

1963 **PART 6. PROPAGATION AND ANALYSIS OF UNCERTAINTY**

1964

1965 Probabilistic descriptions of what is known about some key quantities can have value in their
1966 own right as an input to research planning and in a variety of assessment activities. Often,
1967 however, analysts want to incorporate such probabilistic descriptions in subsequent modeling
1968 and other analysis. A number of closed-form analytical methods exist to perform uncertainty
1969 analysis (Morgan and Henrion, 1990). However, as computing power and speed have continued
1970 to grow, most the standard methods for the propagation of uncertainty through models, and the
1971 analysis of its implications, have come to depend on stochastic simulation.

1972

1973 Such methods are now widely used in environmental, energy and policy research, either
1974 employing standard analysis environments such as @risk[®] <www.atrisk.com>, Crystal Ball[®]
1975 <www.crystalball.com> and Analytica[®] <www.lumina.com/>, or writing special purpose
1976 software to perform such analysis.

1977

1978 While modern computer methods allow investigators to represent all model inputs as uncertain,
1979 and propagate them through the model using stochastic simulation, it is often useful to explore
1980 how much uncertainty in each input variable contributes to the overall uncertainty in the output
1981 of the model. A number of methods are now available to support such an assessment, many of
1982 which have recently been reviewed and critiqued by Borgonovo (2006).

1983

1984 Many studies have used Nordhaus' simple DICE and RICE models (Nordhaus and Boyer, 2000)
1985 to examine optimal emissions abatement policies under uncertainty. In a more recent work,

1986 Keller *et al.* (2005) has used a modified version of the RICE model to examine the implications
1987 of uncertainty about potential abrupt collapse of the North Atlantic Meridian Overturning
1988 Circulation (Gulf Stream).

1989
1990 Other groups, such as the ICAM effort (Dowlatabadi and Morgan, 1993; Morgan and
1991 Dowlatabadi, 1996; Dowlatabadi, 2000) and the MIT Joint Program²⁷, have propagated
1992 uncertainty through more complex integrated assessment models.

1993
1994 A description of the MIT Integrated Global System Model (IGSM) can be found in Sokolov *et*
1995 *al.* (2005) and on the web at <<http://web.mit.edu/globalchange/www/if.html>>. As shown in
1996 Figure 6.1 anthropogenic and natural emissions models are used to provide forcings for a
1997 coupled two-dimensional land- and ocean-resolving model of the atmosphere that is coupled to a
1998 three-dimensional ocean general circulation model. Outputs of that model are used as inputs to a
1999 terrestrial ecosystems model that predicts land vegetation changes, land CO₂ fluxes, and soil
2000 composition. These in turn feed back to the coupled chemistry/climate and natural emissions
2001 models.

2002
2003 Webster *et al.* (2003) used an earlier version of the MIT model to perform a stochastic
2004 simulation that explores the uncertainty associated with a specific policy intervention that
2005 roughly achieves stabilization at 500 ppmv. Results are shown in Figure 6.2.

2006

²⁷For a list of publications from the MIT Joint Program see <<http://web.mit.edu/globalchange/www/reports.html>>.

2007 Using this and similar models, investigators associated with the MIT Joint Center have
2008 conducted a variety of uncertainty analyses. For example, Forest *et al.* (2002, 2006) have used an
2009 optimal fingerprinting method to bound the range of values of climate sensitivity and the rate of
2010 ocean heat uptake that are consistent with their model when matched with the observed climate
2011 record of the 20th century. An example of a recent result is shown in Figure 6.3A.

2012
2013 Using a simple global energy balance model and diffusive ocean, Frame *et al.* (2005) have
2014 conducted studies to constrain possible values of climate sensitivity given plausible values of
2015 effective ocean heat capacity and observed 20th century warming. An example result is shown in
2016 Figure 6.3B. The result shown is for uniform weighting across climate sensitivity. Uniform
2017 weighting across feedbacks yields somewhat different results. The authors note that their results
2018 "fail to obtain a useful upper bound on climate sensitivity unless it is assumed *a priori*."

