

DESCRIPTION OF MAP UNITS SURFICIAL DEPOSITS

Af Artificial fill (Holocene)—Unconsolidated soil, sand, gravel, and crushed rock used for highway and railroad beds and levees

Qac Alluvium of Columbia River floodplain and channel (Holocene and Pleistocene)—Unconsolidated fine sand, silt, and clay that underlies historic floodplain, islands, and bars of the Columbia River at elevations less than 30 ft (10 m), composed largely of quartz, feldspar, and conspicuous muscovite. Sections exposed in banks, excavations, and dredge samples commonly show sand-silt couplets up to 30 cm thick; basal sands locally display ripple cross lamination, planar lamination, and tabular forests, capping silt and clay beds generally massive and bioturbated; couples interpreted as deposits of large annual Columbia River freshets before substantial 20th century river regulation and floodplain dikeing. Unit includes local diatomaceous beds and layers of organic-rich sediment that probably formed in floodplain marshes, ponds, and lakes, as well as thin (less than 2 cm) tephra beds from Holocene eruptions of Mount St. Helens. Regionally, most deposits above modern low-water river level (3 ft (1 m) above sea level in map area) are younger than 2000 years old and most deposits forming Government Island are younger than 500 yr B.P. (table 3). Well logs and seismic reflection profiles show that fine-grained valley fill beneath the historic floodplain, presumably representing river aggradation since the last-glacial sea level low stand of about 15,000 yr B.P. (Baker, 2002), locally extends to 300 ft (90 m) below sea level in the map area (Hoffmeister, 1984; Hartford and McFarland, 1989; Gates, 1994; Pratt and others, 2001; Rapp, 2005). Near Blue Lake, deposits contain locally concentrated ⁷Be, ²¹⁰Pb, ¹³⁷Cs, and ¹³⁴Cs (Gates, 1994) at 45 ft (14 m) below sea level. Near tributary mouths, unit includes rare lenses of cobble to pebble gravel composed largely of andesite and dacite clasts; also includes local unmappped areas of dredge spoils and other fills, and possibly colluvial deposits

Qtr Terrace deposits of lower Washougal River (Holocene and/or Pleistocene)—Unconsolidated sand, gravel, and organic-rich mud along rivers and creeks, in seasonally inundated depressions, and along shore of Lacamas Lake. Deposits along the Washougal and Sandy Rivers consist mostly of stratified sand and cobble to pebble gravel; deposits along smaller watercourses are more variable

Qsr Terrace deposits of Sandy River (Holocene and/or Pleistocene)—Unsorted accumulations of angular basalt blocks and scoria below cliffs on south slope of Prune Hill, includes deposits of both natural and anthropogenic origins

Qtl Terrace deposits of Little Washougal River (Holocene and/or Pleistocene)—Diamictic of unsorted, angular bedrock and surficial material transported down slope on mass. Chiefly deep-seated, semi-oholitic slumps and internally disrupted rockslide, earthflow, and debris-flow deposits. Many mapped slides head at arcuate scars and exhibit subhorizontal spurs, bulbous toes, and hummocky, poorly drained surfaces. Slides at Camas possibly triggered by late Pleistocene Camasia flooding

Qlv Levee deposits (Pleistocene)—Massive unconsolidated deposits of light gray to buff, micaceous, quartzofeldspathic silt and fine sand, commonly contains isolated granules and small pebbles, generally capped with strongly developed red soils. Forms widespread mantle on uplands of map area but mapped only where thick (about 3 to 25 m) and extensive enough to ob-

Basin-fill deposits (Pleistocene)—Sediment deposited by colossal glacier-outburst floods caused by repeated failure of ice dam across Clark Fork River that formed Pleistocene Lake Missoula in western Montana (Breit, 1925, 1959; Bretz and others, 1956; Trimble, 1963; Allison, 1978; Baker and Bunker, 1985; Watt, 1985, 1994, 1996; Atwater, 1986; O'Connor and Baker, 1992; Benito and O'Connor, 2003). The Missoula floods achieved stages of 400–500 ft (120–150 m) as they spread and slowed over the eastern Portland Basin after exiting the western Columbia River Gorge with velocities of 35 m/s at peak discharge (Benito and O'Connor, 2003) and deposited coarse traction load in series of large bars and plains, the Portland delta of Breitz (1925). Hydraulically dammed floodwaters temporarily ponded in Portland Basin and deposited suspended sediment load (Trimble, 1963). Radiocarbon and tephrochronologic data from outside the map area indicate depositional ages between about 17,000 and 13,000 ¹⁴C years B.P. (Watt, 1985, 1994; Atwater, 1986; Benito and O'Connor, 2003; Clague and others, 2003). Coarse bedload deposits and fine sand-water deposits mapped separately

