Prepared in Cooperation with the NATIONAL PARK SERVICE

Chemical Characteristics of Ground-Water Discharge along the South Rim of Grand Canyon in Grand Canyon National Park, Arizona, 2000–2001

Scientific Investigations Report 2004–5146

U.S. Department of the Interior U.S. Geological Survey





Hermit Creek near Hawaii Spring

Cover photograph: View of Palisades of the Desert from near Desert View, Grand Canyon, Arizona. Photograph by Colin McKay.

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*By* Stephen A. Monroe, Ronald C. Antweiler, Robert J. Hart, Howard E. Taylor, Margot Truini, John R. Rihs, *and* Tracey J. Felger

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# Contents

Abstract
Introduction
Purpose and scope 2
Water issues for Grand Canyon National Park.    2
Description of the south rim of Grand Canyon         4
Climate
Hydrogeologic characteristics 6
Previous investigations
Acknowledgments
Investigation methods
Field logistics
Sample collection and processing 8
Laboratory methods
Elemental and constituent analyses of water samples
Isotopic composition analyses of water and rock samples
X-ray diffraction analyses of rock samples
Isotopic analyses
Site descriptions
Chemical characteristics of ground-water discharge at springs
Comparison of water-chemistry data from springs with data from creeks
Consistency of the springs data
Relations to water-quality contaminant standards
Geochemistry of barium and sulfate 22
Chemistry of selected springs
Monument Spring
Lonetree Spring
Salt Creek Spring and Horn Creek
Grapevine Main, Pumphouse, Hawaii, Hermit, and Boucher East Springs
Residence times of ground water discharging at springs
Future studies
Summary and conclusions
References cited
Physical and chemical data
Appendix

## PLATE

Map showing surficial geology, geologic structure, and chemical characteristics of ground-water discharge at selected springs and creeks along the south rim of Grand Canyon in Grand Canyon National Park, Arizona, 2000–2001

In pocket at back of report.

# Figures

1. 2.	Map showing study area and south rim of Grand Canyon, Arizona Stratigraphic section of the south rim of Grand Canyon, Arizona	
3–4.	Photographs showing:	
3.	Water-sample collection and processing.A.Water-sample collection at Miners Spring.B.Water-sample processing at Monument Spring .	
4.	<ul> <li>Photographs showing spring types at sample sites.</li> <li>A. Bedding planes</li> <li>B. Fractures</li> <li>C. Hillslopes</li> <li>D. Channel alluvium</li> <li>E. Travertine deposits</li> </ul>	14 14 14
5–15.	Graphs showing:	
5.	Relative compositions of water from springs and creeks that discharge along the south rim of Grand Canyon, Arizona, 2000–2001	19
6.	Distribution of constituent concentrations of water that discharges from springs and creeks along the south rim of Grand Canyon, Arizona, 2000–2001.	
	<ul> <li>A. Major constituents: calcium, magnesium, sodium, bicarbonate, chloride, and sulfate</li> <li>B. Trace elements (higher concentrations): arsenic, molybdenum,</li> </ul>	20
	<ul> <li>rubidium, selenium, uranium, and zinc</li> <li>C. Trace elements (lower concentrations): neodymium, lead, rhenium, antimony, thallium, and yttrium</li> </ul>	
7.	Concentrations of arsenic in relation to spring or creek locations	20
8.	along the south rim of Grand Canyon, Arizona, 2000–2001 Concentrations of selenium in relation to spring or creek locations	24
9.	along the south rim of Grand Canyon, Arizona, 2000–2001 Concentrations of uranium in relation to spring or creek locations	
10.	along the south rim of Grand Canyon, Arizona, 2000–2001 Relation between concentrations of barium and sulfate in water that	
11.	discharges along the south rim of Grand Canyon, Arizona, 2000–2001 Concentrations of barium in relation to spring or creek locations	
12.	along the south rim of Grand Canyon, Arizona, 2000–2001 Relation between barium and calcium concentrations, demonstrating the clustering of data, for selected springs and creeks that discharge along the south rim of Grand Canyon, Arizona, 2000–2001	
13.	Relation between oxygen-18 ( $\delta^{18}$ O) and deuterium ( $\delta$ D) in water along the south rim of Grand Canyon, Arizona.	
	<ul> <li>A. Precipitation 1989–2002</li> <li>B. Water from springs and creeks, 2000–2001</li> <li>C. Regional trend of oxygen-18 (δ<sup>18</sup>0) values for water from springs and creeks, 2000–2001</li> </ul>	28
14.	Strontium-87/strontium-86 in water from springs and creeks that discharge along the south rim of Grand Canyon, Arizona, and in rocks from near	00
15.	the Bright Angel Fault near the south rim, 2000–2001 Relation between tritium and carbon-14 in water from springs and creeks that discharge along the south rim of Grand Canyon, Arizona, 2000–2001	

iv

# Tables

1.	Locations of springs and creeks sampled in study 3
2.	Spring and creek sample protocols
3.	Chemical and isotopic analytes and corresponding median
	detection limits 10
4.	Geologic characteristics of springs and creeks sampled in study 15
5.	Changes in constituent concentrations from the point of spring
	issuance sample location to the downstream sample location,
	Cottonwood and Monument Creeks, 2000–2001 21
6.	Variability in chemical data for selected constituents for all
	springs sampled more than once during 2000–2001
7.	U.S. Environmental Protection Agency Maximum Contaminant
	Levels (MCLs) for selected constituents
8.	Physical and chemical properties, chemical-constituent
	concentrations, radioactivity, and isotopic data for water
	from springs and creeks that discharge along the south
	rim of Grand Canyon, Arizona, 2000–2001
9.	X-ray diffraction mineralogy and isotope data for selected
	rock samples representing the major stratigraphic units
	present near the Bright Angel Fault in Grand Canyon, Arizona

Multiply	Ву	To obtain
centimeter (cm)	0.3937	inch
millimeter (mm)	0.03937	inch
meter (m)	3.281	foot
kilometer (km)	0.6214	mile
liter (L)	0.2642	gallon
cubic meter per second (m <sup>3</sup> /s)	35.31	cubic foot per second
liter per minute (L/min)	0.2641721	gallon per minute
cubic meter per day (m <sup>3</sup> /d)	264.2	gallon per day
becquerel per liter (Bq/L)	27.027	picocurie per liter

# **Conversion Factors and Datum**

Temperature in degrees Celsius (°C) may be converted to degrees Fahrenheit (°F) as follows: °F =  $(1.8^{\circ}C)+32$ 

#### ABBREVIATED WATER-QUALITY UNITS

Chemical concentration and water temperature are given only in metric units. Chemical concentration in water is given in milligrams per liter (mg/L) or micrograms per liter ( $\mu$ g/L). Milligrams per liter is a unit expressing the solute mass (milligrams) per unit volume (liter) of water. One thousand micrograms per liter is equivalent to 1 milligram per liter. For concentrations less than 7,000 milligrams per liter, the numerical value is about the same as for concentrations in parts per million. Specific conductance is given in microsiemens per centimeter at 25 degrees Celsius ( $\mu$ S/cm at 25 °C). Radioactivity is expressed in picocuries per liter (pCi/L) or picocuries per gram (pCi/g), which is the amount of radioactive decay producing 2.2 disintegrations per minute in a unit volume (liter) of water or mass (gram) of sediment.

#### DATUMS

Vertical coordinate information is referenced to the National Geodetic Vertical Datum of 1929 (NGVD 29)—a geodetic datum derived from a general adjustment of the first-order level nets of both the United States and Canada, formerly called Sea Level Datum of 1929; horizontal coordinate information is referenced to the North American Datum of 1927 (NAD 27). Altitude, as used in this report, refers to distance above or below NGVD 29.

# Chemical Characteristics of Ground-Water Discharge along the South Rim of Grand Canyon in Grand Canyon National Park, Arizona, 2000–2001

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# Abstract

Springs flowing from the south rim of Grand Canyon are an important resource of Grand Canyon National Park, offering refuge to endemic and exotic terrestrial wildlife species and maintaining riparian areas. Population growth on the Coconino Plateau has increased the demand for additional development of ground-water resources, and such development could reduce spring discharge and affect the sustainability of riparian areas within the park. In addition, springs are an important source of drinking water for hikers and are culturally and economically important to Native Americans living in the region.

Water samples were collected from May 2000 to September 2001 from 20 spring and creek sites that discharge water from the Redwall-Muav Limestone aquifer along the south rim of Grand Canyon. Sample collection sites were described and samples were analyzed for major ions, nutrients, trace elements, radioactivity, and selected isotopes, and potential sources of ground-water flow to the springs. Rock samples representing the major stratigraphic units of Grand Canyon were collected near the Bright Angel Fault and analyzed for mineralogy, strontium-87/strontium-86, and carbon-13/carbon-12.

The chemical composition of water samples collected from a given spring did not vary appreciably over the course of the study. Although water at each spring had a temporally constant composition, the composition was chemically distinct from that of every other spring sampled, indicating spatial variability in the groundwater composition. Most samples had a calcium magnesium bicarbonate composition; a few had a substantial sulfate component. Concentrations of arsenic, nitrate, selenium, uranium, and gross alpha approached or exceeded U.S. Environmental Protection Agency Maximum Contaminant Levels in water discharging from some springs. Oxygen and hydrogen isotopic compositions varied little among samples, and for most sites the isotopic data plot close to the global meteoric water line or below the local meteoric water line. Isotopic enrichment indicates fractionation due to evaporation occurs at some sites. The evaporative process may occur prior to recharge or post-discharge. Flow paths are differentiated between the eastern part of the study area where strontium-87/strontium-86 values for water from springs and creeks are more radiogenic than strontium-87/strontium-86 values for water that discharges from sites farther west. Tritium and carbon isotope analyses indicate that residence time of ground-water discharge from springs and creeks ranges from less than 50 years to about 3,400 years. Water with a residence time of less than 50 years is absent at several sites. Discharge of most springs and creeks is a mixture of younger and older waters.

# INTRODUCTION

Spring flow from sedimentary rocks along the south rim of Grand Canyon is an important resource of Grand Canyon National Park (GRCA) and adjoining areas (**fig. 1** and **pl. 1**). Springs offer refuge to endemic and exotic terrestrial wildlife species of Grand Canyon and maintain associated riparian areas. Recent development and ground-water use near the south rim have made environmentalists, commercial developers, municipalities, and State and Federal resource managers increasingly aware of spring resources in the area. Native Americans who inhabit the Grand Canyon region, such as the Havasupai Tribe, are concerned about the effects of ground-water development on spring resources. Springs are economically important and culturally significant to Native Americans.

At the request of the GRCA, the U.S. Geological Survey (USGS) and GRCA began a cooperative study in 1999 to assess spring water quality and determine the sources of ground water discharged at the springs. This study was done under the auspices of the National Park Service (NPS)/USGS Water-Quality Assessment and Monitoring Program (WQAM).

Data collected for this study will provide GRCA resource managers an inventory of spring resources along the south rim of Grand Canyon. The springs result from ground-water discharge from aquifers that are being increasingly developed. As such, the springs are susceptible to increasing ground-water withdrawals on the Coconino Plateau.

### **Purpose and Scope**

The purpose of this report is to present the results of a study of the chemical characteristics of ground-water discharge along the south rim of Grand Canyon in Grand Canyon National Park. The specific objectives of this study were to determine baseline water quality and the residence times and flow paths of water discharging from the Redwall-Muav Limestone aquifer along the south rim of Grand Canyon. Data included in this report were collected from May 2000 to September 2001.

This report describes the aquifers in the study area, the spring and creek sites selected for the study, the variation in water chemistry at the sites, the trends in chemical composition of water among the sites, and the sources and residence times for ground water that discharges at the springs.

Physical, chemical, and water-quality data are presented for 20 sites in the section of Grand Canyon from Red Canyon in the east to Boucher East Spring in the west at which ground water discharges from the Redwall-Muav Limestone aquifer (pl. 1). Sites included are in the core area of GRCA, close to population centers and ground-water withdrawal sites. The 20 sites (table 1) include 14 springs, 5 creeks, and 1 USGS streamflow-gaging station (station number 09403013). The streamflow-gaging station monitors flow from Pumphouse Spring. Two of the creek sites (Cottonwood No. 1 and Cottonwood No. 2) are in the same drainage, and one of the creek sites (Monument Creek No. 1) is in the same drainage as one of the spring sites (Monument Spring).

Data collected include physical descriptions, air and water temperature, pH, specific conductance, dissolvedoxygen concentration, field alkalinity, location, and altitude. Water samples were analyzed for concentrations of major ions, nutrients, and trace elements; for radioactivity; and for stable and radioactive isotope composition. Rock samples representing the major stratigraphic units of Grand Canyon near the Bright Angel Fault (pl. 1) were collected and analyzed for mineralogy and isotope composition.

