

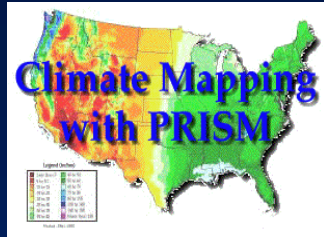
Using Synoptic Climatology and the PRISM Model to Improve Precipitation Assessment and Prediction

George H. Taylor
State Climatologist, Oregon
Director, Oregon Climate Service

Christopher Daly
Director, Spatial Climate Analysis Service

Oregon State University
Corvallis, Oregon, USA

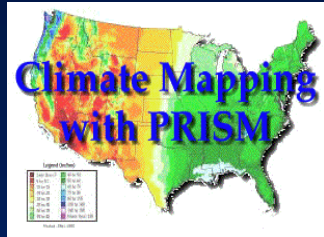




PRISM

Parameter-elevation Regressions on Independent Slopes Model

- Generates gridded estimates of climatic parameters
- Moving-window regression of climate vs. elevation for each grid cell
 - Uses nearby station observations
- Spatial climate knowledge base (KBS) **weights stations** in the regression function by their climatological similarity to the target grid cell



PRISM

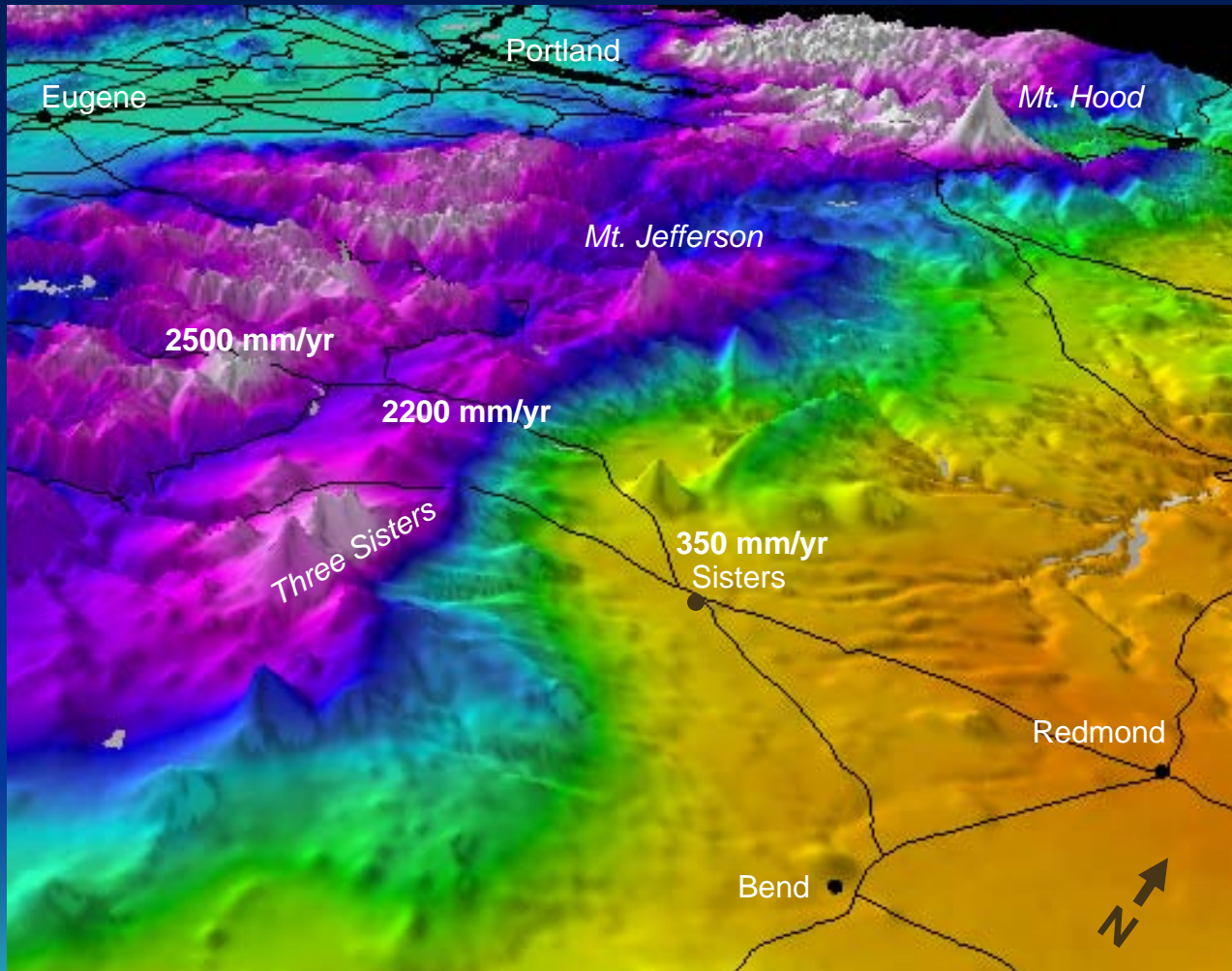
Parameter-elevation Regressions on Independent Slopes Model

PRISM Knowledge Base accounts for spatial variations in climate due to:

- Distance – closer is better
- Elevation – similar elevation is better
- Terrain orientation – same side of mountain is better
- Terrain steepness - same slope steepness is better
- Coastal proximity – similar exposure to coastal effects is better
- Inversion layer – same side (above or below) of inversion is better



Rain Shadows: 1961-90 Mean Annual Precipitation Oregon Cascades



**Dominant
PRISM KBS
Components**

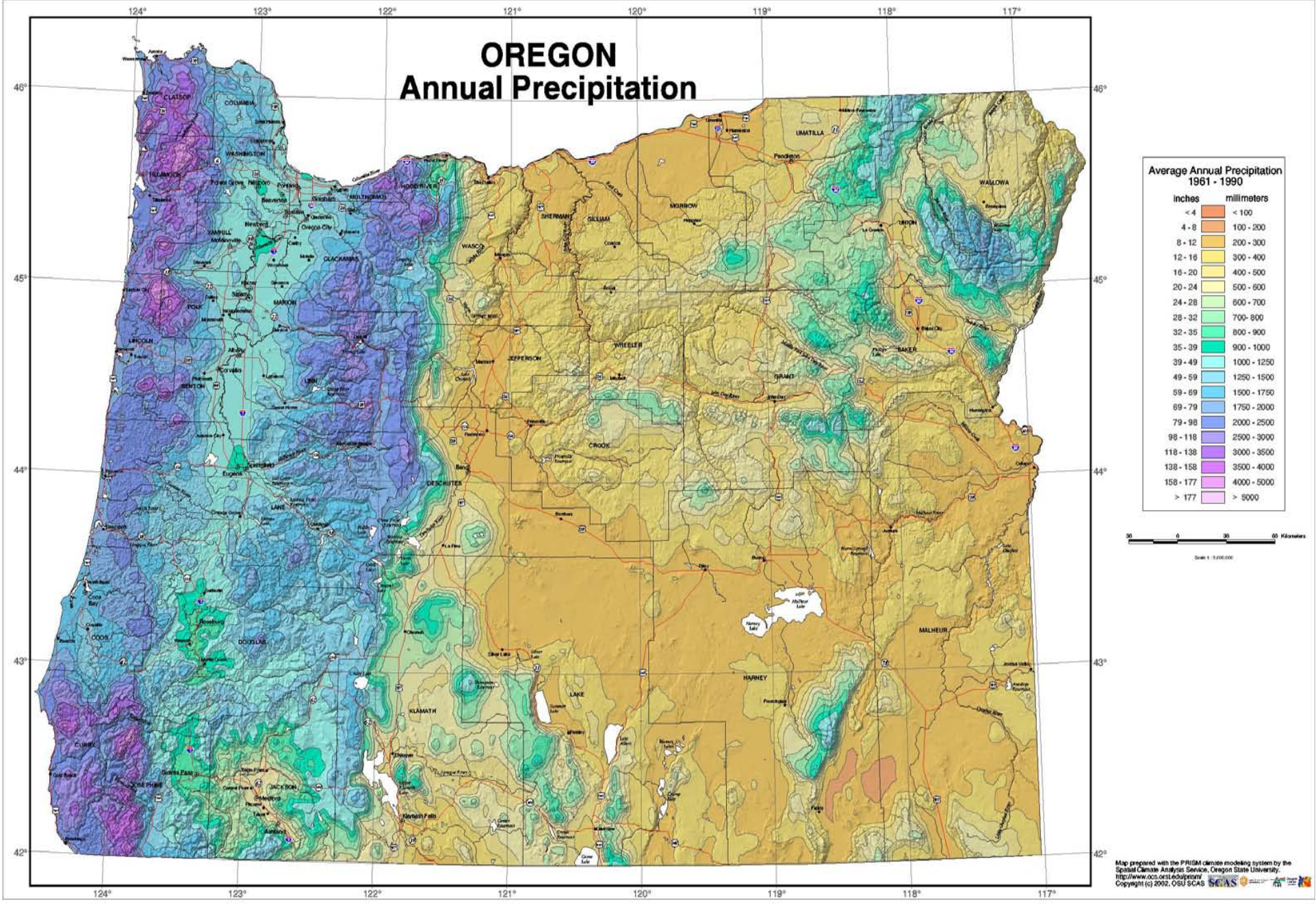
Elevation

Terrain orientation

Terrain steepness



OREGON Annual Precipitation



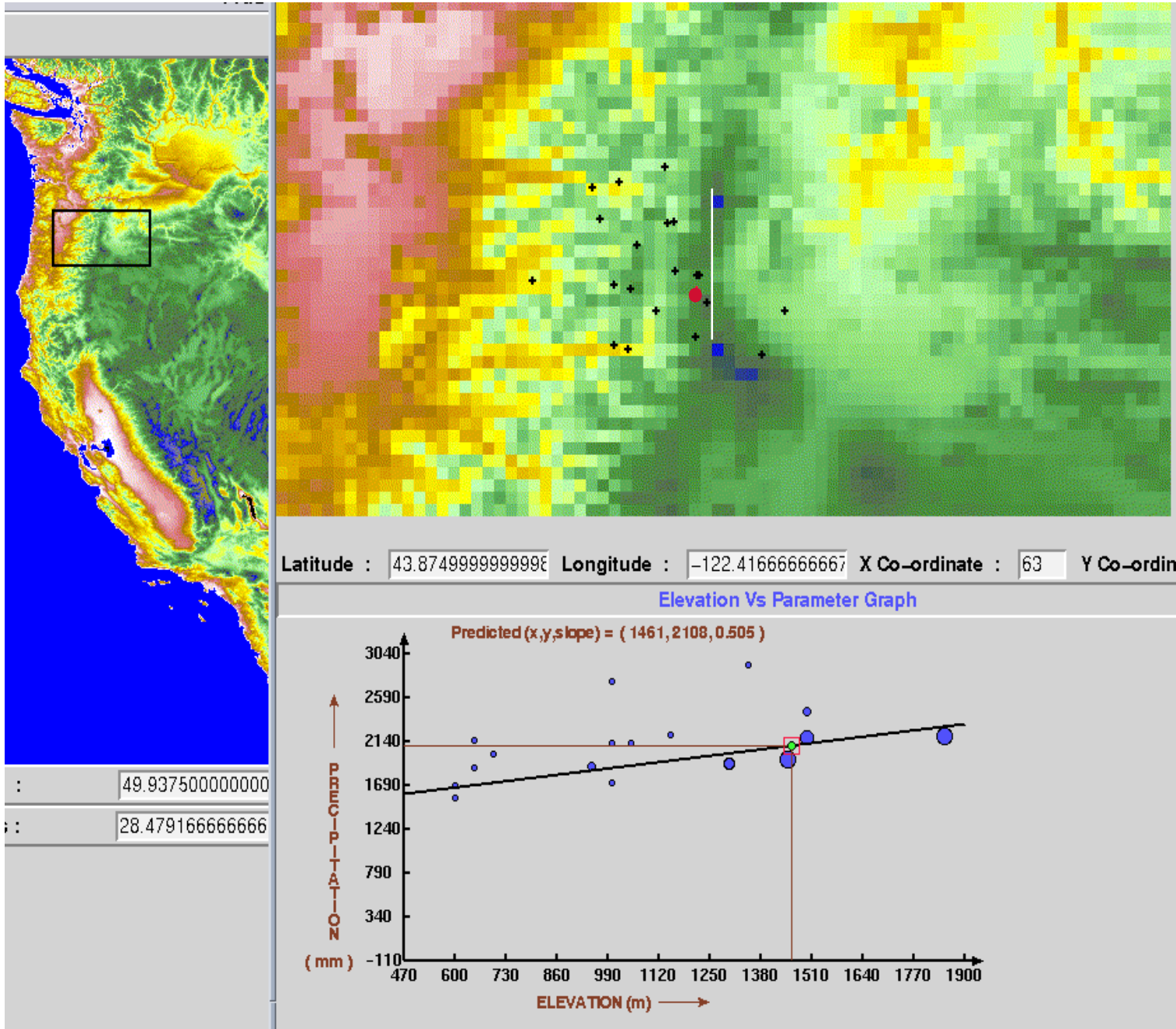
Average Annual Precipitation 1961 - 1990

