

Prepared in cooperation with the State of Hawaii Commission on Water Resource Management

Median and Low-Flow Characteristics for Streams under Natural and Diverted Conditions, Northeast Maui, Hawaii



Scientific Investigations Report 2004-5262

U.S. Department of the Interior U.S. Geological Survey

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By Stephen B. Gingerich

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Multiply	Ву	To obtain
	Length	
inch (in.)	2.54	centimeter (cm)
inch (in.)	25.4	millimeter (mm)
foot (ft)	0.3048	meter (m)
mile (mi)	1.609	kilometer (km)
	Area	
acre	0.004047	square kilometer (km ²)
square foot (ft ²)	0.09290	square meter (m ²)
square mile (mi ²)	2.590	square kilometer (km ²)
	Volume	
million gallons (Mgal)	3,785	cubic meter (m ³)
	Flow rate	
foot per second (ft/s)	0.3048	meter per second (m/s)
cubic foot per second (ft ³ /s)	0.02832	cubic meter per second (m ³ /s)
million gallons per day (Mgal/d)	0.04381	cubic meter per second (m ³ /s)

Vertical coordinate information is referenced relative to local mean sea level.

Horizontal coordinate information is referenced to North American Datum of 1983 (NAD 83).

Altitude, as used in this report, refers to distance above the vertical datum.

Median and Low-Flow Characteristics for Streams under Natural and Diverted Conditions, Northeast Maui, Hawaii

By Stephen B. Gingerich

Abstract

Flow-duration statistics under natural (undiverted) and diverted flow conditions were estimated for gaged and ungaged sites on 21 streams in northeast Maui, Hawaii. The estimates were made using the optimal combination of continuous-record gaging-station data, low-flow measurements, and values determined from regression equations developed as part of this study. Estimated 50- and 95-percent flow duration statistics for streams are presented and the analyses done to develop and evaluate the methods used in estimating the statistics are described. Estimated streamflow statistics are presented for sites where various amounts of streamflow data are available as well as for locations where no data are available.

Daily mean flows were used to determine flow-duration statistics for continuous-record stream-gaging stations in the study area following U.S. Geological Survey established standard methods. Duration discharges of 50- and 95-percent were determined from total flow and base flow for each continuousrecord station. The index-station method was used to adjust all of the streamflow records to a common, long-term period. The gaging station on West Wailuaiki Stream (16518000) was chosen as the index station because of its record length (1914–2003) and favorable geographic location. Adjustments based on the index-station method resulted in decreases to the 50-percent duration total flow, and 95-percent duration base flow, 95-percent duration total flow, and 95-percent duration base flow computed on the basis of short-term records that averaged 7, 3, 4, and 1 percent, respectively.

For the drainage basin of each continuous-record gaged site and selected ungaged sites, morphometric, geologic, soil, and rainfall characteristics were quantified using Geographic Information System techniques. Regression equations relating the non-diverted streamflow statistics to basin characteristics of the gaged basins were developed using ordinary-leastsquares regression analyses. Rainfall rate, maximum basin elevation, and the elongation ratio of the basin were the basin characteristics used in the final regression equations for 50percent duration total flow and base flow. Rainfall rate and maximum basin elevation were used in the final regression equations for the 95-percent duration total flow and base flow. The relative errors between observed and estimated flows ranged from 10 to 20 percent for the 50-percent duration total flow and base flow, and from 29 to 56 percent for the 95-percent duration total flow and base flow.

The regression equations developed for this study were used to determine the 50-percent duration total flow, 50-percent duration base flow, 95-percent duration total flow, and 95-percent duration base flow at selected ungaged diverted and undiverted sites. Estimated streamflow, prediction intervals, and standard errors were determined for 48 ungaged sites in the study area and for three gaged sites west of the study area. Relative errors were determined for sites where measured values of 95-percent duration discharge of total flow were available. East of Keanae Valley, the 95-percent duration discharge equation generally underestimated flow, and within and west of Keanae Valley, the equation generally overestimated flow. Reduction in 50- and 95-percent flow-duration values in stream reaches affected by diversions throughout the study area average 58 to 60 percent.

Introduction

For more than a century, surface-water diversion systems have transported water from the wet, northeastern part of Maui, Hawaii, to the drier, central part of the island, mainly for large-scale sugarcane cultivation. Since the 1930's, the Territory and then the State issued water permits to Alexander and Baldwin, Inc., Hawaiian Commercial and Sugar Co., and East Maui Irrigation Co., Ltd. (EMI), for the diversion of water from streams in northeast Maui. The collection system consists of 388 separate intakes, 24 miles of ditches, and 50 miles of tunnels, as well as numerous small dams, intakes, pipes, and flumes (Wilcox, 1996). With few exceptions, the diversions capture all of the base flow, which represents the ground-water contribution to total streamflow, and an unknown percentage of total streamflow at each stream crossing. During 1925-97, total flow for the diversion systems measured at Honopou Stream, to the west of the study area where records of total diversion-system flow are most complete, averaged about 163 Mgal/d (million gallons per day) (Gingerich, 1999). The highest average flow for an individual ditch system was measured in the Koolau/Wailoa Ditch system, where total flow crossing Honopou Stream averaged 110 Mgal/d for 1924-87. The source of diverted water is a watershed with an area of about 56,000 acres, about two-thirds of which is owned by the State (Wilcox, 1996) and managed by the State Department of Land and Natural Resources.

The Hawaii State Water Code mandates that the Commission on Water Resource Management (CWRM) establish

a statewide instream-use protection program (Chapter 174C-71, Hawaii Revised Statutes). The principal mechanism that CWRM has for the protection of instream uses is establishing instream flow standards. "Each instream flow standard shall describe the flows necessary to protect the public interest in the particular stream. Flows shall be expressed in terms of variable flows of water necessary to protect adequately fishery, wildlife, recreational, aesthetic, scenic, or other beneficial instream uses in the stream in light of existing and potential water developments including the economic impact of restriction of such use" (Chapter 174C-71, Hawaii Revised Statutes). CWRM has recognized certain instream uses as beneficial, including: (1) maintenance of fish and wildlife habitat, (2) outdoor recreational activities, (3) maintenance of ecosystems such as estuaries, wetlands, and stream vegetation, (4) aesthetic values such as waterfalls and scenic waterways, (5) maintenance of water quality, (6) conveyance of irrigation and domestic water supplies to downstream points of diversion, and (7) protection of traditional and customary Hawaiian rights.

The U.S. Geological Survey, in cooperation with CWRM and in collaboration with the Maui Department of Water Supply, the Hawaii State Board of Land and Natural Resources, and East Maui Irrigation Co., Ltd., undertook an investigation to assist in determining equitable, reasonable, and beneficial instream and off-stream uses of the surface-water resources of northeast Maui. The overall objectives of the 3-year study are to (1) assess the effects of existing surface-water diversions on flow characteristics for perennial streams in northeast Maui, (2) characterize the effects of diversions on instream temperature variations, and (3) estimate the effects that streamflow restoration (full or partial) will have on habitat availability for native stream fauna (fish, shrimp, and snails) in northeast Maui. Scientific information generated by the overall study will allow CWRM to complete its work on documenting water rights and uses associated with northeast Maui streams and analyzing the economic effects of curtailing existing uses on the streams, and to then establish technically defensible instream flow standards for those streams.

Purpose and Scope

This report addresses objective 1 described above. This report presents selected estimated flow-duration statistics for streams in northeast Maui, Hawaii, and describes the analyses done to develop and evaluate the methods used in estimating the statistics. Estimated streamflow statistics are presented for sites where various amounts of streamflow data are available and for locations where no data are available. Morphometric, hydrologic, and geologic basin characteristics are provided for each stream basin in the study area. Equations used to estimate the 50- and 95-percent duration flows of total streamflow and stream base flow at ungaged locations are presented. An evaluation of the accuracy of the equations and limitations for their use is also provided. Most-reliable estimates of streamflow statistics for natural (undiverted) and unnatural (diverted) sites on 21 streams and the basis for these estimates are provided.

The statistics for undiverted and diverted flow were compared to assess the effects of existing surface-water diversions on flow characteristics for perennial streams in northeast Maui.

Description of Study Area

The study area lies on the northern flank of the East Maui Volcano (Haleakala), which forms the eastern part of the island of Maui, the second-largest island in the Hawaiian archipelago. The study area, covering about 67 mi2, is bounded to the north by about 11 mi of coastline and lies between (and includes) the drainage basins of Kolea Stream to the west and Makapipi Stream to the east (fig. 1). Land-surface altitudes range from sea level to 10,000 ft at the summit of Haleakala. The topography is gently sloping except for the steep sides of gulches and valleys that were eroded by the numerous streams. The largest valley is Keanae Valley, which extends from the coast to Haleakala Crater where the valley walls are nearly 1,000 ft high. Most of the study area is made up of forest reserves; at intermediate altitudes, rain forests densely cover the slopes up to about 7,000 ft. Grasses and shrubs cover the upper slopes to the north wall of Haleakala Crater. Two small villages (Keanae and Wailua) are at low altitudes along the coast at the mouth of Keanae Valley. Land use around the villages is mainly small-scale agriculture, including wetland taro cultivation.

Streams flow generally south to north from the high altitude flank of Haleakala to the coast. Twenty-two named streams reach the coast in the study area. Access to streams is made difficult by the steep rugged terrain of the incised stream valleys and dense native and non-native vegetation. Rainfall is highly orographic and rates average between about 45 in/yr at the summit of Haleakala to greater than 350 in/yr at about 2,500-ft altitude. Rainfall at the coast ranges from 120 to 160 in/yr (Giambelluca and others, 1986).

Previous Studies

Low-flow duration statistics have not previously been estimated specifically for ungaged streams in the study area of northeast Maui. Fontaine and others (1992) developed regression equations (one for Oahu, Molokai, and Hawaii, and one for Maui and Kauai) to estimate median flows at ungaged, unregulated, perennial streams in the State. Data from gaging stations on ten streams in the northeast Maui study area were used in developing the median-flow equations. Hirashima (1965) and Matsuoka (1983) computed flow statistics for gaging stations throughout the State of Hawaii, including some stations in the study area. Yamanaga (1972) developed regression equations for low-flow frequency to describe annual minimum 7-day and 30-day mean flows at 2- and 20-year recurrence intervals using data from selected windward and leeward gaged basins across the State, including 14 stations in the current study area. The equations and flow-duration statistics presented in this report supersede any previously reported equations or flow-duration statistics. Gingerich (1999)



Figure 1. Northeast Maui study area, island of Maui, Hawaii.

described the ground-water occurrence and contribution to streamflow for northeast Maui covering an area encompassing the current study area. That report detailed the amount of streamflow, base flow, and surface-water diversions in streams and gave detailed descriptions of low-flow measurements made in many of the streams in the current study area.

Numbering System for Surface-Water Gaging Stations

The surface-water gaging stations mentioned in this report are numbered according to the USGS "downstream order" numbering system. Station numbers increase in a downstream direction along the main stream. All stations on a tributary entering upstream from a mainstream station have lower station numbers. A station on a tributary that enters between two mainstream stations is given a number between those two station numbers. In this report, the complete 8-digit downstream-order number for each gaging station has been abbreviated to the middle four digits, for example, 16518000 becomes 5180.

Acknowledgments

The author would like to thank Chiu Yeung and Chien-Hwa Chen for delineating drainage basins and determining the basin characteristics, and Richard Fontaine and Michael Wong for assistance with the statistical analysis used in this report. Layout work on this report was done by Luis Menoyo. East Maui Irrigation Co., Ltd. provided cooperation and assistance in accessing the study area.

Streamflow Characteristics at Continuous-Record Stream-Gaging Stations

Values of daily mean flow are used to determine total streamflow and base flow flow-duration statistics for continuous-record stream-gaging stations. The USGS has operated 16 continuous-record stream-gaging stations for various periods at unregulated sites in or near the northeast Maui study area since 1910 (Fontaine, 1996), although only two of these stations (station 5080 on Hanawi Stream near Nahiku and station 5180 on West Wailuaiki Stream near Keanae) are currently active (Plate 1, table 1). In addition, records from 17 continuous-record stream-gaging stations on regulated sites are available for analysis.

The USGS established standard methods for estimating flow-duration statistics for stream-gaging stations (Searcy, 1959). A flow-duration curve is a graphical representation of the percentage of time streamflows for a given time interval (usually daily) are equaled or exceeded over a specified period at a stream site (fig. 2). Flow-duration curves are constructed by first ranking all of the daily mean discharge values for the period of record at a gaging station, next computing the probability of each value being equaled or exceeded, and then plotting the discharges against their associated exceedance probabilities. Flow-duration statistics are points along the flow-duration curve. For example, the 50-percent duration streamflow (or median streamflow; Q_{50}) has been exceeded 50 percent of the time during the specified period. Flow-duration statistics reflect streamflow conditions only for the period of record for which they were calculated. If the period of record analyzed is sufficiently long, the flow-duration statistics can be considered an indicator of probable future conditions (Searcy, 1959). In an analysis of the data from five long-term stream gages on Oahu, the median discharge determined from a 10-yr streamflow record had a standard error of 15 percent, whereas the standard error for the median discharge determined from a 50-yr record improved to 6 percent (Fontaine, 1996).

At three regulated gaging stations (5090, 5110, and 5210), total unregulated streamflow was calculated by adding the daily flows for the gaging station of interest to the corresponding daily flows for an upstream gaging station on the same stream but above the 1,300-ft (Koolau Ditch) diversion (the flows at stations 5080, 5100, and 5190 plus 5200, respectively). Flow-duration statistics were then calculated using the combined record. This technique is appropriate for estimating the low flows of interest in this study because the diversion captures all low flows much greater than the value of the median total flow, TFQ_{50} , so that the downstream gaging station on each stream measures only that flow gained below the diversion. Combined flows at the downstream gaging station estimated from this technique are incorrect only when the diversion is overtopped, which is generally 20 to 30 percent of the time.

The assumption that the diversion systems within the 1,700- to 1,200-ft altitude interval intercept all low streamflows up to at least the TFQ₅₀ is based on measurements made at four stream diversions in the study area. The overtopping discharges in the four monitored streams were significantly greater than the TFQ₅₀ discharge for each stream. These sites are separated into two categories: (1) sites on West Wailuaiki and Hanawi Streams where water levels at the diversion structures and streamflow were measured concurrently at active continuous-record gaging stations, and (2) sites on Waikamoi and Honomanu Streams, where water levels at the diversion structures were measured concurrently with water levels at staff plates from former continuous-record gaging stations (fig. 3). At all four sites, submersible transducers were used to monitor the water level in the stream every 15 minutes relative to the crests of the diversion dams preventing water from flowing further downstream. When a transducer indicated that the water level was higher than the crest of the diversion dam, it was assumed that water was flowing downstream past the diversion dam. When the transducer indicated that the water level was lower than the crest of the diversion dam, it was assumed that all flow in the stream was captured by the diver
 Table 1.
 Continuous-record surface-water gaging stations operated by the U.S. Geological Survey, northeast Maui, Hawaii.

[[]Abbreviated station numbers are highlighted in **bold**; active station locations are shown in **bold italics**]

Gaging- station number	Station location	Period of record	Station altitude (feet)	Low flow regulated during period of record
16 5070 00	Makapipi Stream	1932-45	920	Yes
16 5080 00	Hanawi Stream	1914-15, 1921-Present	1,318	No
16 5090 00	Hanawi Stream	1932-47, 1992-95	500	Yes
16 5100 00	Kapaula Gulch	1921-63	1,346	No
16 5110 00	Kapaula Gulch	1932-47	540	Yes
16 5130 00	Waiaaka Stream	1932-47	650	Yes
16 5140 00	Paakea Gulch	1932-47	650	Yes
16 5150 00	Waiohue Gulch	1921-63	1,316	No
16 5160 00	Kopiliula Stream	1914-17, 1921-1958	1,292	No
16 5170 00	East Wailuaiki Stream	1914-17, 1922-58	1,329	No
16 5180 00	West Wailuaiki Stream	1914-17, 1921-Present	1,343	No
16 5190 00	West Wailuanui Stream	1914-17, 1921-58	1,268	No
16 5200 00	East Wailuanui Stream	1914-17, 1921-58	1,287	No
16 5210 00	Wailuanui Stream	1932-36, 1938-47	620	Yes
16 5220 00	Palauhulu diversion ditch to Keanae ^a	1934-68	51	Yes
16 5240 00	Honomanu Stream	1921-27, 1932-34, 1962-68	2,900	No
16 5270 00	Honomanu Stream	1914-17, 1921-64	1,733	No
16 5310 00	Kula diversion from Haipuaena Stream	1946-68	4,320	Yes
16 5311 00	Haipuaena Stream	1946-68	4,320	Yes
16 5350 00	Haipuaena diversion ditch to Kolea Stream	1938-60	1,866	Yes
16 5360 00	Haipuaena Stream	1946-60	1,512	Yes
16 5420 00	East Branch Puohokamoa Stream	1921-27, 1931-33	2,800	No
16 5430 00	Middle Branch Puohokamoa Stream	1921-27, 1932-34, 1962-69	2,900	Yes
16 5440 00	West Branch Puohokamoa Stream	1921-28, 1932-34	2,800	Yes
16 5450 00	Puohokamoa Stream	1914-17, 1921-71	1,322	Yes
16 5528 00	Waikamoi Stream	1953-68	4,487	Yes
16 5540 00	Waikamoi Stream	1921-28, 1932-34	3,000	Yes
16 5545 00	East Branch Waikamoi Stream	1921-28, 1932-33	3,020	Yes
16 5550 00	Waikamoi Stream	1922-57	1,294	Yes
16 5560 00	Waikamoi Stream	1914-17, 1921-22	1,150	Yes
16 5570 00	Alo Stream	1914-17, 1921-57	1,248	No
16 5650 00 ^b	Kaaiea Gulch	1921-62	1,310	No
16 5660 00 ^b	Oopuola Stream	1930-57	1,205	No
16 5700 00 ^b	Nailiihaele Stream	1910-11, 1913-75	1,205	No
16 5770 00 ^b	Kailua Stream	1910-11, 1913-58	1,253	No

^aUSGS station previously published as "Taro patch feeder ditch at Keanae, Maui"

^b located west of current study area



Figure 2. Flow-duration curves of total streamflow and base flow at gaging station 5180 on West Wailuaiki Stream, northeast Maui, Hawaii, for period 1914-2002.

sion and the stream was dry immediately downstream of the diversion.

At West Wailuaiki and Hanawi Streams, the water levels were compared with discharge determined from stage measurements made every 15 minutes at the active continuous-record gaging stations upstream (stations 5180 and 5080, respectively). Relation plots of water level and stream discharge show the range of discharges at which the diversion dam is overtopped and flow continues downstream. For West Wailuaiki Stream, the overtopping is initiated at discharge ranging from 20 to 30 ft³/s (TFQ₃₀ to TFQ₂₀ on the basis of the flow-duration plot in fig. 2). In other words, streamflow does not pass the diversion dam at 1,300 ft on West Wailuaiki Stream roughly 70 to 80 percent of the time. On Hanawi Stream, the overtopping discharge ranges from 15 to 30 ft³/s $(TFQ_{25} \text{ to } TFQ_{15} \text{ on the basis of the flow-duration plot for})$ station 5080 discussed later in fig. 7). Streamflow does not pass the dam at 1,300 ft on Hanawi Stream roughly 75 to 85 percent of the time.

At Honomanu and Waikamoi Streams, the water levels relative to the crests of the diversion dams were compared to water levels collected every 15 minutes at staff plates bolted into the bedrock in former gaging station pools upstream from the dams. Ratings for these staff plates are available from the U.S. Geological Survey archives for the time when the stations were discontinued in 1957. Although some erosion on the bedrock and concrete controls at these former gaging stations was observed, the assumption was made that the ratings for staff plates would still provide a reasonable range of flow estimates in the stream to estimate the overtopping discharge. At Honomanu Stream, the overtopping discharge ranges from 15 to 18 ft³/s (TFQ₂₅ to TFQ₂₂ on the basis of the flow-duration plot for station 5270 discussed later in fig. 7). Therefore, streamflow does not pass the diversion dam at 1,720 ft on Honomanu Stream roughly 75 to 78 percent of the time. On Waikamoi Stream, where the water level-discharge relation recorded is more variable, the overtopping discharge is at least 25 ft³/s (TFQ₂₀ on the basis of the flow-duration plot for station 5550 discussed later in fig. 7). Streamflow does not pass the diversion dam at 1,200 ft on Waikamoi Stream roughly 80 percent of the time.