#### Water Issues for Grand Canyon National Park

Grand Canyon National Park is recognized as a place of global importance, and the park contains surface and ground waters and associated habitats that are rare in arid environments. Water availability has been the greatest limitation to development in the south rim area, and historically, water has been expensive to acquire and water supplies have been the subject of contention (U.S. Department of Agriculture, Forest Service, 1999).

One of the highest priorities identified in the park's General and Resource Management Plans is to accurately document locations and water chemistry of springs issuing from within the park boundary. Addressing this issue would benefit GRCA by (1) providing documented water-resource information that can be used to develop a requisite resource management policy; (2) contributing to the understanding of ground-water movement and direction, and estimation of time since recharge of spring water; (3) identifying the presence or absence of effluent and landfill leakage to spring sources; (4) evaluating field logistics for potential long-term monitoring of south rim springs; (5) providing the general public with interpretive information regarding the spring resources of GRCA; (6) identifying future research and monitoring needs; and (7) providing supporting data for concurrent studies including Wild and Scenic River suitability studies, spring-flow monitoring, and biological studies being conducted on Colorado River tributaries in GRCA.

Table I. Locations of springs and creeks sampled in study	Table 1.	Locations of springs and creeks sampled in study
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Site name		sverse Mercator, ters, NAD 27)	Altitude (meters	Rock unit from which water	Dates visited	Measurements or samples
	Easting	Northing	- above NGVD 29)	discharges		collected <sup>1</sup>
Red Canyon Spring	415783.100	3984788.622	1,335	Muav Limestone	8-23-01	А
					9-26-01	A,B,C,D,E,F,G
JT Spring	414289.143	3984642.662	1,329	Muav Limestone	4-8-01	A,B,D,E
					5-11-01	A,B,C,D,E,F,G
Miners Spring	412462.374	3986009.871	1,323	Muav Limestone	5-24-00	A,B,C,D,E,F,G
					11-28-00	A,B,D
					4-7-01	A,B,D,E
Cottonwood Creek No. 1	411097.828	3986905.326	1,186	Bright Angel Shale	5-25-00	A,B,C,D,E,F,G
Cottonwood Creek No. 2	410753.643	3986294.103	1,277	Bright Angel Shale	11-29-00	A,B,C,D,E,F,G
					4-9-01	A,B,D,E
					4-30-01	A,B,C,D,E,F,G
Grapevine Main Spring	409651.041	3985440.995	1,433	Muav Limestone	4-10-01	A,B,D,E
					4-30-01	A,B,C,D,E,F,G
Grapevine East Spring	408867.726	3988909.541	1,113	Bright Angel Shale	5-25-00	A,B,C,D,E,F,G
					11-29-00	A,B
					12-12-00	A,D,F
					4-9-01	A,B,D,E
Lonetree Spring	405689.132	3991639.143	1,305	Muav Limestone	4-11-01	A,B,D,E
					5-1-01	A,B,C,D,E,F,G
Sam Magee Spring	404334.268	3992864.335	1,213	Muav Limestone	4-20-01	A,D,E
Burro Spring	400927.325	3992807.522	1,131	Bright Angel Shale	5-22-00	A,B,C,D,E,F,G
					12-7-00	A,B,D
					4-8-01	A,B,D,E
Pipe Creek	401015.850	3991996.032	1,157	Bright Angel Shale	5-22-00	A,B,C,D,E,F,G
					12-7-00	A,B,D
					4-8-01	A,B,D,E
Pumphouse Spring	398666.261	3992931.487	1,195	Bright Angel Shale	5-22-00	A,B,C,D,E,F,G
					12-7-00	A,B,D
					4-7-01	A,B,D,E
Pumphouse Wash Gage <sup>2</sup>	398645.327	3993014.269	1,160	Bright Angel Shale	12-7-00	A,B
					12-9-00	A,D
Horn Creek	397088.983	3993275.333	1,132	Bright Angel Shale	5-22-00	A,B,C,D,E,F,G
					12-6-00	A,B,D
					4-7-01	A,B,D,E
Salt Creek Spring	395457.355	3992895.375	1,207	Muav Limestone	5-23-00	A,B,C,D,E,F,G
					12-6-00	A,B,D
					4-10-01	A,B,D,E

See footnotes at end of table.

Site name	Universal Transverse Mercator, Zone 12 (meters, NAD 27)		Altitude (meters	Rock unit from which water	Dates visited	Measurements or samples	
	Easting	Northing	above NGVD 29)	discharges		collected <sup>1</sup>	
Monument Creek No. 1	393260.779	3993479.813	914	Tapeats Sandstone	5-24-00	A,B	
Monument Spring	394125.431	3991663.212	1,317	Temple Butte Formation	12-5-00	A,B,C,D,E,F,G	
					4-9-01	A,B,D,E	
Hawaii Spring	390290.367	3992268.383	1,012	Muav Limestone	5-25-00	A,B,C,D,E,F,G	
					12-4-00	A,B,D	
					4-11-01	A,B,D,E	
Hermit Spring	389685.678	3991448.930	1,164	Temple Butte Formation	12-4-00	A,B,D	
					4-11-01	A,B,D	
Boucher East Spring	388629.276	3995826.366	948	Bright Angel Shale	5-26-00	A,B,C,D,E,F,G	
					12-4-00	A,B	
					4-12-01	A,B,D,E	

Table 1.	Locations of springs and creeks sampled in study—Continued
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<sup>1</sup>Measurements or samples collected during site visits: A, field parameters; B, chemical constituents; C, radiochemistry; D, oxygen and hydrogen isotopes; E, tritium; F, carbon isotopes; G, strontium isotopes.

<sup>2</sup>Pumphouse Wash gage is located in close proximity to Pumphouse Spring and monitors total spring discharge from the vicinity.

How existing and proposed ground-water development outside the park will affect the spring resources within the park is unknown. Ground-water flow paths in aquifers of the Coconino Plateau are poorly understood. In addition, little is known about the current and potential effects of human activities, such as wastewater treatment plants, landfills, and historic mining, on the ground-water systems.

#### **Description of the South Rim of Grand Canyon**

The south rim of Grand Canyon is a steep escarpment extending from the Coconino Plateau to the Colorado River. Vertical relief of the south rim is about 1,500 m. The Coconino Plateau is a physiographic feature of the Colorado Plateau that encompasses an area of about 13,000 km<sup>2</sup> in northern Arizona (fig. 1). This study is focused on the part of Grand Canyon National Park centered on the Bright Angel Fault and extending as far east as Red Canyon and as far west as Boucher East Spring (pl. 1). The south rim is characterized by numerous short, steep side canyons and differs from the north rim of Grand Canyon in that it has smaller tributary canyons and no extensively developed karst systems (Huntoon, 2000). Side canyon development is primarily controlled by regional dip of sedimentary rocks and folds and faults (Potochnick and Reynolds, 1990). The regional dip of south rim Paleozoic strata is south, away from Grand Canyon.

Grand Canyon Village, Tusayan, and Valle are small communities in and near Grand Canyon National Park near the south rim (fig. 1). No wells exist in the park; however, deep wells have been developed in Tusayan and Valle. Wastewatertreatment plants and landfills are operated in Grand Canyon Village and Tusayan. Two abandoned mines are in the study area (pl. 1). The Last Chance mine was developed near a breccia pipe at Horseshoe Mesa and produced copper during the early 1900s (Kenny, 2001). The Orphan mine was developed in a breccia pipe in the Horn Creek drainage and produced copper in the early 1900s and uranium during the 1950s and 1960s (Amundson, 2001).

#### Climate

Climatic conditions of the Coconino Plateau vary in relation to altitude. Average precipitation varies from about 48 cm/yr at Grand Canyon Village on the south rim to about 23 cm/yr in the canyon at Phantom Ranch (Sellers and others, 1986). In the canyon, the climate is continental and arid. The greatest rainfall occurs during the warm season (June through October) as a result of the Southwestern monsoon. Snowfall on the rim of the canyon accounts for the greatest part of the total precipitation in the area. Along the canyon rim, vegetation is predominantly ponderosa pine. Altitude decreases to the south of Grand Canyon, and there is a corresponding decrease in precipitation and a transition to desert scrub vegetation. Deep in the canyon, where precipitation is also diminished, desert vegetation predominates.

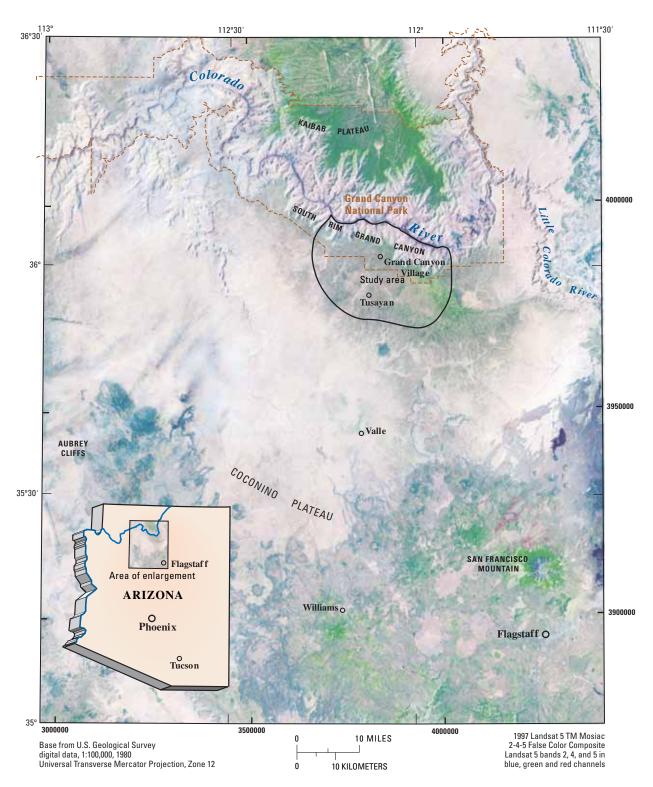


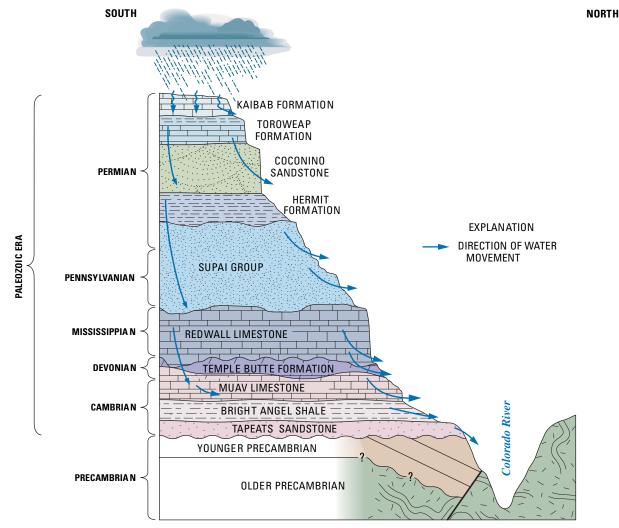
Figure 1. Study area and south rim of Grand Canyon, Arizona.

## Hydrogeologic Characteristics

The primary water-bearing units of the Coconino Plateau near Grand Canyon are the lower Paleozoic carbonate rocks (Metzger, 1961). These rock units constitute the Redwall-Muav Limestone aquifer and are generally found 500 m to 700 m below the land surface of the plateau. Included in the aquifer are the Cambrian Bright Angel Shale, Cambrian Muav Limestone, Devonian Temple Butte Formation, and Mississippian Redwall Limestone (Metzger, 1961; Cooley and others, 1969). The Bright Angel Shale and Muav Limestone intertongue in the study area and thicken towards the west. The Temple Butte Formation is a discontinuous dolomitic limestone that thickens (30 m to 100 m) from east to west in Grand Canyon (Beus, 1990). The Redwall Limestone is a mostly dolomitic limestone that is continuous and massive throughout the study area. In the study area, the aquifer is bounded on the north by the Colorado River, on the west by structural controls near the Aubrey Cliffs, on the east by a

generally defined ground-water divide near the Little Colorado River (E.L. Montgomery and Associates, 1997), and on the south by a generally defined ground-water divide near Williams (Bills and Flynn, 2002).

