

APPENDIX B

CLASSIFICATION OF METAMORPHIC AND OTHER COMPOSITE- GENESIS ROCKS, INCLUDING HYDROTHERMALLY ALTERED, IMPACT-METAMORPHIC, MYLONITIC, AND CATACLASTIC ROCKS

Version 1.0

North American Geologic-map Data Model Steering Committee
Science Language Technical Team (SLTT)
Composite-Genesis Subgroup

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1 1 EXECUTIVE SUMMARY

2 A provisional draft classification system for metamorphic rocks and other composite-genesis
3 rocks is proposed as part of a comprehensive scheme for use in digital geologic maps and
4 computer-based geoscience information systems in North America. The purpose of the rock
5 classification is to identify lithologic constituents of map units or to search digital geologic-map
6 databases for specific rock types.

7 The domain of this classification system includes not only metamorphic rocks as commonly
8 understood, but also a variety of hydrothermally altered, mylonite-series, cataclastic, and impact-
9 metamorphic rocks. We classify these composite-genesis rocks according to descriptive
10 properties that reflect the multiple processes that make the rock composite. This classification is
11 fundamentally descriptive, based mainly on rock composition and fabric; therefore, different
12 observers should be able to classify a rock in the same way. Assignment to a lithologic class
13 implies that certain descriptive criteria are met, and these criteria must be defined in the database.
14 This classification also attempts to follow common usage of terms in the geoscience community,
15 reconciling that usage and database requirements wherever necessary.

16 This report summarizes the logic and rationale for the classification system, which is presented in
17 a series of four diagrammatic decision trees, two class hierarchy diagrams, and a glossary in the
18 form of a Microsoft Access database that has definitions and parent-child relationships for terms
19 in the classification system. Comments from the geoscience community on this proposed
20 classification are encouraged.

21 2 INTRODUCTION

22 This document (hereafter referred to as SLTMM_1.0) proposes a provisional classification system
23 for metamorphic rocks and other composite-genesis rocks as part of a comprehensive scheme for
24 use in digital geologic maps and computer-based geoscience information systems in North
25 America (<http://geology.usgs.gov/dm/>). The purpose of an Earth material terminology system (or
26 controlled vocabulary, see introduction to NADM science language) is to provide a basis for
27 identifying and describing materials that are the substances of the Earth by classifying them using
28 standard terms and definitions. Classification links the normative description (definition) for a
29 term to the material being described. Rock-lithology classification has two important functions in
30 geoscience information systems. First, in the case of rock descriptions, a lithologic constituent of
31 a geologic-map unit may be identified using a standard rock classification in cases for which a
32 detailed description of the rock is unavailable or unnecessary. Second, in cases for which
33 lithologic constituents are described in detail, assignment of the constituent to a standard rock
34 classification allows users to search for standard kinds of rock without having to design queries
35 that analyze the complete description structure. A lithology field in a geoscience database is a
36 place for classification of the constituents used to describe bodies of rock.

37 This report describes the rationale for the accompanying classification system, which is presented
38 in a series of four linked decision trees (flow charts), two class hierarchy diagrams, and an
39 accompanying glossary in the form of a Microsoft Access database that has definitions and
40 parent-child relationships for terms in the classification system for composite-genesis rocks.

41 The domain of this classification system includes composite-genesis rocks, defined as rocks that
42 are the product of more than one rock-forming process. They are not purely igneous or
43 sedimentary. This domain includes metamorphic rocks as commonly understood, as well as
44 impact metamorphic rocks, hydrothermally altered rocks, mylonite-series rocks, and cataclastic
45 rocks. These composite-genesis rocks are classified strictly according to descriptive properties
46 related to the changes that made them ‘composite.’

47 2.1 Composite-genesis rocks

48 2.1.1 Definition

49 *Composite-genesis rock*—Any rock having observable features that document mineralogical,
50 chemical, or structural change of a preexisting earth material essentially in the solid state. It
51 includes metamorphic (*sensu strictu*), hydrothermally altered, cataclastic, and impact-
52 metamorphic rock, but not weathering products or soils.

53 A metamorphic rock has observable features that document change after the original formation of
54 the rock, under physical or chemical conditions that differ from those normally occurring at the
55 surface of the Earth and in zones of cementation and diagenesis below the surface (Smulikowski
56 and others, 1997). Hydrothermally altered rock has fabric and composition indicating solid-state
57 mineralogical and chemical changes in response to hot, mineral-rich waters, and is included as
58 ‘metamorphic rock’ in this classification. Cataclastic rock, impact-metamorphic rock, and
59 composite-genesis melt rocks are treated as special classes. Where possible, metamorphic
60 classification schemes proposed by the British Geological Survey (Robertson, 1999) and
61 preliminary recommendations of the IUGS Subcommittee on the Systematics of Metamorphic
62 Rocks (SCMR) (Schmid and others, 2002) were adapted to meet SLTT database requirements.

63 The composite-genesis rocks are classified along two orthogonal dimensions, fabric and
64 composition (Figs. 1 and 2). Class hierarchy is a directed acyclic graph rather than a tree.
65 Classes that have composition and fabric criteria are ‘children’ of both a ‘composition and fabric’
66 lithology parent, and of a generic composition and a generic fabric parent. Class names have a

67 fabric component (such as granofels) and a compositional component (modifier such as marble or
68 quartzofeldspathic).

69 2.2 Background and purpose

70 A standard North American database model for the input, storage, manipulation, retrieval, and
71 analysis of digital geologic-map information is being developed as a cooperative effort by a
72 consortium of interests including the Association of American State Geologists (AASG), the U.S.
73 Geological Survey (USGS), the Geological Survey of Canada (GSC), and the Canadian
74 Provincial Surveys. The standards for this data model are being developed under the auspices of
75 the multi-constituency North American Data Model Steering Committee (NADMSC,
76 <http://nadm-geo.org>). The NADMSC has commissioned technical teams, including the Science
77 Language Technical Team (<http://nadm-geo.org/slitt>), to develop aspects of the data model. Thus,
78 the purpose of the Science Language Technical Team (SLTT) is to develop standard terminology
79 for classifying and describing geologic materials for digital geologic-map databases produced
80 throughout North America. The Science Language Technical Team includes four subgroups: the
81 Subgroup on Plutonic Rocks, the Subgroup on Volcanic Rocks, the Subgroup on Sedimentary
82 Materials, and the Subgroup on Metamorphic Rocks. This draft report was produced by the
83 Subgroup on Metamorphic Rocks.

84 2.3 Definitions: classification, naming, and description

85 **Classification** is the assignment of an instance (individual object from some point of view) to a
86 group (class) that is defined based on a shared set of properties. It answers the question “what
87 kind of X is Y?” where X represents the domain of the classification. A class of things is defined
88 by the set of the properties shared by members of the class. Boundaries of the class are defined
89 by threshold property values.

90 **Naming** is assignment of an identifier to an instance (object within a class). Ideally, every
91 instance would have a unique name, but of course in the realm of earth and planetary materials,
92 this is not the case. Formalized stratigraphic nomenclature is an example of naming mappable
93 units. Lithologic naming provides an identifier for a particular rock or Earth (and planetary)
94 material that allows geologists to communicate when formal nomenclature is not defined, or
95 when they need to subdivide in more detail than the formal nomenclature allows. Names are also
96 assigned to identify individual classes in a classification system (each class is an instance of a
97 class!).

98 **Description** is a set of statements that characterize the nature of a thing (a class or instance) such
99 that the thing may be identified. The set of shared properties that define a class constitute a
100 description.

101 2.4 Conceptual framework

102 Lithologic classification is the assignment of a rock or unconsolidated material to a named class
103 defined based on its physical properties. The scope of the classification, ‘rock or unconsolidated
104 material’, corresponds to the concept ‘CompoundMaterial’ as defined by the North American
105 Geologic Map Data Model Steering Committee Data Model Design Team (NADMSC, 2003). In
106 this report, the term ‘material’ is used in place of the more precise, but less familiar, term
107 ‘CompoundMaterial’ to mean ‘rock or unconsolidated material’. Note also that this concept may
108 be extended to include not only materials found on Earth (as defined by NADMSC, 2003), but
109 also any planetary material. Lithologic classification is used in a number of contexts. The field
110 geologist classifies rocks as part of the process of defining mappable bodies of rock within a
111 geographic area. A map compiler uses the classification of rocks that compose geologic map
112 units to determine similarity between units on different maps, or as criteria to define composite
113 map units that combine features of more detailed maps. A non-geologist uses the classification

114 system to identify rocks of interest without having to study their descriptions in detail. Earth
115 scientists use standard classification systems to characterize rocks as part of the process of
116 describing them. These applications highlight two sorts of classification process—one aimed at
117 identifying particular bodies of rock in a particular region, and one aimed at grouping similar
118 kinds of rock that may be present in many places.

119 Lithologic classification is scale dependent. Hand-sample-dimension lithologic classification
120 systems are designed to group kinds of material into named classes based on a 1-cm to 30-cm
121 diameter representative volume. A particular rock name, based on a naming scheme such as that
122 proposed by the British Geological Survey (e.g. Robertson, 1999) or by Travis (1955),
123 corresponds to a lithologic class meant to identify a particular kind of rock described on the basis
124 of hand-sample observations. The dimension of the representative volume used for lithologic
125 classification may vary from hand-sample to kilometer scale.

126 An ideal lithologic classification system would assign every material in its scope (domain) to a
127 unique class. However, there are many examples of materials that can be recognized as
128 belonging to more than one class, depending on the criteria used for classification. Examples
129 include low-grade metasedimentary rocks that may be described as metamorphic rocks and as
130 sedimentary rocks, saprolites that may be described as an unconsolidated material and as their
131 bedrock parent, and calc-lithic sandstone that may be classified as both a sandstone and a
132 limestone. Any classification system that attempts to define disjoint (non-overlapping) classes
133 over the entire domain of rocks and unconsolidated materials either must define ad hoc rules that
134 allow ambiguous instances to be assigned to unique classes, or must add numerous new classes
135 that include composite kinds of materials. A better solution is to allow separate classification
136 schemes. Earth scientists who have different geologic interests may use different classification
137 schemes. Individual materials may be classified differently using different schemes. Different
138 lithologic classification schemes have different classification criteria, and may have different
139 domains of classification. The domains of classification for different schemes may overlap, but
140 classes in any particular scheme are disjoint. Lithologic classification over the whole domain of
141 rock and unconsolidated materials is thus overlapping (i.e. not disjoint).

142 To produce a lithologic classification system that allows different observers to classify a given
143 material in the same way, the system must be based on physical properties recognizable by all
144 observers. Strict adherence to this rule would not allow use of genetic interpretations in the
145 classification of a material unless they could be couched in purely descriptive terms (see
146 discussion in Travis, 1955, p. 1). The properties used for field lithologic classification include:

- 147 • modal mineralogy
- 148 • grain size
- 149 • grain shape
- 150 • rock fabric (the arrangement of grains in an aggregate to form the rock)
- 151 • structures in the rock (bedding, layering, etc.)

152 Distinct bodies of rock may be defined based on other physical properties, such as magnetic
153 susceptibility or density, but these are not generally used as field criteria.

154 The approach to a lithologic classification proposed here is fundamentally descriptive
155 [unavoidable genetic classes such as impact-metamorphic rock are subdivided based on
156 descriptive attributes]. Classification of a particular material is based on observable features, and
157 assignment of a material to a lithologic class implies that certain descriptive criteria are met.
158 These criteria must be archived in the database in order to document the classification system.
159 The descriptions that define the lithologic classes also serve to provide default values for
160 properties when a material is assigned to the class but not described in greater detail. The

161 definition of a lithologic class must be part of a classification scheme (or terminology system, or
162 controlled vocabulary) that defines the domain of classification and classification criteria. The
163 definition must state the dimension of the representative volume for the class, the criteria that are
164 sufficient to assign membership in the class, and to the extent possible, a default description of
165 other aspects of materials that are assigned to the class.

166 In order to gain acceptance in the geoscience community, any lithologic classification system
167 needs to be consistent with common usage. This may require some relaxation of the strict
168 adherence to observable physical properties as criteria for classification, because traditional rock
169 classification has always involved some genetic interpretation (e.g., igneous, sedimentary,
170 metamorphic are fundamentally genetic). The operational rule for consistency is that existing
171 terms may be redefined to narrow their meaning, but may not be redefined to include rocks that
172 are not included as part of that class in common usage.

173 2.5 Guides and precedents

174 Pioneering efforts by the British Geological Survey to develop a comprehensive, systematic
175 classification of metamorphic rocks (Robertson, 1999) have proven to be extremely useful as a
176 precedent. In addition, the International Union of Geological Sciences (IUGS) Subcommission
177 on the Systematics of Metamorphic Rocks (SCMR) is following a precedent for igneous rocks
178 (LeMaitre and others, 1989, 2002) in developing standards for the classification and
179 nomenclature of metamorphic rocks (Schmid and others, 2002). Recent SCMR proposals
180 (Appendix 1) include a systematic nomenclature that combines rock composition (mineral
181 components) and structural root terms (gneiss, schist, granofels), as well as standards for non-
182 systematic metamorphic-rock names in common use. Where possible, we have attempted to
183 follow or adapt the SCMR proposals available at this time to meet the SLTT need for a rigorous
184 hierarchical classification compatible with digital databases.

185 Useful glossaries of terms related to metamorphic rocks are available in Miyashiro (1994,
186 Appendix II), Passchier and Trouw (1996), Jackson (1997), Barker (1998), and the IUGS
187 proposals listed in Appendix 1. The IUGS SCMR is currently preparing an authoritative glossary
188 of about 1500 terms for metamorphic rocks. That glossary will be similar to Le Maitre's (1989,
189 2002) glossary for igneous rocks. When completed, it will be released online at
190 <http://www.bgs.ac.uk/SCMR>. The final IUGS glossary will eventually provide authoritative
191 definitions of common metamorphic terms for international use, recommendations for use of
192 uncommon terms, and terms that should be abandoned and replaced by common rock names.

193 2.6 Philosophy of classification (SLTTM_1.0)

194 The classification system proposed here is designed for database applications. This makes it
195 distinct from a system for naming rocks, which provides guidelines for giving a descriptive name
196 to a particular rock or kind of rock for use in discourse. Two naming schemes have recently been
197 proposed (Robertson, 1999; Schmid and others, 2002), and the reader is referred to those
198 documents for more information.

199 The lithologic classification needed for a database is designed to answer the question 'what kind
200 of rock is this?' The flexibility of the database allows a more complete answer to this question
201 that can be encapsulated in a single rock name, and several equally valid answers that depend on
202 the context of the 'kind' question can be provided. Ideally, the database system would also
203 provide a 'free-form rock name' field that allows the geoscientist to give a rock being described
204 whatever name they feel best characterizes the rock.

205 The philosophy of this classification is based (to the degree possible) on features that reflect the
206 transformation of a rock from one state (the protolith) to another (the composite-genesis rock).
207 Features from the original state are not the basis for classification as a metamorphic rock. If such

208 features are still apparent, they are the basis for classification of the lithology of the metamorphic
209 rock protolith, which will also be part of the database description of the rock.

210 Classification Rules:

- 211 • The domain of classification is hand sample (1 cm to 30 cm) scale metamorphic rocks.
- 212 • The scheme should be as descriptive as possible. Genetic considerations are used only to
213 the extent that they can be defined based on features observable in the rock at hand-
214 sample scale.
- 215 • Classification must be complete—if a rock is determined to be a composite-genesis rock,
216 it must have a place in the classification scheme.
- 217 • A rock may be classified in more than one way based on different criteria (fabric,
218 composition), but given a particular set of criteria, classification must be unique.
- 219 • The scheme is for *classifying rocks as metamorphic rocks*, thus interpretation of the
220 protolith of a rock is not used to determine the classification of the rock. [However,
221 where protoliths are known, the combined use of a metamorphic rock classification (as
222 proposed here) and an independent protolith (sedimentary or igneous) rock classification
223 should be permissible.]
- 224 • Modal mineralogy (mineral volume percentage) is described in database fields other than
225 the field containing material names.
- 226 • Geologists can give a rock any name they feel is most appropriate using an informal
227 ‘rock name’ field in the database that is separate from this lithologic classification.
- 228 • Classification must be hierarchical. Superclasses must be included that group the
229 lithologic classes based on fabric or composition criteria. These should include a small
230 number of sub-classes. Each sub-class may be further broken down into smaller
231 subordinate classes, etc.
- 232 • Classification must be useful. The scheme should be sufficiently simple and flexible to
233 facilitate use by workers of varying experience and expertise (Robertson, 1999), both in
234 producing and using digital geologic maps and map databases.

235 The basic elements of the classification are rock fabric and composition (modal mineralogy). If
236 both are known, both the fabric and composition parts of the classification decision tree must be
237 navigated to determine the classification. If only one is known, only one of the two parts will
238 need to be navigated. Fabric may be unknown, in which case the fabric term will just have to be
239 ‘rock’. The sorts of fabric that can be defined for aphanitic or very fine-grained rocks include
240 massive, laminated, and glassy. A wider variety of fabrics is defined for phaneritic rocks,
241 including schistose, gneissic, and granoblastic. Composition may be unknown, either because it
242 is unspecified, in which case the fabric classification tree is used, or because the grain size is too
243 small to determine mineralogy beyond silicic, calcsilicate, argillic, or calcareous.

244 In the class hierarchy, classes that have a composition and a fabric criteria are considered
245 ‘children’ of both a ‘composition and fabric’ lithology parent, and of a generic composition and a
246 generic fabric parent. The class hierarchy is thus a directed acyclic graph, not a tree.
247 Ramifications of this knowledge representation in database implementations should be carefully
248 considered.