Basal fill (Pleistocene)—Unconsolidated light brown to light gray silt, clay, and fine to medium sand. Up to 20 m thick at low elevations in the region but thins toward upper mapable extent at 300–350 ft (90–105 m) elevation in the map area. Upper map limit only approximately placed on basis of topography owing to the difficulty in distinguishing between similar loess and clayey soils that cover most upland surfaces. Most exposures obscure, but rare fresh exposures show multiple, 0.25–to-1.5-m thick, fine-up-up couplets of ripple cross-stratified very fine sand grading up to massive bioturbated clayey silt. Sand composed of quartz, feldspar, and conspicuous muscovite, indicative of a Columbia River provenance. Coarser sand facies contain abundant dark volcanic rock clasts. Interpreted as slack-water sediment and quartzite from temporarily ponded floodwater. May locally include compositionally identical loess

Gravel facies (Pleistocene)—Unconsolidated, gray, stratified, bouldery to cobbly gravel and sand, fine-up-up couplets over older basin-fill deposits on upland areas flanking the Columbia River; organized into prominent large bar and channel complexes. Clasts are in the northwest part of map area and much of east Portland, Fairview, and Troutdale south of the river (Allison, 1978; Minervini and others, 2003). Beneath Holocene floodplain deposits (Qac) northeast of Blue Lake, flood gravel occupies buried late glacial channel floor is at least 240 ft (70 m) below sea level (Hartford and McFarland, 1989; Pratt and others, 2001). Locally, as along southern edge of the historic Columbia River floodplain and the Lacamas Lake trough, Missoula-flood deposits form (<2 m) and discontinuous mantle on older basin-fill and Pleistocene terrace deposits. Clasts are less than 2 m in diameter, and consist of volcanic rocks, locally, such as southwest of Prune Hill, consists of large, angular to subangular basaltic clasts eroded from nearby well-sorted, locally imbricated, and texture varies from compact with sand matrix to loose open-work; clasts subangular to well rounded, include common large (up to 1 m diameter) boulders of hornblende andesite; deposit contains at least one 1-m-thick bed of silt-rich sand inferred to be the distal facies of a Mount Hood lake. Exhibits weakly developed soil profile and weathering rinds on fine-grained volcanic clasts less than 1 mm thick, suggesting relative youth; absence of cover by catclastic (Missoula)-flood deposits indicate that unit postdates flooding; may reflect deposition behind the immense Missoula flood bar extending southwestward from Broughton Hill that temporarily blocked the Sandy River at its confluence with the Columbia River. Equivalent in part to Estacada Formation of Trimble (1963)

Waters Hill Formation (Pleistocene and/or Pleistocene)—Semi-consolidated, deeply weathered, poorly exposed fluvial gravel forming Grant Butte. Consists of well rounded to subrounded cobbles and pebbles of volcanic rocks, predominantly porphyritic andesite derived from Cascade Range to the east; matrix is coarse to fine volcaniclastic sand. Age unknown; topographic position suggests unit is younger than Troutdale Formation (Tm) and older than other gravel units (Qtr, Qg, Qg, Qs) in map area

Basaltic andesite of Prune Hill (Pleistocene)—Light to medium-gray, microwhiskered, olivine-phyric, calc-alkaline basaltic andesite (54–55 wt percent SiO₂) that underlies area directly west of Prune Hill; well developed platy and columnar jointing; up to 65 m thick in Fisher Quarry, where complex jointing patterns and contact with scoriae flow breccia indicate multiple flow lobes. Contains phenocrysts of olivine (3–7 percent; about 0.5 mm, locally to 10 m across); with minute chromian spinel inclusions; plagioclase (0.1 percent, 0.5–1 mm long), and augite (0.1 percent) in trachytic to intergranular, locally microvesicular groundmass of plagioclase, pyroxene, Fe-Ti oxide, and dark-brown interstitial glass; distinguished by presence of prismatic hypersthene microclasts. Some samples contain corroded and sutured plagioclase xenocrysts about 1 mm across. Vent located east of Fisher Quarry in area underlain by scoria deposits (striped pattern) and consistent with a strong positive aeromagnetic anomaly (Snyder and others, 1993). Well sorted to poorly sorted scoria beds consist of black to brick-red, variably vesicular clasts as large as 1 m across that are petrographically and chemically similar to associated lava flows; also commonly contain well rounded pebbles and cobbles of Columbia River Basalt Group and quartzite derived from basin-fill sediments. Normal magnetic polarity (I.T. Hagstrum, written commun., 2000). ⁴⁰Ar/³⁹Ar age of 596±47 ka (table 2) is indistinguishable from K-Ar age of 590±50 ka reported by Conway and others (1998a)