The technique used to determine the combined median and low-flow statistics for a gaging station downstream of the major diversion can be illustrated by using data from Hanawi Stream. From gaging station 5080, upstream of the Koolau diversion, TFQ₅₀ is 7.1 ft³/s (table 2) and from gaging station 5090, downstream of the Koolau diversion and several spring inflows, TFQ₅₀ is 22 ft³/s. Because the Koolau Ditch captures at least 15-30 ft³/s at Hanawi Stream near 1,300 ft altitude, the median flows measured at gaging station 5090 are only those flows gained downstream of the Koolau diversion. The combined TFQ₅₀ at gaging station 5090 is estimated to be 29 ft³/s (7.1 + 22 ft³/s). Median and low-flow statistics for 5090 were generated in two ways, from combined daily flows at the two stations and by adding the statistics calculated from each station's daily flows individually (Table 3). The results of each calculation compare favorably, indicating that, for the study area, flow statistics for sites on gaining streams above and below the major diversion can be estimated by adding the statistics for each site. For the three stations in the study area (5090, 5110, and 5210), the value determined from adding the flow statistics was always less than the value determined



Relation between water level and stream discharge at diversions on selected streams within the 1,700- to 1,200-ft altitude interval, northeast Maui, Hawaii. Figure 3.



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DISCHARGE AT FORMER GAGING STATION 5270 ON HONOMANU STREAM, IN CUBIC FEET PER SECOND





0

DAM CREST, IN FEET

NOISRAVID OT AVITAJAR JAVAJ RATAW

Relation between water level and stream discharge at diversions on selected streams within the 1,700- to 1,200-ft altitude interval, northeast Maui, Hawaii-Continued Figure 3.

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Hawaii.
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Table 2.

[Qxx is the xx-percent flow duration of streamflow; ft³s, cubic feet per second; base period is 1914-17, 1921-2001; active stations are shown in **bold italics**; +, combined with record from indicated station; index station is station 5180; --, no adjustment; NA, not applicable]

	l anoth of	O ₅₀ tota	l flow	0 ₅₀ bas	se flow	Q ₉₅ tot	al flow	Q ₉₅ bas	e flow
Gaging-station number	concurrent record (years)	during concurrent period (ff ³ /s)	adjusted to index station (ft³/s)	during concurrent period (ff³/s)	adjusted to index station (ft³/s)	during concurrent period (ft³/s)	adjusted to index station (ft³/s)	during concurrent period (ft³/s)	adjusted to index station (ft³/s)
5070	13	2.9	2.2	1.6	1.3	0.00	0.00	0.00	0.00
5080	82	7.1	ł	4.6	ł	2.4	1	2.2	1
5090+5080	18	29	28	25	24	19	19	19	19
5100	42	5.2	4.9	2.9	2.8	1.3	1.1	1.0	06.0
5110+5100	15	8.1	7.5	5.3	5.1	3.4	3.3	3.1	3.0
5130	15	06.0	0.86	0.79	0.77	0.54	0.54	0.54	0.53
5140	15	4.2	4.0	3.9	3.8	3.1	3.0	3.1	3.0
5150	42	6.5	6.2	5.0	5.0	3.1	3.0	2.9	2.8
5160	40	9.1	8.0	5.4	5.0	2.6	2.3	2.1	2.0
5170	39	10	9.1	6.1	5.7	3.1	2.8	2.6	2.5
5180	83	10	1	6.0	ł	2.5	1	2.1	1
5190	40	5.1	4.4	2.7	2.4	1.1	1.0	0.93	0.85
5200	40	3.7	3.2	2.2	2.0	1.0	06.0	0.82	0.80
5210+5190+5200	13	9.8	10	5.6	6.1	2.4	2.5	2.1	2.0
5220ª	34	3.4	1	NA	NA	2.4	1	NA	NA
5240	14	2.0	2.2	0.88	0.93	0.31	0.37	0.28	0.30
5270	46	6.2	5.7	3.0	2.8	1.2	1.0	0.82	0.74
$5311 + 5310^{a}$	22	0.46	0.50	0.15	0.17	0.05	0.03	0.03	0.03
5360+5310 ^a +5350	14	6.8	6.8	3.5	3.5	2.1	1.7	1.6	1.3
5420	8	0.93	0.88	0.58	0.53	0.31	0.29	0.19	0.27
5430	15	1.3	1.4	0.62	0.67	0.18	0.21	0.12	0.15
5440	6	1.9	1.9	1.1	1.0	0.46	0.55	0.37	0.44
5450	53	13	12	6.8	6.4	2.8	2.5	2.2	2.1
5528	15	0.12	0.14	0.05	0.06	0.02	0.02	0.02	0.01
5540	7	2.5	2.5	1.3	1.2	0.62	0.69	0.32	0.46
5545	8	1.5	1.3	0.80	0.72	0.46	0.47	0.37	0.43
5550	35	7.9	7.0	3.8	3.5	1.3	1.1	0.89	0.80
5560	4	15	7.1	5.1	3.6	2.0	1.5	1.6	1.0
5570	39	3.1	2.7	1.6	1.4	0.77	0.70	0.62	0.58
^a gaging station on di	version ditch								

 Table 3.
 Comparison of flow statistics computed for stream sites where natural flows are estimated, using data from multiple continuous-record gaging stations, northeast Maui, Hawaii.

Flow statistic	Downstream gaging station	Upstream gagi station	ng Sum of upstream and downstream flow statistic	Flow statistic based on combined up- stream and down- stream daily flows	Relative error, in percent
	5090	5080			
$\mathrm{FQ}_{\mathrm{50}}$	19	7.1	26	28	L-
$\mathrm{FQ}_{\mathrm{50}}$	19	4.6	24	24	0
FQ_{05}	16	2.4	18	19	-5
FQ ₉₅	16	2.2	18	19	Ń
	5110	5100			
${}^{\mathrm{FQ}}_{\mathrm{50}}$	4.7	2.4	7.1	7.5	-5
FQ_{50}	2.8	2.2	5.0	5.1	-2
FQ_{95}	1.1	1.9	3.0	3.3	6-
FQ_{95}	0.90	1.9	2.8	3.0	L-
	5210	5190 521	00		
FQ_{50}	1.6	4.4 3.	2 9.2	10	8-
$\mathrm{FQ}_{\mathrm{50}}$	1.0	2.4 2.	0 5.4	6.1	-11
FQ_{95}	0.39	1.0 0.9	0 2.3	2.5	8-
FQ	0.32	0.85 0.8	0 2.0	2.0	0

by first combining the concurrent daily flows, but the relative error in the value determined from adding the flow statistics was less than or equal to 10 percent in each case.

On Haipuaena Stream (plate 1), gaging station 5310 recorded diverted flow into the Upper Kula Pipeline near 4,300 ft altitude, and gaging station 5311 recorded streamflow past the diversion (Fontaine, 1996). These two records were added together to obtain the total streamflow at this altitude. Further downstream, flow diverted in Kolea Stream was measured at gaging station 5350. This flow was added to that at gaging station 5360 (1,512 ft altitude) upstream of Spreckels Ditch and gaging station 5310 to obtain a value for total streamflow at 1,512 ft altitude.

Gingerich (1999) used a computerized base-flow separation method described by Wahl and Wahl (1995) to estimate the base-flow component of streamflow for northeast Maui streams. Two variables, N (number of days) and f (turningpoint test factor) must be assigned values in the method. The method divides the daily streamflow record into non-overlapping N-day periods and determines the minimum flow within each N-day window. If the minimum flow within a given Nday window is less than f times the minimums in the adjacent N-day windows, then the central window minimum is made a turning point on the base-flow hydrograph. Wahl and Wahl (1995) recommend a value of 0.9 for the turning-point test factor for most applications. The value of N determined for each stream is shown in table 2 of Gingerich (1999). A base-flowduration curve can than be constructed using the daily baseflow data (fig. 2).

Streamflow records generally are adjusted to a common base period for comparison so that differences in flow among stations reflect spatial differences in climate and drainagebasin characteristics and not simply temporal differences in rainfall. Flow-duration curves based on short records are unreliable for predicting the future flow pattern, but they can be made more reliable by adjusting them to represent longer periods. The index-station method described in Searcy (1959) was used to adjust all of the streamflow records used in this analysis to a common period (called the base period).

Index station selection

The two currently active continuous-record gaging stations (5180 on West Wailuaiki Stream and 5080 on Hanawi Stream) are the obvious candidates for index stations in this study. Both have been operated nearly continuously from 1914 to the present (2004), with a break during 1918–1920. The median total (TFQ₅₀) and median base flow (BFQ₅₀) for gaging station 5180 are 10 ft³/s and 6.0 ft³/s, respectively, and the TFQ₅₀ and BFQ₅₀ for gaging station 5080 are 7.1 ft³/s and 4.6 ft³/s, respectively (fig. 4 and table 2).

The records for gaging-stations 5180 and 5080 appear to have long-term downward trends in annual median total flow and base flow during the period 1914–present (fig. 4) and a Kendall's Tau trend test (Helsel and Hirsch, 1992) on

those flow statistics confirms such a trend (table 4). Rather than decreasing monotonically, the flow might actually have decreased in a stepwise manner, and examination of figure 4 indicates that such a stepwise decrease may have occurred about 1941–42. The median total flow at gaging station 5180 during 1914-42 was 14 ft³/s, and after 1942 was 9.6 ft³/s. Analysis of a subset of the data, for the period 1942-2001, shows no statistical trend in median streamflow after the indicated earlier stepped decrease in flow (table 4). No obviously apparent reason exists for such a stepped decrease in median total flow and base flow about 1941-42. Such stepped changes in flow trends are typically caused by changes in a watershed such as reforestation or by the addition of a streamflow diversion but there is little evidence that these factors caused the changes in flow at stations 5180 (on West Wailuaiki Stream) and 5080 (on Hanawi Stream). Helsel and Hirsch (1992) strongly caution against performing step-trend analysis without prior knowledge of an event that would contribute to such a change in streamflow. Additionally, rainfall records for this area do not cover the entire period of streamflow record so it is not possible to determine if rainfall patterns have changed similarly to the flow. Therefore, flow statistics for these stations were calculated on the basis of the entire period of record with no adjustments made for monotonic or step-wise trends in the data.

Gaging station 5180 on West Wailuaiki Stream is favorably located geographically as an index station, because it is near the center of the study area, whereas gaging station 5080, on Hanawi Stream, is at the eastern end of the study area. Therefore, all adjustments of streamflow characteristics to a common base period for the continuous-record stations were made using only gaging station 5180 as the index station. The record for gaging station 5080 was not adjusted because the record was the same length as the index-station base period.

Adjustments to streamflow characteristics for a common period using the index-station method

Relations between the index station and each shorter-term record at the other gaging stations were developed using the following steps (Searcy, 1959):

- 1. Flow-duration curves were developed for the station with a short-term record and the index-station record for concurrent periods of record.
- 2. Discharges for 13 flow-duration points, ranging from 1 to 99 percent, at the short-term station were plotted on logarithmic scales against the same flow-duration points at the index station.
- 3. A line or smooth curve was drawn through the points. The upper part of the line is typically a 45-degree line parallel (assuming 1 log cycle on x-axis is same length as 1 log cycle on y-axis) to lines of equal yield (drainage-area ratio) and equal flow (rainfall ratio) if the



Gaging station number	Annual median streamflow	Period of record	Kendall's Tau	P-level	Slope of trend
5180	Total flow	1914-17, 1921-2001	-0.171	0.022	-0.039
5180	Base flow	1914-17, 1921-2001	-0.210	0.005	-0.025
5080	Total flow	1914-15, 1921-2001	-0.194	0.010	-0.030
5080	Base flow	1914-15, 1921-2001	-0.296	0.000	-0.025
5180	Total flow	1942-2001	0.012	0.893	0.000
5180	Base flow	1942-2001	0.005	0.964	0.001
5080	Total flow	1942-2001	-0.010	0.919	0.000
5080	Base flow	1942-2001	-0.069	0.440	-0.005

 Table 4.
 Trend analysis of annual median flow at active gaging station records for West Wailuaiki (5180) and Hanawi (5080) Streams, northeast Maui, Hawaii.

streams have similar high flow characteristics. Deviation from the line at the lower points is commonly an indication of different geologic characteristics that effect low flows between the stream basins.

4. The adjusted discharges at the various flow-duration points for the short-term station are graphically determined from the plot using the corresponding flowduration points from the index station for the entire period of record (base period) of the index station.

For example, selected flow-duration points for total flow and base flow at gaging station 5170 on East Wailuaiki Stream, which was operated during 1914–17 and 1922–58, are plotted against the flow-duration points for the index station (5180) during the same period on logarithmic axes (fig. 5) and smooth fitting lines are drawn through the points. Matching base period flow-duration points are determined from the relation plot (e.g. base period TFQ₅₀ at station 5180 = 10 ft³/s and adjusted TFQ₅₀ at station 5170 = 9.0 ft³/s; base period TFQ₉₅ at station 5180 = 2.5 ft³/s and adjusted TFQ₉₅ at station 5170 = 2.8 ft³/s) and then plotted to determine the adjusted flow-duration curve for station 5170 (fig. 6). Similar relation plots and adjusted flow-duration curves were computed for 26 continuous-record gaging stations for streams in the study area (fig. 7).

The adjusted Q_{50} and Q_{95} statistics for total flow and base flow for the 26 stations are listed in table 2. Adjustments to TFQ₅₀ ranged from a 17-percent increase (station 5528) to a 53-percent decrease (station 5560), and averaged a 6-percent decrease. Adjustments to TFQ₉₅ ranged from a 20-percent increase (station 5440) to a 40-percent decrease (stations 5310 + 5311) and averaged a 5-percent decrease. Adjustments to BFQ50 ranged from a 20-percent increase (station 5528) to a 29-percent decrease (station 5560) and averaged a 4-percent decrease. Adjustments to BFQ₉₅ ranged from a 44-percent increase (station 5540) to a 50-percent decrease (station 5528) and averaged a 1-percent decrease. In general, the largest adjustments were needed for stations with the shortest record lengths.

The point on the total flow-duration curve at which streamflow is equivalent to median base flow provides added information about the ground-water contribution to streamflow. BFQ₅₀ ranges from 56 percent to 78 percent on the total flow-duration curve for the gaged basins and averages 70 percent (table 5). BFQ $_{95}$ ranges from 95 percent to 98 percent on the total flow duration curve for the gaged basins and averages 96 percent. Gaging stations east of the index station on West Wailuaiki Stream all have the same or higher rainfall-normalized low flow relative to the index station, indicating a higher ground-water contribution to these streams. The mapped springs that contribute to the low flows in these streams are listed in the "Comments" column in table 5. Gaging stations west of the index station generally have the same or lower rainfall-normalized low flows relative to the index station, indicating that the ground-water contribution to streamflow is relatively less to the west. The effects of minor upstream

diversions in this area also are apparent from the relatively lower low-flows measured at the gaging stations.

Estimation of Flow Characteristics of Ungaged Streams

Multiple linear-regression analysis is a standard technique used to develop equations for estimating streamflow statistics for ungaged sites (Koltun and Schwartz, 1986; Vogel and Kroll, 1990; Ludwig and Tasker, 1993; Ries and Friesz, 2000). In this study, a selected streamflow statistic (the duration discharges TFQ₅₀, TFQ₉₅, BFQ₅₀, or BFQ₉₅) at unregulated, gaged sites was the dependent variable and the quantified basin characteristics, rainfall rates, and surficial geologic units were the independent variables used as input in the regression analysis. The regression analysis statistically relates the dependent variable to the independent variables and results in an equation that can be used to estimate the selected streamflow statistic for a site where no streamflow data are available. The goal of the regression algorithm is to minimize the differences between the values of the dependent variable used in the analysis (observed values) and the corresponding values provided by the regression equations (estimated or fitted values).

Drainage-Basin Characteristics

For the drainage basin of each continuous-record gaged site and selected ungaged sites, the morphometric, geologic, and rainfall characteristics were quantified using Geographic Information System (GIS) techniques.

Morphometric characteristics

Drainage basins were delineated on the basis of 10-meter digital elevation model (DEM) data and the GIS program GISWeasel (U.S. Geological Survey, 2004). The drainage basins thus delineated (plate 1) were checked against existing manually determined drainage-basin boundaries for gaged sites to ensure the reliability of the computerized delineation routine. Minor adjustments were made to the boundaries of those drainage basin for which the computer-delineated drainage area differed from the manually determined drainage area by more than 5 percent (Chiu Yeung, U.S. Geological Survey, written commun., 2003). The GIS program Basinsoft (Harvey and Eash, 1996) was used to quantify basin characteristics considered in the regional regression analysis. Basinsoft uses GIS data layers of drainage divides, hydrography, and a digital elevation model to automatically and efficiently quantify 22 morphometric characteristics for each drainage basin selected. These characteristics are described in Appendix A. Computed basin characteristics for the gaged basins are listed in Table 6 and those for selected ungaged basins are listed in table 7.







Figure 6. Unadjusted and adjusted flow-duration curves of total streamflow and base flow at gaging station 5170 on East Wailuaiki Stream, northeast Maui, Hawaii.

The ungaged basins are denoted by an abbreviation for the stream name followed by a U, M, or L for upper, middle, or lower locations on the stream (plate 1). These three locations on each stream roughly correspond to the following settings: (U) upstream of the main diversion ditch at around 1,700- to 1,200-ft altitude; (M) roughly 600- to 500-ft altitude; and (L) near the stream mouth. These locations were chosen to help meet the goals of objective 3 of this study regarding the habitat available in the stream for native species.

Hydrologic and geologic characteristics

Average yearly total rainfall rates for each drainage basin were determined using GIS techniques by overlaying the drainage-basin boundaries on a map of mean annual rainfall isohyets from Giambelluca and others (1986) (fig. 8). The overlying GIS layers created polygons that were assigned rainfall rates, in inches per year, equal to the average of the two bounding isohyets of each polygon. Each polygon area, in square feet, was multiplied by the assigned rainfall rate, after converting the rainfall rate to feet per second, to determine a volume rainfall rate, in cubic feet per second, for each polygon. Final rainfall rates were then determined by summing the rainfall volumes in a basin. In areas of equal rainfall rates, larger drainage basins would have larger total rainfall volumes and smaller drainage basins would have smaller total rainfall volumes. Average annual rainfall rates in the gaged basins ranged from 1.3 to 76 ft³/s (table 6). Average annual rainfall rates in the ungaged basins ranged from 1.9 to 192 ft³/s (table 7).

Distributions of surficial geologic units in each gaged drainage basin were determined from a digital geology map by Sherrod and others (2003) (fig. 9). Three geologic units were considered, the Honomanu Basalt, the Kula Volcanics, and the Hana Volcanics (table 6). The percentage of each unit in the basin was determined by dividing the area of each unit in the basin by the total basin area and multiplying the result by 100. Only two gaged basins contained surficial exposures of Honomanu Basalt, both having less than 1 percent of the total area of each basin, so only Kula and Hana Volcanics were included in the following analysis.

Development of regression equations

Linear equations generated by use of regression analysis have the general form

$$Y_{i} = b_{0} + b_{1}X_{1} + b_{2}X_{2} + \dots + b_{n}X_{n} + \varepsilon_{i}$$
(1)

where Y_i is the estimate of the dependent variable for site i, X_i to X_n are the *n* independent variables, b_0 to b_n are the *n*+1 regression model coefficients, and e_i is the residual error (difference between the observed and estimated value of the dependent variable for site *i*). Regression analysis results must be evaluated to make sure that the following assumptions are met: (1) equation 1 adequately describes the relation between the dependent and independent variables, (2) the mean of e_i is zero, (3) the variance of e_i is constant and independent of the value of X_n , (4) values of e_i are normally distributed, (5) values of e_i are independent of each other, (6) all independent variables selected are statistically significant at the 5-percent level, (7) independent variables are not correlated, and (8) the



Figure 7. Correlation of flows at selected gaging stations in northeast Maui, Hawaii with flow at index station 5180 (West Wailuaiki Stream) based on discharge of equal percentage duration and unadjusted and adjusted duration curves of total streamflow and base flow.





Figure 7. Correlation of flows at selected gaging stations in northeast Maui, Hawaii with flow at index station 5180 (West Wailuaiki Stream) based on discharge of equal percentage duration and unadjusted and adjusted duration curves of total streamflow and base flow–Continued.



Figure 7. Correlation of flows at selected gaging stations in northeast Maui, Hawaii with flow at index station 5180 (West Wailuaiki Stream) based on discharge of equal percentage duration and unadjusted and adjusted duration curves of total streamflow and base flow–Continued.



Figure 7. Correlation of flows at selected gaging stations in northeast Maui, Hawaii with flow at index station 5180 (West Wailuaiki Stream) based on discharge of equal percentage duration and unadjusted and adjusted duration curves of total streamflow and base flow–Continued.



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Figure 7. Correlation of flows at selected gaging stations in northeast Maui, Hawaii with flow at index station 5180 (West Wailuaiki Stream) based on discharge of equal percentage duration and unadjusted and adjusted duration curves of total streamflow and base flow–Continued.

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Figure 7. Correlation of flows at selected gaging stations in northeast Maui, Hawaii with flow at index station 5180 (West Wailuaiki Stream) based on discharge of equal percentage duration and unadjusted and adjusted duration curves of total streamflow and base flow–Continued.



Figure 7. Correlation of flows at selected gaging stations in northeast Maui, Hawaii with flow at index station 5180 (West Wailuaiki Stream) based on discharge of equal percentage duration and unadjusted and adjusted duration curves of total streamflow and base flow–Continued.



Figure 7. Correlation of flows at selected gaging stations in northeast Maui, Hawaii with flow at index station 5180 (West Wailuaiki Stream) based on discharge of equal percentage duration and unadjusted and adjusted duration curves of total streamflow and base flow–Continued.



