



ELSEVIER

Catena 23 (1994) 213–225

CATENA

## The relation between surface rock-fragment cover and semiarid hillslope profile morphology

J.R. Simanton<sup>a</sup>, T.J. Toy<sup>b</sup>

<sup>a</sup>USDA-ARS, Southwest Watershed Research Center, 2000 E. Allen Rd., Tucson, AZ 85719, USA

<sup>b</sup>Dept. of Geography, University of Denver, Denver, CO 80208, USA

### Abstract

Rock-fragment covers veneer hillslopes in many parts of the world and influence the operation of hydrologic and geomorphic processes on those surfaces. The purpose of this investigation was to examine the spatial variability of rock-fragment cover on morphological units of hillslopes in southeastern Arizona, U.S.A. Simple statistical models, based upon hillslope gradient as well as a soil-slope factor (SSF) were developed and validated for the estimation of percent rock-fragment covers (RFc) on these units. Although reasonably accurate estimates were made using either hillslope gradient or the SSF, it appears more efficient to utilize hillslope gradient alone. Likewise, reasonably accurate estimates of RFc can be made using either a composite model for the entire hillslope or unit-specific models; however, it appears more efficient to utilize the gradient composite model.

### 1. Introduction

Recent research demonstrates that rock-fragment cover on hillslope surfaces profoundly influences the operation of hydrologic and geomorphic processes. Poesen et al. (1994) cite the direct effects of such covers as shielding the soil surface from detachment by rainsplash and runoff and entrapping splashed sediment.

#### *1.1. Rock-fragment cover effects on hillslope hydrology and hydraulics*

A number of investigations specifically address the relation between infiltration rates and rock-fragment cover. Results reveal both positive and negative associations between the two variables. Dadkhah and Gifford (1980) found, that on uncompacted soil, infiltration increased as rock-fragment cover increased. However, on compacted soil, there was no significant infiltration increase associated with increasing rock-

These hillslopes are components of alluvial basins within the Basin and Range Physiographic Province of the western U.S.A. (Simanton et al., 1994). The soils are well-drained, calcareous, gravelly loams with rock fragments covering a substantial percentage of the surface. Soil profiles are less than 2 m in depth, with the top 10 cm commonly containing up to 60% rock and gravel. Soils at the sites are classified as thermic *Ustollic Haplargids*; thermic, shallow *Typic Paleorthids*; thermic *Aridic Calcicustolls*; *Lithic Torriorthents*; and *Lithic Haplustolls* (Gelderman, 1970; Hendricks, 1985).

Annual precipitation at Walnut Gulch averages about 300 mm (Simanton et al., 1994). Tucson annual precipitation averages about 285 mm with both winter (November to April) and summer (July to mid-September) receiving about half of the annual. Summer thunderstorms at both locations produce nearly all the annual runoff and erosion from the hillslopes used in this study (Osborn et al., 1979).

Vegetation of Walnut Gulch is typical of the Chihuahuan desert (Simanton et al., 1994). The Sonoran desert vegetation in the Tucson area has similar species as those found at Walnut Gulch but also includes paloverde (*Cercidium microphyllum*) and saguaro cactus (*Carnegiea gigantea*).

### 3. Research methods

#### 3.1. Hillslope selection

Two groups of hillslopes are included in this study. The first consists of 12 hillslopes, within the Walnut Gulch Experimental Watershed, previously examined by Simanton et al. (1994). We graphically reconstructed (from field measurements of hillslope gradient and length) the hillslope profiles and classified them as to convex, straight, and concave units. The data within these groupings were used in the developmental models to be discussed later. The second group consists of 6 additional hillslopes (2 from the Walnut Gulch watershed and 4 from the foothills near Tucson) that were classified in like manner; and used for validation of the developmental models. All hillslopes were rectilinear, valley-side surfaces exhibiting no concentrated flow paths or established rill patterns. They were chosen to represent a range of shapes and gradients characteristic of the study locations.

#### 3.2. Measurements

Once a hillslope was selected, it was divided into morphological units. The mean gradient (%) and field length (m) of each unit was measured with an Abney level and tape from top to the bottom of the unit. Then, the percent rock-fragment cover (RFc) was determined for each unit using the line-point method (Bonham, 1989) at 150 mm intervals along a transect normal to the profile. Rock fragments were defined as particles greater than 5 mm in diameter. This dimension was selected as the lower limit because particles smaller than this can be easily moved from the hillslope by fluvial processes (Poesen, 1987; Parsons et al., 1991). Next, rock fragments were

removed from the surface at a representative location near the mid-point of each transect and two soil samples were taken from the top 50 mm of soil. Laboratory sieve analyses were used to determine the percent (by weight) rock fragments (> 5 mm diameter) in the samples.

