

381 **Executive Summary**

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390

391 **SYNOPSIS**

392 Depletion of the stratospheric ozone layer by human-produced ozone-depleting
393 substances has been recognized as a global environmental issue for three decades, and the
394 international effort to address the issue via the United Nations Montreal Protocol marked
395 its 20-year anniversary in 2007. Scientific understanding underpinned the Protocol at its
396 inception and ever since. As scientific knowledge advanced and evolved, the Protocol
397 evolved through amendment and adjustment. Policy-relevant science has documented the
398 rise, and now the beginning decline, of the atmospheric abundances of many ozone-
399 depleting substances in response to actions taken by the nations of the world. Projections
400 are for a return of ozone-depleting chemicals (compounds containing chlorine and
401 bromine) to their “pre-ozone-depletion” (pre-1980) levels by the middle of this century
402 for the midlatitudes; the polar regions are expected to follow suit within 20 years after

403 that. Global ozone sustained a depletion of about 5% since the 1980s in the midlatitudes
404 of both the Northern Hemisphere and Southern Hemisphere, where most of the Earth's
405 population resides; it is showing signs of turning the corner toward the return to 1980
406 levels. The large seasonal depletions in the polar regions are likely to continue over the
407 next decade but are expected to subside over the next few decades. Ozone-depleting
408 substances should have a negligible effect on ozone in all regions beyond 2070, assuming
409 continued compliance with the Montreal Protocol.

410

411 Large increases in surface ultraviolet (UV) radiation and the associated impacts on
412 human health and ecosystems would have occurred if atmospheric abundances of ozone-
413 depleting substances had continued to grow. Scientific findings regarding the role of
414 ozone-depleting chemicals, projected ozone losses, and the potential UV impacts
415 galvanized international decision making in the 1980s. As a result of the worldwide
416 adherence to the 1987 Montreal Protocol and its amendments and adjustments, the large
417 impacts were avoided, and future UV trends at the surface are expected to be more
418 influenced by factors other than stratospheric ozone depletion (such as changes in clouds,
419 atmospheric fine particles, and air quality in the lower atmosphere).

420

421 Emissions of ozone-depleting substances by the United States have been significant
422 throughout the history of the ozone depletion issue. At the same time, the United States
423 has played a leading role in advancing the scientific understanding, leading the
424 international decision making, and leading industry's actions to reduce usage of ozone-
425 depleting substances. Continued future declines in emissions of ozone-depleting

426 substances from the United States, along with those from other nations, will play a key
427 role in ensuring the ozone layer's recovery.

428

429 Projections of a changing climate have added a new dimension to the issue of the
430 stratospheric ozone layer and its recovery, and scientific knowledge is emerging on the
431 interconnections between these two global issues. Climate change is expected to alter the
432 timing of the recovery of the ozone layer depletion. Ozone-depleting chemicals and
433 ozone depletion are known to influence climate change. The curtailment of the ozone-
434 depleting substances not only helped the ozone layer but also lessened the forcing of
435 climate (*i.e.*, how it alters climate).

436

437 Climate change and ozone layer depletion are coupled; this has led to new scientific and
438 decision-making challenges. The recovery of the ozone layer will occur in an atmosphere
439 that is different from where we started. Our scientific understanding of the connections
440 between climate change and ozone layer depletion is at an early but rapidly advancing
441 stage. That topic will remain a focus for the scientific community's efforts over the next
442 few decades.