Figure 7. Correlation of flows at selected gaging stations in northeast Maui, Hawaii with flow at index station 5180 (West Wailuaiki Stream) based on discharge of equal percentage duration and unadjusted and adjusted duration curves of total streamflow and base flow–Continued.

Table 5. Comparisons of base-flow to total-flow duration statistics and streamflow characteristics at selected stations to those at index gaging station 5180 on West Wailuaiki Stream, northeast Maui, Hawaii

[Qxx is the xx-percent flow duration of streamflow; +, combined with record from indicated station]

	Equivalent tota discharge,	l flow duration- , in percent	Rainfall-normalize index s	d flow relative to tation	
Gaging-station number	0 ₅₀ base flow	0 ₉₅ base flow	Low flow relative to index station	High flow relative to index station	Comments
5070	62	-	Lower	Same	Intermittent flow; dry 30 percent of time; Q ₉₅ not comparable
5080	69	97	Higher	Same	
5090+5080	99	96	Higher	Same	Includes Big and Hanawi Springs
5100	72	98	Same	Lower	
5110+5100	71	97	Higher	Same	Includes Pali Spring
5130	56	95	Higher	Same	Includes unnamed springs
5140	09	95	Higher	Same	Includes numerous unnamed springs
5150	67	97	Higher	Same	
5160	71	98	Higher	Same	
5170	73	97	Higher	Same	
5180	72	97	1	1	Index station
5190	73	98	Same	Same	
5200	72	97	Higher	Same	
5210+5200+5190	71	97	Same	Same	Combination of several gages
5240	78	97	Lower	Same	Lower low flows due to minor diversion at 3,000 ft altitude
5270	75	98	Lower	Same	Lower low flows due to minor diversion at 3,000 ft altitude
5311+5310	76	97	Lower	Same	Combination of two gages
5360+5311+5310+5350	75	98	Same	Lower	Combination of several gages
5420	75	98	Same	Lower	
5430	72	98	Lower	Same	Lower low flows due to diversions at 3,000 ft and 4,300 ft altitude
5440	74	97	Same	Same	Diversion effects not apparent
5450	LL	96	Same	Lower	Lower low flows due to diversions at 3,000 ft and 4,300 ft altitude
5528	74	97	Lower	Same	4,487 ft altitude; intermittent flow; dry 1 percent of time
5540	74	95	Same	Same	
5550	78	96	Lower	Same	Lower low flows due to diversions at 3,000 ft and 4,300 ft altitude
5560	74	98	Same	Same	Gage operated before minor upstream diversions were active
5570	76	97	Same	Lower	
Average	70	96			

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[Abbreviation definitions and units are described in appendix A: gaging stations in **bold italics** were used to develop the final regression equations for estimating flow-duration statistics; precipitation data modified from Giambelluca and others (1986); ft³/s, cubic feet per second; geology data from Sherrod and others (2003); ND, not determined]

	ВР	BR	BS	BW	CCM	CR	DA	H	MAXELEV	MCL	MCS	MCSP	MCSR	MINELEV	ßB	RN
	10.94	4283	1038	0.39	0.17	2.24	1.89	0.32	5219	4.6	887	0.16	0.94	936	10	25400
	14.90	6775	1557	0.57	0.18	2.21	3.62	0.34	8093	6.0	1036	0.19	0.95	1318	6	38798
	17.86	7587	1433	0.74	0.17	2.22	5.15	0.37	8093	7.0	971	0.23	1.00	506	7	45008
	9.19	2968	854	0.13	0.12	3.12	0.69	0.17	4325	4.1	069	0.16	0.76	1357	33	24894
	11.55	3787	1099	0.17	0.14	3.35	0.95	0.20	4325	4.9	745	0.18	0.90	538	25	27527
	2.42	881	1008	0.05	0.09	2.27	0.09	0.21	1587	1.1	922	0.04	0.63	706	24	10264
	4.39	1741	1023	0.11	0.11	2.10	0.35	0.21	2373	2.2	808	0.08	0.69	632	22	15330
	4.70	1413	885	0.15	0.16	1.83	0.53	0.23	2771	2.4	567	0.10	0.69	1358	19	8973
	15.83	6993	1652	0.53	0.15	2.27	3.88	0.30	8372	6.4	956	0.21	0.86	1379	11	46589
	16.02	7186	1527	0.46	0.16	2.41	3.51	0.28	8527	6.2	946	0.20	0.81	1341	13	44113
	17.08	7518	1761	0.44	0.16	2.51	3.69	0.26	8857	6.5	941	0.21	0.78	1339	15	47174
	17.19	6706	1797	0.26	0.17	3.54	1.88	0.22	8055	6.1	741	0.22	0.85	1349	21	38688
	5.92	1424	1162	0.19	0.17	2.34	0.51	0.30	2733	2.5	595	0.10	0.92	1309	11	8531
	19.51	7467	1736	0.30	0.18	3.32	2.75	0.20	8055	7.3	749	0.27	0.80	588	24	41635
	14.47	5420	1619	0.44	0.18	2.55	2.57	0.31	8331	5.4	965	0.18	0.94	2911	10	30319
	18.48	6594	1670	0.41	0.18	2.97	3.07	0.27	8331	6.9	915	0.23	0.93	1737	14	36338
	4.47	1645	1073	0.13	0.15	2.45	0.27	0.29	5979	1.7	952	0.06	0.87	4334	12	10986
	15.14	4468	1577	0.17	0.15	3.90	1.20	0.17	5979	6.0	714	0.22	0.84	1511	33	29733
	2.44	773	1793	0.18	0.16	1.82	0.14	0.55	3645	0.8	802	0.03	1.07	2872	Э	4835
_	7.98	2705	1639	0.14	0.13	3.24	0.48	0.22	5613	3.5	754	0.13	0.99	2908	20	21675
	5.23	1610	1779	0.18	0.17	2.22	0.44	0.31	4470	2.1	844	0.07	0.85	2860	10	9756
	14.99	4304	1795	0.34	0.17	2.78	2.31	0.26	5620	6.4	625	0.26	0.95	1316	15	25914
- `	12.49	4864	1422	0.46	0.18	2.24	2.48	0.33	9329	5.0	849	0.17	0.93	4465	6	27800
	17.39	6340	1399	0.46	0.17	2.77	3.13	0.30	9329	7.1	794	0.25	1.04	2989	12	37028
	24.81	8034	1516	0.38	0.17	3.56	3.87	0.22	9329	9.6	722	0.37	0.98	1295	21	48426
	25.29	8195	1525	0.38	0.17	3.61	3.90	0.22	9329	10.2	717	0.38	0.99	1134	22	49689
	5.38	1270	1551	0.20	0.16	2.22	0.47	0.33	2505	2.2	491	0.10	0.92	1235	6	7836
	7.65	1835	1572	0.19	0.15	2.70	0.64	0.26	3140	3.1	542	0.13	0.91	1305	15	12224
	3.74	784	1451	0.15	0.12	2.16	0.24	0.34	2014	1.5	555	0.06	0.93	1230	6	6313
	16.92	5469	1185	0.38	0.16	2.51	3.61	0.23	6686	7.8	658	0.30	0.82	1217	19	34724
	11.95	3712	1165	0.39	0.20	2.18	2.39	0.29	4975	5.9	571	0.25	0.97	1263	12	18981

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Table 6.

[Abbreviation definitions and units are described in appendix A; gaging stations in bold italics were used to develop the final regression equations for estimating flow-
duration statistics; precipitation data modified from Giambelluca and others (1986); ft ³ /s, cubic feet per second; geology data from Sherrod and others (2003); ND, not
determined]

3aging-						Annual hasin	Ratio of gage to index	Percent cover	rage by surficial (geologic unit
station number	RR	SD	SF	SR	TSL	rainfall (ft³/s)	station annual basin rainfall	Honomanu Basalt	Kula Volcanics	Hana Volcanic:
5070	392	5.93	12.8	0.85	11.2	35	0.71	0	0	100
5080	455	5.73	11.1	0.67	20.7	51	1.03	0	11	89
0609	425	5.93	9.5	0.68	30.6	76	1.55	0	7	92
001	323	8.39	42.2	0.81	5.8	13	0.26	0	0	100
5110	328	7.27	31.2	0.68	6.9	17	0.34	1	0	66
5130	365	11.65	30.5	0.91	1.1	1.3	0.03	0	0	100
5140	397	8.81	28.5	0.79	3.1	5.5	0.11	0	0	100
5150	301	6.35	23.8	0.64	3.3	9.7	0.20	0	6	91
0919	442	6.66	14.0	0.58	25.9	50	1.01	0	58	42
5170	449	6.14	16.4	0.62	21.6	48	0.97	0	100	0
5180	440	6.28	18.8	0.53	23.1	49	1.00	0	93	L
0619	390	5.77	27	0.41	10.8	24	0.81	0	63	37
5200	241	5.99	14.0	0.51	3.0	9.1	0.19	0	98	2
5210	383	5.58	30.6	0.43	15.3	38	0.77	0	74	26
5240	375	5.59	13.1	0.60	14.4	33	0.66	0	100	0
5270	357	5.51	18.0	0.55	16.9	43	0.86	0	100	0
5311	368	6.68	14.9	0.89	1.8	3.3	0.07	0	100	0
360	295	6.65	42.5	0.45	8.0	20	0.41	0	100	0
6420	317	6.25	4.2	0.45	0.9	2.8	0.06	0	100	0
5430	339	8.01	25.4	0.46	3.9	8.0	0.16	0	100	0
5440	308	6.06	13.1	0.47	2.7	7.6	0.15	0	100	0
5450	287	6.02	19.6	0.35	13.9	40	0.82	0	100	0
5528	389	5.72	11.9	0.60	14.1	15	0.30	0	100	0
540	365	5.84	14.7	0.57	18.3	26	0.52	0	100	0
5550	324	6.03	26.4	0.48	23.4	38	0.78	0	100	0
5560	324	6.06	27.4	0.47	23.7	39	0.79	0	100	0
5570	236	6.17	11.7	0.32	2.9	6.8	0.14	0	100	0
5650	240	6.66	18.6	0.35	4.2	9.8	ND	ND	ND	ND
9660	210	8.05	11.0	0.38	1.9	3.1	ND	ND	ND	ND
5700	323	6.35	24.7	0.56	22.9	50	ND	ND	ND	ND
0000	211		[[0						

Table 7. Watershed characteristics for selected ungaged stream drainage basins, northeast Maui, Hawaii

[Basins and abbreviations listed from east to west; L, lower; M, middle; U, upper; watershed characteristic abbreviation definitions and units are described in appendix A; precipitation data

Stream location	BL	BP	BR	BS	BW	CCM	CB	DA	EB	MAXELEV	WCL	MCS	MCSP
Hanawi lower (HwL)	9.02	19.24	8040	1488	0.59	0.17	2.34	5.36	0.29	8093	7.6	947	0.25
Kapaula lower (KL)	5.79	12.46	4151	1186	0.19	0.14	3.40	1.07	0.20	4325	5.3	732	0.20
Waiaaka lower (WaL)	1.85	3.62	1482	1485	0.08	0.10	2.71	0.14	0.23	1546	1.44	1119	0.04
Paakea lower (PaL)	3.10	0.03	2296	1303	0.17	0.14	2.43	0.53	0.27	2372	2.74	866	0.09
Paakea upper (PaU)	1.90	3.61	1059	801	0.12	0.10	2.14	0.23	0.28	2372	1.63	642	0.06
Waiohue lower (WeL)	3.84	7.78	2738	1270	0.24	0.18	2.30	0.91	0.28	2766	3.63	791	0.13
Waiohue middle (WeM)	3.49	6.70	2269	1078	0.21	0.17	2.19	0.75	0.28	2766	3.23	671	0.13
Puakaa middle (PuM)	3.15	6.80	1889	1338	0.13	0.12	3.06	0.39	0.23	2491	3.00	655	0.12
Puakaa upper (PuU)	2.06	4.65	1177	1113	0.13	0.12	2.52	0.27	0.29	2491	2.08	602	0.09
Kopiliula lower (KpL)	8.29	19.80	8343	1699	0.57	0.16	2.58	4.70	0.30	8372	8.0	837	0.28
Kopiliula middle (KpM)	8.95	18.88	7800	1688	0.47	0.15	2.60	4.19	0.26	8372	7.5	855	0.26
East Wailuaiki lower (EWL)	8.92	20.94	8498	1630	0.44	0.17	2.98	3.94	0.25	8527	8.0	812	0.28
East Wailuaiki middle (EWM)	8.35	19.77	8096	1582	0.46	0.17	2.86	3.80	0.26	8527	7.5	843	0.26
West Wailuaiki lower (WWL)	9.46	21.77	8822	1833	0.43	0.16	3.04	4.07	0.24	8857	8.3	819	0.29
West Wailuaiki middle (WWM)	8.97	20.58	8385	1796	0.44	0.16	2.92	3.97	0.25	8857	7.8	839	0.27
Wailuanui lower (WL)	9.17	22.69	8017	1772	0.35	0.19	3.56	3.24	0.22	8055	8.2	LLL	0.29
Waiokomilo lower (WoL)	10.44	23.09	6459	1405	0.25	0.18	4.01	2.63	0.18	6483	9.2	686	0.35
Waiokomilo middle (WoM)	8.41	18.94	5977	1478	0.25	0.17	3.68	2.11	0.20	6483	7.5	828	0.26
Waiokomilo upper (WoU)	7.42	15.75	5108	1314	0.20	0.16	3.66	1.48	0.19	6483	6.2	828	0.22
Ohia lower (OL)	2.10	2.92	388	821	0.11	0.15	1.75	0.22	0.25	413	1.2	358	0.06
Palauhulu lower (PhL)	8.62	20.48	5745	1510	0.32	0.17	3.49	2.74	0.22	5816	8.2	666	0.32
Palauhulu middle (PhM)	7.43	16.57	5299	1587	0.33	0.17	3.01	2.41	0.24	5816	6.7	772	0.24
Kano upper (KoU)	4.99	11.09	3792	1251	0.20	0.13	3.17	0.97	0.22	5816	4.3	827	0.15
Hauoli Wahine upper (HWU)	1.34	3.94	1052	1280	0.27	0.15	1.85	0.36	0.51	3049	1.4	749	0.05
Piinaau lower (PiL)	16.45	36.91	976	1984	1.07	0.22	2.48	17.58	0.29	10011	14.9	620	0.60
Piinaau middle (PiM)	11.92	33.25	9536	1993	1.40	0.21	2.30	16.63	0.39	10011	13.3	652	0.52
Piinaau upper (PiU)	10.47	29.95	8689	2019	1.43	0.21	2.19	14.96	0.42	10011	11.7	664 (00	0.45
Nuaailua lower (NL)	4.25	9.18	2389	1967	0.28	0.23	2.38	1.19	0.29	2409	3.5 2.5	698 212	0.13
Nuaailua middle (NMI)	5.07	10.0	1891	1/70	0.10	0.17	2.00	0.48	CZ.0	2409	0.7	01/	0.10
Nuaailua upper (NU)	1.18	7.80	500	1036	0.10	0.11	2.30	0.12	0.33	2409	1.0	609	0.04
HONOMANU IOWET (HNL)	9.80	16.07	4000	C117	7070	0.19	66.7 00 c	60.C	07.0	1000	0.0	040	67.0
Dunding Inducte (DILVI) Dunalari Jourar (DIL)	0.01 2 80	C0.12	7531	1001	0.49 0.00	0.16	02.7 02.0	40.4 1 8 0	17.0	1000	0.0	900 648	0.20
Punalan middle (PIM)	3.56	7 89	2053	1748	0.22	0.16	2.5.2	0.78	0.28	2566) () (604	0.13
Haipuaena lower (HaL)	9.44	21.05	5709	1758	0.17	0.15	4.70	1.60	0.15	5979	8.4	652	0.33
Haipuaena middle lower (HaML)	8.80	19.65	5509	1771	0.17	0.15	4.59	1.46	0.16	5979	7.8	647	0.31
Haipuaena middle upper (HaMU)	8.04	17.87	5042	1744	0.17	0.15	4.31	1.37	0.16	5979	7.2	675	0.28
Puohokamoa lower (PL)	8.78	19.69	5599	1801	0.36	0.17	3.14	3.12	0.23	5613	8.5	594	0.35
Puohokamoa middle lower (PML)	8.20	18.44	5100	1752	0.36	0.17	3.02	2.97	0.24	5613	7.9	592	0.33
Puohokamoa middle upper (PMU)	7.56	17.01	4700	1799	0.35	0.17	2.95	2.65	0.24	5613	7.3	617	0.30
Wahinepee lower (WpL)	1.56	4.46	1247	1642	0.26	0.19	1.98	0.40	0.46	1320	1.7	619	0.07
Wahinepee middle (WpM)	1.06	3.14	746	1353	0.26	0.17	1.69	0.28	0.56	1320	1.2	603	0.05
Waikamoi lower (WiL)	12.23	29.77	9205	1650	0.39	0.17	3.86	4.74	0.20	9329	12.0	686	0.46
Waikamoi middle lower (WiML)	11.60	28.29	8834	1623	0.40	0.17	3.69	4.67	0.21	9329	11.4	701	0.43
Walkamol middle upper (WiMU)	11.22	27.54	8594	1609	0.41	0.17	3.62	4.61	0.22	9329	1.11	/03	0.42
Kolea lower (KaL)	5.08	1.02	1221	1688	0.20	0.15 21 0	10.2	0.62	0.29	1846 1846	1.7	627	0.11
Kolea midule (Naivi)	CC.7	0.00	1701	1000	U.2U	0.1J	DC.2	1C.U	7C.U	1040	2.2	cno	U.U
Table 7. Watershed characteristics for selected ungaged stream drainage basins, northeast Maui, Hawaii—Continued

[Basins and abbreviations listed from east to west; L, lower; M, middle; U, upper; watershed characteristic abbreviation definitions and units are described in appendix A; precipitation data modified from Giambelluca and others (1986); ft^3/s , cubic feet per second; values in **bold** fall outside the range shown in table 8 used to develop the regression equations]

[Annual hooin
Stream location	MCSR	MINELEV	RB	RN	RR	SD	SF	SR	TSL	rainfall (ft³/s)
Hanawi lower (HwL) Kanaula lower (KI)	0.84 0.91	53 174	12 25	47070 28937	418 333	5.85 6.97	15.20 31.30	0.64	31.4 7.5	79
Waiaaka lower (WaL)	0.78	64	19	15036	410	10.15	24.16	0.75	1.44	1.9
Paakea lower (PaL)	0.89	76	14	16286	366	7.09	18.12	0.67	3.76	7.9
Paakea upper (PaU)	0.86	1313	12	11139	293	10.52	15.84	0.80	2.39	3.8
Waiohue lower (WeL)	0.94	28	13	15317	352	5.59	16.21	0.62	5.10	15
Waiohue middle (WeM)	0.93	497	13	13356	339	5.89	16.26	0.62	4.40	13
Puakaa middle (PuM)	0.95	602	20	15238	278	8.07	25.28	0.49	3.17	6.4
Puakaa upper (PuU)	1.01	1314	12	9734	253	8.27	15.71	0.54	2.24	4.7
Kopiliula lower (KL)	0.97	29	11	54003	421	6.47	14.61	0.49	30.4	62
Kopiliula middle (KM)	0.84	572	15	50530	413	6.48	19.11	0.51	27.1	54
East Wailuaiki lower (EWL)	0.89	29	16	50650	406	5.96	20.21	0.50	23.5	54
East Wailuaiki middle (EWM)	0.90	431 31	4 i	48633	409	6.01	18.35	0.53	22.8	52
West Wailuaiki lower (WWL)	0.87		17	22/22	C04	0.09 6 15	21.98	0.40	24.0 8.7	40 4 6
West wailualki iliiuule (W W M) Wailiianii loiiar (M/I)	0.07	4/7 38	010	01049 11455	407 353	0.1.0 7.1.7	22.02 25.06	0.47	24.4 16.8	C <i>V V</i>
Walluallui 10WEI (WL) Wajokomilo lower (WoI)	0.88	00 74	225	36511	080 080	295	41 47	040	14.0	+ U7
Waiokomilo middle (WoM)	0.80	205	26	34370	316	575	33.53	0.56	12.1	34
Waiokomilo unner (WoU)	0.83	1375	20	31821	324	6.23	37.28	0.63	9.2	24
Ohia lower (OL)	0.57	25	16	2582	133	6.66	19.85	0.44	1.5	2.2
Palauhulu lower (PhL)	0.95	71	21	33102	280	5.76	27.11	0.44	15.8	48
Palauhulu middle (PhM)	0.90	517	18	31149	320	5.88	22.89	0.49	14.2	44
Kano upper (KoU)	0.87	2024	20	28683	342	7.56	25.61	0.66	7.3	18
Hauoli Ŵahine upper (HWU)	1.05	1997	4	6995	267	6.65	4.93	0.59	2.4	7.5
Piinaau lower (PiL)	0.91	35	12	46178	270	4.63	15.38	0.31	81.4	192
Piinaau middle (PiM)	1.12	475	7	44633	287	4.68	8.55	0.33	77.9	181
Piinaau upper (PiU)	1.12	1322	9	41236	290	4.75	7.32	0.33	71.0	152
Nuaailua lower (NL)	0.83	20	12	10462	260	4.38	15.22	0.36	5.2	16_{-}
Nuaailua middle (NM)	0.84	518	16	10927	290	5.78	19.86	0.41	2 . 7 . 7 .	7.1
Nuaailua upper (NU)	0.88	1/56	9 4	5752	234	8.81	11.81	0.59	0.1.0	270
HONOMANU IOWET (HNL)	1.91	17	C1	44700 7777	04/ 74/	0.00 27 7	17.00	0.4.0	1.12	0/
Dunolatid Induic (DIIN) Dunolan louver (DII)	16.0 0.80	35	1 1	15807	400 400	10.0	18.06	10.0	7.47 7.4	c 0 C1
Punalau modele (PIM)	0.80	513	1 [13070	260	6 37	16.26	0.35	5.0	12
Haipuaena lower (HaL)	06.0	270	44	37942	271	6.65	55.84	0.37	10.6	25
Haipuaena middle lower (HaML)	0.89	470	42	36866	280	6.69	52.98	0.37	9.8	24
Haipuaena middle upper (HaMU)	0.89	937	37	33587	282	6.66	47.31	0.39	9.1	23
Puohokamoa lower (PL)	0.96	14	19	33088	284	5.91	24.66	0.33	18.5	50
Puohokamoa middle lower (PML)	0.97	513	18	30273	277	5.94	22.64	0.34	17.6	49
Puohokamoa middle upper (PMU)	0.97	913	17	28077	276	5.97	21.55	0.34	15.8	45
Wahinepee lower (WeL)	1.09	73	S	6702	280	5.37	6.07	0.41	2.2	4.1
Wahinepee middle (WeM)	1.09	574	ς Ω	4335	237	5.81	4.06	0.45	1.6	3.0
Waikamoi lower (WiL)	0.98	124	25	55302	309	6.01	31.55	0.42	28.5	50
Waikamoi middle lower (WiML)	0.98	495	53	52895	312	5.99	28.83	0.43	27.9	46
Waikamoi middle upper (WiMU)	06.0	(30	17	01440	512	66.C	07.17	0.44	0.12	49 0 1
Kolea lower (KaL)	0.89	505	17	06611	107	0.04 777	12.20	0.54	4.c	0.7
Kolea midule (Nami)	U.01	C7C	IU	9724	177	0./0	12.72	0C.U	5. 4	0.1



Figure 8. Mean annual rainfall, east Maui, Hawaii (modified from Giambelluca and others, 1986).

