



Magnetotelluric Survey to Locate the Archean/Proterozoic Suture Zone North of Wells, Nevada

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Introduction

It is important to know whether major mining districts in the Northern Nevada Gold Province are underlain by rocks of the Archean Wyoming craton, which are known to contain orogenic gold deposits (Hausel and Hull, 1990), or by accreted rocks of the Paleoproterozoic Mojave province (Whitmeyer and Karlstrom, 2004). It is also important to know the location and orientation of the Archean/Proterozoic suture zone between these provinces as well as major basement structures within these terranes because they may influence subsequent patterns of sedimentation, deformation, magmatism, and hydrothermal activity (Crafford and Grauch, 2002; Grauch and others, 2003). The Archean was the main gold-mineralization period, and Archean lode-gold deposits were formed at mid-crustal depths along major shear zones (Cameron, 1988).

The nature of the crystalline basement below the Northern Nevada Gold Province and the location of major faults within it are relevant to Rodinian reconstructions, crustal development, and ore deposit models (e.g., Hofstra and Cline, 2000; Grauch and others, 2003). According to Whitmeyer and Karlstrom (2004), the Archean cratons of the northwestern United States and Canada had stabilized as continental lithosphere by 2.5 Ga, and were rifted and assembled into a large continental mass by 1.8 Ga, to which the 1.73-1.68 Ga Mohave province was accreted by 1.65 Ga. The Archean/Proterozoic suture zone has a west-southwest strike where it is exposed (Reed, 1993) at the eastern Utah and southwestern Wyoming border (Cheyenne Belt) where it is characterized by an up to 7-km-thick mylonite zone (Smithson and Boyd, 1998). In the Great Basin, the strike of the Archean/Proterozoic suture zone is poorly constrained because it is largely concealed below a Neoproterozoic-Paleozoic miogeocline and basin fill. East-west and southwest-northeast strikes for the Archean/Proterozoic suture zone have been inferred based on Sr, Nd, and Pb isotopic compositions of granitoid intrusions (Tosdal and others, 2000). To better constrain the location and strike of the Archean/Proterozoic suture zone below cover, three regional north-south magnetotelluric (MT) sounding profiles were acquired in western Utah and northeastern Nevada (Williams and Rodriguez, 2003; 2004; 2005), and one east-west MT sounding profile (**fig. 1**) MT sounding profile was acquired in northeastern Nevada. Resistivity modeling of the MT data can be used to investigate buried structures or sutures that may have influenced subsequent regional fluid flow and localized mineralization. The purpose of this report is to release the MT sounding data collected along the east-west profile in northeastern Nevada; no interpretation of the data is included.

Electrical Rock Properties

Electromagnetic geophysical methods detect variations in the electrical properties of rocks - in particular, electrical resistivity, or its inverse, electrical conductivity. Electrical resistivity can be correlated with geologic units on the surface and at depth using lithologic logs to provide a three-dimensional (3-D) picture of subsurface geology. In the upper crust the resistivity of geologic units is largely dependent upon their fluid content, pore-volume porosity, interconnected fracture porosity, and conductive mineral content (Keller, 1989). While there is not a one-to-one relationship between lithology and resistivity, there are general correlations that can be made using typical values, even though values can be found at other localities that may fall outside of the ranges