2019
2020 Frame *et al.* (2005) conclude that:

2021 ...if the focus is on equilibrium warming, then we cannot rule out high sensitivity, high
2022 heat uptake cases that are consistent with, but non-linearly related to, 20th century
2023 observations. On the other hand, sampling parameters to simulate a uniform distribution
2024 of transient climate response... gives an approximately uniform distribution in much
2025 more immediately policy-relevant variables ... under all SRES emission scenarios. After
2026 weighting for observations ... this approach implies a 5-95% range of uncertainty in S
2027 [the climate sensitivity] of 1.2-5.2°C, with a median of 2.3°C, suggesting traditional
2028 heuristic ranges of uncertainty in S (IPCC WGI, 2001) may have greater relevance to
2029 medium-term policy issues than recent more formal estimates based on explicit uniform
2030 prior distributions in either S or [feedback strength] λ .

2031
2032 Murphy *et al.* (2004) have completed extensive parametric analysis with the HadAM3
2033 atmospheric model coupled to a mixed layer ocean that they report "allows integration to
2034 equilibrium in a few decades." They selected a subset of 29 of the roughly 100 parameters in this

2035 model, which they judged to be most important in determining the model's climate sensitivity,
2036 and then perturbed them one at a time with respect to their standard values, and created 53
2037 different model versions, each of which was used to simulate present and future 2xCO₂ climate.

2038
2039 Placing uniform probability distributions on all these, they conclude that the implied climate
2040 sensitivity has a "median value of 2.9°C with a spread (corresponding to a 5 to 95% probability
2041 range) of 1.9 to 5.3°C." By using some analysis and expert judgment to shape the prior
2042 distributions, they also produce a "likelihood-weighted" distribution which they report "results in
2043 a narrowing of the 5 to 95% probability range to 2.4 to 5.4°C, while the median value increases
2044 to 3.5°C" (Murphy *et al.*, 2004). They report:

2045 Our probability function is constrained by objective estimates of the relative reliability of
2046 different model versions, the choice of model parameters that are varied and their
2047 uncertainty ranges, specified on the basis of expert advice. Our ensemble produces a
2048 range of regional changes much wider than indicated by traditional methods based on
2049 scaling the response patterns of an individual simulation.
2050

2051 One of the most exciting recent developments in exploring the role of uncertainty in climate
2052 modeling has been the use of a large network of personal computers, which run a version of the
2053 HadSM3 model as a background program when machine owners are not making other uses of
2054 their machine. This effort has been spearhead by Myles Allen and colleagues at Oxford (Allen,
2055 1999). Details can be found at <<http://www.climateprediction.net/index.php>>. As of mid-spring
2056 2006, this network involved over 47 thousand participating machines that had completed over
2057 150 thousand runs of a version of the HadSM3 model, for a total of 11.4 million model years of
2058 simulations.

2059

2060 Initial results from this work were reported by Stainforth *et al.* (2005) who summarize their
2061 findings from a study of 2,578 simulations of the model as follows:

2062 We find model versions as realistic as other state-of-the-art climate models but with
2063 climate sensitivities ranging from less than 2K to more than 11K. Models with such
2064 extreme sensitivities are critical for the study of the full range of possible responses of the
2065 climate system to rising greenhouse gas levels, and for assessing the risks associated with
2066 a specific target for stabilizing these levels...
2067

2068 The range of sensitivity across different versions of the same model is more than twice
2069 that found in the GCMs used in the IPCC Third Assessment Report...The possibility of
2070 such high sensitivities has been reported by studies using observations to constrain this
2071 quantity, but this is the first time that GCMs have generated such behavior. (Stainforth *et*
2072 *al.*, 2005)
2073

2074 The frequency distribution in climate sensitivity they report across all model versions is shown in
2075 Figure 6.4.

2076

2077 While the common practice in many problem domains is to build predictive models, or perform
2078 various forms of policy optimization, it is important to ask whether meaningful prediction is
2079 possible. At least in the context of predicting the future evolution of the energy system, which is
2080 responsible for a large fraction of anthropogenic greenhouse gas emissions, Smil (2003) and
2081 Craig *et al.* (2002) have very clearly shown that accurate prediction for more than a few years in
2082 the future, is virtually impossible. Figure 6.5 redrawn from Smil, shows the sorry history of past
2083 forecasts for United States energy consumption. His summary of forecasts of global energy
2084 consumption shows similarly poor performance.

