

Air-sea heat fluxes in the stratocumulus deck / cold tongue / ITCZ complex of the eastern tropical Pacific

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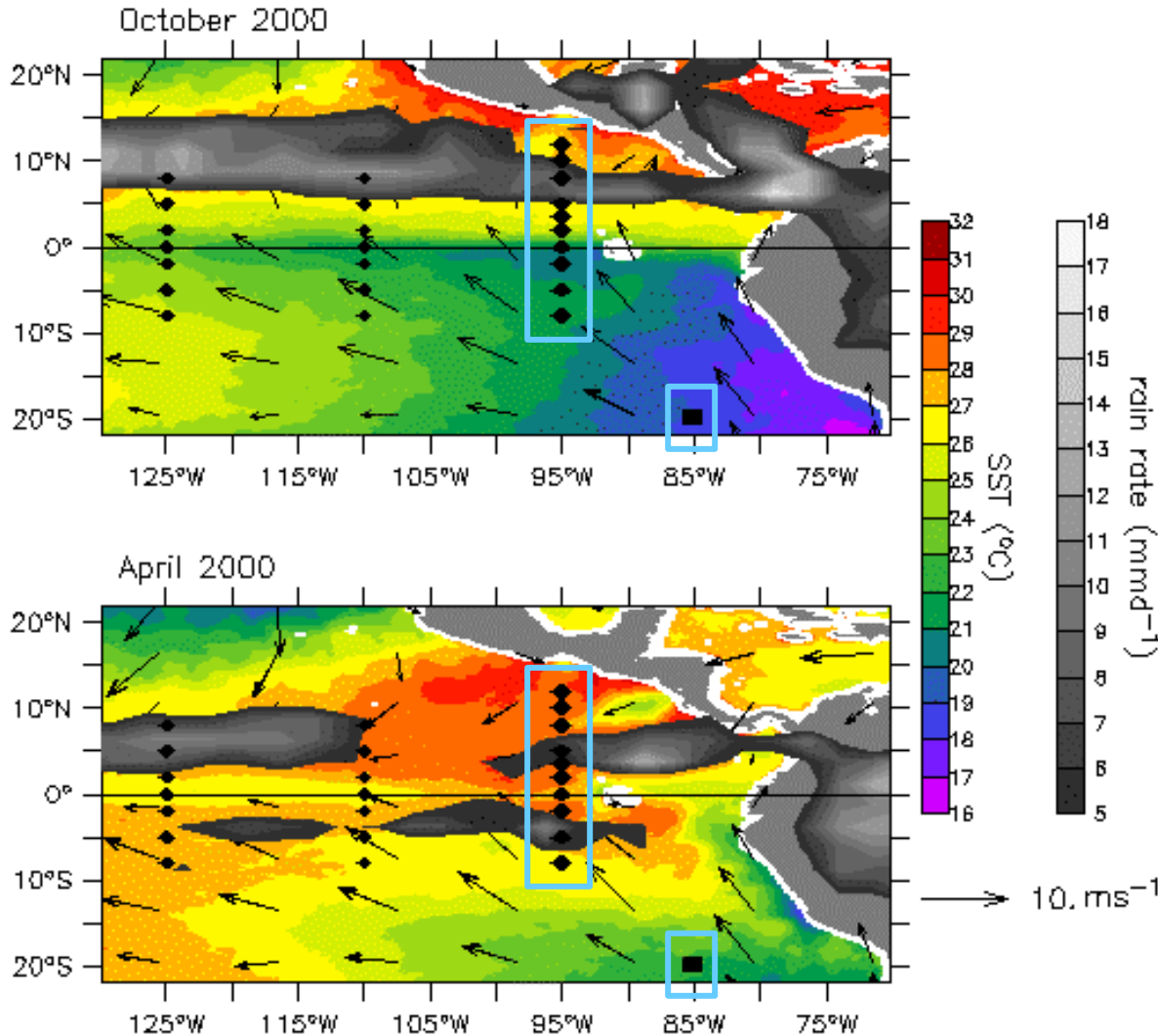
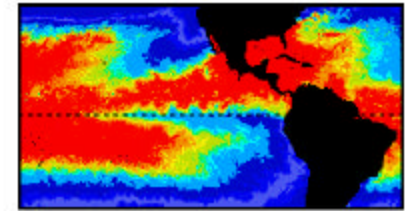
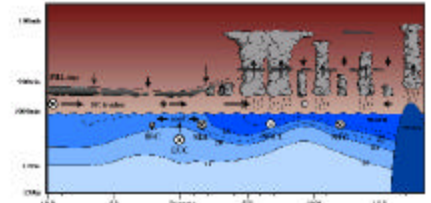
Chris Fairall (NOAA ETL)

Michael J. McPhaden (NOAA PMEL)

Robert Weller (WHOI)

Eastern Pacific Investigation of Climate (EPIC) experiment

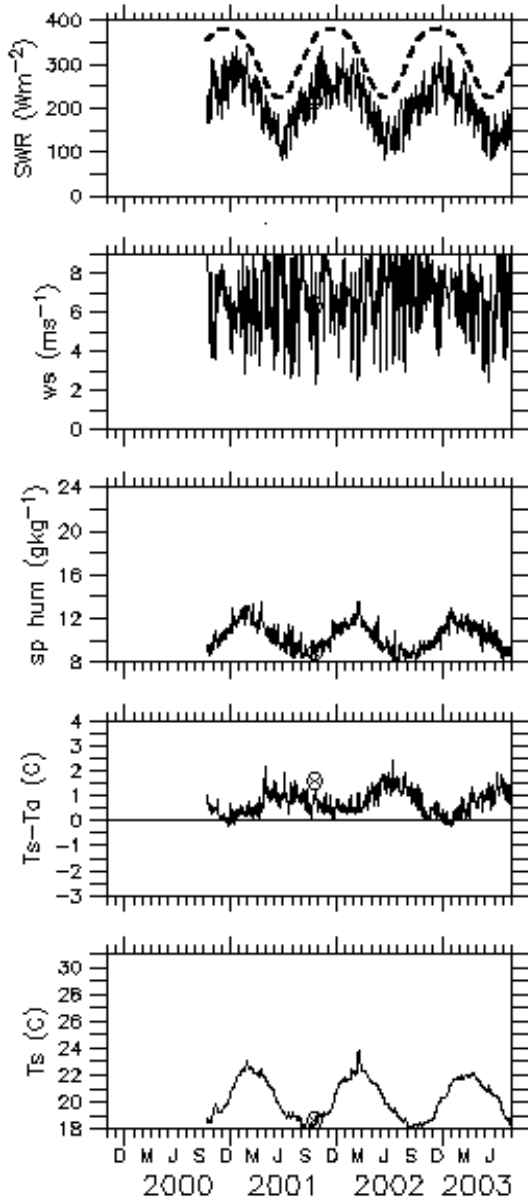
A Science and Implementation Plan for EPIC:
An Eastern Pacific Investigation of Climate Processes
in the Coupled Ocean-Atmosphere System



