Estimating the Magnitude and Frequency of Floods in Rural Basins of North Carolina—Revised

By Benjamin F. Pope, Gary D. Tasker, and Jeanne C. Robbins

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PREFACE

This report revises and supersedes U.S. Geological Survey Water-Resources Investigations Report 99–4114. The revised flood-discharge values are listed in table 2 (Q2, Q5, Q10, Q25, Q50, Q100, Q200, and Q500). The revised flood discharges, for all recurrence intervals, vary by as much as 17 percent from the earlier published values, with 80 percent of all values within 7 percent of the earlier published data. Differences in the values for the 100-, 200-, and 500-year discharges are greater than in the values for the 2-, 5-, 10-, 25-, and 50-year discharges, 80 percent of which are within 3 percent of the earlier published values.

The revised *t*-year discharges were used to update the regional regression equations and the region-of-influence data base, as indicated in revised text tables 5, 6, and 7 and in appendix table 1. The maximum difference in computed results for the regional regression equations was noted for the Coastal Plain equations, where application of the revised equations to small drainage areas, less than 10 square miles, resulted in discharges that are about 3 to 9 percent greater than those values obtained using equations from the previous report. Computed flood discharges using the revised Blue Ridge-Piedmont equations generally were within about 2 percent of the values from the previously published equations, except for results for drainage areas less than 10 square miles, which ranged from about 3 to 7 percent less than the previously published values. Application of the revised regression equations to the Sand Hills hydrologic area shows results in discharges that are up to 3 percent less than those computed using the equations published in the earlier report. The average error of prediction for the revised equations was nearly the same as for the earlier published Blue Ridge-Piedmont equations, lower for the Coastal Plain equations, and higher for the Sand Hills equations.

As in the previous report, the root mean square error (RMSE) for the region-of-influence method was only marginally better than the RMSE reported for the regional regression equations, resulting in neither method being clearly superior. The revised computer program for computing the estimates of floodfrequency discharges, using either the regional regression equations or the region-of-influence method, and the associated site-specific errors of prediction are available at the North Carolina District Web site http://nc.water.usgs.gov/reports/ wri014207.

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Multiply	Ву	To obtain
	Length	
mile (mi)	1.609	kilometer
	Area	
square mile (mi ²)	2.590	square kilometer
	Flow	
foot per mile (ft/mi)	0.1894	meter per kilometer
cubic foot per second (ft^3/s)	0.02832	cubic meter per second

CONVERSION FACTORS and ABBREVIATIONS/ACRONYMS

Abbreviations/Acronyms:

BRP	Blue Ridge-Piedmont hydrologic area
BSLOPE	basin slope
СР	Coastal Plain hydrologic area
CSLOPE	channel slope
DA	drainage area
DEM	digital elevation model
GIS	geographic information system
L	channel length
REG	region variable
RMSE	root mean square error
SH	Sand Hills hydrologic area
SHAPE	basin shape
USGS	U.S. Geological Survey

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ABSTRACT

A statewide study was conducted to develop two methods for estimating the magnitude and frequency of floods in rural ungaged basins in North Carolina. Flood-frequency estimates for gaged sites in North Carolina were computed by fitting the annual peak flows for each site to a log-Pearson Type III distribution. As part of the computation of flood-frequency estimates for gaged sites, new values for generalized skew coefficients were developed. Basin characteristics for these gaged sites were computed by using a geographic information system and automated computer algorithms. Flood-frequency estimates and basin characteristics for 317 gaged sites were combined to form the data base that was used for this analysis.

Regional regression analysis, using generalized least-squares regression, was used to develop a set of predictive equations that can be used to estimate the 2-, 5-, 10-, 25-, 50-, 100-, 200-, and 500-year recurrence interval discharges for rural ungaged basins in the Blue Ridge-Piedmont, Coastal Plain, and Sand Hills hydrologic areas. The predictive equations are all functions of drainage area. Average errors of prediction for these regression equations range from 36 to 65 percent.

A region-of-influence method also was developed that interactively estimates recurrence interval discharges for rural ungaged basins in the Blue Ridge-Piedmont and Coastal Plain hydrologic areas of North Carolina. Regression techniques are used to develop a unique relation between flood discharge and basin characteristics for a subset of gaged sites with similar basin characteristics. This, then, can be used to estimate flood discharges at ungaged sites. Because the computations required for this method are somewhat complex, a computer application was developed that performs the computations and compares the predictive errors for this method. The computer application also includes the option of using the regression equations to compute estimated flood discharges and errors of prediction specific to each ungaged site.

Root mean square errors, computed for each recurrence interval and hydrologic area, are generally only slightly lower for the region-ofinfluence method than for the regression equations and do not provide sufficient basis for recommending one method over the other. In addition, the region-of-influence method is a new method that is still being improved. As a result, the regional regression equations are considered to be the primary method for computing floodfrequency estimates at ungaged sites.

INTRODUCTION

Reliable estimates of the magnitude and frequency of floods are needed by State and local designers and managers. The design of highway and railroad stream crossings, delineation of flood plains and flood-prone areas, management of water-control structures, and management of water supplies are all activities that require estimates of the frequency distribution of flood events. Such estimates can be computed directly by using statistical methods at gaged sites that have at least 10 years of annual peak record; the longer the record of annual peak flows, the more reliable the estimate. It is not feasible, however, to collect 10 years of annual peak record for every location where an estimate of the flood-frequency distribution is needed, nor is it reasonable to wait 10 years for an estimate once a site has been identified.

Estimates that are derived solely from gage records do not provide sufficient spatial coverage to satisfy the need for reliable estimates of the magnitude and frequency of floods. Traditionally, to meet this need, annual peak records at gaged sites have been regionalized, or extended in space. By this process, flood-frequency estimates at gaged sites are related to measurable basin characteristics so that reliable floodfrequency estimates can be made at ungaged sites. In response to the need to improve the accuracy of estimates of flood discharges for ungaged rural basins, the U.S. Geological Survey (USGS), in cooperation with the North Carolina Department of Transportation, initiated an investigation in 1996 to further define the relation between flood discharges of selected recurrence intervals and selected basin characteristics for rural North Carolina basins.

In the past, regionalization was achieved by means of regional regression analysis. Data from gaged sites were used to define a set of relations between selected recurrence interval discharges and drainage area. Once defined, these relations were then used to estimate discharges at selected recurrence intervals for ungaged sites. Often the area of study was subdivided into regions of similar hydrology in order to improve the predictive ability of the equations. Gunter and others (1987) used this approach to develop regional relations for estimating the magnitude and frequency of floods in rural North Carolina basins.

Recently, however, a different approach to regionalization has been developed. This new approach, known as the region-of-influence method, interactively estimates recurrence interval discharges for ungaged sites based on data from gaged sites with similar basin characteristics. For each ungaged site selected, a subset of gaged sites having similar basin characteristics is selected from the entire data base of rural gaged sites. Regression techniques are used to develop a unique relation between flood discharge and basin characteristics for this subset of gaged sites. This relation is then used to estimate flood discharges at the ungaged site. Although computationally intensive, the region-of-influence method is easily automated and performed by a computer application that is discussed later in this report. Because only gaged sites with similar basin characteristics are used to estimate flows at ungaged sites, there is less chance of extrapolation beyond the limits of the explanatory data. Tests of this approach in Texas (Tasker and Slade, 1994) and in Arkansas (Hodge and Tasker, 1995) yielded estimates with lower prediction errors than those produced by using traditional regional regression techniques.

Gunter and others (1987) contains annual peak-flow data collected from gages throughout North Carolina through the 1984 water year¹, whereas this report contains peak-flow data collected through the 1996 water year. Thus, gaged sites that have continued in operation since 1984 have as much as 12 additional years of peak-flow data available for computation of flood-frequency estimates. The 12 intervening years (1985–96) include several years of pronounced drought (1985–88) as well as years in which maximum peaks of record were recorded (1992–93, 1996) for North Carolina streams. In addition, 64 gaged sites that were not used in Gunter and others (1987) are now available for analysis.

Purpose and Scope

This report describes the development, application, and evaluation of two methods for estimating the magnitude and frequency of floods at ungaged, unregulated, rural basins in North Carolina—(1) the regional regression method and (2) the region-of-influence method. A comparison of these two methods, based on their predictive ability and ease of application, also is presented. In order to compare the two methods on an equal basis, each method was applied to the same available data. The regional regression and region-of-influence methods of estimation were applied to the current data base of 317 sites with at least 10 years of unregulated peak-flow record and evaluated.

Approach

A set of eight basin characteristics was computed and compiled for each of 366 gaged rural sites in North Carolina that have peak-flow record. Sites that have

¹Water year is the period October 1 through September 30 and is identified by the year in which it ends.

flows affected by regulation or channelization were identified, and where possible, records for such sites were divided into periods of unregulated and regulated flows. Weighted regional average skew values were used to compute flood-frequency estimates for 317 sites with at least 10 years of unregulated peak-flow record. Flood-frequency estimates and the computed basin characteristics for these 317 sites were combined to form the data base used in the regional analyses.

Generalized least-squares regression analysis was used to develop predictive equations relating the 2-, 5-, 10-, 25-, 50-, 100-, 200-, and 500-year recurrence interval flood discharges to selected basin characteristics for rural basins throughout North Carolina. In addition, a region-of-influence method was developed that interactively estimates the recurrence interval flood discharges for ungaged rural basins in the Blue Ridge-Piedmont and Coastal Plain hydrologic areas.

Computation and compilation of basin characteristics and of the selected recurrence interval discharges are described in the following sections. All aspects of each analysis, including the initial exploratory multiple regression analysis using ordinary least-squares regression, final regional regression using generalized least-squares regression, and the region-of-influence analysis, are described. Finally, a comparison of the results of each method is presented.

Data Compilation

The first step in the regionalization of floodfrequency estimates is the compilation of a list of all gaged sites with annual peak-flow record. Such sites are either continuous-record sites or crest-stage sites. At continuous-record sites, the water-surface elevation, or stage, of the stream is recorded at fixed intervals, typically ranging from 5 to 60 minutes. At crest-stage sites, only the crest, or highest, stages that occur between site visits, usually 6 to 8 weeks, are recorded. Regardless of the type of gage, measurements of discharge are determined throughout the range of recorded stages, and a relation between stage and discharge is developed for the gaged site. Using this stage-discharge relation, or rating, discharges for all recorded stages are determined. The highest peak discharge that occurs during a given year is the annual peak for the year, and the list of annual peaks is the annual peak-flow record. The three hydrologic areas identified and described by Gunter and others (1987),

consisting of (1) the combined Blue Ridge and Piedmont physiographic provinces, (2) the Coastal Plain Province, and (3) a subdivision of the Coastal Plain Province known as the Sand Hills, also were used in this study (fig. 1).

An initial list of 366 rural sites with annual peakflow record was compiled (fig. 1; table 1, p. 19–30). Records for these sites were then examined to determine the extent of available basin characteristic data and to identify sites with flows affected by channelization or regulation. The only consistently available basin characteristics for most sites were drainage area and location. A complete evaluation of all possible relations between flood discharges and other characteristics of rural basins requires a more complete set of basin characteristics. The computation and compilation of the required basin characteristics for all of the 366 initial sites are described in the following section.

Examination of the flow records for the 366 sites revealed 19 sites with record containing only regulated/ channelized flows, 27 sites with record that could be divided into periods of unregulated/unchannelized and regulated/channelized flows, and 320 sites with records unaffected by any known regulation/channelization. Of the 347 sites with at least some period of unregulated flow record, 317 sites had the requisite 10 or more years of record for computation of flood-frequency estimates (table 1). Flood-frequency estimates for these sites were computed and combined with the basin characteristics to form the data base that was used for the regional analyses (table 2, p. 31–42). This data base contained 222 sites in the Blue Ridge-Piedmont hydrologic area, 80 sites in the Coastal Plain hydrologic area, and 15 sites in the Sand Hills hydrologic area (table 2). Of the 46 sites with regulated flow records, floodfrequency estimates were computed for 42 sites with periods of regulated flow longer than 10 years but were not included in either regional analysis.

Acknowledgments

The authors gratefully acknowledge the assistance and support of Mr. Archie Hankins of the North Carolina Department of Transportation. The peak-flow data used in the analyses described herein were collected throughout North Carolina at stream gages operated in cooperation with a variety of Federal, State, and local agencies. The authors also would like to recognize the dedicated work of the USGS field

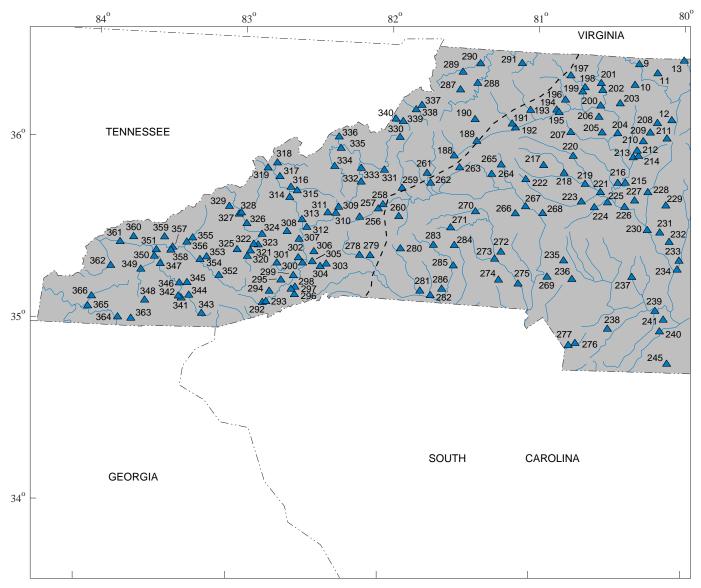
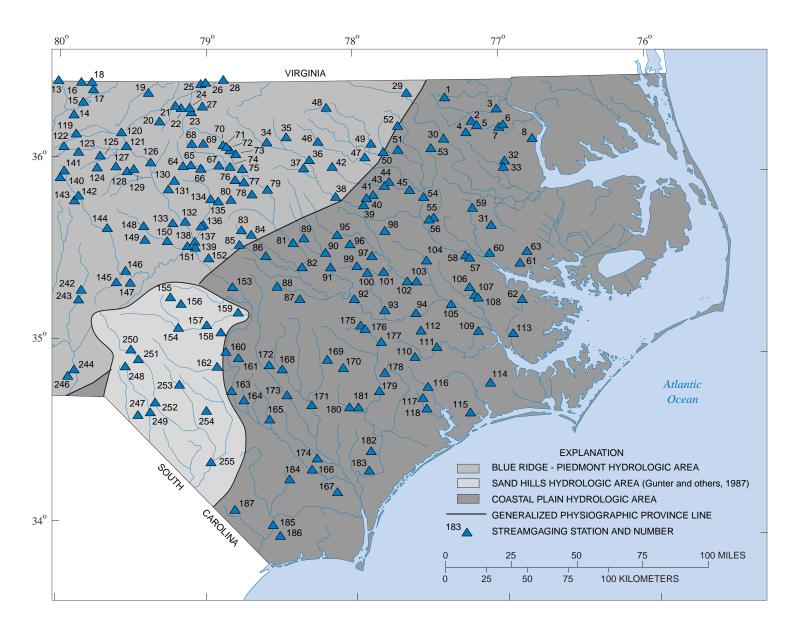


Figure 1. Locations of gaged rural sites in North Carolina.



office staff in collecting, processing, and storing the peak-flow data necessary for the completion of this report.

BASIN CHARACTERISTICS

The annual peak-flow data that were used in this study were collected at gages in rural basins from all areas of the State, representing the wide range of physical and climatic conditions that occur in North Carolina. Eight parameters that characterize the size, shape, relief, and climate of rural basins in North Carolina were computed and compiled for each site used in the study. Physical basin characteristics include drainage area (DA), channel length (L), channel slope (CSLOPE), basin slope (BSLOPE), and basin shape (SHAPE) (table 3). The primary climatic characteristics relevant to flood frequency in each basin are the intensity, duration, and amount of storm rainfall, as well as other meteorologic inputs that control evaporation and transpiration. Lichty and Liscum (1978) suggested the use of a regional climate factor, CF_t , where t = 2-, 25-, and 100-year recurrence intervals, that integrates long-term rainfall and pan evaporation information and represents the effect of these climatic influences on flood frequency. In this study, a refined version of CF_{t} , as developed and described by Lichty and Karlinger (1990), was used to characterize climatic effects of flood frequency. Climate factors, CF_t, for each site were computed by using a computer algorithm that used the maps of climate factor isolines presented in Lichty and Karlinger (1990) and the latitude and longitude of a site to interpolate values for the three climate factors, CF₂, CF₂₅, and CF₁₀₀.

The hydrologic area for each site was determined by examining drainage boundary maps. The appropriate integer value for each site was then assigned to the region variable (REG) (table 3).

Other than drainage area, the physical basin characteristics selected for use in this study were not readily available for most of the basins in the study. In previous studies, drainage area was the primary explanatory variable; thus, there was no prior need to measure or compute the other characteristics. As a result, the other physical basin characteristics had to be computed and compiled. Because of the large number of sites involved and the need for consistent, unbiased methodology in making measurements and computations, a geographic information system (GIS)

Table 3. Basin characteristics that were used in the North Carolina flood-frequency regionalization study

[mi², square mile; mi, mile; ft/mi, foot per mile; ----, a dimensionless characteristic]

Basin	Unit of	Definition			
characteristic	measure				
	•	characteristics			
DA	mi ²	Drainage area, measured area contained within basin divides.			
L mi Channel length, measured fro gage site upstream along main channel to basin divid					
CSLOPE ft/mi Channel slope, computed between points at 10- and percent of the length, measured from the gage si					
BSLOPE ft/mi Basin slope, mean value of slope measured along seve flow paths from basin divi to channel.					
drainage area by the sq		Shape, computed by dividing drainage area by the square of channel length (DA/L ²).			
	Climatic	characteristics			
CF ₂		2-year recurrence interval climate factor			
CF ₂₅		25-year recurrence interval climate factor			
CF ₁₀₀		100-year recurrence interval climate factor			
	Regior	al identifiers			
BRP		1, if site is in Blue Ridge- Piedmont; 0, if not.			
СР		 if site is in Coastal Plain; if not. 			
SH		1, if site is in Sand Hills; 0, if not.			
REG		 if site is in Blue Ridge- Piedmont; if site is in Coastal Plain; if site is in Sand Hills. 			

was used to compute the required physical basin characteristics.

In order to use GIS to develop basin characteristics, a digital elevation model (DEM) was created by combining individual data sets. These data sets included the U.S. Environmental Protection Agency River File 3 (McKay and others, 1994), USGS digital line graph contour lines (U.S. Geological Survey, 1989), and the National Oceanic and Atmospheric Administration shoreline data set (National Oceanic and Atmospheric Administration, 1999). Known drainage basin boundaries were overlain onto the DEM, and a combination of computer and visual interpolation techniques were used to define boundaries between the 366 gage sites and the known drainage boundaries.

Once the DEM was constructed and basin boundaries were delineated for all sites, a set of computer algorithms was developed to automatically compute drainage area, L, CSLOPE, BSLOPE, and SHAPE. Although GIS-computed drainage area was computed, the values used for DA were the drainage areas compiled from site records that were handcomputed and checked when the sites were established. The percent difference between GIS-computed drainage area and DA was automatically computed and used to verify the delineation of basin boundaries and the automated computations. Sites with greater than 10-percent difference between the computed drainage area and DA were flagged and re-examined. Errors in boundary delineation were corrected by comparing USGS 7.5-minute topographic maps with the original hand-delineated basin boundary and by using manual techniques to match the GIS basin boundary to the original. After adjusting basin boundaries, basin characteristics were recomputed and rechecked until satisfactory results were obtained. Several sites with drainage areas less than about 1 square mile (mi²) did not meet the criteria of less that 10-percent difference between computed drainage area and DA because the resolution of the GIS data and computational methods were about one-tenth of a square mile. These sites were examined manually to determine if the automated delineation of basin boundaries was consistent with the hand-drawn boundaries; if not, the boundaries were adjusted accordingly and basin characteristics were recomputed.

ESTIMATION OF FLOOD MAGNITUDE AND FREQUENCY AT GAGED SITES

Flood-frequency estimates for a given stream site are typically presented as a set of exceedance probabilities or, alternatively, recurrence intervals along with the associated discharges. Exceedance probability is defined as the probability of exceeding a specified discharge in a 1-year period and is expressed as decimal fractions less than 1.0 or as percentages less than 100. A discharge with an exceedance probability of 0.10 has a 10-percent chance of being exceeded in any given year. Recurrence interval is defined as the number of years, on average, during which the specified discharge is expected to be exceeded one time and is expressed as number of years. A discharge with a 10-year recurrence interval is one that, on average, will be exceeded once every 10 years. Recurrence interval and exceedance probability are the mathematical inverses of one another; thus, a discharge with an exceedance probability of 0.10 has a recurrence interval of 1/0.10 or 10 years. Conversely, a discharge with a recurrence interval of 10 years has an exceedance probability of one-tenth or 0.10. It is important to remember that recurrence intervals, regardless of length, always refer to the average number of occurrences over a long period of time; for example, a 10-year flood discharge is one that might occur about 10 times in a 100-year period, rather than exactly once every 10 years.