249 2.7 Who developed this document, and how?

250 This document was developed by geoscientists from a variety of American and Canadian
251 geoscience agencies (Table 2.7.1). The group was assembled in early 2000 as the Metamorphic
252 Science Language Technical Team (SLTTM) of the North American Geologic-map Data Model

253 Steering Committee, although the panel soon came to be called the Composite-genesis Technical
 254 Team in recognition of the fact that earth materials they were classifying included more than just
 255 metamorphic rocks as traditionally envisioned. Panel members were appointed in the following
 256 ways:

- 257 (1) Most participants from the U.S. Geological Survey (USGS) were identified by Regional
 258 Geologic Executives from the USGS Western, Central, and Eastern Regions. Some USGS
 259 scientists were appointed by Coordinators of USGS line-item science programs;
 260 (2) Scientists from the Geological Survey of Canada (GSC) were identified by Canadian
 261 members of the North American Geologic-map Data Model Steering Committee;
 262 (3) Scientists from Canadian Provincial geological surveys were identified by Canadian
 263 members of the North American Geologic-map Data Model Steering Committee;
 264 (4) Scientists from academic institutions were selected by the committee co-chairs.

265 The document was written by co-chairs Horton and Richard in consultation with members of the
 266 panel.

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267

Table 2.7.1

268 3 COMPOSITE-GENESIS ROCKS

269 The domain of composite-genesis rocks includes rocks that record the effects of more than one
270 sequential rock-forming process. The primary rock-forming processes are considered to be those
271 related to the formation of sedimentary and igneous rocks. Sedimentary rocks are formed by the
272 accumulation of sediment subsequent lithification by diagenetic processes to form rock. Igneous
273 rocks are formed by the cooling of melted rock (magma) within or derived from the Earth's (or
274 another planetary body's) interior. Composite-genesis rocks are those that record the effects of
275 one or more geologic events subsequent to the primary rock-forming process. These may include
276 burial, heating, ductile or brittle deformation, and open-system chemical changes. A purely
277 descriptive definition of composite-genesis rock is problematic because the definition requires
278 interpreting genetic process, which depends to some extent on the knowledge and skill of the
279 observer.

280 The major classes of composite-genesis rock, based on genetic as well as descriptive criteria, are
281 metamorphic (including hydrothermally altered) rock, cataclastic rock, composite-genesis melt
282 rock, and impact-metamorphic rock. These classes are subdivided and organized based on
283 descriptive criteria into separate hierarchies for rock fabric and composition. The application of
284 this classification to particular rock units requires the dual use of both hierarchies as outlined later
285 in this report.

286 3.1 Metamorphic rock definition and limits (including hydrothermally altered rock)

287 3.1.1 Definition

288 *Metamorphic rock*—any rock derived from pre-existing rock by mineralogical, chemical, or
289 structural changes, essentially in the solid state (based on the *Glossary of Geology*, Jackson,
290 1997).

291 In order to be considered a metamorphic rock, there must be observable features in the rock that
292 document change subsequent to the original formation of the rock, under physical or chemical
293 conditions that differ from those normally occurring at the surface of the Earth (or other planetary
294 bodies) or in zones of cementation and diagenesis below the surface (Smulikowski and others,
295 1997). Robertson (1999) stipulates that the process of metamorphism encompasses all of the
296 solid-state changes that occur between the upper and lower limits of metamorphism.
297 'Metamorphic rock' thus includes hydrothermally altered rock and mylonitic rock, but not melt
298 rock. Although 'cataclastic rock' and 'impact metamorphic rock' could be considered
299 'metamorphic' in the most general sense, they are here distinguished from conventional
300 'metamorphic rock' because their formative processes can be outside the lower and upper limits
301 of metamorphism discussed below. Metamorphism does not include rock weathering or soil
302 formation.

303 3.1.2 Lower limit of metamorphism

304 Because identification of a rock as a metamorphic rock is predicated on observation of features
305 not resulting from the original rock-forming process, identification as a metamorphic rock
306 depends on the skill and interests of the observer, and the kinds of features observed. Distinction
307 of the 'original rock-forming process' as distinct from subsequent 'metamorphic processes' is not
308 always clear. Sediment undergoes a continuous progression of changes from deposition through
309 lithification to form sedimentary rock; this process is known as diagenesis. Changes in a deeply
310 buried sedimentary rock may be continuous from diagenesis into recrystallization to form a
311 metamorphic rock.

312 Robertson (1999) defines the boundary between diagenesis and metamorphism in sedimentary
313 rocks as follows:

314 “...the boundary between diagenesis and metamorphism is somewhat arbitrary and strongly
315 dependent on the lithologies involved. For example changes take place in organic materials
316 at lower temperatures than in rocks dominated by silicate minerals. In mudrocks, a white
317 mica (illite) crystallinity value of $< 0.42D_{2U}$ obtained by X-ray diffraction analysis, is
318 used to define the onset of metamorphism (Kisch, 1991). In this scheme, the first
319 appearance of glaucophane, lawsonite, paragonite, prehnite, pumpellyite or stilpnomelane is
320 taken to indicate the lower limit of metamorphism (Frey and Kisch, 1987; Bucher and Frey,
321 1994; Frey and Robinson, 1998). Most workers agree that such mineral growth starts at
322 $150 \pm 50^\circ \text{C}$ in silicate rocks. Many lithologies may show no change in mineralogy under
323 these conditions and hence the recognition of the onset of metamorphism will vary with
324 bulk composition.”

325 3.1.3 Upper limit of metamorphism

326 At the highest grades of metamorphism, rocks begin to melt. The temperatures and pressures of
327 the onset of melting range from approximately 650°C to more than 1100°C depending on bulk
328 composition and the proportion of water in the fluid phase. Migmatitic rocks grade into igneous
329 rocks as the proportion of granitic melt increases. The upper limit of metamorphism is defined
330 somewhat arbitrarily. Here, a rock mass that consists $\geq 50\%$ (by volume) of rock having igneous
331 composition and texture (crystallized from a melt) is classified as an igneous rock.

332 3.1.4 Hydrothermally altered rock

333 Hydrothermally altered rock has fabric and composition indicating solid-state mineralogical and
334 chemical changes in response to hot, mineral-rich waters. Igneous rocks may continue to undergo
335 mineralogical and chemical changes during cooling even after most or all of the magma has
336 crystallized to produce a ‘solid’ rock. Hydrothermal activity due to residual heat and mineral-rich
337 water from a pluton may change igneous rocks significantly following their crystallization.
338 ***Hydrothermally altered (including metasomatic) rock is treated as ‘metamorphic rock’ in this***
339 ***classification.*** We propose the following criteria to distinguish hydrothermally altered or
340 metasomatic rock from igneous rock. The rock is classified as metamorphic if (1) the texture has
341 been modified such that it can no longer be considered igneous, (2) the bulk composition of the
342 rock is inconsistent with compositions that can be derived purely from a magma and associated
343 processes such as assimilation and differentiation, or (3) minerals inconsistent with magmatic
344 crystallization are present.

345 3.2 Cataclastic rock, composite-genesis melt rock, impact-metamorphic rock

346 The classes of composite-genesis rock, in addition to metamorphic rock, include cataclastic rock,
347 impact-metamorphic rock, and composite-genesis melt rock.

348 *Cataclastic rock*—A rock having more than 10% of its volume consisting of fragments bounded
349 by fractures related to brittle deformation that occurred after the formation of the protolith.
350 Cataclastic rocks are commonly associated with faults, although that genesis is not required by
351 the descriptive definition, and they can form below the limits of metamorphism discussed above.

352 *Composite-genesis melt rock*—A rock that solidified from melt (liquid) of a preexisting rock
353 under conditions (such as contact metamorphism, friction along faults, or meteorite impact)
354 outside the domain of typical igneous rocks.

355 *Impact-metamorphic rock*—A rock formed by the impact of a planetary body (projectile) on a
356 planetary surface (target). While this criterion is unavoidably genetic by definition, it is based on
357 the interpretation of observable features such as shock-induced planar deformation features, high-
358 pressure minerals, shatter cones, and crater structure, and subdivisions are based on descriptive
359 features as discussed in a later section.

360 4 CLASSIFICATION COMPONENTS

361 It is important to remember that the proposed classification of metamorphic rocks is for use in a
362 computer database. This classification system is presented in five diagrams, which include a
363 class hierarchy diagram in two parts (fabric and composition) and four decision trees
364 (metamorphic rock fabric, metamorphic rock composition, mylonitic rock, cataclastic rock). The
365 classification system includes a glossary of definitions, parent-child relationships, and references
366 for terms in the form of a Microsoft Access database. This glossary should be expanded in the
367 future to encompass terms presently missing from the classification, including those to be
368 discouraged on new maps, and recommended usage and suggestions for integrating them into the
369 classification. Future expansion of the glossary for this classification should strive for agreement
370 with the glossary of metamorphic terms under development by the IUGS SCMR
371 [<http://www.bgs.ac.uk/SCMR>] except where precluded by database requirements or common
372 North American usage.

373 The system proposed here classifies a given composite-genesis rock along two orthogonal
374 dimensions, fabric (including texture) and composition. Class names thus have a fabric
375 component (e.g., granofels, schist, gneiss) and a compositional component (e.g., marble,
376 calcsilicate, quartzite, quartzofeldspathic, pelitic, amphibolite, serpentinite). The alternative
377 classification rationale of Robertson (1999, Fig. 3) requires that a rock name is first assigned
378 based on mineralogy, and in the absence of mineralogical information, the classification uses a
379 textural term. In our view, this results in loss of information.

380 5 CLASSIFICATION BY ROCK FABRIC

381 Rock fabric (including texture), as defined in Appendix 2, is the complete spatial and geometrical
382 configuration of all those components that are contained in a rock and that are penetratively and
383 repeatedly developed throughout the volume of rock under consideration. It includes the shapes
384 and characters of individual parts of a rock mass and the manner in which these parts are
385 distributed and oriented in space.

386 5.1 Terminology—phaneritic and aphanitic

387 The IUGS Subcommittee on the Systematics of Metamorphic Rocks (SCMR) extended the use
388 of igneous grain-size terms ‘phaneritic’ (large enough to be distinguished by the unaided eye) and
389 ‘aphanitic’ (too fine grained to be distinguished by the unaided eye) to metamorphic rocks, noting
390 that these qualitative definitions correspond approximately to “aphanitic” (ca. <0.1 mm) and
391 “phaneritic” (ca. >0.1 mm) (Schmid and others, 2002, Table 5). The British Geological Survey
392 rock classification (Gillespie and Styles, 1999, section 3.2, p. 6) applies “phaneritic” and
393 “aphanitic” to metamorphic rocks but interprets the actual grain sizes differently:

394 “Placing the boundary between ‘medium grained’ and ‘fine grained’ crystals in crystalline
395 igneous and metamorphic rocks at 0.25 mm essentially divides aphanitic rocks (in which
396 individual crystals are too fine grained to be distinguished by the naked eye) from phaneritic
397 rocks (in which individual crystals can be distinguished by the naked eye)...This is also the
398 boundary between medium and fine sand for sedimentary clasts.”

399 Others interpret the phaneritic-aphanitic (visible-invisible) distinction as approximately
400 equivalent to the sand-silt grain size distinction for sediments, which is variously considered to be
401 in the range of 0.032 mm (Robertson, 1999, BGS grain size scheme) to 0.062 mm (Jackson,
402 1997) to 0.074 mm (Engineering grain-size scale, ASTM standard D422-63; D643-78). This is
403 smaller than the phaneritic-aphanitic boundary of 0.1 mm proposed by Schmid and others (2002).
404 However, in practical terms, such detailed grain size distinctions are impossible in the field, and
405 very difficult under any conditions. The IUGS usage is endorsed and applied in this
406 classification.

407 5.2 Fabric prototypes

408 Several basic types of rock fabric are recognized in virtually all approaches to composite-origin
409 rocks. For classification purposes, the problem is defining the boundaries between the types
410 based on descriptive criteria. There are terms in common usage (e.g., schist, gneiss, granofels,
411 mylonite, cataclasite) that provide basic prototypes, but there is a great deal of variation in the
412 definition of boundaries between the classes.

413 The basic level of classification is the definition of seven fabric prototypes. These are presented
414 at this stage without assigning names (although the names may be obvious) to emphasize that
415 they represent fundamental *kinds* of rock.

416 The following list of basic fabric classes is meant to partition the domain of composite-origin
417 rocks based on descriptive criteria in a manner that different observers can agree upon. All
418 quantitative boundaries (percentages, ratios, and dimensions) should be considered open to
419 discussion—the qualitative distinctions are more important here. In general, ‘phaneritic’ (having
420 visible grains) is implied except where ‘aphanitic’ is specified.

421 (1) Rocks that consist of angular fragments bounded by fractures. [A threshold for assignment of
422 rocks to this category is proposed where $\geq 10\%$ of the rock consists of fragments bounded by
423 fractures.]

- 424 (2) Aphanitic metamorphic rocks that are too fine-grained to determine mineralogy. [This
425 criterion is meant to separate rocks that can be classified based on modal mineralogy from
426 rocks too fine grained to distinguish mineralogy.]
- 427 (3) Phaneritic metamorphic rocks that have granoblastic fabric and very little or no foliation.
428 [Very little foliation means some foliation may be present, but does not meet the criteria for
429 foliated. Foliated means that $\geq 10\%$ of the mineral grains in the rock are fabric elements in
430 the foliation. To be a foliation-defining fabric element, a mineral grain must have an
431 inequant crystal habit, or an inequant shape due to deformation and an aspect ratios $\geq 1.5:1$
432 between the long and short axis of the deformed grain.]
- 433 (4) Well-foliated rocks characterized by tectonic reduction in grain size and having a foliation
434 defined by the shapes of oriented mineral grains or grain aggregates. [Criteria proposed here
435 for implementation are a foliation defined by the shape of deformed mineral grains or grain
436 aggregates having an aspect ratio $> 1.5:1$, and $>10\%$ 'matrix' showing evidence of tectonic
437 reduction in grain size without loss of material continuity. The matrix consists of new or
438 recrystallized mineral grains that are interpreted to be smaller than the mineral grains in the
439 original, undeformed protolith. The definition is meant to identify a fabric in the rock due to
440 crystal plastic and/or other types of non-cataclastic deformation.]
- 441 (5) Phaneritic rocks that have a well-developed schistosity. [The sticking point is the definition
442 of "well developed" schistosity. We propose that $>50\%$ of rock consists of mineral grains
443 having a tabular, lamellar, or prismatic crystallographic habit that are oriented in a continuous
444 planar or linear fabric (following Jackson, 1997). Continuous is defined on a hand sample-
445 scale, to mean that domains lacking the fabric are <1 cm thick if they are layers, and <5 cm in
446 diameter if they are irregular patches, and constitute $< 25\%$ of the rock. The IUGS SCMR
447 (Brodie and others, 2002) suggests using criteria that rock splits on scale <1 cm, but this
448 criteria depends on the tool used to do the splitting, the skill of the operator doing the
449 splitting, and the degree of weathering or alteration of the rock being split, and is thus not
450 objective and universally applicable. The IUGS criteria could not be used to classify a rock
451 in thin section. The IUGS (Brodie and others, 2002) also proposes to define schistosity as
452 due to inequant mineral grains or grain aggregates, without specifying that their shape is due
453 to crystallographic habit.]
- 454 (6) Foliated, layered phaneritic rocks that lack well developed schistosity and have laterally
455 continuous compositional layering. [We here propose a definition of "laterally continuous
456 compositional layering" as having layers > 5 mm thick that can be traced laterally > 10 cm
457 (length of lateral continuity).]
- 458 (7) Foliated non-layered phaneritic rocks that do not have well developed, continuous schistosity
459 or continuous compositional layering.

