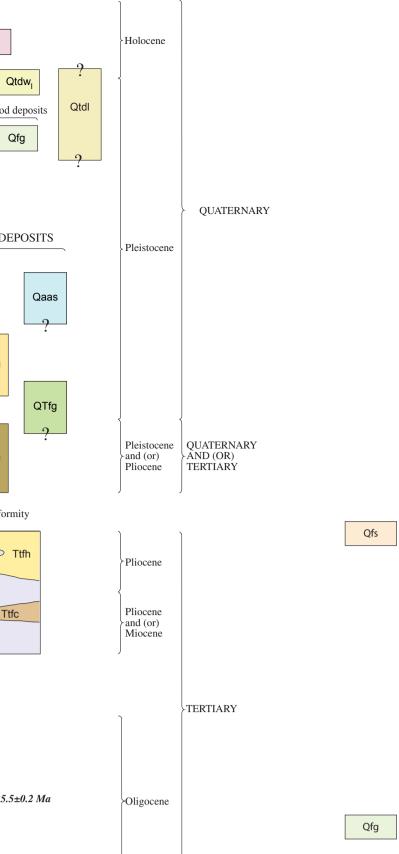


Geologic Map of the Camas Quadrangle, Clark County, Washington, and Multnomah County, Oregon Russell C. Evarts and Jim E. O'Connor



fine upwards; gray except for fine sand beds, which are locally reddish to pinkish gray. Gravel clasts are dominantly fresh, porphyritic, pyroxene and hornblende andesites; sand composed mostly of angular to sub-rounded, gray to red, lithic (andesite) fragments, 5–20 percent clear plagioclase fragments, and 2–10 percent equant hornblende crystals. Diamictons are brown to brownish gray, massive to weakly stratified beds, 0.5 to 5 m thick, composed of 30-50 percent angular andesite fragments up to 10 cm diameter in matrix of sand, silt, and clay (Rapp, 2005). Stratified sand and gravel interpreted to represent fluvial aggradation triggered by highly increased sediment supply following Mount Hood eruptions. Diamictons are interpreted as deposits of Mount Hood lahars. Numerous radiocarbon ages (table 3) and regional stratigraphic relations show that these deposits are younger than 1600 yr B.P. and were produced during and after the Timberline and Old Maid eruptive episodes of Crandell (1980); most are older than 1200 yr B.P. and are related to the Timberline period (Rapp, 2005). Old Maid-age sediment forms several channel-fill deposits, as thick as 5 m, in the Sandy River delta and aggraded in the lower Sandy River valley to elevations as high as or higher than the Timberline-age deposits (Rapp, 2005) Terrace deposits of Sandy River (Holocene and (or)

Pleistocene)—Poorly exposed, unconsolidated, sandy cobble to boulder gravel forming distinct bench with surface elevation between 170 and 190 ft (48 and 58 m) along Sandy River and Beaver Creek. Consist of subhorizontal gravel sheets, 0.5–5 m thick, locally separated by thin sand lenses; poorly sorted, locally imbricated, and texture varies from compact with sand matrix to loose open-work; clasts subangular to well rounded, include common large (up to 1 m diameter) boulders of hornblende andesite; deposit contains at least one 1-m-thick bed of silty lithic-rich sand inferred to be the distal facies of a Mount Hood lahar. Exhibits weakly developed soil profile and weathering rinds on fine-grain volcanic clasts less than 1 mm thick, suggesting relative youth; absence of cover by cataclysmic (Missoula)flood deposits indicate that unit postdates flooding; may reflect deposition behind the immense Missoulaflood bar extending southwestward from Broughton Bluff that temporarily blocked the Sandy River at its confluence with the Columbia River. Equivalent in part to Estacada Formation of Trimble (1963) Terrace deposits of lower Washougal River (Holocene and (or) Pleistocene)—Unconsolidated sandy gravel