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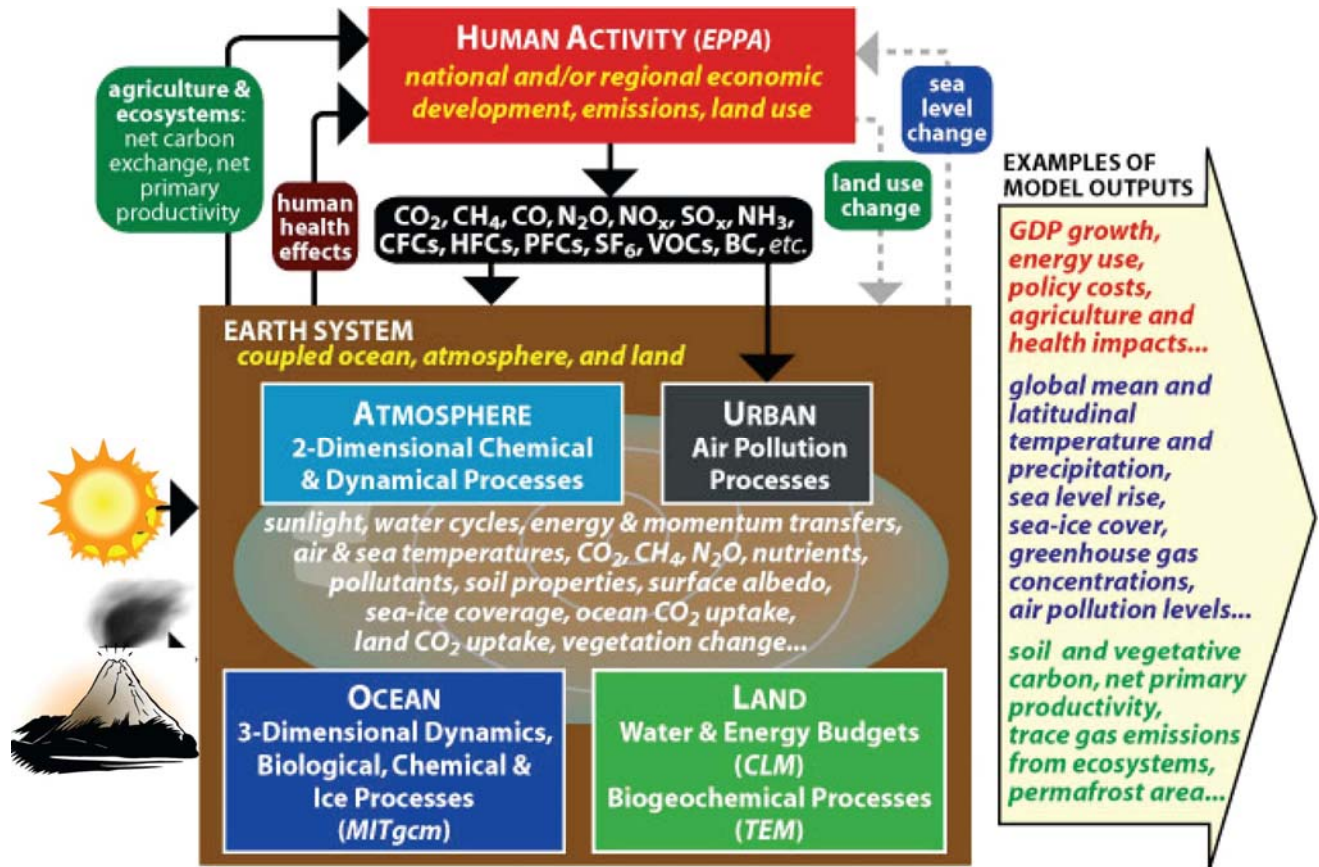
2086 In addition to uncertainties about the long-term evolution of the energy system and hence future
2087 emissions, uncertainties about the likely response of the climate system, and about the possible

2088 impacts of climate change, are so great that a full characterization of coefficient and model
2089 uncertainty in a simulation model can lead to probabilistic results that are so broad that they are
2090 effectively useless (Casman *et al.*, 1999). Similarly, if one does parametric analysis across
2091 different model formulations, one can obtain an enormous range of answers depending on the
2092 model form and other inputs that are chosen. This suggests that there are decided limits to the use
2093 of "predictive models", and "optimization" in many climate assessment and policy settings.