The soil-slope factor (SSF) for each hillslope unit was calculated using the equation reported by Simanton et al. (1994):

$$\text{SSF} = \text{RFp}^{\text{SL}_d^{0.5}} \quad (1)$$

where SSF = soil-slope factor, RFp = % soil profile rock fragment content > 5 mm, and  $\text{SL}_d$  = hillslope gradient (decimal or tangent).

### 3.3. Statistical analyses, model development, and validation

Developmental models for the composite, convex, straight, and concave units were derived using regression techniques that related RFc to both the hillslope gradient and the SSF. Model validation was based on statistical comparison of measured RFc at the 6 validation sites with RFc estimated by the composite and appropriate unit-specific developmental models. For example, the developmental composite model was used to estimate RFc for all validation units regardless of morphology and these estimates were compared to the field measurements by means of regression analysis. To increase the robustness of the RFc predictive models, data from Simanton et al. (1994) and the 6 validation sites were combined to produce final models for composite and specific units.

## 4. Results and discussion

### 4.1. Spatial variability of rock-fragment covers

The sample size, mean, standard deviation and range of the hillslope gradient, RFc, and SSF for the developmental, validation, and final (combined) data sets are provided in Table 1. The values for the validation data are generally similar to those for the developmental data. Hence, the size of the validation data sets appears adequate to reproduce the basic statistical characteristics of the developmental data sets and to test the integrity of the developmental models. These data suggest little difference in RFc among the morphological units.

### 4.2. Developmental models

Models for estimating RFc were developed for the composite hillslope, convex, straight, and concave units. In these regression analyses, hillslope gradient and the SSF served as the independent variables while RFc was the dependent variable.

Table 1

Descriptive statistics: the sample size ( $N$ ); the mean; the standard deviation (STD); and the range for gradient (%), rock-fragment cover (%), and soil-slope factor used in the developmental and final rock-fragment cover models, and of the validation sites

| Model/unit           | $N$ | Slope (%) |      |       | Rock fragment (%) |      |       | Soil-slope factor |     |       |
|----------------------|-----|-----------|------|-------|-------------------|------|-------|-------------------|-----|-------|
|                      |     | Mean      | Std  | Range | Mean              | Std  | Range | Mean              | Std | Range |
| <i>Developmental</i> |     |           |      |       |                   |      |       |                   |     |       |
| Composite            | 61  | 21.2      | 16.2 | 2–61  | 46.3              | 17.1 | 1–72  | 5.8               | 4.2 | 1–19  |
| Convex               | 24  | 17.1      | 11.6 | 2–47  | 48.6              | 14.7 | 24–68 | 4.8               | 2.6 | 2–10  |
| Straight             | 22  | 25.1      | 20.6 | 2–61  | 48.1              | 17.6 | 8–72  | 7.0               | 5.6 | 1–19  |
| Concave              | 15  | 21.9      | 14.4 | 3–50  | 40.0              | 19.4 | 1–67  | 5.7               | 3.8 | 2–14  |
| <i>Validation</i>    |     |           |      |       |                   |      |       |                   |     |       |
| Composite            | 31  | 19.8      | 12.3 | 4–53  | 47.6              | 13.2 | 22–71 | 5.7               | 3.7 | 2–17  |
| Convex               | 7   | 23.6      | 11.4 | 6–26  | 49.0              | 14.3 | 28–64 | 4.2               | 1.6 | 2–6   |
| Straight             | 16  | 20.3      | 14.6 | 4–53  | 47.8              | 14.2 | 22–71 | 6.0               | 4.5 | 2–17  |
| Concave              | 8   | 15.9      | 7.1  | 9–42  | 46.2              | 11.4 | 23–60 | 6.3               | 3.3 | 2–13  |
| <i>Final</i>         |     |           |      |       |                   |      |       |                   |     |       |
| Composite            | 92  | 20.7      | 14.9 | 2–61  | 46.7              | 15.8 | 1–72  | 5.8               | 4.0 | 1–19  |
| Convex               | 31  | 16.6      | 10.7 | 2–47  | 47.9              | 14.0 | 24–68 | 4.7               | 2.4 | 2–10  |
| Straight             | 38  | 23.0      | 18.3 | 2–61  | 47.9              | 16.0 | 8–72  | 6.6               | 5.1 | 1–19  |
| Concave              | 23  | 22.4      | 13.0 | 3–50  | 43.1              | 17.8 | 1–67  | 5.9               | 3.6 | 2–14  |