Inches	millimeters
< 4	< 100
4 - 8	100 - 200
8 - 12	200 - 300
12 - 16	300 - 400
16 - 20	400 - 500
20 - 24	500 - 600
24 - 28	600 - 700
28 - 32	700 - 800
32 - 35	800 - 900
35 - 39	900 - 1000
39 - 49	1000 - 1250
49 - 59	1250 - 1500
59 - 69	1500 - 1750
69 - 79	1750 - 2000
79 - 98	2000 - 2500
98 - 118	2500 - 3000
118 - 138	3000 - 3500
138 - 158	3500 - 4000
158 - 177	4000 - 5000
> 177	> 5000

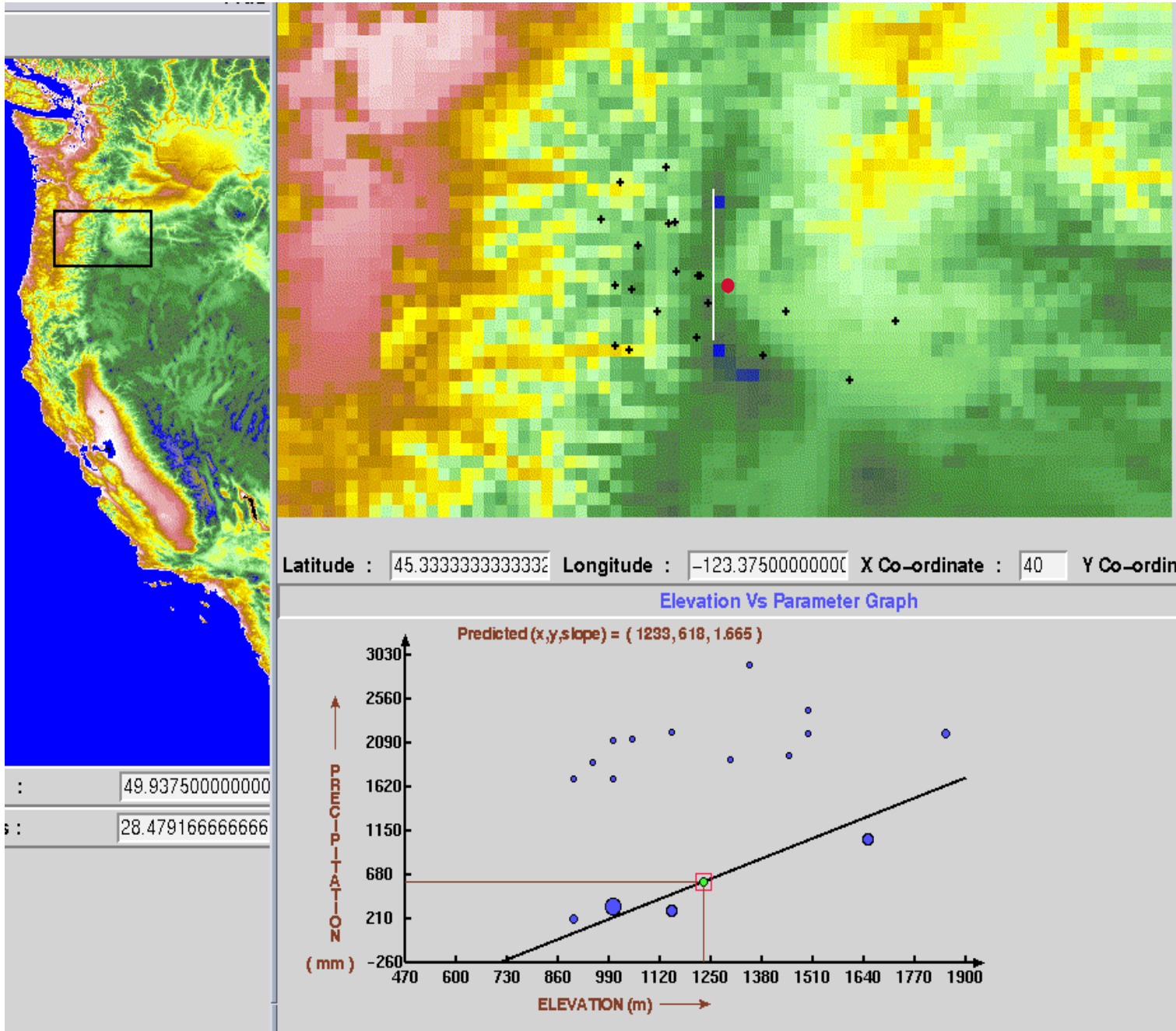


Map prepared with the PRISM climate modeling system by the Spatial Climate Analysis Service, Oregon State University. <http://www.ccl.orst.edu/osu/prism/> Copyright (c) 2002, OSU SCAS

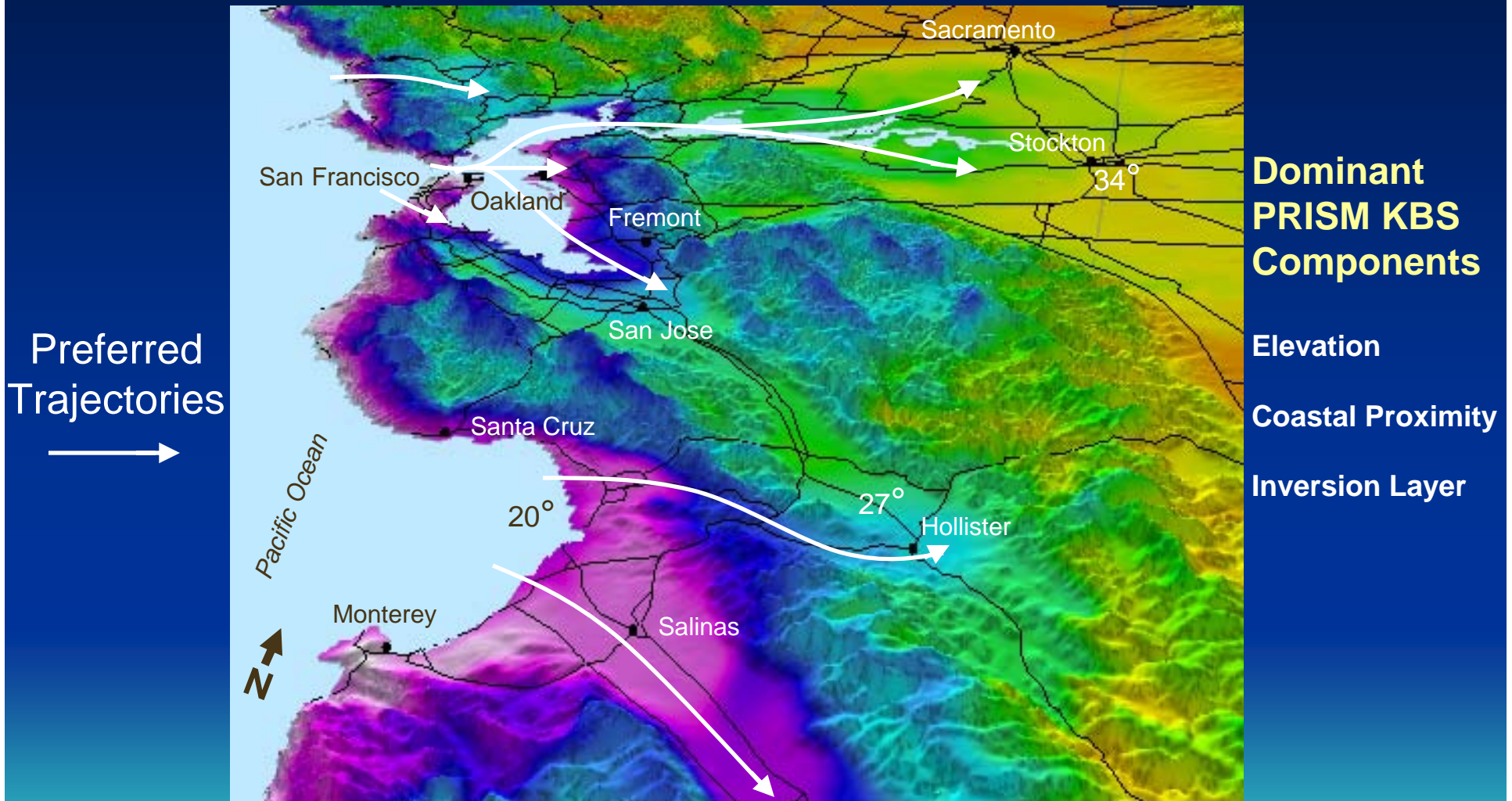
1961-90 Mean Annual Precipitation, Cascade Mtns, OR, USA



