

Figure 10. Schematic vertical section showing the stacked accreted terranes of the Klamath Mountains, and the intrusive relations of the plutons.

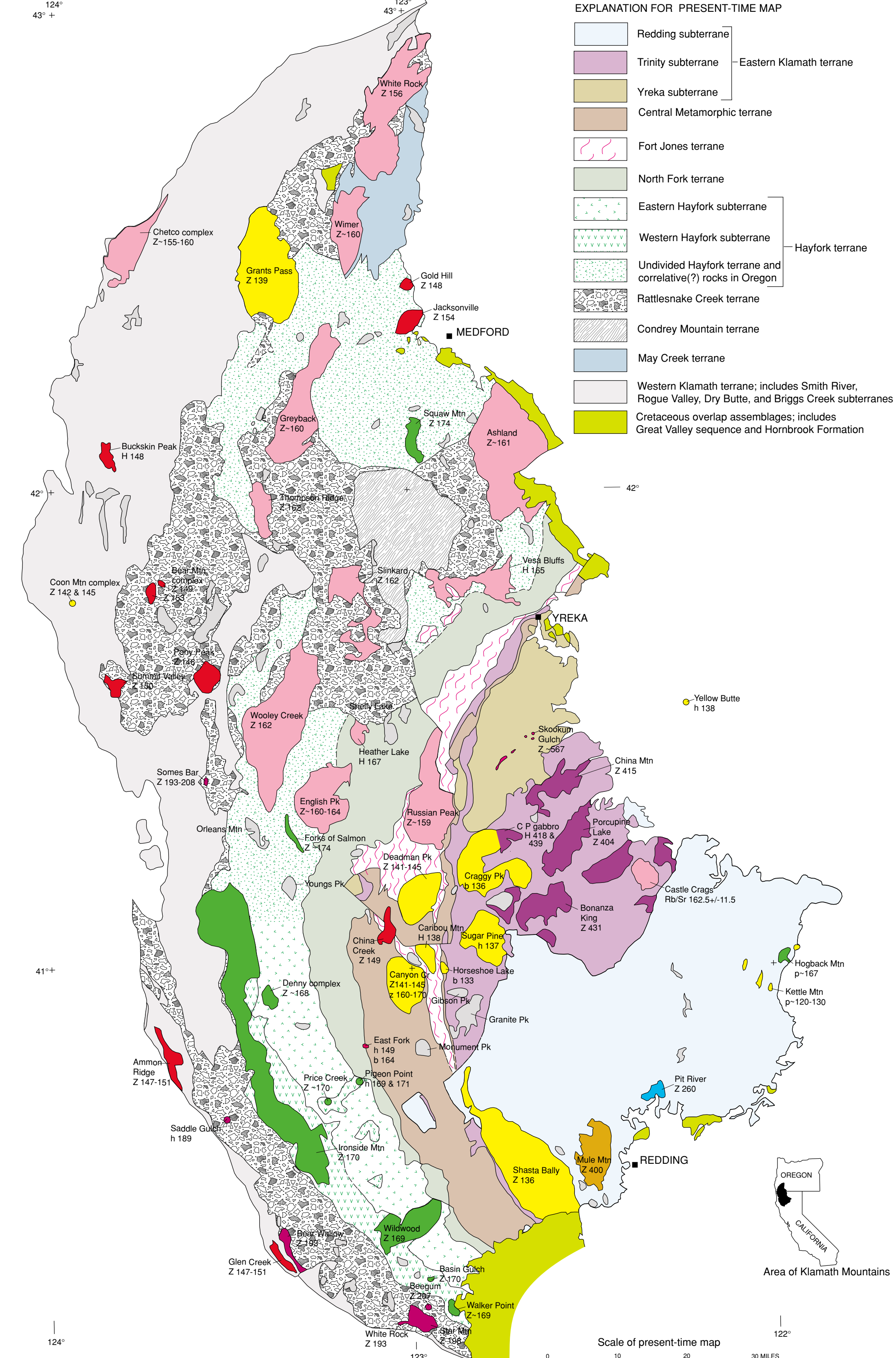
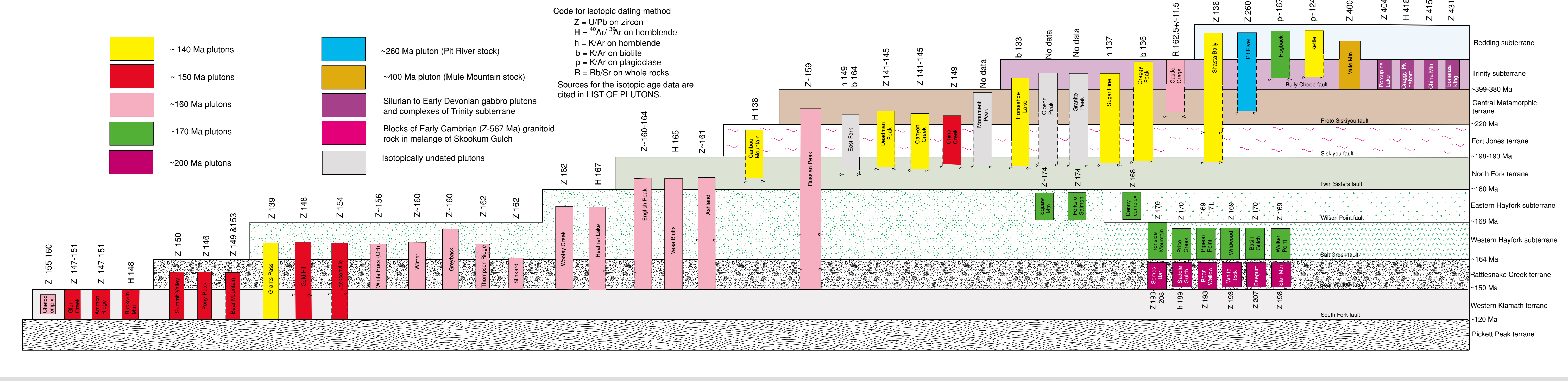