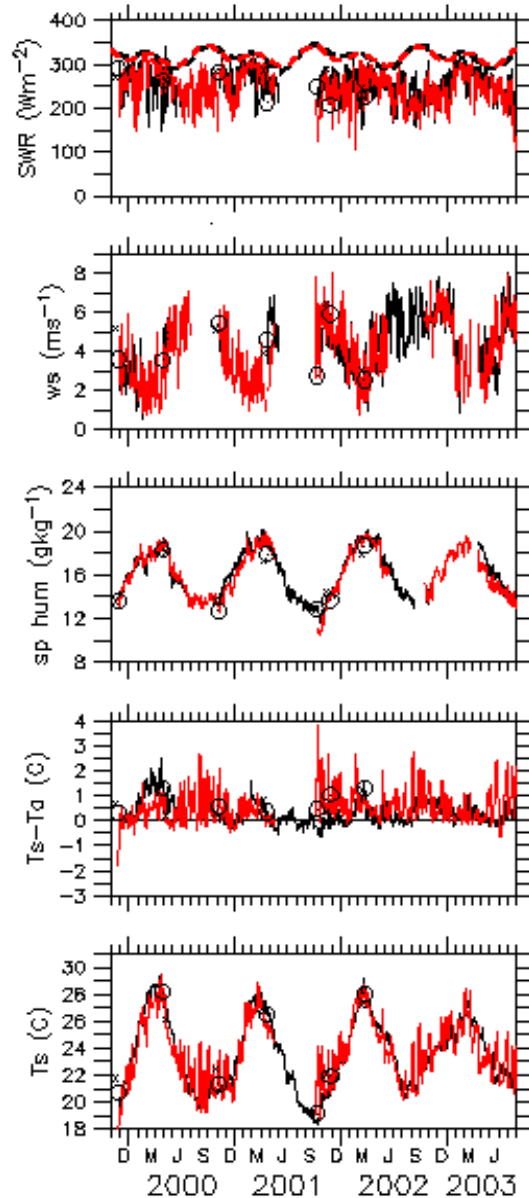
This study uses ship and buoy data along 95W and at 20S, 85W to evaluate air-sea heat fluxes in the stratus deck / cold tongue / ITCZ complex.

Buoy and ship time series

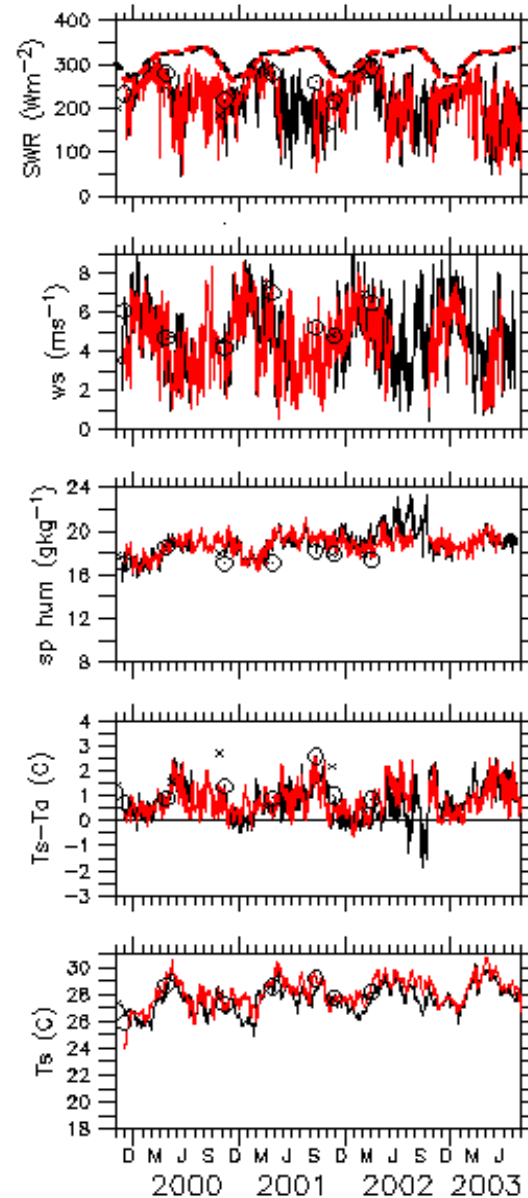
STRATUS
20S85W



COLD TONGUE
2S95W 0N95W

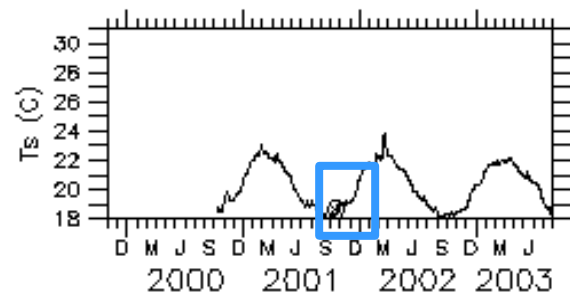
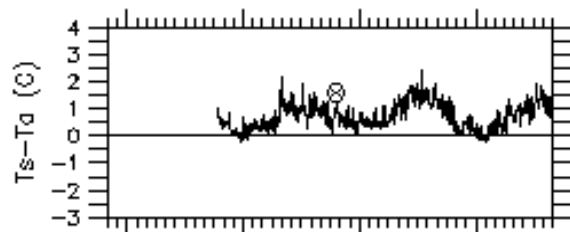
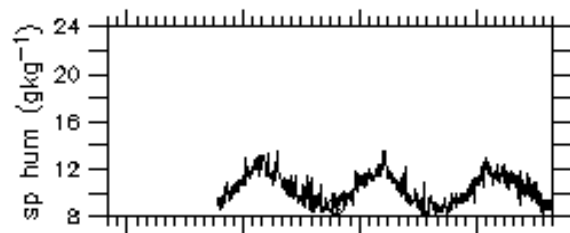
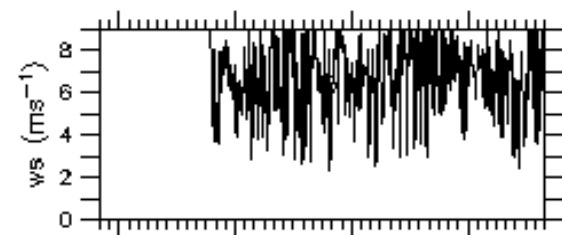
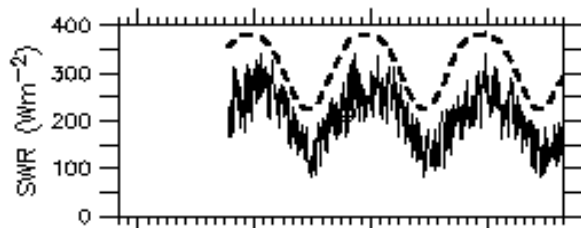