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Figure 9. Generalized surficial geology, northeast Maui, Hawaii (modified from Sherrod and others, 2003).

signs and magnitudes of the coefficients determined for the significant, independent variables are hydrologically reasonable (Fontaine and others, 1992; Iman and Conover, 1983, p. 367).

Streamflow and basin characteristics used in hydrologic regression usually are log-normally distributed; therefore, the variables must be transformed to logarithms to satisfy regression assumption 2. The dependent and independent variables were transformed using log-base 10 units. Where values of percent coverage of Kula Volcanics were 0 percent, they were set to 1 percent before log transformation.

Ordinary-Least-Squares (OLS) regression analysis was used to develop the equations presented in this report. Because streamflow data are correlated spatially and in time, assumption 5 for use of regression is not strictly satisfied. A theoretically more appropriate method, Generalized-Least Squares (GLS) regression, was developed by Tasker and Stedinger (1989) to allow weighting to compensate for length-of-record and spatial correlations. However, Vogel and Kroll (1990) found that the equation parameters $(b_0 \text{ to } b_n)$ were nearly identical when either OLS or GLS was used to develop the equation, even though OLS does not correct for length-ofrecord or spatial differences. Ries and Friesz (2000) used Weighted-Least-Squares (WLS) regression analysis to predict duration flows because WLS can compensate for length-ofrecord differences. They found that equations developed using WLS and GLS methods were nearly identical. Because the streamflow statistics used in the development of the equations in this report were adjusted to equivalent lengths of record, and the spatial correlations between gaged basins in the study area are relatively insignificant, OLS regression analysis was determined to be the most appropriate for this study.

Regression assumption 7 was addressed by removing independent variables having high correlation (> 90 percent) with several other independent variables in the analysis. For example, Drainage Area (DA) was highly correlated with Total Stream Length (98 percent), Basin Relief (94 percent), Rainfall (93 percent), and Basin Width (91 percent), and therefore DA was removed from further analysis. Other basin characteristics removed from further consideration because of high correlations were Basin Perimeter (BP), Rotundity of Basin (RB), and Main Channel Length (MCL).

A variable-selection algorithm was applied to the remaining independent variables to aid in determining which combination of independent variables provides the best estimates of the dependent variables. The algorithm used was a leapsand-bounds implementation with Mallow's Cp as the selection criterion (Insightful Corporation, 2002). Subsets of the independent variables were evaluated and ranked according to the lowest value of Mallow's Cp for each subset of 1, 2, 3 ... n independent variables. The subsets of 1, 2, 3, and 4 independent variables having the lowest Mallow's Cp were then further analyzed using OLS regression to select a final model for each statistic.

During equation development, several gaging stations were eliminated from the analysis because (1) the sites were known outliers on the basis of observed hydrologic differences or (2) plots of Cook's distance (Draper and Smith, 1998) indicated that the flow statistics for a particular gaging station were statistically biasing the results. Gaging stations eliminated for hydrologic reasons prior to the Mallow's *Cp* analysis were stations 5070, 5311, and 5528 (intermittent streams); station 5090 (has anomalous spring input); and stations 5110, 5130, 5140, and 5210 (regulated streams). Gaging stations 5240 and 5420 were eliminated because of high Cook's distance values determined during the analyses. The final number of gaging stations from which data were used to develop the equations was 17 (n = 17).

The final models were selected on the basis of the following parameters: (1) Mallow's *Cp* statistic; (2) R^2 , the proportion of total variation about the mean explained by the regression; (3) *SE*_r, the average standard error of the estimates; and (4) Pr (>|t|), the probability of significance for an independent variable in the regression. Pr (>|t|) had to be lower than 5 percent for each independent variable used in the regression model for that independent variable to be included.

The retransformed regression equations are biased because they predict the median rather than the mean response of the dependent variable. In the case of streamflow data, the median tends to be lower than the mean. Duan's (1983) "smearing estimate", the mean error of the retransformed residuals, was used as the bias-correction factor (BCF) to adjust the retransformed b_0 coefficient. This BCF is advantageous in that it does not require normally distributed regression residuals and is simple to calculate (Ries and Friesz, 2000).

Accuracy and limitations of the regression equations

Regression equations for predicting duration discharges TFQ₅₀, TFQ₉₅, BFQ₅₀, and BFQ₉₅ at unregulated sites were developed using OLS regression as described above. The equations, along with several measures of model adequacy and the BCF for each equation, are presented in table 8. The measures of model adequacy include (1) the coefficient of determination (R^2) ; (2) the average standard error of estimate $(SE_{,})$, in percent; (3) the average standard error of prediction (SE_{\perp}) , in percent; and (4) the median absolute deviation (MAD), in percent. The R^2 is a measure of the proportion of the variation in the dependent variable that is explained by the independent variables. The SE is a measure of the average precision with which the regression equations estimate the streamflow statistics for stations used in the analyses, whereas the SE_{n} indicates the average precision with which the equation can be used to estimate streamflow statistics for ungaged sites with basin characteristics similar to those for the stations used in the regression analyses. About 68 percent of streamflows estimated using the regression equations would have errors within the noted average standard errors. Half of the regression-equation estimates for stations used in the analyses had absolute

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 Table 8.
 Summary of regression equations developed for estimating selected flow-duration statistics of northeast Maui

 streams, Hawaii.
 Streams, Hawaii.

[Statistic: TF is total flow; BF is base flow, Q_{xx} is the xx-percent duration flow; Statistic estimator: Rainfall is area-weighted rainfall rate (cubic feet per second); MAXELEV is maximum drainage-basin elevation (feet); ER is elongation ratio (dimensionless); \mathbf{R}^2 : Coefficient of determination (percent); \mathbf{SE}_p ; Average standard error of estimate and prediction (percent); MAD: Median absolute deviation (percent); BCF: Bias correction factor; n = 17 for all equations]

Statistic	Statistic estimator	R ²	SE,	SEp	MAD	BCF
TFQ ₅₀	3,184(Rainfall) ^{1.338} (MAXELEV) ^{-1.366} (ER) ^{-0.946}	94.9	15.3	20.9	12	1.009
BFQ ₅₀	25,384(Rainfall) ^{1.525} (MAXELEV) ^{-1.735} (ER) ^{-0.937}	91.0	22.5	30.5	17	1.019
TFQ ₉₅	56,267(Rainfall) ^{1.478} (MAXELEV) ^{-1.750}	76.6	38.1	50.3	21	1.059
BFQ ₉₅	409,732(Rainfall) ^{1.620} (MAXELEV) ^{-2.054}	75.3	43.0	56.5	28	1.073

Range used in analysis

	<u>Minimum</u>	Mean	<u>Maximum</u>
Rainfall	6.8	28	51
MAXELEV	2,505	6,602	9,329
ER	0.17	0.26	0.34

errors, in percent, that were greater than the MAD, and half of them were less than the MAD. Scatter plots comparing observed and estimated flow-duration statistics show the fit of the regression equations for sites used in determining the equations (fig. 10).

The rainfall rates for stations used in developing the regression equations ranged from 51 to 6.8 ft³/s and averaged 28 ft³/s (table 8). For all four equations, rainfall has a positive coefficient, thus higher rainfall rates lead to higher estimated flows. The value of MAXELEV ranged from 9,329 to 2,505 ft and averaged 6,602. For all four equations, MAXELEV has a negative coefficient, thus higher elevations lead to lower estimated flows. The value of ER ranged from 0.34 to 0.17 and averaged 0.26. Only the median flow equations. Therefore, short and wide drainage basins (higher ER), will have lower median flows. Estimates of flow-duration statistics derived by using the regression equations for drainage basins with characteristics outside the ranges (table 8) used in equation development could be subject to substantial errors.

Values of estimated streamflow statistics and several measures of estimation error for each continuous-record gaging station in the study area and at four gaged stations (5650, 5660, 5700, and 5770) west of the study area can be compared to the observed statistics to evaluate the performance of the equations (table 9). Statistics were generated for the gaged stations west of the study area to provide an additional indication of the accuracy of the equations for sites where measured streamflows are available. Estimated streamflow, in cubic feet per second, is the value determined by applying the regression equation. Prediction intervals (90 percent lower confidence limit [90% LCL], 90 percent upper confidence limit [90% UCL]), in cubic feet per second, indicate the uncertainty inherent in the use of the regression equations. Assurance is 90 percent that the true value of the streamflow statistic will be within the prediction interval. Standard error, in percent, is a measure of the precision with which the equation estimates the streamflow statistic. Measured flow, in cubic feet per second, is the observed streamflow statistic determined from the continuous record and adjusted to the long-term index station. Relative error, in percent, is calculated from [100(estimated flow – measured flow)/measured flow], and where available, indicates how well the estimated value matches the measured value. The value of MAD for each equation (table 8) is the average of relative errors for all stations used in the development of the regression equations.

At the stations used in the development of the equations, the equations tend to predict more accurately the higher flow statistics, TFQ₅₀ and BFQ₅₀, than the lower flow statistics, TFQ_{05} and BFQ_{05} . At the outlier stations eliminated from the regression analysis, the accuracy of the equation estimates generally is poor, indicating that a factor not considered in the regression analysis is probably affecting the streamflow statistics at these stations (fig. 11). The most likely factor is variable subsurface geology that can control where streams are intermittent and where springs discharge high flow to streams. At the stations on intermittent streams (stations 5070 and 5528), the equations overestimate the flow-duration statistics. At stations on streams with notable springs (stations 5090, 5110, 5130, 5140), the equations underestimate the flow-duration statistics, especially at lower flows. At two (stations 5110 and 5210) of the three regulated stations where streamflow statistics were determined on the basis of combined records of the upstream and downstream gaging stations, the equations provide a reasonably accurate estimate of the flow statistics at the downstream stations. At stations 5110 and 5210, the rela-



Streamflow statistics estimated using regression equations, lower and upper confidence intervals, standard errors, measured flows, and relative errors for continuous-record sites, northeast Maui, Hawaii. Table 9.

duration of base flow; estimated and measured flow and confidence intervals are in cubic feet per second; 90% LCL and 90% UCL is 90-percent lower and upper confidence level; Standard error is in percent; Relative error is the percent difference between the measured statistic and the estimated statistic; measured flows in bold italics are within the 90-percent confidence interval of the computed [Gaging stations in **bold** were used to develop the final regression equations for estimating low-flow statistics; TFQ_{xx} is the xx-percent flow duration of total streamflow; BFQ_{xx} is the xx-percent flow

estimate; NA, not applicable]					
Gaging-station number	Statistic	TFQ ₅₀	BFO ₅₀	TFO ₃₅	BFQ ₉₅
5070	Estimated flow	9.1	6.0	3.4	3.0
	90% LCL	8.0	4.9	2.4	2.2
	90% UCL	10	7.3	4.6	4.3
	Standard error	7.8	11.4	17.6	19.7
	Measured flow	2.2	1.3	0.00	0.00
	Relative error	310	360	NA	NA
5080	Estimated flow	7.7	4.6	2.7	2.2
	90% LCL	6.8	3.8	2.1	1.7
	90% UCL	8.8	5.6	3.5	3.0
	Standard error	7.5	11.1	14.2	15.9
	Measured flow	7.1	4.6	2.4	2.2
	Relative error	8	0	13	0
5090 + 5080	Estimated flow	12	7.9	4.9	4.3
	90% LCL	10	6.1	3.4	2.8
	90% UCL	15	10	7.1	6.6
	Standard error	10.3	15.2	21.5	24.1
	Measured flow	28	24	19	19
	Relative error	-57	-67	-74	-77
5100	Estimated flow	5.7	3.3	1.1	0.89
	90% LCL	4.8	2.6	0.88	0.71
	90% UCL	6.6	4.1	1.3	1.1
	Standard error	8.9	13.0	11.3	12.7
	Measured flow	4.9	2.8	1.1	0.90
	Relative error	16	18	0	-1
5110 + 5100	Estimated flow	6.7	4.1	1.5	1.3
	90% LCL	5.9	3.4	1.2	1.0
	90% UCL	7.6	4.9	1.9	1.7
	Standard error	6.9	10.2	11.8	13.2
	Measured flow	7.5	5.1	3.3	3.0
	Relative error	-11	-20	-55	-57
5130	Estimated flow	0.81	0.44	0.20	0.16
	90% LCL	0.62	0.29	0.10	0.08
	90% UCL	1.0	0.65	0.37	0.33
	Standard error	15.2	22.4	37.4	42.1
	Measured flow	0.86	0.77	0.54	0.53
	Relative error	-6	-43	-63	-70

Table 9.Streamflow statiscontinuous-record sites, nor	tics estimated using regressic rtheast Maui, Hawaii—Contin	n equations, lower and up ued	per confidence intervals, sta	andard errors, measured flo	ws, and relative errors for
[Gaging stations in bold were use duration of base flow; estimated a percent; Relative error is the perc estimate; NA, not applicable]	ed to develop the final regression equand measured flow and confidence i ent difference between the measured	tations for estimating low-flow ntervals are in cubic feet per se I statistic and the estimated sta	statistics; TFQ _{xx} is the xx-percent cond; 90% LCL is und 90% UCL is tistic; measured flows in <i>bold ital</i>	flow duration of total streamflo 90-percent lower and upper conf ics are within the 90-percent con	w; BFQ _{xx} is the xx-percent flow fidence level; Standard error is in fidence interval of the computed
Gaging-station number	Statistic	TFO ₅₀	BFQ50	TFQ ₉₅	BFO ₃₅
5140	Estimated flow	3.3	2.0	0.86	0.76
	90% LCL	2.8	1.6	0.60	0.50
	90% UCL	3.9	2.6	1.2	1.1
	Standard error	9.4	13.8	21.0	23.5
	Measured flow	4.0	3.8	3.0	3.0
	Relative error	-18	-45	-71	-75
5150	Estimated flow	5.2	3.4	1.5	1.4
	90% LCL	4.5	2.7	1.1	0.9
	90% UCL	6.1	4.2	2.1	2.0
	Standard error	8.2	12.1	18.9	21.2
	Measured flow	6.2	5.0	3.0	2.8
	Relative error	-15	-32	-50	-52
5160	Estimated flow	8.1	4.8	2.5	2.0
	90% LCL	7.3	4.1	1.9	1.5
	90% UCL	9.0	5.6	3.1	2.6
	Standard error	6.1	9.0	13.6	15.2
	Measured flow	8.0	5.0	2.3	2.0
	Relative error	1	4-	4	0
5170	Estimated flow	8.0	4.6	2.3	1.8
	90% LCL	7.3	4.0	1.8	1.4
	90% UCL	8.8	5.3	2.8	2.4
	Standard error	5.5	8.1	13.0	14.5
	Measured flow	9.1	5.7	2.8	2.5
	Relative error	-12	-21	-21	-28
5180	Estimated flow	8.5	4.8	2.2	1.8
	90% LCL	7.7	4.2	1.7	1.4
	90% NCL	9.3	5.6	2.8	2.3
	Standard error	5.4	8.0	13.1	14.7
	Measured flow Relative error	-15	0.0 -20	c. 2 -12	<i>2.1</i> -14

Streamflow statistics estimated using regression equations, lower and upper confidence intervals, standard errors, measured flows, and relative errors for continuous-record sites, northeast Maui, Hawaii—Continued Table 9.

duration of base flow; estimated and measured flow and confidence intervals are in cubic feet per second; 90% LCL and 90% UCL is 90-percent lower and upper confidence level; Standard error is in percent; Relative error is the percent difference between the measured statistic and the estimated statistic; measured flows in bold italics are within the 90-percent confidence interval of the computed [Gaging stations in **bold** were used to develop the final regression equations for estimating low-flow statistics; TFQ_{xx} is the xx-percent flow duration of total streamflow; BFQ_{xx} is the xx-percent flow

estimate; NA, not applicable]					
Gaging-station number	Statistic	TFQ50	BFQ50	TF0 ₃₅	BFQ ₉₅
5190	Estimated flow	4.3	2.2	0.89	0.66
	90% LCL	3.9	1.9	0.70	0.51
	90% UCL	4.8	2.6	1.1	0.86
	Standard error	5.7	8.5	13.1	14.6
	Measured flow	4.4	2.4	1.0	0.85
	Relative error	-2	-12	-11	-22
5200	Estimated flow	3.9	2.5	1.4	1.3
	90% LCL	3.3	2.0	1.0	0.89
	90% UCL	4.4	3.1	2.0	1.9
	Standard error	8.1	11.9	18.9	21.2
	Measured flow	3.2	2.0	0.90	0.80
	Relative error	22	25	56	63
5210 + 5200 + 5190	Estimated flow	8.6	4.9	1.8	1.4
	90% LCL	7.7	4.1	1.5	1.1
	90% UCL	9.6	5.7	2.2	1.8
	Standard error	6.2	9.1	11.1	12.4
	Measured flow	10	6.1	2.5	2.0
	Relative error	-14	-20	-28	-30
5240	Estimated flow	4.5	2.4	1.3	1.0
	90% LCL	4.0	2.1	1.1	0.83
	90% UCL	5.0	2.9	1.6	1.3
	Standard error	6.1	9.0	11.1	12.4
	Measured flow	2.2	0.93	0.37	0.30
	Relative error	100	160	250	230
5270	Estimated flow	7.3	4.2	2.0	1.6
	90% LCL	6.9	3.6	1.6	1.2
	90% UCL	8.0	4.7	2.4	2.0
	Standard error	5.0	7.3	11.9	13.3
	Measured flow	5.7	2.8	1.0	0.74
	Kelative error	87	00	82	120

[Gaging stations in bold were use duration of base flow; estimated : percent; Relative error is the perc estimate; NA, not applicable]	d to develop the final regression equations f und measured flow and confidence intervals ent difference between the measured statisti	r estimating low-flow statistics: are in cubic feet per second; 90 ⁶ c and the estimated statistic; mee	. TFQ _{xx} is the xx-percent flow dur & LCL and 90% UCL is 90-perce isured flows in <i>bold italics</i> are w	ration of total streamflow; BFQ _x , ent lower and upper confidence le ithin the 90-percent confidence i	is the xx-percent flow vel; Standard error is in nterval of the computed
Gaging-station number	Statistic	TFQ50	BFQ ₅₀	TFQ ₅₅	BFOs
5311 + 5310	Estimated flow	0.35	0.14	0.08	0.05
	90% LCL	0.25	0.09	0.04	0.02
	90% UCL	0.49	0.23	0.17	0.12
	Standard error	19.3	28.7	45.6	51.6
	Measured flow	0.50	0.17	0.03	0.03
	Relative error	-30	-18	170	67
5360 + 5310 + 5350	Estimated flow	6.6	3.7	1.2	0.94
	90% LCL	5.7	3.0	1.0	0.79
	90% UCL	7.6	4.6	1.4	1.1
	Standard error	8.2	12.1	9.2	10.2
	Measured flow	6.8	3.5	1.7	1.3
	Relative error	-3	9	-29	-28
5420	Estimated flow	0.30	0.14	0.15	0.10
	90% LCL	0.20	0.08	0.08	0.05
	90% UCL	0.45	0.25	0.27	0.20
	Standard error	22.5	33.5	34.9	39.4
	Measured flow	0.88	0.53	0.29	0.27
	Relative error	-99	-74	-48	-63
5430	Estimated flow	1.6	0.78	0.33	0.24
	90% LCL	1.4	0.61	0.22	0.15
	90% UCL	1.9	1.0	0.50	0.37
	Standard error	9.4	13.8	22.9	25.7
	Measured flow	1.4	0.67	0.21	0.15
	Relative error	14	16	57	60

Streamflow statistics estimated using regression equations, lower and upper confidence intervals, standard errors, measured flows, and relative errors for continuous-record sites northeast Maui Hawaii-Continued Table 9.