Figure 11. GEOLOGIC TIME SCALE (Modified from Palmer, 1985), showing the approximate timing of the sequential accretionary episodes that formed the Klamath Mountains of California and Oregon. The generalized stratigraphic column of the Redding subterranean is shown for comparison. The dominantly volcanic formations are shaded. The possibly Silurian part of the Copley Greenstone is not shown.

LIST OF PLUTONS

Name	Isotopic age (Ma)	Reference
Ammon Ridge	Z 147-151	Wright and Fahnestock, 1988
Adrian	Z 140	Vale, 1996
Basin Chert	Z 170	Wright and Fahnestock, 1988
Bear Mtn complex	Z 148 & 153	Safety and Harper, 1993
Bear Valley	Z 163	Wright, 1981
Begun	Z 207	Wright, 1981
Bonanza King	Z 411	Wright and Metcalf, 1996
Buckskin Peak	H 148	Harper and others, 1994
Canby Creek	Z 141-145; 160-170	Wright and Fahnestock, 1988
Castle Crags	H 158	Hecker and Ernst, 1993
Chico complex	Z 155-160	R. W. Kistler; pers. commun., 1999
China Creek	Z 149	J. L. Wooden, This report
China Mountain	Z 415	Wright and Metcalf, 1996
Coon Mtn complex	Z 142 & 145	Safety and others, 1985; Harper and others, 1994
Craggy Peak	H 136	Langphere and others, 1968
Craggy Peak gabbro	H 418 & 439	Langphere and others, 1968
Deadman Peak	Z 141-145	Wright and Fahnestock, 1988
Denny complex	Z 188	Wright and Fahnestock, 1988
East Fork	H 148 & 164	M. A. Langphere, in Irwin, 1985
English Peak	Z 100-164	Wright and Fahnestock, 1988
Forks of Salmon	Z 174	Wright and Fahnestock, 1988
Gibson Peak	No date	Wright and Fahnestock, 1988
Glenn Creek	Z 147-151	J. L. Wooden, This report
Gold Hill	Z 148	J. L. Wooden, This report
Granite Peak	Z 140	Harper and others, 1994
Grant Peak	Z 169	Vale, 1996
Grayback	H 167	Hecker and Ernst, 1993
Hauser Lake	H 167	Renne, 1987
Hogback Mountain	p-167	Iverson and Kistler, 1970
Inside Mountain	H 133	Wright and Fahnestock, 1988
Jacksonville	Z 154	J. L. Wooden, This report
Kettle Mountain	p-120-130	Renne, 1987
Monument Peak	No date	Albers and others, 1981
Mule Mountain	Z 400	Wright and Fahnestock, 1988
Olefin Mountain	No date	Wright and Fahnestock, 1988
Pike Point	H 169 & 171	Wright and Fahnestock, 1988
Pit River	H 167	Wright and Fahnestock, 1988
Pony Peak	Z 140	Safety and Harper, 1993
Provisional Lake	Z 404	Wright and Fahnestock, 1988
Price Creek	Z 176	Wright and Fahnestock, 1988
Russian Peak	Z 139	Wright and Fahnestock, 1988
Saddle Gulch	H 189	M. A. Langphere, in Irwin, 1985
Shasta Baldy	Z 156	Langphere and Jones, 1978
Shelly Lake	No date	
Skookum Gulch	Z 507	Vallin, 1990
Sinkard	Z 182	Wright and Harper, 1993
Somes Bar	Z 192-208	Wright, 1981 (PCK-81)
Squaw Mountain	Z 174	J. L. Wooden, This report
Star Mountain	Z 188	Wright and Fahnestock, 1988
Sugar Pine	H 137	Langphere and others, 1968
Sunnyside Valley	Z 150	Safety and Harper, 1993
Thompson Ridge	Z 162	Vale, 1996
Walker Butte	H 165	Hecker and Ernst, 1993
Walker Point	Z 169	Wright and Fahnestock, 1988
White Rock (OR)	Z 156	J. L. Wooden, This report
White Rock (CA)	Z 165	Wright, 1981 (OR-1)
Wildwood	Z 169	Wright and Fahnestock, 1988
Yreka	Z 160	Vale, 1996
Woolly Creek	Z 162	Safety and Harper, 1993
Woolly Butte	H 138	Holz, 1971