and sand underlying small terraces along Washougal River; generally less than 10 m thick. Gravel is poorly sorted, locally imbricated, and texture varies from compact with sand matrix to loose openwork. Clasts subangular to well rounded, derived from Tertiary volcanic and granitic rocks of Cascade Range, Columbia River Basalt Group, and Troutdale Formation. Deposits are below 50 ft (15 m) elevation and minimally weathered; inferred to represent aggradation behind coarse-grained Missoula Flood bar that now constricts mouth of Washougal River Terrace deposits of Little Washougal River (Holocene and (or) Pleistocene)—Unconsolidated gravel and

sand flanking Little Washougal River near northeastern corner of map area; about 2–5 m thick. Gravel is poorly sorted, composed of subangular to well rounded clasts derived from Tertiary volcanic and granitic rocks of Cascade Range, Columbia River Basalt Group, and Troutdale Formation. Minimal soil development. Formed by aggradation behind downriver landslide complex in adjacent Washougal quadrangle (R.C. Evarts and J.E. O'Connor, unpub. mapping) Loess (Pleistocene)—Massive unconsolidated deposits of light-gray to buff, micaceous, guartzofeldspathic eolian silt and fine sand; commonly contains isolated granules and small pebbles; generally capped with strongly developed red soils. Forms widespread mantle on uplands of map area but mapped only where

scure underlying units. Overlies 596-ka basaltic andesite of Prune Hill, but probably deposited during several episodes throughout late Quaternary time. Below about 300 ft (90 m) elevation, may include slack-water cataclysmic-flood deposits (Qfs) Cataclysmic-flood deposits (Pleistocene)—Sediment deposited by colossal glacier-outburst floods caused by repeated failure of ice dam across Clark Fork River that formed Pleistocene Lake Missoula in western Montana (Bretz, 1925, 1959; Bretz and others, 1956; Trimble, 1963; Allison, 1978; Baker and Bunker, 1985; Waitt, 1985, 1994, 1996; Atwater, 1986; O'Connor and Baker, 1992; Benito and O'Connor, 2003). The Missoula floods achieved stages of 400-500 ft (120-150 m) as they spread and slowed over the eastern Portland Basin after exiting the western Columbia River Gorge with velocities of 35 m/s at peak discharge (Benito and O'Connor, 2003) and deposited coarse traction load in series of large bars and plains, the Portland delta of Bretz (1925). Hydraulically dammed floodwaters temporarily ponded in Portland Basin and deposited suspended sediment load (Trimble, 1963). Radiocarbon and tephrochronologic data from outside the map area indicate depositional ages between about 17,000 and 13,000 <sup>14</sup>C years B.P. (Waitt, 1985, 1994; Atwater, 1986; Benito and O'Connor, 2003; Clague and others, 2003). Coarse bedload deposits and fine slackwater deposits mapped separately

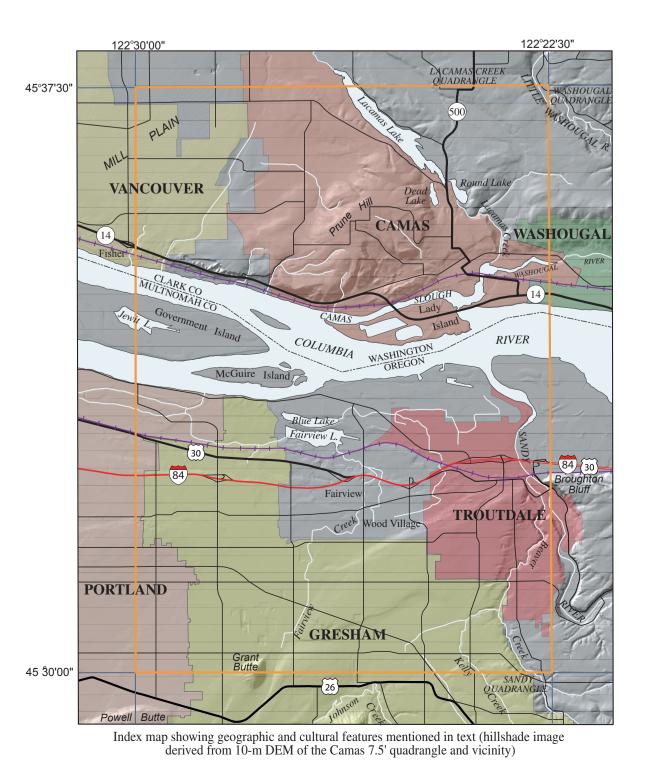
Sand and silt facies—Unconsolidated light brown to light gray silt, clay, and fine to medium sand. Up to 20 m thick at low elevations in the region but thins toward upper mappable extent at 300–350 ft (90–105 m) elevation in the map area. Upper map limit only approximately placed on basis of topography owing to the difficulty in distinguishing from similar loess and clayey soils that cover most upland surfaces. Most exposures obscure, but rare fresh exposures show multiple, 0.25-to-1.5-m thick, fining-up sequences of ripple cross-stratified very fine sand grading up to massive bioturbated clayey silt. Sand composed of quartz, feldspar, and conspicuous muscovite, indicative of a Columbia River provenance. Coarser sand facies contain abundant dark volcanic rock fragments. Interpreted as slack-water sediment settled from temporarily ponded floodwater. May locally include compositionally identical loess