Table 1. Fabric classification terms

Name	Definition	Source
breccia	Generic classification for rock in which penetrative, through going fractures separate visible fragments (>0.1 mm diameter) that form > 30% of rock, and fragments are rotated relative to each other.	Sibson (1977), Snoke and Tullis (1998)
broken rock	Rock consisting of <30% fragments, or rock that consists of >30% fragments, but fragments are not rotated with respect to each other. Because protolith is assumed to be recognizable, do not apply a composition modifier. Composition is indicated by protolith link.	Robertson (1999)
cataclasite	cataclastic rock that maintained primary cohesion during deformation, and consists of 50-90% matrix. Matrix is broken mineral grains or rock fragments that are too small to be discernible (0.1 mm following definition of aphanitic). Equivalent to “mesocataclasite” of Brodie and others (2002)	Sibson (1977), Scholz (1990), Snoke and Tullis (1998), Barker (1998, Appendix II)
cataclasite-series rock	Collective term for cataclastic rocks that maintained primary cohesion during deformation; includes protocataclasite, cataclasite, and ultracataclasite (Sibson, 1977). These subtypes are differentiated based on the fraction of the rock considered ‘matrix’. Matrix is broken mineral grains or rock fragments that are too small to be discernible (0.1 mm following definition of aphanitic).	Sibson (1977), Scholz (1990), Snoke and Tullis (1998)
cataclastic rock	“A rock, such as tectonic breccia, containing angular fragments that have been produced by the crushing and fracturing of preexisting rocks as a result of mechanical forces in the crust.” (Jackson, 1997). Criterion added here for database precision is >10% of rock volume composed of fragments bounded by fractures related to a secondary tectonic event. Use of the term ‘cataclastic’ denotes deformation in the absence of crystal plastic processes. [Mylonite-series rock is not classified here as a cataclastic rock, although it may contain subordinate cataclastic fabric.]	Jackson (1997)
Composite-origin rock	Rock with unspecified fabric, for which insufficient information is available to determine if it is cataclastic, impact-related, glassy, or metamorphic.	this report
contact metamorphic melt rock	Rock in which framework of rock is glassy mineral material that encloses rock and mineral fragments, and there is evidence that melting is related to contact metamorphism.	this report
fault breccia	Cataclastic rock lacking evidence for primary cohesion during deformation, in which fractures separate visible fragments (>= 0.1mm in diameter) that form >30% of rock mass, and are rotated relative to each other, and no evidence of impact metamorphism is observed. Composition modifiers refer to character of comminuted rock between fragments; composition of fragment protolith is assumed identifiable, and composition of that is indicated via protolith relationship.	Snoke and Tullis (1998), Barker (1998, Appendix II)
foliated cataclasite	Cataclastic rock that maintained primary cohesion during deformation, and consists of 50-90% matrix, and has a foliation. Matrix is broken mineral grains or rock fragments.	Snoke and Tullis (1998); Sibson (1977) modified by Scholz (1990)

Name	Definition	Source
foliated gouge	Cataclastic rock lacking evidence for primary cohesion during deformation, in which visible fragments (> 2 0.1 mm in diameter) constitute <30% of the rock mass, and having a foliation. Composition modifier refers to general character of comminuted matrix. Protolith of fragments assumed to be recognizable.	Snoke and Tullis (1998); Sibson (1977) modified by Scholz (1990)
foliated metamorphic rock	Any metamorphic rock that contains a foliation.	this report
foliated protocataclasite	Cataclastic rock that maintained primary cohesion during deformation, and consists of 10-50% matrix, and has a foliation. Matrix is broken mineral grains or rock fragments.	Snoke and Tullis (1998); Sibson (1977) modified by Scholz (1990)
foliated ultracataclasite	Cataclastic rock that maintained primary cohesion during deformation, and consists of 90-100% matrix, and has a foliation. Matrix is broken mineral grains or rock fragments.	Snoke and Tullis (1998); Sibson (1977) modified by Scholz (1990)
gneiss	General term for a foliated, phaneritic rock without well developed, continuous schistosity	Schmid and others (2002), Jackson (1997), Barker (1998, Appendix II)
gneissic mylonite	Mylonite that has continuous compositional layering, >5 mm thick. Continuous means that layers defining the foliation can be traced for >10 cm (length of lateral continuity), and are spaced at a distance \leq the average length of lateral continuity. Monomineralic compositional modifiers do not apply, because if monomineralic, cannot develop gneissoid banding.	revised from IUGS (Brodie and others, 2002) and Robertson (1999)
gneissic protomylonite	Protomylonite that lacks well developed continuous schistosity, but displays continuous compositional banding >5 mm thick.	this report
gouge	Cataclastic rock lacking evidence for primary cohesion during deformation, in which visible fragments (> 0.1 mm in diameter) constitute <30% of the rock mass. Composition modifier refers to general character of comminuted matrix. Protolith of fragments assumed to be recognizable.	Sibson (1977), Snoke and Tullis (1998), Barker (1998, Appendix II)
granoblastic rock	Phaneritic metamorphic rock having granoblastic fabric and very little or no foliation or lineation. Specifically, <10% of the particles in the rock are planar or linear fabric elements with an aspect ratio \square 1.5:1. [Use of generic term granoblastic rock denotes that nothing is known about the rock fabric beyond that it is granoblastic.]	
granofels	Phaneritic metamorphic rock that has little or no foliation or lineation. Specifically, <10% of the particles in the rock are planar or linear fabric elements with an aspect ratio \geq 1.5:1. In this sense, granofels is the phaneritic equivalent of hornfels.	adapted from Goldsmith (1959), Robertson (1999), Barker (1998, Appendix II), and IUGS (Schmid and others, 2002; Brodie and others, 2002)

Name	Definition	Source
hornfels	Aphanitic metamorphic rock that has little or no foliation or lineation. Although hornfels is typically a product of contact metamorphism, that genesis is not a requirement of this descriptive definition. Hornfels, in this sense, is the aphanitic equivalent of granofels.	Winkler (1967) in Jackson (1997), Winkler (1979)
impact breccia	Breccia (as generally defined) containing fragments that show unequivocal evidence of shock metamorphism; typically occurring “around, inside, and below impact craters” (Stöffler, 2001c); subclasses are monomict impact breccia, polymict impact breccia, and suevite.	Stöffler (2001c), Stöffler and Grieve(2001)
impact melt rock	Crystalline, semi-glassy, or glassy rock in which $\geq 50\%$ of the rock volume is solidified from impact melt (implying $< 50\%$ non-melt inclusions). Framework of rock is glassy mineral material and evidence of impact is observed.	Stöffler (2001c), Stöffler and Grieve(2001).
impactite (impact metamorphic rock)	Rock that shows evidence of impact metamorphism. General term for rocks affected by impact resulting from the collision of planetary bodies.	Stöffler (2001c), Stöffler and Grieve (2001).
incohesive cataclastic rock	Cataclastic rock for which evidence for primary cohesion during deformation is lacking or not specified, and evidence for impact metamorphism is lacking or not specified.	Snoke and Tullis (1998); Sibson (1977) modified by Scholz (1990)
layered gneiss	Foliated, phaneritic rock that lacks well developed, continuous schistosity and has laterally continuous compositional layering > 5 mm thick. Laterally continuous means that layers defining the foliation can be traced for > 10 cm (length of lateral continuity), and are spaced at a distance \leq the average length of lateral continuity.	modified from IUGS (Schmid and others, 2002; Brodie and others, 2002), and Robertson (1999)
metamorphic glass	Rock with framework of glassy material that is the product of metamorphic event.	this report
metamorphic rock	A metamorphic rock is any rock derived from pre-existing rock by mineralogical, chemical, or structural changes, essentially in the solid state (based on Jackson, 1997). In order to be considered a metamorphic rock, there must be observable features in the rock that document change after the original formation of the rock as a rock, under physical or chemical conditions that differ from conditions normally occurring at the surface of the Earth and in zones of cementation and diagenesis below the surface (Smulikowski and others, 1997).	Jackson (1997), Smulikowski and others (1997)
migmatite [recommended for text field only]	A heterogeneous composite rock mass consisting of irregular and discontinuous interleaving of leucocratic granitoid material (leucosome) and residual high-grade metamorphic material (restite) (Barker, 1998, Appendix II: Glossary). <i>Migmatite is not used as a root name in this classification for reasons discussed in the text.</i> Here, the adjective “migmatitic” is applied to a fabric term, as in migmatitic gneiss, where the granitoid part makes up $> 10\%$ of the rock volume.	paraphrase of Barker (1998, Appendix II), (Wimmenauer and Bryhni, 2002)

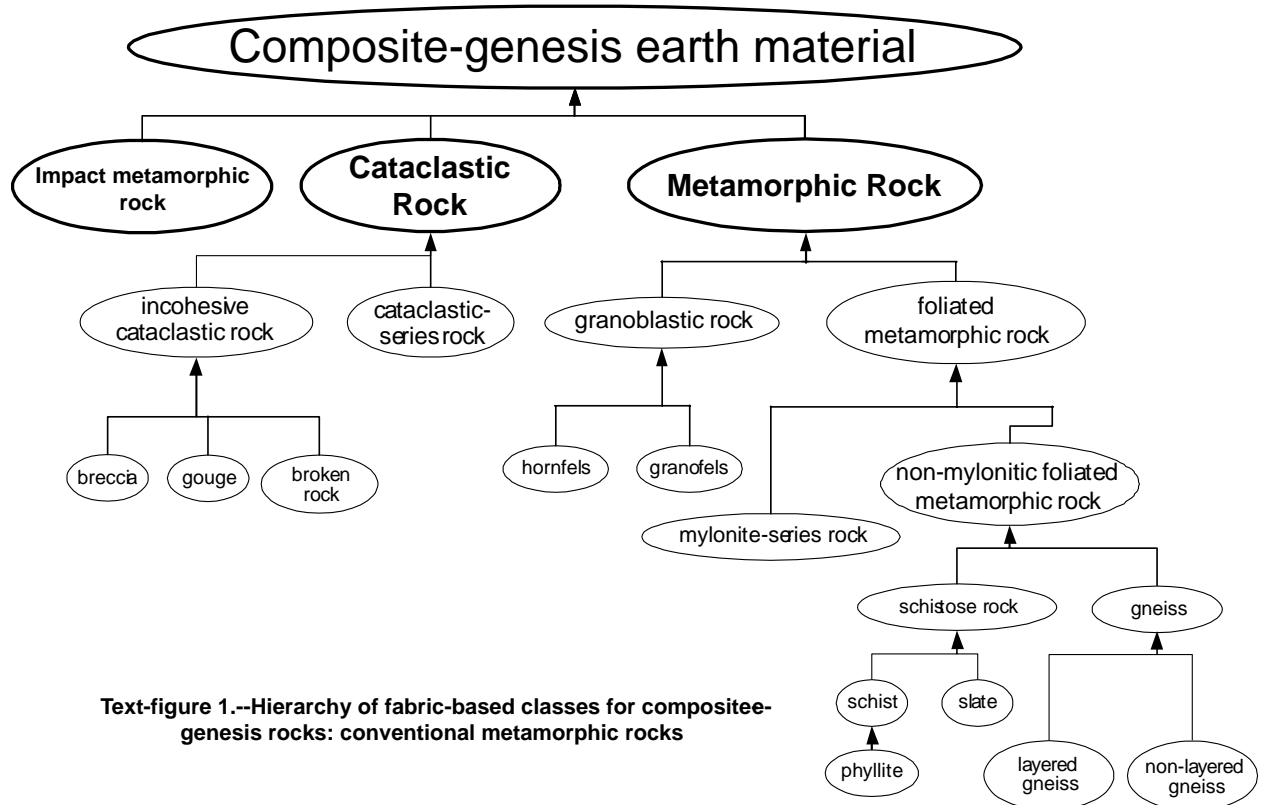
Name	Definition	Source
migmatitic gneiss	Foliated, phaneritic rocks without well developed, continuous schistosity that is megascopically composite, consisting of two or more petrographically different parts, one of which is the country rock in a more or less metamorphic state, the other is of pegmatitic, aplitic, or granitic appearance. The granitoid phase makes up >10% of the rock volume.	composite from Jackson (1997) and Robertson (1999)
migmatitic granofels	Non-foliated, phaneritic rock that is megascopically composite, consisting of two or more petrographically different parts, one of which is the country rock in a more or less metamorphic state, the other is of pegmatitic, aplitic, or granitic appearance. The granitoid phase makes up >10% of the rock volume. Example -- 'ophthalmic' or patch migmatite of Mehnert (1968).	composite from Jackson (1997) and Robertson (1999)
migmatitic layered gneiss	Foliated, phaneritic rock without well developed, continuous schistosity that has continuous compositional layering, > 5 mm thick, and is megascopically composite, consisting of two or more petrographically different parts, one of which is the country rock in a more or less metamorphic state, the other is of pegmatitic, aplitic, or granitic appearance. The granitoid phase makes up >10% of the rock volume.	composite from Jackson (1997) and Robertson (1999)
migmatitic non-layered gneiss	Foliated, phaneritic rock without well developed, continuous schistosity or continuous compositional layering that is megascopically composite, consisting of two or more petrographically different parts, one of which is the country rock in a more or less metamorphic state, the other is of pegmatitic, aplitic, or granitic appearance. The granitoid phase makes up >10% of the rock volume.	composite from Jackson (1997) and Robertson (1999)
migmatitic schist	Foliated, phaneritic rock that has well developed, continuous schistosity that is megascopically composite, consisting of two or more petrographically different parts, one of which is the country rock in a more or less metamorphic state, the other is of pegmatitic, aplitic, or granitic appearance. The granitoid phase makes up >10% of the rock volume.	composite from Jackson (1997) and Robertson (1999)
monomict impact breccia	Impact breccia free of impact melt in which all of the fragments (100%) have essentially the same composition.)	Stöffler (2001c), Stöffler and Grieve(2001).
mylonite	Mylonite-series rock consisting of 50-90% matrix showing evidence of tectonic grain size reduction. Equivalent to "orthomylonite" of Wise and others (1984) and "mesomylonite" of Brodie and others (2002)	Sibson (1977), Hanmer (1987), Passchier and Trouw (1996), Snoke and Tullis (1998)

Name	Definition	Source
mylonite-series rock	A collective term for the “mylonite series” rocks of Sibson (1977), including protomylonite, mylonite, ultramylonite, and phyllonite. Equivalent to Brodie and others’ (2002) broader usage of “mylonite” for “A fault rock which is cohesive and characterized by a well developed [foliation] resulting from tectonic reduction of grain size, and commonly containing rounded porphyroclasts and lithic fragments of similar composition to minerals in the matrix.” Additional criteria introduced here for precision in databases include a foliation defined by deformed mineral grains or grain aggregates having aspect ratio > 1.5:1, and >10% of rock consisting of 'matrix'. Matrix is an aggregate of new mineral grains (not present in the protolith, but may be same mineral species) that are significantly smaller (1% of diameter) than the original size of mineral grains, and are the product of non-cataclastic deformation. Non-cataclastic deformation is deformation in which the material continuity of the deformed volume is maintained on the scale of observation, indicated by the absence of thoroughgoing fractures in the volume. The definition is meant to identify a fabric in the rock due to crystal plastic and/or other types of non-cataclastic deformation.	Sibson (1977), Snoke and Tullis (1998), Brodie and others (2002)
non-layered gneiss	A foliated, phaneritic rock that lacks well developed, continuous schistosity and laterally continuous compositional layering > 5 mm thick. “Laterally continuous” here means that layers defining the foliation can be traced > 10 cm (length of lateral continuity). Foliated means that ≥ 10% of mineral grains in the rock are fabric elements.	IUGS (Schmid and others (2002; Brodie and others, 2002)
non-mylonitic foliated metamorphic rock	Metamorphic rock that is not granoblastic and does not have a mylonitic fabric.	this report
phyllite	Rock that has a well developed, continuous schistosity, an average grain size (excluding porphyroblasts) <0.25 mm and >0.1 mm, and a silvery sheen on cleavage surfaces.	Jackson (1997), Barker (1998, Appendix II)
phyllonite	Phyllosilicate-rich mylonite-series rock of phyllitic appearance (having a silvery sheen on cleavage surfaces); “sometimes (erroneously) used for ultramylonite” (Passchier and Trouw, 1996). Criterion added here for precision is having >40% phyllosilicate minerals (e.g., mica, chlorite).	Sibson (1977), Passchier and Trouw (1996), Snoke and Tullis (1998), Robertson (1999), Brodie and others (2002)
polymict impact breccia	Impact breccia that contains fragments of different composition and is free of impact melt particles.	Stöffler (2001c), Stöffler and Grieve(2001).
protocataclasite	Cataclastic rock that maintained primary cohesion during deformation, and consists of 10-50% matrix. Matrix is broken mineral grains or rock fragments.	Snoke and Tullis (1998): Sibson (1977) modified by Scholz (1990)

Name	Definition	Source
protomylonite	Mylonite-series rock consisting of 10-50% matrix showing evidence of tectonic grain size reduction.	composite from Robertson (1999), Jackson (1997), Snoke and Tullis (1998), Barker (1998, Appendix II), Passchier and Trouw (1996)
pseudotachylite	“Ultrafine-grained vitreous-looking material, usually black and flinty in appearance, occurring as thin planar veins, injection veins or as a matrix to pseudo-conglomerates or breccias, which infills dilation fractures in the host rock” (Brodie and others, 2002). [see “impact pseudotachylite”]	Spray,(1995), Passchier and Trouw (1996), Snoke and Tullis (1998), Brodie and others (2002)
schist	Phaneritic metamorphic rock having a well developed schistosity. Well developed schistosity is defined to mean that >50% of rock consists of mineral grains having a tabular, lamellar, or prismatic crystallographic habit that are oriented in a continuous planar or linear fabric (following Jackson, 1997). Continuous is defined on a hand sample-scale, and in quantitative terms to mean that domains lacking the fabric are <1 cm thick if they are layers, and <5 cm in diameter if they are irregular patches, and constitute < 25% of the rock.	Jackson (1997), Barker (1998, Appendix II)
schistose mylonite	Mylonite that has a well developed continuous schistosity	modified from Snoke and Tullis (1998), Sibson (1977)
schistose protomylonite	Protomylonite that has a well developed schistosity	modified from Snoke and Tullis (1998), Sibson (1977)
shocked rock	non-brecciated rock which shows unequivocal effects of shock metamorphism exclusive of whole-rock melting (Stöffler (2001c). Specifically, <= 30% of the rock is fracture-bounded fragments, or if fragments form >30% of rock, they are not rotated relative to each other.	Stöffler (2001c), Stöffler and Grieve (2001)
slate	Aphanitic rock that has well developed schistosity An average grain size <0.1mm (excluding porphyroblasts) is specified here for precision.	Jackson (1997), Robertson (1999), Brodie and others (2002)
suevite	Impact breccia that contains impact melt particles (>0 and <50%)	Stöffler (2001c), Stöffler and Grieve (2001)
ultracataclasite	Cataclastic rock that maintained primary cohesion during deformation, and consists of 90-100% matrix. Matrix is broken mineral grains or rock fragments.	Snoke and Tullis (1998): Sibson (1977) modified by Scholz (1990)
ultramylonite	Mylonite-series rock consisting of 90-100% matrix showing evidence of tectonic grain size reduction.	composite from Robertson (1999); Jackson (1997), Snoke and Tullis (1998), Passchier and Trouw (1996), Barker (1998, Appendix II)

461 5.3 Metamorphic (including hydrothermally altered) rock

462 SLTTM_1.0 recognizes the following classes of metamorphic rock (including hydrothermally
 463 altered rock) based on fabric: (1) granoblastic rock including granofels and hornfels, and (2) non-
 464 mylonitic foliated metamorphic rock including schistose rock (schist, phyllite, slate) and gneiss
 465 (layered gneiss and non-layered gneiss), and (3) mylonitic rock.