Figure 1. Central Metamorphic episode Late Silurian(?) to Middle Devonian time (>400 Ma—380 Ma) Figure 2. Fort Jones episode Permian—Triassic time (~280—240Ma) Figure 3. North Fork episode Early Jurassic (Pliensbachian) time (~198—193 Ma) Figure 4. Eastern Hayfork episode Early(?) Middle Jurassic (~Bajocian) time (~180 Ma) Figure 5. Western Hayfork episode Late Middle Jurassic (Early Callovian) time (~168 Ma) Figure 6. Rattlesnake Creek episode Late Middle to early Late Jurassic (Callovian to Oxfordian) time (~164 Ma) Figure 7. Western Klamath episode Late Jurassic (Late Kimmeridgian or Early Tithonian) time (~150-2 Ma) Figure 8. Pickett Peak episode Early Cretaceous time (~140-5 Ma) Figure 9. Present time Geology modified from Irwin (1994)

INTRODUCTION
The Klamath Mountains consist of various accreted terranes and include many plutons that range in composition from gabbro to granodiorite. Some of the plutons (preaccretionary plutons) were parts of terranes before the terranes accreted, others (accretionary plutons) intruded during or after the accretion of their host terranes. This report attempts to (1) graphically illustrate how the Klamath Mountains grew by the accretion of allochthonous oceanic terranes during early Paleozoic to Cretaceous times, (2) identify the plutons as either preaccretionary or accretionary, and (3) genetically relate the plutonic intrusions to specific accretionary episodes.

The eight accretionary episodes portrayed in this report are similar to those shown by Irwin and Mankinen (1989) who briefly described the basis for the timing of the episodes and who illustrated the ~110 degrees of clockwise rotation of the Klamath Mountains since Early Devonian time. Each episode is named for the accretionary terrane. In all episodes (Figs. 1-8), the heavy black line represents a fault that separates the accretionary oceanic rocks on the left from earlier accreted terranes on the right. The accretionary plutons are shown within the accretionary oceanic crustal rocks to the left of the heavy black line, and the accretionary plutons in most instances shown intruding previously accreted terranes to the right. Episodes earlier than the Central Metamorphic episode (Fig. 1), and that may have been important in the formation of the early Paleozoic nucleus of the province (the Eastern Klamath terrane), are not known. Also not described in the sequential Figures 1 to 8 are the May Creek and Conroy Mountain terranes. The Present Time distribution of the accreted terranes and plutons is shown at a large scale in Figure 9.