Gravel facies—Unconsolidated, gray, stratified, bouldery to cobbly gravel and sand deposited in thick sheets over older basin-fill deposits on upland areas flanking the Columbia River; organized into prominent large bar and channel complexes on Mill Plain in the northwest part of map area and much of east Portland, Fairview, and Troutdale south of the river (Allison, 1978; Minervini and others, 2003). Beneath Holocene floodplain deposits (Qac) northeast of Blue Lake, flood gravel occupies buried late glacial channel; channel floor is at least 230 ft (70 m) below sea level (Hartford and McFarland, 1989; Pratt and others, 2001). Locally, as along southern edge of the historic Columbia River floodplain and in the Lacamas Lake trough, Missoula-flood deposits form thin ( $\leq 2$ m) and discontinuous mantle on older basin-fill and bedrock units. Numerous active and historic quarry exposures in upland areas reveal more than 15 m of crudely stratified and poorly sorted coarse gravel and sand, commonly deposited in tall (locally >5 m), steeply dipping (up to 35°) foresets. Deposits generally fine westward and away from the Columbia River; in some places, contain immense boulders, some exceeding 4 m in long diameter, particularly near Troutdale and along southwest side of Prune Hill. Texturally and compositionally variable; most clasts derived from Columbia River Basalt Group but includes blocks of Troutdale Formation and Pliocene-Quaternary Cascadian basalt, all probably entrained in the western Columbia River Gorge; locally, such as southwest of Prune Hill, consists of large, angular-to-subangular blocks eroded from nearby valley walls. Sweeping unconformities visible in large fresh exposures likely reflect multiple flood events (Benito and O'Connor, 2003). Unit includes local sand accumulations below 200 ft (60 m) elevation, some of them mapped as sand and silt deposits (Qs) by Trimble (1963), that were probably deposited by smaller late-episode floods confined to the narrow Columbia River valley of late Pleistocene time

**VOLCANIC ROCKS OF THE BORING VOLCANIC** FIELD Basaltic andesite of Prune Hill (Pleistocene)—Light- to

medium-gray, microvesicular, olivine-phyric, calc-alkaline basaltic and esite  $(54-55 \text{ wt percent SiO}_2)$  that underlies area directly west of Prune Hill; well developed platy and columnar jointing; up to 65 m thick in Fisher Quarry, where complex jointing patterns and contacts with scoriaceous flow breccia indicate multiple flow lobes. Contains phenocrysts of olivine (3–7 percent; about 0.5 mm, locally to 1 mm across; with minute chromian spinel inclusions), plagioclase (0.1 percent; 0.5–1 mm long), and augite (0.1 percent) in trachytic to intergranular, locally microvesicular groundmass of plagioclase, pyroxene, Fe-Ti oxide, and dark-brown interstitial glass; distinguished by presence of prismatic hypersthene microlites. Some samples contain corroded and sieved plagioclase xenocrysts about 1 mm across. Vent located east of Fisher Quarry in area underlain by scoria deposits (stipple pattern) and coincident with a strong positive aeromagnetic anomaly (Snyder and others, 1993). Well sorted to poorly sorted scoria beds consist of black to brick-red, variably vesicular clasts as large as 1 m across that are petrographically and chemically similar to associated lava flows; also commonly contain well-rounded pebbles and cobbles of Columbia River Basalt Group and quartzite derived from basin-fill sediments. Normal magnetic polarity (J.T. Hagstrum, written commun., 2000).  $^{40}$ Ar/ $^{39}$ Ar age of 596±47 ka (table 2) is indistinguishable from K-Ar age of 590±50 ka reported by Conrey and others (1996a)