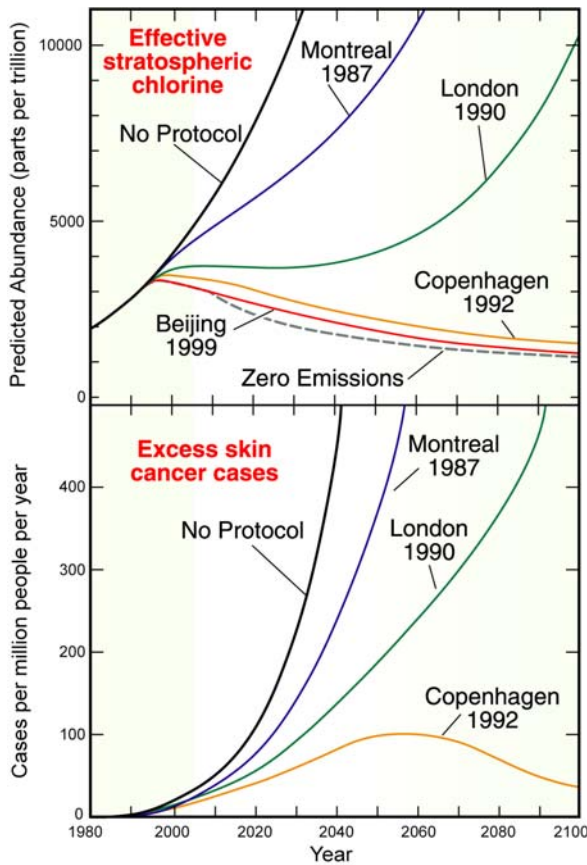
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444 **ES.1 WHAT IS OZONE LAYER DEPLETION AND WHY IS IT A CONCERN?**

445 The stratospheric ozone layer lies in a region of the atmosphere approximately 15 to 45
446 kilometers above Earth's surface. The ozone layer acts as a protective shield, preventing
447 most of the Sun's harmful ultraviolet (UV) radiation from reaching the surface. The
448 depletion of the ozone layer can therefore lead to enhancements of the UV radiation that
449 reaches Earth's surface, with consequences for human health, the Earth's ecosystems, and

450 physical materials. The ozone layer and its changes can also alter the atmosphere's
 451 temperature structure and weather/climate-related circulation patterns.
 452
 453 Research in the 1970s and early 1980s had shown that the ozone-depleting substances
 454 (ODSs), mainly chlorofluorocarbons (CFCs) and certain compounds containing bromine,
 455 would deplete stratospheric ozone. The discovery of the ozone hole in 1985 showed that
 456 ozone depletion was real and occurring at that time, and was not just a prediction for the
 457 future.

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461 **Figure ES.1** Effect of the Montreal Protocol. The top panel gives a measure of the projected future abundance of
 462 ozone-depleting substances in the stratosphere, without and with the Protocol and its various Amendments. The
 463 bottom panel shows similar projections for how excess skin cancer cases might have increased. (From Fahey, 2007)
 464

465 Faced with the scientific consensus that ozone depletion was real and due to human-
466 produced ozone-depleting substances, nations throughout the world agreed to the
467 Montreal Protocol and its subsequent amendments and adjustments. The United States is
468 a signatory to this protocol. The Protocol and its amendments were successfully
469 implemented starting in the late 1980s. Thus, this Protocol was one of the first
470 international agreements to address a global environmental problem. This Protocol has
471 had clear benefits in reducing ozone-depleting substances, placing the ozone layer on a
472 path to recovery, and protecting human health (Figure ES.1).

473

474 Ozone layer depletion, like climate change, is a global issue with regional impacts. The
475 depletion of the ozone layer is caused by the collective emissions of human-produced
476 ozone-depleting substances at Earth's surface from various regions and countries. These
477 ozone-depleting substances persist long enough in the atmosphere to be quite well mixed
478 in the lower atmosphere and then be transported to the stratosphere. Thus, they pose a
479 global threat, irrespective of where on Earth's surface they are emitted. Emissions of
480 ozone-depleting substances arise from their use as coolants, fire-extinguishing chemicals,
481 electronics cleaning agents, and in foam blowing and other applications. The
482 contributions to the global atmospheric burden of these ozone-depleting substances vary
483 by regions and countries. There are large variations in the extent and timing of ozone
484 depletion in various regions, and the impacts are also different. Consequently, the
485 impacts of ozone layer depletion can be different in different regions of the world.

486

487 The findings from this Synthesis and Assessment Product are summarized in three parts.
488 Below, Section ES.2 of this Executive Summary lists the findings to inform the public in
489 general nontechnical terms, and Section ES.3 summarizes findings for those involved in
490 potential policy formulation. The Executive Summary findings are backed up by a more
491 technical set of findings, primarily for scientists and secondarily for those who want to
492 delve more into the details. These technical findings are listed near the beginning of
493 Chapters 2 through 5, and in Chapter 6 on Policy Implications for the United States.
494 Appendix A of this Synthesis and Assessment Product provides extensive background
495 material on the science regarding the ozone layer, ozone-depleting substances, surface
496 ultraviolet radiation, and connections to climate change.

