



# Satellite Remote Sensing of Small Ice Crystals in Cirrus Clouds

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## ABSTRACT

In situ instruments have not been able to characterize the concentrations of small ( $D < 60 \mu\text{m}$ ) ice crystals in cirrus clouds although the 2DS probe appears promising. Some instruments indicate small crystal concentrations are orders of magnitude greater than the larger crystals in a size distribution, while other instruments indicate no anomaly. Therefore we have developed a method for detecting the relative concentration of these small crystals from satellites.

Some in situ measurements indicate small crystals may account for about half of the optical depth of cirrus clouds and the nature of the cirrus cloud feedback in GCM simulations depends strongly on their concentration. Thus simulations of future climate may be highly uncertain until their concentrations are adequately characterized.

## General Approach

Parameterize measurements of the ice particle size distribution (PSD) that include small ice crystals as a function of temperature and IWC. Incorporate this parameterization into a retrieval algorithm that calculates the relative concentration of small mode ice crystals using radiances measured from satellites. Combining in situ and satellite measurements in this way, the PSD is retrieved for a given IWC, even if the PSD is bimodal. Also retrieved are the small-to-large mode number concentration ratio, the IWP and effective size  $D_e$ .

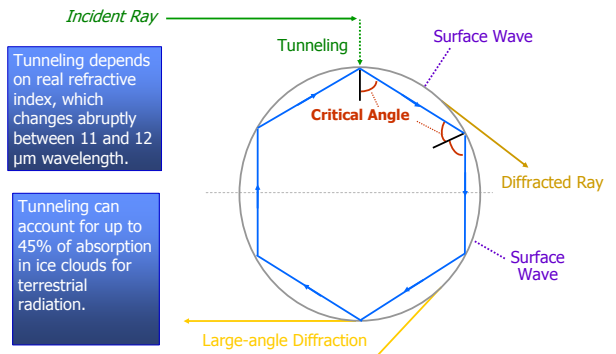
## How Photon Tunneling Can be Used to Remotely Detect Small Ice Crystals in Cirrus

Photon Tunneling is the process by which radiation beyond the physical cross-section of a particle is either absorbed or scattered outside the forward diffraction peak. Tunneling is strongest when:

- 1) Effective size and wavelength are comparable
- 2) Particle shape is spherical or quasi-spherical
- 3) The real refractive index is relatively large

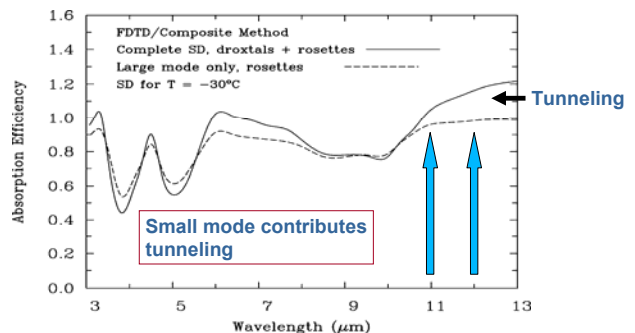
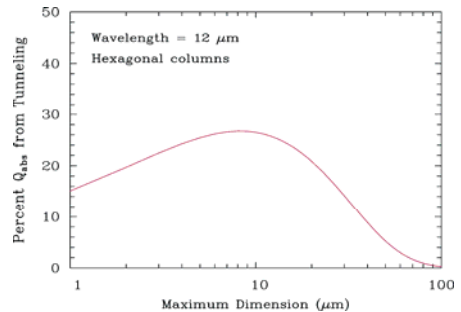
Therefore tunneling contributions at terrestrial wavelengths are greatest for smallest ( $D < 60 \mu\text{m}$ ) ice crystals. To detect the tunneling signal is to detect small ice crystals.

## Tunneling Phenomena



Tunneling depends on real refractive index, which changes abruptly between 11 and 12  $\mu\text{m}$  wavelength.

Tunneling can account for up to 45% of absorption in ice clouds for terrestrial radiation.



