climatology of intense (or major) Atlantic hurricanes. *Monthly Weather Review*, **121**(6), 1710-1713.
- ²²⁹ NOAA's National Climatic Data Center, 2008: *Billion Dollar U.S. Weather Disasters* [Web site] <<http://www.ncdc.noaa.gov/oa/reports/billionz.html>>
- ²³⁰ Larsen, P.H., S. Goldsmith, O. Smith, M.L. Wilson, K. Strzepek, P. Chinowsky, and B. Saylor, 2008: Estimating future costs for Alaska public infrastructure at risk from climate change. *Global Environmental Change*, **18**(3), 442-457.
- ²³¹ U.S. Environmental Protection Agency, 2008: *Inventory of U.S. Greenhouse Gas Emissions and Sinks: 1990-2006*. USEPA 430-R-08-005. U.S. Environmental Protection Agency, Washington, DC, 473 pp. <<http://www.epa.gov/climatechange/emissions/usinventoryreport.html>>
- ²³² National Agricultural Statistics Service, 2002: *2002 Census of Agriculture*. USDA National Agricultural Statistics Service, Washington, DC. <<http://www.agcensus.usda.gov/Publications/2002/index.asp>>
- ²³³ Wolfe, W., L. Ziska, C. Petzoldt, A. Seaman, L. Chase, and K. Hayhoe, 2007: Projected change in climate thresholds in the northeastern U.S.: implications for crops, pests, livestock, and farmers. *Mitigation and Adaptation Strategies for Global Change*, **13**(5-6), 555-575.
- ²³⁴ Frumhoff, P.C., J.J. McCarthy, J.M. Melillo, S.C. Moser, and D.J. Wuebbles, 2007: *Confronting Climate Change in the U.S. Northeast: Science, Impacts and Solutions*. Synthesis report of the Northeast Climate Impacts Assessment. Union of Concerned Scientists, Cambridge, MA, 146 pp.
- ²³⁵ Gu, L., P.J. Hanson, W.M. Post, D.P. Kaiser, B. Yang, R. Nemani, S.G. Pallardy, and T. Meyers, 2008: The 2007 eastern U.S. spring freeze: increased cold damage in a warming world? *BioScience*, **58**(3), 253-262.
- ²³⁶ Inouye, D.W., 2008: Effects of climate change on phenology, frost damage, and floral abundance of montane wildflowers. *Ecology*, **89**(2), 353-362.
- ²³⁷ Either hydrogen dioxide or oxygenated hydrocarbons.
- ²³⁸ Peet, M.M. and D.W. Wolf, 2000: Crop ecosystem responses to climate change: vegetable crops. In: *Climate Change and Global Crop Productivity* [Reddy, K.R. and H.F. Hodges (eds.)]. CABI Publishing, New York, and Wallingford, UK, 472 pp.
- ²³⁹ Bridges, D.C. (ed.), 1992: *Crop Losses Due to Weeds in the United States*. Weed Science Society of America, Champaign, IL, 403 pp.
- ²⁴⁰ Joyce, L.A., G.M. Blate, J.S. Littell, S.G. McNulty, C.I. Millar, S.C. Moser, R.P. Neilson, K. O'Halloran, and D.L. Peterson, 2008: National forests. In: *Preliminary Review of Adaptation Options for Climate-sensitive Ecosystems and Resources*. [Julius, S.H., J.M. West (eds.), J.S. Baron, B. Griffith, L.A. Joyce, P. Kareiva, B.D. Keller, M.A. Palmer, C.H. Peterson, and J.M. Scott (authors)]. Synthesis and Assessment Product 4.4. U.S. Environmental Protection Agency, Washington, DC, USA, pp. 3-1 to 3-127.
- ²⁴¹ Kiely, T., D. Donaldson, and A. Grube, 2004: *Pesticides Industry Sales and Usage: 2000 and 2001 Market Estimates*. U.S. Environmental Protection Agency, Washington DC, 33 pp. <<http://www.epa.gov/oppbead1/pestsales/>>
- ²⁴² Temperatures for the contiguous U.S. are based on data from the U.S. Historical Climatology Network Version 2 (Menne *et al.* 2008). Temperatures for Alaska, Hawaii, and Puerto Rico are based on data from the Cooperative Observers Network adjusted to remove non-climatic influences such as changes in instruments and observer practices and changes in the station environment (Menne and Williams, 2008).
U.S. time series on page 27 calculated with data for the contiguous U.S., Alaska, and Hawaii. U.S. map on page 28 lower left includes observed temperature change in Puerto Rico. Winter temperature trend map in the agriculture section, page 76, is for the contiguous U.S. only.
Menne, M.J., C.N. Williams, and R.S. Vose, 2008: The United States Historical Climatology Network Monthly Temperature Data - Version 2. *Bulletin of the American Meteorological Society*, Early online release, 25 February 2009, doi:10.1175/2008BAMS2613.1.
Menne, M.J. and C.N. Williams Jr., 2008: Homogenization of temperature series via pairwise comparisons. *Journal of Climate*, **22**(7), 1700-1717.
- ²⁴³ Natural Resources Conservation Service, 1997: *1997 Five-Year Natural Resources Inventory*. USDA Natural Resources Conservation Service, Washington, DC. <<http://www.nrcs.usda.gov/technical/NRI/1997/index.html>>
- ²⁴⁴ Ryan, M.G., S.R. Archer, R. Birdsey, C. Dahm, L. Heath, J. Hicke, D. Hollinger, T. Huxman, G. Okin, R. Oren, J. Randerson, and W. Schlesinger, 2008: Land resources. In: *The Effects of Climate Change on Agriculture, Land Resources, Water Resources, and Biodiversity in the United States* [Backlund, P., A. Janetos, D. Schimel, J. Hatfield, K. Boote, P. Fay, L. Hahn, C. Izaurralde, B.A. Kimball, T. Mader, J. Morgan, D. Ort, W. Polley, A. Thomson, D. Wolfe, M.G. Ryan, S.R. Archer, R. Birdsey, C. Dahm, L. Heath, J. Hicke, D. Hollinger, T. Huxman, G. Okin, R. Oren, J. Randerson, W. Schlesinger, D. Lettenmaier, D. Major, L. Poff, S. Running, L. Hansen, D. Inouye, B.P. Kelly, L. Meyerson, B. Peterson, and R. Shaw (eds.)]. Synthesis and Assessment Product 4.3. U.S. Department of Agriculture, Washington, DC, pp. 75-120.
- ²⁴⁵ Parmesan, C., 1996: Climate and species range. *Nature*, **382**(6594), 765-766.
- ²⁴⁶ Hinzman, L.D., N.D. Bettez, W.R. Bolton, F.S. Chapin, M.B. Dyurgerov, C.L. Fastie, B. Griffith, R.D. Hollister, A. Hope, H.P. Huntington, A.M. Jensen, G.J. Jia, T. Jorgenson, D.L. Kane, D.R. Klein, G. Kofinas, A.H. Lynch, A.H. Lloyd, A.D. McGuire, F.E. Nelson, M. Nolan, W.C. Oechel, T.E. Osterkamp, C.H. Racine, V.E. Romanovsky, R.S. Stone, D.A. Stow, M. Sturm, C.E. Tweedie, G.L. Vourlitis, M.D. Walker, D.A. Walker, P.J. Webber, J.M. Welker, K.S. Winker, and K. Yoshikawa, 2005: Evidence and implications of recent climate change in northern Alaska and other Arctic regions. *Climatic Change*, **72**(3), 251-298.
- ²⁴⁷ Mimura, N., L. Nurse, R.F. McLean, J. Agard, L. Briguglio, P. Lefale, R. Payet, and G. Sem, 2007: Small islands. In: *Climate Change 2007: Impacts, Adaptation and Vulnerability*. Contribution of Working Group II to the Fourth Assessment Report of the Intergovernmental Panel on Climate Change [Parry, M.L., O.F. Canziani, J.P. Palutikof, P.J. van der Linden, and C.E. Hanson (eds.)]. Cambridge University Press, Cambridge, UK, and New York, pp. 687-716.

- ²⁴⁸ Millennium Ecosystem Assessment, 2005: *Ecosystems and Human Well-being: Biodiversity Synthesis*. World Resources Institute, Washington, DC, 86 pp. <<http://www.millenniumassessment.org>>
- ²⁴⁹ Fischlin, A., G.F. Midgley, J.T. Price, R. Leemans, B. Gopal, C. Turley, M.D.A. Rounsevell, O.P. Dube, J. Tarazona, and A.A. Velichko, 2007: Ecosystems, their properties, goods, and services. In: *Climate Change 2007: Impacts, Adaptation and Vulnerability*. Contribution of Working Group II to the Fourth Assessment Report of the Intergovernmental Panel on Climate Change [Parry, M.L., O.F. Canziani, J.P. Palutikof, P.J. van der Linden and C.E. Hanson (eds.)]. Cambridge University Press, Cambridge, UK, and New York, pp. 211-272.
- ²⁵⁰ National Interagency Fire Center, [2008]: *Total Wildland Fires and Acres (1960-2007)*. National Interagency Coordination Center, Boise, ID. Data at <http://www.nifc.gov/fire_info/fire_stats.htm>
- ²⁵¹ CCSP, 2009: Case studies. In: *Thresholds of Climate Change in Ecosystems* [Fagre, D.B., C.W. Charles, C.D. Allen, C. Birkeland, F.S. Chapin III, P.M. Groffman, G.R. Guntenspergen, A.K. Knapp, A.D. McGuire, P.J. Mulholland, D.P.C. Peters, D.D. Roby, and G. Sugihara] Synthesis and Assessment Product 4.2. U.S. Geological Survey, Reston, VA, pp. 32-73.
- ²⁵² Pounds, J.A., M.R. Bustamante, L.A. Coloma, J.A. Consuegra, M.P. L. Fogden, P.N. Foster, E. La Marca, K.L. Masters, A. Merino-Viteri, R. Puschendorf, S.R. Ron, G.A. Sánchez-Azofeifa, C.J. Still, and B.E. Young, 2006: Widespread amphibian extinctions from epidemic disease driven by global warming. *Nature*, **439(7073)**, 161-167.
- ²⁵³ Peterson, T.C., D.M. Anderson, S.J. Cohen, M. Cortez-Vázquez, R.J. Murnane, C. Parmesan, D. Phillips, R.S. Pulwarty, and J.M.R. Stone, 2008: Why weather and climate extremes matter. In: *Weather and Climate Extremes in a Changing Climate. Regions of Focus: North America, Hawaii, Caribbean, and U.S. Pacific Islands* [Karl, T.R., G.A. Meehl, C.D. Miller, S.J. Hassol, A.M. Waple, and W.L. Murray (eds.)]. Synthesis and Assessment Product 3.3. U.S. Climate Change Science Program, Washington, DC, pp. 11-34.
- ²⁵⁴ Backlund, P., D. Schimel, A. Janetos, J. Hatfield, M.G. Ryan, S.R. Archer, and D. Lettenmaier, 2008: Introduction. In: *The Effects of Climate Change on Agriculture, Land Resources, Water Resources, and Biodiversity in the United States* [Backlund, P., A. Janetos, D. Schimel, J. Hatfield, K. Boote, P. Fay, L. Hahn, C. Izaurrealde, B.A. Kimball, T. Mader, J. Morgan, D. Ort, W. Polley, A. Thomson, D. Wolfe, M.G. Ryan, S.R. Archer, R. Birdsey, C. Dahm, L. Heath, J. Hicke, D. Hollinger, T. Huxman, G. Okin, R. Oren, J. Randerson, W. Schlesinger, D. Lettenmaier, D. Major, L. Poff, S. Running, L. Hansen, D. Inouye, B.P. Kelly, L. Meyerson, B. Peterson, and R. Shaw (eds.)]. Synthesis and Assessment Product 4.3. U.S. Department of Agriculture, Washington, DC, pp. 11-20.
- ²⁵⁵ Burkett, V.R., D.A. Wilcox, R. Stottlemeyer, W. Barrow, D. Fagre, J. Baron, J. Price, J.L. Nielsen, C.D. Allen, D.L. Peterson, G. Ruggerone, and T. Doyle, 2005: Nonlinear dynamics in ecosystem response to climatic change: case studies and policy implications. *Ecological Complexity*, **2(4)**, 357-394.
- ²⁵⁶ Anderson, K.E., D.R. Cahoon, S.K. Gill, B.T. Gutierrez, E.R. Thieler, J.G. Titus, and S.J. Williams, 2009: Executive summary. In: *Coastal Elevations and Sensitivity to Sea-level Rise: A Focus on the Mid-Atlantic Region* [J.G. Titus (coordinating lead author), K.E. Anderson, D.R. Cahoon, D.B. Gesch, S.K. Gill, B.T. Gutierrez, E.R. Thieler, and S.J. Williams (lead authors)]. Synthesis and Assessment Product 4.1. U.S. Environmental Protection Agency, Washington, DC, pp. 1-8.
- ²⁵⁷ Park, R.A., M.S. Trehan, P.W. Mousel, and R.C. Howe, 1989: *The Effects of Sea Level Rise on U.S. Coastal Wetlands*. U.S. Environmental Protection Agency, Washington, DC, 55 pp.
- ²⁵⁸ Gutierrez, B.T., S.J. Williams, and E.R. Thieler, 2009: Ocean coasts. In: *Coastal Elevations and Sensitivity to Sea-level Rise: A Focus on the Mid-Atlantic Region* [J.G. Titus (coordinating lead author), K.E. Anderson, D.R. Cahoon, D.B. Gesch, S.K. Gill, B.T. Gutierrez, E.R. Thieler, and S.J. Williams (lead authors)]. Synthesis and Assessment Product 4.1. U.S. Environmental Protection Agency, Washington, DC, pp. 43-56.
- ²⁵⁹ Monaco Declaration, 2009: Developed at the *Second International Symposium on the Ocean in a High-CO₂ World*, Monaco, 6-9 October 2008. <<http://ioc3.unesco.org/oanet/Symposium2008/MonacoDeclaration.pdf>>
- ²⁶⁰ Feely, R.A., C.L. Sabine, J.M. Hernandez-Ayon, D. Ianson, and B. Hales, 2008: Evidence for upwelling of corrosive "acidified" water onto the continental shelf. *Science*, **320(5882)**, 1490-1492.
- ²⁶¹ Beckage, B., B. Osborne, D.G. Gavin, C. Pucko, T. Siccama, and T. Perkins, 2008: A rapid upward shift of a forest ecotone during 40 years of warming in the Green Mountains of Vermont. *Proceedings of the National Academy of Sciences*, **105(11)**, 4197-4202.
- ²⁶² Beever, E.A., P.F. Brussard, and J. Berger, 2003: Patterns of apparent extirpation among isolated populations of pikas (*Ochotona princeps*) in the Great Basin. *Journal of Mammology*, **84(1)**, 37-54.
- ²⁶³ Krajick, K., 2004: All downhill from here? *Science*, **303(5664)**, 1600-1602.
- ²⁶⁴ Battin, J., M.W. Wiley, M.H. Ruckelshaus, R.N. Palmer, E. Korb, K.K. Bartz, and H. Imaki, 2007: Projected impacts of climate change on salmon habitat restoration. *Proceedings of the National Academy of Sciences*, **104(16)**, 6720-6725.
- ²⁶⁵ Williams, J.E., A.L. Haak, N.G. Gillespie, H.M. Neville, and W.T. Colyer, 2007: *Healing Troubled Waters: Preparing Trout and Salmon Habitat for a Changing Climate*. Trout Unlimited, Arlington, VA, 12 pp. <<http://www.tu.org/climatechange>>
- ²⁶⁶ Daily, G.C., T. Soderqvist, S. Aniyar, K. Arrow, P. Dasgupta, P.R. Ehrlich, C. Folke, A. Jansson, B.O. Jansson, N. Kautsky, S. Levin, J. Lubchenco, K.G. Maler, D. Simpson, D. Starrett, D. Tilman, and B. Walker, 2000: Ecology - The value of nature and the nature of value. *Science*, **289(5478)**, 395-396.
- ²⁶⁷ Hamilton, J.M. and R.S.J. Tol, 2004: The impact of climate change on tourism and recreation. In: *Human-Induced Climate Change - An Interdisciplinary Assessment* [Schlesinger, M., H.S. Khesghi, J. Smith, F.C. de la Chesnaye, J.M. Reilly, T. Wilson, and C. Kolstad (eds.)]. Cambridge University Press, Cambridge (UK), pp. 147-155.
- ²⁶⁸ Cordell, H.K., B. McDonald, R.J. Teasley, J.C. Bergstrom, J. Martin, J. Bason, and V.R. Leeworthy, 1999: Outdoor recreation participation trends. In: *Outdoor Recreation in American Life: A National Assessment of Demand and Supply Trends* [Cordell, H.K. and S.M. McKinney (eds.)]. Sagamore Publishing, Champaign, IL, pp. 219-321.
- ²⁶⁹ Wall, G., 1998: Implications of global climate change for tourism and recreation in wetland areas. *Climatic Change*, **40(2)**, 371-389.
- ²⁷⁰ Titus, J.G., M. Craghan, S.K. Gill, and S.J. Williams, 2009: Shore protection and retreat. In: *Coastal Elevations and Sensitivity to Sea-level Rise: A Focus on the Mid-Atlantic Region* [J.G. Titus (coordinating lead author), K.E. Anderson, D.R. Cahoon, D.B. Gesch, S.K. Gill, B.T. Gutierrez, E.R. Thieler, and S.J. Williams (lead authors)]. Synthesis and Assessment Product 4.1. U.S. Environmental Protection Agency, Washington, DC, pp. 87-104.
- ²⁷¹ Titus, J.G. and J.E. Neumann, 2009: Implications for decisions. In: *Coastal Elevations and Sensitivity to Sea-level Rise: A Focus on the Mid-Atlantic Region* [J.G. Titus (coordinating lead author), K.E. Anderson, D.R. Cahoon, D.B. Gesch, S.K. Gill, B.T. Gutierrez, E.R. Thieler, and S.J. Williams (lead authors)]. Synthesis and Assessment Product 4.1. U.S. Environmental Protection Agency, Washington, DC, pp. 105-114.