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498 **ES.2 KEY FINDINGS ABOUT THE OZONE LAYER, SURFACE UV, OZONE-**
499 **DEPLETING SUBSTANCES, AND CONNECTIONS TO CLIMATE CHANGE**

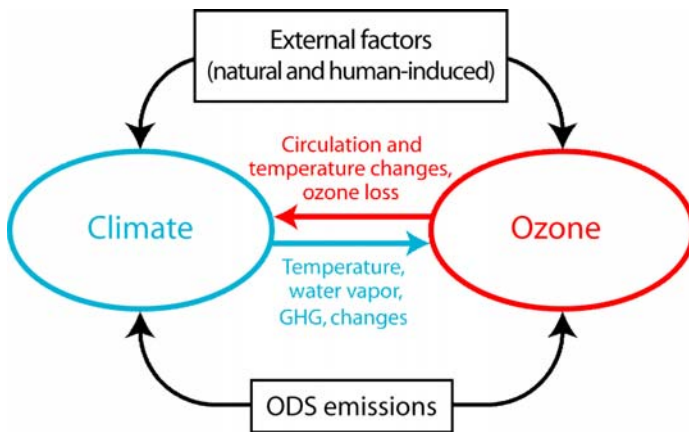
500 **ES.2.1 The Ozone Layer, Ozone-Depleting Substances, and Climate Change: What**
501 **Are the Connections?**

502 *Ozone layer changes caused by ozone-depleting substances are intertwined with the issue*
503 *of climate change, even though the two issues have been distinct in most policy*
504 *formulations.*

505

506 Over the course of the past 20 years, the close connections between stratospheric ozone
507 depletion and climate change issues have become clearer (Figure ES.2).

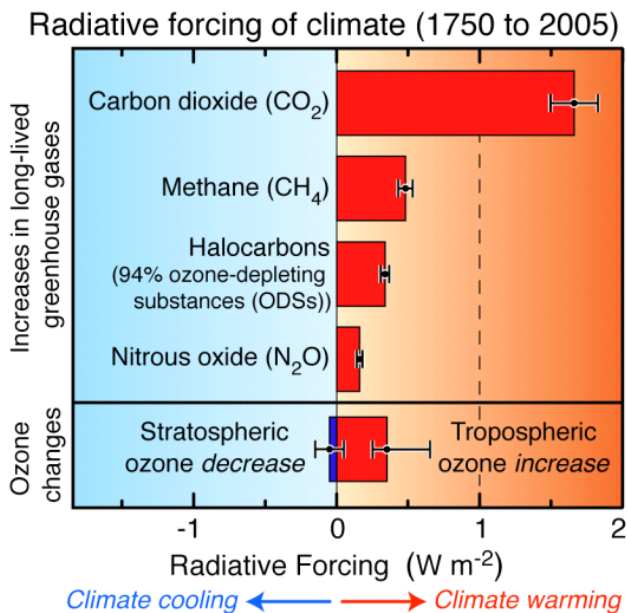
- 508 • Ozone-depleting substances and many of the chemicals being used to replace
- 509 them are potent greenhouse gases that influence the Earth’s climate by trapping
- 510 terrestrial infrared (heat) radiation that would otherwise escape to space.
- 511 • Ozone is itself a greenhouse gas, and the stratospheric ozone layer heats the
- 512 stratosphere and lower atmosphere (troposphere) and is a key component that
- 513 affects climate. Depletion of the ozone layer thereby influences climate.
- 514 • The recovery of the ozone layer is influenced not only by the decreases in ozone-
- 515 depleting substances required by the Montreal Protocol, but also by changes to
- 516 climate and Earth’s atmospheric composition.