Flood-frequency estimates for gaged sites are computed by fitting the series of annual peak flows to some known statistical distribution. For the purposes of this study, estimates of flood-flow frequency are computed by fitting the logarithms (base 10) of the annual peak flows to a log-Pearson Type III distribution, following the guidelines and using the computational methods described in Bulletin 17B of the Hydrology Subcommittee of the Interagency Advisory Committee on Water Data (1982). The equation for fitting the log-Pearson Type III distribution to an observed series of annual peak flows is as follows:

$$Log Q_t = X + KS, \qquad (1)$$

where

- Q_t is the *t*-year recurrence interval discharge in ____ cubic feet per second,
- *X* is the mean of the log-transformed annual peak flows,
- *K* is a factor dependent on recurrence interval and the skew coefficient of the logtransformed annual peak flows, and
- *S* is the standard deviation of the log-transformed annual peak flows.

Values for *K* for a wide range of recurrence intervals and skew coefficients are published in Appendix 3 of Bulletin 17B (Hydrology Subcommittee of the Interagency Advisory Committee on Water Data, 1982).

Fitting the log-Pearson Type III distribution to the general case of a long, well-distributed series of annual peak flows is fairly straightforward. Often, however, a series of peak flows may include low or high outliers, which are extremely low or high peak flows that depart significantly from the trend in the data. The gage record also may frequently include information about maximum peak flows that occurred outside of the period of regularly collected, or systematic, record. Such peak flows, known as historic peaks, are often the maximum peak flows known to have occurred during an extended period of time, longer than the period of collected record. The interpretation of outliers and historic peak information in the fitting process can greatly affect the final flood-frequency estimate. Bulletin 17B (Hydrology Subcommittee of the Interagency Advisory Committee on Water Data, 1982) provides guidelines for detecting and interpreting these data points and provides computational methods for making appropriate corrections to the distribution to account for their presence. In some cases, high or low outliers are excluded from the record, so that the number of systematic peaks may not be equal to the number of years in the period of record.

Statistical measures, such as mean, standard deviation, or skew coefficient, can be described in terms of the sample or computed measure and the population or true measure. In terms of annual peak flows, the period of collected record can be thought of as a sample, or small portion, of the entire record, or population. Statistical measures computed from the sample record are estimates of what the measure would be if the entire population were known and used to compute the given measure. The accuracy of these estimates depends on the nature of the specific measure and the given sample of the population.

Skew coefficient measures the symmetry of the distribution of a set of peak flows about the median of the distribution. A peak-flow distribution with the mean equal to the median is said to have zero skew. A positively skewed distribution has a mean that exceeds the median typically as a result of one or more extremely high peak flows. A negatively skewed distribution has a mean that is less than the median typically because of one or more extremely low peak flows.

The computed skew coefficient for the peak-flow record of a given station is very sensitive to extreme events; therefore, the sample skew coefficient for short records may not provide an accurate estimate of the population skew. This is problematic because the *K*-factor in equation 1 for a given recurrence interval is dependent only on skew coefficient; therefore, an inaccurate skew coefficient will result in a flood-frequency estimate that is not representative of the true, or population, value.

A more accurate estimate of skew coefficient at a site can be obtained by using a weighted average of the sample skew coefficient estimate with a generalized, or regional, skew coefficient. A generalized skew coefficient is obtained by combining skew estimates from nearby, similar sites. A nationwide generalized skew study was conducted for the study documented in Bulletin 17B (Hydrology Subcommittee of the Interagency Advisory Committee on Water Data, 1982). Skew coefficients for long-term gage sites from all over the Nation were computed and used to produce a map of isolines of generalized skew. Gunter and others (1987) used this nationwide generalized skew in their flood-frequency computations. In addition, the USGS in North Carolina has computed other unpublished flood-frequency estimates by using the nationwide generalized skew.

During preliminary computations of floodfrequency estimates for inclusion in the regression analyses, a number of inconsistencies were noted between the computed values of sample skew coefficients at long-term gaging sites in North Carolina and the values obtained from the national generalized skew study. Inconsistencies at long-term sites are of concern because if generalized skew coefficients for a region are accurate estimates of the population skew, then the computed values of sample skew at long-term sites should approach the generalized values. Instead, it was noted that while sample skew coefficients at longterm North Carolina sites were somewhat consistent among themselves, they did not agree with the generalized values obtained from the nationwide generalized skew study. This anecdotal evidence, when considered along with the age and lack of resolution of the national study, was deemed sufficient cause to develop new generalized skew estimates for rural gaging sites in North Carolina.

Bulletin 17B (Hydrology Subcommittee of the Interagency Advisory Committee on Water Data, 1982) describes three methods for performing generalized skew studies using skew coefficients computed from long-term gaging stations—(1) plot computed skew

coefficients on a map and construct skew isolines, (2) use regression techniques to develop a skew prediction equation that would relate station skew coefficients to some set of basin characteristics, or (3) use the arithmetic mean of computed skew coefficients from long-term sites in the area. For the purposes of this report, a modification of the second method initially was decided to be the most likely method to produce satisfactory results. However, rather than using ordinary least-squares regression, a weighted least-squares regression technique was used to determine the relation between the sample skew coefficient and selected basin characteristics. Sample skew estimates were weighted according to their respective record length; sites with long records were assigned greater weight than those with short records. The use of this regression technique in this study made it possible for data from all 347 sites with unregulated flows to be used in developing the estimate.

Multiple regression analysis, using ordinary least-squares regression, was used to determine the best set of basin characteristics to use as explanatory, or independent, variables in the weighted least-squares predictive model. Initial analyses were somewhat disappointing; no combination of basin characteristics accounted for a significant amount of the variance in computed skew. Lacking any significant statewide relationship between sample skew and basin characteristics, three location variables-BRP, CP, and SH, one for each of the three hydrologic areas, Blue Ridge-Piedmont, Coastal Plain, and Sand Hills-were added to the analysis. For a given site, the location variable representing the region of the site was set at 1, and the other two location variables were set at 0 (table 3). When these variables were added to the multiple regression analysis, results were only marginally better. None of the exploratory multiple regression models yielded significant relations between sample skew and the basin characteristics.

Given the lack of satisfactory results in this attempt to develop predictive equations relating skew to a set of basin characteristics, it was decided to apply a modified version of the second method in Bulletin 17B (Hydrology Subcommittee of the Interagency Advisory Committee on Water Data, 1982). A regional regression prediction equation was developed using weighted least-squares regression (Tasker and Stedinger, 1986). Weights were assigned, according to record length, to the computed skews for each site. Because the only statistically significant explanatory variable in the regression analysis was an indicator variable for the Sand Hills hydrologic area, the regression equation predicts one value for all sites in the Sand Hills area and another value for all sites in the remaining hydrologic areas. These predictions are essentially a weighted average of the sites in each of the two areas and, therefore, can be considered a modified version of the third method as well. The two weighted regional average skew values, along with the standard error of prediction and the mean square error of prediction associated with each estimate (table 4), were determined by the methods described in Tasker and Stedinger (1986).

Table 4. Generalized skew coefficient and associated mean square error for rural North Carolina gaging sites

Hydrologic area	Generalized skew coefficient	Standard error	Mean square error
Blue Ridge-Piedmont and Coastal Plain	0.195	0.194	0.038
Sand Hills	0.252	0.250	0.062

As described previously, a weighted skew coefficient is used in order to improve the accuracy of the skew coefficient used to fit peak-flow records to a log-Pearson Type III distribution. The weighted skew coefficient for a given site is computed as the weighted average of the generalized skew coefficient and the site's computed skew coefficient, with weights assigned according to the mean square error of each component skew value. Flood-frequency estimates for all sites with unregulated flow records were computed by using the weighted skew method. Flood-frequency estimates for sites with regulated flow record were computed by fitting the recorded regulated peak flows to the log-Pearson Type III distribution. Computed sample skew coefficients for the regulated flow record were used because regulated peak-flow records typically are not representative of regional or generalized conditions. Although flood-frequency estimates for regulated sites are presented in this report, more detailed, site-specific analyses of flood frequency at many regulated sites are available from the U.S. Army Corps of Engineers.

ESTIMATION OF FLOOD MAGNITUDE AND FREQUENCY AT UNGAGED SITES

Two regional analyses were used to develop methods for estimating flood discharges for ungaged rural basins in North Carolina. The first analysis, a traditional regional regression, required the use of generalized least-squares regression to define a set of predictive equations that relate peak discharges for the 2-, 5-, 10-, 25-, 50-, 100-, 200-, and 500-year recurrence intervals to selected basin characteristics for unregulated rural basins in each of three hydrologic areas of North Carolina (fig. 1). The second analysis, the region-of-influence method, required the development of a computer application to derive, for any given ungaged rural site in the Blue Ridge-Piedmont or Coastal Plain hydrologic areas, unique predictive relations between the 2-, 5-, 10-, 25-, 50-, 100-, 200-, and 500-year recurrence interval discharges and selected basin characteristics. Just as in the traditional regional regression, generalized leastsquares regression is used to develop these predictive relations; however, in the region-of-influence analysis, regression techniques are applied to only a selected subset of gaged sites, rather than the entire data base of gaged sites.

Regional Regression Analysis

Ordinary least-squares regression with flood discharge as the dependent variable was used in exploratory analyses to determine the best regression models for all combinations of the eight basin characteristics that were used as explanatory variables. An additional goal of the exploratory analysis was to determine if the subdivision of the State into three hydrologic areas is supported by current data.

Initially, the regionalization scheme used by Gunter and others (1987), which divided the State into the Blue Ridge-Piedmont, Coastal Plain, and Sand Hills hydrologic areas, was assumed to still be valid. Multiple regression analysis, using Mallow's Cp (Stedinger and Tasker, 1985), adjusted coefficient of determination, and hydrologic judgment as criteria, resulted in one-variable and two-variable models relating flood discharge to basin characteristics for each of the three hydrologic areas. The most significant one-variable models for all three regions included drainage area only. The most significant two-variable models included drainage area and the 25-year climate factor for the Blue Ridge-Piedmont and Sand Hills hydrologic areas; while the best two-variable models for Coastal Plain sites consisted of drainage area and channel length.

The validity of the regionalization scheme was examined by performing additional ordinary leastsquares regression analyses by using the two-variable models determined previously and comparing the coefficients and intercepts for each region's model to those for the rest of the State. In each case, the coefficients and intercepts for each region's model differed from those of the model using the remaining sites in the State. Additionally, a further test was conducted by introducing the location variable (table 3) for each region into the regression model. Each of these variables was set either at 1, if the site was in a particular region, or 0, if not. A five-variable ordinary least-squares regression model, including all available sites and using (1) drainage area, (2) climate factor, (3) location variable, (4) the product of the location variable with drainage area, and (5) the product of the location variable with climate factor as explanatory variables, was constructed for each recurrence interval discharge in each of the three hydrologic area. For a given region's model, a significant coefficient for the location variable indicates a difference in the intercept between sites in that region and sites in the rest of the State; a significant coefficient for either of the terms that are products of a location variable and another variable indicates a difference in the coefficients of the basin characteristic in that term between sites in that region and the rest of the State. In this particular test, a 95-percent confidence level was defined as significant. All three regional models had significant coefficients for at least one of the location variables or location variable product terms. Given the results of these regression tests, the regionalization scheme used by Gunter and others (1987) was accepted.

Ordinary least-squares regression is an appropriate and efficient regression model for use when flow estimates that are used as response variables are independent of each other (no correlation exists between pairs of sites) and when the reliability and variability of flow estimates that are used as response variables are approximately equal. The flow estimates that were used in this regression were generated from peak-flow records at gaging stations in all parts of North Carolina with periods of record ranging from 10 to 101 years. Records from gaging stations on the same stream within the same basin or even in adjacent basins may be highly correlated because the peak flows resulted from the same rainfall events, similar antecedent conditions, and similar basin characteristics. However, records from other sites, in basins remote from each other, have varying degrees of correlation. In general, correlation between pairs of sites can be described as a function of distance between sites. Additionally, the reliability of flow estimates that were used as response variables in this regression is, in general, a function of record length and, as such, cannot be considered equal for all sites in the regression. Variability of the flow estimates, characterized by the standard deviation of the peak-flow record that was used to compute the flow estimate, depends in large part on characteristics of the basin and also cannot be considered equal for all sites used in the regression. For these reasons, ordinary least-squares regression was used only as an exploratory technique in this analysis to identify the best potential regression models and to evaluate the proposed regionalization scheme. The final regression equations were developed by using generalized least-squares regression techniques.

Generalized least-squares regression, as described by Stedinger and Tasker (1985), is a regression technique that takes into account the correlation between, as well as differences in the variability and reliability of, the flow estimates used as dependent, or response, variables. These factors are accounted for in generalized least-squares regression by assigning different weights to each observation of the response variable used in the regression, based on its contribution to the total variance of the sample-flow statistic used as the response variable. In contrast, ordinary least-squares regression assumes equal reliability and variability in flow estimates at all sites and no cross-correlation between flow records at all sites, so that each flow estimate has equal variance and is assigned equal weight in the regression.

The use of generalized least-squares regression techniques to model the relations between peak discharges and basin characteristics of North Carolina rural basins requires estimates of the cross-correlation coefficients and standard deviation of the peak-flow records that were used to compute peak discharges for the selected recurrence intervals. For each of the three hydrologic areas, a scatter-plot of sample correlation coefficients versus distance between sites was constructed for site pairs with long periods (at least

30 years) of concurrent record. A graphical 'best-fit' line to these points was used to define the relation between cross-correlation coefficient and distance between sites. This relation was then used to populate a cross-correlation matrix for the sites contained in each area. Variability of each peak-flow estimate is measured by the standard deviation of the peak-flow record used to compute that estimate. For each hydrologic area, a generalized least-squares regression of the sample standard deviations against drainage area was used to obtain estimates of the standard deviations of the peak-flow records at each site. These regression estimates of the standard deviations were used to assign weights to flow estimates because they are independent of the sample standard deviation estimates used to compute the flow estimate. Finally, length of record at each peak-flow site was used as a direct measure of the relative reliability of the flow estimates computed from those records.

Generalized least-squares regression was used to evaluate the 1- and 2-variable models suggested by preliminary ordinary least-squares regression for each of the three hydrologic areas in North Carolina. The final regression models in all of the regions relate peak discharge to drainage area for each recurrence interval (table 5). The 2-variable model for each region was tested by using generalized least-squares regression, and in each case, the addition of a second variable did not substantially improve the predictive ability of the model.

Table 5. North Carolina rural flood-frequency equations

 [DA, drainage area, in square miles. Result will be in cubic feet per second]

Rural		Hydrologic area	
flood recur- rence interval (years)	Blue Ridge- Piedmont	Coastal Plain	Sand Hills
2	135 DA ^{0.702}	64.7 DA ^{0.673}	33.5 DA ^{0.712}
5	242 DA ^{0.677}	129 DA ^{0.635}	55.5 DA ^{0.701}
10	334 DA ^{0.662}	188 DA ^{0.615}	72.9 DA ^{0.697}
25	476 DA ^{0.645}	281 DA ^{0.593}	98.1 DA ^{0.693}
50	602 DA ^{0.635}	367 DA ^{0.579}	120 DA ^{0.691}
100	745 DA ^{0.625}	468 DA ^{0.566}	143 DA ^{0.688}
200	908 DA ^{0.616}	586 DA ^{0.554}	170 DA ^{0.686}
500	1,160 DA ^{0.605}	773 DA ^{0.539}	210 DA ^{0.684}

Uncertainty in a flow estimate that was predicted for an ungaged site by using the regression equations can be measured by the standard error of prediction, S_p , which is computed as the square root of the mean square error of prediction, MSEp. The MSEp is the sum of two components—the mean square error resulting from the model, γ^2 , and the sampling mean square error, $MSE_{s,i}$, which results from estimating model parameters from samples of the population. The mean square model error, γ^2 , is a characteristic of the model and is a constant for all sites. The mean square sample error, $MSE_{s,i}$, for a given site, however, depends on the values of the explanatory variables (DA) used to develop the flow estimate at that site. The standard error of prediction for a site, *i*, is computed as:

$$S_{p,i} = (\gamma^2 + MSE_{s,i})^{\frac{1}{2}}$$
, (2)

and, therefore, varies from site to site. If the values of the explanatory variables for the gage sites used in the regression are assumed to be a representative sample of all sites in the region, then the average accuracy of prediction for the regression model can be determined by computing the average standard error of prediction:

$$S_{p} = \left\{ \gamma^{2} + \frac{1}{n} \sum_{i=1}^{n} MSE_{s,i} \right\}^{\frac{1}{2}} .$$
 (3)

The standard error of the model $(SE_{(model)})$ can be converted from log (base 10) units to percent error by using the transformation formula,

$$\% SE_{(model)} = 100(10^{2.3026(\gamma^2)} - 1)^{\frac{1}{2}}.$$
 (4)

Similarly, the average standard error of prediction can be transformed from log (base 10) units to percent error by substituting S_p^2 for γ^2 in equation 4. Computation of $S_{p,i}$ for a given ungaged site, *i*, involves fairly complex matrix algebra. Computational procedures and the required matrices are provided in the Appendix.

The standard errors of the model, which measure how well the regression model fits the data used to construct it, ranged from about 34 percent to just over 57 percent. This error term is comparable to errors often cited and referred to as 'model error' or 'standard error of estimate' in earlier studies in which ordinary least-squares regression was used to develop predictive equations. The average standard errors of prediction, which provide a better overall measure of a model's predictive ability, ranged from about 36 percent to about 65 percent (table 6). Another measure of predictive ability is equivalent years of record (Hardison, 1971). Equivalent years of record are the number of years of peak-flow record needed to provide an estimate by using log-Pearson Type III techniques that would be equal in accuracy to an estimate made by using regional methods (table 6).

 Table 6.
 Average predictive errors, in percent, and equivalent years of record associated with North Carolina rural flood-frequency equations

Rural			Hydrolog	jic area			
flood	Blue Ridge-Piedmont		Coastal Plain		Sand Hills		
recurrence interval (years)	Average error of prediction	Equivalent years of record	Average error of prediction	Equivalent years of record	Average error of prediction	Equivalent years of record	
2	41.2	2.0	37.9	2.9	38.4	2.1	
5	41.2	3.0	35.9	4.9	42.6	2.7	
10	42.0	4.1	36.3	6.7	45.6	3.4	
25	43.6	5.4	38.0	8.8	49.8	4.2	
50	45.9	6.4	39.8	10.1	53.1	4.6	
100	47.0	7.2	42.0	11.1	56.6	5.0	
200	48.9	7.9	44.2	11.9	60.2	5.4	
500	51.6	8.7	47.3	12.7	65.1	5.7	

Region-of-Influence Analysis

The region-of-influence method (Tasker and Slade, 1994) estimates flood discharges at ungaged basins by deriving, for a given ungaged rural site, regression relations between the flood discharges and basin characteristics of a unique subset of gaged sites. This unique subset of gaged sites for a given ungaged site, first suggested by Acreman and Wiltshire (1987), was described by Burn (1990a, b) as the region of influence for an ungaged site, hence the name of the method. The unique subset of gaged sites is defined as the N 'nearest' gages to the ungaged site, where distance between sites *i* and *j* is defined by the Euclidean distance metric:

$$d_{ij} = \left(\sum_{k=1}^{p} \left(\frac{x_{ik} - x_{jk}}{sd(X_k)}\right)^2\right)^{\frac{1}{2}},$$
 (5)

where

- d_{ij} is the distance between sites *i* and *j* in terms of basin characteristics,
- p is the number of basin characteristics used to calculate d_{ij} ,
- X_k is the *k*th basin characteristic,
- $sd(X_k)$ is the sample standard deviation for X_k , and x_{ik} is the value of X_k at the *i*th site.

This distance metric is directly analogous to the more familiar equation for distance, D, between two points, (x_1, y_1) and (x_2, y_2) in a 2-dimensional rectangular coordinate system:

$$D = \left[\left(x_2 - x_1 \right)^2 + \left(y_2 - y_1 \right)^2 \right]^{\frac{1}{2}}, \tag{6}$$

where the only difference is the use of sample standard deviation to standardize the different basin characteristics and the slight notational difference of using an additional subscript k rather than changing variable symbols (x, y).

The distances, d_{ij} 's, between a given ungaged site and all the gaged sites are computed and ranked; the N gaging stations with the smallest d_{ij} compose the region of influence for that gaging station. Once determined, generalized least-squares regression techniques are used to develop the unique predictive relations between flood discharge and basin characteristics and estimates of the selected recurrence interval discharge at the ungaged site computed.

The number, p, and identity of the basin characteristics that are used to compute d_{ii} and the number of gaged sites, N, that compose the region of influence are specific to a given set of flood-discharge estimates and basin characteristics. In order to adapt the region-of-influence method to that data set, these parameters must be determined. In addition to these parameters, the set of basin characteristics also must be chosen for use as explanatory variables in the generalized least-squares regression models developed for each region. There is a subtle but important distinction between the two sets of basin characteristics-the first is used to define a region of influence; the second serves as variables in the unique predictive equations that are developed for that region of influence. These two sets of characteristics need not be identical but are in some cases. In other cases, such as in North Carolina, the set of characteristics used as variables is a subset of the set of characteristics used to define the region of influence.