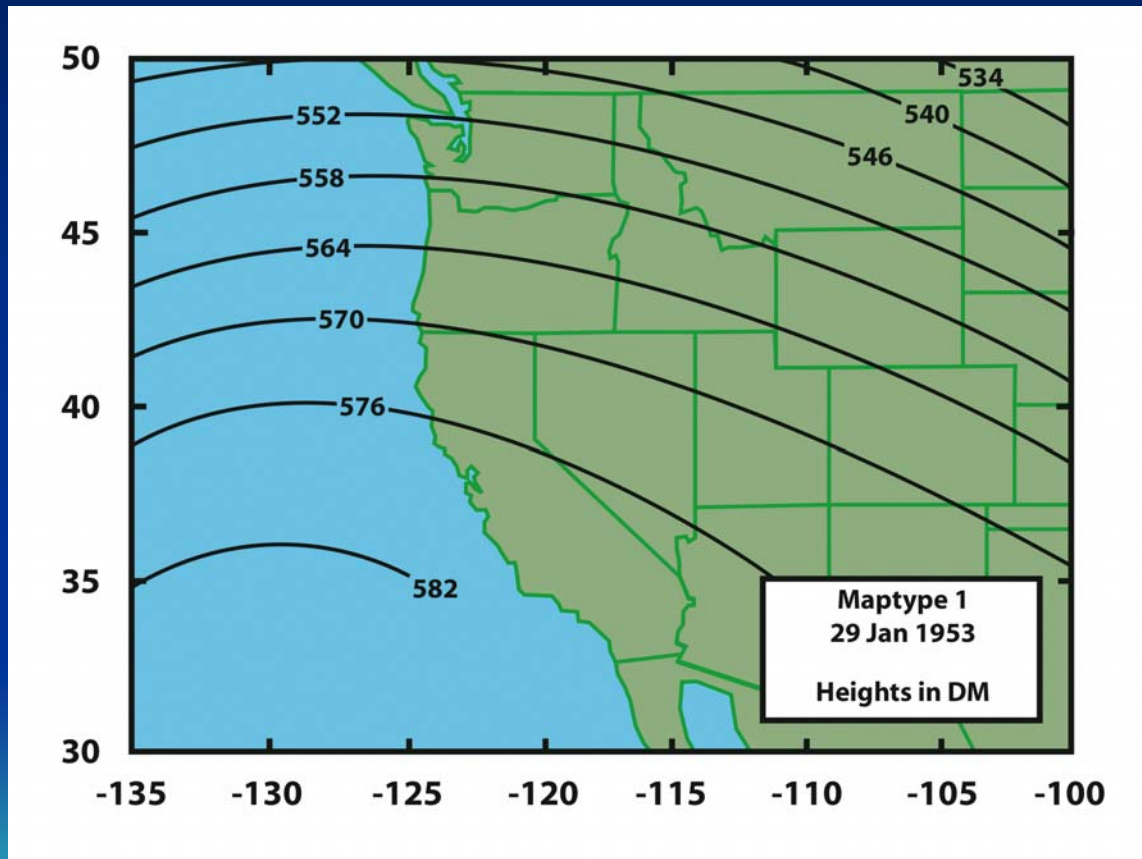
1961-90 Mean Annual Precipitation, Cascade Mtns, OR, USA



Coastal Effects: 1971-00 July Maximum Temperature Central California Coast

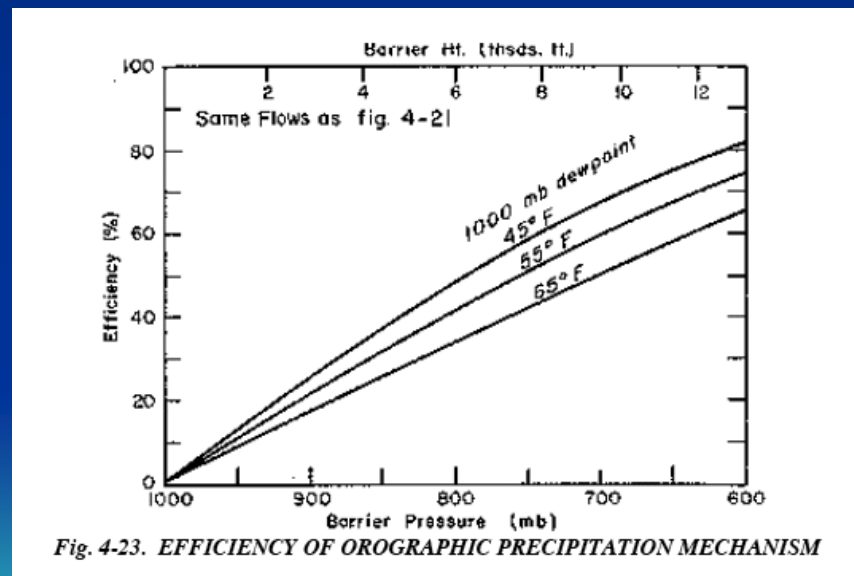


Map-type Climatology



Weaver, 1962 (HMR-37)

METEOROLOGY OF HYDROLOGICALLY CRITICAL STORMS IN CALIFORNIA



Weaver “Map Types”

Low-latitude type. Blocking in, mid-Pacific between longitude 160W and 180W, or about the longitude of the eastern Aleutians. This is primarily a type for precipitation in the northern half of California.

High-latitude type. Blocking in the eastern Pacific east of longitude 160W. This is mainly a storm type for the southern half of California.

Mid-latitude type. Low pressure in the central and eastern Pacific, with varying degrees of blocking over western North America.



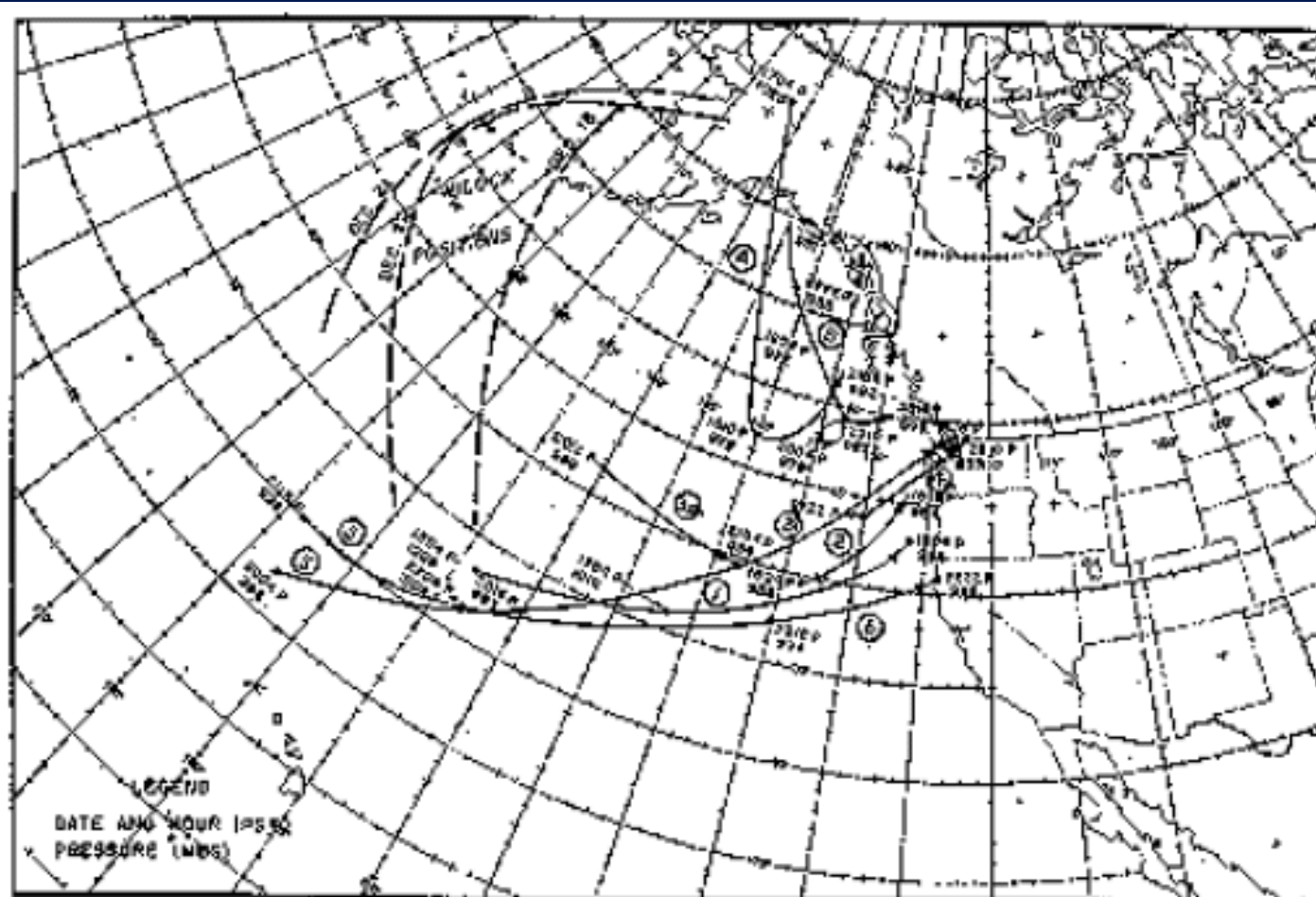
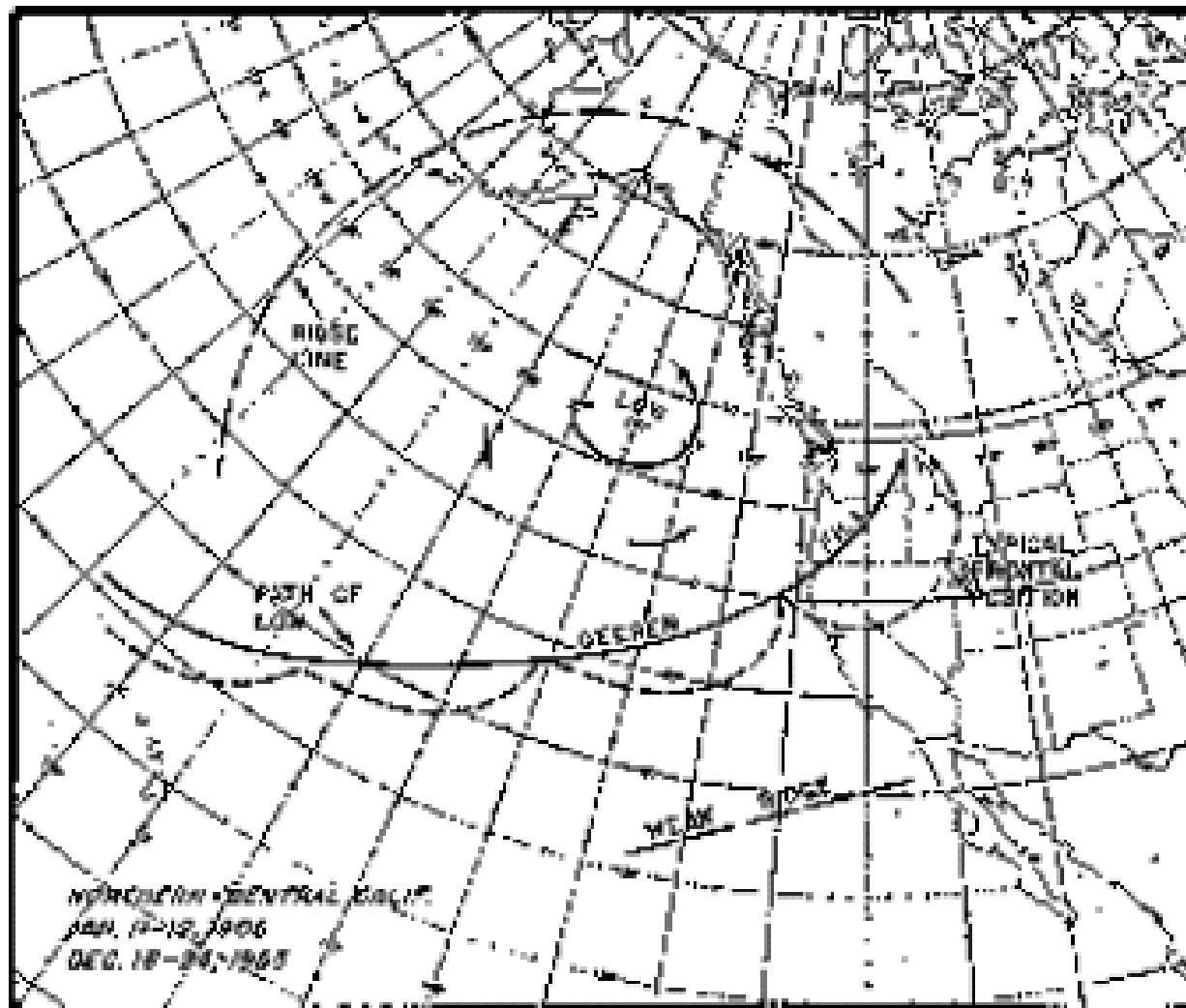


Fig. 2-12. TRACKS OF SURFACE LOWS IN THE DECEMBER 1955 STORM



OSU Map-type Procedure

- a. Use the gridded reanalysis data, which begins in 1946. Most of the period has twice-daily maps. Large-scale patterns are probably best captured using the 700 or 500 mb grids.
- b. Compute the correlations between individual maps of pressure or height; comparisons are restricted to certain ranges of dates (in the current system, calendar days up to 30 days before or after the given date were used for comparison).
- c. Select an “acceptable” correlation coefficient representing the minimum value for establishing similarity between maps. In this case, 0.80 was used.