Qbbb Basaltic andesite of Broughton Bluff (Pleisto**cene**)—Light-gray, olivine-phyric, calc-alkaline basaltic andesite flow  $(52-53 \text{ wt percent SiO}_2)$  (R.C. Evarts and R.M. Conrey, unpub. data); flow caps Broughton Bluff and has well developed columnar jointing; probably erupted from vent at Chamberlain Hill about 3 km to east. Contains phenocrysts and microphenocrysts of olivine (about 9 percent; mostly 0.5 to 1 mm but a few as large as 3 mm across; with inclusions of chromian spinel; locally replaced by iddingsite) in a trachytic to subophitic groundmass of plagioclase, augite, and Fe-Ti oxide. Reversed magnetic polarity (J.T. Hagstrum, written commun., 2002). An  ${}^{40}$ Ar/ ${}^{39}$ Ar age of 1.282±0.014 Ma was obtained from a sample collected just east of the map area (R.J. Fleck, written commun., 2008)



Qaas	BASIN-FILL DEPOSITS Alluvium of ancestral Sandy River (Pleisto
Quus	<b>cene</b> )—Unconsolidated to well-cemented sandy grav el underlying uplands west of the Sandy River. Poor
	ly exposed along Beaver Creek, but well exposed be
	neath 12 to 18 m of cataclysmic-flood gravel (Qfg) is gravel pit 1 km north of Grant Butte; likely underlie
	flood gravel (Qfg) and loess (Qlo) throughout most o terrain south of Interstate 84. Gravel crudely strati
	fied and well sorted, contains abundant well-rounder clasts up to 20 cm across of basaltic and andesiti
	rocks derived from Cascade Range (including Moun Hood volcano) and sparse clasts of Columbia Rive
	Basalt Group and quartzite. Sand lenses, generall
	less than 1 m thick and 10 m long, composed chiefl of volcanic lithic grains; lack quartz and muscovite
	Along Beaver Creek, unit includes poorly sorted de bris-flow deposits that contain subrounded Moun
	Hood-derived pyroxene and hornblende andesit clasts as large as 1 m diameter. Unit at least 50 r
	thick in gravel pit north of Grant Butte; upper 11 to 15 m is distinctly lighter colored than the more oxi
	dized and cemented dark gray sandy gravel compos
	ing lowermost 35 m; surface of upper gravel de scends to north and west, indicating northwesterl
	transport, consistent with clast imbrication. Distribution, sedimentary texture and composition of unit in
	dicate deposition during major, glacially induced(?) aggradational episodes of the ancestral Sandy River
	producing a broad fan or braidplain extending nort and west into the eastern Portland Basin. Interbedde
	debris-flow deposits record lahars associated with eruptive activity at Mount Hood. Age poorly con
	strained; overlies weathered surfaces on the hyalo
	clastic sandstone member of the Troutdale Formatio (Ttfh); weak soil development indicates probable mid
QTc	dle-to-late Pleistocene age Conglomerate (Pleistocene and Plio
	<b>cene?)</b> —Unconsolidated to cemented, thick bedded pebble to boulder conglomerate with minor beds (<
	m thick) and lenses of basaltic and quartzofeldspath
	ic sandstone; up to 90 m thick. Conglomerate varie from well sorted, clast-supported, imbricated and
	crossbedded deposits composed largely of wel rounded clasts of Columbia River Basalt Group an
	quartzite to poorly sorted deposits dominated by sub angular to subrounded clasts of volcanic rocks erod
	ed from Cascade Range. Deeply weathered excep east of Lacamas Lake where top was removed by
	Missoula floods. Age poorly constrained; gradation ally to unconformably overlies hyaloclastic sand
	stone member of Troutdale Formation (Ttfh); proba
Qgwg	bly of latest Pliocene or early Pleistocene age Gravel west of Gresham (Pleistocene)—Unconsolidate
	sandy gravel underlying low hill between Greshar and Grant Butte. Consists of well rounded to sub
	rounded pebbles and cobbles (up to 20 cm diameter of volcanic rocks, chiefly andesites, derived from