Perched water-bearing zones also occur near Grand Canyon. Ground water discharges from these zones at springs and seeps in the Pennsylvanian and Permian Supai Group, the Permian Coconino Sandstone, and the Permian Kaibab Formation (Metzger, 1961). The Supai Group is a series of limestone, mudstone, and sandstone rock units and tends to be highly fractured. The Permian Hermit Formation, overlying the Supai Group, is a siliciclastic siltstone, mudstone, and sandstone (Blakey, 1990). The Coconino Sandstone is a homogeneous silicate sandstone. The Kaibab Limestone is a resistant rock unit that forms a cap throughout most of the Grand Canyon region (fig. 2). Flow at the springs and seeps is generally small and discontinuous, and wells developed in these zones typically have low production rates (McGavock and others, 1986).



Modified from Billingsley and others, 1980

Figure 2. Stratigraphic section of the south rim of Grand Canyon, Arizona.

Water-bearing units near the south rim of Grand Canyon are recharged primarily by precipitation that infiltrates volcanic rocks and the Kaibab Formation in the higher altitudes of the Coconino Plateau (Johnson and Sanderson, 1968). Faults and fractures, many widened by dissolution, create zones of secondary permeability that could be the primary controls of ground-water movement. Ground water in the northern parts of the Coconino Plateau generally flows toward the south rim and discharges at springs and seeps (Metzger, 1961). Springs and seeps in Grand Canyon typically occur at geologic contacts where water-bearing rocks are underlain by less permeable rocks. The larger springs in the canyon are associated with regional faults.

Most of the major springs along the south rim of Grand Canyon discharge from the Muav Limestone, the Temple Butte Formation, or the Redwall Limestone (**fig. 2**) and are associated with faults, fractures, or solution cavities in these rock units (Metzger, 1961; Huntoon, 1974). Springs and seeps discharge along the south rim from the Cambrian Tapeats Sandstone in many tributary canyons. Above the Tapeats Sandstone, the Bright Angel Shale acts as a physical barrier limiting downward movement of ground water from overlying rocks (Metzger, 1961).

Numerous breccia pipes occur on the Coconino Plateau (Billingsley, 2000). The form of breccia pipes found in this region are the result of solution collapse in the Redwall Limestone and a stoping of overlying rock units. The presence of breccia pipes in the study area may be significant because secondary mineralization associated with breccia pipes can potentially influence water chemistry and because breccia pipes can act as conduits for ground-water flow (Wenrich and others, 1994). Breccia pipes are small features, and because they are so numerous in the study area, only those that are close to springs included in this report are shown on illustrations.

The hydrogeology of the Redwall-Muav Limestone aquifer and its hydraulic connection with water-bearing zones in overlying Pennsylvanian and Permian rocks near Grand Canyon is not well understood. Few wells have been developed on the Coconino Plateau, and knowledge of subsurface geology in the region is limited. Little is known about the movement of ground water through the rocks of the Coconino Plateau; therefore, travel times to points of discharge are unknown.

The hydraulic gradient of the potentiometric surface near the south rim is primarily controlled by secondary porosity and is downward towards the canyon (Metzger, 1961); most tributary canyons contain springs or seeps. Depth to ground water is about 900 m. Most springs along the south rim of Grand Canyon are small (discharge less than 15 L/min). There is little or no surface flow on the Coconino Plateau near the south rim, and flow in the side canyons derives from spring discharge or storm runoff.

### **Previous Investigations**

Scientists have visited springs and measured spring flow along the Colorado River of Grand Canyon since the late 1800s; however, most springs have not been inventoried and measured at or near their source. John Wesley Powell, during 1869–72, was the first scientist to describe spring flow in Grand Canyon (Powell, 1874). Johnson and Sanderson (1968) wrote the most recent comprehensive report on spring flow in the Grand Canyon region and summarized previous investigations by LaRue (1925). Johnson and Sanderson (1968) visited major spring sites along the Colorado River in 1960, measured flow at 14 sites close to the river, and estimated flow at several other locations. Miscellaneous spring measurements have been made since 1950 by various investigators, and data for those sites are available in published reports or from the USGS National Water Information System database.

Huntoon (1974) described structurally controlled spring flow for the Kaibab Plateau north of Grand Canyon, and concluded that 65 to 97 percent of the plateau is drained through carbonate rocks 900 m below the north rim of Grand Canyon and Marble Canyon. Huntoon (1981) measured flow at several springs along the Colorado River in Marble Canyon between river miles 30 and 34.

Cooley (1976) described spring flow issuing from pre-Pennsylvanian rocks in Marble Canyon and in the lower 21 km of the Little Colorado River canyon. Cooley and other investigators (1969) visited more than 50 spring sites in the Marble and Little Colorado River canyons during 1966–67. Many of these sites were seeps or had flows estimated at less than 3.8 L/min.

E.L. Montgomery and Associates, Inc. (1997 and 1999) created a numerical model of the ground-water system of the Coconino Plateau as part of an environmental impact statement related to proposed development in the Tusayan area. Wilson (2000) constructed a geologic framework model for the region. This work was expanded by Kessler (2002).

Several researchers have conducted geochemical investigations of springs in Grand Canyon in recent years. Loughlin and Huntoon (1983) compiled water-quality data from various sources. Foust and Hoppe (1985) studied seasonal variability of water chemistry from tributaries of the Colorado River in Grand Canyon. From 1990 to 1995 the USGS (unpublished data) collected data primarily from sites close to the Colorado River and on the Hualapai Indian Reservation in the western part of Grand Canyon (Donald Bills, hydrologist, USGS, written commun., 2002). The University of Nevada-Las Vegas also collected samples from springs in Grand Canyon for chemical analysis (Goings, 1985; Zukosky, 1995; Fitzgerald, 1996). The USGS conducted synoptic water-quality studies of many tributaries of the Colorado River in Grand Canyon in 1990 and 1991 (Taylor and others, 1996) and throughout the Colorado River corridor (Antweiler and Taylor, unpublished data, 1992–93). NPS scientists have compiled several databases on the water resources of GRCA (National Park Service, 1996). The USGS and NPS have monitored discharge from the Redwall-Muav Limestone aquifer of Indian Gardens, Hermit Creek, and Cottonwood Creek since 1994.

Data from springs, creeks, and wells along the south rim of Grand Canyon have been collected by the USGS and others under separate programs subsequent to the WQAM study presented in this report, but are not presented in this report.

#### Acknowledgments

The authors thank the many individuals who helped collect and carry water samples out of Grand Canyon for this project. Personnel from the USGS Arizona District and the USGS National Research Program in Boulder, Colorado, carried heavy backpacks during harsh environmental conditions. Many volunteers also helped carry equipment and samples from remote sites in Grand Canyon. Without the assistance of these individuals, this project would not have succeeded. We also appreciate the assistance of GRCA personnel who provided backcountry permits to our datacollection teams and helped determine the feasibility of accessing many of the spring and creek sites. In addition, we thank Dave Roth, Terry Plowman, and Dale Peart with the USGS National Research Program for laboratory assistance.

# **INVESTIGATION METHODS**

### **Field Logistics**

Many of the springs issuing along the south rim of Grand Canyon (**fig. 2**) are more than 900 m below the canyon rim. Spring sources generally are far from established trails in the headwaters of tributary canyons. Because this area is remote and its terrain rugged, most sample collection trips were made using backpacks to carry equipment and supplies. Some sites were accessed by boat from the Colorado River. Unsafe conditions precluded visits to several potential sampling sites.

Hiking trips varied from 1 to 7 days and required 2 to 6 people depending upon individual site characteristics. Many of the springs were accessible from the Tonto Trail, which traverses the slopes of the Bright Angel Shale (fig. 2 and pl. 1) parallel to the Colorado River. Sampling trips were timed for favorable weather conditions and for discharge conditions that represented base spring flow.

#### Sample Collection and Processing

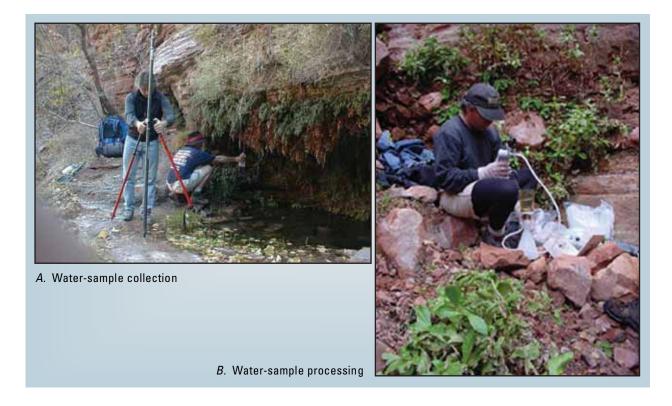
Water samples were collected from each spring at or as near as possible to the point of issuance from the water-bearing rock (fig. 3A). Ground-water discharge generally occurs along bedding planes or through fractures. Most water samples were collected from the area of greatest discharge. Where the point of issuance could not be accessed, water samples were collected at the first accessible location downstream from this point. Efforts were made to minimize contact of spring water with the atmosphere.

Measurements of spring discharge, specific conductance, pH, air and water temperature, dissolved-oxygen concentration, and alkalinity using standard USGS protocols (U.S. Geological Survey, 1997) were made at each site when samples were collected. Specific-conductance and pH meters were calibrated in the field using standard solutions. Alkalinity titrations were done in the field using a pH meter and alkalinity kit. Measurements were made onsite directly in flowing water or from a sample of water. Locations and altitudes of sample sites were determined using Differential Global Positioning System or topographic maps.

Spring discharge was measured volumetrically, with a portable Parshall flume, or with a current meter according to methods described by Rantz and others (1982). Volumetric methods were used when discharge was less than about 25 L/min or where depths were too shallow for use of a current meter.

A distinct sampling protocol, described in this section, was used to collect water samples for laboratory analyses. Because different analytical methods required different sample treatments, each spring sample was collected into 10 distinct bottles for analysis (table 2).

Samples were handled in a proper and consistent manner to minimize contamination. All bottles were pre-cleaned before sample teams entered Grand Canyon. The metals, anions, nutrients, dissolved organic carbon (DOC), and oxygen-18 and hydrogen-2 (D or deuterium) bottles for each site were sequestered into a separate zippered plastic bag, which also contained two pairs of polypropylene gloves, a new 60-mL polystyrene syringe, and four Acrodisk 0.45-micron nominal pore-size Luer-Lock filters (contained in their own zippered plastic bag). Mercury bottles were predosed with 2.5 mL of potassium dichromate/nitric acid and were sealed in their own individual Whirl-pak bags and kept separate from the rest of the sampling equipment and bottles. The remaining four bottles were kept together, but separate from all the other sample bottles. Because the large-volume radiochemistry, carbon-isotope, and strontium-isotope samples also required filtration, a hand-vacuum pump, a filtration manifold, and 0.45-micron polysulfone filters were also packed, but kept separate from the other sampling equipment.



**Figure 3**. Water-sample collection and processing. *A*. Water-sample collection at Miners Spring; *B*, Water-sample processing at Monument Spring.

Table 2.	Spring ar	d creek	sample	protocols

Sample type	Bottle
Anions (chloride, sulfate, carbonate and bicarbonate)	60 mL polyethylene
Nutrients (nitrate, nitrite, phosphate and ammonium)	30 mL amber polyethylene
Metals (major cations and trace elements, except mercury)	125 mL acid-rinsed polyethylene
Dissolved organic carbon	60 mL pre-baked amber glass
Mercury	60 mL acid-rinsed clear glass
Oxygen-18 and hydrogen-2 (deuterium)	60 mL clear glass
Tritium (hydrogen-3)	1 L polyethylene
Radiochemistry (gross alpha and beta)	2 L acid-rinsed polyethylene
Carbon-14, carbon-13/carbon-12	250 mL glass
Strontium-87/strontium-86	250 mL polyethylene

Field teams wore polypropylene gloves for all sample processing and collected samples in a certain order to minimize the potential for contamination. The non-isotope samples were collected and filtered first. A new syringe and Acrodisk filter were used, after field rinsing with spring water, to collect samples at each site. The anions bottle was filled first, followed by the nutrients, metals, and DOC bottles. Because the mercury bottle already contained preservative, it was filled last, after all other non-isotope sample bottles were filled, to minimize the possibility of cross-contaminating the other samples with the potassium dichromate preservative. A second or third Acrodisk filter was used for some nonisotope samples because the original filter became clogged.

The oxygen-18 and deuterium sample was collected unfiltered and care was taken to eliminate head-space within the bottle. Tritium samples were also unfiltered. The radiochemistry, carbon-, and strontium-isotope samples were filtered using a filtration setup that consisted of a hand vacuum pump, which was connected to a polyethylene side-arm Erlenmeyer flask by plastic tubing (fig. 3*B*). A magnetic filterholder on which was mounted a new 47-mm, 0.45-micron polysulfone filter was placed on the side-arm flask. The first 20 mL of the filtered sample was used to rinse the flask. Thereafter the sample was collected in the flask and distributed to either the radiochemistry bottle, the carbon-isotope bottle, or the strontium-isotope bottle.