Selection of the number of gaged sites, N, and the number and identity of the basin characteristics that will define the region of influence for North Carolina was done by trial and error, using a computed root mean square error (RMSE) as the criterion. RMSE was computed by removing one site at a time from the data base and using the remaining sites to compute an estimate of the flow characteristic. Once completed for every site, the RMSE was computed as the square root of the arithmetic mean of the differences between the estimated and computed values at each site. The results of the exploratory multiple regression analyses performed as part of the traditional regional regression analysis were used to provide some insight in selecting initial sets of basin characteristics. The strong evidence for using separate hydrologic areas in the traditional regression analysis led to the decision to restrict a site's region of influence to its hydrologic area. As a result, 15 sites in the Sand Hills region (fig. 1) were not enough to support a valid region-of-influence analysis. For any ungaged site identified as a Sand Hills site, the same set of 15 sites would compose the region of influence, and the unique predictive equation developed would be the same equation developed by using traditional regional regression techniques, as described in previous sections of this report.

Combinations of defining variables that were tested include DA and CF_{25} ; DA and REG; DA, CF_{25} ,

and REG; and DA, CF_{25} , L, and REG. Each set of defining variables was tested by using values of 25, 30, and 35 for N. For all variable combinations, N = 30 provided the best results; and the combination of variables that minimized RMSE for all recurrence intervals was DA, CF_{25} , and REG. For these initial tests, DA and CF_{25} were used as explanatory variables in the unique regression relations. Subsequent testing, after the defining variables and N were determined, indicated that CF_{25} was not significant as an explanatory variable. As a result, only DA is used as an explanatory variable in the final version of the region-of-influence method.

After determining the best combination of variables to define the region of influence and the optimal value for N, the computer application for the region of influence was completed. Equation 5 is used to determine the region of influence for an ungaged site, given the required input variables. Unique predictive equations for the ungaged site are then developed, using a generalized least-squares regression of the sites within the region of influence, and the predicted flood-discharge estimates are computed. In addition, because generalized least-squares regression was used to develop the predictive equations, $S_{p,i}$, the site-specific standard error of prediction is computed for each estimated recurrence interval discharge.

Comparison of Results

Application of the regional regression equations requires one less variable than application of the

region-of-influence method. However, the additional variable, latitude and longitude of the ungaged site, is simple to determine, so that the variable requirements of the methods are nearly equal. The regional regression equations are easily evaluated manually, the region-of-influence method, however, is computationally intensive but is made simpler by the use of a computer application that performs the complex computations.

The average RMSE was computed for each area and recurrence interval (table 7), providing a measure of the predictive ability of the model or method. Average RMSE was computed as the square root of the arithmetic mean of the differences between the floodfrequency estimate determined using the log-Pearson Type III and the flood-frequency estimate computed using either the regression equations or the region-ofinfluence method. RMSE for the region-of-influence method is slightly less than for the traditional regression equations in all cases. A site-specific comparison of predictive error also is possible by using $S_{n,i}$. As discussed previously, the region-of-influence method reports the site-specific standard error of prediction, $S_{p,i}$. The $S_{p,i}$ is not typically computed when evaluating the traditional regression equations manually because of the complexity of the computations involved. Automation of the equations eliminates this concern, and the $S_{p,i}$ is reported along with the flood-discharge estimate for any given site, allowing for comparison of predictive results on a siteby-site basis.

Table 7. Root mean square error, in percent, for the regional regression and region-of-influence methods, presented by hydrologic area and recurrence interval

 [n.a., not applicable]

			Hydrolo	gic area		
Recurrence	Blue Ridge-Piedmont		Coasta	Coastal Plain		Hills
interval	Regional regression	Region of influence	Regional regression	Region of influence	Regional regression	Region of influence
2	43.9	42.9	39.3	34.4	40.9	n.a.
5	45.4	43.3	38.6	34.6	46.1	n.a.
10	47.4	44.7	40.5	37.1	50.3	n.a.
25	50.7	47.3	44.4	41.7	55.9	n.a.
50	53.4	49.5	47.9	45.6	60.3	n.a.
100	56.2	51.9	51.6	49.7	64.7	n.a.
200	59.2	54.4	55.7	53.9	69.3	n.a.
500	63.1	57.9	61.1	59.6	75.4	n.a.

In general, little difference was found in the ease of application or in average predictive abilities between the regional regression equations and the region-ofinfluence method. The region-of-influence method is a new technique and is still being improved. As a result, the region-of-influence method is considered a secondary or alternative method of determining floodfrequency estimates for ungaged rural sites in North Carolina.

Use of Computer Software

As part of the study described by this report, a computer software package was developed that computes (1) estimates of flood-frequency discharges using the region-of-influence method at ungaged rural sites in the Blue Ridge-Piedmont or Coastal Plain hydrologic areas of North Carolina, (2) estimates of flood-frequency discharges using the regional regression equations for ungaged rural sites in each of the three hydrologic areas of North Carolina, and (3) the associated site-specific errors of prediction, $S_{n,i}$, for each method. The complexity of the computations required for the region-of-influence method requires the use of the software for practical application of the method. The regional regression equations can be evaluated manually, but the software allows for easy evaluation of the complex computation of the S_{ni} for the regional regression method.

The computer software package includes an executable program file and four supporting data files. All five files are required for execution of the computer software. The software package and instructions for down loading, installation, and execution of the program currently are available at the North Carolina District home page on the World Wide Web at URL http://nc.water.usgs.gov/reports/wri014207>.

APPLICATION OF METHODS

The methods presented in this report can be used to estimate the 2-, 5-, 10-, 25-, 50-, 100-, 200-, and 500year recurrence interval flood discharges at gaged and ungaged, unregulated, rural sites in North Carolina. Use of either the regional regression equations or the region-of-influence method requires estimates of the input variables. To apply these methods, first locate the ungaged site on a map and identify in which hydrologic area the site is located. An estimate of the latitude and longitude of the site is required for the region-ofinfluence method. Next, delineate the drainage boundaries of the ungaged site and measure the drainage area contained within those boundaries. The corresponding regression equations (table 5) can then be applied to determine an estimate of the flood discharges for the recurrence interval of interest. Alternatively, the region-of-influence computer application can be initiated; it will query the user for an output file name, an identifier for the site of interest, the hydrologic area for the site, the drainage area of the site, and the latitude and longitude of the site. With this information, the computer application computes the climate factor, defines a region of influence, and produces the desired flood-discharge estimates, along with the standard error of prediction, $S_{p,i}$, specific to the ungaged site.

The computer application contains the regression equations and can be used to apply either method. Use of the computer application to evaluate the regression equation provides an automated computation of $S_{p,i}$ for the regression equations as well as for the region-of-influence method. If evaluated manually, $S_{p,i}$ can be computed only by using the rather complex computational procedures described previously and outlined in detail in the Appendix. Although average standard errors of prediction (table 6) give an idea of the relative accuracy of the methods; $S_{p,i}$ is the more precise measure of the accuracy of a specific prediction.

Flood-frequency estimates at gaged sites and ungaged sites on the same stream as a gaged site can be improved by combining the estimate determined by regional methods with the estimate determined by fitting the log-Pearson Type III distribution to the peakflow record at the gaged site. At a gaged site, the best estimate of flood frequency can be determined by

$$Q_t(w) = \frac{Q_t(g)N + Q_t(r)EY}{N + EY},$$
(7)

where

- $Q_t(w)$ is the weighted discharge for recurrence interval *t*;
- $Q_t(g)$ is the discharge for recurrence interval *t* determined using peak-flow record from the gaged site;
- $Q_t(r)$ is the discharge for recurrence interval t determined using regional methods;
 - *N* is the number of systematic peaks in the gaged sites record; and
 - *EY* is the equivalent years of record from table 6.

Flood estimates at an ungaged site that is on the same stream as a gaged site can be determined by using a combination of the regional estimate and the log-Pearson Type III estimate from the nearby gaged site. In order to make the appropriate adjustment, first compute the ratio,

$$R = \frac{Q_t(w)}{Q_t(r)},\tag{8}$$

for the gaged site by using $Q_t(w)$ and $Q_t(r)$ as defined in the preceding paragraph. Next, a correction factor, *R*', is computed as follows:

$$R' = R - \frac{\Delta DA(R-1)}{0.5DA_g}, \qquad (9)$$

where ΔDA is the absolute value of the difference between the drainage areas of the gaged and ungaged sites, and DA_g is the drainage area of the gaged site. If $\Delta DA/DA_g$ is less than 0.5, then the corrected discharge for the ungaged site, $Q_t(\text{corr})$, can be computed by multiplying the correction factor, *R*', by the regional estimate for the ungaged site, $Q_t(r)$. If $\Delta DA/DA_g$ is greater than 0.5, use the results of the regional methods without correction.

At times, flood-frequency estimates may be desired for an ungaged site that is between two gaged sites on the same stream. In this case, select the gaged site for which $\Delta DA/DA_g$ is less than 0.5, compute *R*', and apply as described above. If $\Delta DA/DA_g$ is less than 0.5 for both gaged sites, compute R' for each. If both correction factors are greater than 1.0, use the larger *R*'; if both correction factors are less than 1.0, use the smaller *R*'. If one correction factor is greater than 1.0 and the other smaller than 1.0, an average of the two correction factors should be used.

If the drainage basin for an ungaged site lies within more than one hydrologic area, the computed discharge should be adjusted according to the proportion of the total drainage area that lies within each hydrologic area. The adjusted discharge can be determined by the equation:

$$Q_{t}(\text{adjusted}) = Q_{t}(HA1)x\frac{DA_{1}}{DA_{\text{total}}} + Q_{t}(HA2)x\frac{DA_{2}}{DA_{\text{total}}},$$
 (10)

where Q_t (adjusted) is the adjusted discharge for the *t*-year recurrence interval; Q_t (HA1) and Q_t (HA2) are the discharges computed as if the entire drainage area were within the hydrologic areas, HA1 and HA2; DA₁ and DA₂ are portions of the total drainage area found in the respective hydrologic drainage areas; and DA_{total} is the total drainage area.

SUMMARY

Accurate and reliable estimates of the magnitude and frequency of floods are critical for such activities as bridge design, flood-plain delineation and management, water-supply management, and management of water-control structures, among others. Recognizing the need for accurate estimates of flood frequency at ungaged rural basins, the U.S. Geological Survey, in cooperation with the North Carolina Department of Transportation, conducted a study to further define the relation between flood discharges of selected recurrence intervals and selected physical and climatic characteristics of rural North Carolina basins. This study includes the development of two methods for regionalizing, or extending in space, flood-frequency estimates at gaged sites. In the first method, traditional regional regression analysis, a generalized leastsquares regression analysis is used to develop a set of predictive equations for each of three hydrologic areas in North Carolina-the Blue Ridge-Piedmont, the Coastal Plain, and the Sand Hills. In the second method, the region-of-influence method, floodfrequency estimates for ungaged sites are predicted interactively, based on data from a subset of gaged sites with basin characteristics similar to those of the ungaged site. This report documents the development of both methods, using a data base of flood-discharge estimates and basin characteristics for 317 rural North Carolina gaged sites.

An initial set of 366 gaged sites was determined to have some annual peak-flow record; basin characteristics data were computed and compiled for all of these sites by using a GIS. While the development of the basin characteristics was ongoing, flow records were examined to determine which sites had flows that were affected by regulation or channelization. Of the 366 original sites, 19 sites had only regulated record and 27 sites had periods of unregulated flow record prior to regulation. After basin characteristics were developed and flow records were examined, preliminary computations of flood-frequency estimates were begun. Results of these preliminary computations indicated the need for a generalized skew study for North Carolina basins to replace outdated generalized skews that were based on a nationwide study. After the generalized skew study, flood-frequency estimates for all sites with 10 or more years of record were computed. Flood-frequency estimates were computed for 317 rural, unregulated sites and for 42 rural, regulated sites. The sites with regulated record were excluded from further analysis.

Basin characteristics data and flood-frequency estimates for the 317 rural, unregulated sites were merged to form the data base that was used to develop the regional regression equations and the region-ofinfluence method. Of the 317 total sites, 222 were located in the Blue Ridge-Piedmont hydrologic area, 80 were located in the Coastal Plain hydrologic area, and 15 were located in the Sand Hills hydrologic area. Preliminary multiple regression analyses, using ordinary least-squares regression, were conducted to confirm the validity of the regionalization scheme and to identify the best combination of explanatory variables for inclusion in the generalized least-squares analysis.

Generalized least-squares analysis was used to develop a set of equations for each region that relates the 2-, 5-, 10-, 25-, 50-, 100-, 200-, and 500-year recurrence interval flood discharges to drainage area. Model error and error of prediction for the equations ranged from about 40 percent for the lower recurrence interval equations to more than 50 percent, with two equations for the Sand Hills indicating more than 60 percent.

The region-of-influence method was adapted to the available flood-frequency and basin characteristics data for North Carolina. The drainage area, hydrologic area, and latitude and longitude of an ungaged site in either the Blue Ridge-Piedmont or Coastal Plain hydrologic areas of North Carolina are required to predict the 2-, 5-, 10-, 25-, 50-, 100-, 200-, and 500year recurrence interval flood discharges for a specified ungaged site. The Sand Hills hydrologic area did not have a sufficient number of sites to apply the region-ofinfluence method. Because of the complexity of the computations involved in the region-of-influence method, a computer application is required for the practical use of the method.

A brief comparison of the regional regression and region-of-influence methods, based on ease of

application and RMSE of prediction, resulted in neither method being clearly superior. Both require hydrologic area and drainage area as input variables; the region-ofinfluence method additionally requires latitude and longitude, but these coordinates are fairly simple to determine. The RMSE were, in general, lower for the region-of-influence method, but only slightly. The region-of-influence method is newly developed and still being refined. As a result, the regional regression equations are considered to be the primary method of estimating magnitude and frequency of floods for rural ungaged sites in North Carolina. The region-ofinfluence method can be considered an alternative method.

A computer application is available that automates the complex computations required by the region-of-influence method. This computer application includes the option to compute flood-frequency estimates using the predictive equations developed by the traditional regional regression analysis. The computer application also computes site-specific error of prediction for each method.

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Map dentification number (fig. 1)	Station number	Station name	Latitude	Longitude	Period of analysis	Number o systematic peaks
1	02053110	Wildcat Swamp near Jackson	36°25'48"	77°22'24"	1953–1971	19
2	02053170	Cutawhiskie Creek near Woodland	36°17'54"	77°11'58"	1953-1971	19
3	02053200	Potecasi Creek near Union	36°22'14"	77°01'36"	1929-1996	39
4 ^{nc}	02053400	Ahoskie Creek near Rich Square	36°14'52"	77°14'12"	1965-1973	9
5 ^{nc}	02053450	Ahoskie Creek at Mintons Store	36°16'46"	77°09'28"	1965–1973	9
6	02053500	Ahoskie Creek at Ahoskie	36°16'48"	77°00'00''	1940–1963	13
6^{*r}	02053500*	Ahoskie Creek at Ahoskie (channelized period)	36°16'48"	77°00'00"	1964–1996	33
7	02053510	Ahoskie Creek tributary at Poortown	36°16'29"	77°00'38"	1964–1973	10
8	02053550	Chinkapin Creek near Colerain	36°11'52"	76°47'14"	1953-1971	19
9	02068500	Dan River near Francisco	36°30'53"	80°18'11"	1916–1938	13
9* ^r	02068500*	Dan River near Francisco (regulated period)	36°30'53"	80°18'11"	1939–1996	54
10	02068610	Hog Rock Creek near Moores Springs	36°23'53"	80°19'46"	1955-1971	15
11	02068660	Little Snow Creek near Lawsonville	36°27'54"	80°10'28"	1954-1971	18
12	02069030	Belews Creek near Kernersville	36°12'20"	80°04'25"	1954–1971	17
13	02070500	Mayo River near Price	36°32'05"	79°59'30"	1930–1996	45
14	02070810	Jacobs Creek near Wentworth	36°20'54"	79°53'14"	1954–1973	18
15	02071000	Dan River near Wentworth	36°24'45"	79°49'35"	1908–1996	57
16	02071410	Matrimony Creek near Leaksville	36°31'39"	79°50'08"	1958–1973	15
17 ^{nc}	02071500	Dan River at Leaksville	36°29'00"	79°46'00"	1930–1949	9
18	02074000	Smith River at Eden	36°31'31"	79°45'57"	1940–1949	10
$18*^{r}$	02074000*	Smith River at Eden (regulated period)	36°31'31"	79°45'57"	1950–1996	47
19	02075160	Moon Creek near Yanceyville	36°28'13"	79°23'00"	1954–1989	21
20	02075230	South Country Line Creek near Hightowers	36°19'29"	79°18'20"	1954–1976	23
21	02077200	Hyco Creek near Leasburg	36°23'57"	79°11'50"	1965–1996	30
22	02077210	Kilgore Creek tributary near Leasburg	36°22'38"	79°09'57"	1954–1971	13
23	02077240	Double Creek near Roseville	36°21'44"	79°05'48"	1965-1982	16
24	02077250	South Hyco Creek near Roseville	36°23'09"	79°06'26"	1967–1980	14
$25^{r, nc}$	02077300	Hyco River at McGehees Mill	36°31'02"	79°01'42"	1965–1973	9
26 ^r	02077303	Hyco River below Afterbay Dam near McGehees Mill	36°31'24"	78°59'48"	1974–1996	23
27	02077310	Storys Creek near Roxboro	36°23'48"	79°01'14"	1954–1971	18
28 ^r	02077670	Mayo Creek near Bethel Hill	36°32'26"	78°52'21"	1978–1996	19
29	02080500	Roanoke River at Roanoke Rapids	36°27'37"	77°38'04"	1878–1949	38
29* ^r	02080500*	Roanoke River at Roanoke Rapids (regulated period)	36°27'37"	77°38'04"	1956–1996	41
30	02081000	Roanoke River near Scotland Neck	36°12'34"	77°23'03"	1940–1949	10
31	02081060	Smithwick Creek tributary near Williamston	35°43'51"	77°04'42"	1953–1971	19

[nc, flood-frequency estimates were not computed because the site has less than 10 years of peak-flow record; *, duplicate map identification and station number for sites having separate period of regulated or channelized flows; r, site excluded from regional analysis because flows were affected by regulation or channelization]

Map identification number (fig. 1)	Station number	Station name	Latitude	Longitude	Period of analysis	Number o systematic peaks
32	02081110	White Oak Swamp near Windsor	36°04'46"	76°58'36"	1953–1971	14
33 ^{nc}	0208111310	Cashie River at Secondary Road 1257 near Windsor	36°02'51"	76°59'07"	1988–1996	9
34	02081500	Tar River near Tar River	36°11'41"	78°35'00"	1940–1996	57
35	02081710	Long Creek at Kittrell	36°13'30"	78°27'15"	1954–1976	20
36	02081747	Tar River at U.S. 401 at Louisburg	36°05'34"	78°17'48"	1964–1996	33
37	02081800	Cedar Creek near Louisburg	36°03'14"	78°20'24"	1935-1975	22
38 ^{nc}	02081935	Tar River at Spring Hope	35°55'42"	78°08'53"	1967-1971	5
39	02082000	Tar River near Nashville	35°50'57"	77°55'51"	1919–1970	42
40	02082500	Sapony Creek near Nashville	35°53'10"	77°54'40"	1951-1970	20
41 ^r	02082506	Tar River below Tar River Reservoir near Rocky Mount	35°53'58"	77°51'57"	1973–1996	24
42	02082540	Wildcat Branch near Mapleville	36°03'29"	78°08'39"	1953–1976	11
43 ^r	02082585	Tar River at NC97 at Rocky Mount	35°57'15"	77°47'15"	1977-1996	20
44	02082610	Tar River near Rocky Mount	35°58'38"	77°45'35"	1964-1973	10
45	02082630	Harts Mill Run near Tarboro	35°55'40"	77°37'10"	1953-1971	18
46 ^{nc}	02082731	Devils Cradle Creek near Alert at Secondary Road 1412	36°12'03"	78°14'19"	1993–1996	4
47	02082770	Swift Creek at Hilliardston	36°06'42"	77°55'16"	1924–1996	33
48	02082835	Fishing Creek near Warrenton	36°23'00"	78°10'54"	1954-1976	22
49	02082950	Little Fishing Creek near White Oak	36°11'08"	77°52'34"	1960-1996	37
50 ^{nc}	02082955	Fishing Creek near Glenview	36°08'44"	77°50'31"	1967-1971	5
51	02083000	Fishing Creek near Enfield	36°09'03"	77°41'35"	1915-1996	82
52	02083090	Beaverdam Swamp near Heathsville	36°16'49"	77°41'48"	1953–1971	19
53	02083410	Deep Creek near Scotland Neck	36°09'26"	77°28'24"	1953-1973	21
54	02083500	Tar River at Tarboro	35°53'38"	77°32'00"	1897–1996	95
55	02083800	Conetoe Creek near Bethel	35°46'33"	77°27'45"	1957–1996	40
56 ^{nc}	02083833	Conetoe Creek (tributary 3) near Penny Hill	35°46'00"	77°29'26"	1993–1996	4
57 ^{nc}	02084160	Chicod Creek at Secondary Road 1760 near Simpson	35°33'47"	77°13'43"	1976–1981	6
57* ^r	02084160*	Chicod Creek at Secondary Road 1760 near Simpson (channelized period)	35°33'47"	77°13'43"	1982-1996	11
58 ^{nc}	02084164	Juniper Branch at Secondary Road 1766 near Simpson	35°33'55"	77°14'43"	1976-1978	3
58* ^{r,nc}	02084164*	Juniper Branch at Secondary Road 1766 near Simpson (channelized period)	35°33'55"	77°14'43"	1979–1986	8
59	02084240	Collie Swamp near Everetts	35°49'34"	77°12'03"	1953–1976	24
60	02084500	Herring Run near Washington	35°34'03"	77°01'09"	1946–1980	30
61	02084520	Upper Goose Creek near Yeatsville	35°31'25"	76°53'23"	1953-1973	21
62	02084540	Durham Creek at Edward	35°19'25"	76°52'26"	1966-1992	27
63	02084570	Acre Swamp near Pinetown	35°35'02"	76°50'23"	1953-1969	17
64 ^{nc}	02084909	Sevenmile Creek near Efland	36°03'56"	79°08'39"	1988–1996	9