	Cascade Range to the east, in oxidized silty matrix
	Clasts minimally weathered and weathering rinds of fine-grained volcanic clasts are less than 5 mm thick
	indicating a Pleistocene age. Topographic position suggests this unit is older than the uppermost grave
	of the ancestral Sandy River alluvium (Qaas) expose in excavation immediately to northwest, but it ma
	be an erosional remnant correlative with the litholog ically similar deeper gravel
QTfg	Unnamed fan gravel (Pleistocene and Pliocene?)-
	Deeply weathered unconsolidated sand and cobbl gravel poorly exposed along Beaver Creek valle
	near southern map boundary; underlies extensiv northwest sloping piedmont surface south of the ma
	area and west of Sandy River. Clasts consist chiefl of andesite and other Cascade Range rock types
	weathering rinds on clasts exceed 1 cm thick. Locall includes diamicts, interpreted as lahar and lahar-run
	out deposits, that contain 1-m-diameter clasts of py roxene and hornblende andesite. Capped by as muc
	as (25 m) of strongly oxidized and pedogenically al
	tered micaceous silt, inferred to be loess (QIO). Mor phology and composition of unit indicate deposition
	in broad fan by an ancient Sandy River draining Cas cade Range. Topographic position, degree of weath
	ering, and thick overlying loess mantle imply sub stantial age, probably early Pleistocene or late Plio
	cene. Lahar deposits derived from an ancient volcan ic center, possibly the Sandy Glacier volcano (Wise
	1969; Sherrod and Scott, 1995) located in the Moun Hood area. Mapped as Springwater Formation b
	Trimble (1963), but geomorphic relations outside th map area indicate that it is probably younger than de
	posits in the type area near Springwater, about 30 kr
QTwh	south of map area Walters Hill Formation (Pleistocene and (or) Plio
	<b>cene)</b> —Semiconsolidated, deeply weathered, poorlexposed fluvial gravel forming Grant Butte. Consist
	of well-rounded to subrounded cobbles and pebble of volcanic rocks, predominantly porphyritic ande
	sites derived from Cascade Range to the east; matri
	is coarse to fine volcaniclastic sand. Age unknown topographic position suggests unit is younger that
	Troutdale Formation (Ttfh) and older than other gravel el units (QTc, QTfg, Qgwg, Qaas) in map area
	Troutdale Formation (Pliocene and Miocene)—Semi consolidated to well consolidated conglomerate and
	sandstone. Divided into two informal members sepa
Ttfh	rated by an unconformity <b>Hyaloclastic sandstone member (Pliocene)</b> —Fluvia
	sedimentary strata distinguished by indurated, coars sandstone composed of abundant grains of glassy
	olivine+plagioclase-phyric basalt and conglomerat that contains olivine-bearing basalt clasts; overlie
	and intertongues with micaceous, arkosic sandstone
	siltstone, and claystone of the Sandy River Mudston (Tsr). Well exposed in valley of Lacamas Creek, or
	west shore of Lacamas Lake, and along Sandy Rive east of Troutdale. Hyaloclastic sandstone (equivalen
	to vitric sandstone of Trimble (1963) and Tolan and

to vitric sandstone of Trimble (1963) and Tolan and Beeson (1984)) consists largely to entirely of angular to subrounded fragments, 2 to 6 mm across, of black, generally nonvesicular basalt that contains phenocrysts of olivine (0.5-2 mm) and plagioclase (1-3)mm) in a glassy (sideromelane) to intergranular groundmass; sideromelane partly to completely altered to palagonite, which cements sandstone and imparts a distinctive yellowish-brown color to the originally dark-green rock; many beds contain minor admixed nonvolcanic debris such as quartz, muscovite, hornblende, and potassium feldspar. Sandstone ranges from poorly sorted to well sorted; typically thick-