ITCZ
8N95W 10N95W



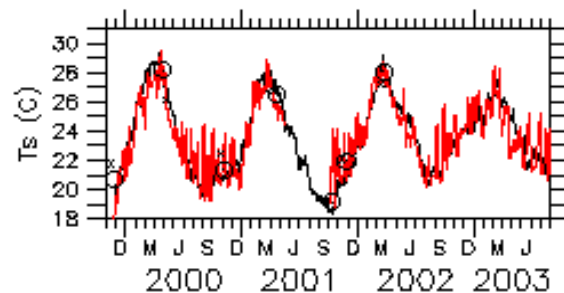
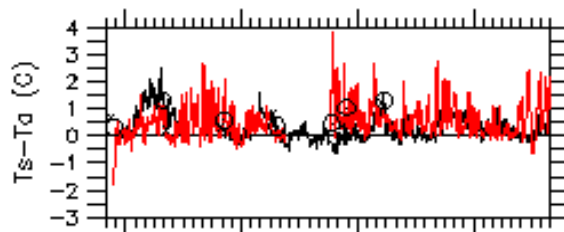
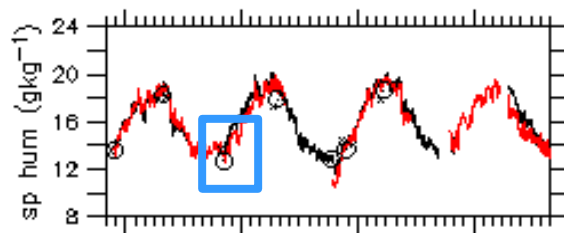
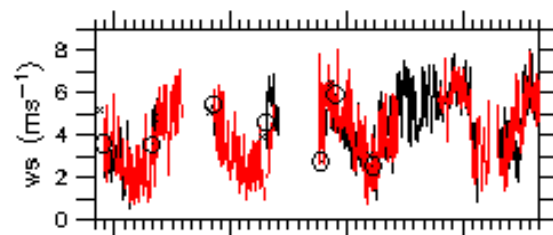
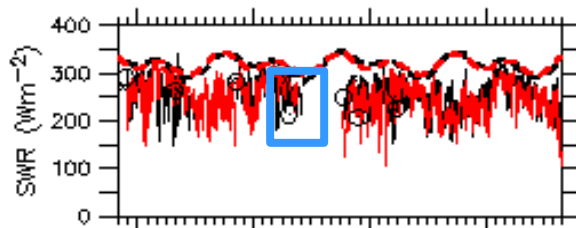
STRATUS

20S85W



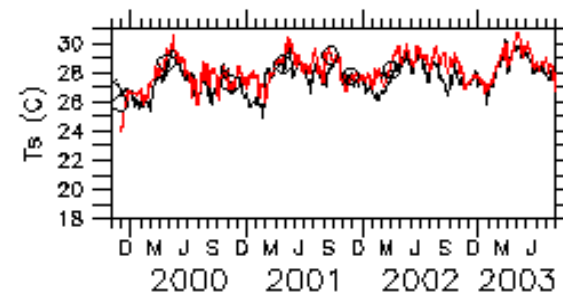
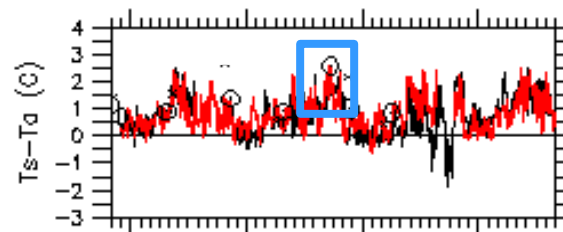
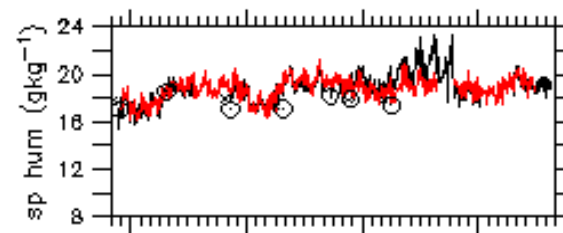
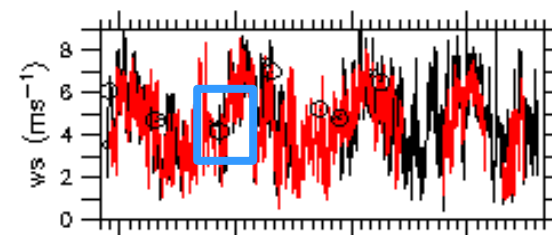
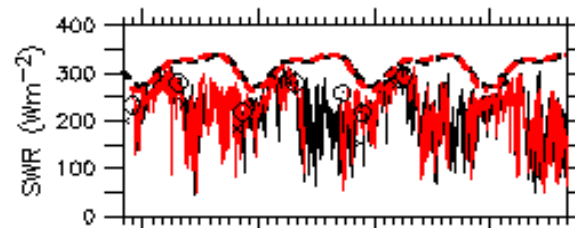
COLD TONGUE

2S95W 0N95W



ITCZ

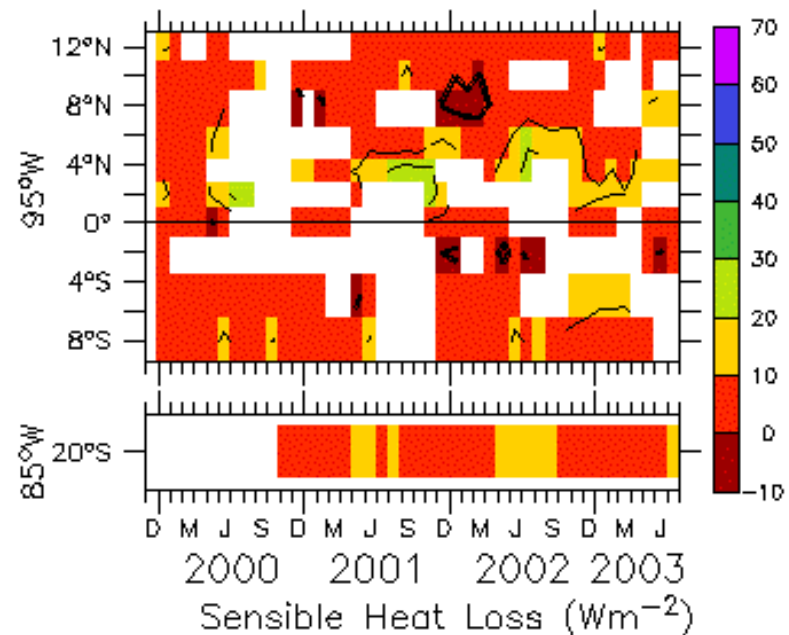
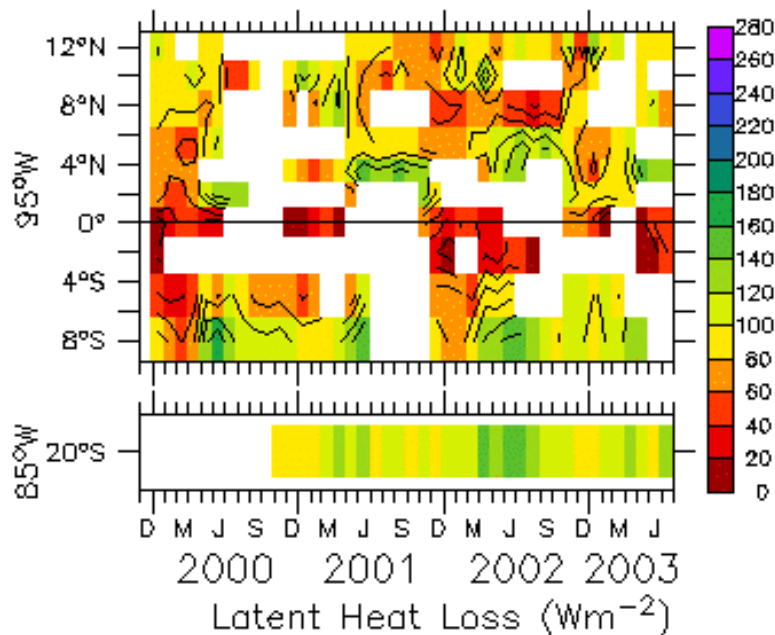
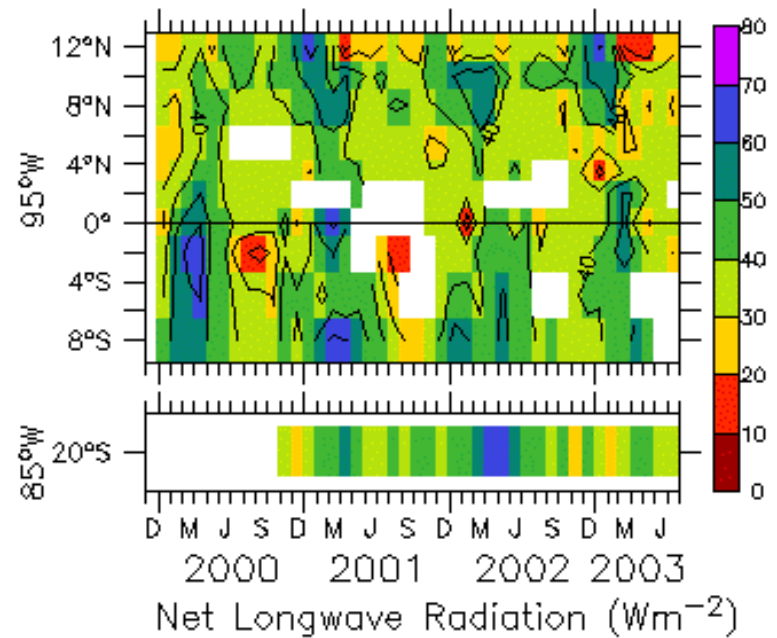
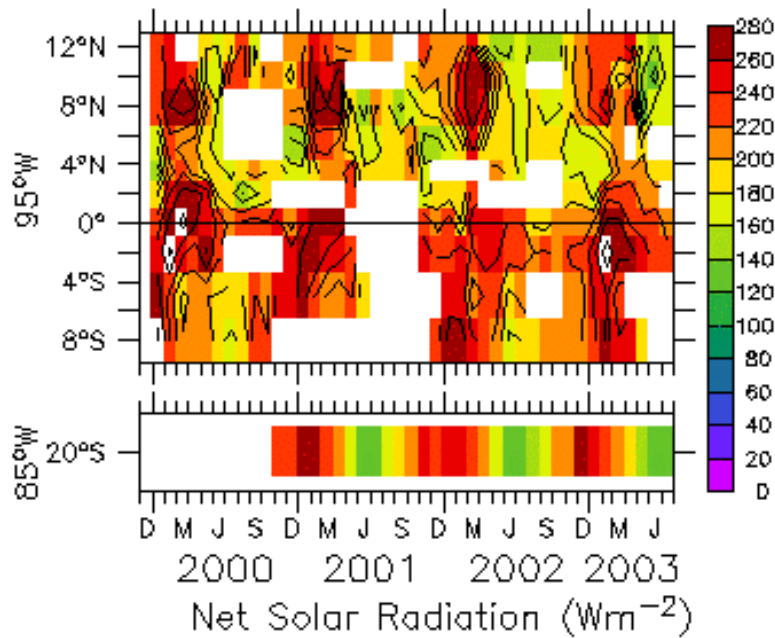
8N95W 10N95W