#### 4.2.1. Hillslope gradient

The analyses for RFc estimation based upon hillslope gradient are presented in Fig. 2. In all these models the natural logarithm of gradient provided the most accurate estimations. The model for composite hillslopes (Fig. 2a), with all unit data taken together, had a standard error of estimate (SE) of 8.84% and a coefficient of determination ( $R^2$ ) of 74%. The model for convex elements alone (Fig. 2b) yielded a slightly lower SE (7.90%) and a slightly lower  $R^2$  (72%) than for the composite model. This indicates little advantage in using a unit-specific equation to estimate RFc of convex hillslope elements. The straight segments model (Fig. 2c) had a SE of 6.35% and a  $R^2$  of 88% indicating some improvement in the accuracy of RFc estimation using the unit-specific model rather than the composite model. The model for concave elements (Fig. 2d) produced a SE of 6.33% and a  $R^2$  of 90%, again indicating a modest increase in the accuracy of RFc estimation by using the unit-specific model rather than the composite model.

#### 4.2.2. Soil-slope factor (SSF)

The analyses for the estimation of RFc based upon the SSF are presented in Fig. 3. For the composite and straight segments, the reciprocal of the SSF provided the most accurate RFc estimations, while for the convex and concave elements, the natural logarithm of SSF provided the most accurate estimations. The composite model (Fig. 3a) had a SE of 7.98% and a  $R^2$  of 79%. The model for convex elements (Fig. 3b) had a SE of 7.19% and a  $R^2$  of 77%. As with the gradient models there is little advantage

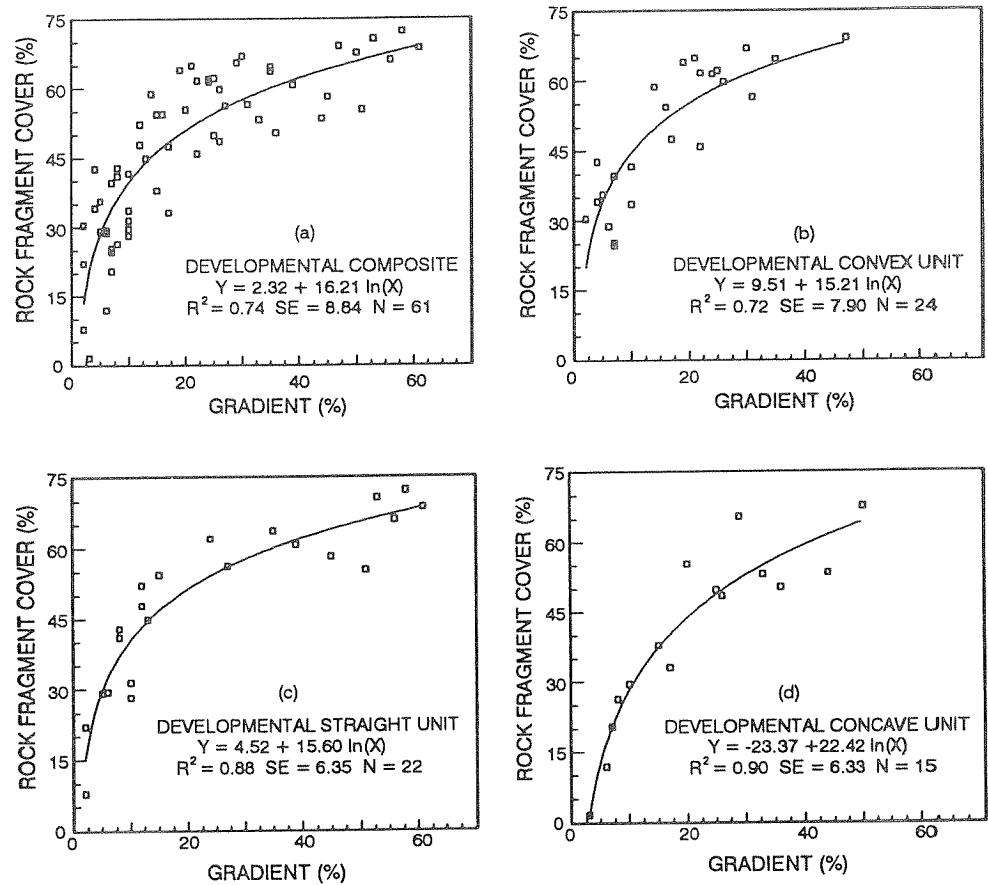


Fig. 2. Relations between hillslope gradient (%) and rock-fragment cover (%) of the composite (a), convex (b), straight (c), and concave (d) units for the developmental models.

in using the unit-specific model to estimate Rfc of convex elements. The straight segment model (Fig. 3c) had a SE of 6.30% and a  $R^2$  of 88%. Compared to the composite model, some increase in the accuracy of estimation might be expected based upon this unit-specific model. The model for concave elements (Fig. 3d) produced a SE of 7.04% and a  $R^2$  of 88%. Again, a modest increase in the accuracy of estimation might be expected based upon the unit-specific model rather than the composite model. In both the hillslope gradient and SSF models, the relation between the residuals from regression and the values of the dependent variable suggested random distributions.