bedded, commonly lenticular, crossbedded, and contains dispersed pebbles and cobbles of olivine-bearing basalt. Interbedded conglomerate mostly wellsorted and clast supported; consists of well rounded to subrounded pebbles and cobbles of variably vesicular, olivine+plagioclase-phyric basalt, basaltic rocks of Columbia River Basalt Group, and generally mi-

nor amounts of quartzite and other rock types; some conglomerate beds poorly sorted, with subangular basalt boulders as large as 1 m across. Angular nonvesicular vitric clasts in hyaloclastic sandstone beds interpreted as debris generated by basalt-water interaction in the Columbia River Gorge and rapidly transported downstream and deposited in eastern Portland Basin (Trimble, 1963; Swanson, 1986, 1988). Vitric clasts and associated basalt cobbles exhibit a low-potassium-tholeiite composition (Swanson, 1986, 1988) Lite; 1992; A.W. Sarna-Wojcicki, written commun. 2005; table 1, no. 34), indicating probable derivation from compositionally similar middle Pliocene basalt flows east of map area (Tolan and Beeson, 1984, Swanson, 1986, 1988; R.C. Evarts and R.M. Conrey, unpub. mapping). Leaf fossils from fine-grained interval within unit below Broughton Bluff date as early Pliocene (Trimble, 1963). Approximately equivalent to informal upper member of Troutdale Formation of Tolan and Beeson (1984)

Conglomerate member (Pliocene? and Miocene)—Indurated, well-sorted, clast-supported, pebble and cobble conglomerate exposed on steep valley wall of Little Washougal River near northeast corner of map area. Composed largely of well-rounded clasts of Columbia River Basalt Group with minor but persistent quartzite and granitic and felsic metamorphic rocks. Sparse interbeds of micaceous quartzofeldspathic sandstone. Equivalent in part to upper member of Troutdale Formation of Mundorff (1964). correlative with lower member of Troutdale Formation of Tolan and Beeson (1984) and with quartziteclast member of Troutdale Formation of Howard (2002) (see fig. 5). Late Miocene to early Pliocene age inferred from stratigraphic relations east of quadrangle (Tolan and Beeson, 1984) Tsr Sandy River Mudstone (Pliocene and Mio-

Ttfc

cene)—Shown in cross sections only. Semiconsolidated, well-bedded, fluvial and lacustrine sandstone, siltstone, claystone, and minor quartzite-bearing conglomerate, pumice-lapilli tuff, and lignite; not exposed in map area but water-well logs indicate this unit may be more than 300 m thick near western edge of quadrangle (Hartford and McFarland, 1989; Swanson and others, 1993). Approximately equivalent to informal lower member of Troutdale Formation as defined by Mundorff (1964). Age poorly known but probably ranges from middle Miocene to middle Pliocene; upper part underlies or interbedded with lower part of hyaloclastic sandstone member of Troutdale Formation, inferred to be about 3 to 4 m.y. old; plant fossils from localities near top of unit along Sandy River to southeast of quadrangle assigned an early Pliocene age (Chaney, 1944; Treasher, 1942; Trimble, 1963; Tolan and Beeson, 1984)