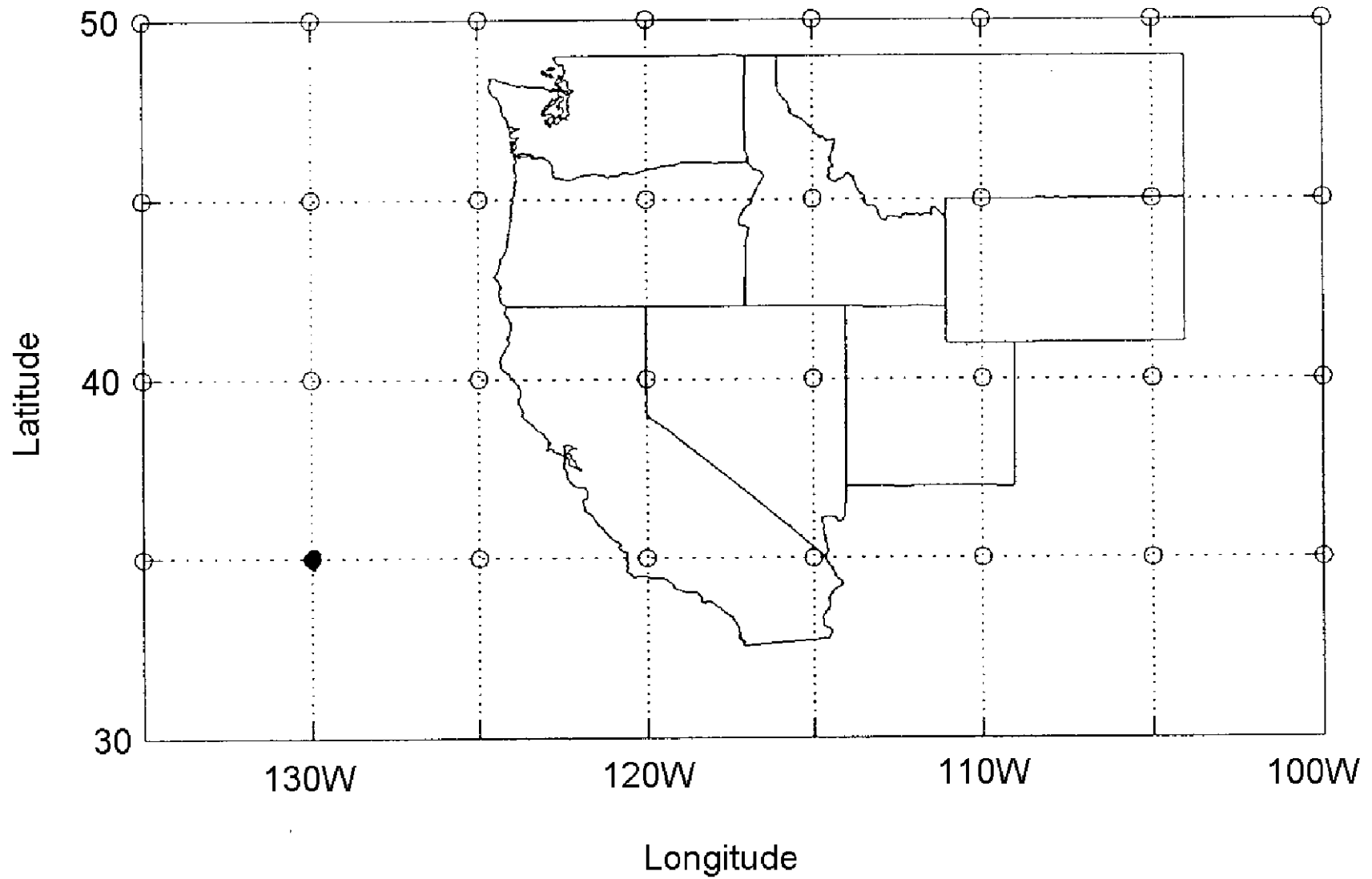
d. Find the map with the largest number of 0.80 correlations with other maps. This will be identified as “Type 1,” and the map selected called the “Key Day” map.

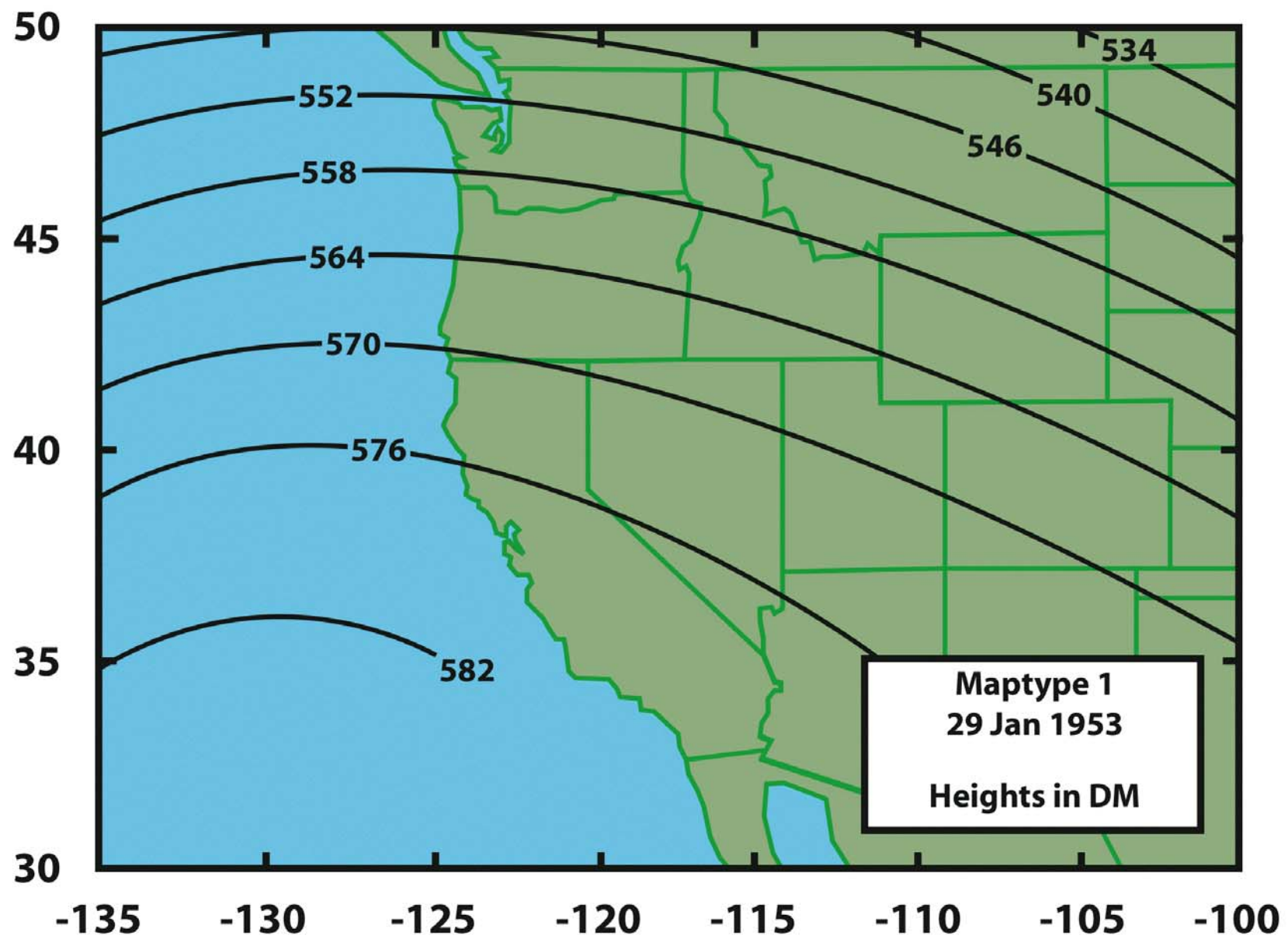
e. Remove the Type 1 maps from the data set and repeat steps c. and d. until some maximum number of types have been identified (typically 12-20 types).

