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## **CCSP Synthesis and Assessment Product 1.2**

### **Past Climate Variability and Change in the Arctic and at High Latitudes**

#### **Chapter 1 — Executive Summary**

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13 **1.1 Introduction**

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15 **Paleoclimate records** play a key role in our understanding of Earth’s past and present  
16 climate system and in our confidence in predicting future climate changes. Paleoclimate data  
17 help to elucidate past and present active mechanisms of climate change by placing the short  
18 instrumental record into a longer term context and by permitting **models** to be tested beyond the  
19 limited time that instrumental measurements have been available.

20 Recent observations in the Arctic have identified large ongoing changes and important  
21 **climate feedback mechanisms** that multiply the effects of global-scale climate changes. Ice is  
22 especially important in these “Arctic amplification” processes, which also involve the ocean, the  
23 atmosphere, and the land surface (vegetation, soils, and water). As discussed in this report,  
24 paleoclimate data show that land and sea ice have grown with cooling temperatures and have  
25 shrunk with warming ones, amplifying temperature changes while causing and responding to  
26 ecosystem shifts and sea-level changes.

27

28 **1.2 Major Questions and Related Findings**

29

30 *How have temperature and precipitation changed in the Arctic in the past? What does this tell*  
31 *us about Arctic climate that can inform projections of future changes?*

32 The Arctic has undergone dramatic changes in temperature and precipitation during the  
33 past 65 million years (m.y.) (the Cenozoic Era) of Earth history. Arctic temperature changes  
34 during this time exceeded global average temperature changes during both warm times and cold  
35 times, supporting the concept of Arctic amplification.

36           At the beginning of the Cenozoic Era, 65 million years ago (**Ma**), there was no sea ice on  
37 the Arctic Ocean, and neither Greenland nor Antarctica supported an ice sheet. General cooling  
38 since that time is attributed mainly to a slow decrease in **greenhouse gases**, especially carbon  
39 dioxide, in the atmosphere. Ice developed during this slow, “bumpy” cooling, first as mountain  
40 glaciers and as seasonal sea ice with the first continental ice sheet forming over Antarctica as  
41 early as 33 Ma ago. Following a global warm period about 3.5 Ma in the middle Pliocene, when  
42 extensive deciduous forests grew in Arctic regions now occupied by **tundra**, further cooling  
43 crossed a threshold about 2.6 Ma, allowing extensive ice to develop on Arctic land areas and thus  
44 initiating the Quaternary ice ages. This ice has responded to persistent features of Earth’s orbit  
45 over tens of thousands of years, growing when sunshine shifted away from the Northern  
46 Hemisphere and melting when northern sunshine returned. These changes were amplified by  
47 feedbacks such as **greenhouse-gas** concentrations that rose and fell as the ice shrank and grew,  
48 and by the greater reflection of sunshine caused by more-extensive ice. Human civilization has  
49 developed during the most recent of the relatively warm **interglacials**, the Holocene (about 11.5  
50 thousand years ago (**ka**) to the present). The penultimate warm interval, about 130–120 ka,  
51 received somewhat more Northern-Hemisphere summer sunshine than the Holocene owing to  
52 differences in Earth’s orbital configuration. Because this more abundant summer sunshine  
53 warmed the Arctic summer about 5°C above recent temperatures, the Greenland Ice Sheet was  
54 substantially smaller than its current size and almost all glaciers melted completely at that time.

55           The last glacial maximum peaked at about 20 ka when the Arctic was about 20°C colder  
56 than at present. Ice recession was well underway by 16 ka, and most of the Northern Hemisphere  
57 ice sheets melted by 7 ka. Summer sunshine rose steadily from 20 ka to a maximum (10% higher  
58 than at present due to the Earth’s orbit) about 11 ka ago, and has been decreasing since then. The