Text-figure 1.--Hierarchy of fabric-based classes for composite-genesis rocks: conventional metamorphic rocks

466
 467 5.3.1 Granoblastic rock

468 5.3.1.1 Granofels

469 Granofels was introduced by Goldsmith (1959) for “medium- to coarse-grained granoblastic
 470 metamorphic rock with little or no foliation or lineation” (Jackson, 1997). The British Geological
 471 Survey (Robertson, 1999) and provisional IUGS nomenclature (Brodie and others, 2002; Schmid
 472 and others, 2002) use the term regardless of grain size.

473 SLTTM_1.0 defines **granofels** as a *phaneritic metamorphic rock that has little or no foliation or*
 474 *lineation*, implying that that <10% of the particles in the rock are fabric elements. To be a fabric
 475 element, a particle must have an inequant shape with an aspect ratio $\geq 1.5:1$, and be aligned with
 476 other particles as part of a fabric. In this sense, granofels is the phaneritic equivalent of hornfels.

477 5.3.1.2 Hornfels

478 The *Glossary of Geology* (Jackson, 1997) cites an earlier edition (1967) of Winkler (1979) in
 479 defining hornfels as “A fine-grained rock composed of a mosaic of equidimensional grains
 480 without preferred orientation and typically formed by contact metamorphism. Porphyroblasts or
 481 relict phenocrysts may be present in the characteristically granoblastic (or decussate) matrix.” A
 482 genetic connotation of contact metamorphism is required by the British Geological Survey
 483 definition (Robertson, 1999, p. 12) of hornfels as “a specific variant of granofels” to be used as a
 484 root name where “features of the original rock have been modified by contact

485 metamorphism...Robertson (1999) states that "a massive, compact, fine-grained rock should be
486 classified as a fine-grained granofels (Goldsmith, 1959) and not a hornfels if there is no direct
487 evidence for contact metamorphism." Alternatively, Winkler (1976, p. 327) states that
488 "Hornfels are typically produced by contact metamorphism...and *occasionally by regional*
489 *metamorphism* [italics added]."

490 SLTTM_1.0 defines hornfels as a non-foliated aphanitic rock having granoblastic fabric. The
491 term hornfels does not necessarily denote a contact metamorphic origin (although that is most
492 commonly the case). This is consistent with the goal of keeping this terminology as descriptive
493 as possible. Hornfels, in this sense, is the aphanitic equivalent of granofels.

494 5.3.2 Non-mylonitic foliated metamorphic rock

495 5.3.2.1 Historical usage of schist and gneiss

496 The term "gneiss" is used inconsistently in the geological literature. Some reports distinguish
497 gneiss and schist based on layering whereas others use mineral percentages (e.g., mica vs.
498 feldspar + quartz). Note the etymology of 'gneiss': "alteration of Middle High German *gneiste*
499 spark" [Merriam-Webster's Dictionary, 2004, <http://www.m-w.com/cgi-bin/dictionary>],
500 suggesting the term originally had more to do with composition than with foliation.

501 Definitions of gneiss:

- 502 1. Medium- to coarse-grained rock having a gneissic fabric, i.e. "it splits parallel to 's'
503 generally along mica or hornblende layers, into plates and angular blocks, a few centimeter
504 to tens of centimeters in thickness, or parallel to 'B' into cylindrical bodies (pencil gneiss).
505 The prevalent light-colored constituents (quartz+feldspar) have interlocking boundaries and
506 provided, as compared to schists, a better coherence and a coarser fissility to the rock;
507 nevertheless the fissility in many cases creates an almost perfect plane" (Wenk, 1963,
508 quoted in Winkler, 1979)
- 509 2. Rock having recognizable parallel structure consisting predominantly of quartz and
510 feldspar (feldspar >20%, and mica >10%) Fristch and others, 1967, quoted in Winkler, 5th
511 edition, 1979, p. 342).
- 512 3. A medium- to coarse-grained irregularly "banded" rock that has fairly poor schistosity
513 because of the preponderance of quartz and feldspar; equivalent or higher regional
514 metamorphic grade than a schist (Hyndman, 1972).
- 515 4. Rock composed chiefly of quartz and feldspar; medium- to coarse-grained phaneritic,
516 granoblastic to lepidoblastic; compositional layering expressed by varying modal
517 proportions of quartz, feldspar, mica and hornblende usually evident, but in some granitic
518 gneisses layering is quite subtle or even absent and a weak foliation is expressed by
519 preferred orientation of inequant mineral grains or grain aggregates. (Best, 2002).
- 520 5. Banded rock that does not break along a preferred plane (Blatt and Tracy, 1996). A gneiss
521 has large grains and is foliated i.e. quartz and feldspar rich layers separated from micaceous
522 or mafic layers (Yardley, 1989). Types of Occurrences: Silicic: light colored minerals such
523 as quartz and feldspar make up the majority of the rock; Intermediate: equal amounts of
524 light and dark minerals; Mafic: ferromagnesian minerals make up most of the rock;
525 Undifferentiated (http://www.geol.lsu.edu/rkd_dir/gneiss.html)
- 526 6. A foliated rock formed by regional metamorphism, in which bands or lenticles of granular
527 minerals alternate with bands or lenticles in which minerals having flaky or elongate
528 prismatic habits predominate; generally <50% of minerals show preferred parallel
529 orientation; although commonly feldspar+quartz rich, mineral composition is not an
530 essential factor in its definition (Jackson, 1997).

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7. A rock (regardless of protolith and modal composition) that is medium- to coarse-grained, inhomogeneous, and characterized by a coarse foliation or layering and some layers >5 mm thick (Robertson, 1999).
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8. Gneiss is a roughly foliated or banded metamorphic rock consisting largely of granular minerals such as quartz (<http://www.cst.cmich.edu/users/dietr1rv/gemrxD-K.htm> , gem rock definitions)
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9. A rock that has alternating bands of granular and flaky (or elongate) minerals. Generally less than 1/2 the minerals show a preferred parallel orientation. (<http://forestry.about.com/library/glossary/blforgll.htm>) forestry terms.
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10. A metamorphic rock displaying a gneissose structure (A type of foliation on the hand specimen scale, defined by (a) irregular or ill-defined layering, or (b) augen and/or lenticular aggregates of mineral grains (augen structure, flaser structure), or (c) inequant mineral grains which are present, however, only in small amounts or which display only a weak preferred orientation, thus defining only a poorly developed schistosity.). The term gneiss may also be applied to rocks displaying a dominant linear fabric rather than a gneissose structure, in which case the expression "lineated gneiss" is applied. The term gneiss is used almost exclusively for rocks containing abundant feldspar +/- quartz, but may be used in exceptional cases for other compositions, in which case the exact mineralogy should be given (e.g. feldspar and quartz- free cordierite-anthophyllite-gneiss). (<http://web.met.unimelb.edu.au/TeachingSupport/625-224/224practerms.pdf> , definitions of technical terms for Roger Powell's Metamorphism class).

552 Definitions of schist:

- 553 1. A strongly foliated crystalline rock, formed by dynamic metamorphism, that can be
554 readily split into thin flakes or slabs due to the well developed parallelism of more than
555 50% of the minerals present, particularly those of lamellar or platy habit, e.g. mica and
556 hornblende. The mineral composition is not an essential factor in its definition unless
557 specifically included in the rock name (Jackson, 1997).
- 558 2. Metamorphic rock “Characterised by parallel alignment of moderately coarse grains,
559 usually clearly visible with the naked eye... This type of fabric is known as schistosity...”
560 (Yardley, 1991, p. 22).
- 561 3. “Metamorphic rock commonly of pelitic composition, with a well developed schistosity.”
562 (Barker, 1998, p. 242).
- 563 4. “A metamorphic rock displaying schistose structure. For phyllosilicate-rich rocks the
564 term schist is reserved for medium- to coarse-grained varieties, whereas finer-grained
565 rocks are termed slates or phyllites.” (Schmid and others, 2002).
- 566 5. The main characteristic of “schist” is a well-developed *schistosity*. Schistosity has been
567 defined as:
- 568 6. “The foliation in schist or other coarse-grained, crystalline rock due to the parallel
569 alignment of platy mineral grains (mica) or inequant crystals of other minerals” (Jackson,
570 1997). [This definition is preferred here because the distinction between foliation related
571 to crystallographic orientation of tabular or elongate crystals (schistosity) and grain-shape
572 fabric is important in interpreting the fabric genesis.]
- 573 7. “A preferred orientation of inequant mineral grains or grain aggregates produced by
574 metamorphic processes.” (Schmid and others, 2002).
- 575 8. “A foliation or lineation which allows the rock to be split easily along planes.
576 Constituent minerals can be seen with the unaided eye.” (Robertson, 1999, p. 13)
- 577 9. A “secondary foliation defined by preferred orientation of inequant fabric elements in a
578 medium to coarse grained rock. Individual foliation-defining elements are visible with
579 the naked eye” (Passchier and Trouw, 1996). [This definition allows a schistosity to be
580 defined by the grains that are inequant due to plastic deformation—thus a pure quartz
581 tectonite that has a grain shape fabric defined by tectonically flattened quartz grains
582 would have a schistosity under this definition.]

583 Based on the various definitions of schist and gneiss, the criteria used to distinguish these rock
584 types are:

- 585 • Presence of compositional banding
- 586 • Thickness of compositional banding
- 587 • Homogeneity of the rock
- 588 • Modal abundance of mica
- 589 • Modal abundance of quartz +feldspar
- 590 • Degree of fissility (how thinly can the rock be parted along the ‘schistosity’)
- 591 • Grain size

592 In the IUGS proposal of Brodie and others (2002), a “schistose structure” is characterized by a
593 “schistosity which is well developed, either uniformly through the rock or in narrowly spaced
594 repetitive zones such that the rock will split on a scale of one cm or less” and a “gneissose

595 structure” is characterized by a “schistosity which is either poorly developed throughout the rock
596 or, if well developed, occurs in broadly spaced zones such that the rock will split on a scale of
597 more than one cm.”

598 In practical terms, spitting may be influenced by extraneous variables such as later foliation-
599 parallel joints and weathering. The thickness of sheets split along the foliation in unweathered
600 rocks cannot be used to classify the weathered rocks commonly encountered in the field.
601 Classification would thus be a function of weathering—an undesirable side effect. Criteria based
602 on the thickness of visible layering would not suffer from dependence on the degree of
603 weathering.

604 5.3.2.2 Schistose rock (schist, phyllite, and slate)

605 *Schistose rock*—Any metamorphic rock that has a well developed, continuous schistosity. “Well
606 developed” schistosity is defined to mean that >50% of the rock consists of mineral grains having
607 a tabular, lamellar, or prismatic crystallographic habit that are oriented in a continuous planar or
608 linear fabric (following Jackson, 1997), and “continuous” schistosity is defined on a hand-sample
609 scale, and in quantitative terms, to mean that domains lacking the fabric are <1 cm thick if they
610 are layers, and <5 cm in diameter if they are irregular patches, and constitute <25% of the rock.

611 SLTTM_1.0 subdivides rocks that have well developed continuous schistosity (schistose rocks)
612 into the following:

613 *Schist*—phaneritic rock having average grain size >.25 mm). Schist may have compositional
614 layering at any scale.

615 *Phyllite*—phaneritic rock but finer-grained than schist, having average grain size <0.25 mm but
616 >0.1 mm and typically a silvery sheen on cleavage surfaces.

617 The term phyllite is widely used and generally understood to describe a fine-grained schistose
618 micaceous rock. Robertson (1999) discusses the term as follows:

619 “The term phyllite has previously been used for rocks possessing a silky or lustrous sheen on
620 foliation surfaces imparted by fine-grained (< 0.1 mm) white mica (including muscovite,
621 paragonite and phengite) orientated parallel to the foliation in the rock. Individual mica flakes
622 can be seen with the naked eye in contrast to slates where they cannot be distinguished. Most
623 are probably derived by the low- to medium-grade metamorphism of mudstones although
624 some rocks that have been described as phyllites may have been confused with
625 phyllonites...Here ‘phyllite’ is classified as a specific variant of schist. It is therefore not
626 permissible as a root name but may be used as a specific qualifier, namely *phyllitic*, for
627 example *phyllitic semipelite*.” (Robertson, 1999, p. 12).

628 *Slate*—aphanitic rock, implying average grain size <0.1 mm. In the system proposed here, the
629 term slate is restricted to schistose rocks that are aphanitic except for porphyroblasts. This
630 definition is essentially equivalent to that of Brodie and others [2002].

631 Because “phyllite” and “slate” are widely used on geologic maps in North America, SLTTM_1.0
632 differs from Robertson (1999) and advocates retaining them as rock names. In order for the terms
633 to be useful and unambiguous, criteria for uniquely distinguishing phyllite from schist and slate
634 must be established.

635 The definition of slate as an aphanitic rock displaying a well developed continuous schistosity is
636 consistent with the provisional IUGS definition (Brodie and others, 2002). In addition, it requires
637 that >50% of the rock consists of mineral grains having a tabular, lamellar, or prismatic
638 crystallographic habit that are oriented in a continuous planar or linear fabric. In an aphanitic
639 rock this determination is generally based on indirect evidence, which is typically the presence of
640 slaty cleavage, and the sheen observed on parting surfaces due to alignment of tiny phyllosilicate

641 grains. An average grain size that is aphanitic (<0.1 mm), except for porphyroblasts, is specified
642 here for precision, and the North American spelling of *slaty* (rather than *slatey*; Jackson, 1997) is
643 preferred for use in North America.

644 5.3.2.3 Layered gneiss and non-layered gneiss

645 Gneiss does not have continuous schistosity on a hand sample scale as defined above, but it may
646 have schistose layers separated by non-schistose layers. The boundary of gneiss and schist in this
647 logic is placed where the mineral grains that define the schistosity are distributed such that the
648 schistosity is deemed “continuous.”. SLTTM_1.0 defines gneissose rocks in the following way:

649 *Gneiss*—a foliated phaneritic metamorphic rock that does not have “well developed” continuous
650 schistosity. This follows the spirit of the proposed IUGS classification scheme (Schmid and
651 others, 2002; Brodie and others, 2002). Classified into **layered gneiss** and **non-layered gneiss**.

652 *Layered gneiss*—a foliated, phaneritic rock that lacks well developed, continuous schistosity and
653 has laterally continuous compositional layering > 5 mm thick.

654 *Non-layered gneiss*—

655 5.3.2.4 Migmatitic rocks as subclasses

656 The IUGS proposed definition of migmatite (Wimmenauer and Bryngi, 2002, p. 2) is “A
657 composite silicate rock, pervasively heterogeneous on a meso- to megascopic scale. It typically
658 consists of darker and lighter parts. The darker parts usually exhibit features of metamorphic
659 rocks while the lighter parts are of plutonic appearance...” Wimmenauer and Bryngi (2002) use
660 “migmatite” as a superclass for any rock that has these characteristics.

661 Numerous terms have been employed for varieties of migmatite (Mehnert, 1968), but most are
662 either genetic or difficult to apply on a hand-specimen scale. In the BGS classification,
663 “Migmatite’ is not permissible as a root name...as it is not a single rock type. However,
664 migmatitic may be used as a specific textural modifier” (Robertson, 1999, p. 11). Another
665 problem is that “the scale of migmatitic structures is such that they mainly require definitions,
666 which refer to rock masses greater than the preferred hand specimen size” (Wimmenauer and
667 Bryngi, 2002). The IUGS (Wimmenauer and Bryngi, 2002) recommends ‘migmatite’ as the
668 “main term” for these rocks as well as “special terms” (equated to recommended rock names) for
669 parts of a migmatite (such as leucosome and restite), genetic types (such as anatexite and veinite),
670 and three descriptive types (agmatite = breccia-like; phlebite = veined, nebulite = having diffuse
671 relics of pre-existing rock). The terms for parts of a migmatite represent roles in a relationship,
672 and are thus not suitable as kinds of rock.

673 The system proposed here specifically avoids use of genetic terminology where possible, and the
674 genetic classification of migmatitic rocks is thus outside the scope of the system. The descriptive
675 terms are logically equivalent to other terms in the classification e.g. agmatite = migmatitic
676 granofels (in most cases), and are thus not included as separate terms. Any of these names could
677 be used in an uncontrolled rock name field in a database.

678 Rather than using “migmatite” as a root term or superclass as proposed by Wimmenauer and
679 Bryngi (2002) for IUGS, the classification here follows the British Geological Survey rationale
680 and precedent (Robertson, 1999), which applies “migmatitic” as an adjective modifier to terms
681 such as gneiss and schist. For precision in database applications, this classification requires that
682 the lighter colored, granitoid phase must form >10% of the rock volume.

683 5.3.3 Mylonite-series rock

684 A rock is here classified as a **mylonite-series rock** if the rock displays a foliation defined by the
685 shapes of deformed mineral grains or grain aggregates having aspect ratios > 1.5:1, >10% of the

686 rock is composed of “matrix” showing evidence of tectonic reduction in grain size, and the
687 foliation and matrix are interpreted to be the product of continuous, crystal-plastic deformation
688 processes. Deformation in faults and shear zones occurs in a continuum of environmental
689 conditions from near the Earth’s surface, where discontinuous/brittle processes dominate, to high
690 temperature conditions deep in the Earth where continuous deformation and crystal-plastic
691 processes dominate (Sibson, 1977; Passchier and Trouw, 1996). Foliated, deformed rocks formed
692 in the crystal-plastic regime are included in the mylonite-series.