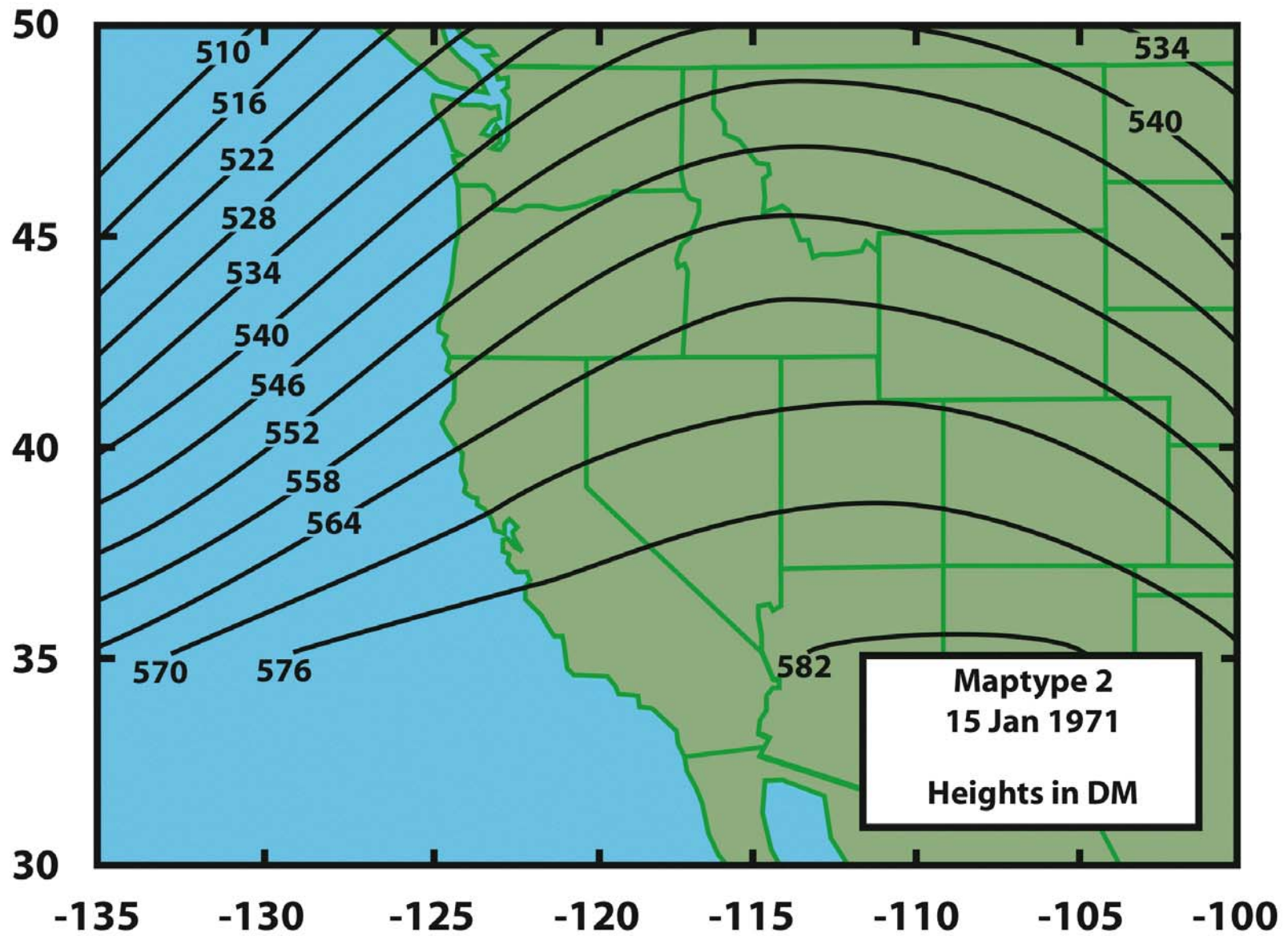
f. Compare the selected maps with the Key Day map for later map types; if the correlation coefficient with another map type is higher, move the selected map to the other map type group. For example, if a map is selected for Type 1, it will be removed from the data set prior to creation of the Type 2, Type 3, etc. data sets. If it fits better with Type 3, it should be moved.