0.35 0.24 0.50 0.50 **0.44** 0.44

0.46 0.33 0.63 0.63 18.4 **0.55** -16

0.78 0.62 0.99 13.4 1.0

1.5 1.3 1.8 9.1 9.1 1.9 -21

Estimated flow 90% LCL 90% UCL Standard error Measured flow Relative error

5440

Streamflow statistics estimated using regression equations, lower and upper confidence intervals, standard errors, measured flows, and relative errors for continuous-record sites, northeast Maui, Hawaii—Continued Table 9.

duration of base flow; estimated and measured flow and confidence intervals are in cubic feet per second; 90% LCL and 90% UCL is 90-percent lower and upper confidence level; Standard error is in percent; Relative error is the percent difference between the measured statistic and the estimated statistic; measured flows in bold italics are within the 90-percent confidence interval of the computed [Gaging stations in **bold** were used to develop the final regression equations for estimating low-flow statistics; TFQ_{xx} is the xx-percent flow duration of total streamflow; BFQ_{xx} is the xx-percent flow

estimate; NA, not applicable]					
Gaging-station number	Statistic	TFQ50	BFO ₅₀	TFO ₃₅	BFO ₅₅
5450	Estimated flow	12	7.9	3.6	3.3
	90% LCL	11	6.5	2.6	2.3
	90% UCL	14	9.6	5.0	4.6
	Standard error	7.5	11.0	18.1	20.3
	Measured flow	12	6.4	2.5	2.1
	Relative error	0	23	44	57
5528	Estimated flow	1.2	0.55	0.33	0.22
	90% LCL	0.98	0.39	0.21	0.13
	90% UCL	1.5	0.78	0.53	0.37
	Standard error	13.2	19.5	26.8	30.1
	Measured flow	0.14	0.06	0.02	0.01
	Relative error	760	820	1600	2100
5540	Estimated flow	2.9	1.4	0.76	0.55
	90% LCL	2.5	1.1	0.57	0.40
	90% UCL	3.3	1.7	1.0	0.76
	Standard error	8.0	11.9	16.4	18.4
	Measured flow	2.5	1.2	0.69	0.46
	Relative error	16	17	10	20
5550	Estimated flow	6.6	3.5	1.4	1.1
	90% LCL	6.0	3.1	1.1	0.83
	90% UCL	7.3	4.1	1.7	1.4
	Standard error	5.6	8.3	12.6	14.0
	Measured flow	7.0	3.5	1.1	0.80
	Relative error	9-	0	27	38
5560	Estimated flow	6.7	3.6	1.4	1.1
	90% LCL	6.1	3.1	1.1	0.84
	90% UCL	7.4	4.2	1.7	1.4
	Standard error	5.6	8.3	12.5	14.0
	Measured flow	7.1	3.6	1.5	1.0
	Relative error	-9	0	-13	10

, lower and upper confidence intervals, standard errors, measured flows, and relative errors for	
s estimated using regression equation	east Maui, Hawaii—Continued
able 9. Streamflow statistics	ontinuous-record sites, northe

[Gaging stations in **bold** were used to develop the final regression equations for estimating low-flow statistics; TFQ_{xx} is the xx-percent flow duration of total streamflow; BFQ_{xx} is the xx-percent flow duration of base flow; estimated and measured flow and confidence intervals are in cubic feet per second; 90% LCL and 90% UCL is 90-percent lower and upper confidence level; Standard error is in percent; Relative error is the percent difference between the measured statistic and the estimated statistic; measured flows in *bold italics* are within the 90-percent confidence interval of the computed percent; Relative error is the percent difference between the measured statistic and the estimated statistic; measured flows in *bold italics* are within the 90-percent confidence interval of the computed percent; Relative error is the percent difference between the measured statistic and the estimated statistic; measured flows in *bold italics* are within the 90-percent confidence interval of the computed percent; Relative error is the percent difference between the measured statistic and the estimated statistic; measured flows in *bold italics* are within the 90-percent confidence interval of the computed percent; Relative error is the percent difference between the measured statistic; measured flows in *bold italics* are within the 90-percent confidence interval of the computed percent.

estimate; NA, not applicable]					
Gaging-station number	Statistic	TFQ ₅₀	BFO ₅₀	TFQ ₉₅	BFO ₅₅
5570	Estimated flow	2.7	1.7	1.1	0.96
	90% LCL	2.3	1.3	0.76	0.65
	90% UCL	3.2	2.2	1.5	1.4
	Standard error	9.3	13.7	19.9	22.3
	Measured flow	2.7	1.4	0.70	0.58
	Relative error	0	13	57	66
5650 ^a	Estimated flow	4.0	2.5	1.2	1.1
	90% LCL	3.5	2.1	0.94	0.80
	90% UCL	4.5	2.9	1.6	1.5
	Standard error	6.6	9.7	16.0	17.9
	Measured flow	2.6	1.4	0.71	0.61
	Relative error	54	79	69	80
5660 ^a	Estimated flow	1.2	0.73	0.50	0.42
	90% LCL	1.0	0.53	0.32	0.26
	90% UCL	1.5	1.0	0.78	0.70
	Standard error	12.2	18.0	26.0	29.2
	Measured flow	1.0	0.42	0.25	0.21
	Relative error	6	74	100	100
5700 ^a	Estimated flow	15	9.3	3.7	3.3
	90% LCL	13	7.5	2.7	2.3
	90% UCL	17	11	5.1	4.6
	Standard error	8.1	11.9	17.8	19.9
	Measured flow	16	9.5	4.5	4.0
	Relative error	9-	-2	-18	-18
5770 ^a	Estimated flow	12	7.9	4.0	3.7
	90% LCL	10	6.4	2.8	2.5
	90% UCL	14	9.8	5.7	5.5
	Standard error	8.3	12.2	20.3	22.7
	Measured flow	8.9	5.0	2.1	1.7
	Relative error	35	58	66	110
^a located west of current st	udy area				





tive errors for TFQ_{50} and BFQ_{50} , range from -10 to -20 percent and for TFQ_{95} and BFQ_{95} , the relative errors range from -28 to -57 percent. At the third regulated station (station 5090) where streamflow statistics were determined on the basis of combined records, the estimates have relative errors ranging from -57 to -77 percent, indicating that the anomalous input from springs to Hanawi Stream is not accounted for by the equations.

The regression equation for TFQ₅₀ estimates that statistic within a relative error of ± 25 percent at 20 of the continuousrecord gaging stations. Flow at 15 stations is underestimated and flow at 12 stations is the same as the measured statistic or overestimated. All but one of the TFQ₅₀ flows at stations east of Keanae Valley and downstream of the Koolau ditch are underestimated, indicating the influence of springs with high discharge volume in this area (fig. 12). The only overestimate east of Keanae Valley is at gaging station 5070 on an intermittent reach of Makapipi Stream. The regression equation for BFQ_{50} estimates that statistic within a relative error of ± 25 percent at 18 of the continuous-record gaging stations. Flow at 13 stations is underestimated and flow at 14 stations is the same as measured or overestimated. Most of the flow at stations east of Keanae Valley is underestimated with the stations downstream of the Koolau ditch having the greatest errors (fig. 13). The regression equation for TFQ_{95} estimates that statistic within a relative error of ± 50 percent at 15 of the continuousrecord gaging stations. The errors are higher for lower flows because, for the same absolute error in flow, the relative error, in percent, increases as the actual flow decreases. For example, at station 5180, an absolute error of 1.0 ft³/s is 10 percent of TFQ_{50} (10 ft³/s) but 17 percent of BFQ_{50} (6.0 ft³/s). Flow at 13 stations is underestimated and flow at 14 stations is the same as the measured statistic or overestimated. The cluster of underestimated stations east of Keanae Valley and downstream from the Koolau ditch is apparent at lower flows (fig. 14). The regression equation for BFQ₀₅ estimates that statistic within a relative error of ± 50 percent at 12 of the continuous-record gaging stations. Flow at 12 stations is underestimated and flow at 15 stations is the same as the measured statistic or overestimated. The cluster of underestimated stations east of Keanae Valley and below the ditch is persistent at this lowest flow statistic (fig. 15).

The results of applying the equation for TFQ_{50} developed in this study can be compared with results using an equation for TFQ_{50} developed by Fontaine and others (1992) for streams on Maui and Kauai:

 $TFQ_{50} = 4.49(DA)^{0.808}(CE)^{-0.641}(P)^{0.985}$, (2)

where: $\mathrm{TFQ}_{\mathrm{50}}$ is median streamflow, in cubic feet per second,

DA is drainage area, in square miles,

 $C\!E$ is mean altitude of the main stream channel, in feet, and

P is mean annual precipitation, in inches.

For this comparison, *CE* was calculated using the relation 0.5[MAXELEV+MINELEV], and *P* was calculated using the relation 13.53719[Rainfall/DA] for each drainage basin from

the data in table 5, because the original values from Fontaine and others (1992) were not available. For the 17 stations from which data were used to develop the regression equations for this study, equation 2 (1992 study) provides a MAD of 20 percent compared with 12 percent determined using the equation for TFQ₅₀ (table 8) developed in this study. The SE_r and SE_p for the newly developed equation are also lower (15.3 and 20.9 percent, respectively) than for the previous equation (46.2 and 54.3 percent, respectively). Therefore, the newly developed equation for TFQ₅₀ is an improvement over the equation determined by Fontaine and others (1992).

Application of the regression equations to ungaged sites

The regression equations developed for this study from the flow statistics in table 2 and the basin characteristics listed in tables 5 and 6 were used to estimate the TFQ_{50} , TFQ_{95} , BFQ_{50} , and BFQ_{95} duration discharges at selected ungaged sites in the study area (plate 1). Estimated streamflow, prediction intervals, standard error, measured flow, and relative error for 47 ungaged sites are listed in Table 10. Where possible, measured flow values were determined using a combination of flow-duration statistics for an upstream gaging-station and low-flow measurements at ungaged sites downstream of the diversions to provide additional basis for evaluating the equation-based estimates.

Generally, an estimate of flow at an ungaged site made on the basis of a flow-duration discharge at an upstream gaging station and a single measurement of flow at the ungaged site results in a large uncertainty in the estimate. For most ungaged sites in the study area, however, these values are all the information that is available. This technique is considered applicable in the study area because the streams are dry immediately downstream of the diversions at least 50 percent of the time, and measured low flow further downstream represents only those gains to the stream downstream of the diversions. Inflow downstream of the diversions from minor tributaries is insignificant.

Low-flow estimates listed in table 10 were derived from low-flow measurements from three sources: (1) measurements made as part of this study in 2002 and 2003; (2) measurements made as part of a previous USGS study in the area during 1995 to 1999 (Gingerich, 1999); and (3) measurements made by EMI in 1928 (reported in Gingerich, 1999). The average of all these low-flow measurements were assumed roughly equal to TFQ₉₅ flow duration. This is because the flow at the index station (5180) during all the USGS measurements (1995–99 and 2002–03) ranged from TFQ_{94} to TFQ_{97} and the flow was at TFQ₉₀ during the 1928 EMI measurements. The difference in flow between TFQ₉₀ (3.3 ft³/s) and TFQ₉₅ (2.5 ft³/s) at 5180 is 0.8 ft³/s, or a relative difference of 32 percent. However, because the low-flow measurements were all made downstream of the diversion ditches during base-flow periods, there is little variation between low flows even as flow at the

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Figure 12. Distribution of relative error between measured and equation-estimated median total flow (TFQ₅₀) at gaging-stations, northeast Maui, Hawaii.



Figure 13. Distribution of relative error between measured and equation-estimated median base flow (BFQ₅₀) at gaging-stations, northeast Maui, Hawaii.

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Figure 14. Distribution of relative error between measured and equation-estimated Ω_{95} total flow (TF Ω_{95}) at gaging-station, northeast Maui, Hawaii.



Figure 15. Distribution of relative error between measured and equation-estimated Ω_{g_5} base flow (BFQ_{g5}) at gaging-station, northeast Maui, Hawaii.

Table 10.	Streamflow stat	tistics	estimé	ated us	sing rec	gressic	on equa	tions, lo	ower a	and up	per co	onfidence	ce inter	vals, s	tandard	errors,	measu	ired flov	v, and I	elative er	rors for	
ungaged t	basins, northeast	Maui,	Hawai	:=																		

LCL and UCL is 90-percent lower and upper confidence level; Standard error is in percent; Relative error is ([estimated flow – measured flow]/measured flow) in percent; L, lower; M, middle; U, upper; Measured flows in *bold italic* fall within the lower and upper 90-percent confidence interval; East Maui Irrigation Co., Ltd (EMI) 1928 measurements from March 16-20 when index station had a [TFQxx is the xx-percent flow duration of total streamflow; BFQxx is the xx-percent flow duration of base flow; estimated and measured flow and confidence intervals are in cubic feet per second; 90%

Q ₉₀ flow (reported in Gingerich, 19	999); >, likely greater than;	<, likely less than;	, not available]			
Stream location	Statistic	TFO ₅₀	BFO ₅₀	TFO ₃₅	BFO ₃₅	Source of measured flow estimates
Hanawi lower (HwL)	Estimated flow	16	10	5.1	4.6	TFQ ₆₀ , BFQ ₆₀ , BFQ ₆₅ ; combined flow statistics
	90% LCL	14	8.3	3.5	3.0	from 5080 and 5090 upstream; TFO : Average of
	90% UCL	19	13	7.6	7.0	flow on entire stream Feb. 22, 1995 $[\mathbf{Q}_{n}]$ (Ging-
	Standard error	9.1	13.4	22.1	24.8	erich, 1999) and combined \mathbf{Q}_{ac} flows from 5080
	Measured flow	> 28	> 24	25	> 19	and 5090 upstream plus flow Nov. 7, 2003 $[\mathbf{Q}_{\mathbf{M}}]$
	Relative error	< -43	< -58	-79	< -76	7
Kapaula lower (KL)	Estimated flow	7.5	4.6	1.8	1.5	Combined flow statistic from 5100 and 5110 up-
	90% LCL	6.6	3.9	1.4	1.2	stream
	90% UCL	8.7	5.6	2.2	2.0	
	Standard error	7.2	10.6	12.6	14.1	
	Measured flow	> 7.5	> 5.1	> 3.3	> 3.0	
	Relative error	0	<-8	<-45	< -50	
Waiaaka lower (WaL)	Estimated flow	1.3	0.80	030	034	Flow statistics from 5130 upstream; unknown
	90% I.CI.	1.1	0.57	0.22	0.18	amount of upstream diversion at Koolau Ditch
	90% IICL	1.8	11	0.67	0.62	
	Standard error	13.0	19.2	31.9	35.9	
	Measured flow	> 0.86	> 0.77	>0.54	> 0.53	
	Relative error	< 63	4 ×	< -28	< -36	
Paakea lower (PaL)	Estimated flow	4.3	28	15	14	Flow statistics from 5140 upstream; unknown
	90% LCL	3.7	2.3	1.0	0.89	amount of upstream diversion at Koolau Ditch
	90% UCL	5.1	3.6	2.1	2.1	
	Standard error	8.9	13.1	217	5.75	
	Measured flow	> 4.0	> 3.8	~ 3 0	2.1.7 2 3 0	
	Relative error	× ×	< -24	<-50	< -53	
Paakea upper (PaU)	Estimated flow	1.5	0.88	0.50	0.42	No data available
	90% LCL	1.3	0.68	0.33	0.27	
	90% UCL	1.8	1.1	0.75	0.66	
	Standard error	10.0	14.7	23.5	26.4	
	Measured flow	1	1	1	1	
	Relative error	-	1	1	-	
Waiohue lower (WeL)	Estimated flow	7.8	5.5	2.9	2.8	Flow statistics from 5150 upstream plus EMI 1928
	90% LCL	9.9	4.3	1.9	1.8	measurement
	90% UCL	9.3	7.1	4.4	4.5	
	Standard error	9.8	14.4	24.0	26.9	
	Measured flow	> 6.8	> 5.6	> 3.6	> 3.5	
	Relative error	< 16	<-2	< -19	< -20	

tervals, standard errors, measured flow, and relative errors	
ssion equations, lower and upper confidence in	
Streamflow statistics estimated using regre-	oasins, northeast Maui, Hawaii—Continued
Table 10.	ungaged t

[TFQx x is the xx-percent flow duration of total streamflow; BFQxx is the xx-percent flow duration of base flow; estimated and measured flow and confidence intervals are in cubic feet per second; 90% LCL and UCL is 90-percent lower and upper confidence level; Standard error is in percent; Relative error is ([estimated flow – measured flow]/measured flow x 100), in percent; L, lower; M, middle; U, upper; Measured flows in *bold italic* fall within the lower and upper 90-percent confidence interval; East Maui Irrigation Co., Ltd (EMI) 1928 measurements from March 16-20 when index station had a

	Source of measured flow estimates	Flow statistics from 5150 upstream plus EMI 1928	measurement					TFQ ₆ ; EMI 1928 measurement; unknown amount	of upstream diversion at Koolau Ditch					No data available						Flow statistics from 5160 upstream plus average of	6 USGS low-flow measurements in 2002-03					Flow statistics from 5160 upstream plus average of	7 USGS low-flow measurements in 2002-03			
	BFO ₉₅	2.2	1.4	3.4	24.6	> 3.5	< -37	0.88	0.60	1.3	22.4	ł	ł	0.54	0.35	0.81	23.9	ł	1	2.9	2.1	4.0	19.0	5.1	-43	2.3	1.7	3.1	16.5	3.0 -23
	TFO ₉₅	2.3	1.6	3.4	21.9	3.6	-36	1.0	0.70	1.4	20.0	> 0.62	< 61	0.63	0.44	0.92	21.3	ł	ł	3.4	2.5	4.6	16.9	5.4	-38	2.8	2.2	3.6	14.8	3.3 -18
not available]	BFO50	4.4	3.5	5.6	13.2	> 5.6	< -21	2.2	1.8	2.8	12.7	1	-	1.1	0.89	1.4	13.3	ł	ł	6.7	5.6	8.1	10.6	> 8.1	< -17	6.2	5.3	7.3	9.0	6.0 < 3
, likely less than;,	TFO50	6.5	5.5	7.6	9.0	> 6.8	4 - ∧	3.6	3.1	4.2	8.6	ł	1	1.9	1.6	2.2	9.1	ł	1	11	9.7	13	7.2	> 11	0	10	9.4	12	6.1	> 9.0 < 11
1999); >, likely greater than; <	Statistic	Estimated flow	90% LCL	90% UCL	Standard error	Measured flow	Relative error	Estimated flow	90% LCL	90% UCL	Standard error	Measured flow	Relative error	Estimated flow	90% LCL	90% UCL	Standard error	Measured flow	Relative error	Estimated flow	90% LCL	90% UCL	Standard error	Measured flow	Relative error	Estimated flow	90% LCL	90% UCL	Standard error	Measured flow Relative error
Q ₉₀ flow (reported in Gingerich,	Stream location	Waiohue middle (WeM)						Puakaa middle (PuM)						Puakaa upper (PuU)						Kopiliula lower (KpL)						Kopiliula middle (KpM)				

Table 10. Streamflow statistics estimated using regression equations, lower and upper confidence intervals, standard errors, measured flow, and relative errors for ungaged basins, northeast Maui, Hawaii—Continued

LCL and UCL is 90-percent lower and upper confidence level; Standard error is in percent; Relative error is ([estimated flow - measured flow]/measured flow x 100), in percent; L, lower; M, middle; U, upper; Measured flows in bold italic fall within the lower and upper 90-percent confidence interval; East Maui Irrigation Co., Ltd (EMI) 1928 measurements from March 16-20 when index station had a [TFQx x is the xx-percent flow duration of total streamflow; BFQxx is the xx-percent flow duration of base flow; estimated and measured flow and confidence intervals are in cubic feet per second; 90%