693 Mylonite-series rocks are commonly linedated. They typically contain fabric elements such as
694 quartz ribbons, mica fish, asymmetric porphyroclasts, S-C composite planar fabrics, and
695 microstructures indicative of crystal-plastic deformation (e.g. recovery and recrystallization, see
696 Sibson, 1977; Passchier and Trouw, 1996; Snoke and Tullis, 1998). The matrix consists of new
697 or recrystallized mineral grains that are interpreted to be smaller than the mineral grains in the
698 original, undeformed protolith. In cases where the protolith was very fine-grained or is unknown,
699 problems in applying this definition may be overcome to some extent by observing the
700 progressive development of fabric and grain-size reduction along the margins of high-strain
701 zones. The distinction between relatively high-temperature cataclastic rocks and mylonite-series
702 rocks is based on the presence of a foliation produced by aligned mineral grains whose shape has
703 been modified by crystal-plastic deformation in mylonite-series rocks.

704 Standard classification schemes for mylonite-series rocks are based on the degree of deformation-
705 related grain-size reduction, which is quantified as the percentage of matrix produced by tectonic
706 reduction in grain size (Sibson, 1977; Wise and others, 1984; Scholz, 1990; Passchier and Trouw,
707 1996; Snoke and Tullis, 1998). Snoke and Tullis (1998, Table 0.1) endorsed Scholz’s (1990,
708 Table 3.1) slightly modified version of Sibson’s (1977) classification of fault rocks. We adopt
709 that classification here with minor modifications for database applications. The mylonite-series
710 rocks, as used here, are equivalent to the “mylonite series” of Sibson (1977), which is subdivided
711 into **protomylonite** (having 10-50% matrix), **mylonite** (having 50-90% matrix), and
712 **ultramylonite** (having 90-100% matrix).

713 **Phyllonite** is a common term for phyllosilicate-rich mylonite-series rock of phyllitic appearance,
714 i.e. having a silvery sheen on cleavage surfaces (Sibson, 1977; Passchier and Trouw, 1996; Snoke
715 and Tullis, 1998; Robertson, 1999, Brodie and others, 2002). SLTTM-1.0 also specifies >40%
716 phyllosilicate minerals (e.g., mica, chlorite) and a *well developed continuous schistosity* as
717 defined in Appendix 2.

718 Under some conditions, recrystallization during deformation may lead to an increase in grain size
719 relative to the protolith of a highly strained rock. Such rocks commonly have equant polygonal
720 mineral grains and a foliation defined by composition or grain-size variations. These rocks are
721 sometimes called “blastomylonite,” but SLTTM_1.0 would classify them as “layered gneiss” or
722 possibly as “non-layered gneiss” if the foliation is defined by flat mineral grains without layering.

723 5.4 Cataclastic rock

724 A rock is here classified as a cataclastic rock if >10% of the volume consists of fragments
725 bounded by fractures related to deformation that occurred after the formation of the protolith.
726 Cataclastic rocks typically include fault rocks formed in the brittle regime, although that origin is
727 not required by the descriptive definition. Deformation in faults and shear zones occurs in a
728 continuum of environmental conditions from near the Earth’s surface, where discontinuous brittle
729 processes dominate, to high temperature conditions deep in the Earth where continuous
730 deformation and crystal-plastic processes dominate (Sibson, 1977; Passchier and Trouw, 1996).
731 Cataclastic rocks are further classified based on the presence or absence of primary cohesion, the
732 percentage of broken fragments large enough to be visible, and the amount of fragmental
733 cataclastic matrix (Sibson, 1977; Snoke and Tullis, 1998).

734 Some cataclastic rocks are foliated, but where this is the case, the foliation is defined by aligned
735 aggregates of rock fragments, rather than oriented mineral grains. Minerals in the same rock may
736 deform by different mechanisms, and loss of material continuity across slip surfaces may vary,
737 resulting some rocks that are transitional between cataclastic rocks and mylonite-series rocks.

738 SLTTM_1.0 adapts Sibson's (1977) widely used classification of cataclastic rocks, as slightly
739 modified by Scholz (1990, Table 3.1) and endorsed by Snoke and Tullis (1998, Table 0.1). The
740 cataclastic rocks are initially subdivided based on the presence or absence of primary cohesion
741 (cohesion during the deformation). Those having evidence of primary cohesion ('**cataclasite**
742 **series**' of Sibson) are subdivided into **protocataclasite** (having 10-50% matrix), **cataclasite**
743 (having 50-90% matrix), and **ultracataclasite** (having 90-100% matrix). Fluid flow along fault
744 zones commonly re-cements rocks that have lost cohesion during deformation, obscuring
745 evidence for the presence or absence of primary cohesion during deformation. **Incohesive**
746 **cataclastic rocks** are those that lack evidence of primary cohesion. These include **fault breccia**
747 (having visible fragments make up >30% of the rock) and **gouge** (having visible fragments <30%
748 of the rock). Cataclastic rocks are further subdivided based on the presence or absence of
749 foliation. **Pseudotachylite**, as discussed below under composite-genesis melt rock, is also
750 regarded as a special type of cataclastic rock. Distinctions between cataclastic rock and **broken**
751 **rock** related to mass wasting (e.g., landslides, rock avalanche, caldera collapse) are based on the
752 geologic setting.

753 5.5 Composite-genesis melt rock

754 5.5.1 Definition

755 *Composite-genesis melt rock*—a rock that solidified from melt (liquid) of a preexisting rock
756 under conditions outside the domain of typical igneous rocks.

757 Examples of composite-genesis melt rock include contact-metamorphic melt rock (melt under
758 contact metamorphic conditions), impact melt rock (also classified as impact-metamorphic rock
759 and discussed under that heading), and pseudotachylite.

760 Pseudotachylite (also spelled pseudotachylyte) is typically a dark, glassy-looking rock that occurs
761 as veins and dike-like injections in fault systems, and it is commonly interpreted as a product of
762 frictional melting along faults (Spray, 1995; Snoke and Tullis, 1998). The IUGS provisional
763 definition for **pseudotachylite** is "Ultrafine-grained vitreous-looking material, usually black and
764 flinty in appearance, occurring as thin planar veins, injection veins or as a matrix to pseudo-
765 conglomerates or breccias, which infills dilation fractures in the host rock" (Brodie and others,
766 2002). SLTTM_1.0 modifies this definition slightly for internal consistency, and to avoid
767 introduction of new terminology:

768 *Pseudotachylite*—an aphanitic, vitreous or flinty material that occupies dilatent fractures
769 associated with fault zones or impact craters, and is interpreted to be related to frictional melting
770 associated with seismic or shock events. It typically contains abundant fractured-rock inclusions
771 and occurs within larger bodies of broken or brecciated rock.

772 "Impact pseudotachylite" is regarded here as a special occurrence of pseudotachylite as discussed
773 below under "impact-metamorphic rock."

774 5.6 Impact-metamorphic rock

775 Impact-metamorphic rock is defined (unavoidably) by genesis and subdivided based on
776 descriptive fabric criteria. In the strictest sense, **impact metamorphism** is "caused by the
777 passage of a shock wave due to impact of a planetary body (projectile or impactor) on a planetary
778 surface (target)" whereas "shock metamorphism" is "caused by shock wave compression due to
779 impact of a solid body or due to the detonation of high-energy chemical or nuclear explosives"

780 (Stöffler, 2001c). Impact metamorphic rocks display evidence of shock metamorphism, such as
781 microscopic planar deformation features within grains or shatter cones (Stöffler and Grieve,
782 2001). Many rocks classified as impact metamorphic rocks are interpreted as such based on field
783 relationships (such as a crater), and such interpretations can be subject to debate.

784 The field classification of impact-metamorphic rocks can be problematic, because of their
785 similarity to some sedimentary or fault-related breccias, and to various fragmental volcanic or
786 pyroclastic rocks in other situations. Where impact metamorphic rocks would meet the
787 descriptive criteria for cataclastic rocks, the distinction of an impact metamorphic rock is based
788 on observable features attributed to shock metamorphism, or by interpretation of field relations at
789 the sample location. The essential feature that identifies an impactite (impact metamorphic rock)
790 is the presence of microscopic planar deformation features that are unequivocally the result of
791 shock metamorphism (Stöffler and Grieve, 2001).

792 Stöffler and Grieve (2001) and Stöffler (2001a, 2001b, 2001c) proposed a threefold IUGS
793 classification of impactites resulting from a single impact event as (i) shocked rock, (ii) impact
794 melt rock, and (iii) impact breccia. The definitions, as slightly modified and adapted for use in
795 this classification, are:

796 **shocked rock**—non-brecciated rock which shows unequivocal effects of shock metamorphism
797 exclusive of whole rock melting.

798 **impact melt rock**—crystalline, semi-glassy, or glassy rock in which $\geq 50\%$ of the rock volume
799 is solidified from impact melt (implying $< 50\%$ non-melt inclusions).

800 **impact breccia**—breccia, as generally defined, that has unequivocal evidence of shock
801 metamorphism.

802 “Impact breccia” is further divided into three subclasses, slightly modified from Stöffler and
803 Grieve (2001) and Stöffler (2001a, 2001b, 2001c), which are: (i) **suevite** [impact breccia that
804 contains impact melt particles (> 0 and $< 50\%$)]; (ii) **polymict impact breccia** [impact breccia that
805 contains fragments of different composition and is free of impact melt particles, and (iii)
806 **monomict impact breccia** [impact breccia free of impact melt in which all of the fragments
807 (100%) have essentially the same composition.”].

808 Pseudotachylite, in addition to being associated with regional faults (see Cataclastic Rocks),
809 occurs in impact structures such as the Vredefort structure in South Africa, where pseudotachylite
810 was first described (discussion and references in Snoke and Tullis, 1998). Pseudotachylite is here
811 classified as a cataclastic rock, whether or not the associated fault or fracture zone happens to be
812 part of an impact structure. “Impact pseudotachylite” as used by Stöffler (2001c) is not
813 distinguished in this classification and, instead, is regarded as a special occurrence of
814 pseudotachylite.

815 Further subdivisions of impact metamorphic rock, including those from multiple impacts as
816 known from the Moon (Stöffler and Grieve, 2001), are not considered here. More specific terms
817 can be used in a text field where desirable. In the case of “shocked rock,” the dual use of a
818 protolith classification (sedimentary, igneous, or metamorphic) can be applied. In a similar
819 fashion, the dual use of an igneous rock classification can provide more information on the
820 composition and texture of “impact melt rock” at places such as Sudbury, Ontario.

821 5.7 Term of last resort in absence of fabric class

822 ‘Rock’ is used as a root name of last resort only where the fabric of the rock is unknown and
823 therefore unspecified. Potential applications may be necessary where geologic map units are
824 inferred from geophysical data or compiled from sources based on other criteria such as age or
825 formation name. If the fabric is known, a more informative term such as schist, gneiss, or

826 granofels should be used. A composition qualifier could be added to rock (e.g., calc-silicate rock,
 827 quartzofeldspathic rock, ultramafic rock).

828 6 CLASSIFICATION BY COMPOSITION (QUALIFIER TERMS)

829 Selection of composition qualifiers is based on the modal mineralogy of a metamorphic rock.
 830 Compositional classification is based on classification of mineral species into groups based on
 831 general chemical similarity. Definitions of the mineral groups that serve to identify the various
 832 composition types are thus fundamental.

833 6.1 Definition of rock-forming mineral groups in this classification

834 **Ferromagnesian minerals.** Omphacite (jadeitic pyroxene), chlorite, dark-colored amphibole,
 835 dark-colored pyroxene, biotite, serpentine, pyrope garnet, talc.

836 **Quartz-feldspar (quartzofeldspathic) minerals:** quartz, plagioclase, K-feldspar.

837 **Calcsilicate minerals:** minerals that contain significant amounts of Ca ± Mg and Si and include
 838 diopside, epidote, grossularite and uvarovite garnet, calcic-amphiboles, titanite, wollastonite,
 839 vesuvianite and calcic plagioclase. Mg-rich minerals such as forsterite and phlogopite are also
 840 common constituents of calcsilicate-rocks. As a general rule, plagioclase may be considered a
 841 calcsilicate mineral if it has >50% anorthite content (Robertson, 1999).

842 **Carbonate minerals:** calcite, dolomite, siderite.

843 **Aluminous minerals:** aluminosilicates, muscovite, kaolinite, garnet (associated with feldspar),
 844 corundum, pyrophyllite.

845 **Phyllosilicate minerals:** mica group, chlorite group

846 Garnets occur in aluminous, ferromagnesian, and calcsilicate rocks. Cordierite, staurolite,
 847 brucite, and periclase are not useful for compositional rock classification without more detailed
 848 knowledge of their mineral associations or compositions.

849 Table 2. Summary of composition terms

Qualifier	definition	source	kind
amphibolite	Rock that consists of >75% green, brown, or black amphibole plus plagioclase (including albite) and amphibole >30% (modal) of whole rock, and amphibole >50% of total mafic constituents.	Coutinho and others (2002)	common
argillic	use for apparently clay-rich aphanitic rocks	based on Jackson, 1997	standard chemical
calcareous	When mineralogy cannot be identified (aphanitic rocks), means that rock reacts to form bubbles when hydrochloric acid is applied.	Jackson (1997); criteria proposed by this report	standard chemical
calcareous quartzofeldspathic	rock consists of 10%-50% carbonate or calcsilicate minerals, and micaceous or aluminous minerals form <40% of non-(carbonate or calcsilicate) minerals and quartz forms <60% of the non-(carbonate or calcsilicate) minerals	Robertson (1999)	standard chemical
calcareous-pelitic	rock consists of 10%-50% carbonate or calcsilicate minerals, and micaceous or aluminous minerals form ≥ 40% of non-(carbonate or calcsilicate) minerals	Robertson (1999)	standard chemical

Qualifier	definition	source	kind
calcareous-quartzite	rock consists of 10%-50% carbonate or calcsilicate minerals, and micaceous or aluminous minerals form <40% of non-(carbonate or calcsilicate) minerals, and quartz forms $\geq 75\%$ of the non-(carbonate + calcsilicate) minerals	Robertson (1999)	standard chemical
calcareous-semi-pelitic	Rock for which the sum of modal quartz+feldspar+mica + aluminous mineral is $\geq 70\%$, and quartz+feldspar < 60% and carbonate + calcsilicate minerals > 10%	Proposed, this report	standard chemical
calcite marble	carbonate minerals form > 75% of rock, and calcite forms >75% of carbonate minerals.	modified from Robertson (1999)	common
calcsilicate	rock consists of $\geq 50\%$ calcsilicate or carbonate minerals and carbonate minerals \leq calcsilicate minerals in mineral mode. When used for aphanitic rocks, indicates quartz or feldspar is significant in the mineral mode, and the rock does not meet calcareous.	Barker (1998, Appendix II), Robertson (1999)	standard chemical
dark-colored	use for dark-colored aphanitic rocks about which other information not available	descriptive	color
dolomite marble	carbonate minerals form > 75% of rock, and dolomite forms >75% of carbonate minerals.	modified from Robertson (1999)	common
eclogite	Rock composed of >75% garnet (almandine-pyrope) and sodic pyroxene (omphacite).	Carswell (1990), Barker (1998, Appendix II), modified from Jackson (1997)	common
epidosite	Rock consisting of >75% epidote, and >50% of non-epidote mineral is quartz.	modified from Jackson (1997)	monomineralic
ferromagnesian	rock having >40% dark ferromagnesian minerals. Standard term defined by Bates and Jackson (1987) to mean "containing iron and magnesium"	Usage proposed, this report	superclass
greisen	Hydrothermal rock consisting of >70% quartz + muscovite or lepidolite, and having accessory (>1%) topaz or tourmaline (or other fluorine-bearing phase). In a hierarchy based purely on description, would be a kind of semi-pelitic rock (10-40% aluminous minerals; <60% quartz + feldspar) or quartz-feldspathic (<40% aluminous and >60% quartz + feldspar)	modified from Jackson (1997)	common
impure marble	rock consists of >50% calcsilicate or carbonate minerals and relative proportion of calcsilicate and carbonate minerals is unknown or not specified	Usage proposed, this report	superclass
light-colored	use for light-colored aphanitic rocks about which other information not available	descriptive	color
mafic	rock consists of $\geq 40\%$ and <90% ferromagnesian minerals.	modified from Robertson (1999)	standard chemical
magnesian	No definition at present., but may be useful?		