#### 4.3. Model validation

Model validation consists of statistical comparisons between field Rfc measure-

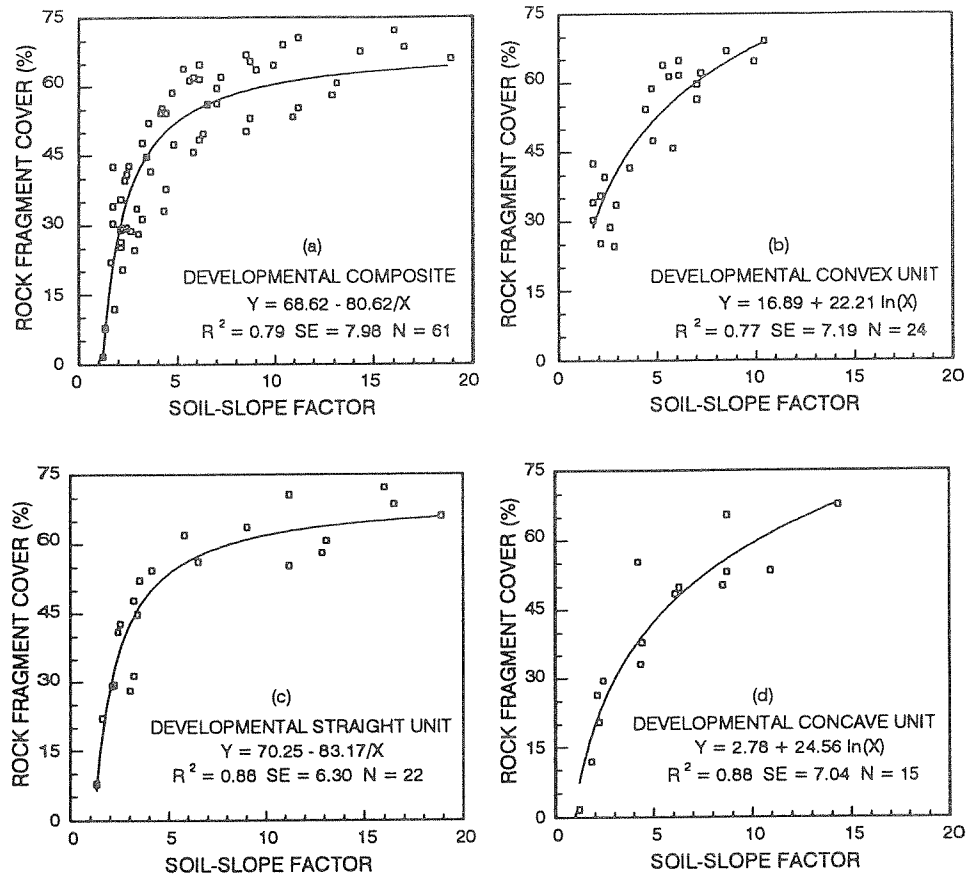


Fig. 3. Relations between the soil-slope factor and rock-fragment cover (%) of the composite (a), convex (b), straight (c), and concave (d) units for the developmental models.

ments at the validation sites and R<sub>Fc</sub> estimates produced by the developmental models based on gradient or the SSF (Fig. 4). Simple regression was used for comparisons; a perfect relation between measured and estimated would result in a  $R^2$  of 100%, a  $Y$ -intercept of "0", and a regression coefficient of 1.0 (1:1 line in Fig. 4). The gradient composite model (Fig. 2a) was used to estimate R<sub>Fc</sub> for all hillslope segments of the validation sites regardless of hillslope morphology. The results are shown in Fig. 4a including a  $R^2$  of 91% and a regression coefficient of 0.77. The R<sub>Fc</sub> field measurements were compared to estimates based on gradient unit-specific models (Fig. 2b, c, and d) as appropriate (i.e. R<sub>Fc</sub> of convex units were estimated using the convex unit-specific model). The results are shown in Fig. 4b including a  $R^2$  of 89% and a regression coefficient of 0.77.