- sis and Assessment Product 4.1. U.S. Environmental Protection Agency, Washington, DC, pp. 141-156.
- ²⁷² Confalonieri, U., B. Menne, R. Akhtar, K.L. Ebi, M. Hauengue, R.S. Kovats, B. Revich, and A. Woodward, 2007: Human health. In: *Climate Change 2007: Impacts, Adaptation and Vulnerability*. Contribution of Working Group II to the Fourth Assessment Report of the Intergovernmental Panel on Climate Change [Parry, M.L., O.F. Canziani, J.P. Palutikof, P.J. van der Linden, and C.E. Hanson (eds.)]. Cambridge University Press, Cambridge, UK, and New York, pp. 391-431.
- ²⁷³ Borden, K.A. and S.L. Cutter, 2008: Spatial patterns of natural hazards mortality in the United States. *International Journal of Health Geographics*, **7**:64, doi:10.1186/1476-072X-7-64.
- ²⁷⁴ Gamble, J.L., K.L. Ebi, A. Grambsch, F.G. Sussman, T.J. Wilbanks, C.E. Reid, K. Hayhoe, J.V. Thomas, and C.P. Weaver, 2008: Introduction. In: *Analyses of the Effects of Global Change on Human Health and Welfare and Human Systems* [Gamble, J.L. (ed.), K.L. Ebi, F.G. Sussman, and T.J. Wilbanks (authors)]. Synthesis and Assessment Product 4.6. U.S. Environmental Protection Agency, Washington, DC, pp. 1-1 to 1-41.
- ²⁷⁵ Brikowski, T.H., Y. Lotan, and M.S. Pearle, 2008: Climate-related increase in the prevalence of urolithiasis in the United States. *Proceedings of the National Academy of Sciences*, **105**(28), 9841-9846.
- ²⁷⁶ Semenza, J.C., 1999: Acute renal failure during heat waves. *American Journal of Preventive Medicine*, **17**(1), 97.
- ²⁷⁷ Luber, G.E. and L.M. Conklin, 2006: Heat-related deaths: United States, 1999–2003. *Morbidity and Mortality Weekly Report*, **55**(29), 796-798.
- ²⁷⁸ Zanobetti, A. and J. Schwartz, 2008: Temperature and mortality in nine US cities. *Epidemiology*, **19**(4), 563-570.
- ²⁷⁹ Davis, R.E., P.C. Knappenberger, P.J. Michaels, and W.M. Novicoff, 2003: Changing heat-related mortality in the United States. *Environmental Health Perspectives*, **111**(14), 1712-1718.
- ²⁸⁰ U.S. Energy Information Administration, 2005: Table HC10.6: Air conditioning characteristics by U.S. census region, 2005. In: *2005 Residential Energy Consumption Survey*. Data available at <http://www.eia.doe.gov/emeu/recs/recs2005/hc2005_tables/detailed_tables2005.html>
- ²⁸¹ U.S. Energy Information Administration, 2005: Table 2: Type of air-conditioning equipment by census region and survey year. In: *2005 Residential Energy Consumption Survey*. Data available at <http://www.eia.doe.gov/emeu/consumptionbriefs/recs/actrends/recs_actrends_table2.html>
- ²⁸² Sheridan, S.C., A.J. Kalkstein, and L.S. Kalkstein, 2008: Trends in heat-related mortality in the United States, 1975–2004. *Natural Hazards*, Online First™, doi:10.1007/s11069-008-9327-2.
- ²⁸³ Hayhoe, K., K. Cherkauer, N. Schlegal, J. VanDorn, S. Vavrus, and D. Wuebbles, 2009: Regional climate change projections for Chicago and the Great Lakes. *Journal of Great Lakes Research*, in press.
- ²⁸⁴ Hayhoe, K., D. Cayan, C.B. Field, P.C. Frumhoff, E.P. Maurer, N.L. Miller, S.C. Moser, S.H. Schneider, K.N. Cahill, E.E. Cleland, L. Dale, R. Drapek, R.M. Hanemann, L.S. Kalkstein, J. Lenihan, C.K. Lunch, R.P. Neilson, S.C. Sheridan, and J.H. Verville, 2004: Emissions pathways, climate change, and impacts on California. *Proceedings of the National Academy of Sciences*, **101**(34), 12422-12427.
- ²⁸⁵ Vavrus, S. and J. van Dorn, 2008: Projected future temperature and precipitation extremes in Chicago. *Journal of Great Lakes Research*, in press.
- ²⁸⁶ Lemmen, D.S. and F.J. Warren (eds.), 2004: *Climate Change Impacts and Adaptation: A Canadian Perspective*. Climate Change Impacts and Adaptation Program, Natural Resources Canada, Ottawa, ON, 174 pp. <http://adaptation.nrcan.gc.ca/perspective/index_e.php>
- ²⁸⁷ Ebi, K.L., J. Smith, I. Burton, and J. Scheraga, 2006: Some lessons learned from public health on the process of adaptation. *Mitigation and Adaptation Strategies for Global Change*, **11**(3), 607-620.
- ²⁸⁸ Ebi, K.L., T.J. Teisberg, L.S. Kalkstein, L. Robinson, and R.F. Weiher, 2004: Heat watch/warning systems save lives: estimated costs and benefits for Philadelphia 1995-98. *Bulletin of the American Meteorological Society*, **85**(8), 1067-1073.
- ²⁸⁹ Medina-Ramon, M. and J. Schwartz, 2007: Temperature, temperature extremes, and mortality: a study of acclimatisation and effect modification in 50 U.S. cities. *Occupational and Environmental Medicine*, **64**(12), 827-833.
- ²⁹⁰ U.S. Environmental Protection Agency, 2008: *National Air Quality: Status and Trends through 2007*. U.S. EPA Air Quality Assessment Division, Research Triangle Park, NC, 48 pp. <<http://www.epa.gov/airtrends/2008/index.html>>
- ²⁹¹ Tao, Z., A. Williams, H.-C. Huang, M. Caughey, and X.-Z. Liang, 2007: Sensitivity of U.S. surface ozone to future emissions and climate changes. *Geophysical Research Letters*, **34**, L08811, doi:10.1029/2007GL029455.
- ²⁹² California Air Resources Board, 2007: *Recent Research Findings: Health Effects of Particulate Matter and Ozone Air Pollution*. [Online fact sheet] California Air Resources Board, [Sacramento], 7 pp. <http://www.arb.ca.gov/research/health/fs/pm_ozone-fs.pdf>
- ²⁹³ California Climate Action Team, 2006: *Climate Action Team Report to Governor Schwarzenegger and the Legislature*. California Environmental Protection Agency, Sacramento, 107 pp. <http://www.climatechange.ca.gov/climate_action_team/reports/index.html>
- ²⁹⁴ Westerling A.L., H.G. Hidalgo, D.R. Cayan, and T.W. Swetnam, 2006: Warming and earlier spring increase western U.S. forest wildfire activity. *Science*, **313**(5789), 940-943.
- ²⁹⁵ Rhode Island Public Transit Authority, Undated: *Air Quality Alert Days Program*. [Web site] <<http://www.ripta.com/content259.html>>
- ²⁹⁶ Friedman, M.S., K.E. Powell, L. Hutwagner, L.M. Graham, and W.G. Teague, 2001: Impact of changes in transportation and commuting behaviors during the 1996 Summer Olympic Games in Atlanta on air quality and childhood asthma. *Journal of the American Medical Association*, **285**(7), 897-905.
- ²⁹⁷ Wuebbles, D.J., H. Lei, and J.-T. Lin, 2007: Intercontinental transport of aerosols and photochemical oxidants from Asia and its consequences. *Environmental Pollution*, **150**, 65-84.
- ²⁹⁸ Bell, M.L., R. Goldberg, C. Hogrefe, P. Kinney, K. Knowlton, B. Lynn, J. Rosenthal, C. Rosenzweig, and J.A. Patz, 2007: Climate change, ambient ozone, and health in 50 U.S. cities. *Climatic Change*, **82**(1-2), 61-76.
- ²⁹⁹ U.S. Environmental Protection Agency, 2008: *National Pollutant Discharge Elimination System (NPDES) Combined Sewer Overflows Demographics*. [Web site] <http://cfpub.epa.gov/npdes/cso/demo.cfm?program_id=5>
- ³⁰⁰ Tibbetts, J., 2005: Combined sewer systems: down, dirty, and out of date. *Environmental Health Perspectives*, **113**(7), A464-A467.
- ³⁰¹ U.S. Environmental Protection Agency, 2008: *Clean Watersheds Needs Survey 2004: Report to Congress*. U.S. Environmental Protection Agency, Washington, DC. <<http://www.epa.gov/cwns/2004rtc/cwns2004rtc.pdf>>
- ³⁰² Patz, J.A., S.J. Vavrus, C.K. Uejio, and S.L. McClellan, 2008: Climate change and waterborne disease risk in the Great Lakes region of the U.S. *American Journal of Preventive Medicine*, **35**(5), 451-458.