Qualifier	definition	source	kind
marble	rock in which carbonate minerals form > 75% of rock	Barker (1998, Appendix II), Robertson (1999)	common
metacarbonate	rock consists of >50% calcsilicate or carbonate minerals and carbonate minerals > calcsilicate minerals in mineral mode.	Robertson (1999)	standard chemical
monomineralic	Rock that consists of >75% of a single mineral species and does not meet any of the other composition terms (e.g. quartzite, calcite marble, dolomite marble, serpentinite...)	for consistency with Robertson (1999)	superclass
None	Composition not specified	this scheme	superclass
pelitic	Rock for which the sum of modal quartz+feldspar+mica + aluminous mineral is $\geq 70\%$, and aluminous mineral + mica content is $\geq 40\%$.	modified from Robertson (1999)	standard chemical
quartz-feldspar-pelitic	Rock for which the sum of modal quartz+feldspar+mica + aluminous mineral is $\geq 70\%$	Usage proposed, this report	superclass
quartzite	Rock that consists of $\geq 75\%$ quartz	Robertson (1999), revised to be consistent with other monomineralic rocks.	common
quartzo-feldspathic	Rock for which the sum of modal quartz+feldspar+mica + aluminous mineral is $\geq 70\%$, and quartz + feldspar >60%.	revised from Robertson (1999)	standard chemical
semi-pelitic	Rock for which the sum of modal quartz+feldspar+mica + aluminous mineral is $\geq 70\%$, and quartz+feldspar < 60%.	revised from Robertson (1999)	standard chemical
serpentinite	Rock that consists of >75% serpentine.	Barker (1998, Appendix II), Jackson (1997), revised to be consistent with other monomineralic rocks.	monomineralic
silicic	Use for apparently siliceous aphanitic rocks. Bates & Jackson (1987) include denotation of igneous origin. For this classification, should be considered to mean "appears to consist largely of quartz and feldspar", generally is aphanitic with hardness ≥ 6 , and the rock does not meet calcareous or calcsilicate.	Usage proposed, this report	standard chemical
skarn	Not recommended. Usage too problematic; distinction from calcsilicate hornfels or calc-silicate granofels not clear.	Einaudi (1982), Einaudi and Burt (1982), Barker (1998, Appendix II)	common

Qualifier	definition	source	kind
special composition	Rock that has a mineral composition that doesn't fit in any defined composition class. A modal mineral description is essential. The rock consists of <40% ferromagnesian minerals and <50% carbonate + calcsilicate minerals and <70% Q+Fs + mica + aluminous minerals.	Usage proposed, this report	superclass
ultramafic	Rock that consists of >90% ferromagnesian minerals.	Robertson (1999)	standard chemical
whiteschist	Not recommended. Use pelitic rock if kyanite + garnet + white mica + quartz >70%, and ky + garnet + white mica must be ≥ 40%, otherwise whiteschist is a kind of special composition metamorphic rock. Whiteschist is defined as a rock consisting of talc, kyanite, white mica (paragonite), garnet, or (sodic whiteschist) jadeite, Mg-glaucophane, kyanite, quartz, and garnet. Desmons and others' (2002) IUGS proposal notes that "Whiteschist is defined in a loose descriptive manner, in a similar way to blueschist... More precise rock terms should be used wherever possible (e.g., kyanite-talc-phengite schist)."	Schreyer (1973, 1977), Meyre and others (1999); Desmons and others (2002)	common

850 6.2 Quartz-feldspar-pelitic rocks

851 Pelite has three definitions in the *Glossary of Geology* (Jackson, 1997): (1) "a sediment or
 852 sedimentary rock composed of clay- or mud-sized particles," (2) "a fine-grained sedimentary rock
 853 composed of more or less hydrated aluminum silicates with which are mingled small particles of
 854 various other minerals; an aluminous sediment," and (3) "the metamorphic derivative of a lutite,
 855 such as the metamorphosed product of a siltstone or mudstone (as commonly used a pelite means
 856 an aluminous sediment metamorphosed, but if used systematically, it means a fine-grained
 857 sediment metamorphosed)." Since Al-rich clays are a major component of most clay-sized
 858 sediment, the distinction between the meaning of "pelite" as a grain size and the compositional
 859 connotation has become blurred. The term "pelite" has a strong connotation of aluminous
 860 composition when used to describe metamorphic rocks, and the British Geological Survey
 861 classification of metamorphic rocks uses it in that sense (Robertson, 1999).

862 SLTTM_1.0 follows the British Geological Survey's lead in emphasizing the aluminous
 863 composition of pelite, but differs in using the adjective form "pelitic-" as a compositional
 864 qualifier prefix to a fabric term as in "pelitic-schist." We use the following classification
 865 categories (Table 2):

866 **Pelitic**—a rock composition having modal percentages of quartz + feldspar + mica + aluminous
 867 minerals ≥ 70% is here classified as if aluminous minerals + mica ≥ 40%.

868 **Semipelitic**—quartz +feldspar <60%

869 **Quartzofeldspathic**—quartz + feldspar >60%.

870 SLTTM_1.0 does not consider biotite mica to be an aluminous or "pelitic" mineral.

871 6.3 Monomineralic rocks

872 Rocks are considered to be monomineralic if they consist predominantly of a single mineral or
 873 closely related mineral group. Common monomineralic rocks such as marble (>75% carbonate),
 874 quartzite (>75% quartz), and serpentinite (>75% serpentine) are defined individually in the

875 classification. Otherwise, “monomineralic” in the SLTTM_1.0 classification system is used as a
876 general prefix to avoid including separate “-ite” names for nearly every mineral. Metamorphic
877 rocks consisting predominantly (>75%) of a single mineral are classified as monomineralic-
878 granofels, monomineralic-hornfels, monomineralic-schist, etc., depending on the fabric. If
879 someone is interested in a specific monomineralic rock, the query would search the modal-
880 mineralogy description for >75% of mineral Z. The user interface could have rock names such as
881 hornblendite, biotitite, and tourmalinite stored in the database as ‘monomineralic metamorphic
882 rock (or granofels, schist, hornfels)’ and specify modal mineralogy >75% of the appropriate
883 mineral.

884 6.4 Calcsilicate and metacarbonate rocks

885 Calcsilicate and metacarbonate rocks are rocks that are composed of $\geq 50\%$ calcsilicate or
886 carbonate minerals. The term **marble** implies that a significant percentage of the rock consists of
887 carbonate mineral. This follows the definition in Jackson (1997): “a metamorphic rock consisting
888 predominantly of fine- to coarse-grained recrystallized calcite or dolomite, usually with a
889 granoblastic, saccharoidal texture”.

890 SLTTM_1.0 defines **marble** as a monomineralic rock in which carbonate minerals form > 75%
891 of rock. For rocks that meet the mineralogical definition of marble, but do not have a
892 granoblastic fabric, the term marble becomes a strictly compositional modifier. Although a
893 marble may consist of several different minerals (calcite, dolomite, siderite...), it is considered
894 monomineralic because they all belong to the same, closely related mineral carbonate minerals
895 group. The name **impure marble** is suggested as a general term for rocks that consist of >50%
896 but $\leq 75\%$ carbonate and calc-silicate minerals, and **metacarbonate rock** for an **impure marble**
897 in which carbonate minerals are more abundant than calc-silicate minerals.

898 6.5 Ferromagnesian rocks

899 Ferromagnesian rocks are those that consist of >40% dark ferromagnesian minerals. Rocks
900 included in this group include eclogite, mafic and ultramafic metamorphic rocks, and
901 amphibolite.

902 The *Glossary of Geology* (Jackson, 1997) defines **amphibolite** as “A crystalloblastic rock
903 consisting mainly of amphibole and plagioclase with little or no quartz. As the quartz content
904 increases, the rock grades into hornblende-plagioclase gneiss.” A wide range of accessory
905 minerals can be present.

906 Most amphibolites are metamorphosed mafic igneous rocks (ortho-amphibolites) but some may
907 be metamorphosed calcareous sediments (para-amphibolites) (Yardley, 1991). Coutinho and
908 others (2002) compiled the modal mineralogy of numerous rocks described as “amphibolite”
909 from different localities, noting that they fit the common definition of amphibolite as a
910 hornblende + plagioclase rock. Their data show that most amphibolites contain >50% amphibole
911 although 30% to 50% is not unusual, that the amphibole + plagioclase is mostly >90% and less
912 commonly as low as 75%, and that most of the rocks described as amphibolite have <10% quartz.
913 Based on these and other data, Coutinho and others (2002) proposed the following IUGS
914 definition:

915 “Amphibolite is a gneissose or granofelsic metamorphic rock mainly consisting of green,
916 brown, or black amphibole and plagioclase (including albite), which combined constitute more
917 than 75% of the rock. The amphibole constitutes more than 50% of the total mafic constituents
918 and is present in an amount more than 30% [*italics added*]. Other common minerals include
919 quartz, clinoproxene, garnet, epidote-group minerals, biotite, titanite, and scapolite.
920 Orthopyroxene is absent.”

921 Members of the SLTT Subgroup on Metamorphic Rocks object to the exclusion of
922 orthopyroxene, which can be present in these rocks at uppermost amphibolite facies, and to the
923 inclusion of scapolite as a "common" mineral. Scapolite may occur in amphibolites of calc-
924 silicate affinity but it is not "common." The main part of Coutinho and others' (2002) proposed
925 IUGS definition shown in italics is adopted here, without requiring the absence of orthopyroxene.

926 SLTTM_1.0 defines **amphibolite** as a metamorphic rock composed mostly of green, brown, or
927 black amphibole and plagioclase (including albite) so that amphibole + plagioclase >75% and
928 amphibole >30% (modal percent) of the whole rock, and amphibole >50% of the total mafic
929 constituents.

930 6.6 Special-composition rocks

931 Special non-systematic rock names such as whiteschist, rodingite, fenite, skarn, gondite, coticule,
932 and greisen are commonly used in some settings. Their definitions typically are based on modal
933 mineralogy, but not fabric, and as such may generally be considered as composition qualifier
934 terms. This classification leans towards a minimum of special rock names for rocks of unusual
935 composition, and relies on the presence of a modal-mineralogy description included in the
936 database to provide a mechanism to describe and search for such rocks. Thus, many of these
937 special rocks would be assigned a composition qualifier 'special composition'. The uncontrolled
938 rock name field in the database is available to assign any special rock name the geologist may
939 prefer.

940 Some of these special-composition rock names may be added as subtypes of special-composition
941 rock, which are identified by decision switches near the top of the classification decision tree.
942 The problem posed by many of these special names is that their definitions overlap other
943 composition-qualifier classes. For example, in the composition-qualifier hierarchy, greisen is a
944 subclass of quartz-feldspar-semipelitic rock but may overlap semipelitic rock, quartzofeldspathic
945 rock, or quartzite, depending on the specific modal mineralogy. Thus, placement of these special
946 terms in the same class hierarchy results in ambiguous classification (should the rock be classified
947 as greisen or as semi-pelitic granofels?), unless a modal mineral analysis description is included.
948 Use of special rock names must be constrained by requiring that a modal-mineral description be
949 provided for any rock for which classification using the special name results in ambiguous
950 placement in the composition qualifier hierarchy. Again, this classification scheme avoids such
951 special terms as much as possible. Such names are better used in the uncontrolled lithology name
952 field of the database.

953 7 CLASSIFICATION BY PROTOLITH

954 Two groups that have recently developed or are in the process of developing terminology for
955 metamorphic rocks (Robertson, 1999; Schmid and others, 2002) include protolith as an important
956 classification criterion in cases where protolith is clearly identifiable. In many metamorphic
957 terranes the protolith for metamorphic rocks is obvious, and in such terranes (e.g. Canadian
958 Shield), rocks are commonly mapped based on protolith, and named by adding the prefix 'meta'
959 to the protolith name to indicate that the rock has been metamorphosed. For many users, such as
960 mineral explorationists and resource appraisers, protolith is the most important (commonly the
961 only important) feature of these rocks, and any useful geologic map database must represent the
962 protolith identification. Why not include protolith-based names (e.g., metabasalt, metagranite,
963 metaconglomerate) in this classification?

964 We acknowledge that, for rock-naming purposes, use of the 'meta' prefix has been and no doubt
965 will continue to be very useful as a way for geologists to (informally) communicate information
966 about metamorphic rocks. However, we are concerned that, for incorporation into geoscience
967 databases where fabric and composition are clearly key attributes of rock description, the tri-fold
968 naming scheme that includes protolith in the rock name (e.g., metabasalt, metaquartzite,
969 metalimestone) clogs up the works and introduces a highly subjective element of genetic
970 interpretation into an exercise that we are trying to keep as descriptive as possible. The
971 classification system proposed here is designed for databases (automated knowledge
972 representation), and in this context, use of the 'meta' prefix is equivalent to dual classification--a
973 protolith classification using the appropriate system for the determined protolith, and
974 classification as 'metamorphic rock' in our proposed composite-genesis classification. The usage
975 tells nothing about what kind of metamorphic rock the material is, beyond whatever connotations
976 the protolith rock name has for composition. Accordingly, SLTTM_1.0 recommends that rock-
977 name terms in the form 'meta(*some rock name*)' are to be placed in an uncontrolled rock name
978 field, designed to allow the geologist to assign the name that is most meaningful for local or
979 traditional use. Such terms also can appear in a user interface for classifying rocks by having
980 underlying software that maps the name assignment to the implied dual classification. To keep
981 the meaning of the classification clear, a metamorphic rock must be classified based on the
982 aspects of the rock that make it a metamorphic rock, i.e., fabric and composition.

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1178 <http://www.bgs.ac.uk/SCMR>
- 1179

1180 10 APPENDIX 2. SELECTED FABRIC TERMS AND DEFINITIONS

1181 **Aphanitic**—“Individual grains not visible with the unaided eye (ca. <0.1 mm)” (Schmid and
1182 others, 2002) [“Said of the texture of an igneous rock in which the grains are too small to
1183 distinguish with the unaided eye...” (Jackson, 1997) and extended to metamorphic rocks
1184 by Schmid and others (2002).]

1185 **Cataclastic**—“Pertaining to the structure produced in a rock by the action of severe mechanical
1186 stress during dynamic metamorphism; characteristic features include the bending,
1187 breaking, and granulation of minerals. Also said of the rocks exhibiting such structures.”
1188 (Jackson, 1997)

1189 **Cataclastic rock** --“A rock, such as tectonic breccia, containing angular fragments that have been
1190 produced by the crushing and fracturing of preexisting rocks as a result of mechanical
1191 forces in the crust.” (Jackson, 1997). Mylonite-series rock is not classified here as a
1192 cataclastic rock, although it may contain subordinate cataclastic fabric.

1193 **Cleavage**—“The property of a rock to split along a regular set of parallel or sub-parallel closely
1194 spaced surfaces” (Brodie and others, 2002).

1195 **Continuous foliation** –Foliation in which the fabric elements are uniformly distributed. For
1196 precision, *continuous* is here defined on a hand sample-scale, and in quantitative terms to
1197 mean that domains lacking the fabric are <1 cm thick if they are layers, and <5 cm in
1198 diameter if they are irregular patches, and constitute <25% of the rock. Distinct from
1199 spaced foliation (Passchier and Trouw, 1996, p. 64-65).

1200 **Compositional layering**--Non-genetic term for foliation defined by layers of different
1201 composition (Passchier and Trouw, 1996). Compositional layering is here considered to
1202 be *laterally continuous* if layers can be traced >10 cm.

1203 **Fabric element**-- Part of a rock fabric such as a foliation, lineation, etc. (Jackson, 1997;
1204 Passchier and Trouw, 1996). Any of the elementary parts of an aggregate whose
1205 orientation and geometry can be described; only structures that are penetrative on the
1206 scale of the domain contribute to the fabric (Turner and Weiss, 1963).

1207 **Fabric**--The complete spatial and geometrical configuration of all those components that are
1208 contained in a rock and that are penetratively and repeatedly developed throughout the
1209 volume of rock under consideration; the shapes and characters of individual parts of a
1210 rock mass and the manner in which these parts are distributed and oriented in space
1211 (paraphrase of Passchier and Trouw, 1996, and Jackson, 1997), both of which cite Hobbs
1212 and others (1976, p. 73). [English translation of German word *Gefüge* used by Sander
1213 (1930) to denote the internal ordering of both geometric and physical spatial data in an
1214 aggregate (Turner and Weiss, 1963). In a non-crystal, the internal geometric
1215 configuration of the elementary parts and of any characteristic features to which the
1216 arrangement of these parts gives rise [paraphrase from, Paterson & Weiss (1961, p. 854)
1217 quoted in Turner & Weiss (1963, p. 19).]

1218 **Foliation**—Planar fabric in a rock that is penetrative on a mesoscopic scale (Turner and Weiss,
1219 1963; Passchier and Trouw, 1996; Brodie and others, 2002).

1220 **Grain-shape fabric**--Fabric defined by shape-preferred orientation of mineral grains or mineral
1221 grain aggregates (Passchier and others, 1990; Passchier and Trouw, 1996, p.74).

1222 **Gneissosity**—“General term for foliation in a gneiss. Use of this term is discouraged because of
1223 its vague connotation; several types of foliation (layering, schistosity) may occur in the
1224 same gneiss.” (Passchier and Trouw, 1996, Glossary)

- 1225 **Granoblastic**--Fabric dominantly formed by equidimensional crystals. An example is a foam
1226 structure...in which platy or acicular mineral shapes are absent, and most grains have
1227 equant shape. (Passchier and Trouw, 1996, Glossary)
- 1228 **Migmatitic**--Having the characteristics of migmatite, which is “A coarse-grained heterogeneous
1229 rock type characteristically with irregular and discontinuous interleaving of leucocratic
1230 granitoid material (leucosome) and residual high-grade metamorphic material (restite).
1231 (Barker, 1998, Appendix II: Glossary)
- 1232 **Mylonite-series rock**--A collective term for the “mylonite series” rocks of Sibson (1977),
1233 including protomylonite, mylonite, ultramylonite, and phyllonite. Equivalent to Brodie
1234 and others’ (2002) broader usage of “mylonite” for “A fault rock which is cohesive and
1235 characterized by a well developed [foliation] resulting from tectonic reduction of grain
1236 size, and commonly containing rounded porphyroclasts and lithic fragments of similar
1237 composition to minerals in the matrix.” By this usage, mylonite-series rock is not a
1238 cataclastic rock; cataclastic fabrics, if present, are subordinate.
- 1239 **Penetrative**—Repeated at distances so small, compared with the scale of the whole...that they
1240 can be considered to pervade it uniformly and be present at every point (paraphrase of
1241 Turner and Weiss, 1963, p. 21).
- 1242 **Penetrative fabric element**--A fabric element that occurs penetratively throughout a rock at the
1243 scale of observation (Passchier and Trouw, 1996).
- 1244 **Phaneritic**—“Individual grains visible with the unaided eye (ca. >0.1 mm)” (Schmid and others,
1245 2002). [“Said of the texture of an igneous rock in which the grains are large enough to
1246 distinguish with the unaided eye, i.e. megascopically crystalline” (Jackson, 1997) and
1247 extended to metamorphic rocks by Schmid and others (2002).]
- 1248 **Schistosity**--Foliation in a rock due to the parallel, planar arrangement of mineral grains of the
1249 platy, prismatic or ellipsoidal types, usually mica (Jackson, 1997). *Well developed*
1250 *schistosity* is here defined to mean that >50% of the rock consists of mineral grains
1251 having a platy, lamellar, tabular, or prismatic crystallographic habit that are oriented in a
1252 continuous planar or linear fabric (following Jackson, 1997). *Continuous* is here defined
1253 on a hand sample-scale, and in quantitative terms to mean that domains lacking the fabric
1254 are <1 cm thick if they are layers, and <5 cm in diameter if they are irregular patches, and
1255 constitute <25% of the rock.
- 1256 **Slaty cleavage**—A “well developed planar schistosity in a rock in which the individual grains are
1257 too small to be seen by the unaided eye and the schistosity is developed on the grain
1258 scale.” (Brodie and others, 2002)
- 1259 **Spaced foliation**—Foliation in which the fabric elements are separated by domains that lack the
1260 foliation; not “continuous” as defined above. Discussion in Passchier and Trouw (1996,
1261 p. 65).