All samples were kept as cool as possible in the absence of ice; these samples were refrigerated as soon as possible after leaving the field, and analyzed as quickly as possible thereafter. The metals and radiochemistry samples were acidified as soon as possible after leaving the field—metals bottles with 1 mL of doubly distilled concentrated nitric acid and the radiochemistry bottles with 5 mL of concentrated hydrochloric acid. The remaining samples did not require any additional preservation procedures.

Rock samples representing the major Paleozoic rock units in Grand Canyon were collected from outcrops near the Bright Angel Fault. These samples were analyzed for mineralogy and for strontium-87/strontium-86 and carbon-13/carbon-12 values.

#### **Laboratory Methods**

Three general categories of chemical analyses are described in the following section: (1) elemental and compound concentration analyses of water samples, (2) isotopic composition analyses of water samples, and (3) X-ray diffraction (XRD) and isotopic composition analyses of rock samples. Methods for these analyses are summarized below.

## Elemental and Constituent Analyses of Water Samples

Elements present at concentration levels in the milligram per liter range, including calcium, potassium, magnesium, sodium, and silica, and some selected trace elements, such as iron, were determined in triplicate by inductively coupled plasma-atomic emission spectrometry (ICP-AES) utilizing a Perkin Elmer Optima 3300DV multichannel emission spectrometer. A general description of the analysis conditions and procedures are reported by Garbarino and Taylor (1979). Details of the operational conditions are described by Mitko and Bebek (1999, 2000).

Trace element determinations (excluding mercury) were performed in triplicate on undiluted field-preserved samples by inductively coupled plasma-mass spectrometer (ICP\_MS) using a Perkin Elmer Elan Model 6000 inductively coupled plasma-mass spectrometer (ICP-MS). Analytes and their respective median detection limits are listed in table 3. These detection limits were computed by the method described by Skogerboe and Grant (1970) and may vary slightly between analysis runs owing to changes in instrument sensitivity or to required dilutions needed to minimize matrix interference effects. Details of the specific analysis techniques, procedures, and instrument settings are described in Garbarino and Taylor (1996) and Taylor (2001).

Table 3. Chemical and isotopic analytes and corresponding median detection limits

[µg/L, micrograms per liter; mg/L, milligrams per liter; mg N/L, milligrams nitrogen per liter; mg P/L, milligrams phosphorus per liter; mg C/L, milligrams carbon per liter; pCi/L, picocuries per liter; pmc, percent modern carbon; ‰, per mil]

Chemical constituent	Symbol	<b>Detection limit</b>	Units	Chemical constituent	Symbol	<b>Detection limit</b>	Units
Bicarbonate	HCO <sub>3</sub>	0.5	mg/L	Antimony	Sb	0.004	μg/L
Carbonate	CO <sub>3</sub>	0.05	mg/L	Arsenic	As	0.05	μg/L
Calcium	Ca	0.001	mg/L	Barium	Ba	0.02	μg/L
Chloride	Cl	0.02	mg/L	Beryllium	Be	0.03	μg/L
Potassium	Κ	0.005	mg/L	Bismuth	Bi	0.003	μg/L
Magnesium	Mg	0.0005	mg/L	Boron	В	11	μg/L
Sodium	Na	0.01	mg/L	Bromide	Br	0.5	μg/L
Sulfate	$SO_4$	0.04	mg/L	Cadmium	Cd	0.009	μg/L
Silica	SiO <sub>2</sub>	0.01	mg/L	Cerium	Ce	0.0005	μg/L
Nitrate	NO <sub>3</sub>	0.01	mg N/L	Cesium	Cs	0.01	μg/L
Nitrite	NO <sub>2</sub>	0.001	mg N/L	Chromium	Cr	0.3	μg/L
Ammonium	NH <sub>4</sub>	0.03	mg N/L	Cobalt	Co	0.007	μg/L
Phosphorus	Р	8	μg/L	Copper	Cu	0.03	μg/L
Phosphate	PO <sub>4</sub>	0.02	mg P/L	Dysprosium	Dy	0.001	μg/L
Dissolved Organic Carbon	DOC	0.1	mg C/L	Erbium	Er	0.0009	μg/L
Aluminum	Al	0.2	μg/L	Europium	Eu	0.0005	μg/L

See footnote at end of table.

Chemical constituent	Symbol	Detection limit	Units	Chemical constituent	Symbol	Detection limit	Units
Gadolinium	Gd	0.001	μg/L	Terbium	Tb	0.0003	μg/L
Holmium	Но	0.0002	μg/L	Thallium	Tl	0.005	μg/L
Iron	Fe	0.5	μg/L	Thorium	Th	0.0006	μg/L
Lanthanum	La	0.0007	μg/L	Thulium	Tm	0.0003	μg/L
Lead	Pb	0.008	μg/L	Tungsten	W	0.001	μg/L
Lithium	Li	0.07	μg/L	Uranium	U	0.001	μg/L
Lutetium	Lu	0.0002	μg/L	Vanadium	V	0.1	μg/L
Manganese	Mn	0.06	μg/L	Ytterbium	Yb	0.0007	μg/L
Mercury	Hg	0.4	ng/L	Yttrium	Y	0.0008	μg/L
Molybdenum	Мо	0.03	μg/L	Zinc	Zn	0.2	μg/L
Neodymium	Nd	0.001	μg/L	Zirconium	Zr	0.003	μg/L
Nickel	Ni	0.03	μg/L	Gross beta	β	4	pCi/L
Praseodymium	Pr	0.0004	μg/L	Gross alpha	α	3	pCi/L
Rhenium	Re	0.001	μg/L	Oxygen-18/oxygen-16	<sup>18</sup> O/ <sup>16</sup> O	$^{1}\pm0.1$	%0
Rubidium	Rb	0.003	μg/L	Hydrogen-2/hydrogen-1	$^{2}H/^{1}H$	$^{1}\pm1$	%0
Samarium	Sm	0.002	μg/L	Strontium-87/strontium-86	<sup>87</sup> Sr/ <sup>86</sup> Sr	10.00002	none
Selenium	Se	0.4	μg/L	Carbon-14	<sup>14</sup> C	<sup>1</sup> ±3	pmc
Strontium	Sr	0.1	μg/L	Carbon-13/carbon-12	<sup>13</sup> C/ <sup>12</sup> C	$^{1}\pm1$	%0
Tellurium	Te	0.02	μg/L	Tritium	<sup>3</sup> H	1	pCi/L

 Table 3.
 Chemical and isotopic analytes and corresponding median detection limits—Continued

<sup>1</sup>Reported precision.

Trace concentrations of dissolved mercury were measured in triplicate using an automated cold-vapor atomic fluorescence spectrometric method and a PS Analytical Millenium System mercury analyzer. Details of the method have been described by Roth (1994) and Roth and others (2001).

Chloride and sulfate were determined by ion exchange chromatography using a modified Dionex 2002i/SP series ion chromatograph. The details of the analytical procedure are described by Brinton and others (1996).

Dissolved nutrients including nitrate, nitrite, ammonium, and orthophosphate were determined in triplicate using methodology described by Antweiler and others (1996). Analyses were performed using an air-segmented continuous flow analyzer (Alpkem RFA-300).

Dissolved organic carbon was determined by oxidation of the organic carbon in the sample to carbon dioxide, which was subsequently measured by an infrared absorption spectrophotometric technique (Wershaw and others, 1983).

Gross-alpha/gross-beta counting is a semiquantitative technique for measuring overall radioactivity in water samples without extensive sample preparation. This method is most often applied for screening purposes. The method has the advantage of being sensitive to activity from a wide range of radionuclides, but does not inherently provide qualitative information about the identity of the radioactive isotopes present. Samples are prepared by evaporation to dryness on a planchet. Gross alpha and beta emissions are measured on the dried residue using a gas-proportional counter by a method specified by the U.S. Environmental Protection Agency (1976). Count rates from the detector are converted to and reported as activities. More details of the methodology can be found in American Public Health Association (1985).

# Isotopic Composition Analyses of Water and Rock Samples

Stable isotopes of oxygen, hydrogen, carbon, and strontium were measured relative to internationally agreed upon standards (International Union of Pure and Applied Chemistry, 1994). For oxygen, hydrogen, and carbon, the deviation of the sample from the standard mean is expressed by the delta ( $\delta$ ) notation, in parts per mil (%), and is computed from equation 1 (appendix). Strontium values are shown as ratios.

Oxygen isotope ratio (<sup>18</sup>O/<sup>16</sup>O) analyses were performed using the carbon dioxide equilibration technique described by Epstein and Mayeda (1953). Hydrogen isotope ratios (<sup>2</sup>H/<sup>1</sup>H) for water were determined using a hydrogen equilibration technique at 30°C to measure <sup>2</sup>H activity (Coplen and others, 1991). Oxygen and hydrogen isotopic results are reported in

per mil relative to Vienna Standard Mean Ocean Water (VSMOW) and normalized (Coplen, 1988, 1994; International Union of Pure and Applied Chemistry, 1994) on scales such that the <sup>18</sup>O/<sup>16</sup>O and <sup>2</sup>H/<sup>1</sup>H values of Standard Light Antarctic Precipitation (SLAP) are -55.5 ‰ and -428 ‰, respectively.

Water samples also were measured for tritium (<sup>3</sup>H), the radioactive isotope of hydrogen, using a liquid scintillation counting technique (Kendall and Caldwell, 1998) after preconcentration by an electrolytic enrichment procedure. Results are reported in pCi/L with accompanying 2 sigma error in pCi/L, and tritium units (TU; 1 TU=1<sup>3</sup>H per 1,018 hydrogen atoms; Fritz and Fontes, 1980; Clark and Fritz, 1997).

Strontium isotope (<sup>87</sup>Sr/<sup>86</sup>Sr) analyses of water and crushed rock samples were performed using solid source mass spectrometry (Taylor, 2000; Bullen and others, 1996). Rock samples were leached in a 0.1 N hydrochloric acid solution before analysis for strontium isotopes. This procedure normalized <sup>87</sup>Sr/<sup>86</sup>Sr results (as ratios) for natural and analytical fractionation to 8.37521.

Carbon isotope (<sup>13</sup>C/<sup>12</sup>C) analyses of water and crushed rock samples were measured by an isotope ratio mass spectrometric technique (Clark and Fritz, 1997) after conversion of inorganic carbon to carbon dioxide by addition of hydrochloric acid. All <sup>13</sup>C results are reported in per mil relative to the Vienna Peedee belemnite standard (Coplen, 1994). These data are necessary to correct carbon-14 (<sup>14</sup>C) results for ground-water dating applications.

Carbon-14 isotope analyses of water samples were measured by accelerator mass spectrometry according to methods described in Beukens (1992). All <sup>14</sup>C determinations are reported in percent modern carbon (pmc) normalized to the 1950 National Bureau of Standards (National Bureau of Standards, 1984) oxalic acid standard (Stuiver and Polach, 1977; Wigley and Muller, 1981), with accompanying 1 sigma error in pmc.

#### X-ray Diffraction Analyses of Rock Samples

Bulk mineralogy and the clay mineral fraction of rock samples representing the major stratigraphic units near the Bright Angel Fault in Grand Canyon were detemined using XRD techniques (Schultz, 1964; Moore and Reynolds, 1997). Physical grain-size techniques were used to determine particlesize distribution for the rock samples. Results of the XRD and grain-size analyses were compared with descriptions of similar rock units in the region (R. Parnell, professor, and T. Loseke, graduate student, Northern Arizona University, written commun., 2002).

#### **Isotopic Analyses**

Stable isotopes of oxygen and hydrogen are used in hydrologic studies to help determine sources of recharge for aquifers. During evaporation, the isotopes of hydrogen and oxygen in water fractionate. The lighter <sup>16</sup>O and <sup>1</sup>H molecules preferentially move from the liquid phase to the gas phase leaving behind <sup>18</sup>O and <sup>2</sup>H molecules. This action will affect the isotopic signature by making it more positive or enriched in the heavier water molecules. Ratios of these isotopes are unaffected by low-temperature geochemical processes in ground-water systems (Clark and Fritz, 1997).

Craig (1961) developed a relation between  $\delta^{18}$ O and  $\delta^{2}$ H (or  $\delta$ D) for meteoric waters that is represented by equation 2 (**appendix**) and is known as the global meteoric water line. Regional variations in  $\delta^{18}$ O and  $\delta$ D may be caused by differences in latitude, altitude, or temperature. For this study, a local meteoric water line was developed using  $\delta^{18}$ O and  $\delta$ D data from precipitation samples collected by the NPS at the National Atmospheric Deposition Program National Trends Network Monitoring Location AZO3, Grand Canyon National Park, Coconino County, Arizona, between 1989 and 2002 (Pendall, 1997; Harvey, 2000; National Atmospheric Deposition Program, 2002). Water samples for  $\delta^{18}$ O and  $\delta$ D analysis were collected at 14 springs and 4 creek sites along the south rim of Grand Canyon.