59 extra energy received in summer in the early Holocene resulted in warmer summers throughout  
60 the Arctic. Summer temperatures were 1°–3°C above 20th century averages, enough to  
61 completely melt many small glaciers in the Arctic and to slightly shrink the ice sheet on  
62 Greenland. Summer sea-ice limits were significantly less than their 20th century average. As  
63 summer sunshine decreased in the second half of the Holocene, glaciers re-established or  
64 advanced, and sea ice became more extensive. Late Holocene cooling reached its nadir during  
65 the Little Ice Age (about 1250–1850 AD), when most Arctic glaciers reached their maximum  
66 Holocene extent. The Little Ice Age temperature minimum may also have been augmented by  
67 multiple large volcanic eruptions that lofted a reflective aerosol layer into the stratosphere at that  
68 time. Subsequent warming during the 19<sup>th</sup> and 20<sup>th</sup> centuries has resulted in Arctic-wide glacier  
69 recession, the northward advance of terrestrial ecosystems, and the reduction of perennial (year-  
70 round) sea ice in the Arctic Ocean. These trends will continue if greenhouse gas concentrations  
71 continue to increase into the future.

72 **Paleoclimate reconstructions** of Arctic temperatures compared with global temperature  
73 changes during four key intervals during the past 4 m.y. allow a quantitative estimate of Arctic  
74 amplification. These data suggest that Arctic temperature change is 3 to 4 times the global  
75 average temperature change during both cold and warm departures.

76

77 *How rapidly have temperature and precipitation changed in the Arctic in the past? What do*  
78 *these past rates of change tell us about Arctic climate that can inform projections of future*  
79 *changes?*

80 As discussed with the previous question, climate changes on numerous time scales for various  
81 reasons, and it has always done so. In general, longer-lived changes are somewhat larger but  
82 much slower than shorter-lived changes.

83

84 Processes linked to **continental drift (plate tectonics)** have affected atmospheric and oceanic  
85 currents and the composition of the atmosphere over tens of millions of years; in the Arctic, a  
86 global cooling trend has switched conditions from being ice-free year-round near sea level to icy  
87 conditions more recently. Within the icy times, variations in Arctic sunshine in response to  
88 features of Earth's orbit have caused regular cycles of warming and cooling over tens of  
89 thousands of years that were roughly half the size of the continental-drift-linked changes. This  
90 "glacial-interglacial" cycling was amplified by colder times bringing reduced greenhouse gases  
91 and greater reflection of sunlight, especially from expanded ice-covered regions. This glacial-  
92 interglacial cycling has been punctuated by sharp-onset, sharp-end (in as little as 1–10 years)  
93 **millennial oscillations**, which near the North Atlantic were roughly half as large as the glacial-  
94 interglacial cycling but which were much smaller Arctic-wide and beyond. The current warm  
95 period of the glacial-interglacial cycling has been influenced by cooling events from single  
96 volcanic eruptions, slower but longer lasting changes from random fluctuations in frequency of  
97 volcanic eruptions and from weak solar variability, and perhaps by other classes of events. Very  
98 recently, human effects have become evident, not yet showing both size and duration that exceed  
99 peak values of natural fluctuations further in the past, but with projections indicating that human  
100 influences could become anomalous in size and duration and, hence, in speed.

101

102 *What does the paleoclimate record tell us about the past size of the Greenland Ice Sheet and*  
103 *its implications for sea level changes?*

104 The paleo-record shows that the *Greenland Ice Sheet* has consistently lost mass and  
105 contributed to sea-level rise when the climate warmed, and has grown and contributed to sea-  
106 level fall when the climate cooled. This occurred even at times when offsetting effects from  
107 elsewhere in the climate system caused the net sea-level change around Greenland to be  
108 negligible, and so these changes in the ice sheet cannot have been caused primarily by sea-level  
109 change. In contrast, no changes in the ice sheet have been documented independent of  
110 temperature changes. Moreover, snowfall has increased with major climate warmings, but the ice  
111 sheet lost mass nonetheless; increased accumulation in the ice sheet center was not sufficient to  
112 counteract increased melt and flow near the edges. Most of the documented changes (of both ice  
113 sheet and **forcings**) spanned multi-millennial periods, but limited data show rapid responses to  
114 rapid forcings have also occurred. In particular, regions near the ice margin have been observed  
115 to respond within a few decades or less. However, major changes of the ice sheet are thought to  
116 take centuries to millennia, and this is supported by the limited data.

