

1 **Chapter 9 — Key Findings and Recommendations**

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10 **9.1 INTRODUCTION**

11 Paleoclimatic data provide a highly informative if incomplete history of Arctic climate.
12 Temperature history is especially well recorded, and it commonly allows researchers to
13 accurately reconstruct changes and rates of changes for particular seasons. Precipitation (rain or
14 snow) and the extent of ice on land and sea are some of the many other climate variables that
15 have also been reconstructed. The data also provide insight to the histories of many possible
16 causes of the climate changes and feedback processes that amplify or reduce the resulting
17 changes. Comparing climate with possible causes allows scientists to generate and test
18 hypotheses, and those hypotheses then become the basis for projections of future changes.

19 Arctic data show changes on numerous time scales and indicate many causes and
20 important feedback processes. Changes in greenhouse gases appear to have been especially
21 important in causing climate changes [sections 4.4; 5.4.1; 5.4.4, 6.4.1; 6.4.2]. Global climate
22 changes have been notably amplified in the Arctic [section 5.5.2], and warmer times have
23 melted ice on land and sea [chapter 8].¹

24

25 **9.2 SUMMARY OF KEY FINDINGS**

26 **Chapter 5 Temperature and Precipitation**

¹ Statistically valid confidence levels often can be attached to scientific findings, but commonly require many independent samples from a large population. Such a standard can be applied to paleoclimatic data in only some cases, whereas in other cases the necessary archives or interpretative tools are not available. However, expert judgment can also be used to assess confidence. The key findings here cannot all be evaluated rigorously using parametric statistics, but on the basis of assessment by the authors, all of the key findings are at least “likely” as used by the Intergovernmental Panel on Climate Change (more than 66% chance of being correct); the authors believe that the most of the findings are “very likely” (more than a 90% chance of being correct).

27 The Arctic of 65 million years ago (Ma) was much warmer than recently; forests grew in
28 all land regions and neither perennial sea ice nor the Greenland Ice Sheet were present. Gradual
29 but bumpy cooling has dominated since, with falling atmospheric CO₂ concentration apparently
30 the most important contributor to the cooling, although with possible additional contributions
31 from changing continental positions and their effect on atmospheric or oceanic circulation.
32 Warm “bumps” during the general cooling trend include the relatively abrupt Paleocene-Eocene
33 Thermal Maximum about 55 Ma, apparently caused by CO₂ release, and a more gradual
34 warming in the middle Pliocene (about 3 Ma) of uncertain cause.

35 Around 2.7 Ma cooling reached the threshold for extensive development of continental
36 ice sheets throughout the North American and Eurasian Arctic. Periodic growth and shrinkage
37 of the ice during tens of thousands of years were directly controlled by periodic changes in
38 northern hemisphere sunshine caused by features of Earth’s orbit. Recent work suggests that, in
39 the absence of human influence, the current interglacial would continue for a few tens of
40 thousands of years before the start of a new ice age. The large temperature differences between
41 glacial and interglacial periods, although driven by Earth’s orbital cycles and the globally
42 synchronous response, reflect the effects of strong positive feedbacks, such as changes in
43 atmospheric CO₂ and other greenhouse gases and in the extent of reflective snow and ice.

44 Interactions among the various orbital cycles have caused small differences between
45 successive interglacials. During the interglacial about 130–120 thousand years ago (ka), the
46 Arctic received more summer sunshine than in the current interglacial, and temperatures in
47 many places were consequently 4° to 6°C warmer than recently, which reduced ice on
48 Greenland (Chapter 7), raised sea level, and melted small glaciers and ice caps.

49 The cooling into and warming out of the most recent glacial were punctuated by
50 numerous abrupt climate changes, with millennial persistence of conditions between jumps
51 requiring years to decades. These events were very large around the North Atlantic but much
52 smaller elsewhere in the Arctic and beyond. Large changes in the extent of sea ice in the North
53 Atlantic were probably responsible, linked to changes in regional and global patterns of ocean
54 circulation. Freshening of the North Atlantic also favored formation of sea-ice.

55 These abrupt changes also occurred in the current interglacial (the Holocene), but they
56 ended as the Laurentide Ice Sheet on Canada melted away. Arctic temperatures in the
57 Holocene broadly responded to orbital changes with warmer temperatures during the early to
58 middle Holocene when there was more summer sunshine. Warming generally led to northward
59 migration of vegetation and to shrinkage of ice on land and sea. Small oscillations in climate
60 during the Holocene, such as the Medieval Climate Anomaly and the Little Ice Age, were
61 linked to variations in the sun-blocking effect of particles from explosive volcanoes and
62 perhaps to small variations in solar output or in ocean circulation or other factors. The warming
63 after the Little Ice Age began for largely natural reasons, but there is now high scientific
64 confidence that human contributions, and especially increasing concentrations of CO₂, have
65 come to dominate the warming (Jansen et al., 2007).