## Wavelength dependence of photon tunneling

## Retrieval of Cirrus Temperature and Emissivity

1. For two  $\text{CO}_2$  channels A and B, cirrus emissivity  $\epsilon$  will be the same due to similar refractive indices. Solving for  $\epsilon$  using the simple non-scattering formulation, and equating  $\epsilon_A = \epsilon_B$ ,

$$I_{\text{OBS}} = (1 - \epsilon) I_{\text{CLR}} + \epsilon B(T_{\text{CLD}})$$

and

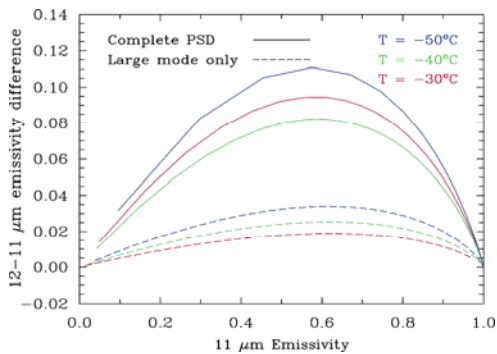
$$\frac{I_{A,\text{OBS}} - I_{A,\text{CLR}}}{B_A(T_{\text{CLD}}) - I_{A,\text{CLR}}} = \frac{I_{B,\text{OBS}} - I_{B,\text{CLR}}}{B_B(T_{\text{CLD}}) - I_{B,\text{CLR}}}$$

We then solve for the one unknown, cirrus temperature  $T_{\text{CLD}}$ .

2. With  $T_{\text{CLD}}$  retrieved and  $I_{\text{CLR}}$  and  $I_{\text{OBS}}$  directly measured, we can solve for  $\epsilon$  at 11 & 12  $\mu\text{m}$ , where  $\epsilon_{11}$  and  $\epsilon_{12}$  are solved independently:

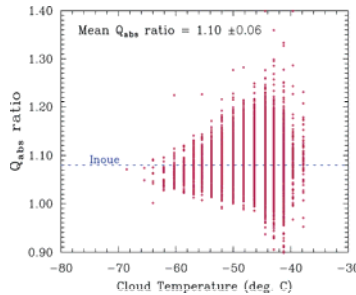
$$\epsilon_{11} = \frac{I_{11,\text{OBS}} - I_{11,\text{CLR}}}{B_{11}(T_{\text{CLD}}) - I_{11,\text{CLR}}} \quad \epsilon_{12} = \frac{I_{12,\text{OBS}} - I_{12,\text{CLR}}}{B_{12}(T_{\text{CLD}}) - I_{12,\text{CLR}}}$$

**Method for Estimating Small Crystal Concentrations**

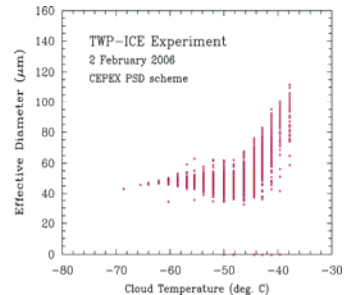


1. Begin with satellite retrievals of cloud temperature and cloud emissivity ( $\epsilon$ ) at 11 and 12 μm wavelength channels.
2. Use retrieved temperature to estimate PSD mean size  $\bar{D}$  and dispersion ( $v$ ) for large and small mode. Difference between the solid and dashed curves above results primarily from differences in contribution of small PSD mode to the ice water content (IWC). This also determines the effective diameter,  $D_e$ . Dispersion  $v$  has little influence on above curves.
3. Locate the retrieved  $\Delta\epsilon$  and the 11 μm  $\epsilon$  by (1) incrementing the modeled ice water path (IWP) to increase  $\epsilon$  (11 μm) and (2) incrementing the small mode contribution to the cloud IWC, which elevates the curve.
4. If retrieved point lies below the "large mode only" curve (e.g. a dashed curve), then systematically decrease  $D$  for large mode until a match is obtained. Negative  $\Delta\epsilon$  values correspond to maximum allowed  $D$  values.
5. This method retrieves IWP,  $D_e$ , the small-to-large mode ice crystal concentration ratio, and the PSD and  $N$  for a given IWC.

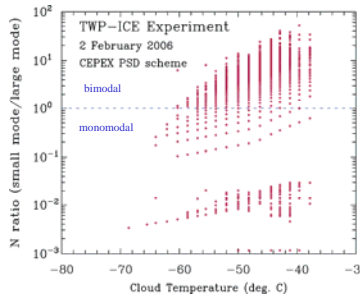
**TWP-ICE Case Study: 2 February 2006 - Continued -**



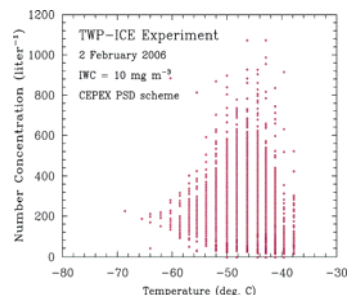
12/11 μm absorption optical depth (AOD) ratios with temperature. Blue-dashed line is mean ratio (1.08) from literature.