Map identification number (fig. 1)	Station number	Station name	Latitude	Longitude	Period of analysis	Number of systematic peaks
65	02085000	Eno River at Hillsborough	36°04'18"	79°05'49"	1928–1996	54
66	02085020	Stony Creek tributary near Hillsboro	36°03'01"	79°02'14"	1953-1971	19
67	02085070	Eno River near Durham	36°04'20"	78°54'30"	1964–1996	33
68	02085190	North Fork Little River tributary near Rougemont	36°11'41"	79°00'52"	1954–1976	23
69	0208521324	Little River at Secondary Road 1461 near Orange Factory	36°08'30"	78°55'10"	1962-1996	35
70	02085500	Flat River at Bahama	36°10'57"	78°52'44"	1926–1996	71
71	02086000	Dial Creek near Bahama	36°10'36"	78°51'24"	1926-1991	47
72 ^r	02086500	Flat River at Dam near Bahama	36°08'55"	78°49'43"	1928-1993	48
73	02086624	Knap Of Reeds Creek near Butner	36°07'40"	78°48'55"	1983-1995	13
74 ^{nc}	02086849	Ellerbe Creek near Gorman	36°03'33"	78°49'58"	1983–1994	8
75	02087000	Neuse River near Northside	36°02'54"	78°44'59"	1928-1980	53
76	0208700780	Little Lick Creek above Secondary Road 1814 near Oak Grove	35°59'11"	78°47'58"	1983-1995	13
77	02087030	Lick Creek near Durham	35°58'50"	78°44'19"	1954-1971	18
78	02087140	Lower Barton Creek tributary near Raleigh	35°54'44"	78°40'55"	1954–1971	18
79	02087183	Neuse River near Falls	35°56'25"	78°34'56"	1945–1980	21
79* ^r	02087183*	Neuse River near Falls (regulated period)	35°56'25"	78°34'56"	1981–1996	16
80	02087240	Stirrup Iron Creek tributary near Nelson	35°53'06"	78°49'37"	1952-1973	20
81	02087500	Neuse River near Clayton	35°38'50"	78°24'22"	1919-1980	53
81* ^r	02087500*	Neuse River near Clayton (regulated period)	35°38'50"	78°24'22"	1981-1996	16
82	02087570	Neuse River at Smithfield	35°30'46"	78°21'00"	1908–1980	48
82* ^r	02087570*	Neuse River at Smithfield (regulated period)	35°30'46"	78°21'00"	1981–1990	10
83	02087580	Swift Creek near Apex	35°43'00"	78°45'00"	1954-1971	18
84 ^{nc}	0208758850	Swift Creek near McCullars Crossroads	35°41'33"	78°41'34"	1992-1996	5
85	02087910	Middle Creek near Holly Springs	35°39'28"	78°48'06"	1954-1971	18
86	02088000	Middle Creek near Clayton	35°34'10"	78°35'30"	1940–1996	56
87	02088140	Stone Creek near Newton Grove	35°20'24"	78°21'54"	1953–1971	19
88	02088210	Hannah Creek near Benson	35°23'36"	78°31'48"	1953-1971	19
89	02088420	Long Branch near Selma	35°38'11"	78°15'06"	1953-1971	19
90	02088470	Little River near Kenly	35°35'20"	78°11'18"	1965-1989	25
91	02088500	Little River near Princeton	35°30'40"	78°09'38"	1919–1996	66
92	02089000	Neuse River near Goldsboro	35°20'14"	77°59'51"	1930–1980	51
92* ^r	02089000*	Neuse River near Goldsboro (regulated period)	35°20'14"	77°59'51"	1984-1996	13
93 ^{nc}	0208925200	Bear Creek at Mays Store	35°16'28"	77°47'40"	1988–1996	9
94	02089500	Neuse River at Kinston	35°15'29"	77°35'09"	1919–1980	53
94* ^r	02089500*	Neuse River at Kinston (regulated period)	35°15'29"	77°35'09"	1981-1996	16

[nc, flood-frequency estimates were not computed because the site has less than 10 years of peak-flow record; *, duplicate map identification and station number for sites having separate period of regulated or channelized flows; r, site excluded from regional analysis because flows were affected by regulation or channelization]

Map identification number (fig. 1)	Station number	Station name	Latitude	Longitude	Period of analysis	Number of systematic peaks
95	02090380	Contentnea Creek near Lucama	35°41'29"	78°06'38"	1965–1976	12
95* ^r	02090380*	Contentnea Creek near Lucama (regulated period)	35°41'29"	78°06'38"	1977–1996	20
96	02090560	Lee Swamp tributary near Lucama	35°38'21"	78°01'37"	1953–1971	19
97	02090625	Turner Swamp near Eureka	35°34'14"	77°52'47"	1969–1987	19
98	02090780	Whiteoak Swamp tributary near Wilson	35°42'24"	77°47'11"	1953–1971	19
99	02090960	Nahunta Swamp near Pikeville	35°30'40"	77°58'56"	1953–1973	19
100 ^{nc}	0209096970	Moccasin Run near Patetown	35°28'46"	77°54'37"	1989–1996	8
101	02091000	Nahunta Swamp near Shine	35°29'20"	77°48'22"	1955–1996	42
102	02091430	Shepherd Run near Snow Hill	35°26'06"	77°38'42"	1953–1971	19
103	02091500	Contentnea Creek at Hookerton	35°25'44"	77°34'59"	1928–1996	68
104	02091700	Little Contentnea Creek near Farmville	35°32'40"	77°30'41"	1957–1987	31
105	02091810	Halfmoon Creek near Fort Barnwell	35°17'58"	77°21'14"	1953–1975	12
106	02091970	Creeping Swamp near Vanceboro	35°23'30"	77°13'46"	1972–1985	14
107	02092000	Swift Creek near Vanceboro	35°20'42"	77°11'45"	1909–1989	39
108	02092020	Palmetto Swamp near Vanceboro	35°20'18"	77°10'16"	1953–1976	24
109 110 111 112 113	02092120 02092290 02092500 02092520 02092520 02092620	Bachelor Creek near New Bern Rattlesnake Branch near Comfort Trent River near Trenton Vine Swamp near Kinston Upper Broad Creek tributary near Grantsboro	35°10'24" 35°00'31" 35°03'54" 35°09'29" 35°08'06"	77°06'14" 77°35'50" 77°27'24" 77°33'16" 76°56'31"	1953–1971 1953–1971 1928–1996 1953–1971 1953–1973	19 19 45 19 21
114	02092720	White Oak River at Belgrade	34°53'30"	77°14'02"	1953–1973	21
115	02092780	Bell Swamp near Hubert	34°42'04"	77°14'01"	1953–1970	18
116	02093000	New River near Gum Branch	34°50'56"	77°31'11"	1908–1996	33
117	02093040	Southwest Creek tributary near Jacksonville	34°47'18"	77°33'08"	1954–1973	19
118	02093070	Southwest Creek near Jacksonville	34°43'56"	77°32'02"	1953–1973	20
119	02093290	Haw River near Summerfield	36°14'32"	79°52'20"	1954–1971	18
120	02093500	Haw River near Benaja	36°15'06"	79°33'55"	1916–1971	43
121 ^{nc}	02093549	Haw River at Altamahaw	36°10'43"	79°30'09"	1968–1973	6
122	02093800	Reedy Fork near Oak Ridge	36°10'22"	79°57'12"	1956–1996	41
123	02094000	Horsepen Creek at Battle Ground	36°08'34"	79°51'40"	1926–1959	30
124	02095000	South Buffalo Creek near Greensboro	36°03'36"	79°43'33"	1929–1958	29
125	02095500	North Buffalo Creek near Greensboro	36°07'13"	79°42'30"	1929–1990	62
126	02096500	Haw River at Haw River	36°05'13"	79°22'02"	1929–1996	68
127	02096660	Rock Creek near Whitsett	36°04'49"	78°47'45"	1954–1971	17
128	02096700	Big Alamance Creek near Elon College	36°02'21"	79°31'29"	1945–1980	23

Map identification number (fig. 1)	Station number	Station name	Latitude	Longitude	Period of analysis	Number of systematic peaks
129	02096740	Gun Branch near Alamance	36°02'58"	79°28'35"	1954–1973	19
130 ^{nc}	02096846	Cane Creek near Orange Grove	35°59'13"	79°12'23"	1989–1996	8
131	02096850	Cane Creek near Teer	35°56'34"	79°14'46"	1960-1973	14
132	02096960	Haw River near Bynum	35°45'48"	79°08'02"	1908-1996	69
133	02097010	Robeson Creek near Pittsboro	35°43'29"	79°12'33"	1954–1976	23
134	02097314	New Hope Creek near Blands	35°53'05"	78°57'58"	1983–1996	14
135	0209741955	Northeast Creek at Secondary Road 1100 near Genlee	35°52'20"	78°54'49"	1983-1996	12
136	02097910	White Oak Creek near Wilsonville	35°44'47"	79°00'44"	1954-1971	18
137	02098000	New Hope River near Pittsboro	35°44'12"	79°01'36"	1908-1973	24
138 ^r	02098198	Haw River below B. Everett Jordan Dam near Moncure	35°39'11"	79°04'03"	1980–1992	13
139 ^{nc}	02098200	Haw River near Haywood	35°39'01"	79°03'59"	1966–1972	7
140	02098500	West Fork Deep River near High Point	36°00'15"	79°58'42"	1924-1966	42
141	02099000	East Fork Deep River near High Point	36°02'15"	79°56'46"	1929–1994	66
142	02099500	Deep River near Randleman	35°54'06"	79°51'05"	1929-1996	66
143 ^{nc}	02100000	Muddy Creek near Archdale	35°52'35"	79°52'43"	1935–1941	7
144	02100500	Deep River at Ramseur	35°43'34"	79°39'20"	1901–1996	73
145	02101000	Bear Creek at Robbins	35°26'03"	79°35'39"	1940-1971	32
146	02101030	Falls Creek near Bennett	35°33'20"	79°29'56"	1954-1973	20
147	02101480	Sugar Creek near Tramway	35°25'28"	79°14'50"	1954-1973	20
148 ^{nc}	0210166029	Rocky River near Crutchfield Crossroads	35°48'25"	79°31'41"	1988–1996	9
149	02101800	Tick Creek near Mount Vernon Springs	35°39'37"	79°24'08"	1959–1996	26
150	02101890	Bear Creek near Goldston	35°37'33"	79°17'54"	1952-1971	19
151	02102000	Deep River at Moncure	35°37'38"	79°06'58"	1931-1996	66
152 ^{nc}	02102192	Buckhorn Creek near Corinth	35°33'34"	78°58'25"	1973-1980	8
152* ^r	02102192*	Buckhorn Creek near Corinth (regulated period)	35°33'34"	78°58'25"	1981–1996	16
153	02102500	Cape Fear River at Lillington	35°24'22"	78°48'48"	1924–1980	57
153* ^r	02102500*	Cape Fear River at Lillington (regulated period)	35°24'22"	78°48'48"	1981-1996	16
154	02102908	Flat Creek near Inverness	35°10'54"	79°10'40"	1969–1996	28
155	02102910	Dunhams Creek tributary near Carthage	35°18'41"	79°22'53"	1954-1971	18
156	02102930	Crane Creek near Vass	35°17'53"	79°16'19"	1954–1971	18
157	02103000	Little River at Manchester	35°11'38"	78°59'14"	1939–1950	11
158	02103390	South Prong Anderson Creek near Lillington	35°15'31"	78°55'27"	1953-1971	19
159	02103500	Little River at Linden	35°15'46"	78°46'35"	1928-1971	44
160	02104000	Cape Fear River at Fayetteville	35°02'49"	78°51'36"	1889–1976	71
161	02104080	Reese Creek near Fayetteville	35°04'49"	78°47'45"	1953-1971	17

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Map identification number (fig. 1)	Station number	Station name	Latitude	Longitude	Period of analysis	Number of systematic peaks
162	02104500	Rockfish Creek near Hope Mills	34°57'57"	78°55'04"	1939–1954	16
163 r	02105500	Cape Fear River at William O. Huske Lock near Tarheel	34°50'05"	78°49'27"	1938–1980	36
163* ^r	02105500*	Cape Fear River at William O. Huske Lock near Tarheel (regulated period)	34°50'05"	78°49'27"	1981–1996	15
164 165	02105570 02105630	Browns Creek near Elizabethtown Turnbull Creek near Elizabethtown	34°36'32" 34°41'32"	78°36'57" 78°35'02"	1953–1973 1949–1971	18 19
166	02105769	Cape Fear River at Lock 1 near Kelly	34°24'15"	78°17'38"	1970–1980	11
166* ^r	02105769*	Cape Fear River at Lock 1 near Kelly (regulated period)	34°24'15"	78°17'38"	1981–1996	16
167	02105900	Hood Creek near Leland	34°16'43"	78°07'34"	1953-1996	24
168	02106000	Little Coharie Creek near Roseboro	34°57'13"	78°29'17"	1924–1991	41
169	02106240	Turkey Creek near Turkey	35°00'11"	78°11'06"	1953–1973	18
170	02106410	Stewarts Creek tributary near Warsaw	34°57'25"	78°04'42"	1955-1971	16
171	02106500	Black River near Tomahawk	34°45'17"	78°17'21"	1928–1996	45
172	02106910	Big Swamp near Roseboro	34°58'38"	78°34'07"	1953–1973	20
173	02107000	South River near Parkersburg	34°48'45"	78°27'26"	1952–1986	35
174	02107500	Colly Creek near Kelly	34°27'48"	78°15'26"	1908–1971	21
175	02107590	Northeast Cape Fear River tributary near Mount Olive	35°11'06"	77°57'34"	1954–1971	18
176	02107600	Northeast Cape Fear River near Seven Springs	35°10'20"	77°55'56"	1959–1975	17
177	02107620	Mathews Creek near Pink Hill	35°05'49"	77°49'10"	1953–1976	16
178	02107980	Limestone Creek near Beulaville	34°45'48"	77°48'15"	1953–1971	19
179	02108000	Northeast Cape Fear River near Chinquapin	34°49'40"	77°50'00"	1941–1996	56
180	02108500	Rockfish Creek near Wallace	34°44'32"	78°02'22"	1955–1981	27
181	02108548	Little Rockfish Creek at Wallace	34°44'02"	77°58'03"	1977-1992	16
182	02108610	Pike Creek near Burgaw	34°30'00"	77°53'58"	1953–1971	18
183	02108630	Turkey Creek near Castle Hayne	34°23'47"	77°54'48"	1953-1971	19
184	02108960	Buckhead Branch near Bolton	34°20'52"	78°26'19"	1953–1971	19
185	02109500	Waccamaw River at Freeland	34°05'43"	78°32'55"	1940–1996	57
186	02109640	Wet Ash Swamp near Ash	34°02'17"	78°30'14"	1953–1971	18
187	02110020	Mill Branch near Tabor City	34°10'59"	78°48'08"	1953–1971	18
188	02111000	Yadkin River at Patterson	35°59'29"	81°33'30"	1940-1996	56
189	02111180	Elk Creek at Elkville	36°04'16"	81°24'13"	1940–1996	31
190	02111340	South Prong Lewis Fork Creek near North Wilkesboro	36°11'23"	81°24'40"	1955–1971	16
191	02111500	Reddies River at North Wilkesboro	36°10'29"	81°10'09"	1940–1995	55
192	02112000	Yadkin River at Wilkesboro	36°09'09"	81°08'45"	1904–1961	48
192* ^r	02112000*	Yadkin River at Wilkesboro (regulated period)	36°09'09"	81°08'45"	1962–1996	35
193	02112120	Roaring River near Roaring River	36°14'59"	81°02'39"	1916–1996	32

Map identification number (fig. 1)	Station number	Station name	Latitude	Longitude	Period of analysis	Number of systematic peaks
194	02112247	Elkin River at Elkin	36°15'12"	80°51'45"	1971–1980	10
195 ^r	02112250	Yadkin River at Elkin	36°14'30"	80°50'49"	1965-1995	31
196	02112360	Mitchell River near State Road	36°18'42"	80°48'26"	1940–1996	32
197	02112410	Fisher River near Bottom	36°26'35"	80°46'12"	1954–1971	16
198	02112500	Fisher River near Dobson	36°23'05"	80°40'20"	1922–1933	12
199	02113000	Fisher River near Copeland	36°21'26"	80°41'10"	1922-1996	74
200 ^r	02113500	Yadkin River at Siloam	35°16'55"	80°33'46"	1977–1987	11
201	02113850	Ararat River at Ararat	36°24'16"	80°33'43"	1947–1996	32
202	02114010	Ararat River at Dam near Pilot Mountain	36°22'00"	80°33'00"	1938–1968	16
203	02114450	Little Yadkin River at Dalton	36°17'56"	80°25'53"	1961–1996	36
204 ^r	02115360	Yadkin River at Enon	36°07'55"	80°26'39"	1965-1996	32
205	02115500	Forbush Creek near Yadkinville	36°08'13"	80°33'09"	1941-1971	31
206	02115520	Logan Creek near Smithtown	36°12'50"	80°33'32"	1954–1971	18
207	02115540	South Deep Creek near Yadkinville	36°08'00"	80°46'00"	1954–1966	13
208 ^{nc}	02115730	Mill Creek near Stanleyville	36°10'49"	80°16'19"	1965–1972	6
209 ^{nc}	02115740	Mill Creek near Oldtown	36°09'06"	80°19'03"	1965-1972	6
210 ^{nc}	02115810	Little Creek near Clemmons	36°02'19"	80°20'46"	1965-1972	6
211	02115830	Smith Creek near Kernersville	36°06'19"	80°06'19"	1954–1971	18
212	02115856	Salem Creek near Atwood	36°02'10"	80°18'35"	1972–1982	11
213	02115860	Muddy Creek near Muddy Creek	36°00'01"	80°20'25"	1965–1991	19
214	02115900	South Fork Muddy Creek near Clemmons	36°00'22"	80°18'07"	1965-1991	19
215	02116500	Yadkin River at Yadkin College	35°51'23"	80°23'14"	1916–1961	33
215* ^r	02116500*	Yadkin River at Yadkin College (regulated period)	35°51'23"	80°23'14"	1962-1996	35
216	02117030	Humpy Creek near Fork	35°51'17"	80°26'24"	1969–1983	15
217	02117410	McClelland Creek near Statesville	35°57'04"	80°56'46"	1954–1976	22
218	02117500	Rocky Creek at Turnersburg	35°54'23"	80°48'34"	1941-1971	31
219	02118000	South Yadkin River near Mocksville	35°50'41"	80°39'34"	1930–1996	58
220	02118500	Hunting Creek near Harmony	36°00'00''	80°44'44"	1952–1996	45
221	02119000	South Yadkin River at Cooleemee	35°48'10"	80°33'22"	1916–1965	37
222 ^r	02119400	Third Creek near Stony Point	35°52'04"	81°04'00"	1957–1969	13
223	02120500	Third Creek at Cleveland	35°45'00"	80°41'00"	1916–1954	14
223* ^r	02120500*	Third Creek at Cleveland (regulated period)	35°45'00"	80°41'00"	1955-1971	17
224	02120780	Second Creek near Barber	35°43'05"	80°35'45"	1980–1996	17
225	02120820	Deal Branch near Salisbury	35°44'43"	80°30'25"	1954–1971	15
226	02121000	Yadkin River near Salisbury	35°43'30"	80°23'50"	1896–1927	30

[nc, flood-frequency estimates were not computed because the site has less than 10 years of peak-flow record; *, duplicate map identification and station number for sites having separate period of regulated or channelized flows; r, site excluded from regional analysis because flows were affected by regulation or channelization]