Latent and sensible heat flux calculations from buoy data

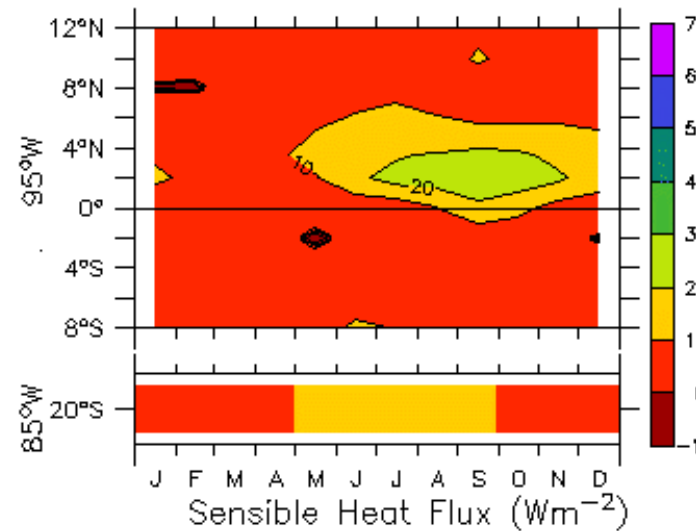
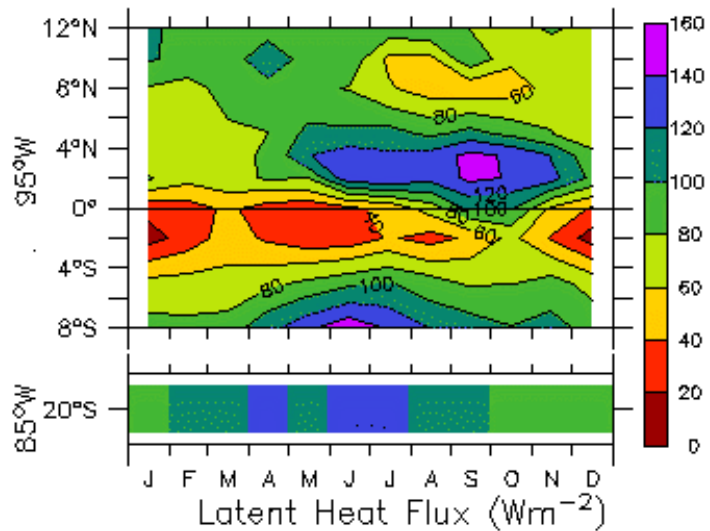
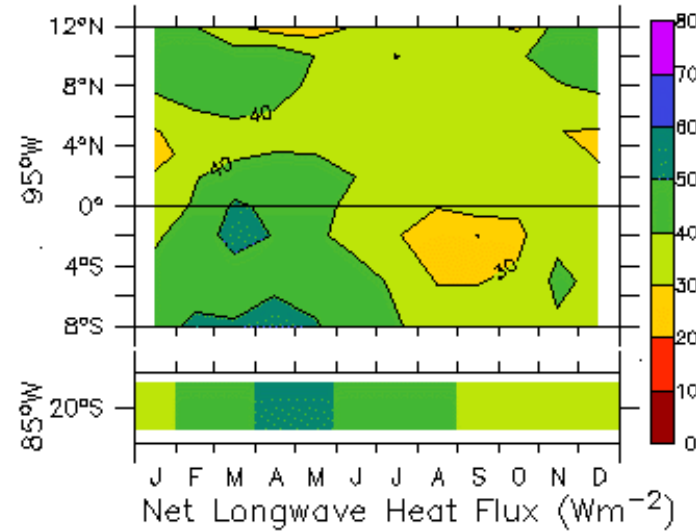
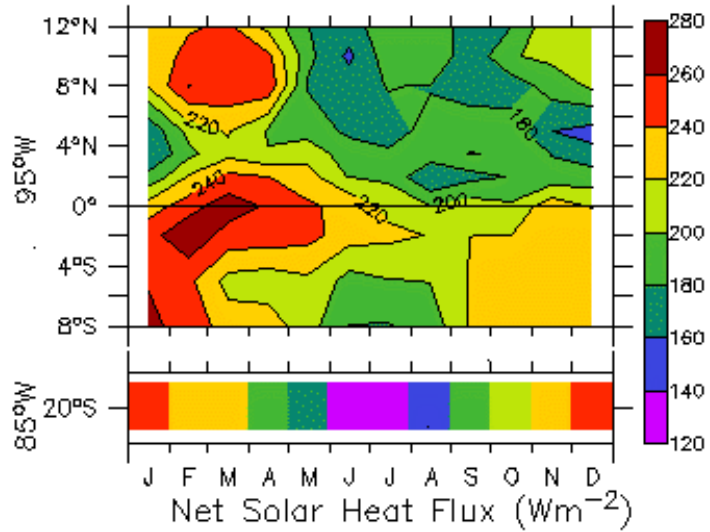
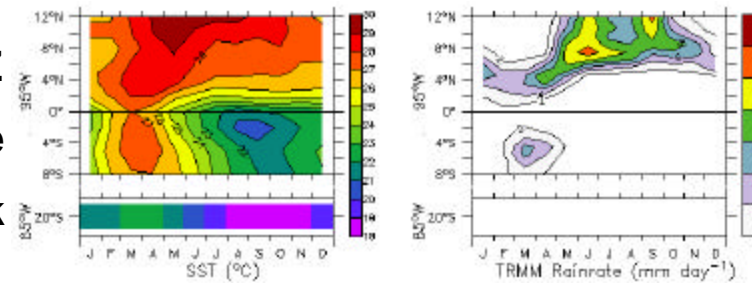
- Used Fairall et al. (2003) v3.0a bulk flux algorithm.
- Used hourly-averaged data (fill gaps with telemetered daily-averaged data).
- Applied warm layer and cool skin corrections to extrapolate 1 m SST to surface (T_{skin}).
- Used winds relative to surface currents.