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 521
 522 **Figure ES.2** Simplified schematic of some of the processes that interconnect the issues of ozone layer
 523 depletion and climate change. (Adapted from Chapter 4 of this report)
 524

525 Ozone-depleting substances are continuing to make a significant contribution to global
 526 climate change, but in the future ODSs are expected to make a smaller and smaller
 527 contribution. The ODS contribution to global climate change between 1750 and 2005, as
 528 measured by a quantity called radiative forcing that is a metric for the ability to force

529 climate change, is approximately 20% of that from carbon dioxide, the largest
 530 anthropogenic contributor to global radiative forcing (Figure ES.3). The combined
 531 radiative forcing from ODSs and substitutes including HFCs is still increasing, but at a
 532 much slower rate than in the 1980s. The total contribution of human-produced ODSs and
 533 substitutes in 2005 was about 15% of the contribution from the major greenhouse gases
 534 (carbon dioxide, methane, and nitrous oxide). The ODS contribution is expected to
 535 decline in coming decades as ODS emissions decline and CO₂ emissions continue to rise.
 536



537
 538 **Figure ES.3** Radiative forcing values for the principal contributions to climate change from atmospheric
 539 gas changes since preindustrial times, including halogen-containing gases such as ODSs, and the cooling
 540 caused by depletion of stratospheric ozone. These climate influences are expressed as radiative forcings, a
 541 metric for the ability to force climate change. (Adapted from IPCC, 2007)
 542
 543

544 Depletion of stratospheric ozone since about 1980 has caused a slight *negative* (cooling)
 545 radiative forcing of climate (approximately -0.05 W per m²). This forcing is small; it is
 546 roughly 15% of, and in the opposite direction to, climate forcing by the ODSs that caused

547 the depletion. Thus, ozone layer depletion currently does not appear to significantly offset
548 the positive (warming) climate forcing by ODSs that caused the depletion (Figure ES.3).

549

550 Climate change will lead to either increases or decreases in ozone abundances depending
551 on the location in the atmosphere and the magnitude of climate change. Observed
552 stratospheric temperature decreases began in the 1960s and are expected to continue. This
553 trend is attributed to ozone depletion, increased carbon dioxide (CO₂), and changes in
554 water vapor. Stratospheric temperatures influence ozone amounts through chemical and
555 transport processes. Stratospheric water vapor influences stratospheric ozone through
556 chemistry, formation of polar stratospheric clouds, and changes in temperature.

557

558 **ES.2.2 Ozone-Depleting Substances: Past, Present, and Future**

559 *The Montreal Protocol has been effective in reducing the use of ozone-depleting*
560 *substances. Assuming continued compliance with the Protocol, the atmospheric*
561 *abundance of ODSs is expected to decline back to its pre-1980 level by the middle of this*
562 *century.*

563

564 Total global production and consumption of ODSs have declined substantially since the
565 late 1980s in response to the Montreal Protocol. By 2005, the annual aggregated
566 production and consumption magnitudes of the ODSs, after accounting for their
567 differences in ozone depletion capabilities, had declined 95% from peak amounts
568 produced and consumed in the late 1980s.

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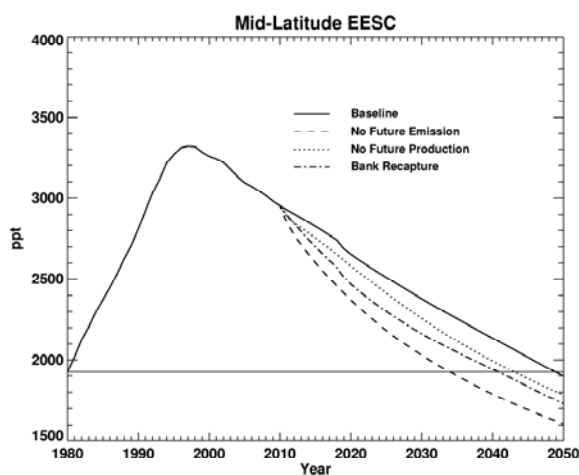
570 In response to these global production and consumption changes, global ODS emissions
571 have declined. Hence, the total amount of ozone-depleting substances in the atmosphere,
572 as measured by their combined ability to deplete the ozone layer, is now decreasing both
573 in the troposphere and stratosphere.

574

575 In this report, future halocarbon emissions are derived using a new bottom-up approach
576 for estimating emissions from the sizes of the banks (ODSs produced but not yet
577 released). The new method gives future CFC emissions that are higher than previously
578 estimated in WMO (2003). There are still some uncertainties in the future abundances of
579 ODSs.