Map identification number (fig. 1)	Station number	Station name	Latitude	Longitude	Period of analysis	Number of systematic peaks
227	02121180	North Potts Creek at Linwood	35°45'28"	80°19'24"	1980–1990	11
228	02121500	Abbotts Creek at Lexington	35°48'23"	80°14'05"	1941–1995	23
229	02121940	Flat Swamp Creek near Lexington	35°43'59"	80°06'37"	1954–1971	18
230 ^{nc}	02122500	Yadkin River at High Rock	35°35'46"	80°13'59"	1916-1927	8
230* ^r	02122500*	Yadkin River at High Rock (regulated period)	35°35'46"	80°13'59"	1942–1961	19
231	02122560	Cabin Creek near Jackson Hill	35°34'57"	80°09'12"	1954–1971	17
232	02122720	Beaverdam Creek tributary near Denton	35°31'57"	80°05'04"	1954–1971	18
233	02123500	Uwharrie River near Eldorado	35°25'47"	80°01'05"	1928–1971	32
234	02123567	Dutchmans Creek near Uwharrie	35°22'05"	80°01'49"	1982–1995	12
235	02124060	North Prong Clarke Creek near Huntersville	35°25'13"	80°47'54"	1954–1973	20
236	02124130	Mallard Creek near Charlotte	35°19'05"	80°44'16"	1954–1971	18
237	02125000	Big Bear Creek near Richfield	35°20'02"	80°20'09"	1955-1996	42
238	02125410	Chinkapin Creek near Monroe	35°02'48"	80°29'33"	1953–1971	18
239	02126000	Rocky River near Norwood	35°08'54"	80°10'33"	1908–1996	67
240	02127000	Brown Creek near Polkton	35°02'10"	80°08'42"	1908–1971	36
241	02127390	Palmetto Branch at Ansonville	35°06'03"	80°07'11"	1953-1971	17
242	02128000	Little River near Star	35°23'11"	79°49'56"	1955–1996	41
243	02128260	Cheek Creek near Pekin	35°12'37"	79°50'49"	1954–1971	18
244 ^r	02129000	Pee Dee River near Rockingham	34°56'46"	79°52'11"	1928–1996	69
245	02129440	South Fork Jones Creek near Morven	34°53'51"	80°00'24"	1954–1971	18
246	02129530	Little Creek tributary near Pee Dee	34°55'07"	79°54'38"	1955-1971	11
247	02132230	Bridge Creek tributary at Johns	34°42'12"	79°26'34"	1953–1973	18
248 ^{nc}	0213228795	Jordan Creek near Silver Hill	34°58'12"	79°31'35"	1985–1993	9
249	02132320	Big Shoe Heel Creek near Laurinburg	34°45'01"	79°23'12"	1987–1996	10
250	02133500	Drowning Creek near Hoffman	35°03'38"	79°29'39"	1940–1996	57
251	02133590	Beaverdam Creek near Aberdeen	35°00'42"	79°26'50"	1953-1971	18
252	02133624	Lumber River near Maxton	34°46'22"	79°19'55"	1987–1996	10
253	02133960	Raft Swamp near Red Springs	34°52'16"	79°10'12"	1953–1971	15
254	02134380	Tenmile Swamp near Lumberton	34°43'34"	78°59'31"	1953–1973	18
255	02134500	Lumber River at Boardman	34°26'32"	78°57'38"	1901–1996	67
256	02137000	Mill Creek at Old Fort	35°37'59"	82°11'14"	1940–1975	15
257	02137727	Catawba River near Pleasant Gardens	35°41'09"	82°03'40"	1981–1996	16
258	02138000	Catawba River near Marion	35°42'26"	82°02'00"	1916–1981	40
259	02138500	Linville River near Nebo	35°47'41"	81°53'25"	1916–1996	74
260	02138680	White Branch near Marion	35°38'46"	81°55'18"	1955–1971	14

Map identification number (fig. 1)	Station number	Station name	Latitude	Longitude	Period of analysis	Number of systematic peaks
261	02140980	Carroll Creek near Collettsville	35°53'21"	81°44'18"	1955–1971	17
262	02140991	Johns River at Arneys Store	35°50'01"	81°42'43"	1986–1996	11
263	02141130	Zacks Fork Creek near Lenoir	35°55'32"	81°31'13"	1967-1976	10
264	02141890	Duck Creek near Taylorsville	35°53'34"	81°18'09"	1954–1971	18
265	02142000	Lower Little River near All Healing Springs	35°56'44"	81°14'13"	1954–1995	42
266	02142480	Hagan Creek near Catawba	35°40'20"	81°08'12"	1954-1971	15
267 ^r	02142500	Catawba River at Catawba	35°43'00"	81°03'59"	1936-1962	30
268	0214253830	Norwood Creek near Troutman	35°40'48"	80°56'44"	1984–1996	13
269	02142900	Long Creek near Paw Creek	35°19'42"	80°54'35"	1966–1996	31
270	02143000	Henry Fork near Henry River	35°41'03"	81°24'10"	1916–1996	59
271	02143040	Jacob Fork at Ramsey	35°35'26"	81°34'02"	1962-1996	35
272	02143310	Lithia Inn Branch near Lincolnton	35°27'47"	81°13'27"	1954-1971	14
273	02143500	Indian Creek near Laboratory	35°25'20"	81°15'52"	1916-1996	45
274	02144000	Long Creek near Bessemer City	35°18'23"	81°14'05"	1954–1996	43
275	02145000	South Fork Catawba River at Lowell	35°17'10"	81°06'00"	1940–1996	42
276	02146890	East Fork Twelve Mile Creek near Waxhaw	34°57'46"	80°42'40"	1954–1972	18
277	02146900	Twelve Mile Creek near Waxhaw	34°57'08"	80°45'21"	1949–1996	36
278 ^r	02148500	Broad River near Chimney Rock	35°25'29"	82°10'54"	1928-1958	31
279	02149000	Cove Creek near Lake Lure	35°25'24"	82°06'42"	1916–1996	45
280	02150420	Camp Creek near Rutherfordton	35°27'47"	81°54'29"	1955-1971	17
281	02151000	Second Broad River at Cliffside	35°14'08"	81°45'57"	1926–1996	71
282	02151500	Broad River near Boiling Springs	35°12'39"	81°41'52"	1926-1996	70
283	02152100	First Broad River near Casar	35°29'35"	81°40'56"	1960-1996	36
284	02152420	Big Knob Creek near Fallston	35°29'34"	81°32'25"	1953-1971	18
285	02152500	First Broad River near Lawndale	35°22'50"	81°32'40"	1916–1980	41
286	02152610	Sugar Branch near Boiling Springs	35°15'00"	81°37'15"	1954–1987	34
287	03160610	Old Field Creek near West Jefferson	36°21'29"	81°31'46"	1955-1971	17
288	03161000	South Fork New River near Jefferson	36°23'35"	81°24'26"	1916–1996	69
289	03162110	Buffalo Creek at Warrensville	36°27'22"	81°30'51"	1940-1971	17
290	03162500	North Fork New River at Crumpler	36°31'04"	81°23'18"	1878–1966	39
291	03162880	Vile Creek near Sparta	36°30'39"	81°06'16"	1955–1971	17
292	03439000	French Broad River at Rosman	35°08'32"	82°49'28"	1908-1996	62
293	03439500	French Broad at Calvert	35°08'55"	82°47'57"	1916-1955	31
294	03440000	Catheys Creek near Brevard	35°12'40"	82°47'00"	1945-1996	21
295	03441000	Davidson River near Brevard	35°16'23"	82°42'21"	1876-1996	73

[nc, flood-frequency estimates were not computed because the site has less than 10 years of peak-flow record; *, duplicate map identification and station number for sites having separate period of regulated or channelized flows; r, site excluded from regional analysis because flows were affected by regulation or channelization]

Map identification number (fig. 1)	Station number	Station name	Latitude	Longitude	Period of analysis	Number of systematic peaks
296	03441440	Little River above High Falls near Cedar Mountain	35°11'32"	82°36'49"	1963–1990	28
297	03441500	Little River near Penrose	35°13'23"	82°38'07"	1916–1973	13
298	03442000	Crab Creek near Penrose	35°14'02"	82°36'39"	1916–1965	13
299	03443000	French Broad River at Blantyre	35°17'56"	82°37'26"	1875-1996	76 13
300	03444000	Boylston Creek near Horseshoe	35°22'10"	82°33'50"	1943–1973	13
301	03444500	South Fork Mills River at The Pink Beds	35°21'59"	82°44'20"	1927-1973	31
302	03446000	Mills River near Mills River	35°23'55"	82°35'42"	1876–1996	64
303	03446410	Laurel Branch near Edneyville	35°22'15"	82°24'10"	1955–1970	12
304	03446500	Clear Creek near Hendersonville	35°21'14"	82°26'40"	1910–1965	10
305	03447000	Mud Creek at Naples	35°22'52"	82°29'54"	1916–1955	17
306	03447500	Cane Creek at Fletcher	35°26'08"	82°29'23"	1916-1973	18
307	03448000	French Broad River at Bent Creek	35°30'07"	82°35'33"	1916–1986	52
308	03448500	Hominy Creek at Candler	35°32'28"	82°40'35"	1940-1978	35
309 ^{nc}	0344894205	North Fork Swannanoa River near Walkertown	35°41'07"	82°19'58"	1990–1996	7
310	03449000	North Fork Swannanoa River near Black Mountain	35°39'11"	82°21'04"	1926–1952	27
310* ^{r,nc}	03449000*	North Fork Swannanoa River near Black Mountain (regulated period)	35°39'11"	82°21'04"	1953–1957	5
311	03450000	Beetree Creek near Swannanoa	35°39'11"	82°24'20"	1927-1996	61
312	03451000	Swannanoa River at Biltmore	35°34'06"	82°32'42"	1791–1979	51
312* ^r	03451000*	Swannanoa River at Biltmore (regulated period)	35°34'06"	82°32'42"	1980–1996	17
313	03451500	French Broad River at Asheville	35°36'33"	82°34'43"	1896–1996	101
314	03452000	Sandymush Creek near Alexander	35°43'49"	82°40'11"	1940–1955	13
315	03453000	Ivy Creek near Marshall	35°46'10"	82°37'16"	1876-1996	42
316	03453500	French Broad River at Marshall	35°47'10"	82°39'39"	1916–1996	54
317	03453880	Brush Creek at Walnut	35°50'40"	82°44'30"	1954–1971	17
318	03454000	Big Laurel Creek near Stackhouse	35°55'12"	82°45'42"	1935–1978	39
319	03454500	French Broad River at Hot Springs	35°53'23"	82°49'16"	1796–1978	15
320	03455500	West Fork Pigeon River above Lake Logan near Hazelwood	35°23'46"	82°56'17"	1955-1996	42
$321^{r,nc}$	0345577330	West Fork Pigeon River near Retreat	35°25'36"	82°55'12"	1989–1996	8
322 ^r	03456100	West Fork Pigeon River at Bethel	35°27'48"	82°54'00"	1955–1996	41
323	03456500	East Fork Pigeon River near Canton	35°27'42"	82°52'13"	1955–1996	42
324	03456991	Pigeon River near Canton	35°31'19"	82°50'53"	1810–1996	71
325	03457500	Allen Creek near Hazelwood	35°25'49"	83°00'30"	1950-1973	24
326 ^{nc}	03458500	Pigeon River nr Crabtree	35°34'37"	82°57'07"	1922-1930	9
327	03459000	Jonathan Creek near Cove Creek	35°37'21"	83°00'25"	1931–1973	43
328	03459500	Pigeon River near Hepco	35°38'05"	82°59'21"	1876–1996	69

Map identification number (fig. 1)	Station number	Station name	Latitude	Longitude	Period of analysis	Number of systematic peaks
329	03460000	Cataloochee Creek near Cataloochee	35°40'02"	83°04'22"	1935-1996	52
330	03461910	North Toe River at Newland	36°05'01"	81°55'45"	1955-1973	19
331	03462000	North Toe River at Altapass	35°53'59"	82°01'50"	1935-1978	24
332	03463300	South Toe River near Celo	35°49'53"	82°11'04"	1958–1996	39
333	03463500	South Toe River at Newdale	35°54'22"	82°11'19"	1916–1978	18
334	03463910	Phipps Creek near Burnsville	35°54'40"	82°22'10"	1957-1973	14
335	03464000	Cane River near Sioux	36°00'52"	82°19'40"	1893-1978	38
336	03464500	Nolichucky River at Poplar	36°04'29"	82°20'41"	1926-1978	30
337	03478910	Cove Creek at Sherwood	36°15'50"	81°47'03"	1940-1972	18
338	03479000	Watauga River near Sugar Grove	36°14'18"	81°49'22"	1916–1996	57
339	03480540	Peavine Branch near Banner Elk	36°10'20"	81°54'42"	1953-1972	11
340	03481000	Elk River near Elk Park	36°11'01"	81°57'45"	1935-1978	21
341	03500000	Little Tennessee River near Prentiss	35°08'59"	83°22'47"	1899–1996	52
342	03500240	Cartoogechaye Creek near Franklin	35°09'31"	83°23'40"	1949–1996	35
343 ^r	03500500	Cullasaja River at Highlands	35°04'14"	83°13'57"	1928-1971	44
344	03501000	Cullasaja River at Cullasaja	35°09'59"	83°19'25"	1908–1976	52
345	03501760	Coon Creek near Franklin	35°14'04"	83°20'28"	1957-1973	17
346	03502000	Little Tennessee River at Iotla	35°13'59"	82°23'32"	1899–1949	17
347	03503000	Little Tennessee River at Needmore	35°20'11"	83°31'37"	1899–1996	51
348	03504000	Nantahala River near Rainbow Springs	35°07'37"	83°37'09"	1940–1996	57
349 ^r	03505500	Nantahala River at Nantahala	35°17'55"	83°39'21"	1943–1982	39
350	03506500	Nantahala River at Almond	35°22'32"	83°33'59"	1923-1941	17
351	03507000	Little Tennessee River at Judson	35°24'30"	83°33'26"	1897–1944	48
352 ^{nc}	03508000	Tuckasegee River at Tuckasegee	35°16'55"	83°07'37"	1840-1940	6
352* ^r	03508000*	Tuckasegee River at Tuckasegee (regulated period)	35°16'55"	83°07'37"	1941–1976	37
353	03509000	Scott Creek above Sylva	35°23'02"	83°12'51"	1929–1995	48
354	03510500	Tuckasegee River At Dillsboro	35°22'00"	83°15'37"	1928-1940	13
354* ^r	03510500*	Tuckasegee River at Dillsboro (regulated period)	35°22'00"	83°15'37"	1941-1982	43
355	03511000	Oconaluftee River at Cherokee	35°29'04"	83°18'56"	1867-1949	28
356	03512000	Oconaluftee River at Birdtown	35°27'41"	83°21'13"	1946–1996	48
357	03513000	Tuckasegee River at Bryson City	35°25'40"	83°26'51"	1898–1940	43
357* ^r	03513000*	Tuckasegee River at Bryson City (regulated period)	35°25'40"	83°26'51"	1941-1995	55
358	03513410	Jenkins Branch tributary at Bryson City	35°24'50"	83°27'20"	1957-1971	13
359	03513500	Noland Creek near Bryson City	35°29'05"	83°30'15"	1936-1971	36
360	03514000	Hazel Creek at Proctor	35°28'38"	83°42'58"	1943-1952	10

Map identification Statior number numbe (fig. 1)		Station name	Latitude	Longitude	Period of analysis	Number of systematic peaks
361 ^{nc}	03515000	Little Tennessee River at Fontana Dam	35°26'45" 83°48'20	83°48'20"	1939–1944	6
361* ^r	03515000*	Little Tennessee River at Fontana Dam (regulated period)	35°26'45"	83°48'20"	1945-1954	10
362	03516000	Snowbird Creek near Robbinsville	35°18'40"	83°51'35"	1943-1952	10
363	03546000	Shooting Creek near Hayesville	35°01'29"	83°42'27"	1923–1955	13
364 ^r	03547000	Hiwassee River below Chatuge Dam near Hayesville	35°01'45"	83°47'45"	1943–1974	32
365	03548500	Hiwassee River above Murphy	35°04'49"	84°00'10"	1897-1941	44
365* ^r	03548500*	Hiwassee River above Murphy (regulated period)	35°04'49"	84°00'10"	1942-1996	55
366	03550000	Valley River at Tomotla	35°08'20"	83°58'50"	1898-1996	86

Table 2. Recurrence interval discharges and basin characteristics for gaged rural sites in North Carolina

[Q, recurrence interval flood discharge for years indicated; DA, drainage area; L, channel length; CSLOPE, channel slope; BSLOPE, basin slope; SHAPE, basin shape; CF, climate factor for recurrence interval years indicated; REG, region: 1, if site is in Blue Ridge-Piedmont; 2, if site is in Coastal Plain; 3, if site is in Sand Hills; nc, flood-frequency estimates were not computed because the site has less than 10 years of peak-flow record; *, duplicate map identification number for sites having separate periods of regulated or channelized flows; r, site excluded from regional analysis because flows were affected by regulation or channelization; n.a., data not available]

Map identification number (fig. 1)	Q ₂	Q ₅	Q ₁₀	Q ₂₅	Q ₅₀	Q ₁₀₀	Q ₂₀₀	Q ₅₀₀	DA (mi ²)	L (mi)	CSLOPE (ft/mi)	BSLOPE (ft/mi)	SHAPE (DA/L ²)	CF ₂	CF ₂₅	CF ₁₀₀	REG
1 2 3 4 ^{nc}	51.7 334 1950	99.7 643 2950	142 914 3690	210 1340 4720	271 1720 5550	343 2160 6430	426 2660 7370	557 3440 8720	0.7 11.8 225 3.7	1.17 5.84 30.86 2.65	8.61 2.93 3.18 2.38	4.33 13.25 16.19 4.10	0.53 .33 .23 .58	2.24 2.25 2.26 2.25	2.89 2.90 2.90 2.90	3.08 3.09 3.10 3.09	2 2 2 2
5 ^{nc}	n.a. n.a.	n.a. n.a.	n.a. n.a.	n.a. n.a.	n.a. n.a.	n.a. n.a.	n.a. n.a.	n.a. n.a.	24.0	8.03	2.18	9.79	.34	2.26	2.90	3.09	2
6 6* ^r 7 8 9	818 963 209 220 4600	1380 1340 253 426 7600	1830 1630 280 612 10000	2500 2070 313 912 13600	3070 2440 337 1190 16700	3700 2860 361 1520 20100	4390 3330 384 1900 24000	5440 4040 415 2510 29800	63.3 63.3 2.6 8.9 129	18.55 18.55 3.07 4.96 46.52	2.34 2.34 7.12 6.71 51.21	8.23 8.23 2.28 1.53 211.90	.18 .18 .25 .35 .06	2.26 2.26 2.26 2.31 2.07	2.90 2.90 2.90 2.94 2.77	3.10 3.10 3.10 3.14 2.95	
9* ^r 10 11 12 13	4210 128 617 879 7020	7370 186 945 1510 11800	9940 227 1190 2030 15700	13700 284 1540 2810 21600	17000 329 1820 3500 26700	20600 377 2120 4280 32600	24600 427 2450 5160 39100	30500 499 2920 6510 49200	129 .3 5.4 14.9 242	46.52 1.96 5.21 6.25 36.02	51.21 393.39 53.24 18.65 21.01	211.90 225.04 154.08 87.83 145.96	.06 .33 .20 .37 .20	2.07 2.08 2.08 2.11 2.10	2.77 2.78 2.77 2.80 2.79	2.95 2.95 2.95 2.97 2.97	1 1 1 1
14 15 16 17 ^{nc} 18	878 18300 960 n.a. 15000	1550 25500 1640 n.a. 24300	2130 30400 2200 n.a. 31600	3020 36600 3030 n.a. 42200	3810 41300 3730 n.a. 51100	4730 46100 4520 n.a. 61000	5770 51000 5410 n.a. 71900	7390 57600 6730 n.a. 88000	16.2 1053 12.0 1150 538	7.98 97.48 9.17 116.76 77.81	40.33 23.62 30.50 20.23 10.76	137.37 148.17 138.34 154.63 194.65	.26 .11 .14 .09 .09	2.11 2.10 2.10 2.10 2.10	2.80 2.80 2.79 2.79 2.79	2.97 2.97 2.97 2.97 2.97	1 1 1 1
18* ^r 19 20 21 22	10300 821 907 1870 44.2	15300 1640 1380 3710 77.6	19000 2390 1740 5370 106	24100 3590 2240 8010 149	28200 4700 2660 10400 188	32500 6010 3120 13200 232	37000 7540 3610 16500 282	43500 9960 4340 21600 359	538 29.9 7.1 45.9 .2	77.81 8.97 4.32 14.66 .39	10.76 20.81 40.20 13.59 19.18	194.65 117.88 103.92 118.79 66.32	.09 .41 .35 .22 .66	2.10 2.11 2.12 2.12 2.12 2.12	2.79 2.80 2.80 2.80 2.80 2.80	2.97 2.98 2.98 2.98 2.98	1 1 1 1
23 24 25 ^{r, nc} 26 ^r 27	724 1960 n.a. 4450 173	1270 3680 n.a. 8400 260	1720 5190 n.a. 10900 322	2410 7570 n.a. 13700 408	3010 9720 n.a. 15600 475	3700 12200 n.a. 17100 547	4480 15100 n.a. 18500 622	5670 19600 n.a. 20100 727	7.5 56.5 191 202 2.0	4.39 14.40 30.26 32.72 2.60	53.25 13.70 6.66 6.28 38.00	94.67 104.89 108.97 108.48 100.05	.39 .27 .22 .19 .28	2.12 2.12 2.11 2.14 2.16	2.80 2.80 2.80 2.82 2.83	2.98 2.98 2.98 3.01 3.01	1 1 1 1
28 ^r 29 29* ^r 30 31	385 77200 22800 51400 70.7	1010 105000 29400 78200 148	1700 125000 35100 99000 220	3030 151000 43900 129000 340	4430 172000 51600 154000 452	6280 194000 60500 181000 585	8690 217000 70800 212000 744	13000 249000 87000 257000 1000	53.5 8386 8386 8671 .9	16.06 280.12 280.12 311.75 1.17	13.21 4.03 4.03 3.83 18.78	108.45 159.83 159.83 155.40 10.29	.21 .11 .11 .09 .38	2.14 2.22 2.22 2.25 2.29	2.82 2.87 2.87 2.89 2.92	3.01 3.06 3.06 3.09 3.12	1 1 1 2 2