117 The paleo-record does not yet give any strong constraints on how rapidly a near-complete loss of  
118 the ice sheet could occur, although the paleo-data indicate that onset of shrinkage will be  
119 essentially immediate after forcings begin. The available evidence suggests such a loss requires  
120 a sustained warming of at least 2-7°C above mean 20th century values, but this threshold is  
121 poorly defined. The paleo-archives are sufficiently sketchy that temporary ice sheet growth in  
122 response to warming, or changes induced by factors other than temperature, could have occurred  
123 without being recorded.

124

125 *What does the paleoclimate record tell us about past changes in Arctic sea ice cover, and what*  
126 *implications does this have for consideration of recent and potential future changes?*

127         Although incomplete, existing data outline the development of Arctic sea-ice cover from  
128 the ice-free conditions of the early Cenozoic. Some data indicate that sea ice has covered at least  
129 part of the Arctic Ocean for the last 13–14 million years, and it has been most extensive during  
130 the last several million years in relationship with Earth’s overall cooler climate. Other data argue  
131 against the development of perennial (year-round) sea ice until the most recent 2 – 3 million  
132 years. Nevertheless, episodes of considerably reduced ice cover, or even a seasonally ice-free  
133 Arctic Ocean, probably punctuated even this latter period. Warmer climates associated with the  
134 **orbitally-paced** interglacials promoted these episodes of diminished ice. Ice cover in the Arctic  
135 began to diminish in the late 19th century and this shrinkage has accelerated during the last  
136 several decades. Shrinkages that were both similarly large and rapid have not been documented  
137 over at least the last few thousand years, although the paleoclimatic record is sufficiently sparse  
138 that similar events might have been missed. Orbital changes have made ice melting less likely  
139 than during the previous millennia since the end of the last ice age, making the recent changes  
140 especially anomalous. Improved reconstructions of sea-ice history would help clarify just how  
141 anomalous these recent changes are.

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### 143 **1.3 Recommendations**

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145         Paleoclimatic data on the Arctic are generated by numerous international investigators  
146 who study a great range of archives throughout the vast reaches of the Arctic. The value of this  
147 diversity is evident in this report. Many of the key results of this report rest especially on the

148 outcomes of community-based syntheses, including the CAPE Project, and multiply replicated,  
149 heavily sampled archives such as the central Greenland deep ice cores. Results from the ACEX  
150 deep coring in Arctic Ocean sediments were appearing as this report was being written. These  
151 results are quite valuable and will become more so with synthesis and replication, including  
152 comparison with land-based and marine records. The number of questions answered, and raised,  
153 by this one new data set shows how sparse the data are on many aspects of Arctic paleoclimatic  
154 change. *Future research should maintain and expand the diversity of investigators,*  
155 *techniques, archives, and geographic locations, while promoting development of community-*  
156 *based syntheses and multiply replicated, heavily sampled archives. Only through breadth and*  
157 *depth can the remaining uncertainties be reduced while confidence in the results is improved.*

158  
159 The questions asked of this study by the CCSP are relevant to public policy and require  
160 answers. The answers provided here are, we hope, useful and informative. However, we  
161 recognize that despite the contributions of many community members to this report, in many  
162 cases a basis was not available in the refereed scientific literature to provide answers with the  
163 accuracy and precision desired by policymakers. *Future research activities in Arctic*  
164 *paleoclimate should address in greater detail the policy-relevant questions motivating this*  
165 *report.*

166  
167 Paleoclimatic data provide very clear evidence of past changes in important aspects of the  
168 Arctic climate system. The ice of the *Greenland Ice Sheet*, smaller glaciers and ice caps, the  
169 Arctic Ocean, and in soils is shown to be vulnerable to warming, and Arctic ecosystems are  
170 strongly affected by changing ice and climate. National and international studies generally



171 project rapid warming in the future. If this warming occurs, the paleoclimatic data indicate that  
172 ice will melt and associated impacts will follow, with implications for ecosystems and  
173 economies. *The results presented here should be utilized by science managers in the design of*  
174 *monitoring, process, and model-projection studies of Arctic change and linked global*  
175 *responses.*