Tritium is commonly used as a tracer for studying groundwater residence time and ground-water mixing at time scales less than 50 years. Low activities of cosmogenic tritium occur as background in natural waters. Anthropogenic tritium was produced by atmospheric thermonuclear tests that began in 1952. Thermonuclear tritium activities peaked in about 1963 before thermonuclear testing was banned (Mazor, 1997).

Tritium has a half-life of 12.3 yr, and radioactive decay of tritium occurs according to equation 3 (appendix), allowing approximation of the date of ground-water recharge. Precipitation samples collected in Flagstaff between 1962 and 1971 were analyzed for tritium activities (International Atomic Energy Agency, 2001). Data range from 2.7 to 53 TU and generally show a decreasing trend over time. In continental regions, tritium activities of less than 0.5 TU indicate recharge prior to 1952. Activities of 0.5 TU to 10 TU indicate possible mixing of pre-1952 and post-bomb waters, and activities greater than 10 TU represent recharge less than 50 yr ago (Mazor, 1997). Water samples were collected for tritium analysis at 13 springs and 4 creek sites for this study.

Ground-water flow paths and water-rock interactions were identified using <sup>87</sup>Sr/<sup>86</sup>Sr values. Strontium-87 is radiogenic and increases in concentration upon the beta-decay of rubidium-87. Geologic units have distinct <sup>87</sup>Sr/<sup>86</sup>Sr, and strontium enters water through dissolution of minerals from these units during chemical weathering. Dissolution rates are dependent on chemical processes and the minerals that are present. The <sup>87</sup>Sr/<sup>86</sup>Sr in ground-water reflects the strontium ratios of the host rock, but may differ slightly if the ground-water flow path includes multiple rock units. This relation enables the use of strontium isotope composition as a tracer in many ground-water systems (Faure, 1986). Water samples collected from 13 springs and 4 creek sites and 16 rock samples collected near the Bright Angel Fault were analyzed for <sup>87</sup>Sr/<sup>86</sup>Sr.

Residence time of ground water can be estimated using the decay rate of naturally occurring <sup>14</sup>C. During 2000–2001 water samples collected at 12 springs and 4 creek sites were analyzed for <sup>13</sup>C/<sup>12</sup>C and <sup>14</sup>C.

Carbon-14 data in pmc were converted (equation 4, **appendix**) to time in years using a method known as Conventional Radiocarbon Age (Stuiver and Polach, 1977) to estimate ground-water residence times.

The Conventional Radiocarbon Age (years before present) term (Stuiver and Polach, 1977) implies

- 1. the use of the 5,568 yr half-life (mean life 8,033 yr);
- 2. the constancy of the <sup>14</sup>C atmospheric level during the past;
- 3. isotopic fractionation normalization of all sample activities to the base  $\delta^{13}C = -25 \%$ ; and
- 4. the year 1950 is automatically the base year, or present, with ages given in years before present.

The Conventional Radiocarbon Age method converts <sup>14</sup>C in pmc to ages in years without considering either isotopic dilution or fractionation from water and carbonate rock interactions. This approach biases the adjusted ground-water age to be artificially older. A better estimate of ground-water residence time using <sup>14</sup>C data requires adjustment to account for the influence of isotopic dilution and fractionation from water and carbonate rock interactions along flow paths. Pearson and Hanshaw (1970) developed equation 5 (appendix), which accounts for isotopic dissolution when combined with the Conventional Radiocarbon Age method developed by Stuiver and Polach (1977). Corrected ground-water residence times were then developed using equation 6 (appendix). The <sup>14</sup>C (adjusted) is then plugged into equation 4 as the pmc value to adjust for isotopic dilution.

Data requirements for <sup>14</sup>C adjustments include  $\delta^{13}$ C from ground water and local carbonate rocks and carbon dioxide from the soil zone (soil gas). The  $\delta^{13}$ C of the soil gas was not available for use in equation 5; therefore, this parameter was estimated. Values of -9 %*c*, -11 %*c*, -13 %*c*, -15 %*c*, and -17 %*c* were used, bracketing known values for soil gas in semiarid regions.

Springs and creeks along the south rim of Grand Canyon represent end points of the regional ground-water flow system. The lack of data from wells close to Grand Canyon, or from wells or springs at other points along the ground-water flow path, limits the ability to accurately estimate ground-water residence time from  $^{14}$ C data.

# SITE DESCRIPTIONS

The 20 spring and creek sites sampled are in tributary drainages to the Colorado River, from Red Canyon in the east to Boucher East Spring in the west (**pl. 1**). Some drainages contain more than one site. Sites where water discharged from bedrock or from above the flood plain were defined as springs. Sites where water discharged from channel alluvium were defined as creeks. In some drainages, debris had buried spring sources and water issued from alluvium. It is not known how

far the issuing water was from the source spring at these sites; however, interpretation of water chemistry data can provide information about this relation.

Site selection criteria developed by the USGS and GRCA were:

- 1. Accessibility. Remoteness of the site, access for sample collection, and safety of field personnel.
- 2. Spatial distribution. Because the study area is large and the boundaries of the Redwall-Muav Limestone aquifer are poorly defined, springs and creeks were selected to provide adequate coverage of the study area.
- 3. Source rock. Sample collection was limited to springs or creeks that discharge from the Redwall Limestone, the Temple Butte Formation, the Muav Limestone, or the Bright Angel Shale.
- 4. Spring discharge. Preference was given to springs having measurable discharges.
- 5. Location. Preference was given to springs and creeks closest to Grand Canyon Village and Tusayan. These are areas of present and likely future ground-water pumping, and the springs and creeks were selected because they could potentially show effects of ground-water pumping sooner than springs farther away from areas of pumping.
- Point of issuance. Whenever possible, sample sites were selected where water emerged directly from bedrock. These springs were frequently several kilometers or more from the Tonto Trail and had not been previously inventoried.

Springs and creeks were categorized by geologic characteristics, which included (1) bedding planes, (2) fractures, (3) hillslopes, (4) stream alluvium, and (5) travertine deposits (**fig. 4** and **table 4**).

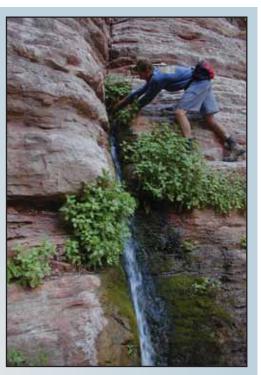
Springs that discharge from bedding planes are typically near the contacts between the Redwall Limestone and the Temple Butte Formation, where present; the Redwall Limestone and Muav Limestone; or the Muav Limestone and the Bright Angel Shale. Springs discharging from hillslopes emerge in alluvial deposits of tributaries immediately below bedrock outcrops. Sites where water discharges from stream alluvium differ from hillslope springs in that ground water discharges far below the head of the canyon or bedrock outcrop and from within the active stream channel. Springs issuing from travertine deposits are along canyon walls. Physical characteristics of each of the springs and creeks are described in the following section and are presented in order from east to west. Data describing physical location and dates of site visits are included in table 1. Water samples were collected at all sites except where noted.



A. Bedding planes—water issues from bedding planes in the Muav Limestone at Salt Creek Spring

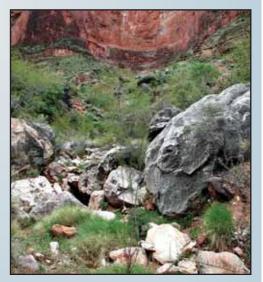


C. Hillslopes—Burro Spring issues from hillslope alluvium at the east side of the Pipe Creek drainage



B. Fractures—vertical fracture at Monument Spring





D. Channel alluvium— water issues from channel alluvium in the upper Horn Creek drainage

E. Travertine deposits—travertine deposits at Boucher East Spring

Figure 4. Spring types at sample sites. A, Bedding planes; B, Fractures; C, Hillslopes; D, Channel alluvium; E, Travertine deposits.

 Table 4.
 Geologic characteristics of springs and creeks sampled in study [see plate 1]

Bedding planes	Fractures	Hillslopes	Channel alluvium	Travertine deposits
Red Canyon Spring	Monument Spring	Burro Spring	Cottonwood Creek No. 1	Boucher East Spring
JT Spring		Pumphouse Spring	Cottonwood Creek No. 2	
Miners Spring		Pumphouse Wash Gage	Pipe Creek	
Grapevine Main Spring			Horn Creek	
Grapevine East Spring			Monument Creek No. 1	
Lonetree Spring				
Sam Magee Spring				
Salt Creek Spring				
Hawaii Spring				
Hermit Spring				

**Red Canyon Spring**.—Red Canyon Spring discharges from bedding planes in the Muav Limestone in the main drainage of Red Canyon and consists of numerous small spring outlets close to one another. The site is at the head of the canyon and at the base of a 300-m-high, near vertical cliff of Redwall Limestone and Muav Limestone (fig. 2). The springs lie along the north-striking Hance Fault (pl. 1). Water samples were collected from the largest of the spring outlets at the east side of the canyon about 75 m downstream from the headwall. Discharge was measured about 300 m downstream where flow in the channel is constricted over a bedrock ledge. Water flowed in the channel of Red Canyon Creek for about 1 km downstream from the point of issuance during all visits. Farther downstream, flow becomes subterranean until it emerges near the mouth of Red Canyon. High cliffs shade the spring outlets much of the time, and maidenhair fern (Adiantum capillusveneris) and other riparian plant species are abundant at the site.

JT Spring.—JT Spring is in an alcove erosional feature that is a tributary to the east branch of Hance Creek. The spring has several outlets issuing from bedding planes in the Muav Limestone (fig. 2). The largest outlet at the time of sampling was at the north side of the amphitheater and was mostly covered by an alluvial slump block. The back wall of the amphitheater is a 300-m-high vertical cliff of Redwall Limestone. The Grandview-Phantom Monocline (pl. 1) trends parallel to the east arm of Hance Creek and is immediately southwest of JT Spring. Samples were collected where the water emerges directly from the slump block. Discharge was measured about 60 m from the spring outlet where all flow from the outlet is in a single channel. Immediately downstream from the confluence of the tributary containing JT Spring and the east branch of Hance Creek, flow becomes subterranean until it emerges near the contact between the Bright Angel Shale and Tapeats Sandstone close to the Tonto Trail (pl. 1).

Miners Spring.-Miners Spring (also known as Page Spring) is just south of Horseshoe Mesa near the head of a small unnamed tributary of Hance Creek (pl. 1). The perennial spring discharges from bedding planes at the cliff face near the contact between the Temple Butte Formation and the Muav Limestone (fig. 2). Spring water drips from beneath a small alcove into a pool in a bedrock depression created by miners in the early 1900s. The spring probably originates from localized faulting along the axis of the northwest-trending Grandview-Phantom Monocline (pl. 1). Water samples were collected by capturing drips coming from the wall of the small alcove. Discharge was measured volumetrically by constricting all flow discharging from the pool into a single conduit. Water leaving the small pool flows about 3 m across the bedrock outcrop and enters the alluvium, and does not reemerge downstream. Aquatic vegetation is extensive at the site, and frogs and water skates were observed in the pool and along its edge during sampling visits.

Cottonwood Creek No. 1.-Cottonwood Creek drains the western side of Horseshoe Mesa (pl. 1) and has three major branches in its upper reaches. The eastern branch, which abuts Horseshoe Mesa, is almost always dry. The central branch has perennial flow in places, and the USGS maintains a streamflow-gaging station (station number 09402450) on this branch about 150 m upstream from the confluence with the eastern branch. There is a small seep in the western branch near the base of the Bright Angel Shale. This seep was not included in this study. Water in the central branch discharges from channel alluvium about 80 m upstream from the streamflowgaging station during most of the year (see site description for Cottonwood Creek No. 2 in the following text). Several springs and seeps discharge from the Temple Butte Formation and Muav Limestone (fig. 2) near the head of this branch; however, these sites were not sampled. The water sample for Cottonwood Creek No. 1 was collected about 100 m

downstream from the gaging station. The discharge was measured near the sample site. Large Fremont cottonwood (*Populus fremontii*) trees and aquatic plant species are present.