	er, 1942; Trimble, 1963; Tolan and Beeson, 1984)
	BEDROCK
Та	Andesite (Oligocene)—Platy porphyritic to seriate ande- site forming Ione Reef in Columbia River. Contains phenocrysts of plagioclase (18 percent; 1–2 mm long) and sparse microphenocrysts of orthopyroxene and Fe-Ti oxide in an intergranular groundmass of plagioclase, pyroxene, Fe-Ti oxide, and minor inter- stitial brown smectite. Age unknown but overlies ba- saltic andesite of Elkhorn Mountain (Tbem)
Td	<ul> <li>Dacite (Oligocene)—Sparsely plagioclase-phyric dacite flow and flow breccia exposed at east end of Lady Island; exhibits pronounced platy parting with widely variable orientation. Contains phenocrysts of plagioclase (1 percent; 0.5–1 mm long), olivine (0.5–1 mm across; 0.1 percent), and clinopyroxene (0.5–1 mm across; 0.1 percent) and microphenocrysts of Fe-Ti oxide and apatite in a pilotaxitic groundmass. Age unknown but overlies basaltic andesite of Elkhorn Mountain (Tbem)</li> </ul>
Tbem	Basaltic andesite of Elkhorn Mountain (Oligo-
Tvs	<ul> <li>cene)—Sequence of lava flows and flow brecia composed of dark-gray to brown, porphyritic to seriate to aphyric tholeiitic basaltic andesite and basalt; individual flows generally about 5 to 8 m thick, exhibit platy to blocky to columnar jointing; unit locally includes minor volcaniclastic rocks too thin or poorly exposed to map. Typical flows contain phenocrysts and glomerocrysts of weakly zoned plagioclase (0–40 percent; 1–6 mm, rarely more than 10 mm long; variably replaced by zeolites and (or) clay) and olivine (&lt;2 percent; 0.5 to 3 mm across; commonly partly resorbed and surrounded by rinds of granular pyroxene and (or) magnetite; rarely contain minute chromian spinel inclusions; almost invariably replaced by some combination of smectite, hematite, carbonate, serpentine, quartz and kaolinite). Some flows also contain phenocrysts of plagioclase, augite, Fe-Ti oxide, and minor to abundant interstitial glass (largely devitrified or replaced by smectite, quartz, or calcite); groundmass textures chiefly intergranular to trachytic, less commonly subophitic or microphyric. All flows in unit are tholeiitic, and many are exceptionally rich in Fe (FeO* as high as 13.3 wt percent) and poor in Sr (&lt;320 ppm) compared to mafic rocks elsewhere in southern Washington Cascade Range (du Bray and others, 2006; R.C. Evarts, unpub. data). <sup>40</sup>Ar/<sup>39</sup>Ar age of 25.5±0.2 Ma obtained for sample near top of unit north of Camas Slough</li> <li>Volcaniclastic sedimentary rocks (Oligocene)—Section of weathered, poorly exposed volcaniclastic rocks interbedded with basaltic andesite flows (Tbem) east of Lacamas Creek; approximately 30 m thick. Primarily massive, light-green to dark-brown, lithic lapilli tuffic</li> </ul>
	and pumice clasts to 5 cm across in sandy to silty matrix; pervasively altered to smectite, kaolinite, and zeolites. Abundance, variety, and subrounded charac- ter of lithic clasts suggests unit consists largely of debris-flow rather than primary pyroclastic-flow de- posits
Tmv	Mafic volcaniclastic rocks (Oligocene)—Indurated, olive-brown, mafic tuff breccia, lapilli tuff, and mi- nor tuff; distinct decimeter-scale stratification de- fined by gradational variations in grain size; moder- ately well sorted to poorly sorted; composed largely of angular fragments as large as 60 cm across of tex- turally variable, lithic to scoriaceous, basalt and ba- saltic andesite. Unit locally contains rounded lithic clasts and blackened wood fragments; one exposure displays matrix-poor, talus-like deposit of angular blocks to 2 m across; cemented by clay, zeolites, cal- cite
	<b>Contact</b> —Dashed where approximately located; short- dashed where inferred; dotted where concealed
<u> </u>	<b>Fault</b> —Dashed where inferred; dotted where concealed, queried where existence or extent uncertain. Ball and bar on downthrown side. Arrows show relative horizontal movement.
	<b>Reverse fault</b> —Dashed where inferred, dotted where concealed. Sawteeth on upper plate
25	Strike and dip of beds
5	Strike and dip of platy parting in lava flows
/	Basaltic andesite dike
● <sup>11</sup>	Sample locality for chemical analysis—See table 1
25.5±0.2	Sample locality for <sup>40</sup> Ar/ <sup>39</sup> Ar age determination; age in Ma (±1-σ error)—See table 2
×	Paleoflora locality
۲	Radiocarbon age locality—See table 3
	Cinder cone
<u></u>	

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