Mean AOD ratio of 1.08 produces retrieved  $D_e$  trend since AOD ratio depends on  $D_e$  and wavelength



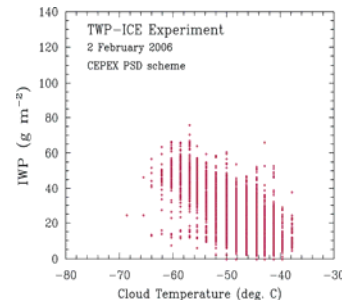
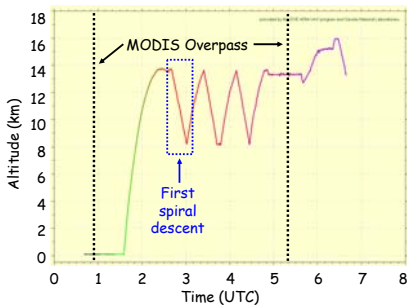
Degree of bimodality in retrieved PSD shown by the small-to-large mode number concentration ratio.



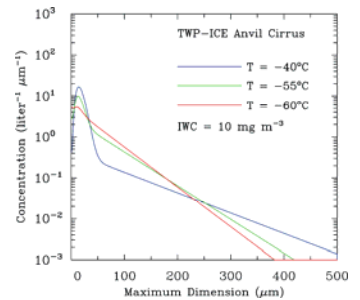
Total number concentrations assuming an IWC of 10 mg m<sup>-3</sup>

**TWP-ICE Case Study: 2 February 2006**

Proteus flight data times, altitudes Thurs 2 Feb 06



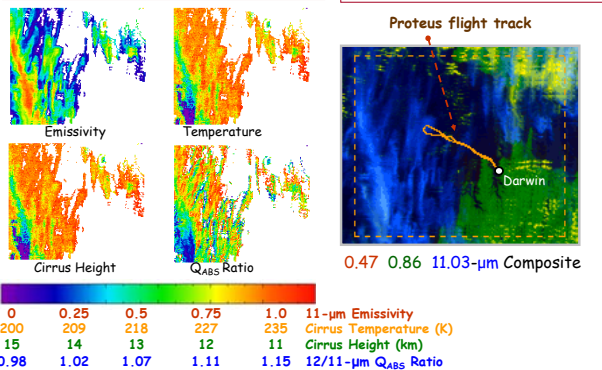
Retrieved IWP, with colder, higher emissivity cirrus having higher IWP



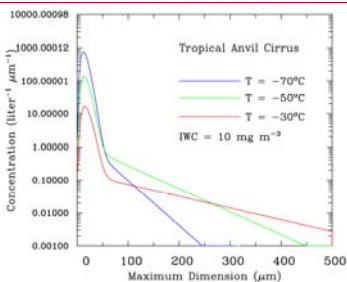
Representative retrieved PSD for the indicated temperatures and assumed IWC.

**Temperature-emissivity related retrievals on MODIS image**

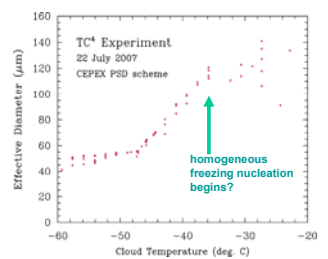
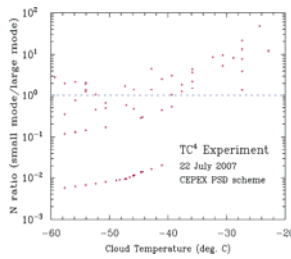
**Sampled region of MODIS image at 0520 UTC**



**PSD scheme based on CEPEX data (anvil cirrus), used in the retrieval algorithm as described above**



**TC4 Case Study: 22 July 2007, Costa Rica**



Retrieval results from TC4 case study and comparisons with in situ measurements involving the new 2DS probe which minimizes ice artifacts from shattering. Note that the  $D_e$ -T relationship is consistent with TWP-ICE.