Cottonwood Creek No. 2.-Cottonwood Creek No. 2 is on the western side of Horseshoe Mesa in the central branch of three branches in the headwaters of Cottonwood Creek (pl. 1). The perennial site discharges from channel alluvium about 600 m upstream from the USGS streamflow-gaging station (station number 09402450). During site visits, water flowed from beneath a large boulder on the east side of the active channel. Flow continued in the channel 10 m before becoming subterranean and re-emerging about 80 m upstream from the streamflow-gaging station. Cottonwood Creek No. 2 is in the Bright Angel Shale (fig. 2). This formation is not exposed near the site; however, outcrops are visible lower and higher in the drainage. The axis of the Grandview-Phantom Monocline trends northwest in this vicinity (pl. 1). Water samples were collected where water flows from a small pool adjacent to the point of issuance. Discharge was measured at the streamflowgaging station. A thick stand of woody riparian species and aquatic vegetation are present in and near the stream channel.

Grapevine Main Spring.—Grapevine Main Spring is in a small canyon that is tributary to the main southeast trending canyon of Grapevine Creek about 3 km upstream from the Tonto Trail (pl. 1). Water discharges from bedding planes at several places in the upper part of the Muav Limestone, just below the Redwall-Muav contact (fig. 2). The springs are at the base of a massive vertical cliff that is predominantly Redwall Limestone. An unnamed fault strikes parallel to the Grapevine Creek drainage near the spring, and the McKee Fault bisects the canyon near the Tonto Trail (pl. 1). The spring is south of the northwest-trending Grandview-Phantom Monocline. Water samples were collected where the water emerges from the bedding planes. Discharge was measured about 80 m downstream from the sample site where all flow is constricted over a bedrock ledge. Flow in Grapevine Creek was continuous from near the side drainage containing the spring to near where the Tonto Trail crosses the creek. Flow in Grapevine Creek is intermittent downstream from the Tonto Trail. The main canyon and the tributary where the spring is located are densely vegetated with woody xeric and riparian species.

*Grapevine East Spring.*—Grapevine East Spring is in a small drainage on the east side of the Grapevine Creek drainage, directly adjacent to the Tonto Trail (pl. 1). The Tonto Trail crosses the drainage near the lower end of a small wetlands. The wetlands are supported by a group of small springs and seeps, most of which are obscured by the vegetation. During the site visit in May 2000, the first appearance of water was in the midst of the wetlands, and water samples were collected at this point. During site visits in

December 2000 and April 2001, water was discharging from bedding planes in the lower part of the Bright Angel Shale (fig. 2) about 30 m upstream from the wetlands, and water samples were collected here. Discharge was measured at a small waterfall about 30 m downstream from the trail. The site is just to the north of the northwest-trending Grandview-Phantom Monocline. Riparian vegetation in the area consists of sawgrass (*Cladium californicum*), Fremont cottonwood, and willow (*Salix* spp.).

Lonetree Spring.-Lonetree Spring is at the upper end of a small steep tributary that trends south from the main Lonetree Canyon drainage about 2 km upstream from the Tonto Trail (pl. 1). A group of small springs discharge along bedding planes in the upper part of the Muav Limestone at the head of the drainage. The discharge points are near the Redwall-Muav contact below a massive vertical wall of Redwall Limestone (fig. 2). The spring group is south of the Grandview-Phantom Monocline (pl. 1). Flow at individual discharge points was too small to fill sample containers; therefore, water samples were collected about 30 m downstream where flow from all spring outlets was combined. Discharge was measured 3 m downstream from the sample site at a bedrock ledge constriction. Just downstream from the discharge-measurement site, spring flow becomes subterranean and re-emerges near the contact between the Bright Angel Shale and Tapeats Sandstone close to the Tonto Trail (pl. 1). During summer visits, however, spring flow did not occur near the trail crossing.

Sam Magee Spring.—Sam Magee Spring is a small spring that emerges from bedding planes at an outcrop of the Muav Limestone (fig. 2), about 500 m above the Tonto Trail, at the eastern side of the Cremation Creek drainage, and just west of Patti Butte (pl. 1). The water flows for about 5 m before it infiltrates the channel alluvium. Water samples were collected where flow emerges from the bedding planes, and discharge was measured about 3 m downstream. During this study, field parameters were measured and water samples were collected for  $\delta^{18}$ O,  $\delta$ D, and tritium analyses only. The spring is northeast of the Grandview-Phantom Monocline (pl. 1).

*Pipe Creek*.—Perennial flow in Pipe Creek discharges from alluvium in the main Pipe Creek channel about 250 m upstream from the Tonto Trail crossing (pl. 1). Water emerges from channel alluvium overlying the lower part of the Bright Angel Shale (fig. 2). The site is near the northwest-striking Pipe Fault (pl. 1). Water samples were collected where flowing water was first evident, and discharge was measured at a permanently installed weir that was originally part of a USGS streamflow-gaging station (Pipe Springs Creek above Tonto Trail, near Grand Canyon, Arizona, station number 09403010; discontinued in 1996). Water flows continuously at the surface of the Pipe Creek channel from the sample site to the Colorado **Burro Spring**.—Burro Spring is along the Tonto Trail at the east side of the Pipe Creek drainage below O'Neill Butte (**pl. 1**). This perennial spring discharges from hillslope alluvium overlying the lower part of the Bright Angel Shale (**fig. 2**). The spring is southwest of an unnamed northweststriking fault (pl. 1) and lies beneath a breccia pipe on the west face of O'Neill Butte (pl. 1). Water discharges diffusely from the hillslope, and sustains an area of thick vegetation, obscuring the points of issuance. Thus, the water sample was collected at the Tonto Trail crossing. Discharge was measured below the trail crossing, just above a large waterfall into Pipe Creek canyon. The area surrounding the spring is covered with sawgrass and there are several large Fremont cottonwood trees.

**Pumphouse Spring**.—Pumphouse Spring (also known as Two Trees Spring) is about 100 m to the east and 50 m up the slope from the NPS Indian Gardens pump station (pl. 1). Spring water first emerges from beneath a large boulder that is about 3 m upslope from two prominent Fremont cottonwood trees. The spring discharges from hillslope alluvium overlying the Bright Angel Shale (fig. 2) and is in the northeast-striking Bright Angel Fault zone (pl. 1). Water samples were collected near the boulder, and discharge was measured volumetrically by constructing a temporary weir directly below that point. The water enters the alluvium immediately below the spring and re-emerges about 5 m above the USGS streamflow-gaging station Pumphouse Wash Spring near Grand Canyon, Arizona, station number 09403013.

**Pumphouse Wash Gage.**—The sample site is a USGS streamflow-gaging station (09403013) in the Garden Creek basin that monitors discharge from Pumphouse Spring. The station is about 40 m northeast and 8 m up the slope from the NPS Indian Gardens pump station and approximately 100 m down the slope from Pumphouse Spring (pl. 1) in a small stream channel developed in hillslope alluvium that overlies the Bright Angel Shale (fig. 2). The site is in the Bright Angel Fault zone (pl. 1). Water samples were collected from the small pool of flowing water at the streamflow-gaging station, and discharge was measured with a Parshall flume. Downstream from the gaging station, water flows continuously in a small channel for about 30 m and then enters Garden Creek, which flows continuously to the Colorado River.

*Horn Creek.*—Horn Creek consists of two primary branches. The west branch is usually dry, and the east branch has perennial flow in places. The sample site is in the east tributary about 500 m upstream from the Tonto Trail crossing (pl. 1) where water discharges from the channel alluvium that overlies the Bright Angel Shale (fig. 2). Water samples were collected at a small waterfall that was formed by boulders in the stream channel. Discharge was measured volumetrically at a small waterfall near the sample site. The spring flow emerges on the downthrown side of the northwest-striking Salt Fault (pl. 1) and flows intermittently to the Colorado River. At the head of the drainage is a breccia pipe and a historic uranium mine. Recent flooding has removed most of the vegetation at the site, leaving a few Fremont cottonwood trees in the nearby channel reach.

Salt Creek Spring.—Salt Creek Spring is at a headwall in the main Salt Creek drainage about 800 m upstream from the Tonto Trail crossing (pl. 1). Water discharges from bedding planes in the Muav Limestone and drips down an 8-m high cliff face of Muav Limestone onto a small talus slope on the west side of the canyon. The point of issuance of water is inaccessible; therefore, water samples were collected at the talus slope using a Visqueen sheet to funnel the flow into a Teflon holding bottle. All water enters the channel alluvium immediately downstream from the talus slope. During all site visits there was no evidence of flowing water in the stream channel from the talus slope to the Tonto Trail crossing. Discharge was measured where water flows over exposed Tapeats Sandstone (fig. 2) surfaces near the Tonto Trail (pl. 1). Water flows intermittently downstream from the Tonto Trail crossing to the Colorado River. No large-scale geologic structures have been identified near the spring. Recent flooding has removed large vegetation from the Salt Creek drainage.

Monument Spring.—Monument Spring is at the headwall of a small tributary on the east side of Monument Creek about 3 km upstream from the Tonto Trail. The head of the Monument Creek drainage is known as The Abyss, and the near-vertical cliffs rise about 1 km to the south rim of Grand Canyon (pl. 1). The spring is in an alcove above a series of waterfalls that support hanging gardens of ferns. There are numerous smaller springs and seeps in the alcove. The main spring discharges from a vertical fracture near the contact between the Redwall Limestone and the Temple Butte Formation (fig. 2). Water samples were collected at the point of issuance. Discharge measurements were made at the bedrock outcrop below the Tonto Trail. An unnamed fault strikes northwest near the spring. Monument Creek is intermittent between the small tributary where the spring is located and the Colorado River. The canyon wall has several hanging gardens that include maidenhair fern, golden columbine (Aquilegia chrysantha), and other aquatic plants.

*Monument Creek No. 1.*—At Monument Creek, water surfaces in the channel alluvium about 50 m upstream from the Tonto Trail crossing (pl. 1) and flows through a series of pools and riffles. A water sample was collected from Monument Creek about 30 m upstream from the trail. In this reach of Monument Creek, the channel alluvium overlies the Tapeats Sandstone (fig. 2). Discharge was measured about 300 m downstream near the NPS campground at a point where water emerges from the alluvium and flows in a small channel over an outcrop of Precambrian bedrock. Water flows intermittently in the stream channel from the discharge measurement site to the Colorado River. Several large debris flows removed vegetation from the stream channel during the 1980s and 1990s.

Hawaii Spring.—Hawaii Spring is about 1.8 km upstream from where the Tonto Trail crosses the Hermit Creek drainage (pl. 1). The perennial spring discharges from bedding planes in the Muav Limestone (fig. 2) in a small grotto to the west of Hermit Creek. There are numerous springs and seeps in the Hermit Creek drainage, and Hawaii Spring is among the largest of the springs. Water samples were collected from the largest of many seeps in the grotto. It was not possible to measure the discharge of Hawaii Spring. Instead, discharge was measured at the USGS streamflow-gaging station on Hermit Creek (Hermit Creek above Tonto Trail near Grand Canyon, Arizona, station number 09403043) approximately 325 m upstream from the Tonto Trail and the NPS campground. The station monitors total spring flow in Hermit Creek including flow from Hawaii and Hermit Springs. The north-striking Hermit Fault, adjacent to Hermit Creek (pl. 1), is a prominent structural feature in this part of Grand Canyon and probably influences ground-water movement. Hanging gardens near the spring include maidenhair fern and golden columbine; sawgrass grows in a dense stand outside the grotto.

*Hermit Spring*.—Hermit Spring is about 3.6 km upstream from where the Tonto Trail crosses the Hermit Creek drainage (pl. 1). Spring water discharges from bedding planes at the base of Redwall Limestone cliffs near the contact with the Temple Butte Formation (fig. 2). The north-striking Hermit Fault, adjacent to Hermit Creek (pl. 1) is a prominent structural feature in this part of Grand Canyon. The water sample was collected at the point of issuance. Flow in Hermit Creek is continuous from the spring to the Colorado River. The USGS streamflow-gaging station (station number 09403043) was used to determine discharge (pl. 1). The vegetation in the Hermit Creek drainage is dense and includes western redbud (*Cercis occidentalis*) and tamarisk (*Tamarix chinensis*).

**Boucher East Spring**.—Boucher East Spring is in the southeast branch of Boucher Creek (pl. 1). The perennial spring discharges from a large travertine deposit associated with a slump block in the upper part of the canyon. The travertine and the slump block overlie the Bright Angel Shale (fig. 2). The strike of the Crazy Jug Monocline (pl. 1) is northwest and parallel to the trend of the southeast branch of Boucher Creek. The actual point of spring issuance is obscured; however, water first appears in the Bright Angel Shale about midway up the slump block. The point of issuance is inaccessible; therefore, water samples were collected near

the base of the travertine deposit and discharge was measured in a small channel about 5 m downstream. The southeast branch of Boucher Creek is intermittent between the spring and the confluence with the main stem of Boucher Creek. The spring supports hanging gardens.