$$Q_0 = Q_{sw} - Q_{lw} - Q_{lat} - Q_{sen}$$



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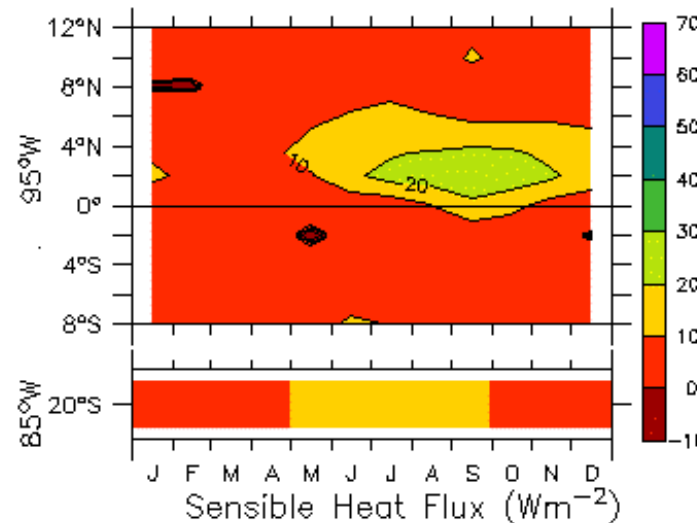
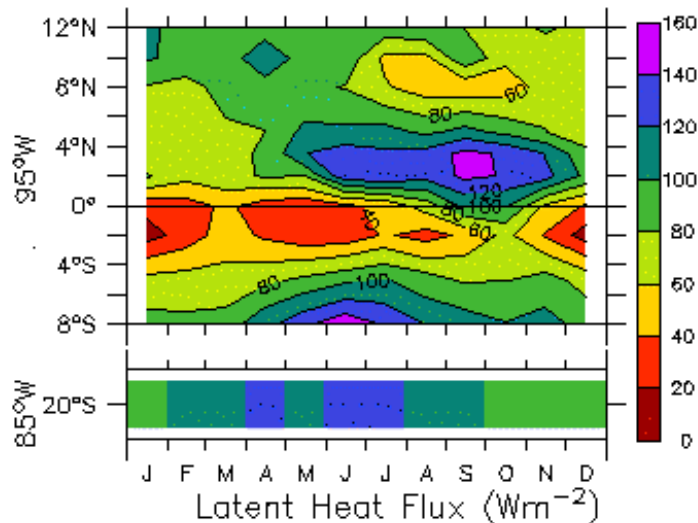
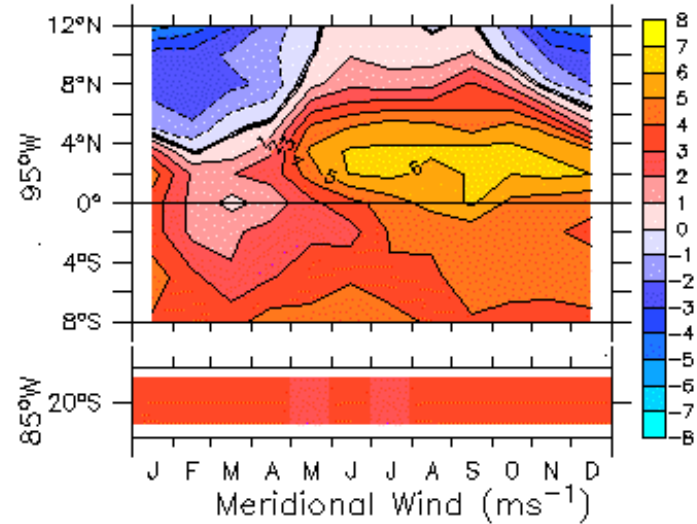
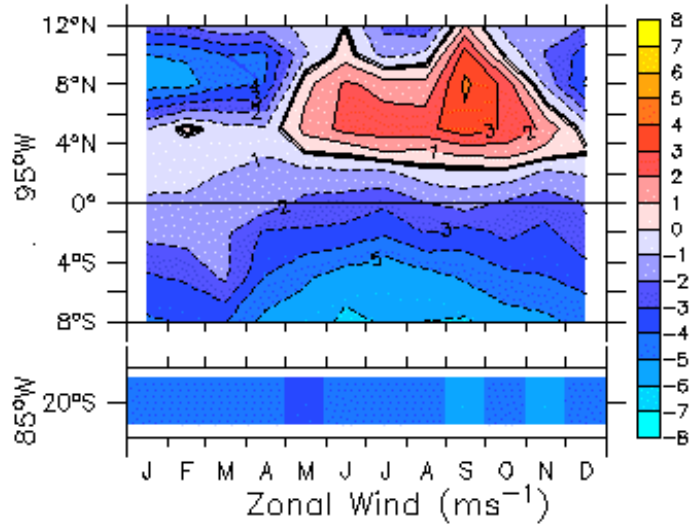
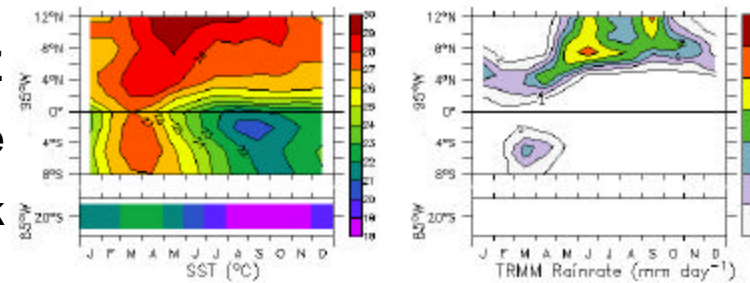
ITCZ
cold tongue
stratus deck



- $Q_0 \sim Q_{sw} - Q_{lat}$
- Q_{lat} and Q_{sen} max over frontal region
- Q_{lat} near zero over cold tongue and is weakest during warm season.