- ³⁰³ Centers for Disease Control and Prevention, 2009: *West Nile Virus: Statistics, Surveillance, and Control*. [Web site] <<http://www.cdc.gov/ncidod/dvbid/westnile/surv&control.htm>>
- ³⁰⁴ Kilpatrick, A.M., M.A. Meola, R.M. Moudy, and L.D. Kramer, 2008: Temperature, viral genetics, and the transmission of West Nile virus by *Culex pipiens* mosquitoes. *PLoS Pathogens*, **4(6)**, e1000092. doi:10.1371/journal.ppat.1000092
- ³⁰⁵ Sussman, F.G., M.L. Cropper, H. Galbraith, D. Godschalk, J. Loomis, G. Luber, M. McGeehin, J.E. Neumann, W.D. Shaw, A. Vedlitz, and S. Zahran, 2008: Effects of global change on human welfare. In: *Analyses of the Effects of Global Change on Human Health and Welfare and Human Systems* [Gamble, J.L. (ed.), K.L. Ebi, F.G. Sussman, and T.J. Wilbanks (authors)]. Synthesis and Assessment Product 4.6. U.S. Environmental Protection Agency, Washington, DC, pp. 4-1 to 4-74.
- ³⁰⁶ Smith, J.B. and D. Tirpak (eds.), 1989: *The Potential Effects of Global Climate Change on the United States*. U.S. Environmental Protection Agency, Washington, DC, 413 pp.
- ³⁰⁷ Ziska, L.H. and F.A. Caulfield, 2000: Rising CO₂ and pollen production of common ragweed (*Ambrosia artemisiifolia* L.), a known allergy-inducing species: implications for public health. *Australian Journal of Plant Physiology*, **27(10)**, 893-898.
- ³⁰⁸ Mohan J.E., L.H. Ziska, W.H. Schlesinger, R.B. Thomas, R.C. Sicher, K. George, and J.S. Clark, 2006: Biomass and toxicity responses of poison ivy (*Toxicodendron radicans*) to elevated atmospheric CO₂. *Proceedings of the National Academy of Sciences*, **103(24)**, 9086-9089.
- ³⁰⁹ Hunt, R., D.W. Hand, M.A. Hannah, and A.M. Neal, 1991: Response to CO₂ enrichment in 27 herbaceous species. *Functional Ecology*, **5(3)**, 410-421.
- ³¹⁰ Ziska, L.H., 2003: Evaluation of the growth response of six invasive species to past, present and future atmospheric carbon dioxide. *Journal of Experimental Botany*, **54(381)**, 395-404.
- ³¹¹ CDC, 2007: *2007 National Diabetes Fact Sheet*. Centers for Disease Control and Prevention, Atlanta, GA, 14 pp. <<http://www.cdc.gov/diabetes/pubs/factsheet07.htm>>
- ³¹² Patz, J.A., H.K. Gibbs, J.A. Foley, J.V. Rogers, and K.R. Smith, 2007: Climate change and global health: quantifying a growing ethical crisis. *EcoHealth*, **4(4)**, 397-405.
- ³¹³ Wilbanks, T.J., P. Kirshen, D. Quattrocchi, P. Romero-Lankao, C. Rosenzweig, M. Ruth, W. Solecki, J. Tarr, P. Larsen, and B. Stone, 2008: Effects of global change on human settlements. In: *Analyses of the Effects of Global Change on Human Health and Welfare and Human Systems* [Gamble, J.L. (ed.), K.L. Ebi, F.G. Sussman, and T.J. Wilbanks (authors)]. Synthesis and Assessment Product 4.6. U.S. Environmental Protection Agency, Washington, DC, pp. 3-1 to 3-31.
- ³¹⁴ U.S. Census Bureau, 2005: *Domestic Net Migration in the United States: 2000 to 2004: Population Estimates and Projections*. <<http://www.census.gov/prod/2006pubs/p25-1135.pdf>>
- ³¹⁵ U.S. General Accounting Office, 2003: *Alaska Native Villages: Most Are Affected by Flooding and Erosion, but Few Qualify for Federal Assistance*. GAO-04-142. U.S. General Accounting Office, Washington, DC, 82 pp. <<http://purl.access.gpo.gov/GPO/LPS42077>>
- ³¹⁶ Cordova, R., M. Gelobter, A. Hoerner, J. Love, A. Miller, C. Saenger, and D. Zaidi, 2006: *Climate Change in California: Health, Economic and Equity Impacts*. Redefining Progress, Oakland, CA, 119 pp. <http://www.rprogress.org/publications/2006/CARB_Full_0306.pdf>
- ³¹⁷ Congressional Black Caucus Foundation, 2004: *African Americans and Climate Change: An Unequal Burden*. Redefining Progress, Washington, DC. <http://www.rprogress.org/publications/2004/CBCF_REPORT_F.pdf>
- ³¹⁸ Hoerner, J.A. and N. Robinson, 2008: *A Climate of Change: African Americans, Global Warming, and a Just Climate Policy for the U.S.* A report of the Environmental Justice and Climate Initiative. Redefining Progress, Oakland, CA, 69 pp. <<http://www.rprogress.org/publications/2008/climateofchange.pdf>>
- ³¹⁹ Borden, K.A., M.C. Schmidlein, C.T. Emrich, W.W. Piegorsch, and S.L. Cutter, 2007: Vulnerability of U.S. cities to environmental hazards. *Journal of Homeland Security and Emergency Management*, **4(2)**, article 5. <<http://www.bepress.com/jhsem/vol4/iss2/5>>
- ³²⁰ van Kamp, I., K. Leidelmeijer, G. Marsman, and A. de Hollander, 2003: Urban environmental quality and human well-being: towards a conceptual framework and demarcation of concepts; a literature study. *Landscape and Urban Planning*, **65(1-2)**, 5-18.
- ³²¹ Grimmond, S., 2007: Urbanization and global environmental change: local effects of urban warming. *Geographical Journal*, **173(1)**, 83-88.
- ³²² Anderson, C.A., 2001: Heat and violence. *Current Directions in Psychological Science*, **10(1)**, 33-38.
- ³²³ Anderson, C.A., B.J. Bushman, and R.W. Groom, 1997: Hot years and serious and deadly assault: empirical test of the heat hypothesis. *Journal of Personality and Social Psychology*, **73(6)**, 1213-1223.
- ³²⁴ Milly, P.C.D., R.T. Wetherald, K.A. Dunne, and T.L. Delworth, 2002: Increasing risk of great floods in a changing climate. *Nature*, **415(6871)**, 514-517.
- ³²⁵ Miller, N.L., K. Hayhoe, J. Jin, and M. Auffhammer, 2008: Climate, extreme heat, and electricity demand in California. *Journal of Applied Meteorology and Climatology*, **47(6)**, 1834-1844.
- ³²⁶ Wang, J.X.L. and J.K. Angell, 1999: *Air Stagnation Climatology for the United States (1948-1998)*. NOAA/Air Resources Laboratory atlas no.1. NOAA Air Resources Laboratory, Silver Spring, MD, 74 pp.
- ³²⁷ Riebsame, W.E., S.A. Changnon Jr., and T.R. Karl, 1991: *Drought and Natural Resources Management in the United States: Impacts and Implications of the 1987-89 Drought*. Westview Press, Boulder, CO, 174 pp.
- ³²⁸ NOAA's National Climatic Data Center, 2009: *NCDC Climate Monitoring: U.S. Records* [Web site] <<http://www.ncdc.noaa.gov/oa/climate/research/records/>>
- ³²⁹ NOAA's National Climatic Data Center, 2009: *U.S. Percent Area Wet or Dry*. [Web site] <http://www.ncdc.noaa.gov/img/climate/research/2008/ann/Reg110_wet-dry_bar01001208-mod_pg.gif>
- ³³⁰ Leung L.R. and W.I. Gustafson Jr., 2005: Potential regional climate change and implications to U.S. air quality. *Geophysical Research Letters*, **32**, L16711, doi:10.1029/2005GL022911.
- ³³¹ Several U.S. Department of Energy reports document the increase in electricity demand to provide air-conditioning in hotter summers in many regions of the country: Chapter 2: Carbon dioxide emissions. In: *Emissions of Greenhouse Gases in the United States 2002*. Report released 2004. <<http://www.eia.doe.gov/oiaf/1605/archive/gg03rpt/index.html>>; *South Atlantic Household Electricity Report*. Release date: 2006. <http://www.eia.doe.gov/emeu/reps/enduse/er01_so-atl.html>; *South Central Appliance Report 2001*. Updated 2005. <http://www.eia.doe.gov/emeu/reps/appli/w_s_c.html>
- ³³² CCSP, 2007: *Effects of Climate Change on Energy Production and Use in the United States*. [Wilbanks, T.J., V. Bhatt, D.E. Bilello, S.R. Bull, J. Ekmann, W.C. Horak, Y.J. Huang, M.D. Levine, M.J. Sale, D.K. Schmalzer, and M.J. Scott (eds.)]. Synthesis and Assessment Product 4.5. Department of Energy, Office of Biological & Environmental Research, Washington, DC, 160 pp. See chapters 2 and 3.

- ³³³ Daily data were used for both air stagnation and heat waves:
1. Heat waves:
 - The GHCN-Daily dataset from NCDC was used <<http://www.ncdc.noaa.gov/oa/climate/ghcn-daily/>>
 - Data from 979 U.S. stations having long periods of record and high quality.
 - At each station, a day was considered hot if the maximum temperature for that day was at or above the 90% of daily maximum temperatures at that station.
 2. Air stagnation:
 - For each day in summer and at each air-stagnation grid point, it was determined if that location had stagnant air:
 - The stagnation index was formulated by Wang, J.X.L. and J.K. Angell, 1999: *Air Stagnation Climatology for the United States (1948-1998)*. NOAA/Air Resources Laboratory atlas no.1 NOAA Air Resources Laboratory, Silver Spring, MD, 74 pp. <<http://www.arl.noaa.gov/documents/reports/atlas.pdf>>
 - Operational implementation of this index is described at <<http://www.ncdc.noaa.gov/oa/climate/research/stagnation/index.php>>
 - Note:* Although Wang and Angell used a criteria of four day stagnation periods, single stagnation days were used for this analysis.
 3. For each location in the air stagnation grid, the nearest station (of the aforementioned 979 U.S. stations) was used to determine the coincidence of summer days having stagnant air and excessive heat as a percentage of the number of days having excessive heat.
- ³³⁴ Solecki, W.D. and C. Rosenzweig, 2006: Climate change and the city: observations from metropolitan New York. In: *Cities and Environmental Change* [Bai, X. (ed.)]. Yale University Press, New York.
- ³³⁵ Rosenzweig, C. and W. Solecki (eds.), 2001: *Climate Change and a Global City: The Potential Consequences of Climate Variability and Change – Metro East Coast*. Columbia Earth Institute, New York. <http://metroeast_climate.ciesin.columbia.edu/>
- ³³⁶ Kirshen, P., M. Ruth, W. Anderson, T.R. Lakshmanan, S. Chapra, W. Chudyk, L. Edgers, D. Gute, M. Sanayei, and R. Vogel, 2004: *Climate's Long-term Impacts on Metro Boston (CLIMB) Final Report*. Civil and Environmental Engineering Department, Tufts University, 165 pp. <http://www.clf.org/uploadedFiles/CLIMB_Final_Report.pdf>
- ³³⁷ Grimm, N.B., S.H. Faeth, N.E. Golubiewski, C.L. Redman, J. Wu, X. Bai, and J.M. Briggs, 2008: Global change and the ecology of cities. *Science*, **319(5864)**, 756-760.
- ³³⁸ Kuo, F.E. and W.C. Sullivan, 2001: Environment and crime in the inner city: Does vegetation reduce crime? *Environment and Behavior*, **33(3)**, 343-367.
- ³³⁹ Julius, S.H., J.M. West, G. Blate, J.S. Baron, B. Griffith, L.A. Joyce, P. Kareiva, B.D. Keller, M. Palmer, C. Peterson, and J.M. Scott, 2008: Executive summary. In: *Preliminary Review of Adaptation Options for Climate-sensitive Ecosystems and Resources* [Julius, S.H. and J.M. West (eds.), J.S. Baron, B. Griffith, L.A. Joyce, P. Kareiva, B.D. Keller, M.A. Palmer, C.H. Peterson, and J.M. Scott (authors)]. Synthesis and Assessment Product 4.4. U.S. Environmental Protection Agency, Washington, DC, pp. 1-1 to 1-6.
- ³⁴⁰ Baker, L.A., A.J. Brazel, N. Selover, C. Martin, N. McIntyre, F.R. Steiner, A. Nelson, and L. Musacchio, 2002: Urbanization and warming of Phoenix (Arizona, USA): impacts, feedbacks, and mitigation. *Urban Ecosystems*, **6(3)**, 183-203.
- ³⁴¹ LoVecchio, F., J.S. Stapczynski, J. Hill, A.F. Haffer, J.A. Skindlov, D. Engelthaler, C. Mrela, G.E. Luber, M. Straetemans, and Z. Duprey, 2005: Heat-related mortality – Arizona, 1993-2002, and United States, 1979-2002. *Morbidity and Mortality Weekly Report*, **54(25)**, 628-630.
- ³⁴² Scott, D., J. Dawson, and B. Jones, 2008: Climate change vulnerability of the US Northeast winter recreation–tourism sector. *Mitigation and Adaptation Strategies for Global Change*, **13(5-6)**, 577-596.
- ³⁴³ Bin, O., C. Dumas, B. Poulter, and J. Whitehead, 2007: *Measuring the Impacts of Climate Change on North Carolina Coastal Resources*. [Appalachian State University, Boone, NC], 91 pp. <<http://econ.appstate.edu/climate/>>
- ³⁴⁴ Mills, E., 2005: Insurance in a climate of change. *Science*, **309(5737)**, 1040-1044.
- ³⁴⁵ Adapted from U.S. Government Accountability Office, 2007: *Climate Change: Financial Risks to Federal and Private Insurers in Coming Decades are Potentially Significant*. Government Accountability Office, Washington, DC, 68 pp. <<http://purl.access.gpo.gov/GPO/LPS89701>> Data shown are not adjusted for inflation.
- ³⁴⁶ Pielke Jr., R.A., J. Gratz, C.W. Landsea, D. Collins, M. Saunders, and R. Musulin, 2008: Normalized hurricane damages in the United States: 1900-2005. *Natural Hazards Review*, **9(1)**, 29-42.
- ³⁴⁷ Rosenzweig, C., G. Casassa, D.J. Karoly, A. Imeson, C. Liu, A. Menzel, S. Rawlins, T.L. Root, B. Seguin, and P. Tryjanowski, 2007: Assessment of observed changes and responses in natural and managed systems. In: *Climate Change 2007: Impacts, Adaptation and Vulnerability*. Contribution of Working Group II to the Fourth Assessment Report of the Intergovernmental Panel on Climate Change [Parry, M.L., O.F. Canziani, J.P. Palutikof, P.J. van der Linden, and C.E. Hanson, (eds.)]. Cambridge University Press, Cambridge, UK, and New York, pp. 79-131.
- ³⁴⁸ Pielke Jr., R.A., 2005: Response to: “Attribution of Disaster Losses” by E. Mills. *Science*, **310(5754)**, 1615.
- ³⁴⁹ Mills, E., 2009: A global review of insurance industry responses to climate change. *The Geneva Papers on Risk and Insurance—Issues and Practice*, in press.
- ³⁵⁰ Meehl, G.A. and C. Tebaldi, 2004: More intense, more frequent, and longer lasting heat waves in the 21st century. *Science*, **305(5686)**, 994-997.
- ³⁵¹ Mills, E., 2006: Synergisms between climate change mitigation and adaptation: an insurance perspective. *Mitigation and Adaptation Strategies for Global Change*, **12(5)**, 809-842.
- ³⁵² U.S. Government Accountability Office, 2007 [data]; assembled by Evan Mills, Lawrence Berkeley National Laboratory.
- ³⁵³ Reeve, N. and R. Toumi, 1999: Lightning activity as an indicator of climate change. *Quarterly Journal of the Royal Meteorological Society*, **125(555)**, 893-903.
- ³⁵⁴ Price, C. and D. Rind, 1994: Possible implications of global climate change on global lightning distributions and frequencies. *Journal of Geophysical Research*, **99(D5)**, 10823-10831.
- ³⁵⁵ Ross, C., E. Mills, and S. Hecht, 2007: Limiting liability in the greenhouse: insurance risk-management in the context of global climate change. *Stanford Environmental Law Journal and Stanford Journal of International Law, Symposium on Climate Change Risk*, **26A/43A**, 251-334.
- ³⁵⁶ Nutter, F.W., 1996: Insurance and the natural sciences: partners in the public interest. *Research Review: Journal of the Society of Insurance Research*, **Fall**, 15-19.
- ³⁵⁷ Federal Emergency Management Agency, 2008: *Emergency Management Institute*. [Web site] <<http://training.fema.gov/EMIWeb/CRS/>>
- ³⁵⁸ Bernstein, L., J. Roy, K.C. Delhotal, J. Harnisch, R. Matsuhashi, L. Price, K. Tanaka, E. Worrell, F. Yamba, and Z. Fengqi, 2007: Industry. In: *Climate Change 2007: Mitigation of Climate Change*. Contribution of Working Group III to the Fourth Assessment Report of the Intergovernmental Panel on Climate Change [Metz, B., O.R. Davidson, P.R. Bosch, R. Dave, and L.A. Meyer (eds.)].