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Map identification number (fig. 1)	Q ₂	Q ₅	Q ₁₀	Q ₂₅	Q ₅₀	Q ₁₀₀	Q ₂₀₀	Q ₅₀₀	DA (mi ²)	L (mi)	CSLOPE (ft/mi)	BSLOPE (ft/mi)	SHAPE (DA/L ²)	CF2	CF ₂₅	CF ₁₀₀	REG
32 33 ^{nc} 34 35 36	637 n.a. 5040 348 5920	1010 n.a. 8190 685 9130	1290 n.a. 10600 997 11500	1700 n.a. 14000 1510 14900	2030 n.a. 16800 1990 17600	2400 n.a. 19800 2570 20500	2810 n.a. 23100 3270 23600	3400 n.a. 27800 4390 28100	17.1 108 167 3.3 427	6.76 24.98 30.12 5.41 52.26	5.83 2.45 11.33 35.92 7.45	9.42 9.75 77.62 104.99 93.14	0.42 .19 .18 .22 .16	2.27 2.36 2.17 2.18 2.19	2.91 2.98 2.84 2.84 2.85	3.10 3.16 3.02 3.02 3.03	2 2 1 1 1
37 38 ^{nc} 39 40 41 ^r	1190 n.a. 6690 913 7530	1990 n.a. 9690 1630 10000	2610 n.a. 11900 2220 11200	3500 n.a. 15000 3120 12400	4240 n.a. 17400 3900 13100	5050 n.a. 20000 4790 13700	5940 n.a. 22800 5790 14100	7230 n.a. 26900 7300 14600	47.8 660 701 64.8 777	14.33 76.62 87.60 19.85 95.12	14.77 4.14 4.17 4.83 3.81	93.74 90.51 88.94 44.34 84.58	.23 .11 .09 .17 .09	2.19 2.21 2.25 2.25 2.25	2.85 2.87 2.90 2.90 2.90	3.03 3.05 3.09 3.09 3.09	1 1 1 1 2
42 43 ^r 44 45 46 ^{nc}	56.2 8570 7700 266 n.a.	115 11300 10100 399 n.a.	170 12400 11700 498 n.a.	261 13400 13800 635 n.a.	347 13900 15300 746 n.a.	450 14300 16900 865 n.a.	572 14500 18500 993 n.a.	769 14800 20700 1180 n.a.	.3 925 930 8.6 13.4	.53 102.35 104.89 5.49 7.79	98.69 3.42 3.71 10.82 24.43	39.08 80.30 79.30 37.10 81.63	.76 .09 .09 .29 .22	2.20 2.25 2.25 2.26 2.18	2.86 2.90 2.90 2.90 2.85	3.04 3.09 3.09 3.09 3.03	2 2 2 2 2 2
47 48 49 50 ^{nc} 51	1820 1150 2410 n.a. 4590	3000 2170 3980 n.a. 7460	3930 3080 5250 n.a. 9720	5280 4540 7160 n.a. 13000	6430 5860 8810 n.a. 15700	7690 7420 10700 n.a. 18700	9090 9250 12800 n.a. 22000	11200 12100 15900 n.a. 26900	166 45.0 177 440 526	40.80 11.36 30.93 53.86 57.85	5.81 11.65 6.85 3.65 3.97	87.37 85.77 76.60 82.13 77.03	.10 .36 .19 .16 .16	2.24 2.17 2.24 2.24 2.24	2.89 2.84 2.89 2.89 2.89 2.89	3.08 3.02 3.08 3.08 3.08	2 2 2 2 2 2
52 53 54 55 56 ^{nc}	210 386 13900 819 n.a.	422 808 20300 1250 n.a.	619 1210 24900 1570 n.a.	946 1890 31200 2000 n.a.	1250 2540 36300 2350 n.a.	1630 3330 41700 2710 n.a.	2070 4290 47500 3100 n.a.	2790 5850 55700 3650 n.a.	9.4 11.7 2183 78.1 11.0	5.22 5.95 148.95 16.59 9.76	15.08 4.60 2.63 2.14 2.24	36.16 10.38 66.04 8.86 3.45	.37 .36 .10 .26 .13	2.23 2.25 2.26 2.27 2.26	2.88 2.89 2.90 2.91 2.91	3.07 3.09 3.09 3.10 3.10	2 2 2 2 2 2
57 ^{nc} 57* ^r 58 ^{nc} 58* ^{r,nc} 59	n.a. 1460 n.a. 671	n.a. 2260 n.a. n.a. 1160	n.a. 2690 n.a. n.a. 1570	n.a. 3110 n.a. 2190	n.a. 3370 n.a. 2720	n.a. 3570 n.a. 3320	n.a. 3740 n.a. n.a. 4000	n.a. 3920 n.a. 5020	45.0 45.0 7.5 7.5 29.0	10.82 10.82 4.90 4.90 9.75	3.64 3.64 9.66 9.66 3.43	12.34 12.34 13.23 13.23 16.82	.38 .38 .34 .34 .33	2.32 2.32 2.32 2.32 2.32 2.28	2.94 2.94 2.94 2.94 2.91	3.12 3.12 3.12 3.12 3.12 3.11	2 2 2 2 2 2
60 61 62 63 64 ^{nc}	244 103 443 623 n.a.	397 192 750 1150 n.a.	517 269 1000 1610 n.a.	687 388 1380 2330 n.a.	828 494 1710 2980 n.a.	982 615 2090 3720 n.a.	1150 754 2510 4580 n.a.	1390 968 3150 5920 n.a.	9.6 1.5 26.0 32.2 14.1	4.76 2.37 9.15 7.85 6.11	8.21 5.55 2.07 3.90 26.95	8.18 3.21 16.80 5.16 79.86	.43 .29 .28 .53 .38	2.32 2.32 2.37 2.32 2.15	2.95 2.95 2.98 2.95 2.82	3.14 3.14 3.17 3.14 2.99	2 2 2 2 1

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Map identification number (fig. 1)	Q ₂	Q 5	Q ₁₀	Q ₂₅	Q ₅₀	Q ₁₀₀	0 ₂₀₀	Q ₅₀₀	DA (mi ²)	L (mi)	CSLOPE (ft/mi)	BSLOPE (ft/mi)	SHAPE (DA/L ²)	CF ₂	CF ₂₅	CF ₁₀₀	REG
65 66 67 68 69	2770 87.2 4840 156 3470	4340 165 8500 273 6180	5550 232 11400 371 8410	7270 339 15800 519 11700	8680 434 19400 649 14600	10200 545 23400 796 17800	11900 673 27800 962 21400	14300 872 34400 1220 26700	66.0 .8 141 1.0 78.2	15.62 1.11 33.53 1.34 18.92	11.78 126.53 11.00 39.06 15.26	81.54 135.53 92.00 63.54 76.96	0.27 .72 .13 .52 .22	2.17 2.17 2.17 2.17 2.17 2.17	2.84 2.84 2.84 2.84 2.84	3.01 3.01 3.01 3.01 3.01	1 1 1 1
70 71 72 ^r 73 74 ^{nc}	6820 337 6650 2210 n.a.	10800 606 10200 3980 n.a.	13800 835 12400 5470 n.a.	18000 1190 15000 7750 n.a.	21400 1510 16800 9770 n.a.	25100 1870 18400 12100 n.a.	29100 2280 19900 14700 n.a.	34800 2930 21700 18700 n.a.	149 4.8 168 43.0 21.9	24.50 5.13 29.13 14.25 12.15	12.08 31.73 11.83 17.05 14.29	79.98 112.77 82.17 85.63 84.86	.25 .18 .20 .21 .15	2.17 2.17 2.17 2.17 2.17 2.17	2.84 2.84 2.84 2.84 2.84	3.01 3.01 3.01 3.02 3.01	1 1 1 1
75 76 77 78 79	8500 868 696 109 6960	13200 1330 827 195 9670	16900 1670 908 267 11600	22100 2140 1000 377 14300	26500 2520 1070 472 16400	31200 2920 1140 581 18600	36400 3350 1210 704 21000	44000 3960 1300 891 24400	535 10.1 13.8 .7 771	47.55 5.01 7.04 1.06 64.87	9.18 16.67 15.49 98.62 7.08	83.44 91.21 100.58 100.04 87.30	.24 .40 .28 .60 .18	2.18 2.18 2.18 2.18 2.18 2.19	2.84 2.84 2.84 2.85 2.85	3.02 3.02 3.02 3.02 3.03	1 1 1 1
79* ^r 80 81 81* ^r 82	4820 49.4 9710 7090 8810	5900 85.2 13200 10000 11400	6620 115 15700 12300 13300	7540 159 18900 15600 15600	8230 197 21400 18300 17500	8930 240 23900 21300 19300	9640 289 26600 24600 21300	10600 363 30300 29600 24000	771 .3 1150 1150 1206	64.87 .73 94.85 94.85 108.53	7.08 75.75 4.90 4.90 4.37	87.30 111.31 87.62 87.62 86.24	.18 .48 .13 .13 .10	2.19 2.18 2.25 2.25 2.26	2.85 2.84 2.88 2.88 2.88	3.03 3.02 3.05 3.05 3.05	1 1 1 1
82* ^r 83 84 ^{nc} 85 86	7630 1390 n.a. 535 1410	10300 2100 n.a. 980 2620	11700 2630 n.a. 1360 3670	13000 3370 n.a. 1960 5310	13700 3970 n.a. 2490 6790	14300 4610 n.a. 3100 8510	14700 5300 n.a. 3800 10500	15200 6280 n.a. 4890 13600	1206 19.5 35.8 8.2 83.5	108.53 7.18 11.55 6.53 22.16	4.37 21.38 13.54 23.77 9.08	86.24 89.93 89.20 82.99 81.46	.10 .37 .27 .19 .17	2.26 2.23 2.24 2.23 2.25	2.88 2.87 2.87 2.87 2.88	3.05 3.03 3.04 3.03 3.04	1 1 1 1
87 88 89 90 91	575 139 487 1630 2320	1120 296 1030 2620 3500	1610 449 1530 3390 4350	2430 711 2370 4500 5500	3180 965 3160 5420 6410	4090 1280 4100 6430 7360	5160 1660 5230 7530 8350	6890 2300 7040 9150 9740	27.9 2.6 7.6 191 232	9.77 3.45 5.36 39.18 49.28	11.67 34.25 21.29 5.87 5.32	46.15 80.30 45.76 56.06 50.49	.29 .23 .27 .12 .09	2.26 2.26 2.25 2.25 2.25 2.26	2.88 2.88 2.89 2.89 2.89 2.89	3.05 3.04 3.06 3.06 3.06	2 2 2 2 2 2
92 92* ^r 93 ^{nc} 94 94* ^r	12700 10200 n.a. 13500 10800	18300 15700 n.a. 19800 15700	22500 19700 n.a. 24400 18900	28200 25100 n.a. 30800 22800	32700 29500 n.a. 35900 25500	37600 34000 n.a. 41300 28200	42800 38900 n.a. 47200 30900	50200 45700 n.a. 55500 34200	2399 2399 57.7 2692 2692	169.34 169.34 15.33 203.19 203.19	2.78 2.78 4.61 2.10 2.10	68.54 68.54 19.67 63.32 63.32	.08 .08 .25 .07 .07	2.31 2.31 2.32 2.33 2.33	2.93 2.93 2.93 2.94 2.94	3.09 3.09 3.10 3.11 3.11	2 2 2 2 2 2

Tables

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Map identification number (fig. 1)	Q ₂	Q ₅	Q ₁₀	Q ₂₅	Q ₅₀	Q ₁₀₀	Q ₂₀₀	Q ₅₀₀	DA (mi ²)	L (mi)	CSLOPE (ft/mi)	BSLOPE (ft/mi)	SHAPE (DA/L ²)	CF ₂	CF ₂₅	CF ₁₀₀	REG
95 95* ^r 96 97 98	1760 2160 171 121 126	2860 3250 293 252 258	3740 3900 392 376 381	5010 4620 538 584 584	6080 5090 662 782 774	7260 5520 799 1020 1000	8570 5900 952 1310 1270	10500 6350 1180 1780 1710	161 161 2.8 2.1 2.6	28.02 28.02 3.74 2.12 2.51	6.14 6.14 11.87 15.71 17.67	55.52 55.52 30.38 20.07 8.17	0.20 .20 .21 .43 .44	2.25 2.25 2.28 2.29 2.29	2.90 2.90 2.91 2.92 2.92	3.09 3.09 3.09 3.10 3.10	2 2 2
99 100 ^{nc} 101 102 103	386 n.a. 1090 70 4000	656 n.a. 1740 125 6410	879 n.a. 2250 191 8220	1220 n.a. 3000 301 10800	1510 n.a. 3620 407 12800	1840 n.a. 4310 535 15000	2210 n.a. 5060 688 17400	2790 n.a. 6190 937 20800	18.6 1.9 80.4 1.5 733	10.51 3.15 22.82 2.11 71.80	7.48 13.58 4.03 43.56 2.84	38.55 10.80 28.29 39.01 35.09	.16 .23 .15 .34 .14	2.30 2.30 2.33 2.32 2.32	2.92 2.92 2.95 2.93 2.93	3.09 3.10 3.13 3.10 3.11	2 2 2
104 105 106 107 108	1420 340 501 1930 517	2150 711 1000 3090 1140	2690 1060 1460 3990 1760	3450 1650 2210 5250 2830	4060 2200 2900 6290 3870	4730 2870 3730 7420 5140	5440 3670 4710 8640 6710	6480 4970 6290 10400 9300	93.3 4.9 27.0 182 24.0	18.49 3.71 8.46 27.24 6.29	3.08 7.90 5.17 2.11 3.48	17.46 21.89 11.26 9.57 16.11	.28 .35 .41 .25 .57	2.32 2.33 2.33 2.33 2.33	2.93 2.94 2.94 2.94 2.94	3.11 3.11 3.12 3.12 3.12	2
109 110 111 112 113	869 217 1760 230 141	1590 390 2910 435 367	2220 539 3830 613 620	3220 772 5180 892 1110	4120 982 6340 1140 1630	5170 1220 7640 1430 2330	6390 1500 9080 1770 3240	8310 1940 11300 2290 4890	33.6 2.5 168 6.3 3.3	8.56 3.23 32.46 3.77 3.60	2.56 5.47 2.04 11.66 5.93	6.86 11.12 17.33 7.75 2.05	.51 .44 .16 .46 .28	2.34 2.34 2.34 2.33 2.39	2.95 2.94 2.94 2.94 2.99	3.12 3.11 3.11 3.11 3.17	2 2 2
114 115 116 117 118	606 122 1570 109 769	1300 263 2760 214 1570	1990 403 3750 308 2320	3190 645 5260 455 3580	4360 883 6590 588 4760	5810 1180 8090 743 6190	7610 1540 9800 921 7910	10600 2150 12400 1200 10700	53.3 4.9 94 1.0 26.9	15.10 2.55 15.74 1.21 10.28	2.22 10.28 4.07 31.68 6.16	5.60 15.94 24.27 22.23 22.31	.26 .71 .33 .38 .25	2.36 2.40 2.35 2.35 2.40	2.96 2.99 2.95 2.95 2.95 2.97	3.13 3.15 3.11 3.11 3.11	
119 120 121 ^{nc} 122 123	469 1670 n.a. 899 662	758 3020 n.a. 1630 1130	982 4180 n.a. 2240 1520	1300 6020 n.a. 3150 2120	1560 7670 n.a. 3940 2640	1850 9590 n.a. 4820 3240	2160 11800 n.a. 5800 3930	2620 15300 n.a. 7280 4980	26.3 168 188 20.6 15.9	12.57 32.93 42.64 9.17 7.48	10.60 6.66 5.17 19.56 21.30	85.79 89.48 90.27 88.69 81.00	.17 .16 .10 .25 .29	2.11 2.12 2.12 2.11 2.11 2.12	2.80 2.80 2.81 2.80 2.80	2.97 2.98 2.98 2.97 2.98	1 1
124 125 126 127 128	1660 2110 11500 1260 3740	2900 3600 18200 2330 5340	3950 4830 23300 3250 6460	5580 6690 30600 4680 7920	7030 8320 36500 5950 9050	8700 10200 42900 7410 10200	10600 12200 49800 9090 11400	13600 15400 59800 11700 13100	33.6 37.1 606 14.6 116	13.47 12.83 52.96 6.27 20.29	12.21 12.47 6.84 21.48 12.27	74.59 73.20 85.86 86.92 88.54	.19 .23 .22 .37 .28	2.12 2.12 2.14 2.17 2.13	2.80 2.80 2.81 2.84 2.81	2.98 2.98 2.98 3.02 2.98	1 1 1

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Map identification number (fig. 1)	Q ₂	Q ₅	Q ₁₀	Q ₂₅	Q ₅₀	Q ₁₀₀	Q ₂₀₀	Q ₅₀₀	DA (mi ²)	L (mi)	CSLOPE (ft/mi)	BSLOPE (ft/mi)	SHAPE (DA/L ²)	CF ₂	CF ₂₅	CF ₁₀₀	REG
129 130 ^{nc}	221 n.a.	566 n.a.	945 n.a.	1660 n.a.	2410 n.a.	3390 n.a.	4650 n.a.	6880 n.a.	5.0 7.5	4.16 4.66	22.72 25.70	65.39 94.32	0.22 .35	2.13 2.18	2.81 2.83	2.98 2.99	1 1
131	1820	2860	3650	4760	5670	6650	7710	9240	33.7	9.75	18.51	93.90	.35	2.18	2.83	2.99	1
132	25000	36600	45000	56300	65300	74800	84800	99000	1275	82.36	6.20	85.28	.19	2.21	2.85	3.01	1
133	175	313	428	602	755	927	1120	1420	1.1	1.93	40.95	95.44	.41	2.19	2.83	3.00	1
134	2370	4060	5380	7250	8790	10400	12200	14800	75.9	21.38	18.18	93.48	.17	2.21	2.85	3.01	1
135	1300	2260	3060	4270	5330	6540	7900	9990 2650	21.1	8.79	11.61	78.16	.27	2.21	2.85	3.02	1
136 137	828 3880	1250 5610	1550 6850	1980 8500	2330 9800	2690 11200	3090 12600	3650 14600	23.6 285	11.69 34.81	16.49 11.44	90.08 91.08	.19 .23	2.22 2.22	2.86 2.86	3.02 3.02	1
137 138^{r}	14800	17000	17600	18000	18200	18200	12000	14000	1689	91.01	6.68	87.29	.23	2.22	2.80	3.02	1
139 ^{nc}	n.a.	n.a.	n.a.	n.a.	n.a.	n.a.	n.a.	n.a.	1689	91.01	6.68	87.29	.20	.00	.22	.86	1
140	1590	2590	3390	4570	5580	6690	7930	9790	32.1	9.71	15.85	66.01	.34	2.16	2.82	2.99	1
141	1660	2770	3630	4860	5890	7000	8210	9960	14.8	6.54	18.72	79.92	.34	2.12	2.80	2.98	1
142 nc	4760	7290	9170	11800	13900	16200	18600	22100	125	23.34	10.90	83.89	.23	2.17	2.82	2.99	1
143 ^{nc}	n.a.	n.a.	n.a.	n.a.	n.a.	n.a.	n.a.	n.a.	16.7	9.61	19.55	79.42	.18	2.17	2.82	2.99	1
144	12100	17400	21200	26400	30600	35100	39700	46400	349	45.45	8.79	94.27	.17	2.18	2.83	3.00	1
145	6400	11700	16400	23900	30700	38700	48000	62700	137	18.60	17.28	90.65	.40	2.20	2.84	3.01	1
146 147	521 143	860 266	1130 371	1520 533	1850 675	2220 838	2620 1020	3230 1310	3.0 .9	2.38 1.60	36.12 68.37	70.12 91.52	.48 .32	2.19 2.24	2.84 2.87	3.00 3.03	1
147 148 ^{nc}	143 n.a.	200 n.a.	571 n.a.	555 n.a.	075 n.a.	030 n.a.	n.a.	n.a.	.9 7.4	6.09	22.99	69.92	.32	2.24	2.87	3.00	1
149	1090	2050	2870	4160	5300	6620	8140	10500	15.5	8.06	24.42	85.61	.24	2.19	2.83	3.00	1
150	2920	4480	5650	7310	8660	10100	11700	14000	43.2	16.66	9.64	61.78	.15	2.19	2.83	3.00	1
151	21800	29600	35200	42500	48300	54300	60600	69500	1434	115.73	5.50	91.10	.11	2.23	2.86	3.02	1
152 ^{nc}	n.a.	n.a.	n.a.	n.a.	n.a.	n.a.	n.a.	n.a.	76.3	15.61	9.64	103.60	.31	2.23	2.87	3.03	1
152* ^r	746	1470	2170	3380	4570	6070	7930	11100	76.3	15.61	9.64	103.60	.31	2.23	2.87	3.03	1
153	42100	57300	68000	82300	93500	105000	118000	135000	3464	119.72	6.22	93.60	.24	2.25	2.87	3.04	1
153* ^r	28700	36300	41100	46700	50800	54700	58600	63600	3464	119.72	6.22	93.60	.24	2.25	2.87	3.04	1
154 155	146 90.9	229 168	293 236	385 347	462 449	546 571	638 715	772 945	7.6 2.2	5.87 2.10	42.55 63.92	86.78 88.21	.22 .48	2.25 2.21	2.87 2.84	3.03 3.01	3 3
155	90.9 863	168	230	2770	449 3420	4160	4980	945 6220	2.2 32.4	2.10	63.92 17.11	88.21 86.34	.48 .27	2.21	2.84 2.87	3.01	3 3
157	2760	3700	4340	5190	5850	6530	7250	8230	348	40.74	7.27		.21	2.25	2.87	3.03	
157	120	3700 199	4340 262	356	5850 436	527	628	8230 781	548 7.6	40.74	17.00	84.77 77.76	.21	2.25 2.25	2.87	3.03 3.04	3 3
159	3570	5540	7100	9370	11300	13400	15800	19300	459	57.53	5.33	80.68	.14	2.25	2.88	3.04	3
160	46000	63400	75700	92000	105000	118000	132000	151000	4395	156.35	5.15	89.44	.18	2.26	2.88	3.04	2
161	175	319	442	634	806	1000	1230	1590	7.9	4.81	7.27	23.11	.43	2.26	2.88	3.04	2