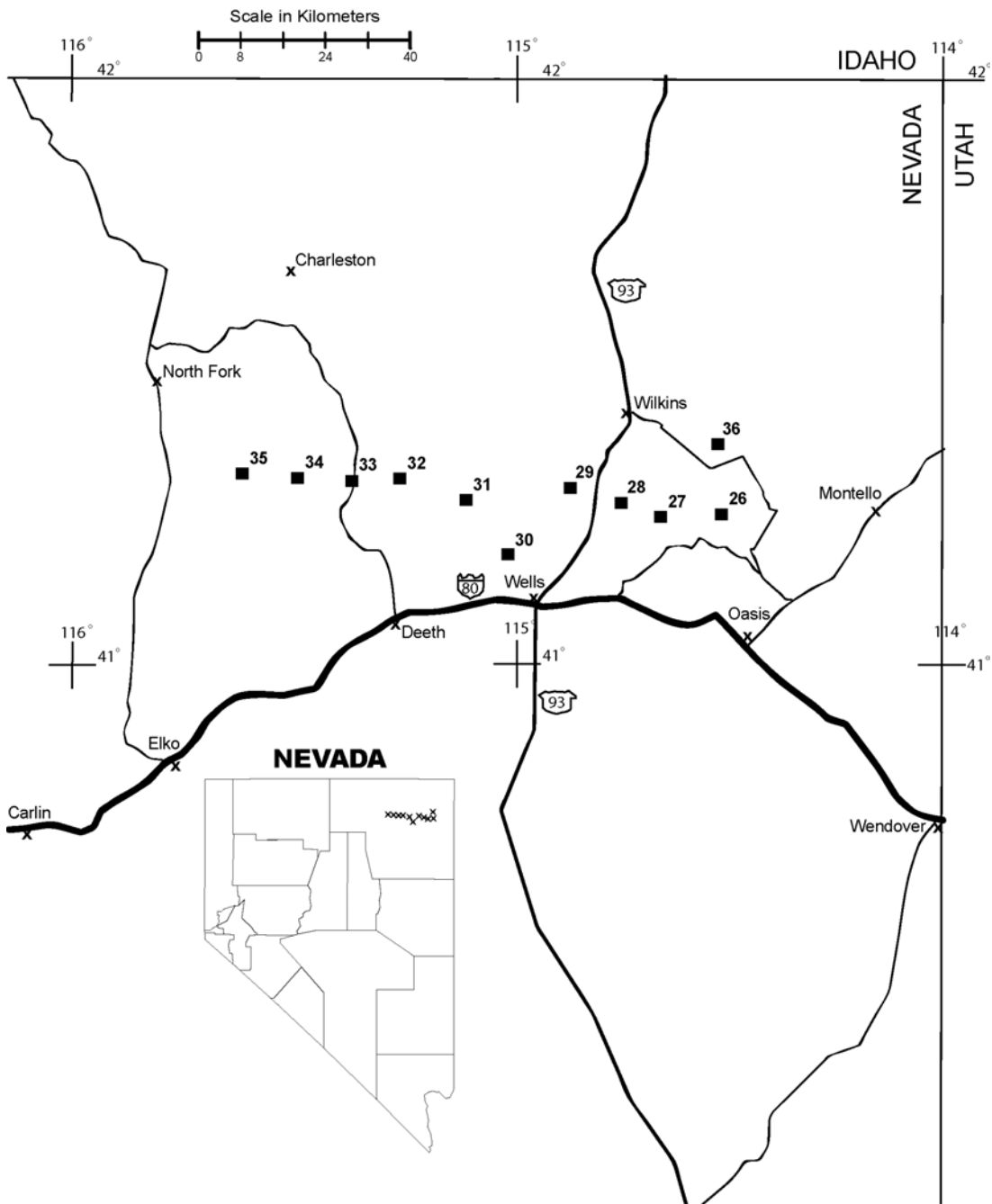


Figure 1. Magnetotelluric (MT) profile index map in north eastern Nevada. MT stations acquired in July 2005 are labeled 26-36. Base map modified from Nevada Bureau of Mines and Geology Topographic Map of Nevada, scale 1:1,000,000 (1972).