The schematic vertical section (Fig. 10) depicts the terranes as a stack of horizontal slabs that include or are intruded by vertical plutons. All of the slabs are shown as extending to the right-hand edge of the schematic section, but this is not meant to imply that there is evidence that the structurally lower slabs extend to beneath the eastern limits of the uppermost slabs. Note that their base is the ~170 Ma preaccretionary plutons of the Western Hayfork subterranean are truncated by the ~160 Ma Salk Creek accretionary fault, the ~160 Ma accretionary plutons are truncated by the ~150 Ma Bear Willow accretionary fault, and the ~150 Ma accretionary plutons (and questionably the Grants Pass ~139 Ma accretionary pluton) are truncated by the ~140 Ma South Fork fault.

The isotopic ages shown for the plutons are mostly from various published sources (see LIST OF PLUTONS). Most K-Ar/Pb ages measured on zircon, which are thought to most reliably represent the crystallization ages of the plutons. Isotopic ages measured on hornblende, biotite, or plagioclase by other methods are used where U/Pb ages are not available. K. W. Kistler generously provided a Rb-Sr isotopic age for Castle Crags pluton. New U/Pb isotopic ages for this study were measured on zircon from five plutons by J. L. Wooden. Many small plutons in the Klamath Mountains are not isotopically dated.

The Eastern Klamath terrane, which consists of the Redding, Trinity, and Yreka subterraneans, was the last to be accreted. The stratigraphy of the Redding subterranean is a remarkable record of sedimentation, volcanism, and plutonism from Silurian-Devonian to Jurassic time, and is the longest span of time of virtually continuous stratigraphy of any terrane in the Klamath Mountains. Notably, the principal volcanic plutonic units of the Redding subterranean are dated to generally coincide in time with accretionary episodes, which suggests that the accretionary episodes caused coeval volcanism and plutonism in the Redding subterranean. The Copley Greenstone-Balakala Rhyolite volcanics, intruded by the coeval Mule Mountain pluton, are correlative with the Central Metamorphic episode; the Dekkas Andesite-Bully Hill Rhyolite volcanics and Pit River stock (and Redding dike) probably correlate with the Fort Jones episode; and the Arvonson Fin-Bagley Andesite volcanics and Hogback Mountain intrusives correlate with the North Fork episode (see Fig. 11). Both the early Paleozoic Balakala Rhyolite and the late Paleozoic Bully Hill Rhyolite were important in the formation of the Redding subterranean. The Redding stratigraphic column spans from Silurian-Devonian to Middle Jurassic time (Fig. 11). During this ~230 m.y. life-span as an intermittently active volcanic arc the Redding subterranean was intruded by only two large granitoid plutons, the Mule Mountain stock (400 Ma) and Pit River stock (260 Ma), and several smaller intrusives (Fig. 9). It was later that the large Cretaceous granitoids such as Shasta Baldy batholith (136 Ma) intruded the Redding and other subterraneans.

THE ACCRETIONARY EPISODES
Central Metamorphic episode (Fig. 1). Protoliths of the Central Metamorphic terrane were subducted beneath the Eastern Klamath terrane. The Eastern Klamath terrane was the nucleus of the accretionary episode. The accretionary plutons were intruded by the ~160 Ma Salk Creek accretionary fault, the ~160 Ma accretionary plutons are truncated by the ~150 Ma Bear Willow accretionary fault, and the ~150 Ma accretionary plutons (and questionably the Grants Pass ~139 Ma accretionary pluton) are truncated by the ~140 Ma South Fork fault.

The Klamath Mountains in Silurian-Devonian time and consisted of the Redding, Trinity, and Yreka subterraneans. During the Central Metamorphic episode, the Redding subterranean developed as a Silurian-Devonian volcanic arc, built on peridotite upper mantle of the Trinity subterranean. [Alternatively, Vallin and Metcalf (1988) consider these rocks a supra-subduction zone ophiolite. The volcanic rocks of the arc are Copley Greenstone and Balakala Rhyolite (gray area, Fig. 1). They were intruded by a coesitic granitoid pluton, the Mule Mountain stock (Z=400 Ma). Several large gabbroic plutons and complexes intrude the peridotite of the Trinity subterranean. Fort Jones accretionary episode was originally considered Late Triassic based on the ~200 Ma K-Ar isotopic age of the blueschist (Irwin and Mankinen, 1989). However, the episode is herein considered Permian-Triassic(?) based on a probable coeval relation with the Pit River-Redding dike intrusives and the Nosoqi Dekkas-Bully Hill volcanics of the Redding subterranean.