- Cambridge University Press, Cambridge, UK, and New York, pp. 447-496.
- ³⁵⁹ Hayhoe, K., C.P. Wake, B. Anderson, X.-Z. Liang, E. Maurer, J. Zhu, J. Bradbury, A. DeGaetano, A. Hertel, and D. Wuebbles, 2008: Regional climate change projections for the northeast U.S. *Mitigation and Adaptation Strategies for Global Change*, **13(5-6)**, 425-436.
- ³⁶⁰ New York City Department of Health and Mental Hygiene, 2006: Deaths associated with heat waves in 2006. In: *NYC Vital Signs: Investigation Report, Special Report*. Department of Health and Mental Hygiene, New York, 4 pp. <<http://www.nyc.gov/html/doh/downloads/pdf/survey/survey-2006heatdeaths.pdf>>
- ³⁶¹ Kunkel, K.E., H.-C. Huang, X.-Z. Liang, J.-T. Lin, D. Wuebbles, Z. Tao, A. Williams, M. Caughey, J. Zhu, and K. Hayhoe, 2008: Sensitivity of future ozone concentrations in the northeast U.S. to regional climate change. *Mitigation and Adaptation Strategies for Global Change*, **13(5-6)**, 597-606.
- ³⁶² Hauagge, R. and J.N. Cummins, 1991: Phenotypic variation of length of bud dormancy in apple cultivars and related malus species. *Journal of the American Society for Horticultural Science*, **116(1)**, 100-106.
- ³⁶³ DeMoranville, C., 2007: Personal communication from May 29, 2008. Experts at the University of Massachusetts Cranberry Station estimate cranberry chilling requirements to be around 1,200-1,400 hours, but they advise growers to seek 1,500 hours to avoid crop failure. There are 4-5 commonly grown cultivars but no low-chill varieties. Dr. Carolyn DeMoranville is the director of the UMass Cranberry Station, a research and extension center of UMass-Amherst.
- ³⁶⁴ Iverson, L., A. Prasad, and S. Matthews, 2008: Potential changes in suitable habitat for 134 tree species in the northeastern United States. *Mitigation and Adaptation Strategies for Global Change*, **13(5-6)**, 487-516.
- ³⁶⁵ U.S. Department of Agriculture (USDA) National Agriculture Statistics Service (NASS), 2002: *Statistics by State*. [Web site] <http://www.nass.usda.gov/Statistics_by_State/>
- ³⁶⁶ St. Pierre, N.R., B. Cobanov, and G. Schnitkey, 2003: Economic losses from heat stress by U.S. livestock industries. *Journal of Dairy Science*, **86(E Sup)**, E52- E77.
- ³⁶⁷ Gornitz, V., S. Couch, and E.K. Hartig, 2001: Impacts of sea level rise in the New York City metropolitan area. *Global and Planetary Change*, **32(1)**, 61-88.
- ³⁶⁸ AIR Worldwide Corporation, 2008: *The Coastline at Risk: 2008 Update to the Estimated Insured Value of U.S. Coastal Properties*. AIR Worldwide Corporation, Boston, MA, 3 pp. <<http://www.air-worldwide.com/download.aspx?c=388&id=15836>>
- ³⁶⁹ Kirshen, P., C. Watson, E. Douglas, A. Gontz, J. Lee, and Y. Tian, 2008: Coastal flooding in the northeastern United States due to climate change. *Mitigation and Adaptation Strategies for Global Change*, **13(5-6)**, 437-451.
- ³⁷⁰ Bowman, M., D. Hill, F. Buonaiuto, B. Colle, R. Flood, R. Wilson, R. Hunter, and J. Wang, 2008: Threats and responses associated with rapid climate change in metropolitan New York. In: *Sudden and Disruptive Climate Change: Exploring the Real Risks and How We Can Avoid Them*. [MacCracken, M.C., F. Moore, and J.C. Topping Jr. (eds.)]. Earthscan, London and Sterling, VA, pp. 119-142.
- ³⁷¹ Titus, J.G., 2009: Ongoing adaptation. In: *Coastal Elevations and Sensitivity to Sea-level Rise: A Focus on the Mid-Atlantic Region* [J.G. Titus (coordinating lead author), K.E. Anderson, D.R. Cahoon, D.B. Gesch, S.K. Gill, B.T. Gutierrez, E.R. Thieler, and S.J. Williams (lead authors)]. Synthesis and Assessment Product 4.1. U.S. Environmental Protection Agency, Washington, DC, pp. 157-162.
- ³⁷² International Snowmobile Manufacturers Association, 2006: *International Snowmobile Industry Facts and Figures*. [Web site] <http://www.snowmobile.org/pr_snowfacts.asp>
- ³⁷³ Northeast Climate Impact Assessment (NECIA), 2006: *Climate Change in the U.S. Northeast: A Report of the Northeast Climate Impacts Assessment*. Union of Concerned Scientists, Cambridge, MA, 35 pp.
- ³⁷⁴ Atlantic States Marine Fisheries Commission, 2005: *American Lobster*. [Web site] <<http://www.asmf.org/americanLobster.htm>>
- ³⁷⁵ Fogarty, M.J., 1995: Populations, fisheries, and management. In: *The Biology of the American Lobster Homarus americanus*. [Factor, J.R. (ed.)]. Academic Press, San Diego, CA, pp. 111-137.
- ³⁷⁶ Glenn, R.P. and T.L. Pugh, 2006: Epizootic shell disease in American lobster (*Homarus americanus*) in Massachusetts coastal waters: interactions of temperature, maturity, and intermolt duration. *Journal of Crustacean Biology*, **26(4)**, 639-645.
- ³⁷⁷ Fogarty, M., L. Incze, R. Wahle, D. Mountain, A. Robinson, A. Pershing, K. Hayhoe, A. Richards, and J. Manning, 2008: Potential climate change impacts on Atlantic cod (*Gadus morhua*) off the northeastern United States. *Mitigation and Adaptation Strategies for Global Change*, **13(5-6)**, 453-466.
- ³⁷⁸ Dutil, J.-D. and K. Brander, 2003: Comparing productivity of North Atlantic cod (*Gadus morhua*) stocks and limits to growth production. *Fisheries Oceanography*, **12(4-5)**, 502-512.
- ³⁷⁹ Drinkwater, K.F., 2005: The response of Atlantic cod (*Gadus morhua*) to future climate change. *ICES Journal of Marine Science*, **62(7)**, 1327-1337.
- ³⁸⁰ Karl, T.R. and R.W. Knight, 1998: Secular trends of precipitation amount, frequency, and intensity in the United States. *Bulletin of the American Meteorological Society*, **79(2)**, 231-241.
- ³⁸¹ Keim, B.D., 1997: Preliminary analysis of the temporal patterns of heavy rainfall across the southeastern United States. *Professional Geographer*, **49(1)**, 94-104.
- ³⁸² Observed changes in precipitation for the Southeast were calculated from the US Historical Climatology Network Version 2. See Menne, M.J., C.N. Williams, and R.S. Vose, 2008: The United States Historical Climatology Network Monthly Temperature Data - Version 2. *Bulletin of the American Meteorological Society*, Early online release, 25 February 2009, doi:10.1175/2008BAMS2613.1
- ³⁸³ Temperature:
Menne, M.J., C.N. Williams, and R.S. Vose, 2008: The United States Historical Climatology Network Monthly Temperature Data - Version 2. *Bulletin of the American Meteorological Society*, Early online release, 25 February 2009, doi:10.1175/2008BAMS2613.1
- Precipitation:
NOAA's National Climatic Data Center, 2008: *The USHCN Version 2 Serial Monthly Dataset*. [Web site] <<http://www.ncdc.noaa.gov/oa/climate/research/ushcn/>>
- ³⁸⁴ Hoyos, C.D., P.A. Agudelo, P.J. Webster, and J.A. Curry, 2006: Deconvolution of the factors contributing to the increase in global hurricane intensity. *Science*, **312(577)**, 94-97.
- ³⁸⁵ Mann, M.E. and K.A. Emanuel, 2006: Atlantic hurricane trends linked to climate change. *EOS Transactions of the American Geophysical Union*, **87(24)**, 233, 244.
- ³⁸⁶ Trenberth, K.E. and D.J. Shea, 2006: Atlantic hurricanes and natural variability in 2005. *Geophysical Research Letters*, **33**, L12704, doi:10.1029/2006GL026894.
- ³⁸⁷ Webster, P.J., G.J. Holland, J.A. Curry, and H.-R. Chang, 2005: Changes in tropical cyclone number, duration, and intensity in a warming environment. *Science*, **309(5742)**, 1844-1846.
- ³⁸⁸ Komar, P.D. and J.C. Allan, 2007: Higher waves along U.S. East Coast linked to hurricanes. *EOS Transactions of the American Geophysical Union*, **88(30)**, 301.