presented below (Palacky, 1987). Fluids within the pore spaces and fracture openings, especially if saline, can reduce resistivities in what would otherwise be a resistive rock matrix. Resistivity can also be lowered by the presence of electrically conductive clay minerals, graphitic carbon, and metallic mineralization. It is common, for example, for altered volcanic rocks to contain replacement minerals that have resistivities ten times lower than those of the surrounding rocks (Nelson and Anderson, 1992). Fine-grained sediments, such as clay-rich alluvium, marine shales, and other mudstones, are normally conductive from a few ohm-m to a few tens of ohm-m (Keller, 1987; Palacky, 1987). Metamorphic rocks (non-graphitic) and unaltered, unfractured igneous rocks are normally moderately to highly resistive (a few hundreds to thousands of ohm-m). Carbonate rocks can have similarly high resistivities depending on their fluid content, porosity, and impurities (Keller, 1987; Palacky, 1987). Fault zones may be moderately conductive (tens of ohm-m) when comprised of rocks fractured enough to have hosted fluid transport and consequent mineralogical alteration (Eberhart-Phillips and others, 1995). Higher subsurface temperatures cause higher ionic mobility that reduces rock resistivities (Keller, 1987; Palacky, 1987). Tables of electrical resistivity for a variety of rocks, minerals and geological environments may be found in Keller (1987) and Palacky (1987).

Magnetotelluric Method

The MT method is a passive surface electromagnetic geophysical technique that measures variations in the Earth's natural electromagnetic fields to investigate the electrical resistivity structure of the subsurface from depths of tens of meters to tens of kilometers (Vozoff, 1991). Worldwide lightning activity at frequencies of 10,000 to 1 Hertz (Hz) and geomagnetic micro-pulsations at frequencies of 1 to 0.001 Hz provide the majority of natural signal used by the MT method. The natural electromagnetic wave propagates vertically in the earth due to the large resistivity contrast between the air and the earth causing a vertical refraction of the electromagnetic wave transmitted into the Earth (Vozoff, 1972).

The natural electric and magnetic fields are recorded in two orthogonal, horizontal directions at the Earth's surface (the vertical magnetic field is also recorded). The resulting time-series signals are used to derive tensor apparent resistivities and phases by first converting them to complex cross-spectra using Fourier transform techniques. Least squares, cross-spectral analysis (Bendat and Piersol, 1971) are used to solve for a tensor transfer function. Prior to conversion to apparent resistivity and phase, the tensor is normally rotated into principal directions that usually correspond to the direction of maximum and minimum apparent resistivity. For a two-dimensional (2-D) Earth, in which Earth's resistivity structure varies with depth and in one lateral direction, the MT fields can be decoupled into transverse-electric (TE) and transverse-magnetic (TM) modes; 2-D resistivity modeling is generally computed to fit both modes. When the geology satisfies the 2-D assumption, the MT data for the TE mode represents electric fields oriented parallel to geologic strike, and the data for the TM mode represents electric fields oriented perpendicular to strike. The MT method is well suited for studying complicated geological environments because the electric and magnetic fields are sensitive to vertical and horizontal variations in resistivity. The method is capable of establishing whether the electromagnetic fields are responding to subsurface rock bodies

of effectively 1, 2, or 3 dimensions. An introduction to the MT method and references for a more advanced understanding are contained in Dobrin and Savit (1988) and Vozoff (1991).

Magnetotelluric Survey

Eleven MT soundings were collected in July 2005 along a 90-km-long profile in northeastern Nevada (**fig. 1**) starting just west of Toano Well No. 2, northeast of Wells, Nevada, near the Pequop Siding railroad stop, then continuing westerly and ending just northeast of Contact Ranch. The profile location was selected to cross the Archean/Proterozoic suture zone on the basis of hypothesized projections from MT data to the southeast (Williams and Rodriguez, 2004) and regional interpretations of geologic mapping (Reed, 1993).

Station locations were chosen for proximity to roads and to avoid electrical noise such as power lines. All data at the stations were collected with a portable Electromagnetic Instruments, Inc. (EMI), MT-1 system (EMI, Inc., 1996). Horizontal electric fields were recorded using copper sulfate porous pots placed in an L-shaped, three-electrode array with dipole lengths of 30 meters (m). The orthogonal, horizontal magnetic fields in the direction of the electric-field measurement array were sensed using high magnetic permeability, mu-metal-cored induction coils. Frequencies were sampled from 0.009 to 70Hz, using single-station recordings of the orthogonal, horizontal components of the electric and magnetic fields and the vertical magnetic field.