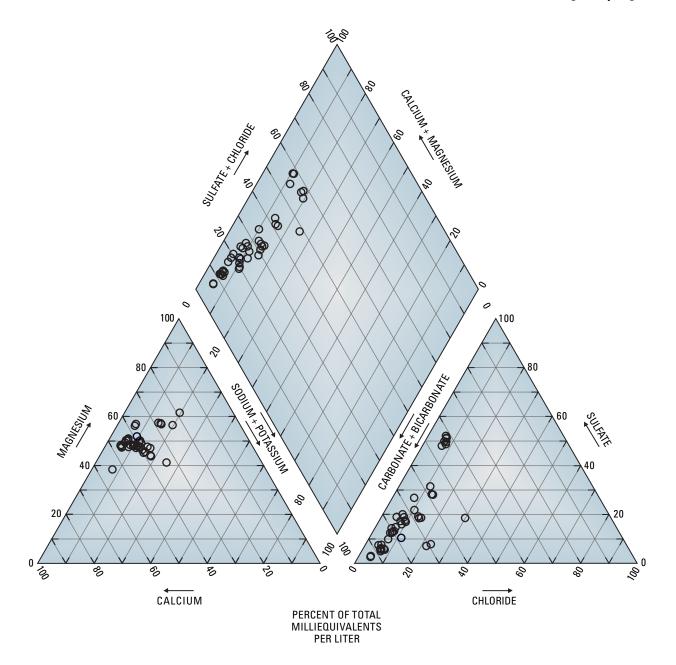
# CHEMICAL CHARACTERISTICS OF GROUND-WATER DISCHARGE AT SPRINGS

Water sampled from springs during this study was generally a calcium magnesium bicarbonate type; a few samples had a substantial sulfate component (**fig. 5** and pl. 1). To give a brief picture of the overall chemical characteristics of the springs, the concentration distribution of 18 different constituents is shown in three panels (**fig. 6**). For example, sulfate ranged from 6 to 340 mg/L, and the median value was 37 mg/L; sodium ranged from 4 to 53 mg/L, and the median value was 11 mg/L (fig. 6*A*); zinc ranged from 0.4 to 57 µg/L, and the median value was 3.7 µg/L (fig. 6*B*); and lead ranged from near zero to 260 ng/L, and the median concentration was 15 ng/L (fig. 6*C*).

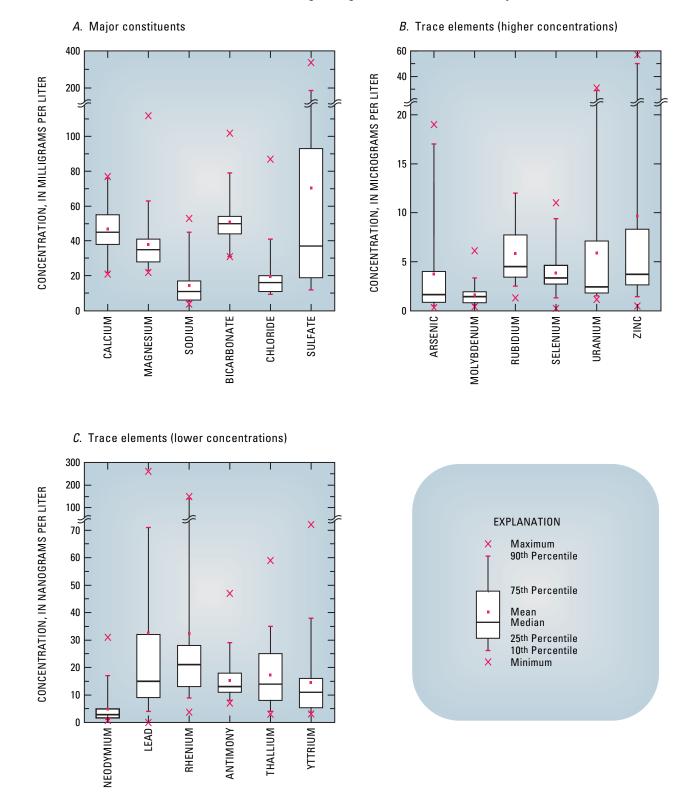
## Comparison of Water-Chemistry Data from Springs with Data from Creeks

Although collection of samples from the point of issuance was emphasized, samples were collected both from the point of spring issuance and downstream from the point of issuance in the alluvial channel at Cottonwood and Monument Creeks. Water-chemistry data for samples collected from the point of issuance differ considerably from data for samples collected from the downstream sample site (table 5).

The observed changes in chemistry in Cottonwood and Monument Creeks indicate that the chemical and physical processes occurring during the transmittal of water through the alluvium are dynamic and variable. The increase in concentration of some constituents can be attributed to dissolution of rock materials through interaction with water, and the decrease in concentration of some constituents is probably mostly due to adsorption or precipitation. The change in concentrations at Cottonwood and Monument Creeks reiterates the importance of collecting samples from the point of issuance in order to provide relevant information about the source of ground-water discharged at the spring.



**Figure 5**. Relative compositions of water from springs and creeks that discharge along the south rim of Grand Canyon, Arizona, 2000–2001.



**Figure 6**. Distribution of constituent concentrations of water that discharges from springs and creeks along the south rim of Grand Canyon, Arizona, 2000–2001. *A*, Major constituents: calcium, magnesium, sodium, bicarbonate, chloride, and sulfate; *B*, Trace elements (higher concentrations): arsenic, molybdenum, rubidium, selenium, uranium, and zinc; *C*, Trace elements (lower concentrations): neodymium, lead, rhenium, antimony, thallium, and yttrium.

#### Chemical Characteristics of Ground-Water Discharge at Springs 21

**Table 5**. Changes in constituent concentrations from the point ofspring issuance sample location to the downstream sample location,Cottonwood and Monument Creeks, 2000–2001

Changes in constituent concentrations	Constituent
Chemical constituents that increased in concentration downstream from the point of issuance	Bicarbonate, Boron, Bromide, Calcium, Chloride, DOC, Lithium, Magnesium, Molybdenum, Potassium,
Chemical constituents that decreased in concentration downstream from the point of issuance	Sodium, Strontium, Sulfate Arsenic, Cesium, Chromium, Copper, Lead, Mercury, Nitrate, Phosphorus, Vanadium
Chemical constituents that did not change in concentration between sampling sites	Rhenium, Uranium
Chemical constituents for which concentrations exhibited no common trend between sample sites	Barium, Rubidium, Selenium, Silica

### Consistency of the Springs Data

The data indicate that there is a temporal consistency of results for individual springs. For those springs that were sampled more than once, the general rule is that concentrations were consistent over the duration of the study, regardless of seasonality, weather conditions, or any other factor. For example, Miners Spring was sampled three times, in May and November 2000 and in April 2001. In these samples, the calcium concentrations ranged between 21 and 22 mg/L and arsenic concentrations ranged between 17 and 19 µg/L. Similarly, Boucher East Spring was sampled in May and December 2000 and April 2001. Sodium concentrations in these samples ranged between 6.1 and 7.0 mg/L. Salt Creek Spring was sampled during the same months, and uranium concentrations in those samples ranged between 29 and 31 µg/L. The average variability (in percent relative standard deviation) in all chemical constituents sufficiently above their respective detection limits for all springs that were sampled more than once was less than 8 percent for all but two springs, JT Spring and Grapevine East Spring (table 6). The primary conclusion from these data is that spring-water chemistry appears to be consistent over short-term (year) periods. There are two major implications that follow from this conclusion. The first is that the ground water discharging from the sites sampled had a constant chemistry during this study that can be used to characterize it. The second is that the lack of agreement among constituent data for any given site suggests that at least some of the samples from that site probably do not represent spring water at the point of issuance. Data from Grapevine East and JT Springs had far greater variability than data from other springs and fall into this category. Conversely, data from Horn Creek, Cottonwood Creek No. 2, and Pipe Creek all demonstrated consistency and indicate that although the

samples may have been collected from the channel alluvium, they likely represent water issuing from bedrock close to the sample site.

Grapevine East Spring was sampled three times (May and November 2000 and April 2001), and constituent concentrations generally varied by 40 percent or more among samples. The spring is on the slopes of a large unnamed promontory that is connected to the regional land mass of the south rim only by a narrow strip of land. Pennsylvanian and Permian rock units are mostly absent from the promontory above the spring. This spring is unique in these characteristics among the springs sampled during this study. As noted in the site descriptions, the May 2000 sample was collected in the midst of a small wetlands, unlike the samples collected later. When the data from this wetlands sample are not included, the variability improves to about 11 percent relative standard deviation, but even this value is substantially greater than the variability in data for other springs. This suggests that samples from Grapevine East Spring probably do not represent spring waters sampled at the point of issuance, or that this spring has at least two source flow paths of water that do not contribute consistently the same proportions of water. Because of this, chemical data from Grapevine East Spring were excluded from discussion in the following sections titled "Relations to Water-Quality Contaminant Standards" and "Chemistry of Selected Springs."

**Table 6**.
 Variability in chemical data for selected constituents for all springs sampled more than once during 2000–2001

[Selected constituents are arsenic, barium, bicarbonate, boron, bromide, calcium, chloride, lithium, magnesium, molybdenum, potassium, rhenium, rubidium, selenium, silica, sodium, strontium, sulfate, and uranium]

Spring	Average percent relative standard deviation for selected chemical constituents	Number of times sampled
JT Spring	23.1	2
Miners Spring	4.2	3
Cottonwood Creek No. 2	2.1	2
Grapevine Main Spring	1.2	2
Grapevine East Spring	45.3	3
Grapevine East Spring <sup>1</sup>	10.8	2
Lonetree Spring	3.2	2
Burro Spring	4.5	3
Pipe Creek	6.9	3
Pumphouse Spring	5.0	3
Horn Creek	5.3	3
Salt Creek Spring	4.0	3
Monument Spring	1.3	2
Hawaii Spring	3.1	3
Hermit Spring	2.3	2
Boucher East Spring	7.2	3

<sup>1</sup>The second entry for Grapevine East Spring excludes the sample collected in the wetlands.

JT Spring was sampled in April and May 2001, and constituent concentrations varied by about 23 percent between samples. In general, the April 2001 sample had higher constituent concentrations than the May 2001 sample.

Immediately prior to collection of the sample in April 2001, the site received sustained rainfall. This sample, therefore, could have been affected by local surface-water runoff. Consequently, this sample was also excluded from the ensuing discussion.

These factors underscore the need to collect samples at more than one time from each spring. Good agreement of results suggests that samples represent the ground water at or near the point of issuance, whereas lack of agreement suggests either that one or more of the samples was compromised, or that the samples do not represent the ground water at the point of issuance.

# Relations to Water-Quality Contaminant Standards

Concentrations of several constituents at some springs were higher with respect to the other springs in the study area and (or) exceeded U.S. Environmental Protection Agency (USEPA) Maximum Contaminant Levels for drinking water (MCLs; table 7). MCLs are based on quantity of consumption and are intended for public drinking-water supply systems and not the occasional user (U.S. Environmental Protection Agency, 2003). Nevertheless, values above the MCLs represent a concern for water users.

Arsenic concentrations in samples from Red Canyon Spring, JT Spring, and Miners Spring exceeded the MCL of 10  $\mu$ g/L (tables 7 and 8), averaging between 14 and 19  $\mu$ g/L. Cottonwood Creek No. 2 and Grapevine Main Spring had concentrations of about 5  $\mu$ g/L. All these springs are in the easternmost part of the study area (pl. 1). Other springs in the study area averaged about 2  $\mu$ g/L of arsenic or less, with the exception of Monument Spring, which averaged 4  $\mu$ g/L (fig. 7). This pattern of the easternmost springs having the highest concentrations was also seen for cesium, thallium, tungsten, vanadium, and to a lesser extent, antimony and nitrate.

Concentrations of selenium ranged from 1.3 to about 11  $\mu$ g/L; most springs had concentrations of approximately 4  $\mu$ g/L (**fig. 8**). Although these values are well below the MCL (50  $\mu$ g/L), many other countries have adopted a standard of 10  $\mu$ g/L (National Health and Medical Research Council, and Agriculture and Resource Management Council of Australia and New Zealand, 1996; Safe Drinking Water Foundation, 2003; Drinking Water Inspectorate, 2003). Concentrations of selenium were highest at Salt Creek Spring (9.4 to 11  $\mu$ g/L). With the exception of data from Salt Creek Spring, there is a decreasing trend in concentrations east to west from Lonetree Spring (6  $\mu$ g/L) to Boucher East Spring (about 3  $\mu$ g/L).