The validation of the developmental models using the SSF as the independent

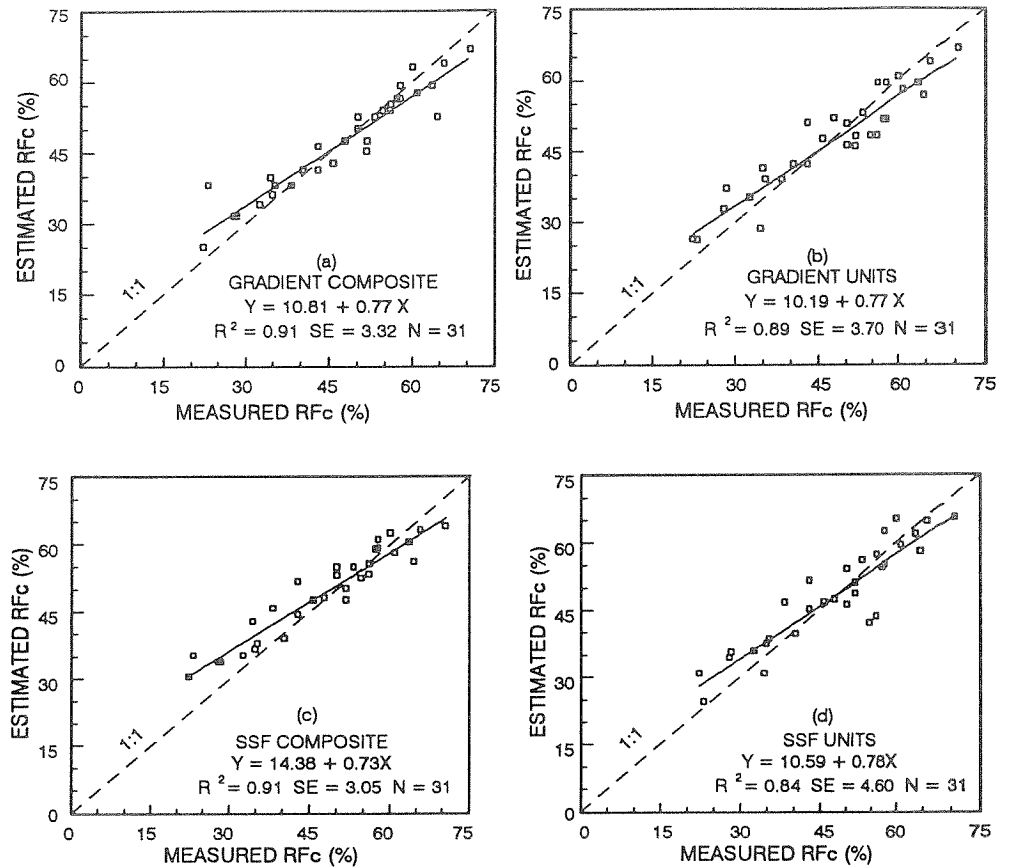


Fig. 4. Measured versus estimated rock-fragment cover based on hillslope gradient for the validation composite (a), units (b); and the soil-slope factor estimated rock-fragment cover for the validation composite (c), and units (d).

variable is also presented in Fig. 4. The comparison of field measurements and R<sub>Fc</sub> estimates based upon the composite model (Fig. 3a) yielded a  $R^2$  of 91% and a regression coefficient of 0.73 (Fig. 4c). The comparison of field measurements and estimates based on the unit-specific models (Figs. 3b, c, d) yielded a  $R^2$  of 84% and regression coefficient of 0.78 (Fig. 4d).

The results of the validation tests indicate that either gradient or the SSF can be used to estimate R<sub>Fc</sub> for the composite or unit-specific components of hillslopes. The SE for the various developmental models differ by only a few percent, although the differences in the  $R^2$  are somewhat larger. The greater than zero regression intercept values (Fig. 4) suggest some measure of bias inherent in the models. We cannot explain this phenomena because the measurements were taken in exactly the same manner for all of the data collected.