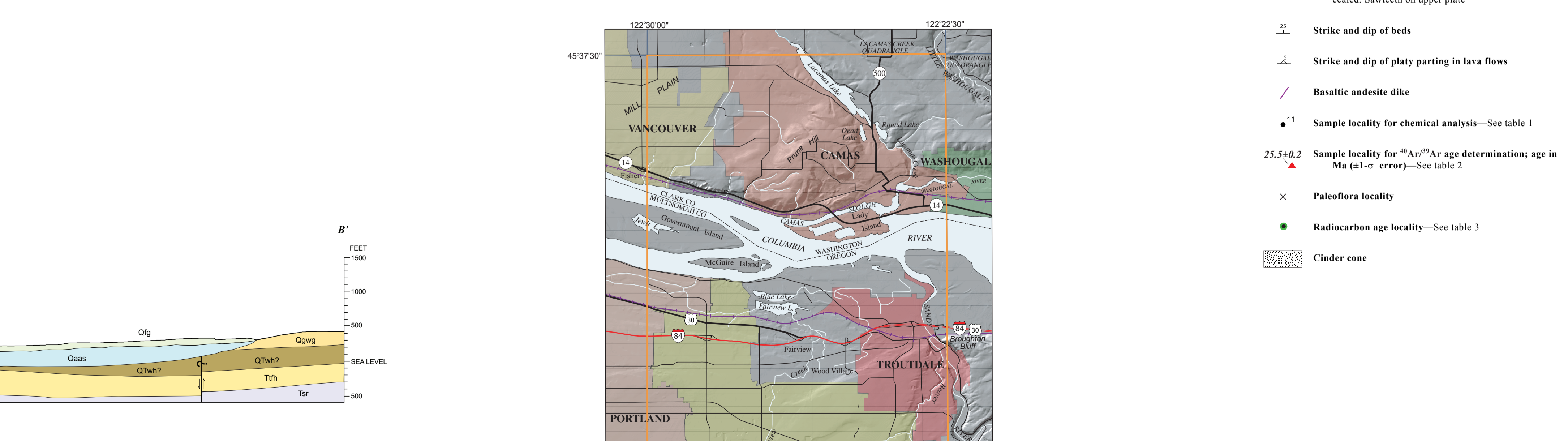
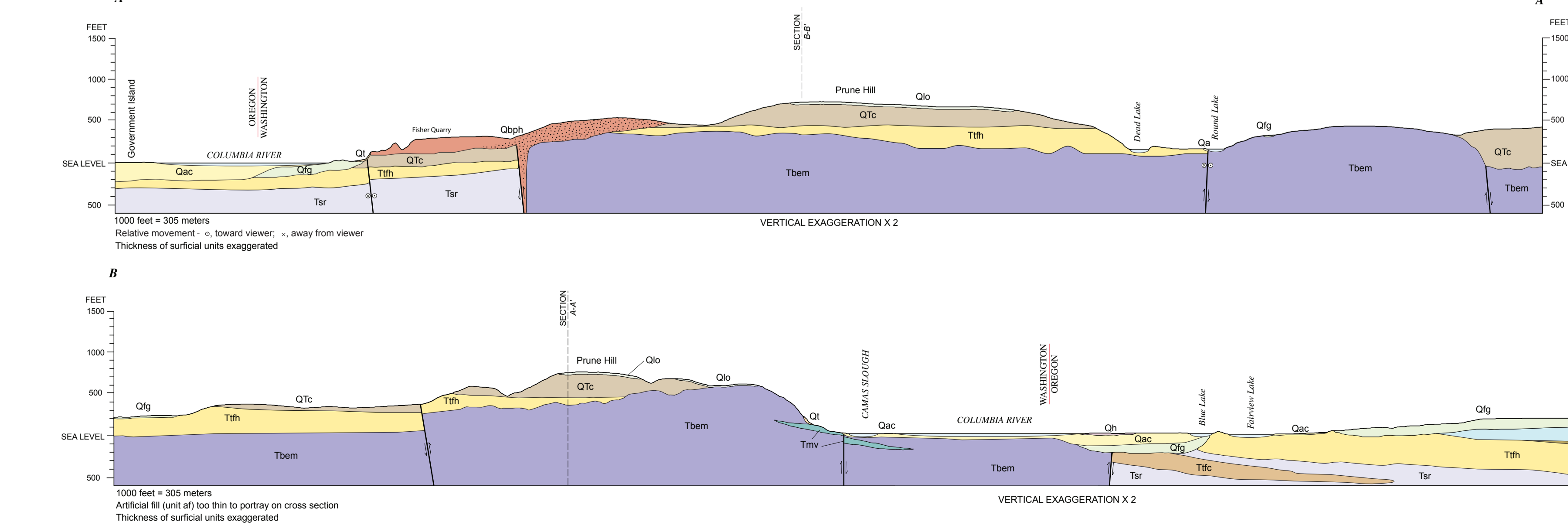
Basaltic andesite of Broughton Hill (Pleistocene)—Light-gray, olivine-phyric, calc-alkaline basaltic andesite flow (52–53 wt percent SiO₂) (R.C. Everts and R.M. Conroy, unpub. data). Flow caps Broughton Hill and has well developed columnar jointing; probably erupted from vent at Chamberlain Hill about 3 km to east. Contains phenocrysts and microphenocrysts of olivine (about 9 percent; mostly 0.5 to 1 mm but a few as large as 3 mm across); with inclusions of chromian spinel; locally replaced by iddingsite) in a trachytic to subophitic groundmass of plagioclase, augite, and Fe-Ti oxide. Reversed magnetic polarity (I.T. Hagstrum, written commun., 2002). An ⁴⁰Ar/³⁹Ar age of 1,282±0.014 Ma was obtained from a sample collected just east of the map area (R.J. Fleck, written commun., 2008)