	BFQ ₅₅ Source of measured flow estimates	2.2 Flow statistics from 5170 upstream plus EMI	1.7 1928 measurement	2.9	16.2	> 2.8	< -21	2.1 Flow statistics from 5170 upstream plus EMI	1.6 1928 measurement	2.8	15.7	2.8	< -25	2.1 Flow statistics from 5180 upstream plus EMI	1.6 1928 measurement	2.7	15.9	> 2.9	< -28	2.0 Flow statistics from 5180 upstream plus EMI	1.5 1928 measurement	2.6	15.6	2.9	-31	1.8 Combined flow statistics from 5190, 5200,	1.4 and 5210 upstream plus average of 2 USGS	^{2.2} low-flow measurements in 2002-03: unknown	13.7 amount of taro diversion and return flow		2.1 <
	TFQ ₉₅	2.7	2.1	3.4	14.4	> 3.2	<-16	2.5	2.0	3.3	14.0	3.2	-22	2.5	2.0	3.3	14.2	> 3.3	< -24	2.4	1.9	3.1	13.9	3.3	-27	2.2	1.7	2.7	12.2		7.7
not available]	BFO ₅₀	6.1	5.2	7.1	8.8	> 6.1	0	5.6	4.8	6.4	8.8	> 6.1	<- 8-	6.0	5.2	7.1	8.9	> 6.7	< -10	5.6	4.8	6.5	8.5	> 6.7	< -16	5.5	4.7	6.4	8.5	0 4	
, likely less than;,	TFO ₅₀	10	9.3	12	6.0	> 9.4	< 6	9.5	8.5	10	5.8	> 9.4	<1	10	9.3	12	6.0	> 11	6->	9.7	8.7	11	5.8	> 11	<-12	9.5	8.6	10.5	5.8	202	C.K <
9); >, likely greater than; <	Statistic	Estimated flow	90% LCL	90% UCL	Standard error	Measured flow	Relative error	Estimated flow	90% LCL	90% UCL	Standard error	Measured flow	Relative error	Estimated flow	90% LCL	90% UCL	Standard error	Measured flow	Relative error	Estimated flow	90% LCL	90% UCL	Standard error	Measured flow	Relative error	Estimated flow	90% LCL	90% UCL	Standard error	Maccurad flour	MOLLOC INCLU
Q ₉₀ flow (reported in Gingerich, 1999	Stream location	East Wailuaiki lower (EWL)						East Wailuaiki middle	(EWM)					West Wailuaiki lower	(WWL)	~				West Wailuaiki middle	(MMM)					Wailuanui lower (WL)					

DUL and UUL is SUPPERCENTIONET and upper; Measured flows in <i>bold italic</i> Q ₉₀ flow (reported in Gingerich, 199	of all within the lower revert of the lower and up of (); >, likely greater than; <,	tandard error is in pe pper 90-percent confi likely less than;, n	itent; relative end idence interval; Eas iot available]	or 1s (lesumateu 110 st Maui Irrigation C	w – measureu 110 20., Ltd (EMI) 192	w//measured flow x 100), in percent, L, lower; M, inituate, U, 28 measurements from March 16-20 when index station had a
Stream location	Statistic	TFO ₅₀	BFO ₅₀	TFO ₉₅	BFO ₉₅	Source of measured flow estimates
Waiokomilo lower (WoL)	Estimated flow	14	8.7	2.8	2.4	TFQ _{ac} : 1999 USGS low-flow measurement
	90% LCL	12	6.8	2.2	1.8	(Gingerich, 1999)
	90% UCL	17	11	3.6	3.2	
	Standard error	9.8	14.4	14.2	15.9	
	Measured flow	1	1	5.7	1	
	Relative error	1	-	-51	1	
Waiokomilo middle (WoM)	Estimated flow	10	6.1	2.2	1.8	TFO : Average of EMI 1928 measurement
	90% LCL	9.0	5.0	1.8	1.4	and 1999 USGS low-flow measurement
	90% UCL	12	7.4	2.6	2.3	(Gingerich 1000)
	Standard error	7.3	10.7	11.4	12.8	
	Measured flow	ł	ł	4.9	1	
	Relative error	ł	ł	-55	1	
Waiokomilo upper (WoU)	Estimated flow	7.0	3.9	1.3	1.1	No data available
	90% LCL	6.1	3.3	1.1	06.0	
	90% UCL	7.8	4.7	1.6	1.3	
	Standard error	6.9	10.2	9.1	10.1	
	Measured flow	ł	ł	ł	ł	
	Relative error	1	1	1	1	
Ohia lower (OL)	Estimated flow	8.7	8.6	4.6	6.1	Ohia Spring average flow is 4.7 cubic feet per
~	90% LCL	5.5	4.3	1.5	1.8	second (Stearns and Macdonald, 1942) but
	90% UCL	14	17	14	21	little flow reaches the ocean due to infiltration
	Standard error	26.6	40.0	69.1	79.3	and aoricultural evaportansmiration losses
	Measured flow	1	1	1	1	ania agus annana s'abou ang hu annan 100000
	Relative error	1	1	1	1	
Palauhulu lower (PhL)	Estimated flow	17	11	4.4	4.0	TFO : TFO flow at 5220, measuring taro
~	90% LCL	14	8.9	3.1	2.7	diversion from stream; losing stream therefore
	90% UCL	20	15	6.3	6.0	effects of natural flow addition are unknown
	Standard error	9.6	14.2	20.5	23.0	
	Measured flow	1	ł	> 2.4	1	
	Relative error		1	< 83	1	

[TFQx x is the xx-percent flow duration of total streamflow; BFQxx is the xx-percent flow duration of base flow; estimated and measured flow and confidence intervals are in cubic feet per second; 90% 1.C. and 11CL is 90-nercent lower and inner confidence level. Standard error is in nercent. Relative error is (festimated flow – measured flow)/measured flow x 100) in nercent. Lower: M. middle: 11. Streamflow statistics estimated using regression equations, lower and upper confidence intervals, standard errors, measured flow, and relative errors for ungaged basins, northeast Maui, Hawaii—Continued Table 10.

Table 10. Streamflow statistics estimated using regression equations, lower and upper confidence intervals, standard errors, measured flow, and relative errors for ungaged basins, northeast Maui, Hawaii—Continued

LCL and UCL is 90-percent lower and upper confidence level; Standard error is in percent; Relative error is ([estimated flow - measured flow]/measured flow x 100), in percent; L, lower; M, middle; U, upper; Measured flows in bold italic fall within the lower and upper 90-percent confidence interval; East Maui Irrigation Co., Ltd (EMI) 1928 measurements from March 16-20 when index station had a [TFOx x is the xx-percent flow duration of total streamflow; BFOxx is the xx-percent flow duration of base flow; estimated and measured flow and confidence intervals are in cubic feet per second; 90%

Q ₉₀ flow (reported in Gingerich, 19	999); >, likely greater than; <,	likely less than;, 1	not available]			
Stream location	Statistic	TFO ₅₀	BFO ₅₀	TFO ₉₅	BFO ₃₅	Source of measured flow estimates
Palauhulu middle (PhM)	Estimated flow	14	9.3	3.9	3.5	Plunkett Spring average flow is 2.7 cubic feet
	90% LCL	12	7.5	2.8	2.4	per second (Stearns and Macdonald, 1942) but
	90% UCL	16	12	5.4	5.1	stream poes dry due to infiltration losses so
	Standard error	8.3	12.2	18.9	21.1	effects of natural flow addition are unknown
	Measured flow	1	ł	ł	1	
	Relative error	1	1	1	ł	
Kano upper (KoU)	Estimated flow	4.5	2.5	1.0	0.82	No data available
	90% LCL	4.2	2.2	0.87	0.68	
	90% UCL	4.9	2.8	1.2	0.99	
	Standard error	4.5	6.6	9.7	10.9	
	Measured flow	1	1	1	1	
	Relative error	1	1	1	1	
Hauoli Wahine upper	Estimated flow	1.5	0.93	0.88	0.75	No data available
(IIMH)	90% LCL	1.2	0.64	0.66	0.54	
	90% UCL	2.0	1.4	1.2	1.0	
	Standard error	14.7	21.8	16.7	18.7	
	Measured flow	1	1	1	1	
	Relative error	1	1	1	1	
Piinaau lower (PiL)	Estimated flow	40	28	13	13	TFO: EMI 1928 measurement. unknown
	90% LCL	31	19	7.1	6.3	amount of unstream diversion at Koolan Ditch
	90% UCL	52	41	25	25	
	Standard error	14.6	21.6	36.3	40.9	
	Measured flow	1	1	> 0.47	1	
	Relative error	ł	1	< 2700	1	
Piinaau middle (PiM)	Estimated flow	28	20	12	11	TFQ _n : EMI 1928 measurement, unknown
	90% LCL	22	13	6.7	5.8	amount of upstream diversion at Koolau Ditch
	90% UCL	36	29	22	22	
	Standard error	15.0	22.2	34.9	39.3	
	Measured flow	ł	ł	> 0.47	1	
	Relative error	1	ł	< 2500	1	

LCL and UCL is 90-percent lower upper; Measured flows in <i>bold ital</i> . Q ₉₀ flow (reported in Gingerich, 19	and upper confidence level; <i>lic</i> fall within the lower and u 999); >, likely greater than; <	Standard error is in p upper 90-percent con: c, likely less than;,	percent; Relative er fidence interval; E not available]	ror is ([estimated f. ast Maui Irrigation	low – measured flo Co., Ltd (EMI) 192	w]/measured flow x 100), in percent; L, lower; M, middle; U, 28 measurements from March 16-20 when index station had a
Stream location	Statistic	TFO ₅₀	BFO50	TFO ₉₅	BFO ₃₅	Source of measured flow estimates
Piinaau upper (PiU)	Estimated flow	21	14	9.4	8.5	No data available
	90% LCL	16	9.7	5.5	4.7	
	90% UCL	27	20	16	16	
	Standard error	14.4	21.3	31.1	34.9	
	Measured flow	1	1	1	1	
	Relative error	ł	ł	1	ł	
Nuaailua lower (NL)	Estimated flow	6.6	7.4	4.0	4.1	TFO: EMI 1928 measurement: unknown
	90% LCL	8.0	5.4	2.4	2.3	amount of unstream diversion at Spreckels
	90% UCL	12	10	6.7	7.2	Ditch: effects of natural flow addition are
	Standard error	12.1	17.8	29.8	33.5	untrous because of natural 110% autilities of
	Measured flow	ł	ł	> 0.31	ł	unknown uccause su cann may uc gammig ur
	Relative error	ł	ł	< 1200	ł	IOSIIIg
Municipation and dis (MM)	Estimated flow	3.0	чc	- -	1	TEO · EMI 1038 magnimum finanti
				7.1	1.1	$1\mathbf{r}\mathbf{V}_{95}$. LIVII 1720 IIICASUICIIICUILUUII
	90% LCL	5.5	7.0	0.80	c/.0	amount of upstream diversion at Spreckels
	90% UCL	4.5	3.1	1.8	1.7	Ditch; effects of natural flow addition are
	Standard error	8.6	12.6	20.8	23.3	unknown because stream may be gaining or
	Measured flow	1	1	> 0.31	1	losino
	Relative error	1	1	< 290	1	
Nuaailua upper (NU)	Estimated flow	0.56	0.28	0.19	0.15	No data available
4	90% LCL	0.43	0.19	0.11	0.08	
	90% UCL	0.73	0.42	0.34	0.27	
	Standard error	15.1	22.4	33.3	37.5	
	Measured flow	;	1	1	1	
	Relative error	1	ł	1	1	
Honomanu lower (HnL)	Estimated flow	16	10	4.6	4.0	Values are maximum estimates using flow
	90% LCL	14	8.3	3.2	2.7	statistics from 5270 upstream because stream
	90% UCL	19	13	6.6	6.0	infiltration losses are unknown; stream goes
	Standard error	8.5	12.6	20.7	23.2	dry at low flow
	Measured flow	< 5.7	< 2.8	< 1.0	< 0.74	
	Relative error	> 180	> 260	> 320	> 440	

[TFQxx is the xx-percent flow duration of total streamflow; BFQxx is the xx-percent flow duration of base flow; estimated and measured flow and confidence intervals are in cubic feet per second; 90% Table 10. Streamflow statistics estimated using regression equations, lower and upper confidence intervals, standard errors, measured flow, and relative errors for ungaged basins, northeast Maui, Hawaii—Continued

Table 10. Streamflow statistics estimated using regression equations, lower and upper confidence intervals, standard errors, measured flow, and relative errors for ungaged basins, northeast Maui, Hawaii—Continued

LCL and UCL is 90-percent lower and upper confidence level; Standard error is in percent; Relative error is ([estimated flow - measured flow]/measured flow x 100), in percent; L, lower; M, middle; U, upper; Measured flows in bold italic fall within the lower and upper 90-percent confidence interval; East Maui Irrigation Co., Ltd (EMI) 1928 measurements from March 16-20 when index station had a [TFOx x is the xx-percent flow duration of total streamflow; BFOxx is the xx-percent flow duration of base flow; estimated and measured flow and confidence intervals are in cubic feet per second; 90%

BFQ Source of measured flow estimates	3.1 Values are maximum estimates using flow	2.2 statistics from 5270 upstream because stream	4.4 infiltration losses are unknown: stream ones	19.9 dry at low flow	< 0.74 us the matrice < 0.74	320	2.1 No data available	1.4	3.4	25.7	:	-	2.0 TFQ _{n≤} : EMI 1928 measurement; unknown	1.3 amount of upstream diversion at Spreckels	3.1 Ditch: effects of natural flow addition are	25.0 Inchnown heraitee stream may he raining or	0.00 locing: stream cool day of low flow	NA IUMIIIS, SUEGIII BOES ULY ALIUW IIUW	1.3 Flow statistics from 5360 upstream plus EMI	1.1 1928 measurement: effects of natural flow	1.6 addition are unknown because stream is losino	10.4 etrasm more drug at loug flour	1.3 sucan gues up at row now	0	1.2 Flow statistics from 5360 upstream plus EMI	1.0 1928 measurement; effects of natural flow	^{1.4} addition are unknown because stream is losing.	10.2 stream poes drv at low flow	1.3 Junit Second and an and and a
TFQ.,	3.7	2.7	5.0	17.8	< 1.0	> 240 >>	2.2	1.5	3.4	22.9	1	-	2.1	1.4	3.1	22.3	0.00	NA	1.6	1.3	1.9	9.3	1.7	9	1.5	1.2	1.7	9.1	1.7
not available] BF0	8.0	6.6	9.7	10.8	< 2.8	> 190	4.5	3.5	5.7	13.8	-	ł	3.9	3.1	5.0	13.5	1	ł	5.6	4.2	7.4	15.7	3.5	60	5.0	3.9	6.5	14.8	3.5
<, likely less than;, TFQ	13	11	15	7.3	< 5.7	> 130	6.5	5.5	7.6	9.4	1	ł	5.8	4.9	6.8	9.1	ł	1	9.7	8.0	12	10.7	6.8	43	8.8	7.3	10	10.1	6.8
9); >, likely greater than; < Statistic	Estimated flow	90% LCL	90% UCL	Standard error	Measured flow	Relative error	Estimated flow	90% LCL	90% UCL	Standard error	Measured flow	Relative error	Estimated flow	90% LCL	90% UCL	Standard error	Measured flow	Relative error	Estimated flow	90% LCL	90% UCL	Standard error	Measured flow	Relative error	Estimated flow	90% LCL	90% UCL	Standard error	Measured flow
Q ₉₀ flow (reported in Gingerich, 195 Stream location	Honomanu middle (HnM)	~					Punalau lower (PIL)						Punalau middle (PIM)						Haipuaena lower (HaL)	~ 4					Haipuaena middle lower	(HaML)	~		

LCL and UCL is 90-percent lower a upper; Measured flows in <i>bold itali</i> , Q ₉₀ flow (reported in Gingerich, 199	and upper confidence level; <i>c</i> fall within the lower and u 99); >, likely greater than; <	Standard error is in p upper 90-percent coni c, likely less than;,	ercent; Relative eri ödence interval; Ea not available]	ror is ([estimated f] ast Maui Irrigation	low – measured flc Co., Ltd (EMI) 193	w]/measured flow x 100), in percent; L, lower; M, middle; U, 28 measurements from March 16-20 when index station had a
Stream location	Statistic	TFO ₅₀	BFO ₅₀	TFO ₉₅	BFQ	Source of measured flow estimates
Haipuaena middle upper	Estimated flow	7.8	4.5	1.4	1.1	Flow statistics from 5360 upstream; effects
(HaMU)	90% LCL	6.7	3.5	1.2	0.93	of natural flow addition are unknown because
	90% UCL	9.2	5.6	1.6	1.3	stream may be gaining or losing
	Standard error	0.6	13.2	9.0	10.0	Survey to Survey of fair unong
	Measured flow	6.8	3.5	1.7	1.3	
	Relative error	15	29	-18	-15	
Puohokamoa lower (PL)	Estimated flow	18	12	5.0	4.7	Flow statistics from 5450 upstream plus EMI
	90% LCL	15	9.6	3.4	3.0	1928 measurement: effects of natural flow
	90% UCL	22	16	7.5	7.3	addition are unknown hecause stream may he
	Standard error	10.1	14.9	22.7	25.4	autitation are annual accurate account may accurate and accurate and accurate accu
	Measured flow	> 12	> 6.6	> 2.7	> 2.3	
	Relative error	< 50	< 82	< 85	< 100	
Puohokamoa middle lower	Estimated flow	17	11	4.8	4.4	Flow statistics from 5450 upstream plus EMI
(PML)	90% LCL	14	9.0	3.3	2.9	1928 measurement; effects of natural flow
~	90% UCL	20	15	7.0	6.8	addition are unknown because stream may be
	Standard error	9.6	14.1	22.0	24.7	ogining or losing
	Measured flow	> 12	> 6.6	2.7	> 2.3	
	Relative error	< 42	< 67	78	< 91	
Puohokamoa middle upper	Estimated flow	15	9.9	4.3	3.9	Flow statistics from 5450 upstream; effects
(PMU)	90% LCL	13	7.9	3.0	2.6	of natural flow addition are unknown because
	90% UCL	17	12.4	6.1	5.8	stream may be gaining or losing
	Standard error	8.7	12.8	20.3	22.8	
	Measured flow	> 11	> 6.4	> 2.5	> 2.1	
	Relative error	< 25	< 55	< 72	< 86	
Wahinepee lower (WeL)	Estimated flow	2.4	1.8	1.6	1.6	TFQ: EMI 1928 measurement; unknown
4	90% LCL	1.8	1.1	0.88	0.83	amount of upstream diversion at Wailoa Ditch;
	90% UCL	3.2	2.7	2.8	3.0	effects of natural flow addition are unknown
	Standard error	16.6	24.6	33.9	38.1	because stream may be gaining or losing
	Measured flow	1	1	> 0.46	1	
	Relative error	1	1	< 250	1	

[TFQxx is the xx-percent flow duration of total streamflow; BFQxx is the xx-percent flow duration of base flow; estimated and measured flow and confidence intervals are in cubic feet per second; 90% Streamflow statistics estimated using regression equations, lower and upper confidence intervals, standard errors, measured flow, and relative errors for ungaged basins, northeast Maui, Hawaii—Continued Table 10.