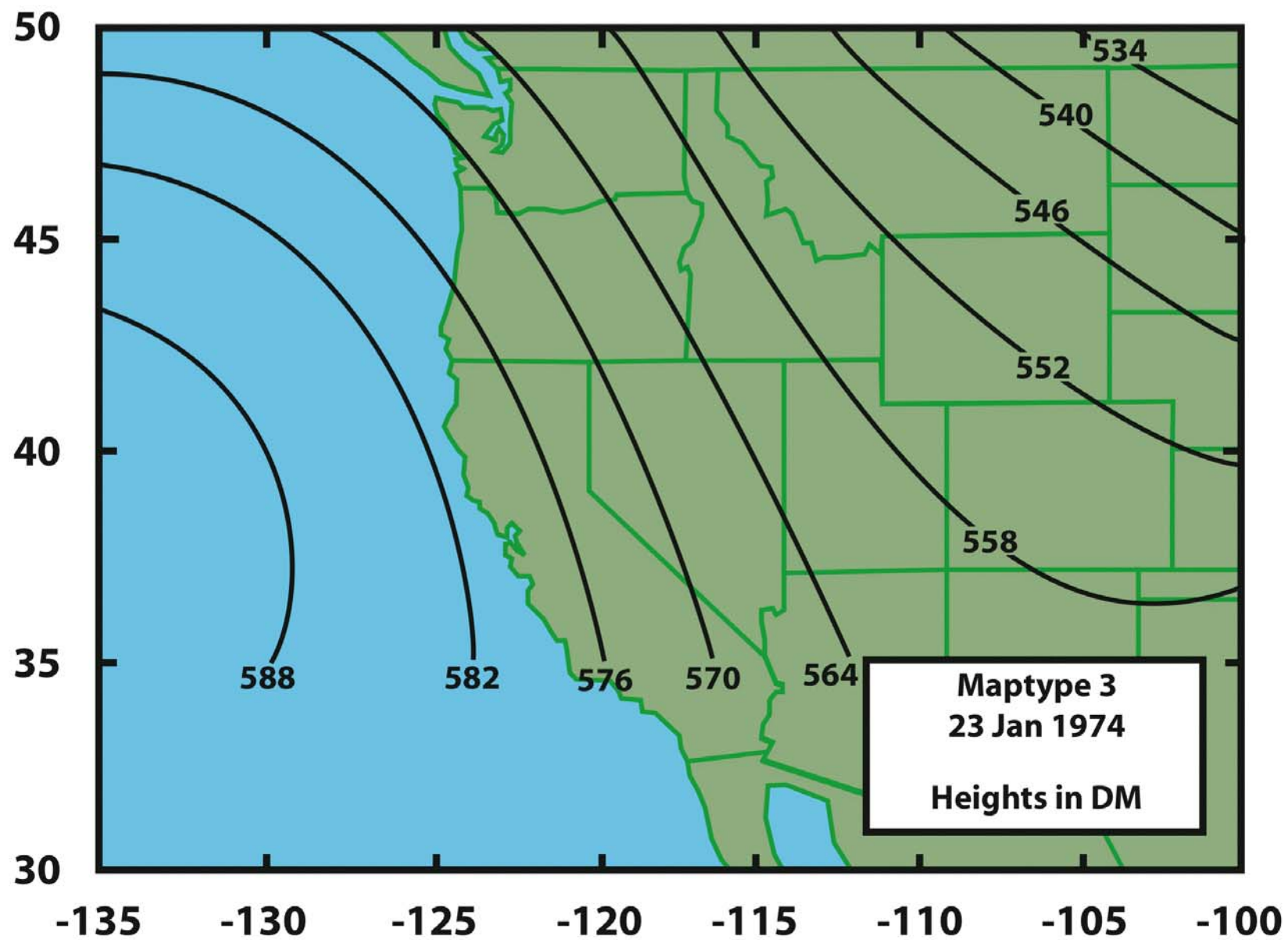


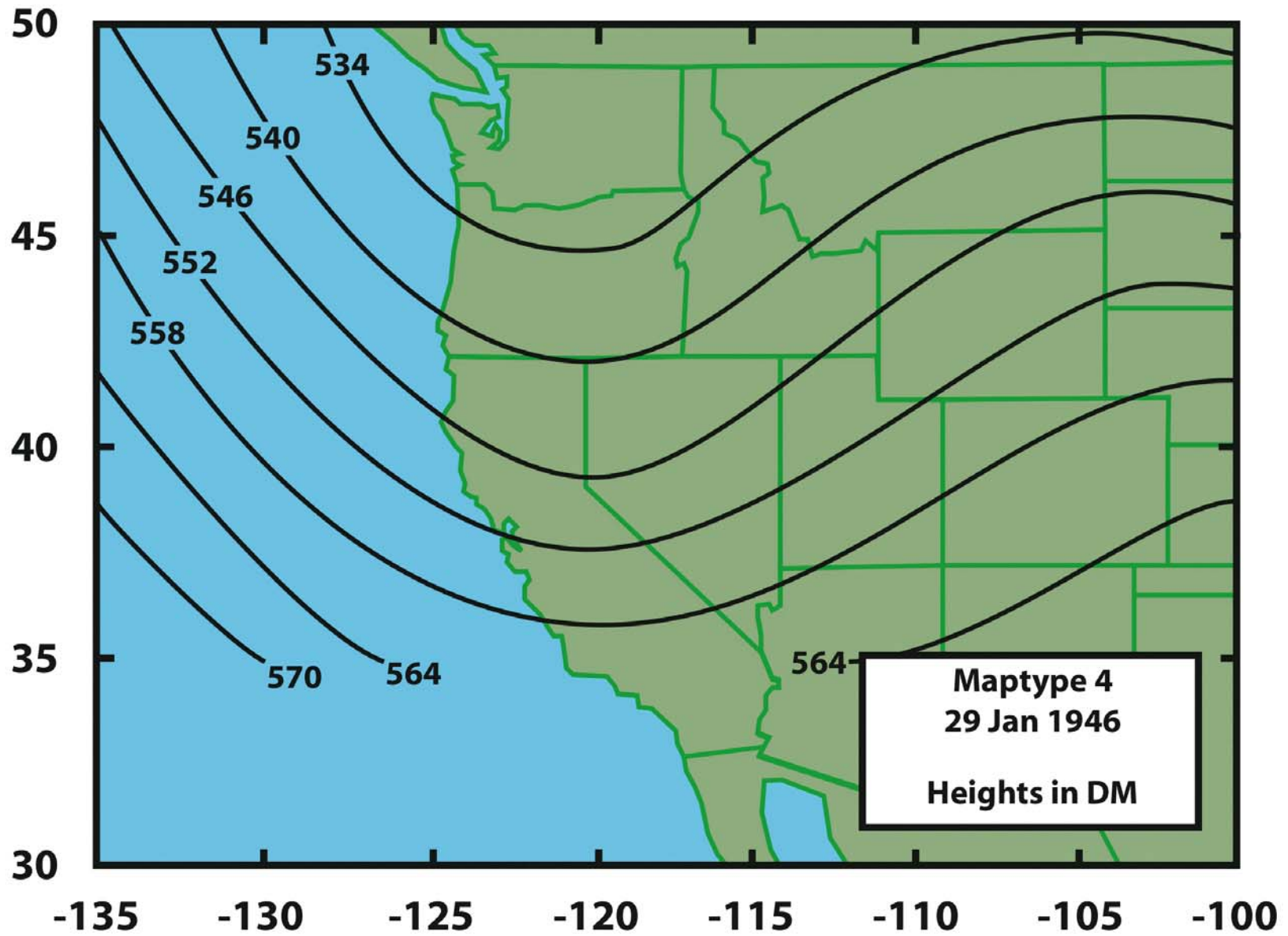
Location of Data Points







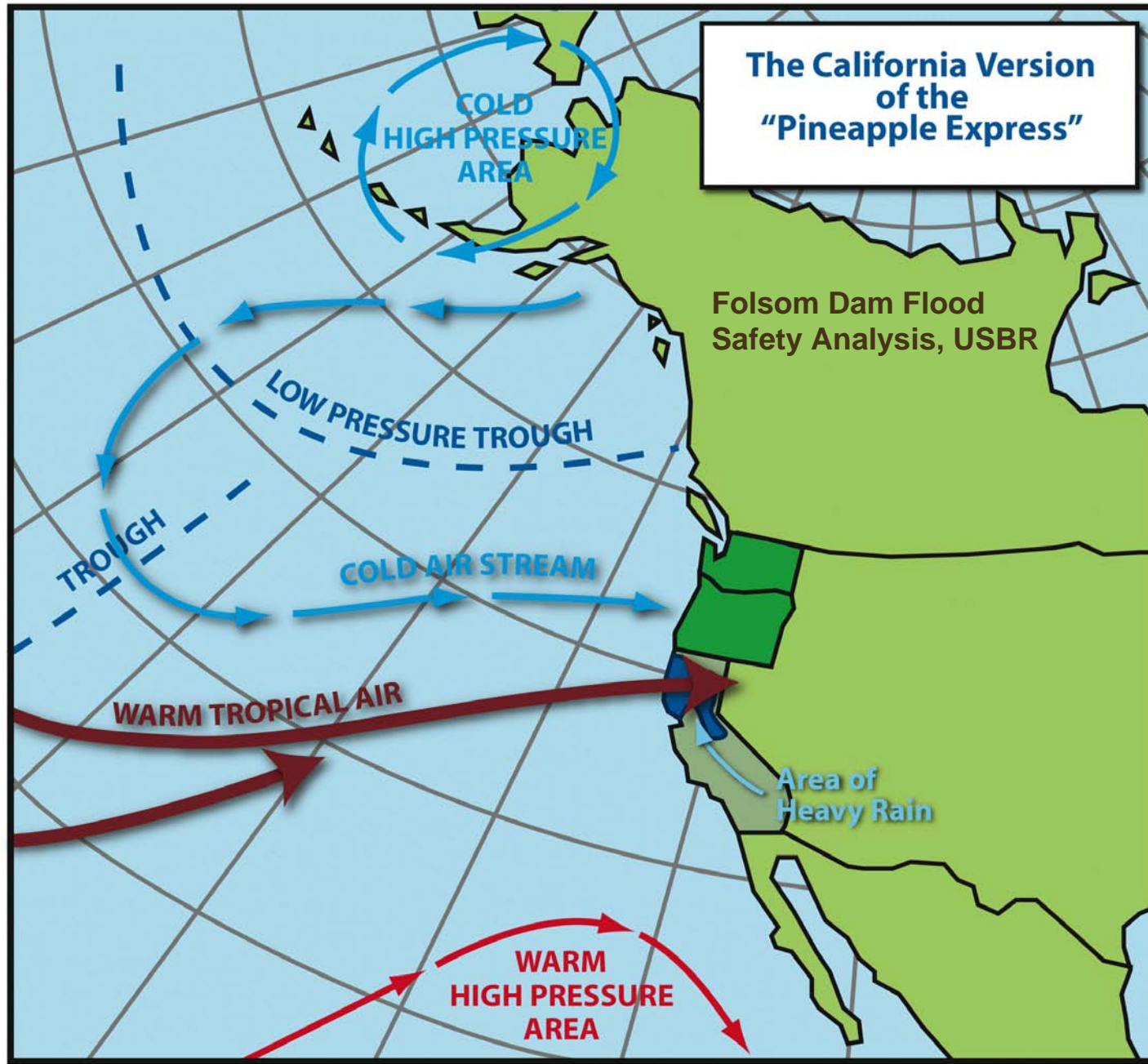




Comparing Extreme Storms in The Pacific Northwest and California



**The California Version
of the
"Pineapple Express"**



Folsom Dam Flood
Safety Analysis, USBR

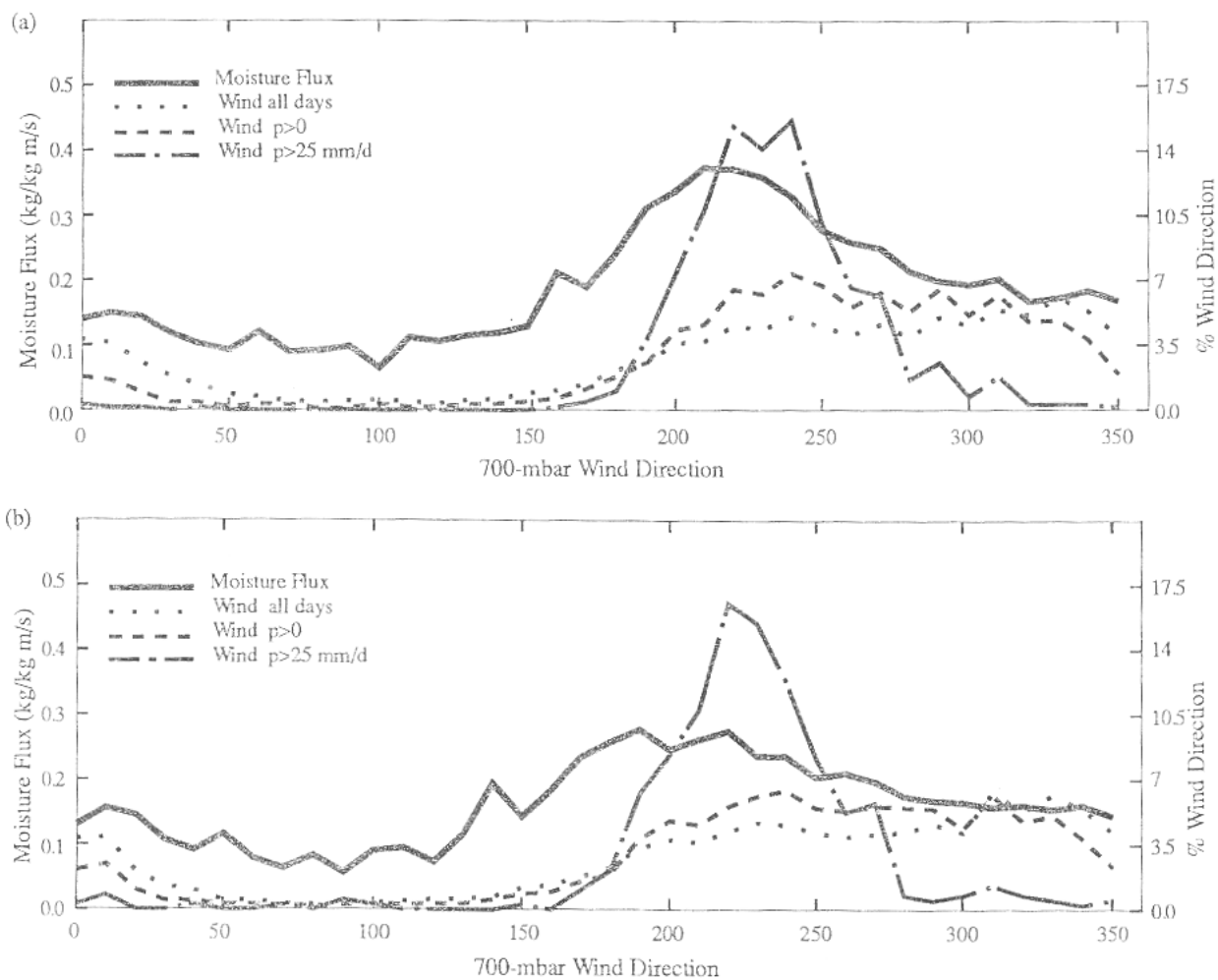
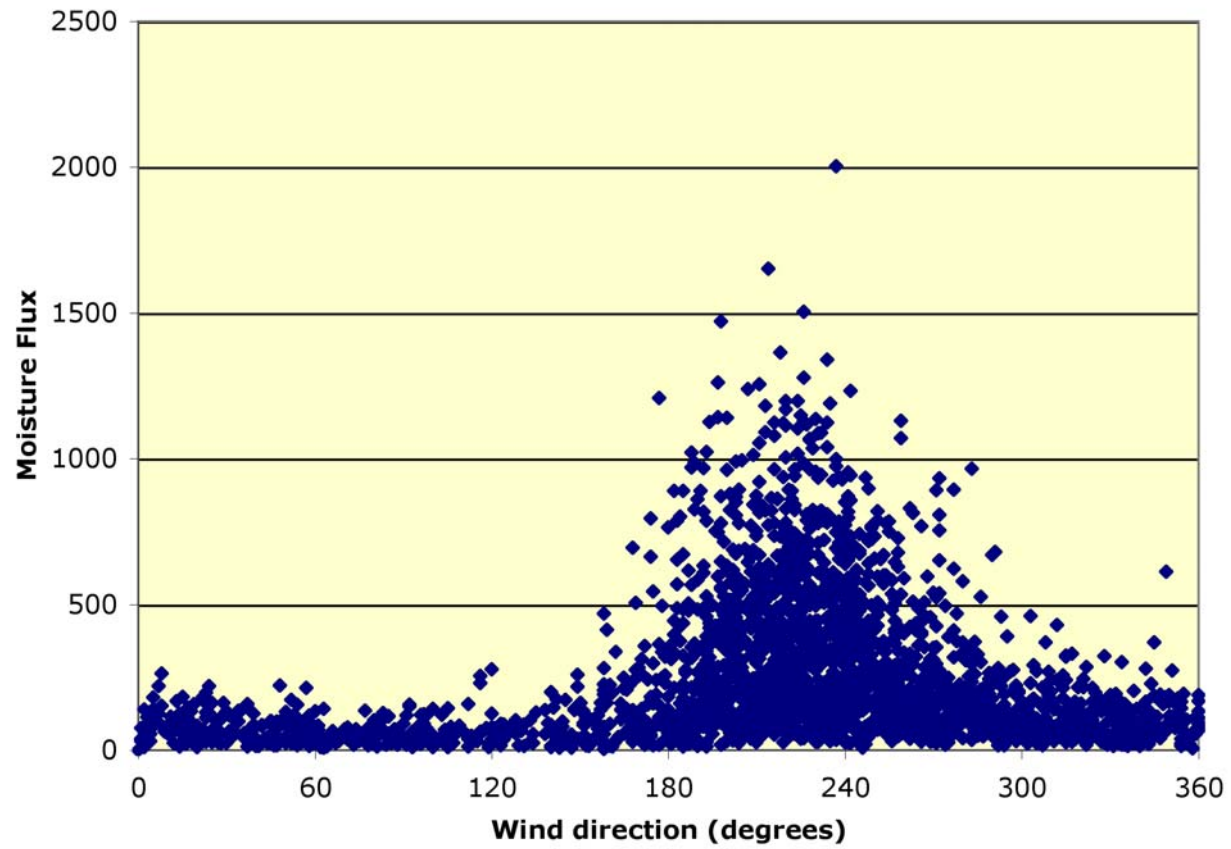
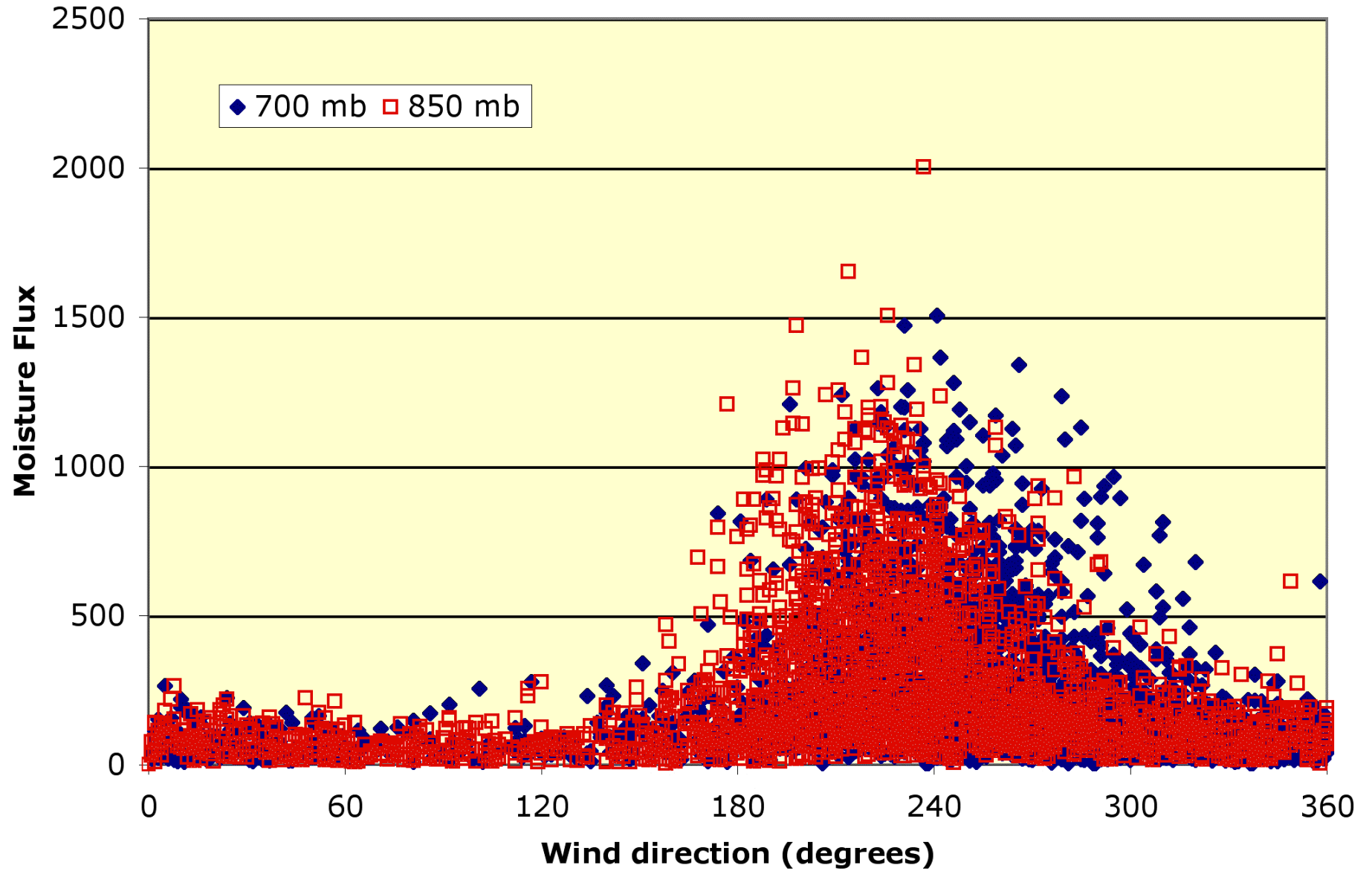