- ³⁸⁹ Change calculated from daily minimum temperatures from NCDC's Global Historical Climatology Network - Daily data set. <<http://www.ncdc.noaa.gov/oa/climate/ghcn-daily/>>
- ³⁹⁰ Nicholls, R.J., P.P. Wong, V.R. Burkett, J.O. Codignotto, J.E. Hay, R.F. McLean, S. Ragoonaden, and C.D. Woodroffe, 2007: Coastal systems and low-lying areas. In: *Climate Change 2007: Impacts, Adaptations and Vulnerability*. Contribution of Working Group II to the Fourth Assessment Report of the Intergovernmental Panel on Climate Change [Parry, M.L., O.F. Canziani, J.P. Palutikof, P.J. van der Linden, and C.E. Hanson (eds.)]. Cambridge University Press, Cambridge, UK, and New York, pp. 316-356.
- ³⁹¹ Boyles, S., 2008: *Heat Stress and Beef Cattle*. Ohio State University Extension Service. <<http://beef.osu.edu/library/heat.html>>
- ³⁹² Convention on Biological Diversity, 2006: *Guidance for Promoting Synergy Among Activities Addressing Biological Diversity, Desertification, Land Degradation and Climate Change*. CBD technical series 25. Secretariat of the Convention on Biological Diversity, Montreal, Canada, 43 pp. <<http://www.biodiv.org/doc/publications/cbd-ts-25.pdf>>
- ³⁹³ Burkett, V., 2008: The northern Gulf of Mexico coast: human development patterns, declining ecosystems, and escalating vulnerability to storms and sea level rise. In: *Sudden and Disruptive Climate Change: Exploring the Real Risks and How We Can Avoid Them*. [MacCracken, M.C., F. Moore, and J.C. Topping (eds.)]. Earthscan Publications, London [UK], and Sterling, VA, pp. 101-118.
- ³⁹⁴ Twilley, R.R., E. Barron, H.L. Gholz, M.A. Harwell, R.L. Miller, D.J. Reed, J.B. Rose, E. Siemann, R.G. Welzel, and R.J. Zimmerman, 2001: *Confronting Climate Change in the Gulf Coast Region: Prospects for Sustaining Our Ecological Heritage*. Union of Concerned Scientists, Cambridge, MA, and Ecological Society of America, Washington, DC, 82 pp.
- ³⁹⁵ USGS, photos and images:
Photos on the left:
 Tihansky, A.B., 2005: Before-and-after aerial photographs show coastal impacts of Hurricane Katrina. *Sound Waves*, September. <<http://soundwaves.usgs.gov/2005/09/fieldwork2.html>>
Images on the right:
 Adapted from Fauver, L., 2007: Predicting flooding and coastal hazards: USGS hydrologists and geologists team up at the National Hurricane Conference to highlight data collection. *Sounds Waves*, June. <<http://soundwaves.usgs.gov/2007/06/meetings.html>>
- ³⁹⁶ Barras, J.A., 2006: *Land Area Change in Coastal Louisiana After the 2005 Hurricanes: A Series of Three Maps*. U.S. Geological Survey open-file report 2006-1274. <<http://pubs.usgs.gov/of/2006/1274/>>
- ³⁹⁷ Main Development Region of the Atlantic Ocean is defined in: Bell, G.D., E. Blake, C.W. Landsea, S.B. Goldenberg, R. Pasch, and T. Kimberlain, 2008: Tropical cyclones: Atlantic basin. In: Chapter 4: The Tropics, of *State of the Climate in 2007* [Levinson, D.H. and J.H. Lawrimore (eds.)]. *Bulletin of the American Meteorological Society*, **89**(Supplement), S68-S71.
- ³⁹⁸ Williams, K.L., K.C. Ewel, R.P. Stumpf, F.E. Putz, and T.W. Workman, 1999: Sea-level rise and coastal forest retreat on the west coast of Florida. *Ecology*, **80**(6), 2045-2063.
- ³⁹⁹ McNulty, S.G., J.M. Vose, and W.T. Swank, 1996: Potential climate change affects on loblolly pine productivity and hydrology across the southern United States. *Ambio*, **25**(7), 449-453.
- ⁴⁰⁰ Zimmerman, R.J., T.J. Minello, and L.P. Rozas, 2002: Salt marsh linkages to productivity of penaeid shrimps and blue crabs in the northern Gulf of Mexico. In: *Concepts and Controversies in Tidal Marsh Ecology* [Weinstein, M.P. and D.A. Kreeger (eds.)]. Kluwer, Dordrecht and Boston, pp. 293-314.
- ⁴⁰¹ U.S. Army Corps of Engineers, 2009: *Risk Reduction Plan: Levees/Floodwalls/Armoring*. [Web site] U.S. Army Corps of Engineers New Orleans District. <http://www.mvn.usace.army.mil/hps2/hps_levees_flood_armor.asp>
- ⁴⁰² Kling, G.W., K. Hayhoe, L.B. Johnson, J.J. Magnuson, S. Polasky, S.K. Robinson, B.J. Shuter, M.M. Wander, D.J. Wuebbles, and D.R. Zak, 2003: *Confronting Climate Change in the Great Lakes Region: Impacts on Our Communities and Ecosystems*. Union of Concerned Scientists, Cambridge, MA, and Ecological Society of America, Washington, DC, 92 pp. <http://www.ucsusa.org/great_lakes/>
- ⁴⁰³ Hayhoe, K., D. Wuebbles, and the Climate Science Team, 2008: *Climate Change and Chicago: Projections and Potential Impacts*. City of Chicago, [175 pp.] <<http://www.chicagoclimatereaction.org/>>
- ⁴⁰⁴ Wuebbles, D.J. and K. Hayhoe, 2004: Climate change projections for the United States Midwest. *Mitigation and Adaptation Strategies for Global Change*, **9**(4), 335-363.
- ⁴⁰⁵ Ebi, K.L. and G.A. Meehl, 2007: The heat is on: climate change and heat waves in the Midwest. In: *Regional Impacts of Climate Change: Four Case Studies in the United States*. Pew Center on Global Climate Change, Arlington, VA, pp. 8-21. <http://www.pewclimate.org/regional_impacts>
- ⁴⁰⁶ Kalkstein, L.S., J.S. Green, D.M. Mills, A.D. Perrin, J.P. Samenow, and J.-C. Cohen, 2008: Analog European heat waves for U.S. cities to analyze impacts on heat-related mortality. *Bulletin of the American Meteorological Society*, **89**(1), 75-85.
- ⁴⁰⁷ Hayhoe, K., S. Sheridan, J.S. Greene and L. Kalkstein, 2009: Climate change, heat waves, and mortality projections for Chicago. *Journal of Great Lakes Research*, in press.
- ⁴⁰⁸ Lin, J.-T., K.O. Patten, X.-Z. Liang, and D.J. Wuebbles, 2008: Climate change effects on ozone air quality in the United States and China with constant precursor emissions. *Journal of Applied Meteorology and Climatology*, **47**(7), 1888-1909.
- ⁴⁰⁹ Holloway, T., S.N. Spak, D. Barker, M. Bretl, K. Hayhoe, J. Van Dorn, and D. Wuebbles, 2008: Change in ozone air pollution over Chicago associated with global climate change. *Journal of Geophysical Research*, **113**, D22306, doi:10.1029/2007JD009775.
- ⁴¹⁰ Hedegaard, G.B., J. Brandt, J.H. Christensen, L.M. Frohn, C. Geels, K.M. Hansen, and M. Stendel, 2008: Impacts of climate change on air pollution levels in the Northern Hemisphere with special focus on Europe and the Arctic. *Atmospheric Chemistry and Physics*, **8**(12), 3337-3367.
- ⁴¹¹ Several different city of Chicago analyses have substantiated the finding of up to 77 degree difference – for example, see City of Chicago, Department of Environment, Undated: *City Hall Rooftop Garden*. [Web site] <http://egov.cityofchicago.org/city/webportal/portalDeptCategoryAction.do?deptCategoryOID=-536889314&contentType=COC_EDITORIAL&topChannelName=Dept&entityName=Environment&deptMainCategoryOID=-536887205>
- ⁴¹² Angel, J. and K. Kunkel, 2009: The response of Great Lakes water levels to future climate scenarios with an emphasis of Lake Michigan. *Journal of Great Lakes Research*, in press.
- ⁴¹³ Annin, P., 2006: *The Great Lakes Water Wars*. Island Press, Washington, DC, 303 pp.
- ⁴¹⁴ Assel, R.A., 2003: *An Electronic Atlas of Great Lakes Ice Cover, Winters: 1973-2002*. NOAA Great Lakes Environmental Research Laboratory, Ann Arbor, MI, 2 CD-ROM set or DVD. <<http://www.glerl.noaa.gov/data/ice/atlas>>
- ⁴¹⁵ Changnon, S.A. (ed.), 1996: *The Great Flood of 1993: Causes, Impacts and Responses*. Westview Press, Boulder, CO, 321 pp.
- ⁴¹⁶ Changnon, S.A. and K.E. Kunkel, 2006: *Severe Storms in the Midwest*. Illinois State Water Survey, Champaign, IL, 74 pp. <<http://www.sws.uiuc.edu/pubdoc/IEM/ISWSIEM2006-06.pdf>>

- ⁴¹⁷ Kunkel, K.E., K. Andsager, G. Conner, W.L. Decker, H.J. Hilaker Jr., P.N. Knox, F.V. Nurnberger, J.C. Rogers, K. Scheeringa, W.M. Wendland, J. Zandlo, and J.R. Angel, 1998: An expanded digital daily database for climatic resources applications in the mid-western United States. *Bulletin of the American Meteorological Society*, **79(7)**, 1357-1366.
- ⁴¹⁸ Arbor Day Foundation, 1990 and 2006: *Hardiness Zones*. [Web site] The Arbor Day Foundation, Nebraska City, NE. <<http://www.arborday.org/media/zones.cfm>>
- ⁴¹⁹ Woodhouse, C.A. and J.T. Overpeck, 1998: 2000 years of drought variability in the central United States. *Bulletin of the American Meteorological Society*, **79(12)**, 2693-2714.
- ⁴²⁰ DeGaetano, A.T. and R.J. Allen, 2002: Trends in twentieth-century temperature extremes across the United States. *Journal of Climate*, **15(22)**, 3188-3205.
- ⁴²¹ Garbrecht, J., M. Van Liew, and G.O. Brown, 2004: Trends in precipitation, streamflow, and evapotranspiration in the Great Plains of the United States. *Journal of Hydrologic Engineering*, **9(5)**, 360-367.
- ⁴²² McGuire, V., 2007: *Water-level Changes in the High Plains Aquifer, Predevelopment to 2005 and 2003 to 2005*. U.S. Geological Survey scientific investigations report 2006-5324. U.S. Geological Survey, Reston, VA, 7 pp. <<http://pubs.usgs.gov/sir/2006/5324/>>
- ⁴²³ Dennehy, K. 2000: *High Plains Regional Ground-Water Study*. U.S. Geological Survey fact sheet FS-091-00, 6 pp. <<http://pubs.er.usgs.gov/usgspubs/fs/fs09100>>
- ⁴²⁴ Gurdak, J.J., R.T. Hanson, P.B. McMahon, B.W. Bruce, J.E. McCray, G.D. Thyne, and R.C. Reedy, 2007: Climate variability controls on unsaturated water and chemical movement, High Plains aquifer, USA. *Vadose Zone Journal*, **6(3)**, 533-547.
- ⁴²⁵ Green, T.R., M. Taniguchi, and H. Kooi, 2007: Potential impacts of climate change and human activity on subsurface water resources. *Vadose Zone Journal*, **6(3)**, 531-532.
- ⁴²⁶ Data from the PRISM Group, Oregon State University <<http://www.prismclimate.org>>; Map created by National Climatic Data Center, March 2009.
- ⁴²⁷ Mahmood, R., S.A. Foster, T. Keeling, K.G. Hubbard, C. Carlson, and R. Leeper, 2006: Impacts of irrigation on 20th century temperature in the northern Great Plains. *Global and Planetary Change*, **54(1-2)**, 1-18.
- ⁴²⁸ Moore, N. and S. Rojstaczer, 2002: Irrigation's influence on precipitation: Texas High Plains, USA. *Geophysical Research Letters*, **29(16)**, 1755, doi:10.1029/2002GL014940.
- ⁴²⁹ Schubert, S.D., M.J. Suarez, P.J. Pegion, R.D. Koster, and J.T. Bacmeister, 2004: On the cause of the 1930s Dust Bowl. *Science*, **303(5665)**, 1855-1859.
- ⁴³⁰ Motha, R.P. and W. Baier, 2005: Impacts of present and future climate change and climate variability on agriculture in the temperate regions: North America. *Climatic Change*, **70(1-2)**, 137-164.
- ⁴³¹ Izaurrealde, R.C., N.J. Rosenberg, R.A. Brown, and A.M. Thomson, 2003: Integrated assessment of Hadley Center (HadCM2) climate-change impacts on agricultural productivity and irrigation water supply in the conterminous United States: Part II. Regional agricultural production in 2030 and 2095. *Agricultural and Forest Meteorology*, **117(1-2)**, 97-122.
- ⁴³² Ziska, L. and K. George, 2004: Rising carbon dioxide and invasive, noxious plants: potential threats and consequences. *World Resource Review*, **16**, 427-447.
- ⁴³³ Parton, W., M. Gutmann, and D. Ojima, 2007: Long-term trends in population, farm income, and crop production in the Great Plains. *Bioscience*, **57(9)**, 737-747.
- ⁴³⁴ Reilly, J., F. Tubiello, B. McCarl, D. Abler, R. Darwin, K. Fuglie, S. Hollinger, C. Izaurrealde, S. Jagtap, J. Jones, L. Mearns, D. Ojima, E. Paul, K. Paustian, S. Riha, N. Rosenberg, and C. Rosenzweig, 2003: U.S. agriculture and climate change: new results. *Climatic Change*, **57(1-2)**, 43-69.
- ⁴³⁵ Allen, V.G., C.P. Brown, R. Kellison, E. Segarra, T. Wheeler, P.A. Dotray, J.C. Conkwright, C.J. Green, and V. Acosta-Martinez, 2005: Integrating cotton and beef production to reduce water withdrawal from the Ogallala aquifer in the southern high plains. *Agronomy Journal*, **97(2)**, 556-567.
- ⁴³⁶ Hanson, J.D., M.A. Liebig, S.D. Merrill, D.L. Tanaka, J.M. Krupinsky, and D.E. Stott, 2007: Dynamic cropping systems: increasing adaptability amid an uncertain future. *Agronomy Journal*, **99(4)**, 939-943.
- ⁴³⁷ Cameron, G.N. and D. Scheel, 2001: Getting warmer: effect of global climate change on distribution of rodents in Texas. *Journal of Mammalogy*, **82(3)**, 652-680.
- ⁴³⁸ Levia, D.F. and E.E. Frost, 2004: Assessment of climatic suitability for the expansion of *Solenopsis invicta* Buren in Oklahoma using three general circulation models. *Theoretical and Applied Climatology*, **79(1-2)**, 23-30.
- ⁴³⁹ Peterson, A.T., 2003: Projected climate change effects on Rocky Mountain and Great Plains birds: generalities of biodiversity consequences. *Global Change Biology*, **9(5)**, 647-655.
- ⁴⁴⁰ Niemuth, N.D. and J.W. Solberg, 2003: Response of waterbirds to number of wetlands in the prairie pothole region of North Dakota, USA. *Waterbirds*, **26(2)**, 233-238.
- ⁴⁴¹ Conway, W.C., L.M. Smith, and J.D. Ray, 2005: Shorebird breeding biology in wetlands of the Playa Lakes, Texas, USA. *Waterbirds*, **28(2)**, 129-138.
- ⁴⁴² Scanlon, B., R. Reedy, and J. Tachovsky, 2007: Semiarid unsaturated zone chloride profiles: archives of past land use change impacts on water resources in the southern High Plains, United States. *Water Resources Research*, **43**, W06423, doi:10.1029/2006WR005769.
- ⁴⁴³ Haukos, D.A. and L.M. Smith, 2003: Past and future impacts of wetland regulations on playa ecology in the southern Great Plains. *Wetlands*, **23(3)**, 577-589.
- ⁴⁴⁴ Matthews, J., 2008: *Anthropogenic Climate Change in the Playa Lakes Joint Venture Region: Understanding Impacts, Discerning Trends, and Developing Responses*. Playa Lakes Joint Venture, Lafayette, CO, 40 pp. <<http://www.pljv.org/cms/climate-change>>
- ⁴⁴⁵ Playa Lakes Joint Venture, Undated: *Playas and the Ogallala Aquifer – What's the Connection?* Playa Lakes Joint Venture, Lafayette, CO, 2 pp. <<http://www.pljv.org/assets/Media/Recharge.pdf>>
- ⁴⁴⁶ U.S. Census Bureau, 2007: *Census Bureau Announces Most Populous Cities*. Press release June 28, 2007. <<http://www.census.gov/Press-Release/www/releases/archives/population/010315.html>>
- ⁴⁴⁷ Ebi, K.L., D.M. Mills, J.B. Smith, and A. Grambsch, 2006: Climate change and human health impacts in the United States: an update on the results of the US national assessment. *Environmental Health Perspectives*, **114(9)**, 1318-1324.
- ⁴⁴⁸ Nielsen, D.C., P.W. Unger, and P.R. Miller, 2005: Efficient water use in dryland cropping systems in the Great Plains. *Agronomy Journal*, **97(2)**, 364-372.
- ⁴⁴⁹ NOAA's National Climatic Data Center, 2008: Southwest region Palmer Hydrological Drought index (PHDI). In: *Climate of 2008 – October: Southwest Region Moisture Status*. [Web site] NOAA National Climatic Data Center, Asheville, NC. <http://www.ncdc.noaa.gov/img/climate/research/prelim/drought/Reg107Dv00_palm06_pg.gif>
- ⁴⁵⁰ Guhathakurta, S. and P. Gober, 2007: The impact of the Phoenix urban heat island on residential water use. *Journal of the American Planning Association*, **73(3)**, 317-329.
- ⁴⁵¹ Gleick, P.H., 1988: Regional hydrologic consequences of increases in atmospheric CO₂ and other trace gases. *Climatic Change*, **10(2)**, 137-160.