66 Comparison of summertime temperature anomalies for the Arctic and for lower
67 latitudes, averaged over at least millennia for key climatic intervals of the past, shows that
68 Arctic changes were threefold to fourfold larger than those in lower latitudes. This more
69 pronounced response applies to intervals that were both warmer and colder than in recent
70 decades. Arctic amplification of temperature changes thus appears to be a consistent feature of
71 the Earth system.

72

73 **Chapter 6 Rates of Change**

74 Climate changes have many causes and occur at different rates sustained for different
75 intervals. The changing atmospheric composition, and atmospheric and oceanic circulation
76 linked to tectonic processes during tens of millions of years have shifted the Arctic from ice-
77 free winters to icy summers. Features of Earth's orbit acting for tens of thousands of years have
78 rearranged sunshine on the planet and paced the growth and shrinkage of great ice-age ice
79 sheets. Anomalously cold single years have resulted from the influence of large, explosive
80 volcanoes, with slightly anomalous decades in response to the random variations in the
81 frequency of occurrence of such explosive volcanoes..

82 The local effects of these changes, as observed in Greenland or more generally around
83 the Arctic, yield trends such that more-persistent forcings have produced larger changes at a
84 lower average rate. Relative to these general trends, abrupt climate changes linked to shifts in
85 oceanic conditions of the North Atlantic produced anomalously large and rapid temperature
86 changes near the North Atlantic but relatively small average global temperature changes. And,
87 relative to these general trends, human-linked perturbations of the most recent decades do not
88 appear anomalously rapid or large, but changes projected as a part of the IPCC process can
89 become anomalously large and rapid.

90 Interpretation of these observations is complicated by lack of a generally accepted way
91 of formally assessing the effects or importance of size versus rate versus persistence of climate
92 change. The report here relied much more heavily on ice-core data from Greenland than would
93 be ideal in assessing Arctic-wide changes. Great opportunities exist for generation and synthesis
94 of other data sets to improve and extend the results here, using the techniques described in this

95 report. If widely applied, such research could remove the over-reliance on Greenland data.

96

97 **Chapter 7 The Greenland Ice Sheet**

98 Paleoclimate data show that the Greenland Ice Sheet has changed greatly with time and
99 has affected global sea level. Physical understanding indicates that many environmental factors
100 can force changes in ice-sheet size. Comparing histories of important forcings with ice-sheet
101 size implicates cooling as causing ice-sheet growth, warming as causing shrinkage, and
102 sufficiently large warming as causing loss. The evidence for temperature control is clearest for
103 temperatures similar to or warmer than those occurring in the last few millennia. The available
104 evidence shows less ice when snowfall was higher, indicating that snowfall rate is not the
105 leading control on ice-sheet size. Rising sea level tends to float marginal regions of ice sheets
106 and force their retreat, so the generally positive relation between sea level and temperature
107 means that, typically, both have pushed the ice sheet in the same direction. However, for some
108 small changes during the most recent millennia, marginal fluctuations in the ice sheet have been
109 opposed to those expected from local relative sea-level forcing but in the direction expected
110 from temperature forcing. This, plus the tendency for shrinkage to pull ice-sheet margins out of
111 the ocean, indicate that sea-level change has not been the dominant forcing at least for
112 temperatures similar to or greater than those of the last few millennia.

113 Histories of ice-sheet volume in fine time detail are not available, but the limited
114 paleoclimatic data at least agree that short-term and long-term response to temperature change
115 have been in the same direction. The best estimate from paleoclimatic data is thus that warming
116 shrinks the Greenland Ice Sheet, and warming of a few degrees is sufficient to cause ice-sheet
117 loss. Figure 7.13 shows a threshold for ice-sheet removal from sustained warming of 5°C, with

118 a range of uncertainties from 2° to 7°C, but tightly constrained numerical estimates are not
119 available, nor are rigorous error bounds, and the available data poorly constrain the rate of loss.
120 Numerous opportunities exist for additional data collection and analyses that would reduce the
121 uncertainties.

122

123 **Chapter 8 Arctic Sea Ice**

124 Geological data indicate that the history of Arctic sea ice is closely linked with
125 temperature changes. Sea ice in the Arctic Ocean may have appeared in response to long-term
126 cooling as early as 46 Ma. Year-round sea ice in the Arctic possibly developed as early as 13–
127 14 Ma, before the opening of the Bering Strait at 5.5 Ma. Nevertheless, extended seasonally ice-
128 free periods probably occurred until about 2.5 Ma. They ended with a large increase in the
129 extent and duration of sea-ice cover that more or less coincided with the onset of extensive
130 glaciation on land (within the considerable dating uncertainties). Some data suggest that ice
131 reductions marked subsequent interglacials and that the Arctic Ocean may have been seasonally
132 ice-free during the warmest events. For example, reduced-ice conditions are inferred for the last
133 interglacial and the onset of the current interglacial, about 130 and 10 ka .