The following table lists the 11 MT station locations. Remote reference stations were not used because stations were located remotely from man-made noise sources. Coordinates are referenced to the 1866 Clarke spheroid and North American 1927 Western United States datum. Longitude and latitude format is degrees:minutes:seconds. Universal Transverse Mercator (UTM) units and station elevations are in meters. The accuracy of the north and east component is ± 5 m and is ± 10 m for the elevation. X direction is in degrees clockwise from True North.

Table 1

Station	X Dir.	Longitude	Latitude	North(m)	East(m)	Elevation
35	270	-115:37:03	41:19:31	4,575,574	11,615,704	1724
34	180	-115:29:34	41:19:06	4,574,981	11,626,160	1908
33	90	-115:22:15	41:18:49	4,574,636	11,636,356	2094
32	0	-115:15:46	41:19:01	4,575,175	11,645,417	1735
31	90	-115:06:47	41:16:49	4,571,384	11,658,025	1739
30	180	-115:01:17	41:11:16	4,561,283	11,665,936	1716
29	180	-114:52:53	41:18:06	4,574,203	11,677,371	1866
28	90	-114:46:02	41:16:33	4,571,553	11,687,006	1962
27	270	-114:40:42	41:15:09	4,569,181	11,694,513	1899
26	180	-114:32:28	41:15:23	4,570,132	11,705,926	1783
36	270	-114:32:55	41:22:33	4,583,177	11,704,999	1709

Magnetotelluric Data

The recorded time-series data were converted to the frequency domain and processed to determine both the impedance tensor, which is used to derive apparent

resistivities, and phase tensor at each site. Rotation of the impedance tensor to maximum and minimum directions allowed for decoupling into the TE and TM modes.

Although remote reference techniques were not used in the study, cross-power files were sorted to select optimal signal-to-noise time-series data sets (see Appendix 1).

The effects of near-surface resistivity anomalies can give rise to “static shifts” in the data (Sternberg and others, 1988) that can be modeled with two-dimensional forward modeling in the near surface. Static shifts were significant at stations **33**, **29**, **28**, **27**, and **36**. Cultural features can affect the response of the MT system. Fences, pipelines, communication lines, moving vehicles and trains, and other manmade sources of electromagnetic noise can contaminate the responses.

The figures in Appendix 1 represent the field-processed MT data for each station after the time-series data were converted to the frequency domain, and after the tensor-transfer function was rotated into principal directions as described in the “Magnetotelluric Method” section of this report.

For each station, nine separate plots are given:

1. Apparent Resistivity (x and o symbols are xy and yx components)
2. Impedance Phase (x and o symbols are xy and yx components)
3. Impedance Skew
4. Multiple Coherency (x and o symbols are xy and yx components)
5. Impedance Polar Plots
6. Tipper Magnitude
7. Tipper Strike
8. HzHx (x symbol) and HzHy (o symbol) Coherency

Error bars (],]) on the Apparent Resistivity, Impedance Phase, Skew, Tipper Magnitude, and Tipper Strike plots represent probable errors within one standard deviation of the sample variance (Gamble and others, 1979).

Apparent resistivity is the ratio of the electric field strength magnitude over the magnetic field strength magnitude for a given frequency. The impedance phase is proportional to the slope of the apparent-resistivity curve on a log-log plot, relative to a baseline at -45 degrees (Vozoff, 1991). A measure of the dimensionality for MT data is provided by the impedance skew of the impedance tensor (Vozoff, 1972). If the effective, measured resistivity response to the geology beneath an MT station is truly one- or two-dimensional, then the skew will be zero. Both instrument and environmental sources of noise contribute to non-zero skew values but are typically small (about 0.1) for relatively low-noise-level recordings. Higher skews (more than 0.2) indicate either the resistivity response to 3-D geology or higher levels of noise.