Figure 5. Climatological distribution of incoming moisture flux as a function of wind direction and fraction of time the wind is blowing from each direction at Oakland in (a) early winter (NDJ) and (b) late winter (FMA). The incoming moisture fluxes (solid line) are marked on the left y axis. The distributions of winds are marked on the right y axis and are shown for all days (dotted line), for days with any measurable precipitation ($p > 0$, dashed line), and for days when precipitation exceeds 25 mm/d (dashed-dotted line).

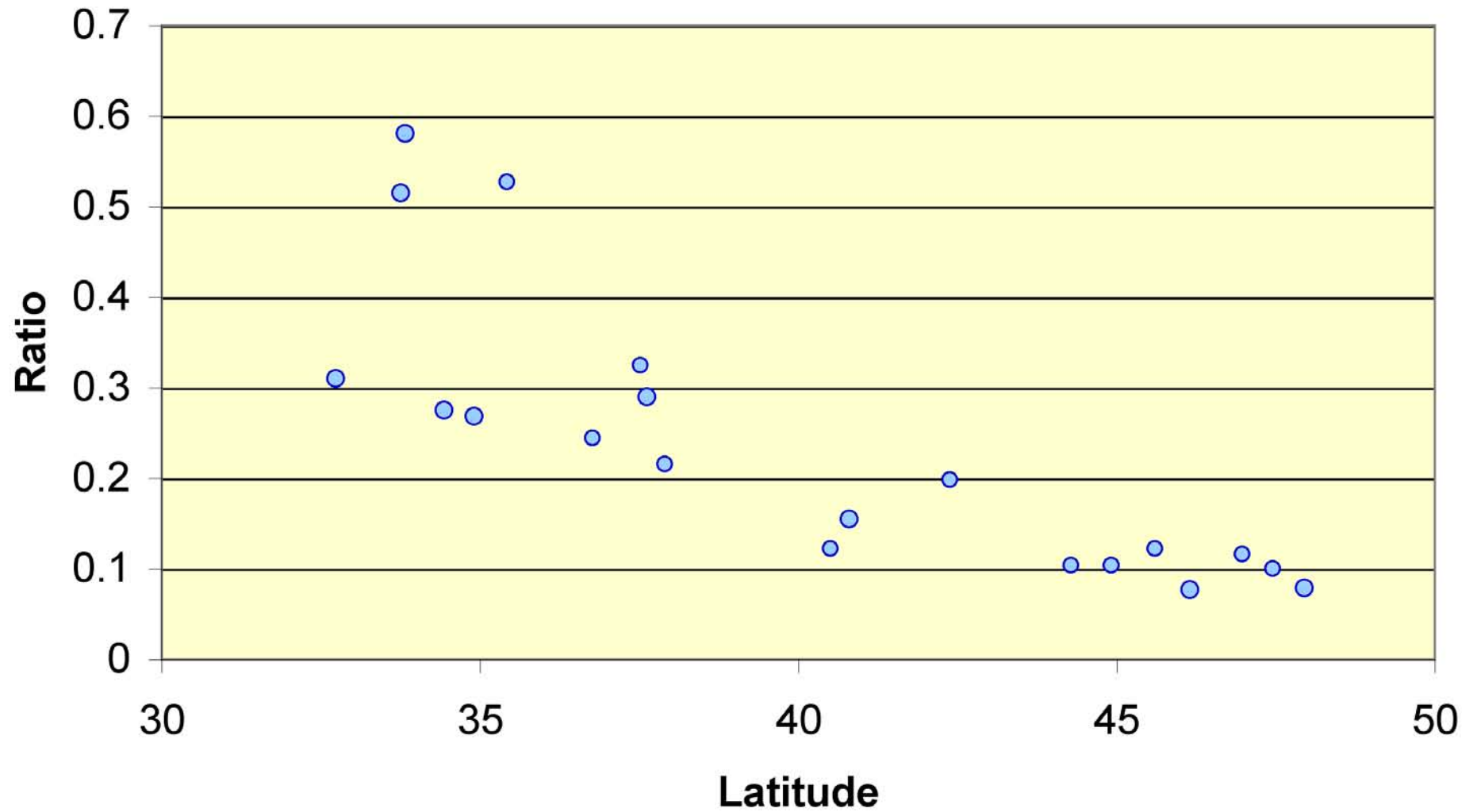
Salem 850 mb Wind Direction vs. Moisture Flux, January 1956-96