1262 11 APPENDIX 3. METAMORPHIC FACIES AND TYPES

1263 This appendix presents recommendations (initially proposed by Subgroup members Hoisch and
1264 Williams) for the classification of metamorphic facies and types of metamorphism. These
1265 recommendations assume as a prerequisite that any rock should fit into only one facies and only
1266 one type of metamorphism. We exclude the hornfels and sanidinite facies, as done by Spear
1267 (1993) because the mineral assemblages are no different than in the higher-pressure facies. The
1268 only distinction appears to be textural, and that will appear in other attribute lists (rocks called
1269 hornfels or hornfelsic). "Sanidinite facies" is essentially synonymous with pyrometamorphism
1270 and consequently is not considered a true facies. We suggest including it in the granulite facies.
1271 Common usage for pyrometamorphism includes ultra-high temperature contact metamorphism,
1272 which poses the problem of a non-unique definition, unless it is confined to metamorphism by
1273 burning coal seams. That seems appropriate, since the root "pyro" means fire. It is an uncommon
1274 but spectacular type of metamorphism. The term "ultrametamorphism" means partially melted,
1275 but this overlaps both the upper amphibolite and granulite facies, and so we exclude it here.
1276 Lithologic terminology for partially melted rocks should follow the IUGS recommendations of
1277 Wimmenauer and Bryhni (2002) as much as possible.

1278 We do not define the metamorphic facies, at this stage, because that would involve numerous
1279 debates on their boundaries. If we absolutely must define them, then some general statements can
1280 be added.

1281 The lists below are recommendations for consideration by others. There is quite a range of
1282 published opinion. We tried not to make arbitrary decisions, but it is unlikely that any set of
1283 recommendations will please everyone.

1284 **Types of metamorphism:**

- 1285 1) **Contact metamorphism**--Metamorphism of country rock at the contact of an igneous body.
1286 2) **Regional metamorphism**--Metamorphism not obviously localized along contacts of igneous
1287 bodies; includes burial metamorphism and ocean ridge metamorphism.
1288 3) **Hydrothermal metamorphism (metasomatism)**--Metamorphism involving significant
1289 changes in a rock's bulk chemistry as a result of interaction with chemically reactive fluids.
1290 4) **Shock (impact) metamorphism**—Shock metamorphism is “caused by shock wave
1291 compression due to impact of a solid body or due to the detonation of high-energy chemical or
1292 nuclear explosives” (Stöffler, 2001c). Impact metamorphism is “caused by the passage of a shock
1293 wave due to impact of a planetary body (projectile or impactor) on a planetary surface (target). It
1294 includes melting and vaporization of the target rock(s)” (Stöffler, 2001c).
1295 5) **Pyrometamorphism**--Ultra-high temperature metamorphism at shallow depths caused by
1296 burning coal seams.

1297 Metamorphic facies:

- 1298 1) Zeolite
1299 2) Prehnite-pumpellyite
1300 3) Greenschist
1301 4) Epidote amphibolite
1302 5) Amphibolite
1303 6) Granulite
1304 7) Blueschist (glaucophane schist)

1305 8) Eclogite

1306 9) Ultra-high pressure

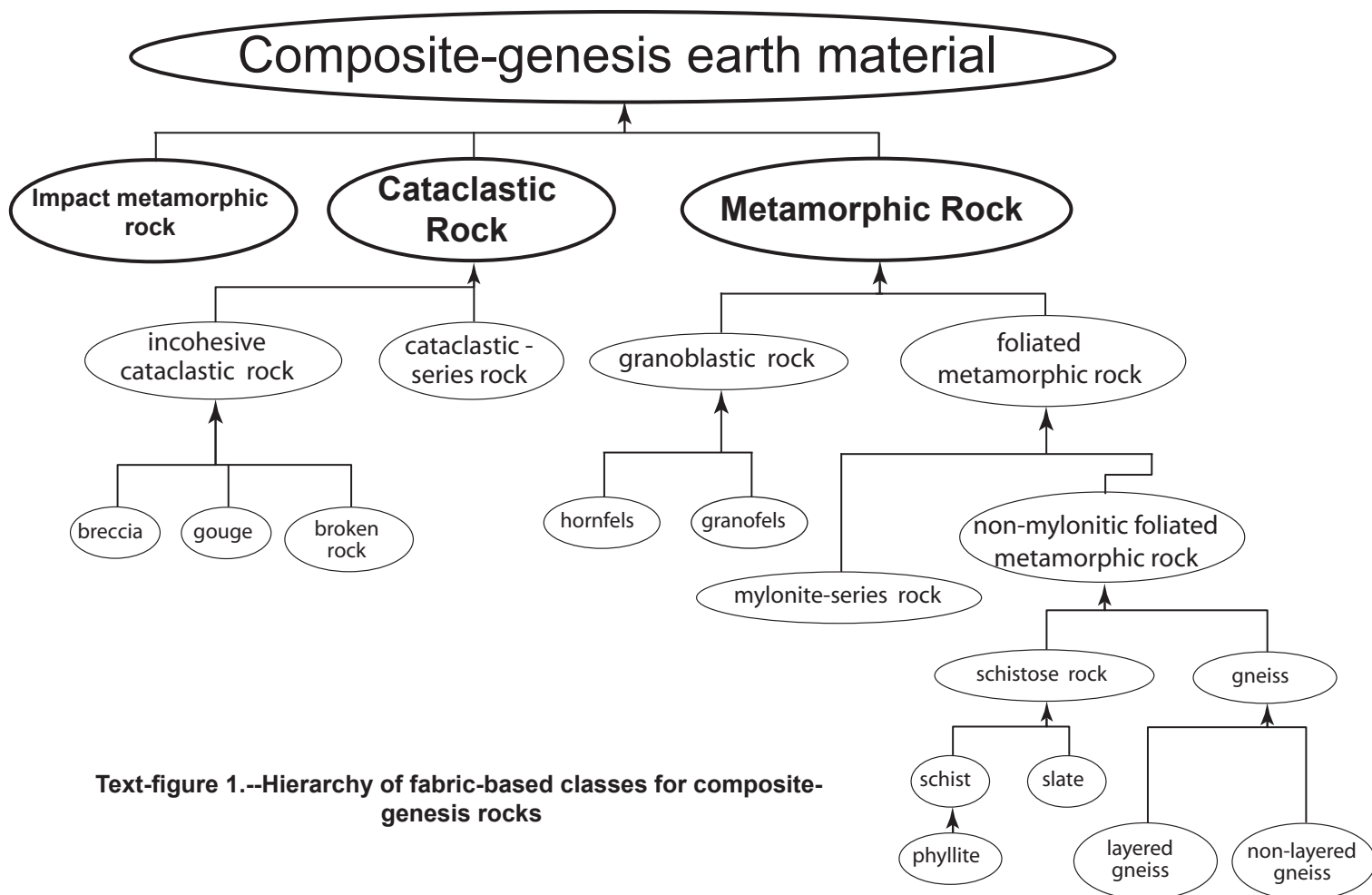
1307 This proposed list of metamorphic facies differs from the IUGS recommendations of
1308 Smulikowski and others (1997), who also include the sanidinite facies (here included in granulite
1309 facies) and pyroxene hornfels facies. The glaucophane schist facies as used by Smulikowski and
1310 others (1997) is another name for the blueschist facies as listed here.

1311 12 APPENDIX 4. GLOSSARY OF METAMORPHIC ROCK TERMS AND PARENT-CHILD
1312 RELATIONSHIPS (DATABASE)

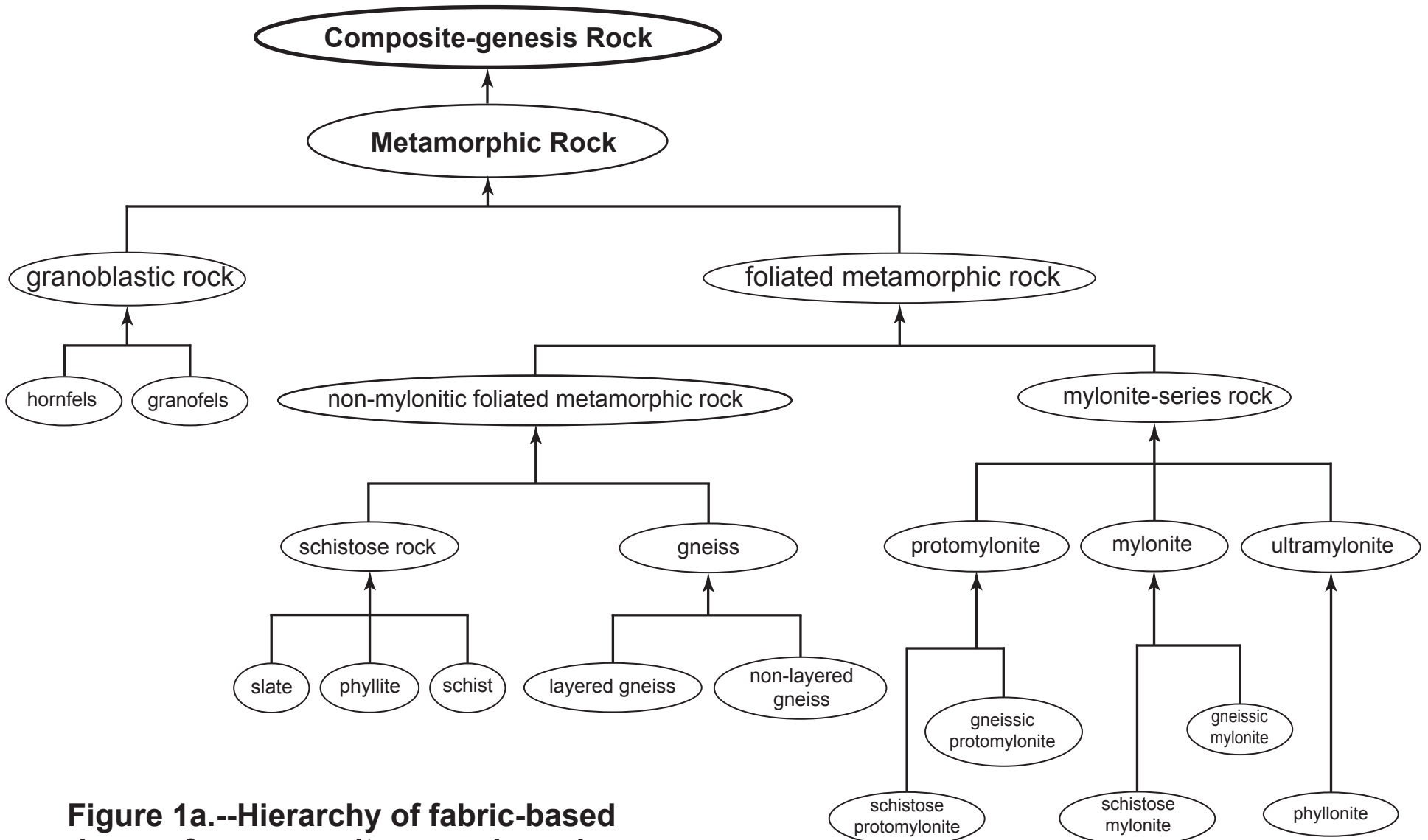
1313

1314 Glossary is included in accompanying Microsoft Access database.

Figure 1



Text-figure 1.--Hierarchy of fabric-based classes for composite-genesis rocks



- 50 -

Figure 1a.--Hierarchy of fabric-based classes for composite-genesis rocks: conventional metamorphic rocks

Figure 1b

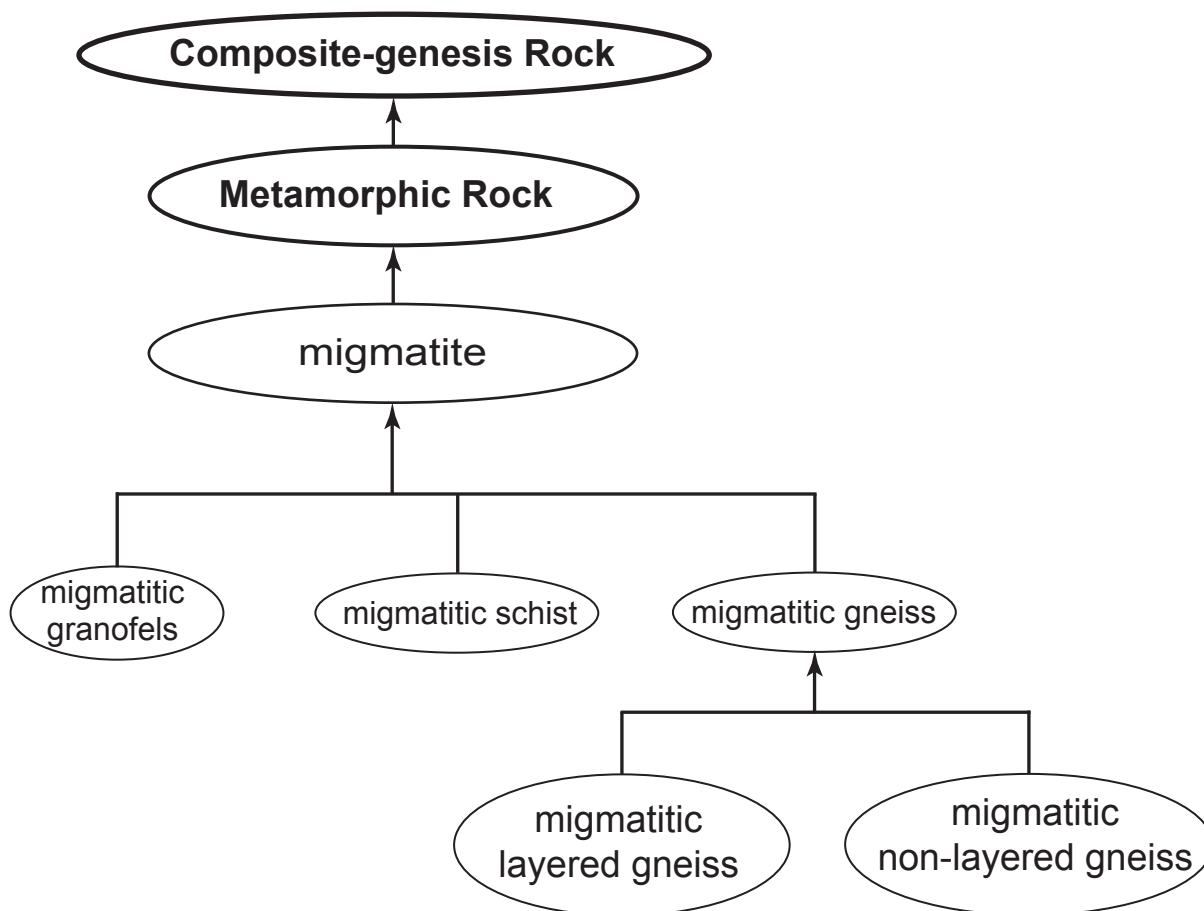


Figure 1b.--Hierarchy of fabric-based classes for composite-genesis rocks: migmatitic metamorphic rocks