- ⁴⁵² Woodhouse, C.A., S.T. Gray, and D.M. Meko, 2006: Updated stream-flow reconstructions for the upper Colorado River basin. *Water Resources Research*, **42**, W05415, doi:10.1029/2005WR004455.
- ⁴⁵³ Meko, D.M., C.A. Woodhouse, C.A. Basisan, T. Knight, J.J. Lukas, M.K. Hughes, and M.W. Salzer, 2007: Medieval drought in the upper Colorado River basin. *Geophysical Research Letters*, **34**, L10705, doi:10.1029/2007GL029988.
- ⁴⁵⁴ Stine, S., 1994: Extreme and persistent drought in California and Patagonia during mediaeval time. *Nature*, **369(6481)**, 546-549.
- ⁴⁵⁵ Breshears, D.D., N.S. Cobb, P.M. Rich, K.P. Price, C.D. Allen, R.G. Balice, W.H. Romme, J.H. Hastens, M.L. Floyd, J. Belnap, J.J. Anderson, O.B. Myers, and C.W. Meyer, 2005: Regional vegetation die-off in response to global-change drought. *Proceedings of the National Academy of Sciences*, **102(42)**, 15144-15148.
- ⁴⁵⁶ Archer, C.L. and K. Caldiera, 2008: Historical trends in the jet streams. *Geophysical Research Letters*, **35**, L08803, doi:10.1029/2008GL033614.
- ⁴⁵⁷ McAfee, S.A. and J.L. Russell, 2008: Northern annular mode impact on spring climate in the western United States. *Geophysical Research Letters*, **35**, L17701, doi:10.1029/2008GL034828.
- ⁴⁵⁸ Westerling, A.L., H.G. Hidalgo, D.R. Cayan, and T.W. Swetnam, 2006: Warming and earlier spring increase western U.S. forest wildfire activity. *Science*, **313(5789)**, 940-943.
- ⁴⁵⁹ Lenihan, J.M., D. Bachelet, R.P. Neilson, and R. Drapek, 2008: Response of vegetation distribution, ecosystem productivity, and fire to climate change scenarios for California. *Climatic Change*, **87(Supplement 1)**, S215-S230.
- ⁴⁶⁰ Westerling, A.L. and B.P. Bryant, 2008: Climate change and wildfire in California. *Climatic Change*, **87(Supplement 1)**, S231-S249.
- ⁴⁶¹ Fried, J.S., J.K. Gilles, W.J. Riley, T.J. Moody, C.S. de Blas, K. Hayhoe, M. Moritz, S. Stephens and M. Torn, 2008: Predicting the effect of climate change on wildfire behavior and initial attack success. *Climatic Change*, **87(Supplement 1)**, S251-S264.
- ⁴⁶² Moritz, M.A. and S.L. Stephens, 2008: Fire and sustainability: considerations for California's altered future climate. *Climatic Change*, **87(Supplement 1)**, S265-S271.
- ⁴⁶³ Rehfeldt, G.E., N.L. Crookston, M.V. Warwell, and J.S. Evans, 2006: Empirical analyses of plant-climate relationships for the western United States. *International Journal of Plant Sciences*, **167(6)**, 1123-1150.
- ⁴⁶⁴ Weiss, J. and J.T. Overpeck, 2005: Is the Sonoran Desert losing its cool? *Global Change Biology*, **11(12)**, 2065-2077.
- ⁴⁶⁵ Dole, K.P., M.E. Loik, and L.C. Sloan, 2003: The relative importance of climate change and the physiological effects CO₂ on freezing tolerance for the future distribution of *Yucca brevifolia*. *Global and Planetary Change*, **36(1-2)**, 137-146.
- ⁴⁶⁶ Loarie, S.R., B.E. Carter, K. Hayhoe, S. McMahon, R. Moe, C.A. Knight, and D.D. Ackerley, 2008: Climate change and the future of California's endemic flora. *PLoS ONE*, **3(6)**, e2502, doi:10.1371/journal.pone.0002502.
- ⁴⁶⁷ Myers, N., R.A. Mittermeier, C.G. Mittermeier, G.A.B. daFonseca, and J. Kent, 2000: Biodiversity hotspots for conservation priorities. *Nature*, **403(6772)**, 853-858.
- ⁴⁶⁸ Mittermeier R.A., P. Robles Gil, M. Hoffman, J. Pilgrim, T. Brooks, C. Goettsch Mittermeier, J. Lamoreux, and G.A.B. da Fonseca, 2005: *Hotspots Revisited: Earth's Biologically Richest and Most Endangered Terrestrial Ecoregions*. Conservation International, Washington DC, 392 pp.
- ⁴⁶⁹ Farjon, A. and C.N. Page (eds.), 1999: *Conifers: Status Survey and Conservation Action Plan*. IUCN/SSC Conifer Special Group. International Union for Conservation of Nature and Natural Resources, Gland, Switzerland, and Cambridge, UK, 121 pp.
- ⁴⁷⁰ Nixon, K.C., 1993: The genus *Quercus* in Mexico. In: *Biological Diversity of Mexico: Origins and Distribution* [Ramamoorthy, T.P., R. Bye, A. Lot, and J. Fa (eds.)]. Oxford University Press, New York, 812 pp.
- ⁴⁷¹ Brower, L.P., G. Castilleja, A. Peralta, J. López-García, L. Bojórquez-Tapia, S. Díaz, D. Melgarejo, and M. Missrie, 2002: Quantitative changes in forest quality in a principal overwintering area of the monarch butterfly in Mexico, 1971-1999. *Conservation Biology*, **16(2)**, 346-359.
- ⁴⁷² Goodrich, G. and A. Ellis, 2008: Climatic controls and hydrologic impacts of a recent extreme seasonal precipitation reversal in Arizona. *Journal of Applied Meteorology and Climatology*, **47(2)**, 498-508.
- ⁴⁷³ Allan, R.P. and B.J. Soden, 2008: Atmospheric warming and the amplifications of precipitation extremes. *Science*, **321(5895)**, 1481-1484.
- ⁴⁷⁴ Bales, R.C., N.P. Molotch, T.H. Painter, M.D. Dettinger, R. Rice, and J. Dozier, 2006: Mountain hydrology of the western United States. *Water Resources Research*, **42**, W08432, doi:10.1029/2005WR004387.
- ⁴⁷⁵ Delta Risk Management Strategy, 2008: Section 2: Sacramento/San Joaquin Delta and Suisun Marsh. In: *Phase I Report: Risk Analysis*. California Department of Water Resources, [13 pp.] <<http://www.drms.water.ca.gov>>
- ⁴⁷⁶ Delta Risk Management Strategy, 2008: Summary report. In: *Phase I Report: Risk Analysis*. California Department of Water Resources, [42 pp.] <<http://www.drms.water.ca.gov>>
- ⁴⁷⁷ Zimmerman, G., C. O'Brady, and B. Hurlbutt, 2006: Climate change: modeling a warmer Rockies and assessing the implications. In: *The 2006 State of the Rockies Report Card*. Colorado College, Colorado Springs, pp. 89-102. <<http://www.coloradocollege.edu/stateoftherockies/06ReportCard.html>>
- ⁴⁷⁸ Lazar, B. and M. Williams, 2008: Climate change in western ski areas: potential changes in the timing of wet avalanches and snow quality for the Aspen ski areas in the years 2030 and 2100. *Cold Regions Science and Technology*, **51(2-3)**, 219-228.
- ⁴⁷⁹ Kleeman, M.J., 2008: A preliminary assessment of the sensitivity of air quality in California to global change. *Climatic Change*, **87(Supplement 1)**, S273-S292.
- ⁴⁸⁰ Vicuna, S., R. Leonardson, M.W. Hanemann, L.L. Dale, and J.A. Dracup, 2008: Climate change impacts on high elevation hydro-power generation in California's Sierra Nevada: a case study in the upper American River. *Climatic Change*, **87(Supplement 1)**, S123-S137.
- ⁴⁸¹ Medellín-Azuara, J., J.J. Harou, M.A. Olivares, K. Madani, J.R. Lund, R.E. Howitt, S.K. Tanaka, M.W. Jenkins, and T. Zhu, 2008: Adaptability and adaptations of California's water supply system to dry climate warming. *Climatic Change*, **87(Supplement 1)**, S75-S90.
- ⁴⁸² Baldocchi, D. and S. Wong, 2008: Accumulated winter chill is decreasing in the fruit growing regions of California. *Climatic Change*, **87(Supplement 1)**, S153-S166.
- ⁴⁸³ Lobell, D., C. Field, K. Nicholas Cahill, and C. Bonfils, 2006: Impacts of future climate change on California perennial crop yields: model projections with climate and crop uncertainties. *Agricultural and Forest Meteorology*, **141(2-4)**, 208-218.
- ⁴⁸⁴ Purkey, D.R., B. Joyce, S. Vicuna, M.W. Hanemann, L.L. Dale, D. Yates, and J.A. Dracup, 2008: Robust analysis of future climate change impacts on water for agriculture and other sectors: a case study in the Sacramento Valley. *Climatic Change*, **87(Supplement 1)**, S109-S122.
- ⁴⁸⁵ Mote, P.W., 2003: Trends in temperature and precipitation in the Pacific Northwest during the twentieth century. *Northwest Science*, **77(4)**, 271-282.