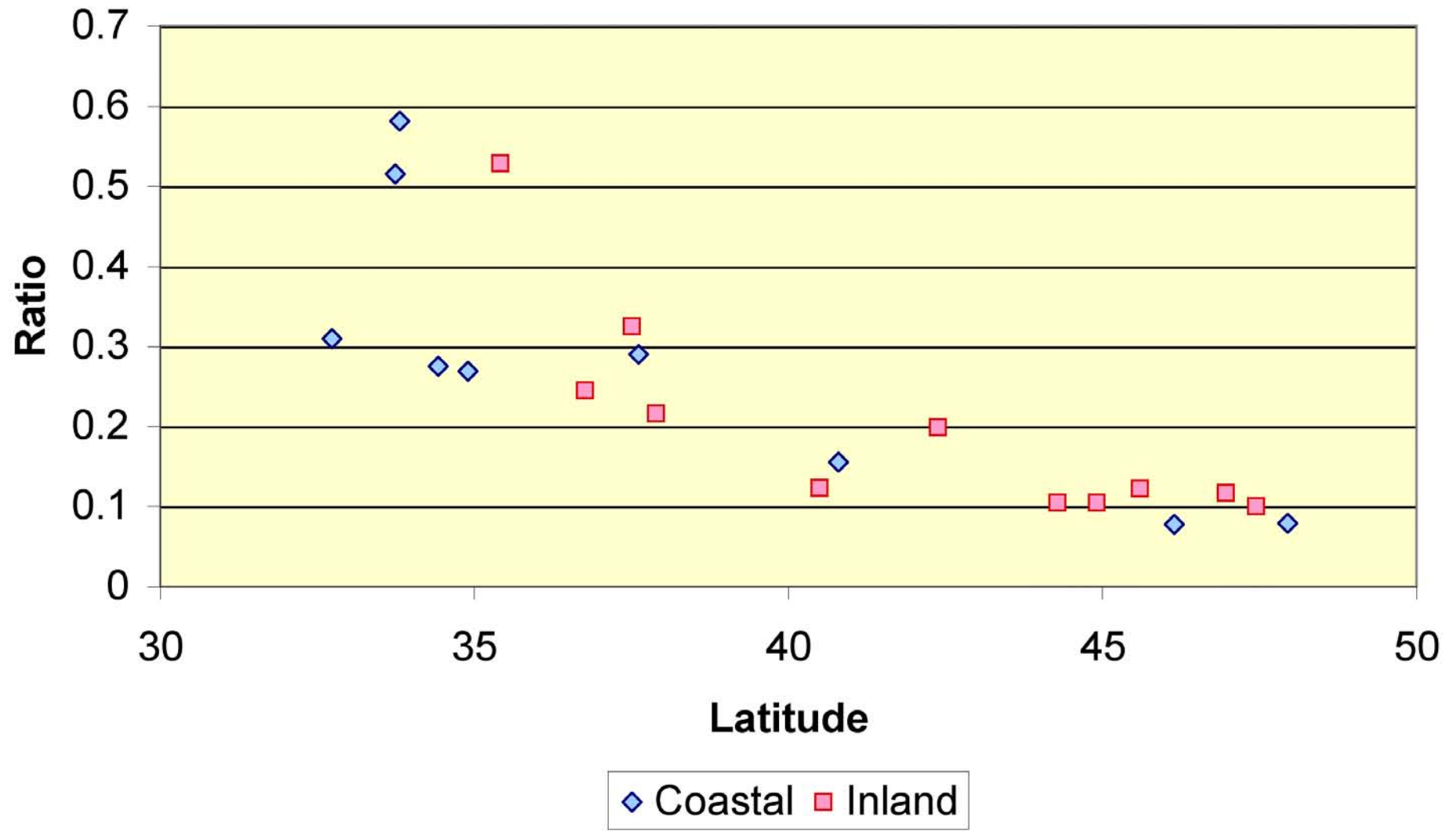
Moisture Flux, January 1956-96



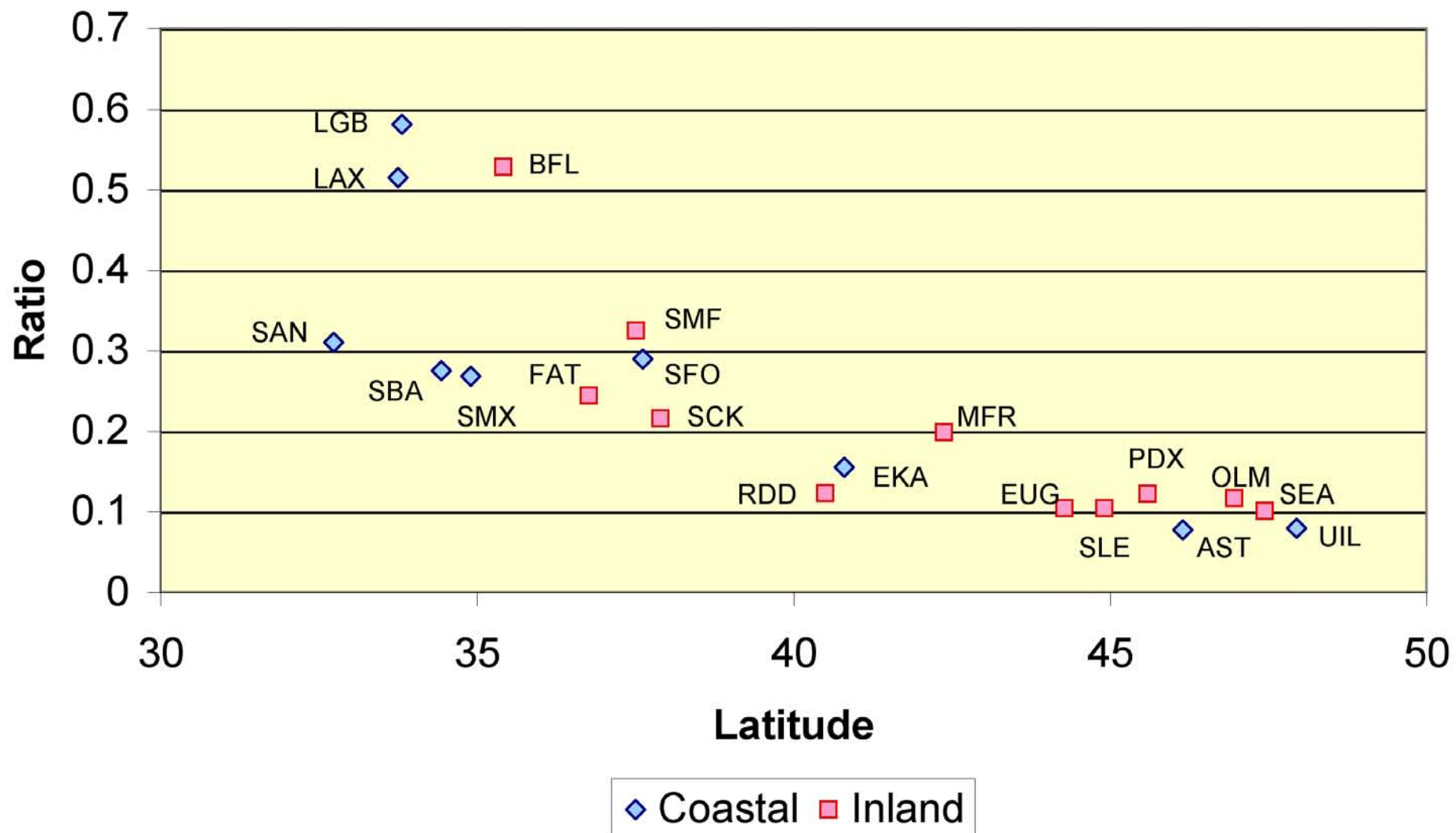
Ratio of Maximum 24-hour Precipitation to Mean Annual



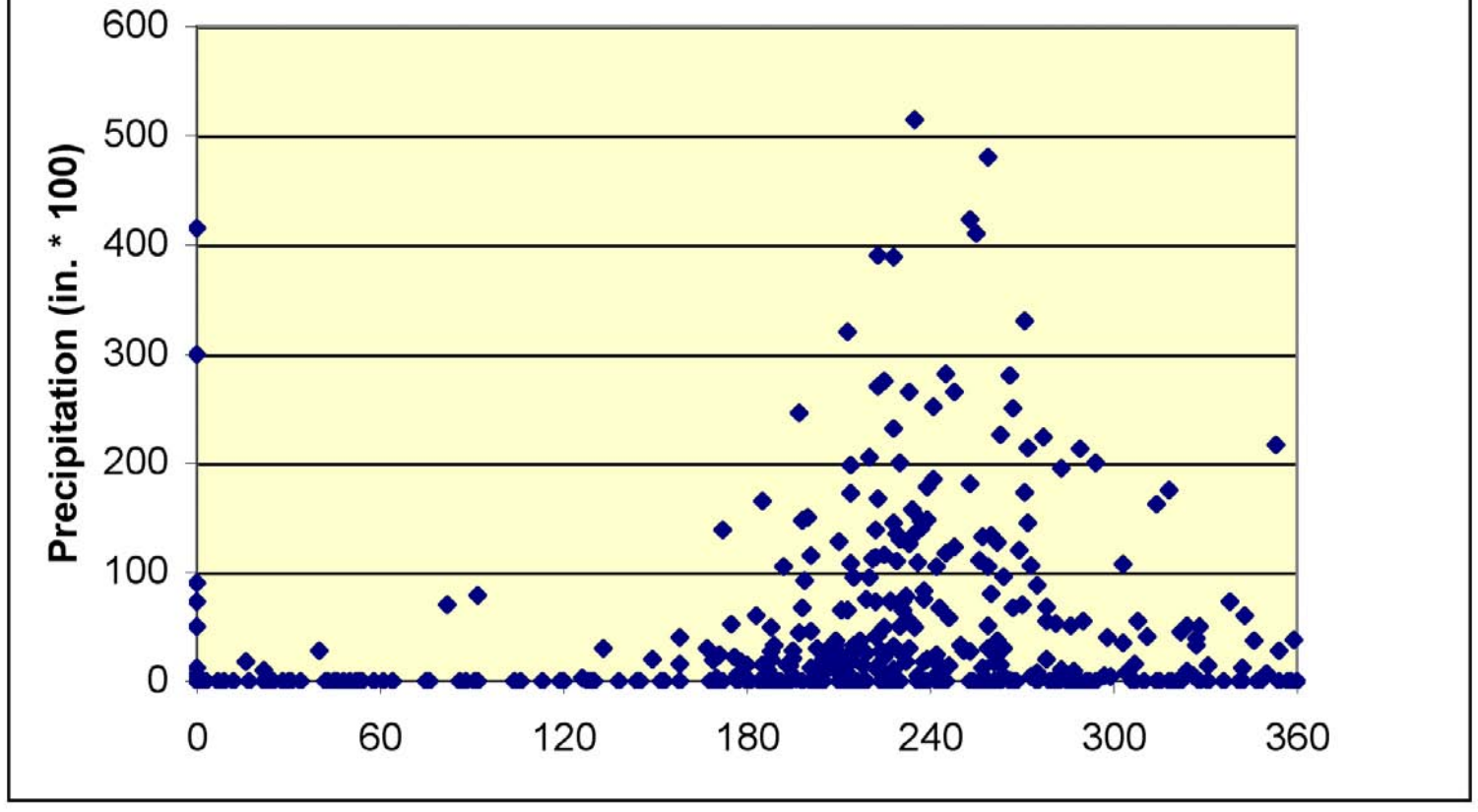
Ratio of Maximum 24-hour Precipitation to Mean Annual



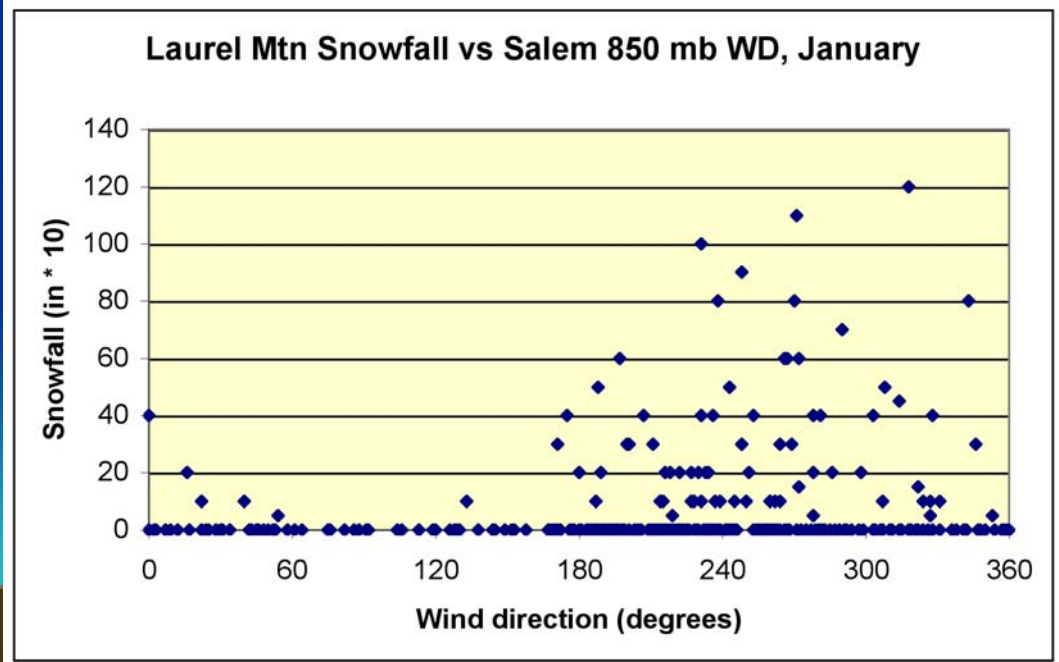
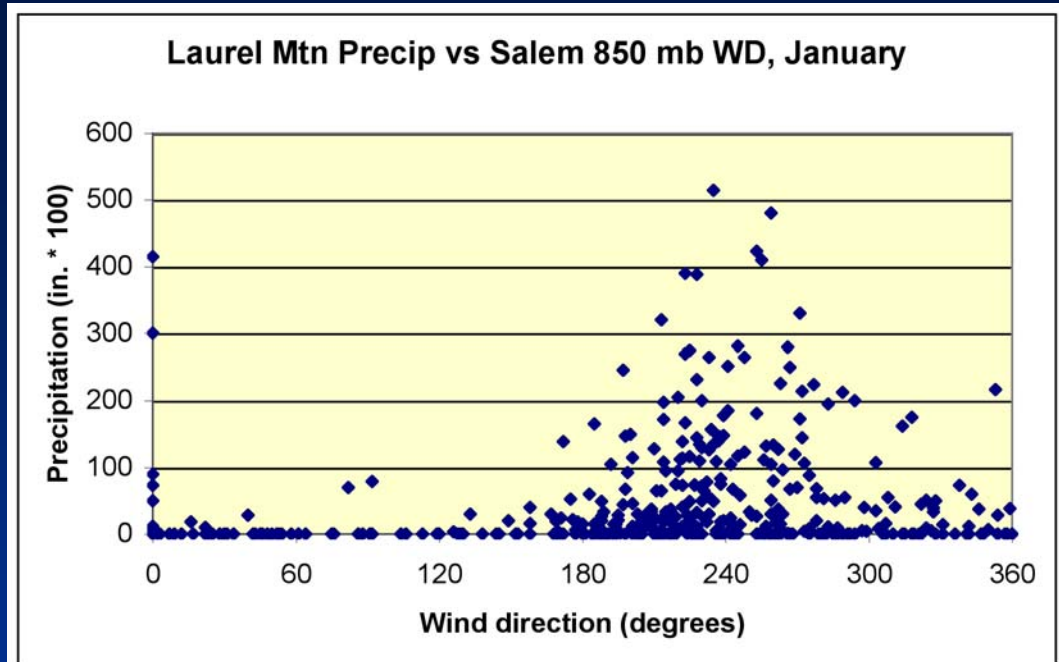
Ratio of Maximum 24-hour Precipitation to Mean Annual



Laurel Mtn Precip vs Salem 850 mb WD, January



Coast Range Precip and snowfall as functions of WD



The Northwest Version
of the
"Pineapple Express"