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Map identification number (fig. 1)	Q ₂	Q ₅	Q ₁₀	Q ₂₅	Q ₅₀	Q ₁₀₀	Q ₂₀₀	Q ₅₀₀	DA (mi ²)	L (mi)	CSLOPE (ft/mi)	BSLOPE (ft/mi)	SHAPE (DA/L ²)	CF2	CF ₂₅	CF ₁₀₀	REG
162 163 163* ^r 164 165	2090 36500 28500 147 490	3600 45500 33500 385 1000	4870 51200 36200 654 1490	6790 58300 39100 1170 2310	8480 63500 41000 1730 3100	10400 68600 42700 2480 4060	12600 73800 44200 3460 5210	15900 80600 46100 5230 7120	292 4852 4852 14.1 60.1	39.33 176.67 176.67 7.56 19.22	5.14 4.67 4.67 12.15 3.19	66.05 86.19 86.19 24.86 11.61	0.18 .16 .16 .30 .16	2.29 2.30 2.30 2.31 2.32	2.89 2.90 2.90 2.90 2.90 2.90	3.05 3.05 3.05 3.05 3.05	2 2
166 166* ^r 167 168 169	32500 25100 657 861 395	47400 33900 1200 1480 764	58200 39800 1660 1990 1090	72900 47400 2360 2760 1600	84600 53100 2970 3420 2050	96900 58800 3660 4170 2580	110000 64700 4440 5010 3190	128000 72600 5630 6290 4130	5255 5255 21.6 92.8 15.7	229.76 229.76 7.51 24.69 5.78	3.55 3.55 8.09 4.25 6.99	81.22 81.22 12.90 22.86 19.12	.10 .10 .38 .16 .45	2.38 2.38 2.39 2.29 2.33	2.95 2.95 2.95 2.89 2.92	3.09 3.09 3.09 3.04 3.07	
170 171 172 173 174	56.9 4140 555 1970 488	96.8 6860 1110 2990 732	129 9050 1640 3740 912	178 12300 2510 4790 1160	219 15000 3330 5630 1360	265 18100 4330 6530 1580	317 21500 5540 7500 1810	395 26700 7510 8880 2140	.5 676 32.3 379 103	.86 52.96 11.64 65.20 29.39	8.83 2.15 7.19 2.72 1.79	41.15 31.48 23.35 25.51 6.03	.58 .24 .25 .09 .12	2.33 2.33 2.28 2.32 2.38	2.92 2.91 2.89 2.90 2.95	3.07 3.05 3.04 3.05 3.09	2 2 2 2 2 2
175 176 177 178 179	31.2 940 139 1070 4810	68.7 1590 310 2100 7600	105 2110 482 3040 9780	166 2880 788 4570 12900	224 3530 1090 5980 15600	295 4260 1470 7660 18400	379 5060 1950 9640 21600	517 6270 2760 12800 26300	.6 47.5 8.6 49.7 599	1.41 10.83 4.84 15.46 47.76	31.94 5.65 19.02 4.84 2.47	10.32 18.88 36.42 24.90 26.63	.33 .41 .36 .22 .27	2.32 2.32 2.33 2.38 2.34	2.93 2.93 2.93 2.95 2.93	3.09 3.09 3.10 3.09 3.09	2 2
180 181 182 183 184	1510 280 129 326 411	2600 345 331 731 700	3490 386 550 1140 936	4820 436 953 1870 1290	5980 472 1370 2600 1590	7270 508 1900 3520 1920	8720 543 2570 4670 2300	10900 589 3720 6630 2870	69.3 7.8 1.1 10.2 15.3	13.50 5.35 1.27 4.40 4.35	5.64 4.79 3.71 6.96 5.03	26.50 12.77 .81 7.38 6.70	.37 .29 .70 .49 .75	2.37 2.38 2.39 2.40 2.32	2.94 2.94 2.95 2.96 2.91	3.08 3.08 3.10 3.10 3.06	2 2 2
185 186 187 188 189	3870 416 150 1380 4070	6090 796 323 2730 8150	7780 1130 489 4000 11900	10100 1660 772 6120 18100	12100 2140 1040 8150 23900	14200 2700 1370 10600 30800	16400 3350 1770 13600 39000	19600 4370 2420 18500 52300	680 16.0 3.8 28.8 48.1	38.60 5.02 3.57 15.93 20.64	.87 4.36 9.32 89.91 48.76	7.67 .72 17.28 323.81 314.89	.47 .66 .30 .12 .12	2.33 2.35 2.32 2.15 2.14	2.92 2.93 2.91 2.78 2.78	3.07 3.08 3.06 2.94 2.95	
190 191 192 192* ^r 193	456 3640 12700 7340 6280	768 6000 20900 9530 12000	1020 7830 27600 11000 17000	1400 10500 37800 12700 24700	1730 12700 46700 14100 31500	2100 15000 56800 15400 39300	2520 17600 68300 16700 48200	3150 21400 85800 18500 61800	11.0 89.2 504 504 128	6.66 20.61 48.21 48.21 20.23	386.26 32.65 21.29 21.29 33.91	325.50 265.44 243.78 243.78 254.90	.26 .20 .22 .22 .31	2.11 2.13 2.13 2.13 2.09	2.77 2.79 2.79 2.79 2.79 2.77	2.95 2.96 2.96 2.96 2.95	1 1 1

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Map identification number (fig. 1)	Q ₂	Q ₅	Q ₁₀	Q ₂₅	Q ₅₀	Q ₁₀₀	Q ₂₀₀	Q ₅₀₀	DA (mi ²)	L (mi)	CSLOPE (ft/mi)	BSLOPE (ft/mi)	SHAPE (DA/L ²)	CF ₂	CF ₂₅	CF ₁₀₀	REG
194 195 ^r 196 197 198	3250 16200 3270 1800 4710	4640 21900 5300 2650 6530	5630 25000 6850 3260 7810	6960 28300 9020 4090 9490	8010 30400 10800 4750 10800	9110 32200 12700 5440 12100	10300 33800 14700 6180 13500	11900 35700 17600 7220 15500	35.5 869 78.8 44.7 109	14.74 70.50 22.18 15.33 27.61	32.24 12.10 31.52 27.18 14.56	168.32 218.10 281.93 257.85 202.20	0.16 .17 .16 .21 .15	2.08 2.08 2.08 2.08 2.08	2.77 2.77 2.77 2.77 2.77	2.95 2.95 2.95 2.95 2.95 2.95	1 1 1 1
199 200 ^r 201 202 203	5450 28000 6730 7760 3240	9070 35300 11200 13200 5400	12000 39200 14800 17500 7100	16500 43300 20000 23700 9550	20300 45800 24300 28800 11600	24700 48100 29100 34300 13800	29600 50100 34300 40400 16300	37000 52500 42000 49200 19900	128 1226 231 287 42.8	30.68 93.13 28.32 31.47 13.55	13.23 6.98 19.16 17.57 22.17	198.38 208.67 173.69 165.49 145.31	.13 .14 .29 .29 .25	2.08 2.20 2.08 2.08 2.09	2.77 2.82 2.77 2.77 2.78	2.95 2.99 2.95 2.95 2.95 2.95	1 1 1 1
204 ^r 205 206 207 208 ^{nc}	39200 1150 217 1570 n.a.	53100 1700 357 2900 n.a.	59000 2100 467 4040 n.a.	63800 2650 625 5830 n.a.	66100 3090 756 7430 n.a.	67700 3560 898 9270 n.a.	68800 4050 1050 11400 n.a.	69800 4760 1280 14700 n.a.	1694 22.1 .9 19.5 10.2	110.42 11.48 1.13 8.29 6.32	5.25 21.49 85.26 38.11 28.92	193.78 105.01 88.49 134.61 66.74	.14 .17 .64 .26 .28	2.10 2.09 2.09 2.14 2.10	2.78 2.78 2.78 2.80 2.79	2.96 2.96 2.95 2.97 2.97	1 1 1 1
209 ^{nc} 210 ^{nc} 211 212 213	n.a. n.a. 208 2660 4210	n.a. n.a. 398 3280 7040	n.a. n.a. 570 3680 9300	n.a. n.a. 850 4160 12600	n.a. n.a. 1110 4510 15400	n.a. n.a. 1420 4860 18500	n.a. n.a. 1780 5200 21900	n.a. n.a. 2360 5650 27000	27.8 6.8 2.2 65.6 186	10.14 6.67 1.69 16.39 25.82	19.34 31.49 47.86 13.47 10.45	69.36 60.00 66.44 86.79 90.93	.27 .16 .73 .24 .27	2.10 2.12 2.12 2.12 2.12 2.13	2.79 2.80 2.80 2.80 2.80 2.80	2.97 2.98 2.98 2.98 2.98	1 1 1 1
214 215 215* ^r 216 217	1250 30200 32700 93.2 227	1950 43600 46400 180 316	2470 53500 54600 258 378	3210 67200 64100 382 458	3820 78300 70700 494 520	4460 90100 76700 625 583	5160 103000 82400 778 648	6170 121000 89400 1020 737	42.9 2280 2280 1.0 1.6	14.12 149.50 149.50 1.60 2.80	15.63 4.38 4.38 56.18 78.01	92.00 169.42 169.42 100.24 132.30	.21 .10 .10 .40 .19	2.12 2.15 2.15 2.15 2.15 2.15	2.80 2.80 2.80 2.80 2.80 2.80	2.98 2.98 2.98 2.98 2.98 2.97	1 1 1 1
218 219 220 221 222 ^r	2660 4060 5080 6940 63.2	4240 6610 7620 10200 68.3	5450 8570 9450 12500 70.9	7140 11300 11900 15900 73.6	8520 13600 13900 18500 75.3	10000 16100 16000 21400 76.9	11600 18700 18100 24500 78.2	13900 22500 21200 29000 79.9	101 306 155 569 4.8	34.64 48.80 30.06 56.83 4.56	18.90 11.36 13.21 10.23 29.75	138.79 120.89 148.05 125.33 52.42	.08 .13 .17 .17 .23	2.15 2.15 2.15 2.16 2.15	2.80 2.80 2.80 2.80 2.80 2.80	2.97 2.98 2.97 2.98 2.97	1 1 1 1
223 223* ^r 224 225 226	1430 1490 2750 560 53300	2080 1830 4040 1040 79500	2550 2050 4980 1440 98500	3200 2310 6300 2070 124000	3710 2490 7360 2620 145000	4260 2680 8490 3260 166000	4830 2860 9710 3980 189000	5650 3100 11500 5090 221000	87.4 87.4 118 3.9 3450	33.20 33.20 16.59 3.17 170.84	11.75 11.75 12.07 28.98 3.90	91.45 91.45 96.08 99.43 170.74	.08 .08 .43 .36 .12	2.16 2.16 2.16 2.16 2.16	2.80 2.80 2.81 2.81 2.81	2.98 2.98 2.98 2.98 2.98 2.98	1 1 1 1

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Map identification number (fig. 1)	0 ₂	0 ₅	Q ₁₀	Q ₂₅	0 ₅₀	Q ₁₀₀	Q ₂₀₀	Q ₅₀₀	DA (mi ²)	L (mi)	CSLOPE (ft/mi)	BSLOPE (ft/mi)	SHAPE (DA/L ²)	CF ₂	CF ₂₅	CF ₁₀₀	REG
227 228 229 230 ^{nc} 230* ^r	566 4640 449 n.a. 41300	934 6540 685 n.a. 51800	1230 7870 862 n.a. 58000	1650 9640 1110 n.a. 65200	2010 11000 1310 n.a. 70200	2400 12400 1530 n.a. 74900	2840 13900 1760 n.a. 79400	3480 16000 2100 n.a. 85000	9.6 174 6.6 4000 4000	5.98 29.85 6.14 186.48 186.48	17.12 9.77 16.73 3.43 3.43	96.67 97.43 89.50 160.97 160.97	0.27 .20 .17 .11 .11	2.15 2.17 2.17 2.18 2.18	2.81 2.82 2.82 2.83 2.83	2.98 2.99 3.00 3.00 3.00	1 1 1 1
231 232 233 234 235	852 397 7730 425 571	955 715 10500 719 1080	1020 990 12500 959 1520	1090 1420 15100 1320 2230	1130 1800 17000 1630 2880	1180 2250 19100 1970 3630	1220 2760 21200 2360 4510	1280 3570 24200 2940 5890	13.7 2.9 342 3.4 3.6	6.59 2.80 51.53 4.41 3.55	21.96 45.36 6.63 72.33 35.95	86.01 113.22 131.78 173.68 95.22	.31 .37 .13 .18 .28	2.19 2.19 2.19 2.20 2.18	2.83 2.83 2.83 2.83 2.83 2.81	3.00 3.00 3.00 3.00 2.98	1 1 1 1
236 237 238 239 240	1670 4630 1400 33200 2190	2420 7020 2320 47300 4120	2960 8770 3040 56800 5850	3700 11100 4090 69100 8630	4290 13000 4970 78400 11200	4920 15000 5930 87900 14200	5580 17100 6990 97500 17800	6520 20000 8540 111000 23600	20.7 55.6 8.5 1372 110	7.02 14.09 5.26 81.80 26.68	29.69 21.06 27.59 5.70 7.72	93.64 90.60 55.44 84.63 83.54	.42 .28 .28 .21 .16	2.18 2.20 2.22 2.23 2.24	2.81 2.83 2.83 2.85 2.85	2.99 3.00 3.00 3.01 3.01	1 1 1 1
241 242 243 244 ^r 245	176 4490 1030 73500 845	284 6350 1750 106000 1220	367 7650 2350 132000 1480	486 9360 3270 170000 1850	584 10700 4070 202000 2140	690 12000 4970 237000 2440	806 13500 6000 276000 2760	975 15400 7570 335000 3210	.9 106 15.4 6863 16.7	1.06 29.55 9.24 255.33 5.91	89.92 12.14 47.53 4.59 33.41	65.70 108.15 142.49 134.56 75.65	.73 .12 .18 .11 .47	2.24 2.20 2.20 2.25 2.24	2.85 2.84 2.84 2.86 2.86	3.01 3.00 3.01 3.02 3.02	1 1 1 1
246 247 248 ^{nc} 249 250	16.8 107 n.a. 363 1350	28.7 195 n.a. 528 2430	38.6 272 n.a. 646 3400	53.4 393 n.a. 806 4990	66.2 502 n.a. 932 6460	80.7 630 n.a. 1060 8230	97.1 778 n.a. 1200 10300	122 1010 n.a. 1400 13800	.1 6.2 .1 83.3 183	1.03 5.46 .37 20.22 26.23	114.65 11.97 119.76 7.30 10.26	86.40 23.56 58.22 40.58 76.68	.14 .19 1.45 .20 .27	2.25 2.26 2.25 2.29 2.23	2.86 2.87 2.86 2.89 2.85	3.02 3.03 3.03 3.04 3.02	1 3 3 3 3
251 252 253 254 255	77 1450 467 232 4890	116 2200 648 345 7610	144 2760 774 428 9660	184 3530 942 541 12500	215 4160 1070 633 14900	248 4830 1210 729 17400	284 5560 1350 832 20100	334 6600 1550 979 24000	4.7 365 39.8 16.1 1228	4.32 64.79 14.30 6.85 130.44	33.83 4.11 7.65 3.83 2.09	66.99 63.07 24.83 7.32 27.34	.24 .09 .21 .36 .07	2.25 2.29 2.28 2.30 2.31	2.86 2.89 2.88 2.89 2.90	3.03 3.04 3.04 3.05 3.05	3 3 3 3 3
256 257 258 259 260	1120 6270 7050 4580 57.3	1810 9440 12500 8750 93.4	2370 11800 17200 12500 122	3200 14900 24600 18600 163	3920 17500 31300 24200 197	4710 20200 39200 30900 234	5610 23000 48300 38700 275	6950 27100 62700 51300 336	20.7 127 172 66.7 0.5	8.95 20.20 23.12 34.93 1.07	268.53 67.00 52.06 86.54 526.60	407.99 299.40 299.74 285.37 243.48	.26 .31 .32 .05 .48	2.19 2.18 2.18 2.17 2.18	2.78 2.78 2.78 2.78 2.78 2.78	2.94 2.94 2.94 2.94 2.95	1 1 1 1