2094
2095 The difficulties, or sometimes even impossibility, of performing meaningful predictive analysis
2096 under conditions of what has been called "deep" or "irreducible" uncertainty have led some
2097 investigators to pursue a different approach based on two key ideas: describing uncertainty about
2098 the system relevant to a decision with multiple representations, as opposed to a single best-
2099 estimate joint probability distribution, and using a robustness, as opposed to an optimality, as the
2100 criteria for evaluating alternative policy options. We turn to a more detailed discussion of these
2101 approaches in the latter parts of the next section.

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2106 **Figure 6.1** Simplified block diagram of the MIT Integrated Global System Model (IGSM) Version 2. Source: MIT
 2107 Global Change Joint Program. Reprinted with permission.

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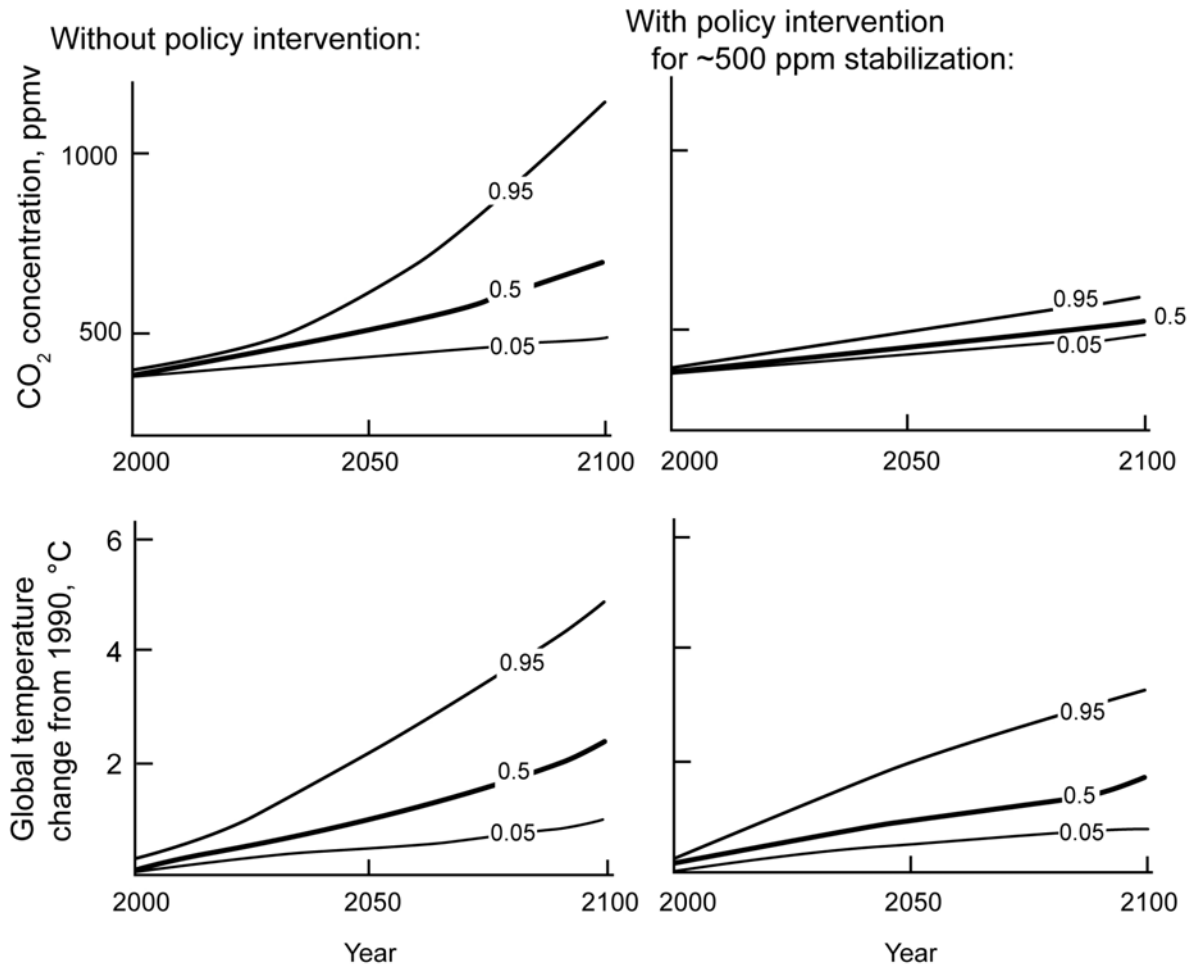
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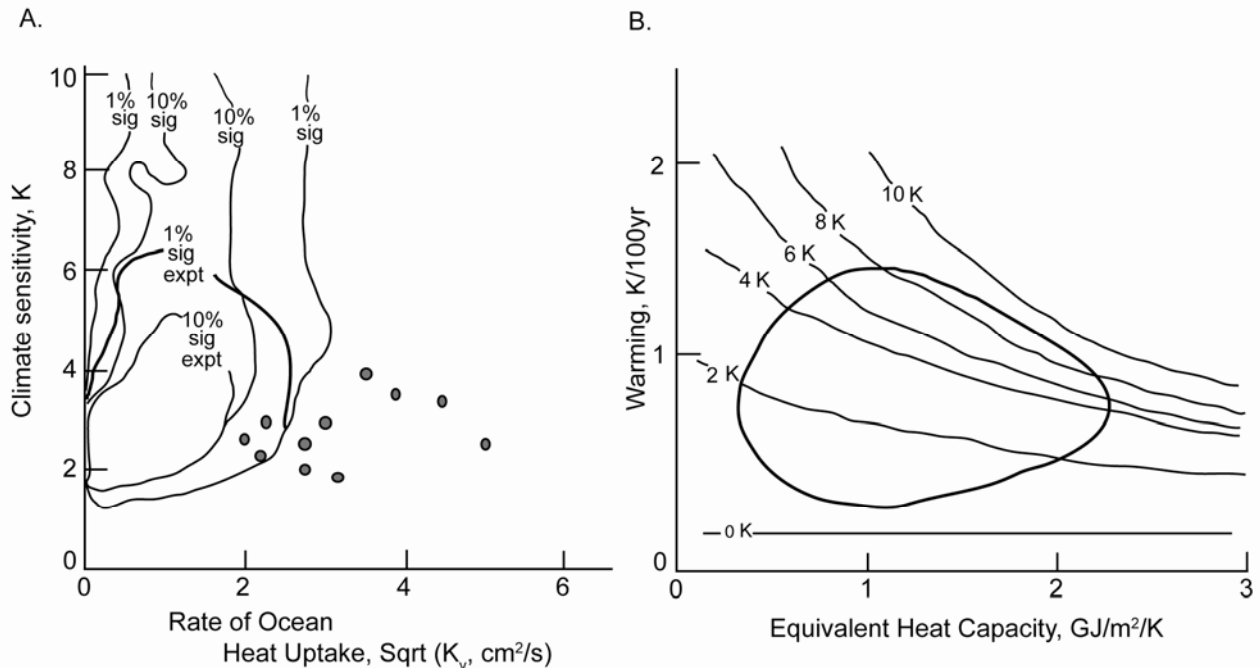
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2116 **Figure 6.2** Results of simulation conducted by Webster *et al.* (2003) which use an earlier version of the MIT IGSM
 2117 model with probability distributions on model inputs that are constrained by past performance of the climate system.
 2118 Results on the left are the authors' projection for no policy intervention and on the right for a specific policy
 2119 intervention that roughly achieves stabilization at 500 ppmv. Heavy curves show median results from the
 2120 simulations. Light curves show 0.05 and 0.95 confidence intervals. [Redrawn from Webster *et al.* (2003).]