134 Limited data suggest poorly understood variability in ice circulation for centuries to
135 millennia, but without strong periodic behavior on these time scales. Historical observations
136 suggest that ice cover has been shrinking since the late 19th century, and that the decline has
137 accelerated during the last several decades. This accelerated rate exceeds natural declines
138 typical of at least the most recent few millennia. This ice loss appears to be unrelated to natural
139 climatic and hydrographic variability on decadal time scales or to multi-millennial orbital
140 insolation changes.

141

142 **9.3 RECOMMENDATIONS**

143 Paleoclimatic data on the Arctic are generated by numerous international investigators
144 who study a great range of archives throughout the vast reaches of the Arctic. The value of this
145 diversity is evident in this report. Many of the key results of this report rest especially on the
146 outcomes of community-based syntheses, such as the CAPE Project, and on multiply replicated
147 and heavily sampled archives, such as the central Greenland deep ice cores. Results from the
148 ACEX deep coring in Arctic Ocean sediments were appearing as this report was being written;
149 these results were quite valuable and will become more so with synthesis and replication,
150 including comparison with land-based as well as marine records. The number of questions
151 answered, and raised, by this one new data set shows how sparse the data are on many aspects
152 of Arctic paleoclimate change. *We recommend that future research maintain and expand the*
153 *diversity of investigators, techniques, archives, and geographic locations, while promoting*
154 *development of community-based syntheses and multiply-replicated, heavily-sampled*
155 *archives; only through breadth and depth can the remaining uncertainties be reduced while*
156 *confidence in the results is improved.*

157

158 The questions asked of this study by the CCSP are relevant to public policy and require
159 answers. The answers provided here are, we hope, useful and informative. However, we
160 recognize that despite the contributions of numerous community members to this report, in
161 many cases a basis was not available in the refereed scientific literature to provide answers with
162 the accuracy and precision desired by policymakers. *We recommend that members of the*

163 *Arctic paleoclimatic community formulate future research activities to address in greater*
164 *detail the policy-relevant questions that motivated this report.*

165

166 Paleoclimatic data provide very clear evidence of past changes in important aspects of
167 the Arctic climate system. The ice of the Greenland ice sheet, smaller glaciers and ice caps, the
168 Arctic Ocean, and soils are shown to be vulnerable to warming, and Arctic ecosystems are
169 strongly affected by changing ice and climate. National and international studies generally
170 project rapid warming in the future. If this warming occurs, the paleoclimatic data indicate that
171 melting of ice and associated effects will follow, with implications for ecosystems and
172 economies. *We recommend that policymakers and science managers use the results presented*
173 *here in design of monitoring, process, and model-projection studies of Arctic change and*
174 *linked global responses.*

175

176 **Highlights of Key Findings**

- 177 • Arctic climate has changed greatly during the last 65 million years and
178 before, at highly varying rates and in response to many causes, with changing
179 atmospheric carbon-dioxide concentrations especially important in controlling
180 temperature.
- 181 • Arctic temperature changes have been larger than correlative globally
182 averaged changes, by approximately threefold to fourfold in both warmer and colder
183 times, in response to processes still active in the Arctic.
- 184 • Arctic temperatures have changed greatly but slowly in response to long-
185 lasting causes and by lesser amounts but more rapidly in response to other causes.

186 Human-forced changes of the most recent decades do not appear notably anomalous in
187 rate or size for their duration when they are compared with these natural changes, but
188 projections for future human-caused changes include the possibility of anomalously
189 large and rapid changes.

190 • The Greenland Ice Sheet has consistently grown with cooling and shrunk
191 with warming, and a warming of a few degrees (about 5°C, with uncertainties between
192 about 2° and 7°C) or more has been sufficient to completely or almost completely
193 remove the ice sheet if maintained long enough; the rate of that removal is poorly
194 known. Reduction in the size of the Greenland Ice Sheet in the past has resulted in sea
195 level rise.

196 • Warming has decreased sea ice, which in turn strongly magnifies
197 warming, and seasonally ice-free conditions and even year-round ice-free conditions
198 have occurred in response to sufficiently large but poorly quantified forcing.

199 • Although major climate changes have typically affected the whole Arctic,
200 important regional differences have been common; a full understanding of Arctic
201 climatology and paleoclimatology requires regionally-resolved studies.

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