PBL stabilized over cold tongue and destabilized over frontal region

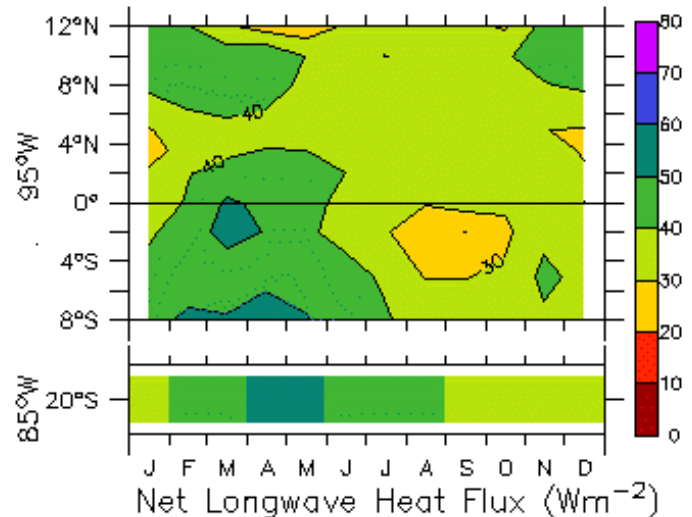
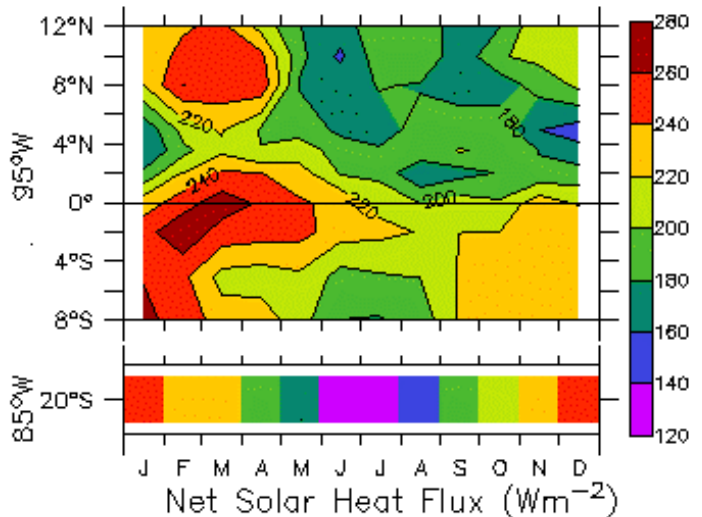
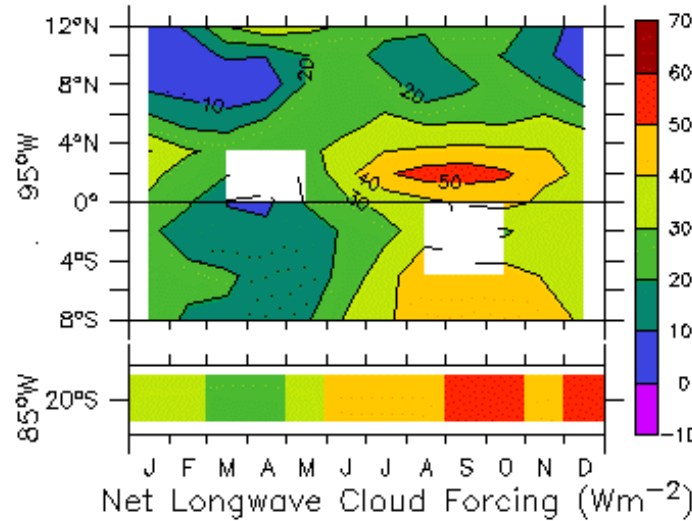
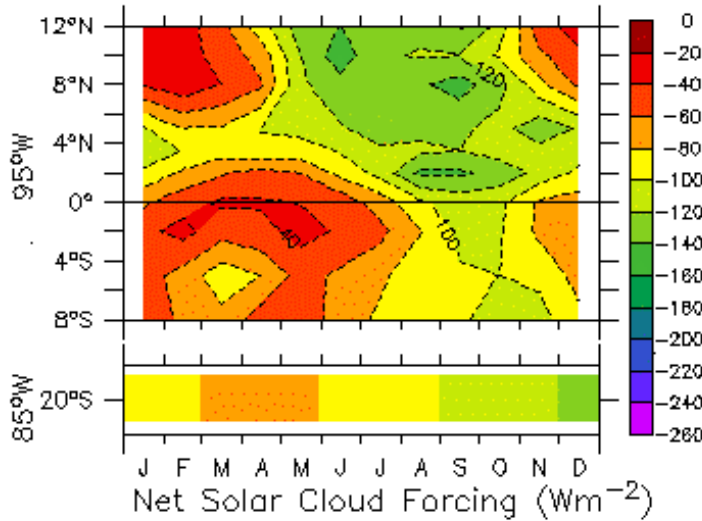
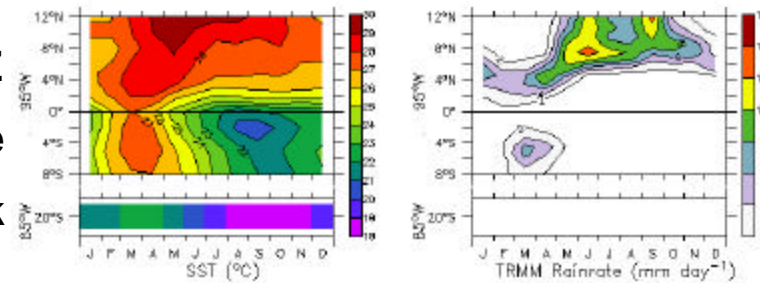
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Cloud Forcing is reduction or enhancement in surface radiation caused by clouds