- ⁴⁸⁶ Mote, P., E. Salathé, V. Dulière, and E. Jump, 2008: *Scenarios of Future Climate for the Pacific Northwest*. Climate Impacts Group, University of Washington, Seattle, 12 pp. <<http://cses.washington.edu/db/pubs/abstract628.shtml>>
- ⁴⁸⁷ Hamlet, A.F., P.W. Mote, M. Clark, and D.P. Lettenmaier, 2005: Effects of temperature and precipitation variability on snowpack trends in the western United States. *Journal of Climate*, **18(21)**, 4545-4561.
- ⁴⁸⁸ Mote, P.W., 2006: Climate-driven variability and trends in mountain snowpack in western North America. *Journal of Climate*, **19(23)**, 6209-6220.
- ⁴⁸⁹ Payne, J.T., A.W. Wood, A.F. Hamlet, R.N. Palmer, and D.P. Lettenmaier, 2004: Mitigating the effects of climate change on the water resources of the Columbia River basin. *Climatic Change*, **62(1-3)**, 233-256.
- ⁴⁹⁰ Figure provided by the Climate Impacts Group, University of Washington, Seattle. <<http://cses.washington.edu/cig/>>
- ⁴⁹¹ Mote, P.W., A.F. Hamlet, M.P. Clark, and D.P. Lettenmaier, 2005: Declining mountain snowpack in western North America. *Bulletin of the American Meteorological Society*, **86(1)**, 39-49.
- ⁴⁹² Hamlet, A.F. and D.P. Lettenmaier, 2007: Effects of 20th century warming and climate variability on flood risk in the western U.S. *Water Resources Research*, **43**, W06427, doi:10.1029/2006WR005099.
- ⁴⁹³ Bonneville Power Administration, 2001: *The Columbia River System Inside Story*. Internal report DOE/BP-3372. Bonneville Power Administration, Portland OR, 2nd edition, 78 pp. <http://www.bpa.gov/corporate/Power_of_Learning/docs/columbia_river_in_side_story.pdf>
- ⁴⁹⁴ Casola, J.H., J.E. Kay, A.K. Snover, R.A. Norheim, L.C. Whitely Binder, and Climate Impacts Group, 2005: *Climate Impacts on Washington's Hydropower, Water Supply, Forests, Fish, and Agriculture*. Center for Science in the Earth System, Joint Institute for the Study of the Atmosphere and Ocean, University of Washington, Seattle, 43 pp. <<http://cses.washington.edu/db/pdf/kc05whitepaper459.pdf>>
- ⁴⁹⁵ Ministry of the Environment, British Columbia, Canada, 2007: *Environmental Trends 2007: The Mountain Pine Beetle in British Columbia*. <<http://www.env.gov.bc.ca/soe/et07/pinebeetle.html>>
- ⁴⁹⁶ Francis, R.C. and N.J. Mantua, 2003: Climatic influences on salmon populations in the northeast Pacific. In: *Assessing Extinction Risk for West Coast Salmon* [MacCall, A.D. and T.C. Wainwright (eds.)]. NOAA technical memo NMFS-NWFSC-56. National Marine Fisheries Service, [Washington, DC], pp. 37-67. <http://www.nwfsc.noaa.gov/assets/25/3946_06162004_130044_tm56.pdf>
- ⁴⁹⁷ Salathé, E.P., 2005: Downscaling simulations of future global climate with application to hydrologic modelling. *International Journal of Climatology*, **25(4)**, 419-436.
- ⁴⁹⁸ Mote, P.W., A. Petersen, S. Reeder, H. Shipman, and L. Whitely Binder, 2008: *Sea Level Rise Scenarios for Washington State*. Center for Science in the Earth System, Joint Institute for the Study of the Atmosphere and Oceans, University of Washington, Seattle, and Washington Department of Ecology, Lacey, 11 pp. <<http://www.cses.washington.edu/db/pdf/moteetalslr579.pdf>>
- ⁴⁹⁹ Petersen, A.W., 2007: *Anticipating Sea Level Rise Response in Puget Sound*. M.M.A. thesis, School of Marine Affairs. University of Washington, Seattle, 73 pp.
- ⁵⁰⁰ Snover, A.K., L. Whitely Binder, J. Lopez, E. Willmott, J. Kay, D. Howell, and J. Simmonds, 2007: *Preparing for Climate Change: A Guidebook for Local, Regional, and State Governments*. The Climate Impacts Group, University of Washington, and King County, Washington, in association with and published by ICLEI – Local Governments for Sustainability, Oakland, CA, 172 pp.
- ⁵⁰¹ Fitzpatrick, J., R.B. Alley, J. Brigham-Grette, G.H. Miller, L. Polyak, and M. Serreze, 2008: Preface: Why and how to use this synthesis and assessment report. In: *Past Climate Variability and Change in the Arctic and at High Latitude*. Synthesis and Assessment Product 1.2. U.S. Geological Survey, Reston, VA, pp. 8-21.
- ⁵⁰² Data provided by Dr. Glenn Juday, School of Natural Resources and Agricultural Science, Agricultural and Forestry Experiment Station, University of Alaska, Fairbanks.
- ⁵⁰³ Euskirchen, E.S., A.D. McGuire, D.W. Kicklighter, Q. Zhuang, J.S. Klein, R.J. Dargaville, D.G. Dye, J.S. Kimball, K.C. McDonald, V.E. Melillo, V.E. Romanovsky, and N.V. Smith, 2006: Importance of recent shifts in soil thermal dynamics on growing season length, productivity, and carbon sequestration in terrestrial high-latitude ecosystems. *Global Change Biology*, **12(4)**, 731-750.
- ⁵⁰⁴ Euskirchen, E.S., A.D. McGuire, and F.S. Chapin III, 2007: Energy feedbacks of northern high-latitude ecosystems to the climate system due to reduced snow cover during 20th century warming. *Global Change Biology*, **13(11)**, 2425-2438.
- ⁵⁰⁵ Barber, V.A., G.P. Juday, and B.P. Finney, 2000: Reduced growth of Alaskan white spruce in the twentieth century from temperature-induced drought stress. *Nature*, **405 (6787)**, 668-673.
- ⁵⁰⁶ Juday, G.P., V. Barber, P. Duffy, H. Linderholm, T.S. Rupp, S. Sparrow, E. Vaganov, and J. Yarie, 2005: Forests, land management, and agriculture. In: *Arctic Climate Impact Assessment*. Cambridge University Press, Cambridge, UK, and New York, pp. 781-862. <<http://www.acia.uaf.edu/pages/scientific.html>>
- ⁵⁰⁷ Balshi, M.S., A.D. McGuire, P. Duffy, M. Flannigan, J. Walsh, and J.M. Melillo, 2008: Assessing the response of area burned to changing climate in western boreal North America using a Multivariate Adaptive Regression Splines (MARS) approach. *Global Change Biology*, **15(3)**, 578-600.
- ⁵⁰⁸ Berman, M., G.P. Juday, and R. Burnside, 1999: Climate change and Alaska's forests: people, problems, and policies. In: *Assessing the Consequences of Climate Change in Alaska and the Bering Sea Region*. Proceedings of a workshop at the University of Alaska Fairbanks, 29-30 October 1998. Center for Global Change and Arctic System Research, University of Alaska, pp. 21-42.
- ⁵⁰⁹ Fleming, R.A. and W.J.A. Volney, 1995: Effects of climate change on insect defoliator population processes in Canada's boreal forest: some plausible scenarios. *Water, Soil, and Air Pollution*, **82(1-2)**, 445-454.
- ⁵¹⁰ Kasischke, E.S. and M.R. Turetsky, 2006: Recent changes in the fire regime across the North American boreal region - spatial and temporal patterns of burning across Canada and Alaska. *Geophysical Research Letters*, **33**, L09703, doi:10.1029/2006GL025677.
- ⁵¹¹ Chapin, F.S., III, S.F. Trainor, O. Huntington, A.L. Lovecraft, E. Zavaleta, D.C. Natcher, A.D. McGuire, J.L. Nelson, L. Ray, M. Calef, N. Fresco, H. Huntington, T.S. Rupp, L. DeWilde, and R.A. Naylor, 2008: Increasing wildfire in Alaska's boreal forest: pathways to potential solutions of a wicked problem. *BioScience*, **58(6)**, 531-540.
- ⁵¹² Berner, J. and C. Furgal, 2005: Human health. In: *Arctic Climate Impact Assessment*. Cambridge University Press, Cambridge, UK, and New York, pp. 863-906. <<http://www.acia.uaf.edu/pages/scientific.html>>
- ⁵¹³ Klein, E., E.E. Berg, and R. Dial, 2005: Wetland drying and succession across the Kenai Peninsula Lowlands, south-central Alaska. *Canadian Journal of Forest Research*, **35(8)**, 1931-1941.
- ⁵¹⁴ Riordan, B., D. Verbyla, and A.D. McGuire, 2006: Shrinking ponds in subarctic Alaska based on 1950–2002 remotely sensed images. *Journal of Geophysical Research*, **111**, G04002, doi:10.1029/2005JG000150.

- ⁵¹⁵ Osterkamp, T., 2007: Characteristics of the recent warming of permafrost in Alaska. *Journal of Geophysical Research*, **112**, F02S02, doi:10.1029/2006JF000578.
- ⁵¹⁶ Instanes, A., O. Anisimov, L. Brigham, D. Goering, L.N. Khrustalev, B. Ladanyi, and J.O. Larsen, 2005: Infrastructure: buildings, support systems, and industrial facilities. In: *Arctic Climate Impact Assessment*. Cambridge University Press, Cambridge, UK, and New York, pp. 907-944. <<http://www.acia.uaf.edu/pages/scientific.html>>
- ⁵¹⁷ Brown, J. and V.E. Romanovsky, 2008: Report from the International Permafrost Association: State of permafrost in the first decade of the 21st century. *Permafrost and Periglacial Processes*, **19(2)**, 255-260.
- ⁵¹⁸ Busey, R.C., L.D. Hinzman, J.J. Cassano, and E. Cassano, 2008: Permafrost distributions on the Seward Peninsula: past, present, and future. In: *Proceedings of the Ninth International Conference on Permafrost*, University of Alaska Fairbanks, June 29-July 3, 2008 [Kane, D.L. and K.M. Hinkel (eds.)]. Institute of Northern Engineering, University of Alaska, Fairbanks, pp. 215-220.
- ⁵¹⁹ Shumaker, V.O., 1979: *The Alaska Pipeline*. Julian Messner, New York, 64 pp.
- ⁵²⁰ Bengtsson, L., K.I. Hodges, and E. Roeckner, 2006: Storm tracks and climate change. *Journal of Climate*, **19(15)**, 3518-3543.
- ⁵²¹ U.S. Army Corps of Engineers, 2008: *Newtok Evacuation Center, Mertarvik, Nelson Island, Alaska*. Revised environmental assessment. U.S. Army Corps of Engineers, Alaska District, [Elmendorf AFB, 64 pp.]
- ⁵²² Jones, B.M., C.D. Arp, M.T. Jorgenson, K.M. Hinkel, J.A. Schmutz, and P.L. Flint, 2009: Increase in the rate and uniformity of coastline erosion in Arctic Alaska. *Geophysical Research Letters*, **36**, L03503, doi:10.1029/2008GL036205.
- ⁵²³ Yin, J.H., 2005: A consistent poleward shift of the storm tracks in simulations of 21st century climate. *Geophysical Research Letters*, **32**, L18701, doi:10.1029/2005GL023684.
- ⁵²⁴ Salathé, E.P., Jr., 2006: Influences of a shift in North Pacific storm tracks on western North American precipitation under global warming. *Geophysical Research Letters*, **33**, L19820, doi:10.1029/2006GL026882.
- ⁵²⁵ Data provided by Dr. David Atkinson, International Arctic Research Center, University of Alaska, Fairbanks.
- ⁵²⁶ Feely, R.A., C.L. Sabine, K. Lee, W. Berelson, J. Kleypas, V.J. Fabry and F.J. Millero, 2004: Impact of anthropogenic CO₂ on CaCO₃ system in the oceans. *Science*, **305(5682)**, 362-366.
- ⁵²⁷ Grebmeier, J.M., J.E. Overland, S.E. Moore, E.V. Farley, E.C. Carmack, L.W. Cooper, K.E. Frey, J.H. Helle, F.A. McLaughlin, and S.L. McNutt, 2006: A major ecosystem shift in the northern Bering Sea. *Science*, **311(5766)**, 1461-1464.
- ⁵²⁸ Ray, G.C., J. McCormick-Ray, P. Berg, and H.E. Epstein, 2006: Pacific walrus: benthic bioturbator of Beringia. *Journal of Experimental Marine Biology and Ecology*, **330(1)**, 403-419.
- ⁵²⁹ Mueter, F.J. and M.A. Litzow, 2007: Sea ice retreat alters the biogeography of the Bering Sea continental shelf. *Ecological Applications*, **18(2)**, 309-320.
- ⁵³⁰ Shea, E.L., G. Dolcemascolo, C.L. Anderson, A. Barnston, C.P. Guard, M.P. Hamnett, S.T. Kubota, N. Lewis, J. Loschnigg, and G. Meehl, 2001: *Preparing for a Changing Climate: The Potential Consequences of Climate Variability and Change: Pacific Islands*. East-West Center, Honolulu, HI, 102 pp. <<http://www2.eastwestcenter.org/climate/assessment/>>
- ⁵³¹ Guidry, M.W. and F.T. Mackenzie, 2006: *Climate Change, Water Resources, and Sustainability in the Pacific Basin: Emphasis on O'ahu, Hawai'i and Majuro Atoll, Republic of the Marshall Islands*. University of Hawaii Sea Grant Program, Honolulu, 100 pp. <<http://nsgl.gso.uri.edu/hawau/hawaut06001.pdf>>
- ⁵³² CEO, 2004: *Caribbean Environmental Outlook*. [Heileman, S., L.J. Walling, C. Douglas, M. Mason, and M. Chevannes-Creary (eds.)]. United Nations Environmental Programme, Kingston, Jamaica, 114 pp. <http://www.unep.org/geo/pdfs/caribbean_eo.pdf>
- ⁵³³ Church, J.A., N.J. White, and J.R. Hunter, 2006: Sea-level rise at tropical Pacific and Indian Ocean islands. *Global and Planetary Change*, **53(3)**, 155-168.
- ⁵³⁴ Burns, W.C.G., 2002: Pacific island developing country water resources and climate change. In: *The World's Water* [Gleick, P. (ed.)]. Island Press, Washington, DC, 3rd edition, pp. 113-132.
- ⁵³⁵ Baker, J.D., C.L. Littnan, and D.W. Johnston, 2006: Potential effects of sea level rise on the terrestrial habitats of endangered and endemic megafauna in the northwestern Hawaiian Islands. *Endangered Species Research*, **2**, 21-30.
- ⁵³⁶ Firing, Y. and M.A. Merrifield, 2004: Extreme sea level events at Hawaii: influence of mesoscale eddies. *Geophysical Research Letters*, **31**, L24306, doi:10.1029/2004GL021539.
- ⁵³⁷ Scott, D., M. Overmars, T. Falkland, and C. Carpenter, 2003: *Pacific Dialogue on Water and Climate, Synthesis Report*. South Pacific Applied Geoscience Commission, Fiji Islands, 28 pp. <<http://www.oas.org/CDWC/Documents/SIDS%20Paper/Pacific%20Report%20-%20Final.pdf>>
- ⁵³⁸ Hay, J., N. Mimura, J. Campbell, S. Fifita, K. Koshy, R.F. McLean, T. Nakalevu, P. Nunn, and N. deWet, 2003: *Climate Variability and Change and Sea Level Rise in the Pacific Islands Regions: A Resource Book for Policy and Decision Makers, Educators and Other Stakeholders*. South Pacific Regional Environmental Programme (SPREP), Apia, Samoa, 94 pp. <http://www.sprep.org/publication/pub_detail.asp?id=181>
- ⁵³⁹ Gillespie, R.G., E.M. Claridge, and G.K. Roderick, 2008: Biodiversity dynamics in isolated island communities: interaction between natural and human-mediated processes. *Molecular Ecology*, **17(1)**, 45-57.
- ⁵⁴⁰ Becken, S. and J.E. Hay, 2007: *Tourism and Climate Change: Risks and Opportunities*. Channel View Publications, Clevedon, UK, and Buffalo, NY, 329 pp.
- ⁵⁴¹ State of Hawaii, Division of Business, Economic Development, and Tourism, 2008: *Facts and Figures: State of Hawaii*. [Web site] <<http://hawaii.gov/dbedt/info/economic/library/facts/state>>
- ⁵⁴² Cesar, H.S.F. and F.H. van Beukering, 2004: Economic valuation of the coral reefs of Hawaii. *Pacific Science*, **58(2)**, 231-242.
- ⁵⁴³ Hoegh-Guldberg, O., P.J. Mumby, A.J. Hooten, R.S. Steneck, P. Greenfield, E. Gomez, C.D. Harvell, P.F. Sale, A.J. Edwards, K. Caldeira, N. Knowlton, C.M. Eakin, R. Iglesias-Prieto, N. Muthiga, R.H. Bradbury, A. Dubi, and M.E. Hatzitolos, 2007: Coral reefs under rapid climate change and ocean acidification. *Science*, **318(5857)**, 1737-1742.
- ⁵⁴⁴ Donner, S.D., W.J. Skirving, C.M. Little, M. Oppenheimer, and O. Hoegh-Guldberg, 2005: Global assessment of coral bleaching and required rates of adaptation under climate change. *Global Change Biology*, **11(12)**, 2251-2265.
- ⁵⁴⁵ Pauly, G., L. Kirkendale, G. Lambert, and C. Meyer, 2002: Anthropogenic biotic interchange in a coral reef ecosystem: a case study from Guam. *Pacific Science*, **56(4)**, 403-422.
- ⁵⁴⁶ Hotta, M., 2000: The sustainable contribution of fisheries to food security in the Asia and Pacific region: regional synthesis. In: *Sustainable Contribution of Fisheries to Food Security*. Food and Agriculture Organization of the United Nations, Bangkok, Thailand, pp. 1-28. <<http://www.fao.org/DOCREP/003/X6956E/x6956e02.htm>>
- ⁵⁴⁷ Graham, N.A.J., S.K. Wilson, S. Jennings, N.V.C. Polunin, J.P. Bijoux, and J. Robinson, 2006: Dynamic fragility of oceanic coral reef ecosystems. *Proceedings of the National Academy of Sciences*, **103(22)**, 8425-8429.