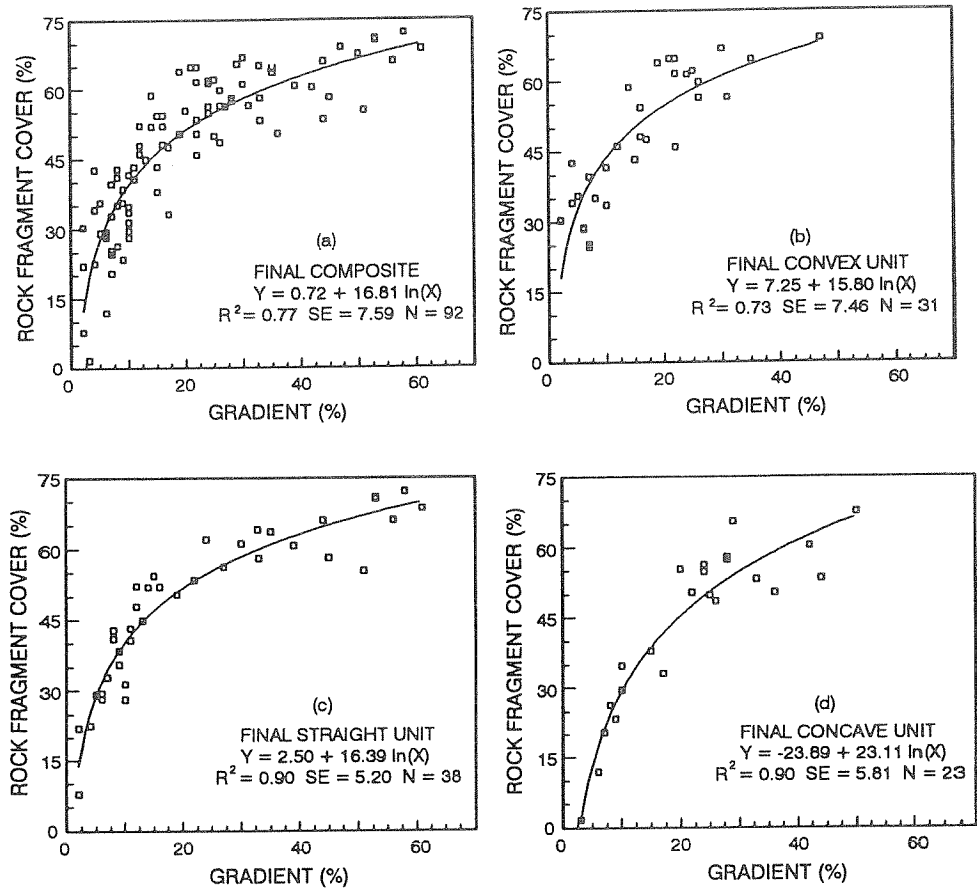


Fig. 5. Relations between hillslope gradient and rock-fragment cover of the composite (a), convex (b), straight (c), and concave (d) units for the final models.

#### 4.4. Final models

The data used for model development and validation were combined to construct the final models depicted in Figs. 5 and 6.

##### 4.4.1. Hillslope gradient

The natural logarithm of the hillslope gradient was the independent variable in each of these final models. The final composite model (Fig. 5a) was developed from data for 92 segments and yielded a SE of 7.59% and a  $R^2$  of 77%. The final model for convex elements (Fig. 5b) has a slightly lower SE (7.46%) and a slightly lower  $R^2$  (73%) than the final composite model (Fig. 5a). The SE (5.20%) of the final model for straight segments (Fig. 5c) was lower and the  $R^2$  (90%) was higher than for the final composite (Fig. 5a). The final model for the concave element (Fig. 5d) had a slightly