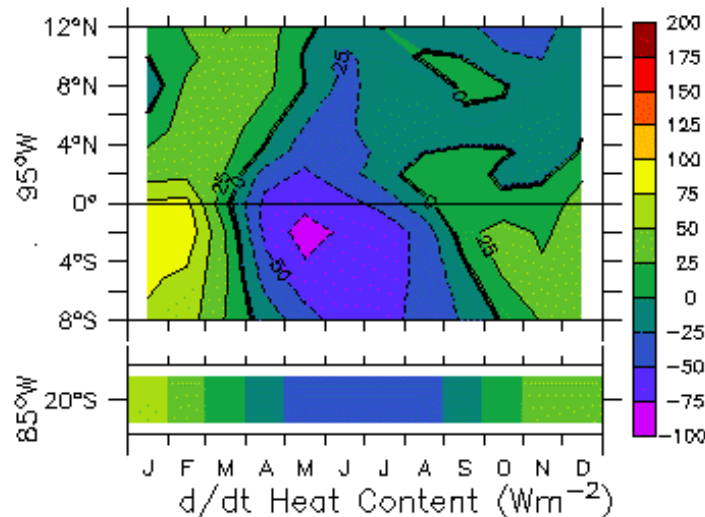
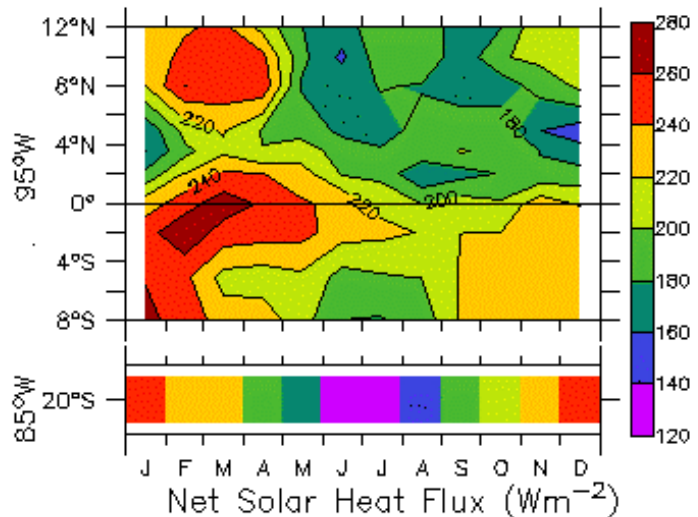
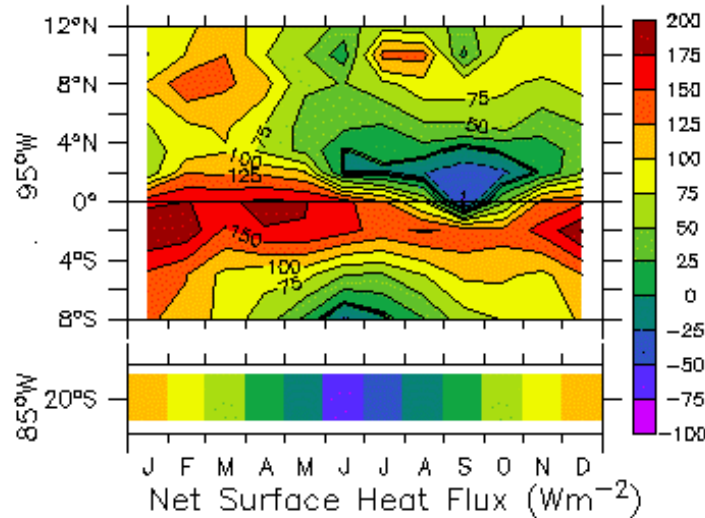
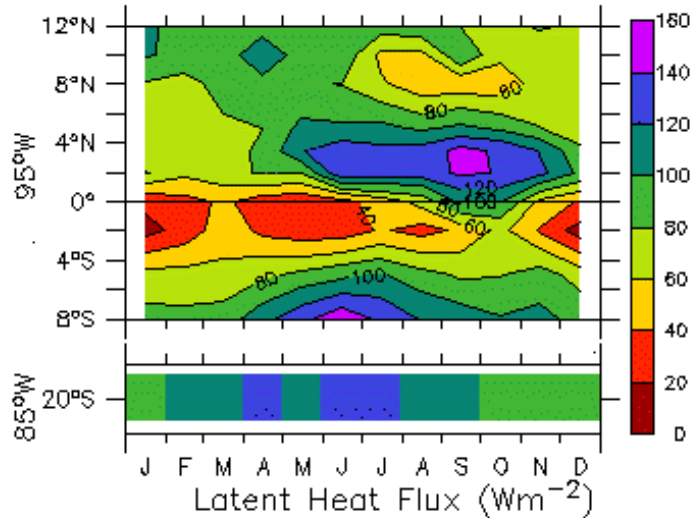
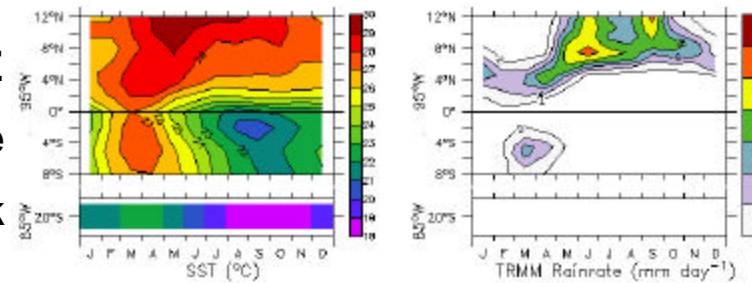
ITCZ
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- $Q_0 \sim Q_{sw} - Q_{lat}$
- Q_{lat} and Q_{sen} max over frontal region
- Q_{lat} near zero over cold tongue and is weakest during warm season.
- Q_{sw} and Q_{lw} modulated by clouds.

$Q_0' \sim Q_{sw} - Q_{lat} \sim ? d(HC)/dt$
 $HC = 30 \text{ m Heat Content}$

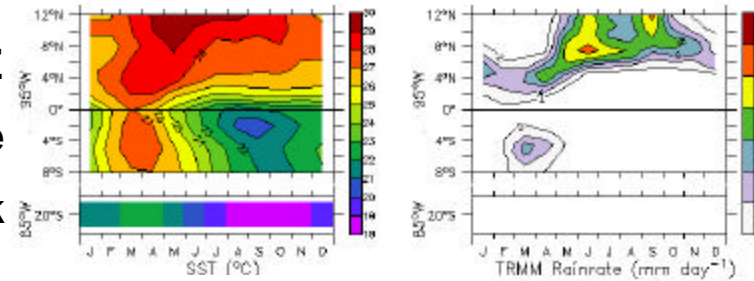
ITCZ
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 stratus deck



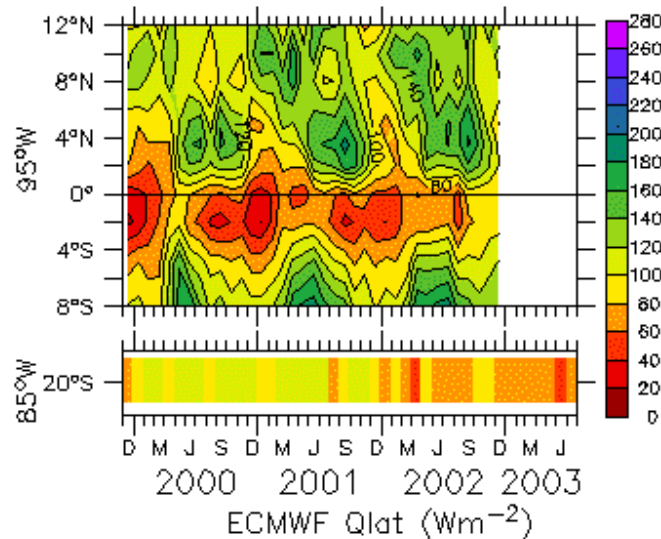
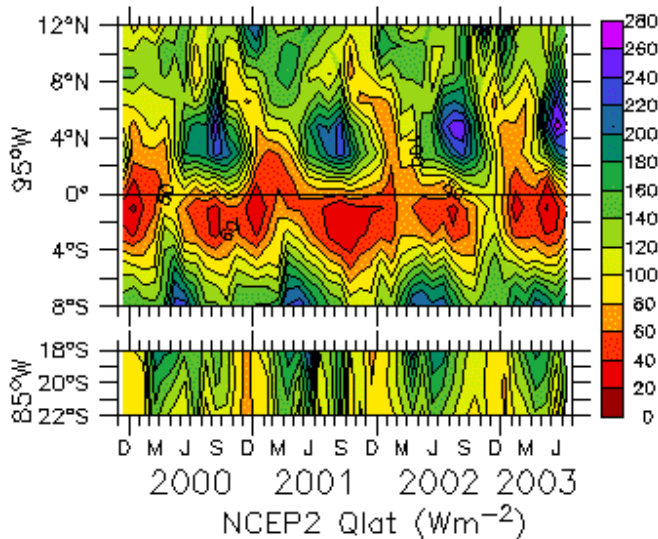
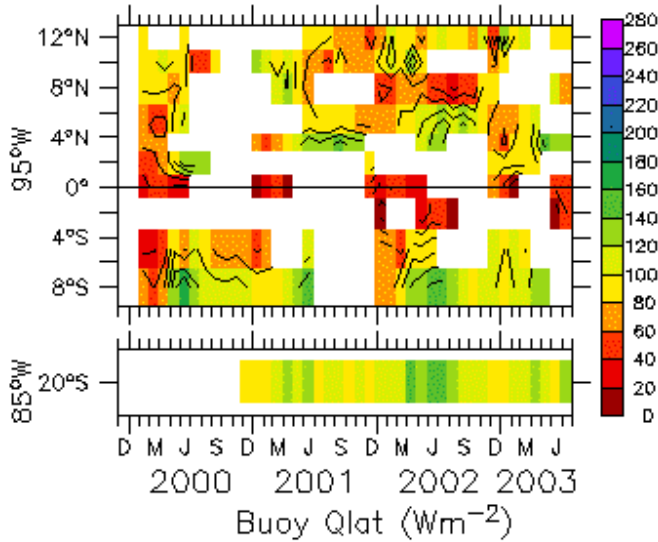
- $Q_0 \sim Q_{sw} - Q_{lat}$
- Q_{lat} and Q_{sen} max over frontal region
- Q_{lat} near zero over cold tongue and is weakest during warm season.
- Q_{sw} and Q_{lw} modulated by clouds.
- Ocean processes also important

How do NWP latent heat fluxes compare to buoy?