Table 10. Streamflow statistics estimated using regression equations, lower and upper confidence intervals, standard errors, measured flow, and relative errors for ungaged basins, northeast Maui, Hawaii—Continued

LCL and UCL is 90-percent lower and upper confidence level; Standard error is in percent; Relative error is ([estimated flow - measured flow]/measured flow x 100), in percent; L, lower; M, middle; U, upper; Measured flows in bold italic fall within the lower and upper 90-percent confidence interval; East Maui Irrigation Co., Ltd (EMI) 1928 measurements from March 16-20 when index station had a [TFQx x is the xx-percent flow duration of total streamflow; BFQxx is the xx-percent flow duration of base flow; estimated and measured flow and confidence intervals are in cubic feet per second; 90%

Q ₉₀ flow (reported in Gingerich, 19	199); >, likely greater than; <	, likely less than;, I	not available]			
Stream location	Statistic	TFO ₅₀	BFO50	TFO ₃₅	BFO ₉₅	Source of measured flow estimates
Wahinepee middle (WeM)	Estimated flow	1.3	0.89	0.97	0.94	TFQ.5: EMI 1928 measurement; unknown
	90% LCL	0.92	0.54	0.56	0.50	amount of upstream diversion at Wailoa Ditch;
	90% UCL	1.8	1.4	1.7	1.7	effects of natural flow addition are unknown
	Standard error	19.2	28.5	32.4	36.5	hecanse stream may he oaining or losing
	Measured flow	1	1	0.46	1	Surger to Surray of this man of surray
	Relative error	1	1	110	1	
Waikamoi lower (WiL)	Estimated flow	10	5.7	2.0	1.6	Flow statistics from 5550 and 5570 upstream
~	90% LCL	9.0	4.8	1.6	1.2	plus Waikamoi middle upper and middle
	90% UCL	12	6.9	2.6	2.1	lower sites: effects of natural flow addition are
	Standard error	7.1	10.4	13.2	14.8	unknown herause stream is losing: stream mes
	Measured flow	< 11	< 5.9	< 2.8	< 2.4	diminum occause succin is rosmig, succini goes
	Relative error	6	<- <	> -29	> -33	ury at row mow
Waikamoi middle lower	Estimated flow	9.6	T.	Ċ	t.	Flow statistics from 5550 and 5570 upstream
(WiML)	90% LCL	8.6	4. v 4. v	7.0	0.1	plus Waikamoi middle upper plus average of 9
	90% UCL	11	C.4 ∠ ∧	0.1 2 c	1.2	USGS low-flow measurements in 2002-03
	Standard error	6.6	0.4 0	C.7 1 2 1	2.U 1 / 7	
	Measured flow	> 11	~ 50	2.61	74.	
	Relative error	-13	φ	-29	-33	
Waikamoi middle upper	Estimated flow	9.2	5.2	2.0	1.6	Flow statistics from 5550 and 5570
(WiMI)	90% LCL	8.3	4.4	1.6	1.2	unstream plus average of 10 USGS low-flow
	90% UCL	10	6.1	2.5	2.0	measurements in 2002-03
	Standard error	6.2	9.2	13.1	14.6	
	Measured flow	> 10.5	> 5.7	2.6	> 2.2	
	Relative error	<-16	6->	-23	< -27	
Kolea lower (KaL)	Estimated flow	4.8	3.4	1.9	1.9	TFO: EMI 1928 measurement; unknown
×.	90% LCL	3.9	2.5	1.2	1.1	amount of unstream diversion at Wailoa Ditch:
	90% UCL	5.9	4.6	3.1	3.2	effects of natural flow addition are unknown
	Standard error	11.4	16.9	28.1	31.5	hecalise stream may be vaining or losing
	Measured flow	ł	1	> 0.2	ł	Surger to Suring and Initi time no aconnego
	Relative error	ł	ł	< 1100	ł	

ow statistics estimated using regression equations, lower and upper confidence intervals, standard errors, measured flow, and relative errors for	cheast Maui, Hawaii—Continued
e 10. Streamflo	aged basins, nort
Tabl	nng;

[TFQx is the xx-percent flow duration of total streamflow; BFQxx is the xx-percent flow duration of base flow; estimated and measured flow and confidence intervals are in cubic feet per second; 90% LCL and UCL is 90-percent lower and upper confidence level; Standard error is in percent; Relative error is ([estimated flow – measured flow]/measured flow); in percent; L, lower; M, middle; U, upper; Measured flows in *bold italic* fall within the lower and upper 90-percent confidence interval; East Maui Irrigation Co., Ltd (EMI) 1928 measurements from March 16-20 when index station had a O., flow (reported in Gingerich, 1999); >, likely greater than; <, likely less than; -, not available]

	Source of measured flow estimates	TFQ _{as} : EMI 1928 measurement; unknown	amount of upstream diversion at Wailoa Ditch;	effects of natural flow addition are unknown	because stream may be vaining or losing			
	BFO ₉₅	1.5	0.89	2.5	30.2	ł	ł	
	TFQ ₉₅	1.6	0.98	2.5	26.9	> 0.16	< 900	
	BFO50	2.5	1.9	3.4	16.7	ł	ł	
, HNCLY ICSS UIGH,, II	TFO ₅₀	3.6	3.0	4.4	11.3	ł	ł	
, 1999), 2, IINCIJ BICARLI UIAII, 2,	Statistic	Estimated flow	90% LCL	90% UCL	Standard error	Measured flow	Relative error	
Q ₉₀ LIOW (LEPOTICU III OIIIGUICI),	Stream location	Kolea middle (KaM)						

index station varied between TFQ₉₀ and TFQ₉₅. Flow-duration values for five gaging stations (5090, 5110, 5130, 5140, and 5210) operated downstream of the Koolau Diversion indicate an average relative difference between TFQ₉₀ and TFQ₉₅ of 11 percent, and absolute differences ranging from 0.05 ft³/s at station 5130 to 1.0 ft³/s at station 5090. In general, base flows sustained by springs downstream from the diversion ditch exhibit less variability. Where measured TFQ₅₀ and BFQ₅₀ flow durations are listed in table 10, they are based on a combination of upstream flow statistics and the measured TFQ₉₅ low flows, but are preceded by a greater than symbol (>) indicating that the TFQ₉₅ flow is a minimum value and the actual flow is higher by an unknown amount.

Application of the regression equations to estimate flow at some of the sites shown in table 10 violates the assumptions of the regression analysis because they are intermittent-flow sites or have basin characteristics that are outside the range used to develop the equations. Some of the poor results shown in table 10 can be explained by these violations. A plot of the spatial distribution of relative-error values for TFQ_{0s} , the only duration discharge for which some measured flow values are available, shows two distinct groupings of results. East of Keanae Valley, the TFQ₉₅ equation generally underestimates flow, and within and west of Keanae Valley, the equation generally overestimates flow (fig. 16). This grouping reflects the pattern for ground-water occurrence east and west of Keanae Valley described by Gingerich (1999). West of Keanae Valley, ground water occurs at high elevations in the low-permeability Kula Volcanics and at low elevations in the higher permeability Honomanu Basalt. Streams at high elevations gain base flow from the upper perched zone and lose water to the underlying unsaturated zone nearer the coast. Where streams lose water at lower elevations, the regression equations generally overestimate the amount of water in the stream. East of Keanae Valley, the discharge of ground water from a vertically extensive freshwater lens causes streams to gain flow all the way to the coast. The regression equations generally underestimate the additional streamflow gained from springs. Within Keanae Valley, streamflow gain from the freshwater lens is lost to the veneering lava flows of the Hana Volcanics; hence, the equations overestimate streamflow here as well.

Ordinarily, regression equations would be developed for each distinct hydrologic regime (east or west of Keanae Valley) to better account for the basin characteristics controlling streamflow. However, dividing the gaging stations used to generate the regression equations (n = 17) into two groups would result in the number of stations in each group (n = 8 and n = 9) being too small for significant statistical analyses.

At sites where flow is underestimated by the regression equation, the relative errors range from 7 to 79 percent and average 30 percent. At sites where flow is overestimated, the relative errors are not always meaningful because many of the streams are dry. Therefore, any flow estimate greater than zero means that relative error for that site cannot be calculated.

Most-reliable Estimates of Natural Flow-Duration Statistics

Most-reliable estimates of flow-duration statistics for natural (undiverted) streamflow at ungaged sites on 21 streams in the study area were made using a combination of continuous record gaging-station data, low-flow measurements, and values determined from the regression equations developed as part of this study (Table 11). No estimates were made for Piinaau Stream because no flow data were available and the regression equations were not applicable to this intermittent stream. Furthermore, all three of the basin characteristics for Piinaau Stream that are used in the regression equations fall outside the range of values used to develop the equations (table 8). The reliability of the estimates depends on the combination of data used to develop the estimates. Estimates of flow statistics developed on the basis of data from continuous-record gaging stations are deemed the most reliable; those estimates developed from the regression equations alone are considered the least reliable. Estimates that are developed for sites downstream from gaging stations and adjusted on the basis of low-flow measurements or the regression equations are considered to be of intermediate reliability.

The various data combinations used to develop the most-reliable estimates can be described using Hanawi Stream as an example. At the Hanawi upper site, the TFQ₅₀, TFQ₉₅, BFQ₅₀, and BFQ₉₅ duration discharges were estimated on the basis of data available from gaging station 5080. At the Hanawi middle site, which is downstream from the Koolau ditch, duration discharges were estimated from a combination of daily flows at gaging stations 5090 at the middle site and 5080 at the upper site. At the Hanawi lower site, TFQ₉₅ was developed using the estimated TFQ₉₅ at the middle site (19 ft³/s), plus the average of gains in streamflow downstream from the middle site, which was determined from two pairs of low-flow measurements [2.8 ft³/s (TFQ₀₇ at index site); and 3.4 ft³/s (TFQ₉₄ at index site)]. TFQ₅₀ at the Hanawi lower site was estimated using the estimated TFQ_{50} from the middle site (28 ft³/s), plus the estimated gain between the lower and middle sites, using the TFQ₅₀ regression equation (16 ft³/s - 12 ft³/s =4.0 ft³/s). BFQ₅₀ and BFQ₉₅ were estimated in a similar manner to TFQ_{50} .

Discussion of results for selected streams

For some streams, the decisions made in developing the most-reliable estimates of flow-duration statistics require further discussion. Estimates for middle and lower sites on Makapipi Stream were not attempted due to lack of flow data at these sites and the inapplicability of the regression equations to this intermittent/losing stream.

The TFQ_{95} estimate at the middle site on Waiokomilo Stream was based on flow measurements upstream that demonstrated net gains of about 4.1 ft³/s (Gingerich, 1999)



Figure 16. Distribution of relative error between measured and equation-estimated Ω_{95} total flow (TF Ω_{95}) at ungaged sites in the study area and at selected gages west of the study area, northeast Maui, Hawaii.

Table 11. Estimates of natural (undiverted) streamflow statistics for gaged and ungaged basins, northeast Maui, Hawaii.

[TFQx x is the xx-percent flow duration of total streamflow; BFQx x is the xx-percent flow duration of base flow; all flows are in cubic feet per second; numbers in **bold italic** are considered maximums at sites downstream of unquantified but known losing reaches; g.s., gaging station; adjustment]

Stream location	TFQ	BFO	TFO	BFQ	Source of estimate
Makapipi unner (5070)	с С	7 7			Continuous record agains station
Hanawi	1.1	L.	0.0	0.0	
lower (HwL)	32	26	22	19	Middle site estimate plus equation adj.; TFQ ₉₅ : Middle site estimate plus low-flow measurements
middle (5090)	28	24	19	19	Continuous record gaging station plus upper site estimate
(ngnc) raddn	7.1	4.6	2.4	2.2	Continuous record gaging station
Kapaula	0 2	L 2	2 5	((المنافع منابع مدابستان سمسمينا مستعلمات مناطر
nower (NL) middle (5110)	0.0	1.0	 	7.0	Mucute sue estimate prior regionsion equation auj.
$\frac{111000}{1000}$		0.1	0.0 0	0.0	Continuous record gaging station plus upper site estimate
Wainaka	4.9	7.8	1.2	1.1	Continuous record gaging station
lower (WaL)	1.4	1.1	0.73	0.70	Middle site estimate plus regression equation adj.
middle (5130)	0.86	0.77	0.54	0.52	Continuous record gaging station
Paakea	2 9	v v	L L		Middla cita actimota nhuc varracción acuntion adi
nower (Fall) middle (5140)	0.0 V	0 4 0 1	v.	9.7 7	Antonio suo sontinao pue regissioni equationi auj. Osnitinios no contina totojon altro inversa feta estistante
upper (PaU)	0.0 V	0.00	0.50	1.0	Collutinous tector gaging station plus upper site contrate Removerion equiption
Waiohue	C: 1	0000	0000		
lower (WeL)	8.8	7.5	4.5	3.6	Middle site estimate plus measurement plus regression equation adj.
middle (WeM)	7.4	6.0	3.9	3.0	Upper site estimate plus equation adj.; TFQ _n ; Upper site estimate plus low-flow measurement
upper (5150)	6.2	5.0	3.0	2.9	Continuous record gaging station
Kopiliula		ц С	L L	с с	
lower (KpL)	c1 ;	C.Y	0.0	3.8 0.0	Middle sites estimates plus equation adj. $\mathbf{F}\mathbf{Q}_{35}$ Middle sites estimates plus low-riow measurements
Middle (KpM)	10	6.5	3.4	2.3	Upper site estimate plus equation adj.; $\mathbf{TFQ}_{\mathbf{0s}}$: Upper site estimate plus low-flow measurements
middle (PuM)	3.6	2.2	1.1	0.90	Regression equation; TFQ ₉₅ : Upper site estimate plus low-flow measurement
upper (5160)	8.0	5.0	2.4	2.0	Continuous record gaging station
upper (ruu)	1.9	1.1	0.60	0.50	Regression equation
East Wailuaiki lower (FWL)	11	7.2	3.4	2.9	Middle site estimate plus equation adi.: TFO : Middle site estimate plus low-flow measurement
middle (EWM)	11	6.8	3.2	2.8	Upper site estimate plus equation adj.; TFQ: Upper site estimate plus low-flow measurement
upper (5170)	9.1	5.8	2.9	2.5	Continuous record gaging station
West Wailuaiki	10	<i>C L</i>	۲ ۲	Ч с	. Middle cite estimate alus equation adi · TRO · Middle cite estimate alus low-flow measurement
IOWEL (W W L) middle (W/W/M)	11	1.1) () (t c i c	$T_{1} = \sum_{i=1}^{n} (i \in \mathbb{N}) $
$\frac{1111000}{10000000000000000000000000000$	11	0.0	0.0 1	C.2	Opper sue estimate plus equation adj. $\mathbf{1r}\mathbf{Q}_{s}$. Opper sue estimate plus low-mon measurement
(por c) rada	10	0.0	C .7	7.1	Continuous record gaging station
wailuanui lower (WL)	11	6.7	2.7	2.3	Middle site estimate plus equation adj.; TFQ: Middle site estimate plus low-flow measurements
middle (5210)	10	6.1	2.5	2.0	Continuous record gaging station plus upper sites estimates
upper (5200)	4.4	2.5	1.0	0.90	Continuous record gaging station
(URIC) Topped	3.2	2.0	0.90	0.80	Continuous record gaging station

at sites downstream of unquan	tified but know	n losing reach	es; g.s., gagin	g station; adj.	tow duration of base frow, an frows are in cubic feet per second, numbers in out tunt . are considered maximums adjustment]
Stream location	TFO ₅₀	BFO50	TFO ₉₅	BFO ₉₅	Source of estimate
Waiokomilo lower (WoL)	14	8.7	6.8	6.8	Regression equation: TFO :measurement plus upper site estimate
middle (WoM)	10	6.1	5.4	5.4	Regression equation; TFQ _a ::measurement plus upper site estimate
upper (WoU)	7.0	3.9	1.3	1.1	Regression equation
lower (OL)	4.7	4.7	4.7	4.7	Spring measurements
Palauhulu	17	11	4 3	4.0	Recreasesion equation: TEO – REO - measurements at caring station also more sites estimates
Iower (5220) middle (PhM)	14	9.3	1.9	1.6	Acgression equation: 17°C₉₅, DTC ₉₅ , incasurements at gazing station puts upper sites contributes Regression equation: TFO BFO measurements plus upper sites estimates
upper (KoU)	4.5	2.5	1.0	0.82	Regression equation
upper (HWU)	1.6	0.93	0.88	0.75	Regression equation
lower (PiL) middle (PiM)	No estima	tes availabl	e due to lac	k of data	
upper (P1U) Nuaailua					
lower (NL)	9.9	7.4	3.3	3.3	Regression equation; \mathbf{TFQ}_{95} , \mathbf{BFQ}_{95} : middle site estimate plus equation adj.
middle (NM) unner (NU)	3.9	2.5	0.50	0.40	Regression equation; TFQ_{05} , BFQ ₉₅ : upper site estimate plus low-flow measurement
appen (100)	00.0	0.28	0.19	CI.U	Kegression equation
Honomanu lower (HnL)	15	9.0	1.1	0.70	Middle site estimate plus equation adj.; TFQ ₉₅ ; BFQ ₉₅ : Middle site estimate plus low-flow measure-
Middle (HnM)	11	6.7	1.1	0.70	Upper site estimate plus equation $adj_{:}$; TFQ_{95} ; BFQ_{95} ; Upper site estimate plus low-flow measure-
upper (5270)	5.7	2.8	1.1	0.70	ment Continuous record gaging station
Punalau		2		0	
lower (PIL) middle (PIM)	6.5 5.8	4.5 3.9	2.3 2.1	2.2 2.0	Regression equation Regression equation; TFQ_s: Jow-flow measurement
Haipuaena lower (HaL)	9.9	5.5	2.0	1.6	Middle-lower site estimate plus equation adj.; TFQ ₉₅ , BFQ ₉₅ : Middle-lower site estimate plus low-
middle lower (HaML)	8.9	4.9	2.0	1.6	Itow measurement Middle-upper site estimate plus equation ad_j ; TFQ_{95} , BFQ_{95} ; Middle-upper site estimate plus low-
middle upper (HaMU)	8.0	43	1.9	1.5	flow measurement Unner site estimate plus equation adi.
upper (5360)	6.8	3.6	1.7	1.3	Continuous record gaging station plus upstream gaging station
Puohokamoa lower (PL)	17	11	3.6	3.1	Middle-lower site estimate plus equation adj.; TFQ ₉₅ , BFQ ₉₅ : Middle-lower site estimate plus low-
middle lower (PML)	16	10	3.6	3.1	Middle-upper site estimate plus equation adj.; TFQ_{95} , BFQ_{95} ; Middle-upper site estimate plus low-
middle upper (PMU)	7	6 1	3 1		TIOW measurement TImmer site astimute admotion adi
upper (5450)	12	6.4	2.5	2.1	Continuous record gaging station
Wahinepee lower (WpL)	2.4	1.8	1.1	1.1	Middle site estimate plus equation adi.
middle (ŴpM)	1.3	06.0	0.50	0.50	Regression equation; TFQ ₅₅ , BFQ ₅₅ ; low-flow measurement

Most-reliable Estimates of Natural Flow-Duration Statistics 61

Table 11. Estimates of natural (undiverted) streamflow statistics for gaged and ungaged basins, northeast Maui, Hawaii—Continued

[TFQx is the xx-percent flow duration of total streamflow; BFQxx is the xx-percent flow duration of base flow; all flows are in cubic feet per second; numbers in **bold italic** are considered maximums at sites downstream of unquantified but known losing reaches; g.s., gaging station; adj., adjustment]

	bs Source of estimate	Middle-lower site estimate plus equation $adj.$; TFQ_{95} , BFQ_{95} : Middle-lower site estimate plus low-flow measurements	Middle-upper site estimate plus equation $adj.$; TFQ_{95} ; Middle-upper site estimate plus low-flow measurements	Upper sites estimates plus low-flow measurements	Continuous record gaging station	Continuous record gaging station	Regression equation; \mathbf{TFQ}_{95} , \mathbf{BFQ}_{95} : Middle site estimate plus equation adj.	Regression equation; TFQ ₉₅ , BFQ ₉₅ : low-flow measurement
	BFQ ₉₅	2.0	1.9	1.9	0.60	0.80	0.60	0.20
0.0.0.	TFQ ₉₅	2.8	2.8	2.6	0.70	1.1	09.0	0.20
0	BFO ₅₀	7.0	6.7	6.6	1.5	3.5	3.4	2.5
	TFQ ₅₀	13	13	12	2.7	7.0	4.8	3.6
	Stream location	Waikamoi lower (WiL)	middle lower (WiML)	middle upper (WiMU)	upper (5550)	(accc) raddn	Kolea lower (KaL)	middle (KaM)

combined with the estimate developed from the regression equation at the upper site $(1.3 + 5.9 - 1.8 = 5.4 \text{ ft}^3/\text{s})$. At the lower site, additional gains $(0.12 \text{ ft}^3/\text{s})$ and a tributary inflow $(1.3 \text{ ft}^3/\text{s})$ increase the undiverted flow estimate to 6.8 ft}^3/\text{s} [5.2 ft}^3/\text{s}) were diverted from the stream for taro at the time of the measurements (Gingerich, 1999)]. The estimate for TFQ₉₅ at the upper site appears high on the basis of reconnaissance observations on March 5, 2003, when the stream at the upper site was observed to be nearly dry, yet the average daily flow at the index station was 5.8 ft}^3/\text{s} (about BFQ₅₀). However, because no other measurements are available to support this observation, the estimate from the regression equation was not adjusted.

Ohia Stream is fed almost entirely by Ohia Spring, which discharges about 4.7 ft³/s (Stearns and Macdonald, 1942), yet during reconnaissance on March 12, 2003 (about TFQ₈₅ at index station), the stream was nearly dry at the mouth. Streamflow is lost to evapotranspiration through watercress agriculture and through infiltration to the subsurface where stream channel modifications have filled the natural channel with soil and vegetation. The regression-equation-derived estimates for this stream basin are not reliable because rainfall (2.2 ft³/s) and MAXELEV (413 ft) values for this basin are outside the ranges used to develop the equations (table 8). Therefore, the most-reliable estimates are based on the average spring discharge, and are considered maximum values because of the unquantified streamflow losses between the spring and the coast.

Palauhulu Stream also has sections of losing channel and was dry between 800 and 300 ft altitude during reconnaissance on March 12–13, 2003 (about TFQ_{85} at index station). Therefore, estimates of TFQ_{95} and BFQ_{95} at the middle site (about 500 ft altitude) are maximum values, assuming all flow at the upper sites reaches the middle site. Estimates of TFQ_{95} and BFQ_{95} at the lower site are based on a continuous record gaging station (5220), which measured diversions from the stream, plus flow from the middle and upper sites assuming that all flow from those sites would reach the lower site.