Figure 2

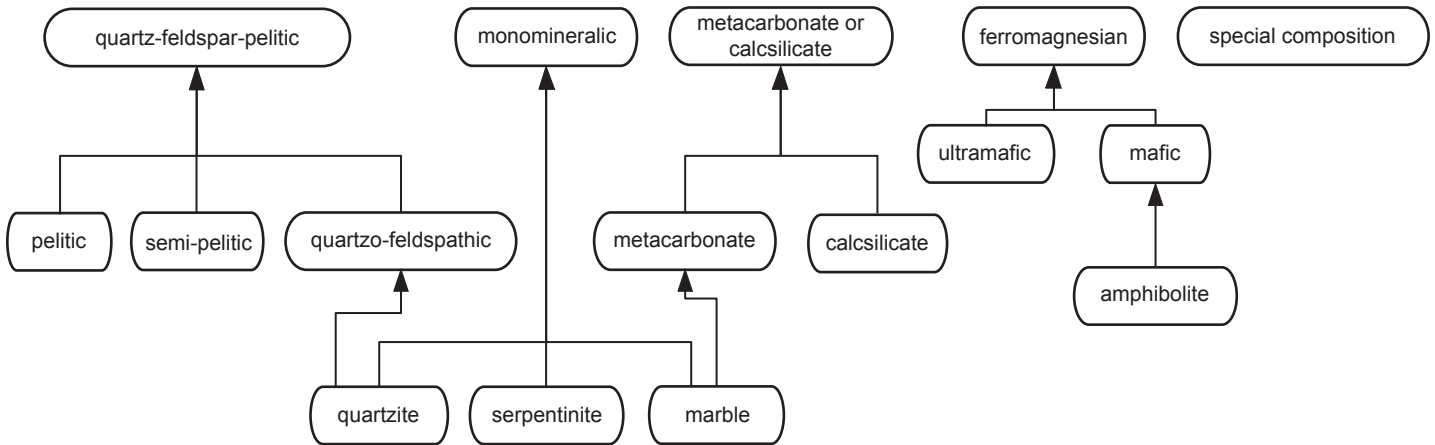


Figure 2.--Hierarchy of composition qualifier terms for composite genesis rocks

Figure 3

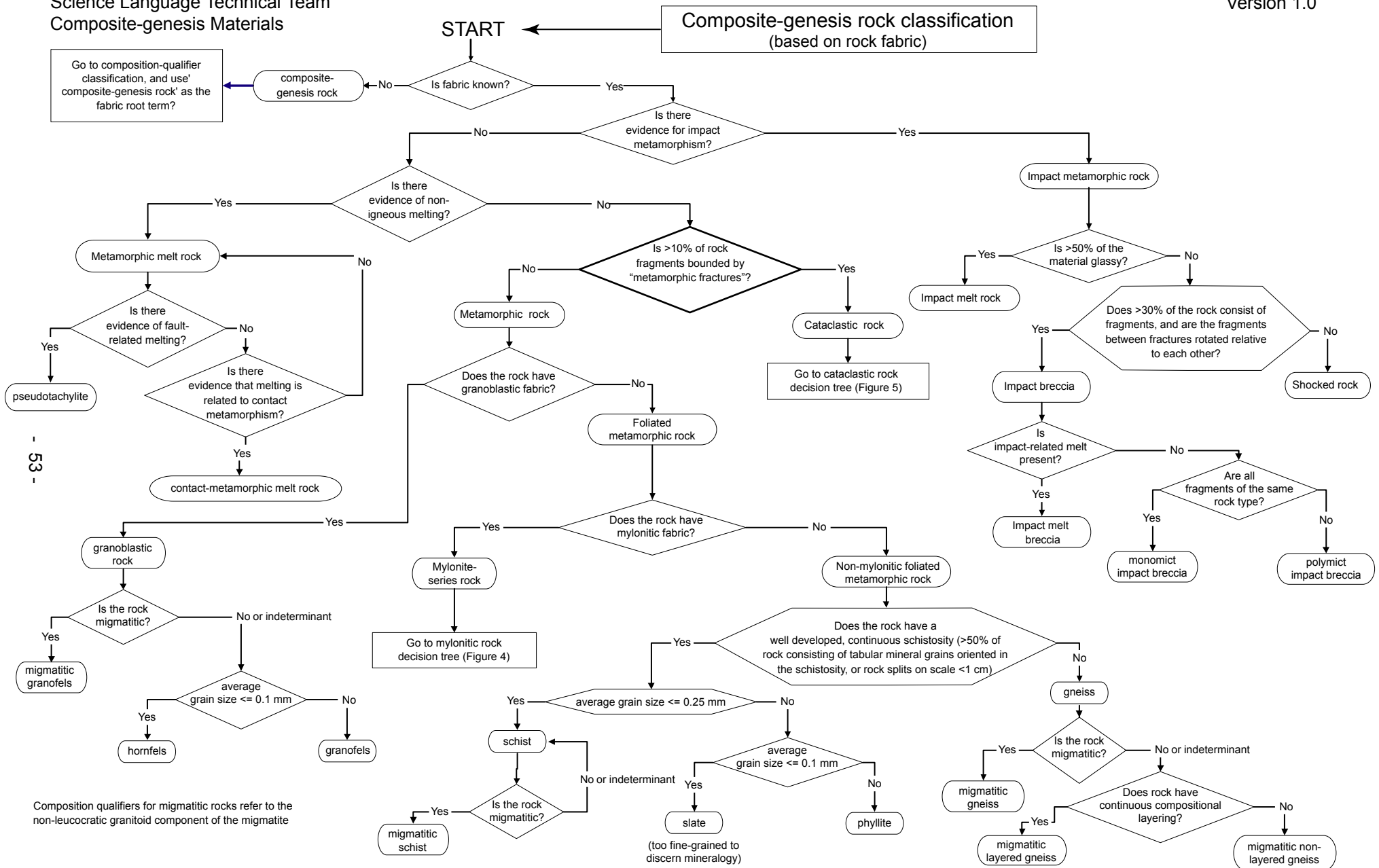


Figure 3.--Decision tree for classifying composite-genesis rocks based on fabric

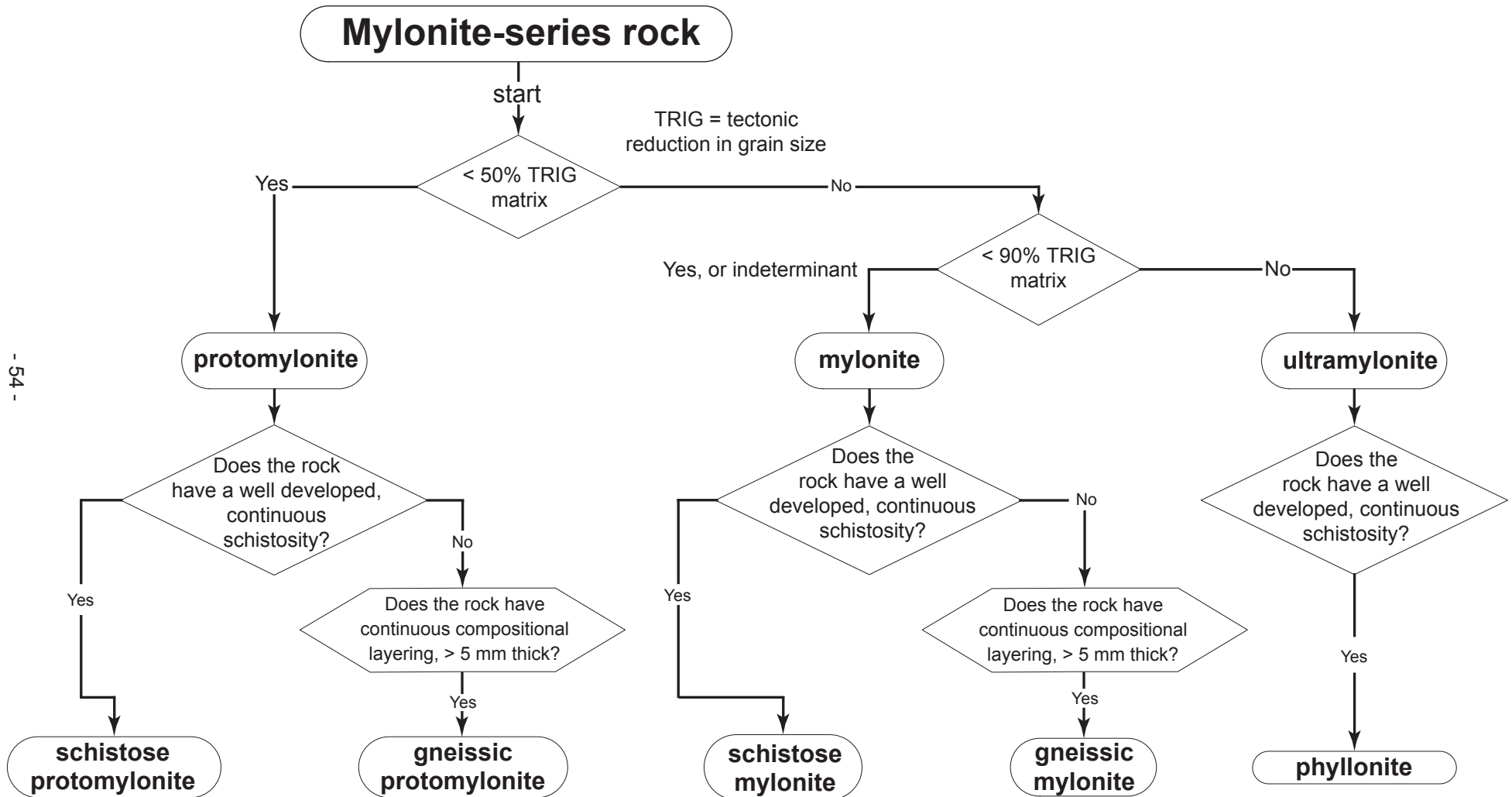


Figure 4.--Decision tree for mylonite-series rocks

Figure 5

These rocks need a special treatment for composition. The composition will in general be inherited from the protolith, which should also be classified, except in cases (gouge, ultracataclasite) where the protolith cannot be identified. Thus, compositional modifiers are only necessary for these. The fragmental rocks may have cement that is made of mineral materials not present in the protolith.

- 55 -

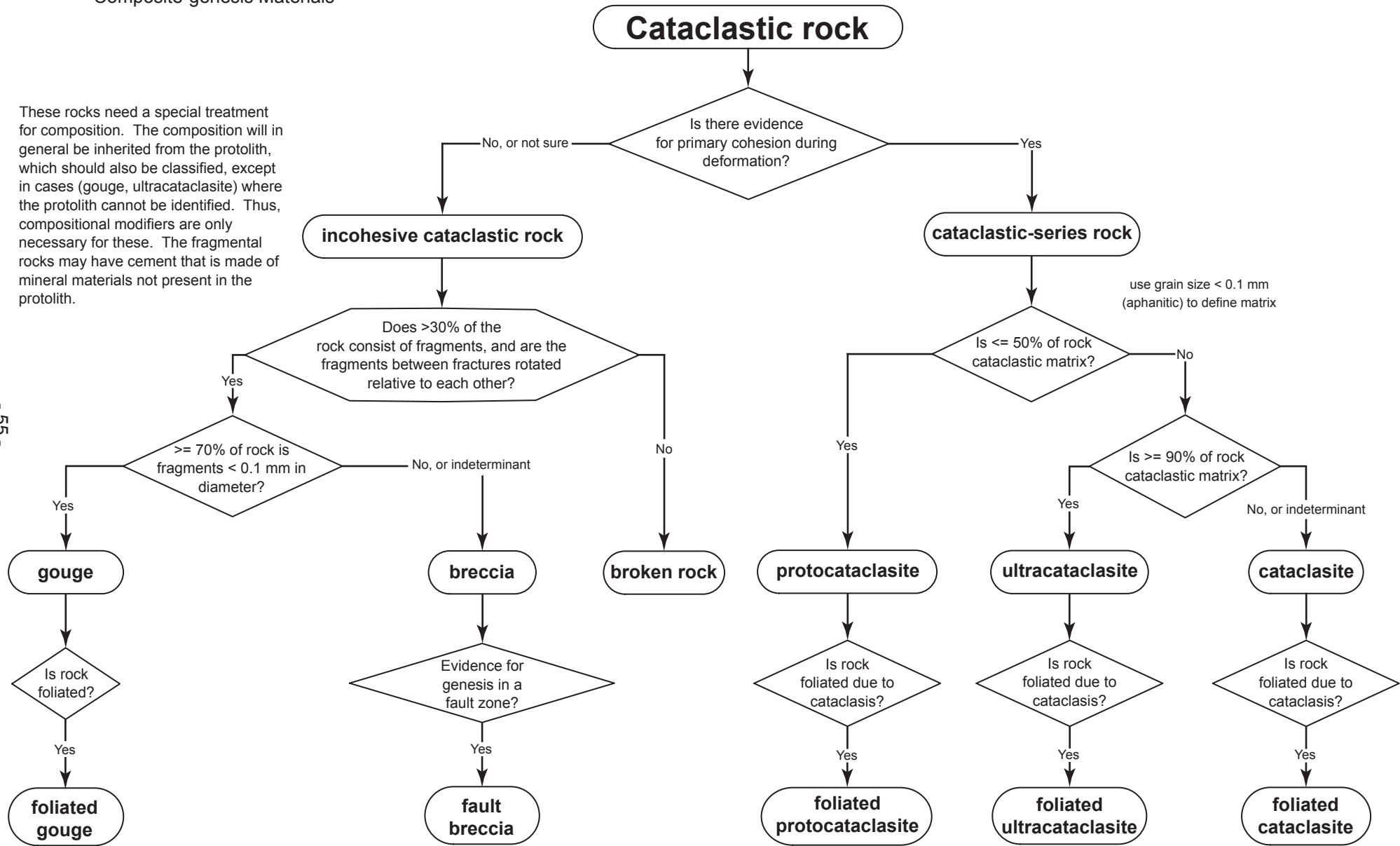


Figure 5.--Decision tree for cataclastic rocks

For rocks that have special names, if the special name denotes the rock fabric (typically granoblastic) then that rock name is used as a composition and fabric term; if a rock of that mineralogy has a different fabric from the denoted fabric, or the name does not denote a fabric, then use the special name as the composition modifier for an appropriate fabric term (e.g., marble gneiss, amphibolite schist, epidosite gneiss).

Select composition qualifier for metamorphic rock

FOR ROCKS ON THIS SIDE OF THE DECISION TREE:
 Use compound name composed of fabric root name and compositional modifier prefix, unless there is an equivalent, mineralogy-based "standard common rock name".
 If the "standard common rock name" does not have a fabric denotation, then the "standard name" is used a prefix for the appropriate textural root name (e.g., marble mylonite, amphibolite gneiss, amphibolite schist, etc.).

75% is chosen arbitrarily to quantify *predominantly* in "...consisting predominantly of fine- to coarse-grained calcite and/or dolomite..." (from the definition of marble in Jackson [1997])

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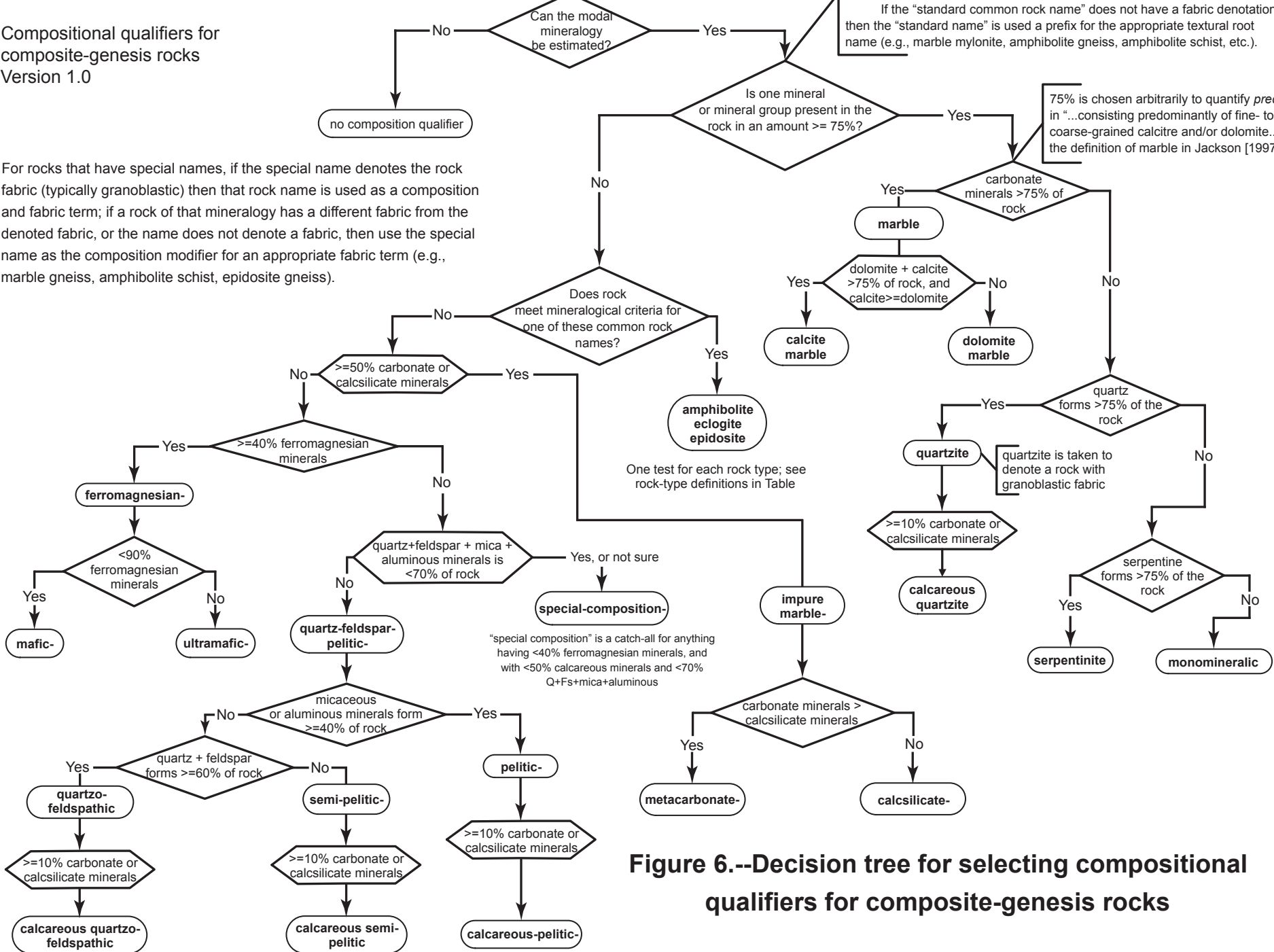


Figure 6.--Decision tree for selecting compositional qualifiers for composite-genesis rocks