In the study area, noise from a number of small power lines and small moving vehicles was negligible at distances of 0.4 km and farther from the noise source. Power-line amplitude levels were measured at each site and were typically less than 20 percent of the maximum recordable signals. Noise from larger power lines, power generators, pipelines, and trains was negligible at distances more than 5 km. Local lightning, wind, and rainstorms also can degrade data quality, but these were avoided by not recording during active thunderstorm periods. Burying the magnetic induction coils and keeping the electric dipole wires flat on the ground helped to minimize wind noise.

Predicted values of the electric field can be computed from the measured values of the magnetic field (Vozoff, 1991). The coherence of the predicted electric field with the measured electric field is a measure of the signal-to-noise ratio provided in the multiple coherency plots. Values are normalized between 0 and 1, where values at 0.5 signify signal levels equal to noise levels. For this data set, coherencies generally were at an acceptable level, except at times in the frequency “dead band” (0.01 to 5 Hz) (Dobrin and Savit, 1988).

The figures in Appendix 1 represent the field-processed MT data at each station, and include some data scatter and poor signal-to-noise ratios. The only effort aimed at removing noisy data points was to visually inspect and select the best signal-to-noise field data to combine into the final data plots.

The impedance polar plots provide a measure of the MT data dimensionality (Reddy and others, 1977). For 1-D resistivity structures, the principal impedance (off-diagonal elements) polar diagram (dashed line) is a circle. For 2-D or 3-D resistivity structures, the principal impedance polar diagram (dashed line) elongates either parallel to or perpendicular to strike direction. Over resistors, the principal impedance polar diagram elongates perpendicular to strike direction, while over conductors, the principal impedance polar diagram elongates parallel to strike direction. For 2-D resistivity structures, the additional impedance polar diagram (solid line) attains the shape of a symmetric clover leaf. For 3-D resistivity structures, the additional impedance polar diagram (solid line) elongates in one direction, and its amplitude is comparable to that of the principal impedance polar diagram (dashed line), although high noise levels can produce the same effect on the polar diagram. A 3-D analysis of polar plots at each frequency should also take into account the corresponding coherence and skew values along with their associated error levels. The polar plots computed for our data show station **35** was moderately 3-D below 0.05 Hz. Stations **33** and **32**, were weakly 3-D below 0.05 Hz. Station **31** was moderately 3-D below 0.1 Hz. Station **29** was weakly 3-D below 0.2 Hz. Because of noise, station **27** is difficult to interpret but is possibly 3-D below 1 Hz. Station **36** was moderately 3-D below 0.2 Hz. Station **28** was 3-D over all frequencies.

The tipper can be calculated from the vertical component of the magnetic field. The tipper magnitude is a measure of the tipping of the magnetic field out of the horizontal plane (Vozoff, 1991). The magnitude is zero for the 1-D case, typically increases between 0.1 to 0.5, and rarely is as great as 1 as it responds to vertical and subvertical structures. The tipper strike typically is used to help resolve the 90-degree ambiguity in the impedance rotation angle. The tipper magnitude of these stations typically ranged between 0.1 and 0.6 over the lower frequencies, indicating some vertical structure at depth. The HzHx and HzHy coherency is a measure of the signal-to-noise ratio of the vertical magnetic field with respect to each of the orthogonal, horizontal magnetic field directions. Values are normalized between 0 and 1, where values of 0.5 signify signal levels equal to noise levels. These three components of magnetic-field coherence provide a check on the quality of the measured values in the tipper magnitude and tipper strike plots.

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Appendix 1 Magnetotelluric Data Plots

There are nine separate plots for each station:

1. Apparent Resistivity for the unrotated xy (x symbol) and yx (o symbol) modes
2. Impedance Phase for the unrotated xy (x symbol) and yx (o symbol) modes
3. Impedance Skew for the impedance tensor
4. Multiple Coherency for the xy (x symbol) and minimum (o symbol) modes of the electric field
5. Impedance Polar Plots (at 12 selected frequencies)
6. Tipper Magnitude for the vertical magnetic field
7. Tipper Strike for the vertical magnetic field
8. HzHx (x symbol) and HzHy (o symbol) Coherency

Refer to the “Magnetotelluric Data” section in this report for an explanation of these plots.