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584 **Figure ES.4** Estimates of the effective sum of ozone-depleting chlorine and bromine in the stratosphere
585 (called Equivalent Effective Stratospheric Chlorine, EESC), a metric that accounts for the differences in
586 ozone depletion capabilities of chlorine and bromine. Estimates in the past are based upon observations,
587 and ones in the future are based upon a baseline scenario and three comparative test cases. The horizontal
588 line represents the 1980 (“pre-ozone-depletion”) level of EESC. (adapted from WMO, 2007; see also
589 chapter 5)

590

591 The effective sum of chlorine and bromine in the stratosphere, with bromine weighted by
592 its larger per-atom efficiency in depleting ozone, is estimated to recover to the 1980 value
593 between 2040 and 2050 in the midlatitudes (Figure ES.4), and between 2060 and 2070 in
594 the polar regions.

595

596 **ES.2.3 Ozone in the Stratosphere: Past, Present, and Future**

597 *Total global ozone, as well as seasonal springtime ozone in both southern and northern*
598 *polar regions, exhibited declines since the early 1980s, but recent observations show that*
599 *ozone depletion is not worsening and in some atmospheric regions is showing signs of*
600 *beginning recovery. Ozone in the future is projected to recover as the atmospheric*
601 *amounts of ODSs decline over the next few decades (with recovery above midlatitudes*
602 *and the Arctic preceding Antarctic recovery). With continued adherence to the Montreal*
603 *Protocol, ozone-depleting substances identified in the Protocol should have a negligible*
604 *effect on ozone in all regions beyond 2070.*

605

606 Total global ozone declined by roughly 5% since the early 1980s but has remained
607 relatively constant over the last four years (2002 to 2006). Northern midlatitude ozone
608 reached a minimum in 1993, and has increased somewhat since then. The 1993 minimum
609 largely resulted from the increase of particles in the stratosphere present due to the
610 eruption of Mt. Pinatubo. Southern midlatitude ozone decreased until the late 1990s, and
611 has been constant since. There are no significant ozone trends over the tropics.

612

613 Ozone depletion in the upper stratosphere, where the influence of chlorine is easiest to
614 detect, has slowed, and has closely followed the trends in the sum of chlorine plus
615 bromine. The slowdown of the negative (or decreasing) trend may be attributed to the
616 fact that ozone-depleting chlorine and bromine are leveling off in this region of the
617 stratosphere.

618

619 Antarctic ozone depletion can be measured in different ways, *e.g.*, total amount of ozone
620 lost (called mass deficit), minimum values of ozone observed, geographical area of the
621 ozone hole, *etc.* Over the last decade (1995 to 2006), the Antarctic ozone depletion by
622 these measures has not worsened. The ozone hole area and ozone mass deficit were
623 observed to be below average in some recent winter years while higher minimum column
624 amounts have also been recorded. This variability results from the strong influence of
625 meteorological variability on ozone amounts, and not from any changes in the amounts of
626 chlorine and bromine available for ozone depletion.

627

628 Arctic spring total ozone values over the last decade were lower than values observed in
629 the 1980s. In addition, spring Arctic ozone is highly variable depending on
630 meteorological conditions. For current halogen levels, human-caused chemical loss and
631 variability in ozone transport are about equally important for year-to-year Arctic ozone
632 variability. Colder-than-average vortex conditions result in larger halogen-driven
633 chemical ozone losses.

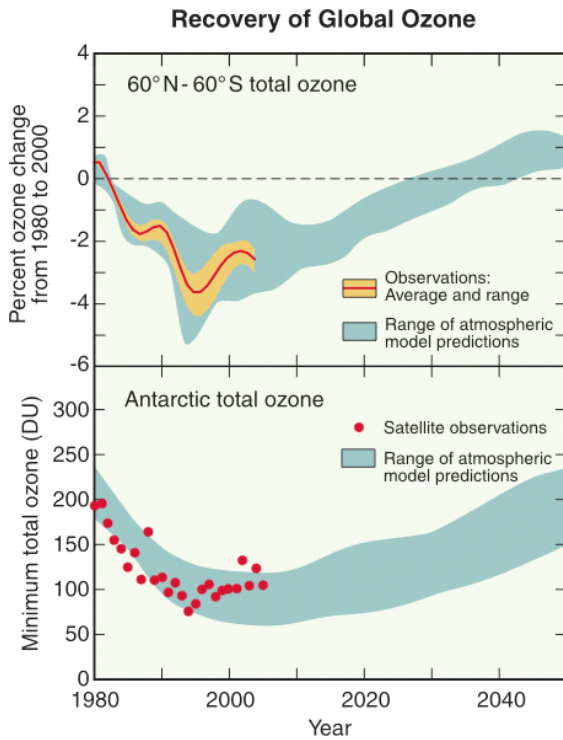
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635 If explosive volcanic eruptions were to occur in the coming decades, they are expected to
636 cause major temperature and circulation changes in the stratosphere as have occurred
637 after past eruptions. The changes are caused by the large increases in fine particles
638 formed from sulfur dioxide injected into the stratosphere following such eruptions. The
639 increases result in only a short-term shift in stratospheric ozone levels and climate
640 because natural processes gradually remove the additional sulfate particles within a few
641 years after the eruption.