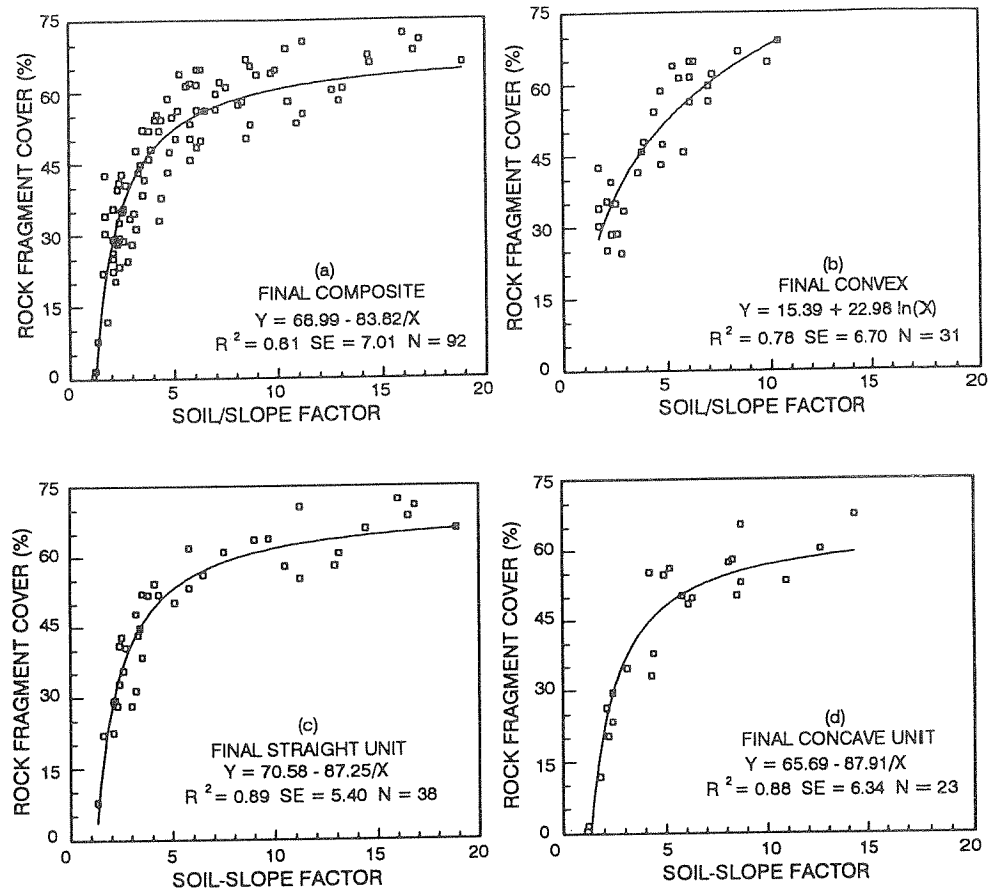


Fig. 6. Relations between the soil-slope factor and rock-fragment cover of the composite (a), convex (b), straight (c), and concave (d) units for the final models.

lower SE (5.81%) than the final composite (Fig. 5a) and the  $R^2$  (90%) was higher than for the final composite model (Fig. 5a).

#### 4.4.2. Soil-slope factor (SSF)

Similar increases and decreases in standard errors of estimates and coefficients of determination were obtained for the SSF final models. However, the final composite, straight, and concave models (Fig. 6a, c, and d) employed the reciprocal of the SSF as the independent variable. The final model for convex elements (Fig. 6b) used the natural logarithm of the SSF as the independent variable and yielded a slightly lower SE (6.70%) and a slightly higher  $R^2$  (78%) than for the final composite (Fig. 6a). Both the final straight and concave models (Fig. 6c and 6d) had lower standard errors of estimates and larger coefficients of determination than for the SSF composite final model (Fig. 6a).

## 5. Conclusion and observation

The results of the foregoing analyses support the following conclusions and observations:

(1) There is little difference in the RFc on the various morphological units of hillslopes in the environments from which the data were collected. These results, similar to Simanton et al. (1994), indicate that even when morphological units of hillslopes are considered individually that the suite of hydrologic and geomorphic processes operating on hillslopes is likely to be similar in frequency and magnitude, or at least insufficiently different to produce concomitant differences in RFc.

(2) The differences in RFc that exist between morphological units appear to be closely related to differences in hillslope gradient or the SSF. Either of these variables can be used to make reasonable estimates of RFc. Because of the similarity in the estimate errors associated with the use of these two variables, there is little to warrant the additional time and effort necessary to collect and analyze the soil samples as required for the computation of the SSF. Hence, we recommend that estimates of RFc utilize hillslope gradient alone.

(3) There is also similarity in the estimate errors associated with the use of the composite and unit-specific models. Unless a probable reduction in estimate error of very few percent is important for a particular application, we recommend that RFc estimates utilize a composite model. This would conserve time and effort necessary to determine hillslope unit morphology.

(4) Collectively the best choice of models for most purposes is the final composite model presented in figure 5a. Soil erosion models, such as RUSLE (Renard et al., 1991) or WEPP (Lane and Nearing, 1989) that seek to account for the effects of rock-fragment covers on erosion rates may estimate the RFc on the basis of hillslope gradient alone. This model (Fig. 5a) is most appropriate for areas with environmental conditions similar to those from which the data were collected.

(5) Abrahams and Parsons (1992, p. 419) note that “models of hillslope development must recognize the critical role of surface gravel covers in controlling resistance to flow and, hence, the rate of erosion; where gravel size or concentration varies in a systematic manner down a hillslope, form resistance and erosion rate will also vary, and this systematic variation may have a significant effect on the shape of the hillslope as it evolves through time.” The results of our study suggest that in these environments, the downslope variation in RFc is largely a function of changes in hillslope gradient. Consequently, systematic changes in one should be linked to systematic changes in the other.

## Acknowledgement

We wish to thank Drs. A.D. Abrahams; J.E. Box, Jr.; W.E. Emmerich; G.F. Gifford; and K.G. Renard for their review and constructive comments of this paper. Also, thanks to C.B. Escapule for the timely and efficient reduction of the field samples.