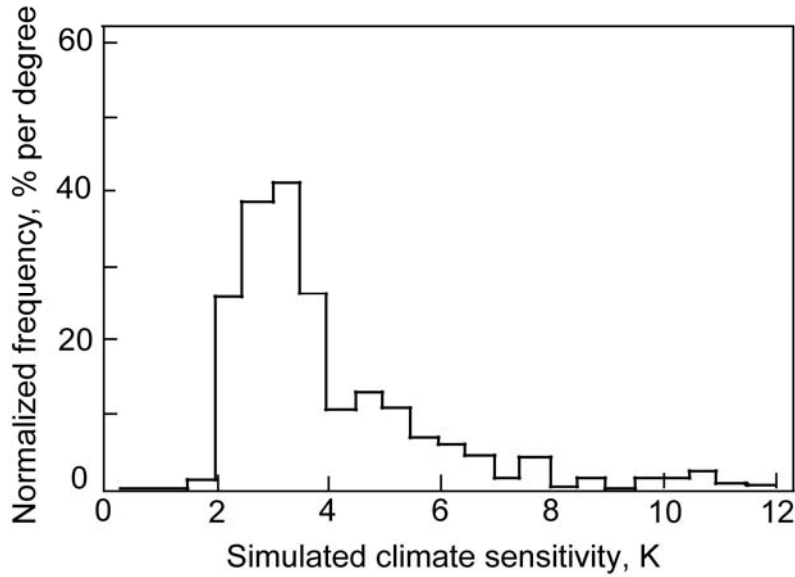
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2123 **Figure 6.3** Two examples of recent efforts to bound sensitivity and heat uptake or heat capacity by combining
 2124 expert judgment and model simulations.
 2125 **A.** (redrawn from Forest *et al.*, 2006) shows the marginal posterior probability density function obtained when using
 2126 uniform probability distributions across all relevant forcings and matching outputs from the ocean and atmospheric
 2127 portion of the MIT IGSM model. Light contours bound the 10% and 1% significance regions. Similarly, the two
 2128 dark contours are for an expert PDF on climate sensitivity. Dots show outputs from a range of leading GCMs all of
 2129 which lie to the right of the high-probability region, suggesting that if Forest *et al.* (2006) are correct, these models
 2130 may be mixing heat into the deep ocean too efficiently.
 2131 **B** (redrawn from Frame *et al.*, 2005) shows the relationship between climate sensitivity, shown as light contours,
 2132 effective ocean heat capacity, and 20th century warming for the case of uniform sampling of climate sensitivity (not
 2133 shown are similar results for uniform sampling across feedback strength). The dark contour shows the region
 2134 consistent with observations at the 5% level. Note: We have roughly extrapolated the climate sensitivity contours
 2135 from colored points in the original diagram that report each of many of hundreds of individual model runs. In this
 2136 diagram, they are only qualitatively correct.
 2137 Note that neither of these analyses account for the issue of structural uncertainty.
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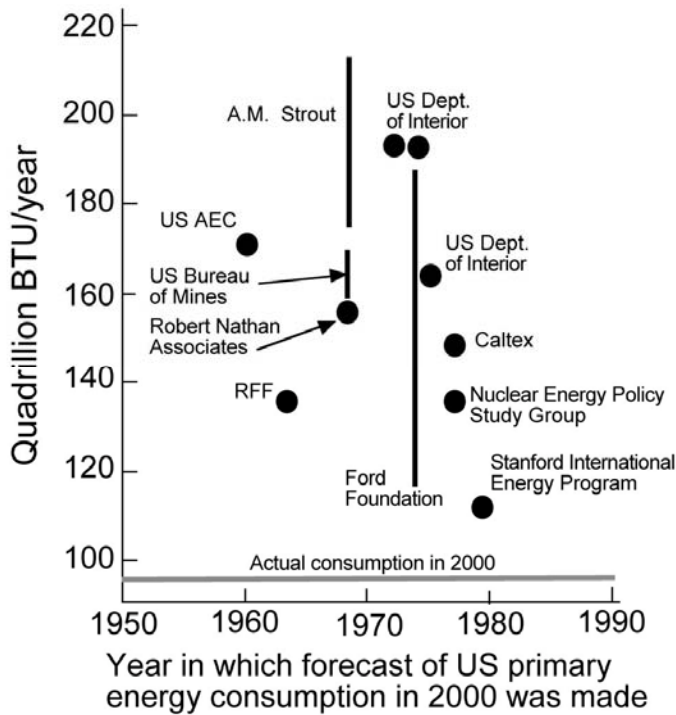


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2143 **Figure 6.4** Histogram (redrawn) of climate sensitivities found by Stainforth *et al.* (2005) in their simulation of
 2144 2,578 versions of the HadSM3 GCM model.

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2148 **Figure 6.5** Summary of forecasts of United States primary energy consumption compiled by Smil (2003) as a
 2149 function of the date on which they were made. [Figure redrawn from Smil (2003).]

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