- ⁵⁴⁸ Lehodey, O., M. Bertignac, J. Hampton, A. Lewis, and J. Picaut, 1997: El Niño Southern Oscillation and tuna in the western Pacific. *Nature*, **389(6652)**, 715-718.
- ⁵⁴⁹ Crowell, M., S. Edelman, K. Coulton, and S. McAfee, 2007: How many people live in coastal areas? *Journal of Coastal Research*, **23(5)**, iii-vi.
- ⁵⁵⁰ Gill, S.K., R. Wright, J.G. Titus, R. Kafalenos, and K. Wright, 2009: Population, land use, and infrastructure. In: *Coastal Elevations and Sensitivity to Sea-level Rise: A Focus on the Mid-Atlantic Region* [J.G. Titus (coordinating lead author), K.E. Anderson, D.R. Cahoon, D.B. Gesch, S.K. Gill, B.T. Gutierrez, E.R. Thieler, and S.J. Williams (lead authors)]. Synthesis and Assessment Product 4.1. U.S. Environmental Protection Agency, Washington, DC, pp. 105-116.
- ⁵⁵¹ U.S. Commission on Ocean Policy, 2004: *An Ocean Blueprint for the 21st Century*. U.S. Commission on Ocean Policy, Washington, DC. <http://www.oceancommission.gov/documents/full_color_rpt/welcome.html>
- ⁵⁵² Day, J.W., Jr., D.F. Boesch, E.J. Clairain, G.P. Kemp, S.B. Laska, W.J. Mitsch, K. Orth, H. Mashriqui, D.J. Reed, L. Shabman, C.A. Simenstad, B.J. Streever, R.R. Twilley, C.C. Watson, J.T. Wells, and D.F. Whigham, 2007: Restoration of the Mississippi Delta: lessons from hurricanes Katrina and Rita. *Science*, **315(5819)**, 1679-1684.
- ⁵⁵³ Boesch, D.F. (ed.), 2008: *Global Warming and the Free State: Comprehensive Assessment of Climate Change Impacts in Maryland*. University of Maryland Center for Environmental Science, Cambridge, MA, 85 pp. <<http://www.umces.edu/climateimpacts/>>
- ⁵⁵⁴ Cazenave, A., K. Dominh, S. Guinehut, E. Berthier, W. Llovel, G. Ramillien, M. Ablain, and G. Larnicol, 2009: Sea level budget over 2003–2008: a reevaluation from GRACE space gravimetry, satellite altimetry and Argo. *Global and Planetary Change*, **65(1-2)**, 83-88.
- ⁵⁵⁵ Vaughan, D.G. and R. Arthern, 2007: Why is it hard to predict the future of ice sheets? *Science*, **315(5818)**, 1503-1504.
- ⁵⁵⁶ Justić, D., N.N. Rabalais, and R.E. Turner, 2003: Simulated responses of the Gulf of Mexico hypoxia to variations in climate and anthropogenic nutrient loading. *Journal of Marine Systems*, **42(2-3)**, 115-126.
- ⁵⁵⁷ Boesch, D.F., V.J. Coles, D.G. Kimmel, and W.D. Miller, 2007: Coastal dead zones and global climate change: ramifications of climate change for Chesapeake Bay hypoxia. In: *Regional Impacts of Climate Change: Four Case Studies in the United States*. Pew Center for Global Climate Change, Arlington, VA, pp. 57-70. <http://www.pewclimate.org/regional_impacts>
- ⁵⁵⁸ Wicks, C., D. Jasinski, and B. Longstaff, 2007: *Breath of Life: Dissolved Oxygen in Chesapeake Bay*. EcoCheck, Oxford, MD, 4 pp. <http://www.eco-check.org/pdfs/do_letter.pdf>
- ⁵⁵⁹ Kimmerer, W.J., E. Garside, and J.J. Orsi, 1994: Predation by an introduced clam as the likely cause of substantial declines in zooplankton of San Francisco Bay. *Marine Ecology Progress Series*, **113**, 81-93.
- ⁵⁶⁰ Carpenter, K.E., M. Abrar, G. Aeby, R.B. Aronson, S. Banks, A. Bruckner, A. Chiriboga, J. Cortés, J.C. Delbeek, L. DeVantier, G.J. Edgar, A.J. Edwards, D. Fenner, H.M. Guzmán, B.W. Hoeksema, G. Hodgson, O. Johan, W.Y. Licuanan, S.R. Livingstone, E.R. Lovell, J.A. Moore, D.O. Obura, D. Ochavillo, B.A. Polidoro, W.F. Precht, M.C. Quibilan, C. Reboton, Z.T. Richards, A.D. Rogers, J. Sanciangco, A. Sheppard, C. Sheppard, J. Smith, S. Stuart, E. Turak, J.E.N. Veron, C. Wallace, E. Weil, and E. Wood, 2008: One-third of reef-building corals face elevated extinction risk from climate change and local impacts. *Science*, **321(5888)**, 560-563.
- ⁵⁶¹ Chan, F., J.A. Barth, J. Lubchenco, A. Kirincich, H. Weeks, W.T. Peterson, and B.A. Menge, 2008: Emergence of anoxia in the California current large marine ecosystem. *Science*, **319(5865)**, 920.
- ⁵⁶² Data for hypoxia distribution maps made available by Dr. Bill Peterson, (NOAA), the Bonneville Power Administration, Drs. Francis Chan and Jane Lubchenco, (PISCO, Partnership for Interdisciplinary Studies of Coastal Oceans), Drs. Jack Barth and Steve Pierce (Oregon State University/College of Oceanic and Atmospheric Sciences) and NOAA Northwest Fisheries Science Center.

PHOTOGRAPHY CREDITS

Inside Cover: USA Map, NOAA/NCDC

Table of Contents: Global, planet earth, NASA Water

Glass, ©iStockphotos.com/luchschen; Tractor

trailer ©iStockphotos.com/claylib; Hospital scene, ©iStockphotos.com/Jeffrey Smith; San Francisco, Grant Goodge, STG Inc.;

Page 13: Planet Earth, NASA

Page 38: Lake Effect snow, Dave Kelly, NOAA/NSSL

Page 41: Water Glass, ©iStockphotos.com/luchschen; Skagit River, Philip Mote PhD, JISAO/CSES Climate Impacts Group

Page 47: Turbidity in New York City's Ashokan, Thurlough Smyth, Climate Change Planning, New York City Department of Environmental Protection, Bureau of Environmental Planning & Assessment

Page 48: Flood water road damage, Grant Goodge, STG Inc.

Page 51: Two Lake Powell photos, John C. Dohrenwend

Page 52: Water and energy corner schematic, Electric Power Research Institute, Palo Alto, CA and Tetra Tech Inc., Lafayette, CA

Page 59: St. Lucie power plant, CCSP SAP 4.5

Page 60: Hydro-electric dam, Philip Mote PhD, JISAO/CSES Climate Impacts Group; Ship, NOAA

Page 61: Tractor trailer ©iStockphotos.com/claylib; Houses against bridge, Associated Press/Jeff Roberson

Page 65: Kosrae Island photo, John E. Hay

Page 68: Broken bridge, Joe Furr

Page 75: CO₂ plant photos, Lewis Ziska, USDA ARS

Page 80: Edith's checkerspot butterfly, Camille Parmesan, University of Texas at Austin

Page 83: Rob Wu, CCSP SAP 4.3

Page 85: Polar bear, Mike Dunn, North Carolina Museum of Natural Sciences

Page 86: Pika, ©iStockphotos.com/Global Exposure; Forest species shift, Adapted from James Abundis, Boston Globe

Page 89: Hospital scene, ©iStockphotos.com/Jeffrey Smith;

Page 92: New York and Atlanta smog, Paul Grabhorn

Page 93: Los Angeles smog, ©iStockphotos.com/Stein Photo

Page 95: Mosquito, ©iStockphotos.com/Julie Le de Leseleuc; Virus cycle illustration, Deborah Misch, STG Inc.

Page 97: Poison Ivy plant and girl scratching, Sara Veasey, NCDC

Page 98: EMS officer and woman, Sara Veasey, NOAA/NCDC

Page 99: Chapter and thumbnail photo, Grant Goodge, STG Inc., Midwest flood, Associated Press/Jeff Roberson

Page 101: Person in floodwater, Associated Press/Dave Martin; Aerial flood, USGS

Page 109: Boston Deer Island sewage plant, Massachusetts Water Resources Authority

Page 113: Dry Lake Lanier, Kent Frantz, SRH NOAA

Page 114: Barrier islands, USGS

Page 115: Fish plant, Barataria Terrebonne; Flooded area, Dr. Kerry St. Pé of the Barataria-Terrebonne National Estuary program

Page 116: Flood levee, USA COE

Page 118: Chicago map, City of Chicago; Thermal image, Chicago Police Department and the Cook County Sheriff's Department

Page 120: Flooded highway, Missouri Highway & Transportation
Department/USGS; Flood of 1993, NOAA
Page 125: Dust bowl, NOAA
Page 128: Seedlings, USDA/NRCS
Page 126: Mallard duck, ©iStockphotos.com/Andy Gehrig
Page 131: Playa lakes, Katharine Hayhoe, Texas Tech University
Page 143: Alaska pipeline, SSgt. Joshua Strang, USAF
Page 147: El Niño Billboard, US NWS Pacific Region Headquarters
Page 148: Micronesia, Eric L. Gilman
Page 150: Ghost swamp, USGS, National Wetlands Research Center
Page 152: Sea wall, Jack Pellette, NWS

DRAFT

DRAFT

Contact Information

Global Change Research Information Office
c/o U.S. Global Change Research Program Office
1717 Pennsylvania Avenue, NW
Suite 250
Washington, DC 20006
202-223-6262 (voice)
202-223-3065 (fax)

To obtain a copy of this document, place
an order at the Global Change Research
Information Office (GCRI) web site:
<http://www.gcric.org/orders>

U.S. Global Change Research Program and the Subcommittee on Global Change Research

Jack Kaye, Vice Chair,
National Aeronautics and Space Administration,
Acting Director, U.S. Global Change Research
Program

Allen Dearry
Department of Health and Human Services

Anna Palmisano
Department of Energy

Mary Glackin
National Oceanic and Atmospheric Administration

Patricia Gruber
Department of Defense

William Hohenstein
Department of Agriculture

Linda Lawson
Department of Transportation

Mark Myers
U.S. Geological Survey

Tim Killeen
National Science Foundation

Patrick Neale
Smithsonian Institution **Jacqueline Schafer**
U.S. Agency for International Development

Joel Scheraga
Environmental Protection Agency

Harlan Watson
Department of State

EXECUTIVE OFFICE AND OTHER LIAISONS

Robert Marlay
Climate Change Technology Program

Katharine Gebbie
National Institute of Standards & Technology

Stuart Levenbach
Office of Management and Budget

Margaret McCalla
Office of the Federal Coordinator for Meteorology

1
2
3
4
5
6
7
8
9
10
11
12
13
14
15
16
17
18
19
20
21
22
23
24
25
26
27
28
29
30
31
32
33
34
35
36
37
38
39
40
41
42
43
44
45
46
47
48
49
50
51
52

1
2
3
4
5
6
7
8
9
10
11
12
13
14
15
16
17
18
19
20
21
22
23
24
25
26
27
28
29
30
31
32
33
34
35
36
37
38
39
40
41
42
43
44
45
46
47
48
49
50
51
52

DRAFT



U.S. Global Change Research Program
1717 Pennsylvania Avenue, NW • Suite 250 • Washington, DC 20006 USA
1-202-223-6262 (voice) • 1-202-223-3065 (fax)
<http://www.climate-science.gov>