## References

- Abrahams, A.D. and Parsons, A.J., 1991. Resistance to overland flow on desert pavement and its implications for sediment transport modeling. *Water Resour. Res.*, 27: 1827–1836.
- Abrahams, A.D. and Parsons, A.J., 1992. Overland flow on semiarid Arizona hillslopes. In: D.G. Janelle (Editor), *Geographical Snapshots of North America*. Guilford Press, New York, NY, pp. 416–419.
- Bonham, C.D., 1989. *Measurements for Terrestrial Vegetation*. Wiley, New York, 338 pp.
- Cooke, R.U., 1970. Stone pavements in deserts. *Ann. Assoc. Am. Geogr.*, 60: 560–577.
- Cooke, R.U. and Warren, A., 1973. *Geomorphology in Deserts*. Batsford, London, 374 pp.
- Dadkhah, M and Gifford, G.F., 1980. Influence of vegetation, rock cover, and trampling on infiltration rates and sediment production. *Water Resour. Bull.*, 16: 979–986.
- De Ploey, J., 1981. The ambivalent effects of some factors of erosion. *Mem. Inst. Geol. Louvain*, 31: 171–181.
- Gelderman, F.W., 1970. Soil Survey, Walnut Gulch experimental watershed. Arizona Special Report, USDA-SCS. 55 pp.
- Hadley, R.F. and Toy, T.J., 1977. Relation of surficial erosion on hillslope to profile geometry. *J. Res. U.S. Geol. Surv.*, 5: 487–490.
- Hendricks, D.M., 1985. *Arizona Soils*. College of Agriculture, Univ. of Arizona, Tucson, AZ, 244 pp.
- Lane, L.J. and Nearing, M.A., 1989. USDA-Water Erosion Prediction Project: Hillslope profile model documentation. NSERL Report No. 2, USDA-ARS National Soil Erosion Research Laboratory, West Lafayette, IN, 274 pp.
- Meyer, L.D., Foster, G.R. and Romkens, M.J.M., 1975. Sources of soil eroded by water from upland slopes. Present and Prospective Technology for Predicting Sediment Yields and Sources. U.S.D.A., Agricultural Research Service, ARS-S-40, pp. 177–189.
- Osborn, H.B., Renard, K.G. and Simanton, J.R., 1979. Dense networks to measure convective rainfall in the Southwestern United States. *Water Resources Research*, 15: 1701–11.
- Parsons, A.J. and Abrahams, A.D., 1987. Gradient-particle size relations on quartz monzonite debris slopes in the Mojave desert. *J. Geol.*, 95: 423–452.
- Parsons, A.J., Abrahams, A.D. and Luk, S., 1991. Size characteristics of sediment in interrill overland flow on a semiarid hillslope, southern Arizona. *Earth Surf. Process. Landforms*, 16: 143–152.
- Poesen, J., 1987. Transport of rock fragments by rill flow — a field study. In: R.B. Bryan (Editor), *Rill Erosion*. *Catena Suppl.*, 8: 35–54.
- Poesen, J., Ingelmo-Sanchez, F. and Mucher, H., 1990. The hydrological response of soil surfaces to rainfall as affected by cover and position of rock fragments in the top layer. *Earth Surf. Process. Landforms*, 15: 653–671.
- Poesen, J., Torri, D. and Bunte, K., 1994. Effects of rock fragments on soil erosion by water at different spatial scales: A review. *Catena*, 23: 141–166.
- Renard, K.G., Foster, G.R., Weesies, G.A. and Porter, J.P., 1991. RUSLE: Revised universal soil loss equation. *J. Soil Water Conserv.*, 46: 30–33.
- Shaw, C.F., 1929. Erosion pavement. *Geogr. Rev.*, 19: 638–641.
- Simanton, J.R., Rawitz, E. and Shirley, E.D., 1984. The effects of rock fragments on erosion of semiarid rangeland soils. In: *Erosion and Productivity of Soils Containing Rock Fragments*. SSSA Special Pub. 13, Soil Science Society of America, Madison, WI, Chpt. 7, pp. 65–72.
- Simanton, J.R., Renard, K.G., Christiansen, C.M. and Lane, L.J., 1994. Spatial distribution of surface rock fragments along catenas in semiarid Arizona and Nevada, USA. *Catena*, 23: 29–42.
- Young, A., 1972. *Slopes*. Longman, London, 288 pp.
- Young, R.A. and Mutchler, C.K., 1969. Effect of slope shape on erosion and runoff. *Trans. Am. Soc. Agric. Eng.*, 12: 231–239.

.....

.....

.....