642

643 Based on the projected changes in ozone-depleting substances and changes in the major
644 climate-relevant trace gases, assuming no changes in activities such as a volcanic
645 eruption, and using modeling calculations, the following are predicted for the future of
646 the ozone layer (Figure ES.5):

- 647 • The ozone content between 60°N – 60°S, between now and 2020, will increase in
648 response to decreases in halogen loading.
- 649 • Global ozone is expected to return to its 1980 value up to 15 years earlier than the
650 halogen recovery date because of stratospheric cooling and changes in circulation
651 associated with greenhouse gas emissions.
- 652 • Global ozone abundances are expected to be 2 to 5% above the 1980 values by
653 2100 for the assumed scenario for greenhouse gases noted in this report.
- 654 • The minimum ozone value for Antarctic ozone is projected to start increasing
655 after 2010 in several model calculations, while another measure of ozone
656 depletion (the ozone mass deficit, the total amount lost in a season) begins
657 decreasing around 2005 in most models.



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 661

Figure ES.5 Global ozone recovery predictions. (From Fahey, 2007)

- 662
- Model simulations show that the ozone amount in the Antarctic will reach the
- 663
- 1980 values 10 to 20 years earlier than the 2060 to 2070 time frame of when the
- 664
- ozone-depleting substances reach their 1980 levels in polar regions.
- 665
- Ozone in the Arctic region is expected to increase as ODSs decline in the
- 666
- atmosphere. Because of large interannual variability, the simulated results do not
- 667
- show a smooth monotonic recovery of Arctic ozone. The dates of the minimum
- 668
- ozone from different models occur between 1997 and 2015.
- 669
- Most climate chemistry models show Arctic ozone values by 2050 larger than the
- 670
- 1980 values, with the recovery date between 2020 and 2040.
- 671

672 The above projections are based on currently available models. As our scientific
673 understanding and modeling capabilities continue to evolve, our best predictions of the
674 timing and extent of ozone layer recovery will also evolve.

675

676 **ES.2.4 Surface Ultraviolet Radiation: Past, Present, and Future**

677 *The Montreal Protocol and its amendments have prevented large increases in global*
678 *surface UV radiation. As the stratospheric ozone layer recovers over the next few*
679 *decades, factors such as changes in clouds, atmospheric fine particles, and air quality in*
680 *the lower atmosphere will be the dominant factors influencing future UV changes.*

681

682 Surface UV changes resulting from ozone depletion over Antarctica in early Austral
683 Spring have been very large. Changes in the surface UV due to ozone depletion in most
684 other locations of the world have not been clearly discernable, because the effects have
685 been much smaller compared with changes due to other factors. For example, trends in
686 UV exposure changes at ground level in the midlatitude United States are difficult to
687 discern and are dependent on changes in clouds and pollution from suspended fine
688 particles in the air in addition to ozone changes. What is clear is that in the absence of the
689 Montreal Protocol, ozone depletion would have caused increases in surface UV by 2010
690 over most of the world, to such an extent that other factors (*e.g.*, clouds, atmospheric fine
691 particles, air quality) would have been of relatively minor importance.

692

693 The future UV trend at the surface is likely to be influenced more by changes in cloud,
694 aerosols, and lower atmosphere air quality than by ozone layer depletion.