North Fork episode (Fig. 3). Protoliths of the North Fork terrane were overridden along the Sikkiyou and correlative faults by the Fort Jones, Central Metamorphic, and Eastern Klamath terranes. Note that south of Deadman Peak pluton the Fort Jones terrane is exposed only through an antiformal window in Central Metamorphic terrane (Figs. 3 and 9). The protoliths of the North Fork terrane were a subduction complex that included dimembered ophiolite, mafic volcanic and sedimentary rocks, radiolarian chert, and minor limestone. They range from late Paleozoic to Jurassic. The youngest fossiliferous rocks are radiolarian chert of Early Jurassic (Pliensbachian) age (Blome and Irwin, 1983), which is the basis for the approximate age of 198-193 Ma assigned to the North Fork accretionary episode. The North Fork terrane contains no preaccretionary plutons, but is intruded by several large ~160 Ma plutons related to a younger (Rattlesnake Creek) accretionary episode. However, in the Redding subterranean, the plutonism is represented by only two large granitoid plutons, the Mule Mountain stock (400 Ma) and Pit River stock (260 Ma), and several smaller intrusives (Fig. 9). It was later that the large Cretaceous granitoids such as Shasta Baldy batholith (136 Ma) intruded the Redding and other subterraneans.

Fort Jones episode (Fig. 2). The Eastern Klamath and Central Metamorphic terranes overrode a subduction complex along the proto Sikkiyou fault and its northern extensions, forming the Fort Jones (aka Stuart Fork) terrane. South of Deadman Peak pluton, the Fort Jones terrane is exposed only in an antiformal window in Central Metamorphic terrane. Note the small structural outliers (see Fig. 8) of Eastern Klamath terrane resting on Central Metamorphic terrane. The age of the Fort Jones terrane is not clearly known. Fossils are not found except for vestiges of radiolarians in some of the least metamorphosed chert. Whole-rock samples yielded K-Ar isotopic ages of 133 and 158 Ma (Langphere and others, 1988). K-Ar isotopic ages of ~220 Ma (Late Triassic) were measured on blueschist facies metamorphic rocks in the northern part of the Fort Jones terrane (Holtz and others, 1977), but these and especially the younger K-Ar ages are suspect. No preaccretionary plutons are recognized in the accretionary rocks. The oldest pluton intruding the Fort Jones terrane is the Russian Peak pluton (Z=159 Ma). However, in Redding subterranean the accretionary Pit River stock (Z=260 Ma) intrudes Bald Fin (Miss. Penn.), and a chain of irregular linear intrusions of quartz-satellite diorite, the Redding dike of Hinks (1933), extends northward 15 or more miles (Fig. 9). The Redding dike extends mostly along the contact between the Bald Fin and McCoold Limestone, locally engulfing the McCoold Limestone and locally intruding the Nosoqi Fin and Dekkas Andesite (Hinks, 1933). Although not isotopically dated, the Redding dike is evidently Permian or younger and may well be the feeder for the Nosoqi-Dekkas-Bully Hill (Permian-Triassic?) volcanic formations (gray area, Fig. 3). The age of the accretionary episode is poorly constrained and is estimated to be ~180 Ma (Bajocian), approximately midway between the preceding and following episodes. No plutonism for this episode is recognized.