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Map identification number (fig. 1)	Q ₂	Q ₅	Q ₁₀	Q ₂₅	Q ₅₀	Q ₁₀₀	Q ₂₀₀	Q ₅₀₀	DA (mi ²)	L (mi)	CSLOPE (ft/mi)	BSLOPE (ft/mi)	SHAPE (DA/L ²)	CF ₂	CF ₂₅	CF ₁₀₀	REG
261 262 263 264	229 11900 449 945	420 21200 776 1560	585 29000 1040 2040	840 41000 1440 2760	1070 51600 1790 3370	1330 63600 2170 4040	1630 77400 2600 4790	2100 98500 3250 5910	2.4 201 9.1 18.4	3.09 34.09 6.71 9.83	572.66 20.52 31.80 44.78	225.03 277.87 227.46 189.99	.19	2.16 2.17 2.16 2.15	2.78 2.78 2.78 2.79	2.94 2.95 2.95 2.97	1 1 1 1
265 266 267 ^r 268 269	1520 853 20500 568 1330	2660 1430 39400 1010 2040	3610 1900 58400 1380 2590	5020 2590 92700 1940 3360	6240 3190 128000 2420 4010	7610 3850 174000 2970 4700	9140 4600 233000 3590 5460	11400 5720 338000 4540 6570	28.2 7.8 1535 7.2 16.4	8.08 4.30 99.51 4.34 7.16	47.94 54.76 5.45 31.23 19.53	173.16 63.08 174.39 139.86 80.95	.44 .16 .35 .31	2.15 2.17 2.16 2.16 2.18	2.79 2.80 2.80 2.80 2.81	2.97 2.97 2.97 2.98 2.98	1 1 1 1
270 271 272 273 274 275	5020 2280 199 2100 1400 9490	8550 3550 413 3600 2350 14300	11400 4510 615 4840 3140 17900	15600 5860 952 6700 4310 22800	19100 6960 1270 8330 5320 26700	23100 8160 1660 10200 6470 30900	27500 9450 2120 12200 7760 35400	34000 11300 2880 15400 9730 41800	83.2 25.7 1.0 69.2 31.8 628	30.41 8.74 1.49 21.98 10.79 76.32	35.01 112.47 81.90 15.60 20.61 9.82	165.65 223.96 43.94 72.28 79.15 88.95	.09 .33 .46 .14 .26 .11	 2.17 2.18 2.18 2.18 2.19 2.18 	2.80 2.79 2.80 2.80 2.81 2.81	2.97 2.95 2.98 2.98 2.98 2.98 2.98	1 1 1 1 1 1
276 277 278 ^r 279 280	2370 3130 3020 3040 598	3200 4720 7990 4760 967	3780 5910 13100 6030 1250	4540 7560 22100 7760 1650	5130 8900 30900 9140 1980	5740 10300 41500 10600 2340	6380 11900 54300 12100 2730	7260 14100 74900 14300 3290	41.8 76.5 97.0 79.0 13.0	9.79 13.19 21.78 19.99 8.05	12.37 9.34 96.11 22.73 40.64	65.58 75.36 394.11 295.91 280.73	.43 .44 .20 .20 .20	2.23 2.23 2.19 2.19 2.19	2.83 2.83 2.78 2.78 2.79	2.99 2.99 2.94 2.95 2.95	1 1 1 1
281 282 283 284 285	4860 16900 3070 1070 6990	7620 25900 4900 1680 10200	9700 32500 6290 2150 12700	12600 41800 8230 2820 16100	15000 49300 9820 3380 18900	17600 57200 11500 3990 21900	20300 65800 13400 4660 25200	24300 78000 16000 5640 29900	220 875 60.5 16.4 200	42.35 64.54 20.27 7.85 40.97	15.04 27.19 25.87 50.61 13.36	212.62 223.48 305.26 91.82 212.84	.12 .21 .15 .25 .12	2.23 2.23 2.19 2.19 2.21	2.81 2.81 2.79 2.79 2.80	2.97 2.97 2.96 2.96 2.97	1 1 1 1
286 287 288 289 290	348 104 5060 1160 5860	601 148 8850 1920 9590	805 179 12100 2550 12700	1100 221 17300 3480 17400	1360 254 21900 4290 21500	1640 289 27300 5200 26100	1940 326 33600 6230 31500	2400 378 43500 7790 39700	1.4 2.4 205 21.8 277	2.29 1.98 66.19 8.58 41.20	72.02 159.21 10.53 74.20 19.42	100.64 286.39 314.07 430.02 386.04	.30 .58 .05 .29 .16	2.22 2.10 2.09 2.09 2.09	2.81 2.76 2.76 2.75 2.75	2.97 2.93 2.93 2.92 2.92	1 1 1 1
291 292 293 294 295	168 4150 4680 698 2760	276 6150 7170 1250 4310	363 7610 9060 1710 5460	492 9610 11700 2420 7050	603 11200 13900 3030 8320	726 12900 16300 3730 9680	865 14700 18800 4530 11100	1070 17300 22500 5740 13200	2.1 67.9 103 11.7 40.4	2.32 14.21 19.57 4.85 13.36	219.54 67.39 79.36 205.76 131.60	266.66 407.60 373.79 446.05 502.95	.38 .34 .27 .49 .22	2.07 2.27 2.26 2.26 2.25	2.77 2.80 2.80 2.80 2.80	2.94 2.95 2.95 2.95 2.95 2.95	1 1 1 1

[Q, recurrence interval flood discharge for years indicated; DA, drainage area; L, channel length; CSLOPE, channel slope; BSLOPE, basin slope; SHAPE, basin shape; CF, climate factor for recurrence interval years indicated; REG, region: 1, if site is in Blue Ridge-Piedmont; 2, if site is in Coastal Plain; 3, if site is in Sand Hills; nc, flood-frequency estimates were not computed because the site has less than 10 years of peak-flow record; *, duplicate map identification number for sites having separate periods of regulated or channelized flows; r, site excluded from regional analysis because flows were affected by regulation or channelization; n.a., data not available]

Map identification number (fig. 1)	Q ₂	Q ₅	Q ₁₀	Q ₂₅	Q ₅₀	Q ₁₀₀	Q ₂₀₀	Q ₅₀₀	DA (mi ²)	L (mi)	CSLOPE (ft/mi)	BSLOPE (ft/mi)	SHAPE (DA/L ²)	CF ₂	CF ₂₅	CF ₁₀₀	REG
296	1510	2430	3140	4170	5020	5960	6980	8490	26.8	11.33	37.09	239.70	0.21	2.24	2.80	2.96	1
297	1680	2540	3160	4000	4670	5370	6100	7130	41.4	13.96	44.43	242.70	.22	2.24	2.80	2.96	1
298	556	935	1250	1710	2120	2580	3100	3890	10.9	4.91	338.78	285.36		2.24	2.80	2.96	1
299	6970	10900	14000	18500	22300	26600	31300	38300	296	36.20	3.88	343.47	.23	2.24	2.80	2.95	1
300	479	715	893	1140	1350	1570	1820	2170	14.8	11.44	50.40	263.91	.11	2.23	2.79	2.95	1
301	658	1120	1500	2080	2590	3170	3820	4820	10.0	4.78	179.84	387.36	.45	2.25	2.79	2.95	1
302	2490	3840	4880	6360	7600	8940	10400	12600	66.7	27.47	50.49	440.37	.09	2.23	2.78	2.94	1
303	80.1	109	129	154	174	194	215	243	.6	.88	436.52	199.55	.72	2.24	2.80	2.96	1
304	1450	2420	3200	4360	5360	6480	7730	9630	42.2	13.01	24.24	226.31	.25	2.24	2.80	2.96	1
305	3220	6340	9240	14100	18600	24200	30900	41800	109	15.41	10.25	187.68	.46	2.23	2.79	2.95	1
306	1790	2920	3840	5230	6440	7800	9360	11700	63.1	15.97	28.17	307.71	.25	2.23	2.79	2.95	1
307	10700	15300	18700	23300	26900	30700	34700	40400	676	61.46	2.71	300.03	.18	2.23	2.78	2.94	1
308	2030	3610	4990	7170	9160	11500	14200	18500	79.8	13.83	57.53	415.17	.42	2.22	2.78	2.93	1
309 ^{nc}	n.a.	n.a.	n.a.	n.a.	n.a.	n.a.	n.a.	n.a.	14.5	4.76	503.19	641.91	.66	2.19	2.78	2.94	1
310	1780	3100	4230	5990	7550	9360	11400	14700	23.8	7.84	282.05	600.92	.39	2.19	2.78	2.94	1
310* ^{r,nc}	n.a.	n.a.	n.a.	n.a.	n.a.	n.a.	n.a.	n.a.	23.8	7.84	282.05	600.92	.39	2.19	2.78	2.94	1
311	247	427	575	794	983	1190	1430	1790	5.5	3.89	574.31	591.79	.36	2.19	2.78	2.94	1
312	3020	5220	7130	10100	12800	16000	19700	25600	130	23.03	21.11	426.04	.24	2.21	2.78	2.94	1
312* ^r	2900	4940	6670	9370	11800	14600	17800	22900	130	23.03	21.11	426.04	.24	2.21	2.78	2.94	1
313	14900	22000	27300	34600	40700	47100	54100	64200	945	72.49	2.59	320.73	.18	2.20	2.78	2.94	1
314	2050	3080	3850	4930	5810	6750	7770	9250	79.5	19.44	34.12	403.22	.21	2.21	2.77	2.92	1
315	4080	6670	8740	11800	14300	17200	20300	25000	158	24.08	47.85	440.28	.27	2.21	2.77	2.92	1
316	19800	30400	37900	48000	55900	64100	72600	84500	1332	90.72	4.09	337.83	.16	2.21	2.77	2.92	1
317	626	942	1170	1490	1740	2010	2300	2700	8.0	5.34	142.20	400.62	.27	2.20	2.77	2.92	1
318	3350	5410	7020	9340	11300	13400	15700	19100	126	28.76	72.67	508.25	.15	2.20	2.76	2.92	1
319	23300	37000	48000	64000	77700	92900	110000	135000	1567	106.96	6.30	358.29	.14	2.20	2.77	2.92	1
320	4100	5830	7050	8690	9970	11300	12700	14600	27.6	10.24	259.95	565.57	.26	2.26	2.80	2.95	1
$321^{r,nc}$	n.a.	n.a.	n.a.	n.a.	n.a.	n.a.	n.a.	n.a.	33.5	12.63	213.40	568.89	.21	2.26	2.79	2.95	1
322^{r}	4400	6590	8210	10500	12300	14200	16300	19300	58.4	17.24	172.62	565.25	.19	2.25	2.79	2.95	1
323	4320	6990	9070	12100	14600	17300	20200	24600	51.5	19.40	155.96	569.36	.14	2.25	2.79	2.95	1
324	7690	11800	15000	19500	23300	27400	31800	38400	130	24.00	125.34	533.76	.22	2.22	2.77	2.93	1
325	761	1110	1370	1720	2010	2310	2630	3090	14.4	5.16	556.54	589.19	.54	2.26	2.80	2.95	1
326 ^{nc}	n.a.	n.a.	n.a.	n.a.	n.a.	n.a.	n.a.	n.a.	243	35.84	79.37	478.74	.19	2.21	2.77	2.92	1
327	1930	2710	3250	3960	4510	5070	5660	6470	65.3	17.58	74.41	510.37	.21	2.21	2.77	2.92	1
328	11200	16900	21200	27000	31700	36800	42200	49800	350	42.50	51.32	482.01	.19	2.21	2.77	2.92	1

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Map identification number (fig. 1)	Q ₂	Q ₅	Q ₁₀	Q ₂₅	Q ₅₀	Q ₁₀₀	Q ₂₀₀	Q ₅₀₀	DA (mi ²)	L (mi)	CSLOPE (ft/mi)	BSLOPE (ft/mi)	SHAPE (DA/L ²)	CF ₂	CF ₂₅	CF ₁₀₀	REG
329 330 331 332 333	1910 368 2840 5400 5550	2980 430 4540 9260 9390	3770 468 5880 12500 12600	4850 513 7860 17300 17400	5720 546 9540 21600 21700	6640 578 11400 26400 26500	7620 609 13500 31900 32000	9000 650 16600 40300 40400	49.2 9.2 104 43.3 60.8	10.68 4.90 37.92 14.34 27.05	183.96 127.36 34.80 158.20 50.38	530.96 313.06 410.63 520.68 484.24	0.43 .39 .07 .21 .08	2.21 2.16 2.17 2.18 2.17	2.77 2.77 2.78 2.78 2.78 2.77	2.92 2.94 2.94 2.94 2.94	1 1 1 1
334 335 336 337 338	145 4950 16000 1050 5940	224 9080 26300 2040 10500	283 12800 34600 2940 14300	367 18800 47000 4390 20200	436 24300 57600 5730 25400	510 30900 69600 7330 31400	590 38600 83100 9200 38300	707 51100 104000 12200 48900	1.6 157 608 23.1 92.1	3.08 39.09 73.77 9.70 19.38	519.11 34.98 23.71 118.81 84.54	405.77 531.07 452.57 398.27 376.69	.17 .10 .11 .25 .24	2.17 2.17 2.17 2.15 2.15	2.77 2.77 2.77 2.77 2.77	2.93 2.93 2.93 2.94 2.94	1 1 1 1 1
339 340 341 342 343 ^r	15.7 2260 3340 1960 975	23.5 3750 5170 2870 1600	29.2 4970 6580 3530 2160	37 6830 8580 4420 3070	43.3 8450 10200 5140 3920	50 10300 12000 5890 4950	57.1 12400 14000 6690 6200	67.3 15600 16900 7830 8250	.5 42.0 140 57.1 14.9	1.72 11.66 25.26 16.11 7.41	997.21 104.61 6.86 31.33 98.65	415.55 406.56 375.26 436.73 246.33	.18 .31 .23 .22 .26	2.16 2.16 2.28 2.28 2.28	2.77 2.77 2.80 2.80 2.81	2.94 2.93 2.95 2.95 2.95	1 1 1 1
344 345 346 347 348	2900 124 6160 9490 2500	4580 256 8870 13500 3450	5930 375 10800 16300 4090	7900 568 13500 20000 4940	9590 746 15700 22800 5590	11500 954 18000 25700 6250	13600 1200 20400 28600 6940	16700 1580 23900 32700 7880	86.5 1.6 323 436 51.9	20.22 2.64 34.17 52.12 16.68	123.81 491.80 5.76 6.48 63.83	435.46 467.80 394.79 408.05 456.46	.21 .25 .28 .16 .19	2.28 2.27 2.24 2.27 2.28	2.80 2.80 2.80 2.80 2.80 2.80	2.95 2.95 2.96 2.94 2.95	1 1 1 1
349 ^r 350 351 352 ^{nc} 352 ^{*r}	2800 5120 13600 n.a. 3990	4370 7330 21400 n.a. 6940	5530 8910 27400 n.a. 9190	7120 11000 36000 n.a. 12300	8390 12700 43100 n.a. 14800	9740 14500 50800 n.a. 17500	11200 16400 59200 n.a. 20300	13200 19000 71500 n.a. 24300	144 174 664 143 143	38.21 49.80 63.31 21.61 21.61	40.37 42.39 8.43 112.60 112.60	464.09 475.41 430.10 411.88 411.88	.10 .07 .17 .31 .31	2.27 2.27 2.27 2.27 2.27 2.27	2.80 2.79 2.79 2.80 2.80	2.95 2.94 2.94 2.95 2.95	1 1 1 1
353 354 354* ^r 355 356	1470 6780 7470 5370 8720	2070 9790 11800 7370 11500	2490 12100 15000 8720 13300	3050 15200 19300 10500 15600	3480 17800 22700 11800 17400	3930 20700 26200 13200 19100	4400 23700 29900 14600 20900	5040 28200 35000 16500 23300	51.0 347 347 131 184	15.75 37.92 37.92 22.52 26.94	167.42 56.29 56.29 106.56 85.10	520.46 464.22 464.22 613.81 589.97	.21 .24 .24 .26 .25	2.27 2.27 2.27 2.27 2.27 2.27	2.80 2.80 2.80 2.80 2.80 2.80	2.95 2.95 2.95 2.94 2.94	1 1 1 1
357 357* ^r 358 359 360	16900 16600 19.7 944 2310	26600 23100 36.8 1290 3640	33800 27200 51.8 1530 4660	43800 32200 75.5 1840 6120	51700 35800 97 2090 7330	60200 39300 122 2340 8640	69200 42700 151 2590 10100	81900 47200 198 2950 12200	655 655 .5 13.8 44.4	57.17 57.17 1.78 8.53 12.93	37.57 37.57 595.52 346.36 173.28	502.95 502.95 474.30 601.47 606.86	.20 .20 .20 .19 .26	2.27 2.27 2.27 2.27 2.27 2.26	2.79 2.79 2.80 2.79 2.79	2.94 2.94 2.94 2.94 2.93	1 1 1 1

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Map identification number (fig. 1)	Q ₂	Q ₅	Q ₁₀	Q ₂₅	Q ₅₀	Q ₁₀₀	Q ₂₀₀	Q ₅₀₀	DA (mi ²)	L (mi)	CSLOPE (ft/mi)	BSLOPE (ft/mi)	SHAPE (DA/L ²)	CF2	CF ₂₅	CF ₁₀₀	REG
361 ^{nc}	n.a.	n.a.	n.a.	n.a.	n.a.	n.a.	n.a.	n.a.	1571	85.95	21.06	472.39	0.21	2.28	2.79	2.94	1
361* ^r	10800	17500	22400	29100	34500	40100	46100	54400	1571	85.95	21.06	472.39	.21	2.28	2.79	2.94	1
362	2920	4140	5020	6200	7140	8130	9180	10700	42.0	19.16	86.51	480.21	.11	2.29	2.80	2.95	1
363	1440	2410	3200	4380	5400	6550	7830	9780	37.6	7.91	150.30	474.21	.61	2.29	2.81	2.96	1
364 ^r	1640	2090	2470	3070	3600	4220	4940	6070	190	26.23	50.04	384.09	.28	2.29	2.81	2.96	1
365	11300	15800	18800	22800	25800	28900	32100	36500	406	47.41	15.70	368.28	.18	2.31	2.81	2.96	1
365* ^r	8960	12600	14800	17400	19200	20800	22400	24300	406	47.41	15.70	368.28	.18	2.31	2.81	2.96	1
366	4190	6510	8320	10900	13000	15400	18000	21800	104	22.63	50.23	435.97	.20	2.31	2.81	2.95	1

APPENDIX

The value of the mean square error (MSE_s) at a specific site can be estimated as follows: Denote the column vector of *n* logarithms of observed peak-discharge characteristics at *n* sites in a region by *Y*. For example,

in which, $Q_{50,i}$, represents the observed 50-year peak at the *i*th gaging station in the region. Further, let X represent a (*n* by *p*) matrix of *p*-1 basin characteristics augmented by a column of ones at *n* gaging stations and *B* represent a column vector of *p* regression coefficients.

For example,

$$X = \begin{bmatrix} 1 & \log(DA_1) & \log(IA_1) & \log(RQ50_1) \\ 1 & \log(DA_2) & \log(IA_2) & \log(RQ50_2) \\ & & & & \\ 1 & \log(DA_n) & \log(IA_n) & \log(RQ50_n) \end{bmatrix} \text{ and } B = \begin{bmatrix} a \\ b_1 \\ b_2 \\ b_4 \end{bmatrix}.$$

The linear model can be written as

$$Y = XB$$
.

The mean square sampling error, $MSE_{s,0}$, for an ungaged site with basin characteristics given by the row vector $x_0=[1 \log (DA_0) \log (IA_0) \log (RQ50_0)]$, for example, is calculated as

$$MSE_{s,0} = x_0 \{X^T \Lambda^{-1} X\}^{-1} x_0^T$$

in which Λ is the (*n* by *n*) covariance matrix associated with *Y*. The diagonal elements of Λ are model error variance, γ^2 , plus the time-sampling error for each site i (i=1,2,3,...n), which is estimated as a function of a regional estimate of the standard deviation of annual peaks at site *i*, the recurrence interval of the dependent variable and the number of years of record at site *i*. The off-diagonal elements of Λ are the sample covariance of the estimated *t*-year peaks at sites *i* and *j*. These offdiagonal elements are estimated as a function of a regional estimate of the standard deviation of annual peaks at sites *i* and *j*, the recurrence interval of the dependent variable and the number of concurrent years of record at sites *i* and *j* (Tasker and Stedinger, 1989). The (*p* by *p*) matrix $\{X^T \Lambda^{-1} X\}^{-1}$ for each equation is given in Appendix table 1. The mean square error of a prediction, in log (base 10) units, at specific ungaged sites can be estimated as

$$MSE_{p,0} = (\gamma^2 + MSE_{s,0}).$$

The standard error of a prediction, $SE_{\text{prediction}}$, in percent, can be calculated as

$$SE_{\text{prediction}} = 100 \{ e^{5.302 \times (MSEp, 0)} - 1 \}^{0.5}$$

Appendix Table 1. Matrix ${\{X}^T\Lambda^{-1}X\}^{-1}$ for the equations in table 5 (p. 11)

[These matrices can be used to compute the standard error of prediction and prediction intervals as explained in the text. Numbers are given in scientific notation, for example, $0.43958E-01 = 0.43958 \times 10^{-1} = 0.043958$]

		Hydrol	ogic area		
Blue Ridge	-Piedmont	Coasta	ıl Plain	San	1 Hills
		2-year recu	irrence interval		
0.14072E-02	-0.49612E-03	0.31893E-02	-0.98276E-03	0.10200E-01	-0.44648E-02
-0.49612E-03	0.24350E-03	-0.98276E-03	0.43777E-03	-0.44648E-02	0.25105E-02
		5-year recu	irrence interval		
0.16431E-02	-0.55517E-03	0.37147E-02	-0.10840E-02	0.12971E-01	-0.56042E-02
-0.55517E-03	0.26322E-03	-0.10840E-02	0.45031E-03	-0.56042E-02	0.31270E-02
		10-year rec	urrence interval		
0.18985E-02	-0.62424E-03	0.43917E-02	-0.12495E-02	0.15456E-01	-0.65943E-02
-0.62424E-03	0.28912E-03	-0.12495E-02	0.50084E-03	-0.65943E-02	0.36522E-02
		25-year rec	urrence interval		
0.22833E-02	-0.73140E-03	0.54496E-02	-0.15204E-02	0.19118E-01	-0.80470E-02
-0.73140E-03	0.33094E-03	-0.15204E-02	0.59196E-03	-0.80470E-02	0.44209E-02
		50-year rec	urrence interval		
0.25999E-02	-0.82124E-03	0.63333E-02	-0.17517E-02	0.22136E-01	-0.92451E-02
-0.82124E-03	0.36687E-03	-0.17517E-02	0.67303E-03	-0.92451E-02	0.50554E-02
		100-year rec	currence interval		
0.29342E-02	-0.91725E-03	0.72726E-02	-0.20005E-02	0.25348E-01	-0.10523E-01
-0.91725E-03	0.40581E-03	-0.20005E-02	0.76200E-03	-0.10523E-01	0.57330E-02
		200-year rec	currence interval		
0.32839E-02	-0.10186E-02	0.82596E-02	-0.22640E-02	0.28740E-01	-0.11875E-01
-0.10186E-02	0.44737E-03	-0.22640E-02	0.85763E-03	-0.11875E-01	0.64516E-02
		500-year rec	currence interval		
0.37671E-02	-0.11600E-02	0.96272E-02	-0.26319E-02	0.33482E-01	-0.13772E-01
-0.11600E-02	0.50586E-03	-0.26319E-02	0.99272E-03	-0.13772E-01	0.74616E-02