Estimates of flow-duration statistics for Piinaau Stream determined from the regression equations are the highest of any sites in the study area (table 10), yet the flow observations, although scarce, indicate that flows are much lower than estimated. The stream channel was dry between 1,200 ft and 600 ft altitude during reconnaissance on March 14, 2003 (about TFQ₉₁ at index station), and only a trickle of flow was observed upstream of the 1,300-ft diversion. A recent (2001) large landslide, which covered the stream at about 1,000 ft altitude and filled most of the stream channel downstream to 600 ft altitude with gravel, cobbles, and boulders, complicates flow in the stream. This basin has the highest rainfall and MAXELEV in the study area and both are above the range of characteristics used to develop the flow-duration regression equations. Because the regression equations are not valid for this stream and reliable flow measurements are lacking, no estimates of stream statistics were made for Piinaau Stream sites.

Information about gaining or losing reaches on Nuaailua or Punalau Streams is not available and therefore estimates were made using the regression equations and single low-flow measurements.

Honomanu Stream is known to lose flow and be mainly dry downstream of the diversion at 1,700 ft altitude (Gingerich, 1999). However, the amount of streamflow lost has not been quantified. Therefore, the estimates of TFQ_{95} and BFQ_{95} at the middle and lower sites (table 11) are maximum values, assuming all natural flow at the upper site (station 5270) reaches the middle and lower sites. The estimates of TFQ_{50} and BFQ_{50} are determined solely from the regression equations and are expected to be overestimates but by an unknown amount.

Losses of flow were observed in the lower reaches of Waikamoi Stream and are expected but have not been quantified in the lower reaches of Haipuaena, Puohokamoa, Wahinepee, and Kolea Streams. Therefore, estimates at ungaged sites in the lower reaches of these streams are expected to be high.

Estimates of Flow-Duration Statistics under Diverted Conditions

Estimates of flow-duration statistics for diverted streams were made for gaged and ungaged sites on 21 streams in the study area downstream from the main diversion systems on the basis of a combination of continuous-record gaging-station data, low-flow measurements, and values determined from the regression equations developed as part of this study (Table 12). It is assumed that the diversion systems remove all flow lower than TFQ₅₀ above 1,200 ft altitude in all diverted streams. The flow-duration statistics for the streams can be easily calculated by subtracting the flows above the diversion system from the estimated undiverted flows below the diversion system. For example, TFQ₅₀ at the lower site on West Wailuaiki Stream is calculated by subtracting the TFQ_{50} flow (from table 2) at the upper site (station 5180) from the estimated undiverted TFQ_{50} flow (from table 11) at the lower site (12 ft³/s - 10 ft³/s = 2 ft³/s). In this example, it is assumed that all of the 10 ft³/s median flow at the upper site is removed from the stream by the diversion just downstream from the upper site. Values of BFQ₅₀, TFQ₉₅, and BFQ₀₅ are similarly calculated to be 1.2, 1.0, and 0.3 ft³/s, respectively. West of Keanae Valley, where another diversion ditch at lower altitude (Spreckels) captures additional low flows, the flow statistics for diverted streams are even lower. For example, TFQ₅₀ at the lower site on Waikamoi Stream is calculated by subtracting the TFQ₅₀ regressionequation estimate at the middle upper site from the estimated TFQ_{50} flow at the lower site (10 ft³/s - 9.2 ft³/s = 0.80 ft³/s) to account for the diversion of 9.2 ft³/s by Spreckels Ditch. Values of BFQ₅₀, TFQ₉₅, and BFQ₉₅ are similarly calculated to be 0.50, 0.20, and 0.00 ft³/s, respectively. Flow-duration curves of natural and diverted flow for median and lower flows at the lower West Wailuaiki and lower Waikamoi sites illustrate the

Table 12. Estimates of diverted streamflow statistics and percent flow reduction for gaged and ungaged basins, northeast Maui, Hawaii.

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	HI	Q ₅₀	BF	Q ₅₀		Q <u>.</u>	BF	Q _%	
Stream location	Estimate	Percent reduction	Estimate	Percent reduction	Estimate	Percent reduction	Estimate	Percent reduction	Comments
Hanawi									
lower (HwL)	25	22	21	24	20	6	17	11	Diverted at Koolau Ditch
middle (5090)	19	33	19	21	16	16	16	16	Diverted at Koolau Ditch
upper (5080)	7.1	0	4.6	0	2.4	0	2.2	0	Not diverted
Kapaula								ć	
lower (KL)	3.2	61	2.6	54	2.2	37	2.1	45 7 1	Diverted at Koolau Ditch
middle (5110)	2.4	68	2.1	59	1.9	42	1.9	<u>, c</u>	Diverted at Koolau Ditch
upper (5100)	4.9	0	2.8	0	1.2	0	1.1	þ	Not diverted
Waiaaka									
lower (WaL)	1.4	0	1.1	0	0.73	0	0.70	0	Diverted at Koolau Ditch
middle (5130)	0.86	0	0.77	0	0.54	0	0.52	0	Diverted at Koolau Ditch
Paakea									
lower (PaL)	5.0	23	4.6	16	3.6	12	3.6	10	Diverted at Koolau Ditch
middle (5140)	4.0	27	3.8	19	3.0	14	3.0	12	Diverted at Koolau Ditch
upper (PaU)	1.5	0	0.90	0	0.50	0	0.40	0	Not diverted
Waiohue									
lower (WeL)	2.6	70	2.1	72	1.4	69	1.3	64	Diverted at Koolau Ditch
middle (WeM)	1.3	82	1.0	83	0.80	<i>4</i>	0.70	77	Diverted at Koolau Ditch
upper (5150)	6.2	0	5.0	0	3.0	0	2.9	0	Not diverted
Kopiliula									
lower (KpL)	4.7	69	2.8	71	1.7	69	1.3	00 70	Diverted at Koolau Ditch
middle (KpM)	2.0	80	1.2	82	0.50	85	0.50	0/	Diverted at Noolau Duch
middle (PuM)	1.7	53	1.1	50	0.60	45	0.34	70	Diverticu at Noviau Duch Not diverted
upper (5160)	8.0	0	5.0	0	2.4	0	2.0		Not diverted
upper (PuU)	1.9	0	1.1	0	0.60	0	0.50	>	
East Wailuaiki									
lower (EWL)	2.4	78	1.5	79	0.40	88	0.30	90	Diverted at Koolau Ditch
middle (EWM)	1.5	86	1.0	85	0.20	94	0.20	93	Diverted at Koolau Ditch
upper (5170)	9.1	0	5.8	0	2.9	0	2.5	0	Not diverted

	TF	Q_{50}	BF	Q_{50}	TF	Q ₉₅	BF	Q ₉₅	
Stream location	Estimate	Percent reduction	Estimate	Percent reduction	Estimate	Percent reduction	Estimate	Percent reduction	Comments
West Wailuaiki									
lower (WWL)	1.9	84	1.2	83	0.30	91	0.30	87	Diverted at Koolau Ditch
middle (WWM)	1.2	89	0.80	88	0.20	94	0.20	91	Diverted at Koolau Ditch
upper (5180)	10	0	6.0	0	2.5	0	2.1	0	Not diverted
Wailuanui								t	
lower (WL)	1.7	85	1.1	85	0.81	70	0.56	9/	Diverted at Koolau Ditch
middle (5210)	1.6	84	1.0	84	0.39	84	0.32	84	Diverted at Koolau Ditch
upper (5200)	4.4	0	2.5	0	1.0	0	06.0		
upper (5190)	3.2	0	2.0	0	06.0	0	0.80	D	inot diverted
Ohia									
lower (OL)	No esti	imates available	e due to lack of	data					
Palauhulu	t	;		ì		ì	,		-
IOWET (5220) middle (DhM)	7.0	55 44	4.8	56 36	1.9	56 100	1.6	60 100	Taro diversion Locing stream
upper (KoU)	4.5	0	5.9	0	1.0	0	0.82	0	Not diverted
upper (HWU)	1.6	0	2.5 0.93	0	0.88	0	0.75	0	Not diverted
Diinaan									
lower (PiL) middle (PiM)	No esti	mates available	due to lack of	data					
upper (P1U)									
Nuaailua	ć	,	Ţ	-	-				
IOWET (INL) middle (NM)	9.5 3.3	0 15	7.1 2.2	4 21	$\frac{3.1}{0.30}$	0 40	$\frac{3.1}{0.25}$	0 37	Diverted at Spreckels Ditch
upper (NU)	0.56	0	0.28	0	0.19	0	0.15	0	Not diverted
11									
nonomanu lower (HnL)	87	64	5 8	36	0.00	100	000	100	Diverted at Snreckels Ditch
middle (HnM)	5.7	1 84	3.8	5 6 4	0.00	100	0.00	100	Diverted at Spreckels Ditch
upper (5270)	5.7	0	2.8	0	1.1	0	0.70	0	Minor upstream diversion

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Table 12. Estimates of diverted streamflow statistics and percent flow reduction for gaged and ungaged basins, northeast Maui, Hawaii—Continued

[TFQxx is the xx-percent flow duration of total streamflow; BFQxx is the xx-percent flow duration of base flow; percent reduction is relative to undiverted flow at the same location; all flows are in cubic feet per second; numbers in bold italic are considered maximums at sites downstream of unquantified but known losing reaches]

						0	F ₂		
	HL	Q_{s0}	BF	Q_{s0}	TF	Q ₅₅	BF	Q. ₅	
Stream location	Estimate	Percent reduction	Estimate	Percent reduction	Estimate	Percent reduction	Estimate	Percent reduction	Comments
Punalau lower (PIL) middle (PIM)	0.80 5.8	88 0	0.60 3.9	87 0	0.20 2.1	94 0	0.20 2.0	91 0	Diverted at Manuel Luis Ditch
Haipuaena lower (HaL) middle lower (HaML) middle upper (HaMU) upper (5360)	1.9 1.0 6.8	81 89 85	1.1 0.50 0.80 3.6	80 90 81 0	0.10 0.10 0.20 1.7	95 95 0	0.10 0.10 0.20 1.3	94 94 0	Diverted at Manuel Luis Ditch Diverted at Spreckels Ditch Diverted at Wailoa Ditch Minor upstream diversion
Puohokamoa lower (PL) middle lower (PML) middle upper (PMU) upper (5450)	3.0 2.0 3.0	82 87 0	2.1 1.1 2.0 6.4	81 89 76 0	0.40 0.40 0.70 2.5	88 77 0	0.40 0.40 0.60 2.1	87 78 0	Diverted at Manuel Luis Ditch Diverted at Spreckels Ditch Diverted at Wailoa Ditch Minor upstream diversion
Wahinepee lower (WpL) middle (WpM)	1.1 1.3	54 0	06.0 06.0	50 0	0.60	83 0	0.60	83 0	Diverted at Manuel Luis Ditch
Waikamoi lower (WiL) middle lower (WiML) middle upper (WiMU) upper (5570) upper (5550)	0.8 0.4 2.3 7.0	94 97 0 0	0.5 0.2 1.6 3.5	93 97 0 0	0.2 0.2 0.80 0.70 1.1	93 93 0	0.00 0.00 0.50 0.60 0.80	100 100 74 0	Diverted at Manuel Luis Ditch Diverted at Manuel Luis Ditch Diverted at Wailoa Ditch Not diverted Minor upstream diversion
Kolea lower (KaL) middle (KaM) Average of diverted	1.2 3.6	75 0 58	0.9 2.5	71 0 55	0.40	33 0 60	0.40 0.20	33 0 58	Diverted at Manuel Luis Ditch
SILES									



Figure 17. Estimated low-flow duration curves of natural and diverted streamflow at lower West Wailuaiki and Waikamoi Streams, northeast Maui, Hawaii.

significant effects of the diversion on streamflow statistics (fig. 17).

Estimated flow-duration curves are made by connecting the four estimated flow statistics (TFQ_{50} , BFQ_{50} , TFQ_{95} , and BFQ_{95}) with a line assuming a linear relation. The values of BFQ_{50} and BFQ_{95} are plotted at the average equivalent total-flow durations from table 5 (72 and 97 percent, respectively). Inspection of the flow-duration curves in fig. 7 shows that nearly all the curves can be represented by a straight line through TFQ_{50} , BFQ_{50} , TFQ_{95} , and BFQ_{95} so that it is reasonable to synthesize the estimated curves using a straight line to estimate the missing flow-duration points.

In addition to diversions above 1,200 ft altitude, three streams in the study area (Wailuanui, Waiokomilo, and Palauhulu Streams) also are diverted at lower altitudes for taro cultivation. On August 8, 2002, the taro diversion at about 200 ft altitude on Wailuanui Stream was observed to divert all water from the stream and then subsequently allow a significant amount of the diverted water to return to the stream channel downstream. An additional unmeasured amount of flow was observed to enter the Wailuanui stream channel downstream of the taro cultivation. Therefore, the divertedflow statistics for the lower site at Wailuanui were estimated by assuming 50 percent of the diverted flow at about 200 ft altitude is lost to the stream although measurements to confirm this assumption are not available.

Three taro diversions take water from Waiokomilo Stream downstream of 540 ft altitude including one at 440 ft altitude that takes all low flows (Gingerich, 1999). Flow at the middle Waiokomilo Stream site is affected by one diversion of relatively constant volume of about 0.40 ft³/s. Additionally, flow at the lower site is reduced by at least 3.7 ft³/s at the 440ft altitude diversion, 1.1 ft³/s at the 220-ft altitude diversion, and an unknown amount of flow is "lost" to individual private water diversions at about 250 ft altitude. These estimates of diverted volumes are determined from a set of low-flow measurements reported in Gingerich (1999). Diverted flow statistics at the lower site were calculated on the assumption that all low flows are removed at 1,300 ft and again at 440 ft altitude (table 12). Because the equations for TFQ_{50} and BFQ_{50} estimate less flow than has been measured for TFQ₉₅, the values for TFQ_{50} and BFQ_{50} were increased to the value of TFQ_{95} and should be considered a minimum estimate.

One taro diversion takes water from Palauhulu Stream at about 50 ft altitude several hundred feet upstream from the stream mouth. A continuous-record gaging station (5220) was operated on the diversion during 1934-68 and during that time, the statistics for flow in the diversion were 3.4, 3.0, 2.4, and 2.3 ft³/s for TFQ₅₀, BFQ₅₀, TFQ₉₅, and BFQ₉₅, respectively. These values were subtracted from the estimated natural flow statistics at the lower Palauhulu Stream site to obtain estimates of diverted flow statistics.

Estimated total reductions in streamflow due to diversions in the study area average 58 percent for TFQ_{50} , 55 percent for BFQ_{50} , 60 percent for TFQ_{95} , and 60 percent for BFQ_{95} . The streams with the lowest relative reduction in

streamflow are mostly those in the eastern side of the study area, where springs discharge below the main diversion (fig. 18 and table 12).

Needs for Additional Data

Additional data are needed to improve and confirm the estimates of median and low-flow statistics in the study area. Estimates of flow lost in losing streams would allow better definition of flow statistics in these streams. In reaches where streams go dry, additional water is needed in the reaches to permit accurate measurements of flow losses. Such additional water sources include allowing natural streamflow to bypass the diversions or releasing water from a diversion system into the stream. Streams for which the reliability of flow statistics would be most improved by the data provided by additional measurements include Ohia, Palauhulu, and Piinaau Streams in Keanae Valley; Honomanu Stream; and the lower reaches of streams west of Honomanu.

Summary and Conclusions

Median and low-flow statistics were estimated for streams in northeast Maui, Hawaii, and analyses were made to develop and evaluate the methods used to estimate the statistics. Estimated flow statistics are presented for continuousrecord gaging sites and for other sites where various amounts of streamflow data are available, as well as for locations where no data are available.

Records of daily mean flows were used to determine flow-duration, low-flow frequency, and base flow statistics for continuous-record stream-gaging stations in the study area following US Geological Survey established standard methods. Duration discharges of 50- and 95-percent were determined from total-flow and base-flow data for each continuous record. In order to compare streamflow records to each other, records were adjusted to concurrent periods, so that differences between the records were due to differences in climatic or drainage-basin characteristics and not to the fact that the records cover different times. The index-station method was used to adjust all of the streamflow records to a common period with the gaging station on West Wailuaiki Stream (5180), which was chosen as the index station because of its record length (1914–2003) and favorable geographic location near the middle of the study area. Adjusting the record length resulted in average decreases to the 50-percent duration total flow, 50-percent duration base flow, 95-percent duration total flow, and 95-percent duration base flow of 7, 3, 4, and 1 percent, respectively. In general, the largest adjustments were needed for the records with the shortest lengths.

For the drainage basin of each continuous-record gaged site and selected ungaged sites, morphometric, geologic, soil, and rainfall characteristics were quantified using GIS tech-



Figure 18. Reduction in Q₉₅ total flow (TFQ₉₅) due to diversions at selected ungaged and gaged sites in the study area, northeast Maui, Hawaii.

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niques. Regression equations relating the streamflow statistics to basin characteristics of the gaged basins were developed using ordinary-least-squares regression analyses. Rainfall rate, maximum basin elevation, and the elongation ratio of the basin were the basin characteristics used in the final regression equations for 50-percent duration total flow and base flow. Rainfall rate and maximum basin elevation were used in the final regression equations for the 95-percent duration total flow and base flow. The proportion of the variation in the dependent variable that is explained by the independent variables (R^2) ranged from 94.9 to 75.3 percent, with the highest flows having the highest R^2 . Standard errors of prediction ranged from 20.9 to 56.5 percent, with the highest flows having the lowest errors. The relative errors between observed and estimated flows ranged from 11 to 20 percent for the 50-percent duration total flow and from 29 to 56 percent for the 95-percent duration total flow and base flow.

The regression equations developed for this study were used to determine the 50-percent duration total flow, 50-percent duration base flow, 95-percent duration total flow, and 95-percent duration base flow at selected ungaged sites within the study area and at three gaging stations west of the study area using the appropriate basin characteristics. Estimated streamflow, prediction intervals, and standard errors were determined for 47 ungaged sites in the study area and four gaging stations west of the study area. Relative errors were determined for sites for which observed values of 95-percent duration discharge of total flow were available. East of Keanae Valley, the 95-percent duration discharge equation generally underestimated flow, and within and west of Keanae Valley, the equation generally overestimated flow.

Finally, most-reliable estimates of natural (undiverted) and diverted streamflow flow-duration statistics at gaged and ungaged sites on 21 streams in the study area were made using a combination of continuous-record gaging-station data, low-flow measurements, and values determined from the regression equations developed as part of this study. Average reduction in the low flow of streams due to diversions ranges from 55 to 60 percent.

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Appendix A: Selected Drainage-Basin Characteristics Quantified Using Basinsoft

[Descriptions modified from Harvey and Eash (1996) for purposes of regression analyses of streamflow data from northeast Maui, Hawaii]

BL-Basin length, in miles, measured along a line areally centered through the drainage divide from basin outlet to where the main channel extended meets the basin divide.

BP-Basin perimeter, in miles, measured along entire drainage-basin divide.

- **BR**-Basin relief, in feet, measured as the difference between the elevation of the highest grid cell and the elevation of the grid cell at the basin outlet, BR = MAXELEV-MINELEV
- **BS**-Average basin slope, in feet per mile, measured by the "contour-band" method, within the drainage area (DA). BS = (total length of all selected elevation contours) (contour interval) / DA.
- **BW**-Effective basin width, in miles, BW = DA/ BL.
- CCM-constant of channel maintenance, in square miles per mile, CCM=DA/TSL=1/SD.
- **CR**-Compactness ratio, dimensionless, the ratio of the perimeter of the basin to the circumference of a circle of equal area, $CR = BP/2 (\pi DA)^{0.5}$.

DA-Drainage area, in square miles.

- DF- Drainage frequency, in number of first-order streams per square mile
- **ER**-Elongation ratio, dimensionless, ratio of (1) the diameter of a circle of area equal to that of the basin to (2) the length of the basin, ER = $[4 \text{ DA}/\pi \text{ (BL)}^2]^{0.5} = 1.13 \text{ (1/SF)}^{0.5}$
- MAXELEV- maximum basin elevation, in feet
- MCL-Main channel length, in miles, measured along the main channel from the basin outlet to where the main channel, if extended, meets the basin divide.
- **MCS**-Main-channel slope, in feet per mile, an index of the slope of the main channel computed from the difference in streambed elevation at points 10 percent (E_{10}) and 85 percent (E_{85}) of the distances along the main channel from the basin outlet to the basin divide. MCS = ($E_{85} E_{10}$) / (0.75 MCL).

MCSP-Main channel slope proportion, dimensionless, MCSP=MCL/(MCS)^{0.5}.

MCSR-Main-channel sinuosity ratio, dimensionless, MCSR = MCL / BL.

MINELEV- minimum basin elevation, in feet

RB-Rotundity of basin, dimensionless, $RB = [\pi (BL)^2]/[4 DA] = 0.785$ SF.

RN- Ruggedness number, in feet per mile, RN = (TSL)(BR)/DA

RR-Relative relief, in feet per mile, RR = BR/BP.

RSD- Relative stream density, dimensionless, $RF = DF/(SF)^2$

SD-Stream density, in miles per square mile, SD = TSL / DA.

SF-Shape factor, dimensionless, ratio of basin length to effective basin width, SF = BL/BW.

SR- Slope ratio of main-channel slope to basin slope, dimensionless, SR = MCS/BS

TSL-Total stream length, in miles, computed by summing the length of all stream segments within the DA.

Reference

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