Western Hayfork episode (Fig. 5). Protoliths of the Western Hayfork subterranean were subducted beneath the Eastern Hayfork subterranean along the Wilson Point and correlative faults. The protoliths were a volcanic arc consisting mainly of the Hayfork Baldy Meta-andesite and the Ironside Mountain batholith and other ophiolite plutons. These preaccretionary rocks range in isotopic ages from 188 to 177 Ma (Fahnestock, 1982; Langphere and others, 1968; Wright, 1981), the youngest of which is assumed to approximate the age of subduction. The Denny Complex (Fig. 5) is assumed to be accretionary to the Eastern Hayfork subterranean. In the central and northern parts of the Klamath Mountains, the Eastern Hayfork and Western Hayfork terranes are not clearly distinguished and are shown together as undivided Hayfork terrane (see Figs. 9 and 10). As previously mentioned in Figure 5, the Hogback Mountain suite of small plutons (p-168 and 167 Ma) is here shown in the Jurassic strata of the Redding subterranean, and if the isotopic ages are correct, would be considered accretionary plutons of the Western Hayfork episode.

Rattlesnake Creek episode (Fig. 6). Protoliths of the Rattlesnake Creek terrane were subducted beneath the Western Hayfork subterranean along the Salk Creek and correlative faults. The protoliths were a mélange of mainly ophiolitic rocks, bodies of limestone, radiolarian chert, minor blocks of amphibolite, and Early Jurassic (~200 Ma) preaccretionary plutons. The limestone and chert contain late Paleozoic, Triassic, and Early to Middle Jurassic fossils. The chert is mostly Late Triassic and Early to Middle Jurassic, some of which may be as young as Bathonian. The time of subduction is broadly constrained to Callovian-Oxfordian (~164 Ma) by Middle Jurassic fossils and by ~160 Ma accretionary plutons that cut the subduction zone fault.

Western Klamath episode (Fig. 7). Protoliths of the Western Klamath terrane were subducted beneath the Rattlesnake Creek terrane along the Bear Willow and correlative faults. They consisted mainly of the Josephine ophiolite, the Rogue and Galice Formations, and several other radiolarian chert, limestone, and scattered blocks of schist. Fossils in the chert and limestone are mainly Late Permian and Triassic. None is clearly Jurassic. Some of the Late Permian fossils are of Thelyon faunas that are unknown to most of North America. The age of the accretionary episode is poorly constrained and is estimated to be ~180 Ma (Bajocian), approximately midway between the preceding and following episodes. No plutonism for this episode is recognized.

Western Klamath episode (Fig. 8). Protoliths of the Western Hayfork subterranean were subducted beneath the Eastern Hayfork subterranean along the Wilson Point and correlative faults. The protoliths were a volcanic arc consisting mainly of the Hayfork Baldy Meta-andesite and the Ironside Mountain batholith and other ophiolite plutons. These preaccretionary rocks range in isotopic ages from 188 to 177 Ma (Fahnestock, 1982; Langphere and others, 1968; Wright, 1981), the youngest of which is assumed to approximate the age of subduction. The Denny Complex (Fig. 5) is assumed to be accretionary to the Eastern Hayfork subterranean. In the central and northern parts of the Klamath Mountains, the Eastern Hayfork and Western Hayfork terranes are not clearly distinguished and are shown together as undivided Hayfork terrane (see Figs. 9 and 10). As previously mentioned in Figure 5, the Hogback Mountain suite of small plutons (p-168 and 167 Ma) is here shown in the Jurassic strata of the Redding subterranean, and if the isotopic ages are correct, would be considered accretionary plutons of the Western Hayfork episode.

Pickett Peak episode (Fig. 9). Protoliths of the Pickett Peak terrane were subducted beneath the Western Klamath terrane along the South Fork and correlative faults. The Western Klamath terrane overrode westward the Franciscan(?) sedimentary and volcanic rocks along the South Fork fault, forming the South Fork Mountain Schist and Choptaquim Metabasalt of the Pickett Peak terrane of the California Coast Ranges. Structural outliers of correlative schist (Colebrook Schist) lie west of the Klamath Mountains in the Coast Ranges of Oregon. The isotopic (K-Ar) age of the South Fork Mountain Schist is ~120-115 Ma (Langphere and others, 1978), but a metamorphic episode that age is too young to have generated the ~140 Ma plutons. Thus the K-Ar isotopic age is suspect and the Pickett Peak episode is tentatively considered early Early Cretaceous (~140-5 Ma). Alternatively, the ~140 Ma plutons may represent either a late phase of the Western Klamath episode or an unrecognized accretionary episode.

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PLUTONS AND ACCRETIONARY EPISODES OF THE KLAMATH MOUNTAINS, CALIFORNIA AND OREGON

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