695

696 **ES.3 IMPLICATIONS FOR THE UNITED STATES: IMPACTS,**697 **ACCOUNTABILITY, AND POTENTIAL MANAGEMENT OPTIONS**

698 It is not possible to make a simple connection between emissions of ozone-depleting
699 substances from the United States with the depletion of ozone above the country. This is
700 because ODSs persist long enough in the atmosphere to be quite well mixed in the global
701 lower atmosphere, before transport to the stratosphere occurs. Thus, ODSs pose a global
702 threat, irrespective of where on Earth's surface they are emitted. However, the depletion
703 of stratospheric ozone over the various United States regions, and the contribution of
704 emissions from the United States to the global burden of ozone-depleting substances, can
705 be quantified.

706

707 **Impacts: Changes in Ozone and Surface Ultraviolet Radiation Over the United**708 **States**

709 Ozone depletion above the continental United States (*i.e.*, the midlatitudes) has
710 essentially followed the depletion occurring over the northern midlatitude regions: a
711 decrease to a minimum around the mid-1990s and a slight increase since that time. The
712 minimum total column ozone amounts over the continental United States, reached in
713 1993, were about 5-8% below the amounts present prior to 1980. The ozone increase
714 since 1993 has diminished the ozone deficit to about 2-5% below the pre-1980 amounts.
715 These midlatitude ozone changes are estimated to have a significant contribution from the
716 ozone depletion that occurs in the Arctic during springtime.

717

718 Ozone over Northern high latitudes, such as over northern Alaska, is most influenced by
719 Arctic springtime total ozone values, which in recent years have been lower than those
720 observed in the 1980s. The springtime ozone depletions are highly variable from year to
721 year.

722

723 Calculations based on satellite observations of column ozone and reflectivity of the
724 surface suggest that the averaged erythemal irradiance (which is a weighted combination
725 of UVA and UVB based on skin sensitivity) over the United States had increased roughly
726 by about 7% at the time when the ozone minimum was reached in 1993 and is now about
727 4% higher than in 1979. Direct surface-based observations do not show significant trends
728 in UV levels over the United States over the past three decades because effects of clouds
729 and aerosol have likely masked the increase in UV due to ozone depletion over this
730 region.

731

732 **Accountability: U.S. Contributions to Ozone-Depleting Substances**

733 The contributions of the United States to the emission of ODSs to date have been
734 significant. For example, in terms of the regulated uses of ODSs, emissions from the
735 United States accounted for between 15 and 39% of the overall atmospheric abundance
736 of ODSs measured during 1994 and 2004. The United States has also contributed
737 significantly to emission reductions of ODSs, thereby helping efforts to achieve the
738 expected recovery of the ozone layer and prevent large surface UV changes.

739

740

741 Future Options

742 United States emissions of ODS in the future, like those from other developed nations,
743 will be determined to a large extent by the size of “banks of ODSs,” *i.e.*, those ODSs that
744 are already produced but not yet released to the atmosphere. The expected future
745 declining emissions of ODSs from the United States will also aid in reducing the climate
746 forcing from these substances.

747

748 While the Montreal Protocol has had a large beneficial effect on current and projected
749 ozone depletion, there remain options for the United States, and other countries as well,
750 to reduce ozone depletion arising from ozone-depleting substances over the coming
751 decades. The greatest reduction possible would be obtained from the hypothetical
752 cessation of all future emissions of ozone-depleting substances (including emissions from
753 banks and future production). If such a cessation had been implemented in 2007, the
754 anticipated return of the ozone-depleting substances to their 1980 level would be
755 advanced by about 15 years.

756

757 Methyl bromide is a potent ODS that has significant unregulated quarantine and pre-
758 shipment uses, and critical use exemptions that are large compared to current regulated
759 uses. The importance of human-emitted methyl bromide to future ozone depletion will
760 depend on the magnitude of these future unregulated uses and of the critical use
761 exemptions. Reducing such unregulated emissions would benefit the ozone layer.

762

763

764 **The World Avoided**

765 Without the Montreal Protocol regulations, the levels of ODSs around 2010 likely would
766 have been more than 50% larger than currently predicted (Figure ES.1). The abundances
767 in the remaining 21st century would have depended on the specific actions taken by
768 humankind. The increases in ODSs would have caused a corresponding substantially
769 greater global ozone depletion. The Antarctic ozone hole would have persisted longer and
770 may have been even larger than what has been observed to date.

771

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