Basaltic andesite of Elkhorn Mountain (Oligocene)—Sequence of lava flows and flow breccia composed of dark-gray to brown, porphyritic to sericitic olivine-basaltic andesite and Fe-Ti oxide in an intergranular groundmass of plagioclase, Fe-Ti oxide, and minor interstitial brown smectite. Age unknown but overlies basal andesite of Elkhorn Mountain (Tm)

Dacite (Oligocene)—Sparsely plagioclase-phyric dacite flow and flow breccia exposed at east end of Lady Island; exhibits pronounced platy parting with widely variable orientation. Contains phenocrysts of sericitic plagioclase (18 percent; 1–2 mm long) and sparse microphenocrysts of orthopyroxene and Fe-Ti oxide in an intergranular groundmass of plagioclase, Fe-Ti oxide, and minor interstitial brown smectite. Age unknown but overlies basal andesite of Elkhorn Mountain (Tm)

Basaltic andesite of Elkhorn Mountain (Oligocene)—Sequence of lava flows and flow breccia composed of dark-gray to brown, porphyritic to sericitic olivine-basaltic andesite and Fe-Ti oxide in an intergranular groundmass of plagioclase, augite, Fe-Ti oxide, and minor to abundant interstitial glass (largely devitrified or replaced by smectite, quartz, or calcite); groundmass textures chiefly intergranular to trachytic, less commonly subophitic or microphyric. All flows in unit are diolitic, and many are exceptionally rich in Fe (Fe⁰ as high as 13.3 wt percent and poor in Sr (<30 ppm) compared to mafic rocks elsewhere in southern Washington Cascade Range (de Bray and others, 2006; R.C. Everts, unpub. data). ⁴⁰Ar/³⁹Ar age of 23,510±2 Ma obtained for sample near top of unit north of Camas Slough

Volcaniclastic sedimentary rocks (Oligocene)—Section of weathered, poorly exposed volcaniclastic rocks interbedded with basaltic andesite flows (Tm) east of Lacamas Creek; approximately 30 m thick. Primary matrix of olivine-basaltic andesite and conglomerate that contains olivine-bearing basaltic clasts; overlies and intertongues with micaceous, arkosic sandstone, siltstone, and claystone of the Sandy River Mudstone (Tr). Well exposed in valley of Lacamas Creek, on west shore of Lacamas Lake, and along Sandy River east of Troutdale. Hydraulic sandstone equivalent to vitreous sandstone of Trimble (1963) and Tolan and Beeson (1984) consists largely of entirely of angular to subrounded fragments, 2 to 6 mm across, of black, generally nonvesicular basalt that contains phenocrysts of olivine (0.5–2 mm) and plagioclase (1–3 mm) in a glassy (sideromelane) to intergranular groundmass; sideromelane partly to completely altered to palagonite, which cements sandstone and imparts a distinctive yellowish-brown color to the originally dark-green rock; many beds contain minor admixed nonvolcanic debris such as quartz, muscovite, hornblende, and potassium feldspar. Sandstone ranges from poorly sorted to well sorted; typically thick-bedded, commonly lenticular, crossbedded, and contains dispersed pebbles and cobbles of olivine-bearing basalt. Interbedded conglomerate mostly well-sorted and clast supported, consists of well rounded to subrounded pebbles and cobbles of variably vesicular, olivine-plagioclase-phyric basalt, basaltic rocks of Columbia River Basalt Group, and generally mi-



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