ITCZ
cold tongue
stratus deck

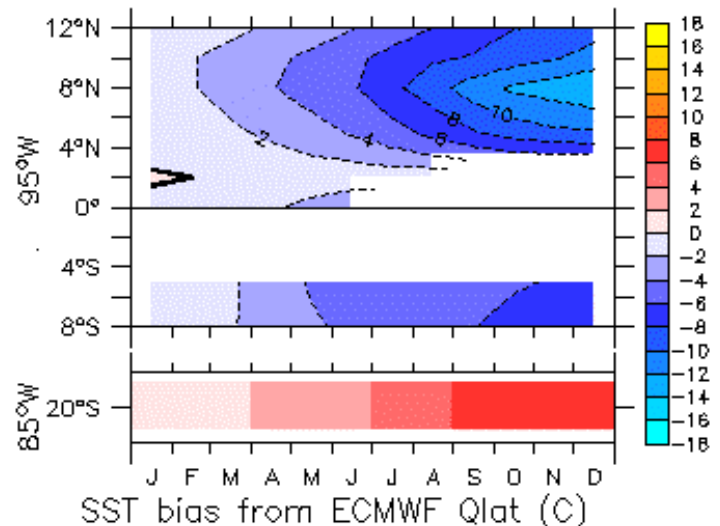
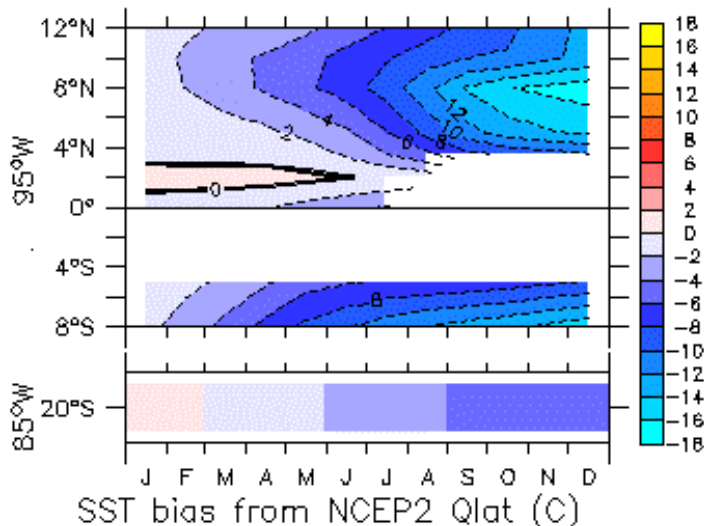
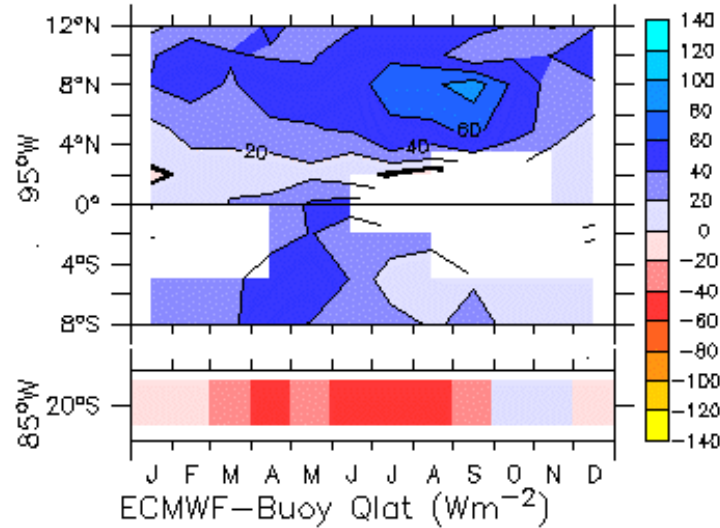
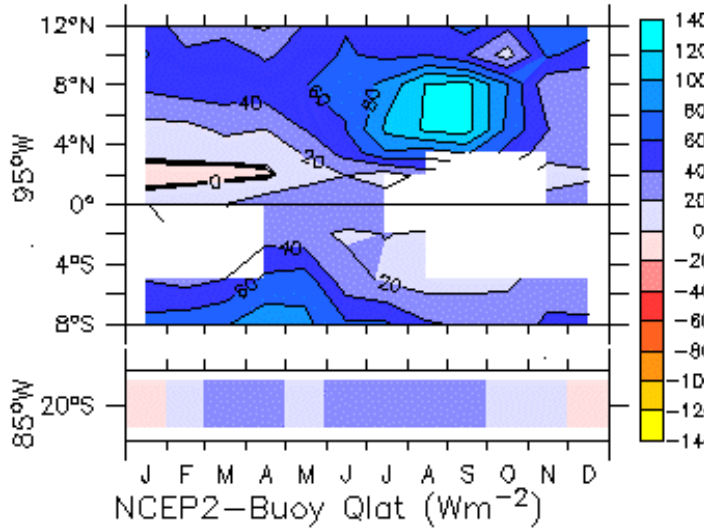
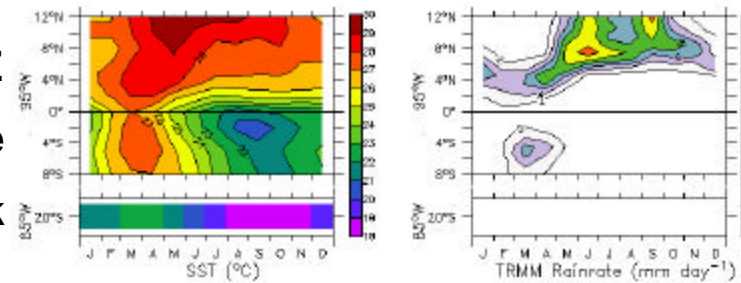


- NCEP2 latent heat flux is much stronger than buoy (cold bias)
- ECMWF is more realistic



What are SST biases associated with NWP latent heat flux “bias”?

ITCZ
cold tongue
stratus deck



- SST biases are **HUGE** when integrated over a year.
- ECMWF has a warm bias in stratus deck region and cold bias in cold tongue / ITCZ complex.
- NCEP2 has a cold bias everywhere.

Conclusions

- **Stabilized PBL causes low latent heat loss over cold tongue; and destabilized PBL causes maximum latent heat loss over frontal region.**
- **Solar radiation, latent heat loss and ocean processes control SST variability in the east Pacific stratus deck / cold tongue / ITCZ complex.**
- **ECMWF latent heat loss appears to be more realistic than NCEP2.**
- **ECMWF has a warm bias (due to latent heat flux) in stratus deck region and cold bias in cold tongue / ITCZ complex. NCEP2 has a cold bias everywhere.**
- **SST biases associated with latent heat biases are HUGE when integrated over a year.**