Concentrations of uranium did not show a regional pattern; however, values at some sites were significantly higher than at others (fig. 9). Uranium concentrations were about 30  $\mu$ g/L (equal to the MCL) in all samples from Salt Creek Spring (table 7). Salt Creek Spring also had the highest

dissolved gross alpha radioactivity (22 pCi/L, based on  $^{230}$ Th). Only one sample from Horn Creek (April 2001) was near the MCL for uranium; the other two samples had values of about 9 µg/L. Uranium mining has occurred in the Horn Creek drainage, and unacceptably high levels of uranium were reported for samples collected previously (Fitzgerald, 1996). In Fitzgerald's study, samples were taken not at the point of spring issuance, but where the Tonto Trail crosses the drainage. The uranium in south rim springs may be related to the presence of nearby breccia pipes and faults.

### Geochemistry of Barium and Sulfate

Concentrations of sulfate varied widely—between 6 and 190 mg/L when suspected compromised samples have been eliminated (figs. 5 and 6). This variation is not seen for the other major ions. When sulfate data are plotted against barium data in molar concentrations (**fig. 10**), an inverse relation is observed, indicating that barium and sulfate concentrations are controlling each other. Barium sulfate—barite (BaSO<sub>4</sub>)— solubility is typically a control on sulfate concentration in natural waters (Hem, 1985). Figure 10 indicates that sulfate concentrations are limited by barite solubility for the samples in this study.

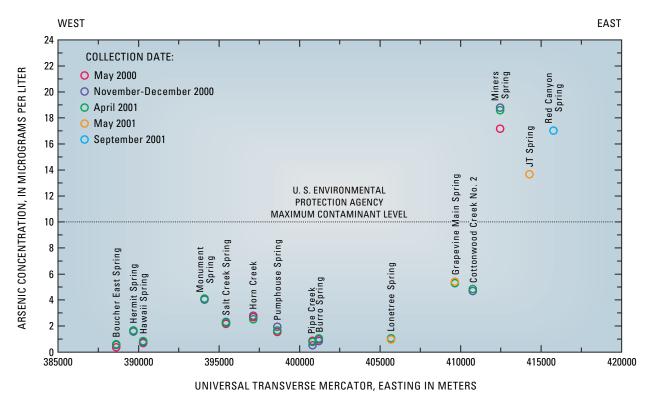
A plot of the spatial distribution of barium concentrations at sites in the study area reveals two populations that are spatially intermixed but can be distinguished from each other by concentrations (fig. 11). One group contains concentrations lower than about 130 µg/L and consists of JT Spring, Miners Spring, Lonetree Spring, Burro Spring, Pipe Creek, Horn Creek, Salt Creek Spring, Hawaii Spring, and Boucher East Spring. The other population consists of those springs whose barium concentrations are higher than about 200 µg/L and contains Red Canyon Spring, Cottonwood Creek No. 2, Grapevine Main Spring, Pumphouse Spring, Monument Spring, and Hermit Spring. If the easternmost four sites (Red Canyon Spring, JT Spring, Miners Spring, and Cottonwood Creek No. 2) are separated from these two populations because of the arsenic, cesium, thallium, vanadium, and tungsten data (see previous discussion on arsenic) to create a third population, a set of clear intrapopulation relations becomes apparent. Population 1 consists of Lonetree Spring, Burro Spring, Pipe Creek, Horn Creek, Salt Creek Spring, Hawaii Spring, and Boucher East Spring, and shows a linear increase in barium to the west from about 40 µg/L at Lonetree Spring to about 110 µg/L at Boucher East Spring. Population 2 consists of Grapevine Main Spring, Pumphouse Spring, Monument Spring, and Hermit Spring, and shows a linear decrease in barium to the west from about 350 µg/L at Grapevine Main Spring to about 220 µg/L at Hermit Spring. Population 3, consisting of Red Canyon Spring, JT Spring, Miners Spring, and Cottonwood Creek No. 2, shows no spatial barium relation, but shows a strong spatial arsenic relation (fig. 7).

#### Table 7. U.S. Environmental Protection Agency Maximum Contaminant Levels (MCLs) for selected constituents

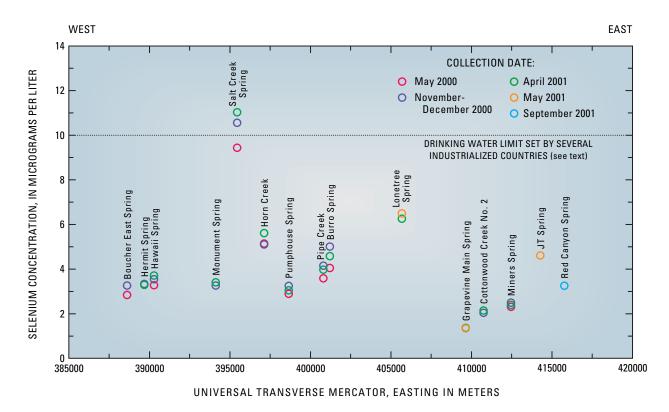
[Values exceeding the Maximum Contaminant Level (MCL) are italicized. Source of MCLs: U.S. Environmental Protection Agency (2003)]

Trace element	MCL, micrograms per liter	Potential health effects from ingestion of water	Maximum concentrations measured in this study
Antimony	6	Increase in blood cholesterol; decrease in blood sugar	0.047 in JT Spring
Arsenic	10	Skin damage; problems with circulatory systems; increased risk of cancer	19 in Miners Spring; 17 in Red Canyon Spring; 14 in JT Spring; 5.4 in Grapevine Main Spring; 4.8 in Cottonwood Creek No. 2; 4.0 in Monument Spring
Barium	2,000	Increase in blood pressure	359 in Grapevine Main Spring
Beryllium	4	Intestinal lesions	0.008 in Horn Creek and Salt Creek Spring
Cadmium	5	Kidney damage	0.020 in Pipe Creek
Chromium	100	Allergic dermatitis	0.9 in Red Canyon Spring
Lead <sup>1</sup>	15	Infants and children: delays in physical and mental development; children could show slight deficits in attention span and learning disabilities	0.26 in Monument Spring
Mercury	2	Kidney damage	0.002 in Monument Spring
Nitrate as nitrogen	10,000	Infants below the age of six months who drink water in excess of MCL could become seriously ill and if untreated may die. Symptoms include shortness of breath and blue-baby syndrome	6,800 in Monument Spring; 4,500 in Monument Creek No. 1
Selenium	50	Hair or fingernail loss; numbness in fingers or toes; circulatory problems	11 in Salt Creek Spring
Thallium	2	Hair loss; changes in blood; kidney, intestine or liver problems	0.059 in Monument Spring
Uranium	30	Increased risk of cancer; kidney toxicity	31 in Salt Creek Spring; 29 in Horn Creek

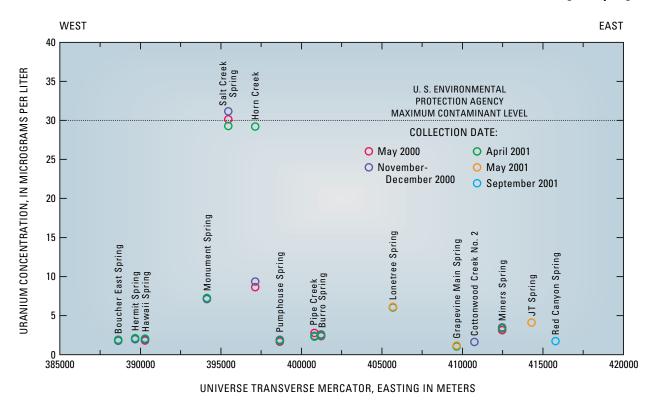
<sup>1</sup>Action level. For public water systems, lead is regulated by a treatment technique that requires systems to control the corrosiveness of their water.



**Figure 7**. Concentrations of arsenic in relation to spring or creek locations along the south rim of Grand Canyon, Arizona, 2000–2001.



**Figure 8**. Concentrations of selenium in relation to spring or creek locations along the south rim of Grand Canyon, Arizona, 2000–2001.



**Figure 9**. Concentrations of uranium in relation to spring or creek locations along the south rim of Grand Canyon, Arizona, 2000–2001.

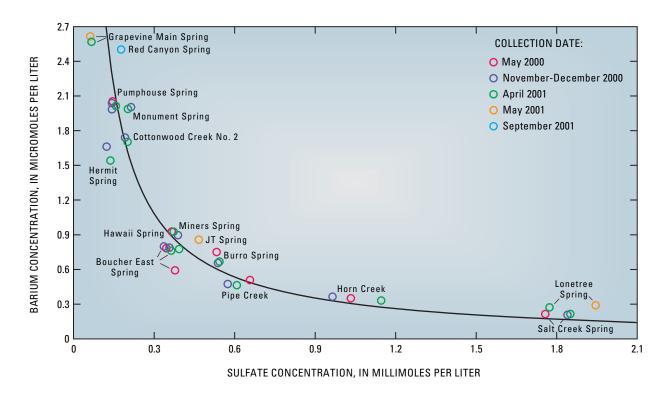


Figure 10. Relation between concentrations of barium and sulfate in water that discharges along the south rim of Grand Canyon, Arizona, 2000–2001.

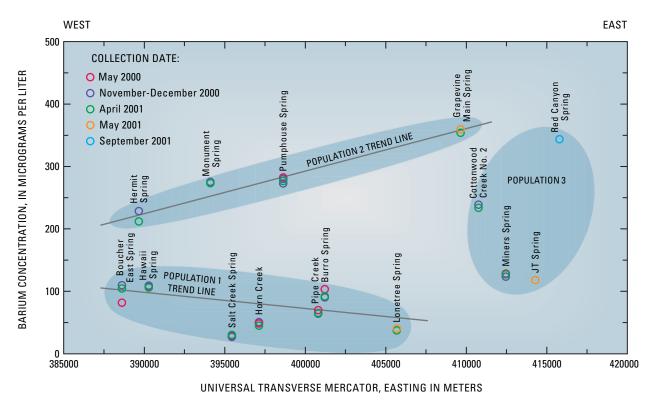


Figure 11. Concentrations of barium in relation to spring or creek locations along the south rim of Grand Canyon, Arizona, 2000–2001.

The division of the springs into these three populations (based on barium and arsenic concentrations) is also observed, although less clearly, for calcium, magnesium, sulfate, and potassium. If Monument Spring is eliminated from this discussion, an action justified by the fact that it is anomalous with respect to the other springs in this study in many ways (see below), chloride and sodium also demonstrate the division.

## **Chemistry of Selected Springs**

#### Monument Spring

Although much of the chemistry of Monument Spring is similar to that of surrounding springs, concentrations of several constituents are distinctly different. At the point of issuance, concentrations of nitrate ranged from 5.7 to 6.8 mg/L as nitrogen, which are less than the MCL of 10 mg/L (table 7) but much higher than the concentrations at any other springs. The source of the nitrate in the spring is unknown. Monument Spring also had higher concentrations of chloride and several heavy metals, including copper, mercury, lead, arsenic, and thallium, than other springs. In addition, semiquantitative analyses of these samples indicated concentrations of dissolved gold that were an order of magnitude greater than concentrations at other springs. These patterns indicate that ground water discharging at Monument Spring is influenced either by wastewater contamination on the south rim of Grand Canyon or by exposure to rocks and (or) geologic features in the subsurface (for example, breccia pipes or faults) not encountered by water discharging at the other springs.

# Lonetree Spring

Water issuing from Lonetree Spring is unusual in that it has much higher concentrations of the rare earth elements (REE) than the other springs. The average concentration for lanthanum in all waters analyzed was about 0.004  $\mu$ g/L, yet for Lonetree Spring it was 0.023  $\mu$ g/L; similarly, the average concentration for holmium was about 0.0003  $\mu$ g/L, yet for Lonetree Spring it was 0.0016  $\mu$ g/L. In addition to the REE, Lonetree Spring was unusually high in sodium, calcium, lithium, rubidium, strontium, boron, and sulfate, and was unusually low in antimony and barium. There are no known sources of REE in the vicinity of Lonetree Spring.

## Salt Creek Spring and Horn Creek

Salt Creek Spring had high concentrations of uranium and selenium, and had a high gross alpha radioactivity, and one of the Horn Creek samples had a high concentration of uranium. These springs also had anomalously high concentrations (relative to the other springs) of magnesium, potassium, sulfate, lithium, bromide, molybdenum, and

#### Chemical Characteristics of Ground-Water Discharge at Springs 27

rhenium. These sites are near mapped breccia pipes and an abandoned copper and uranium mine (**pl. 1**), which may influence the chemistry of water that discharges at the two springs.

# Grapevine Main, Pumphouse, Hawaii, Hermit, and Boucher East Springs

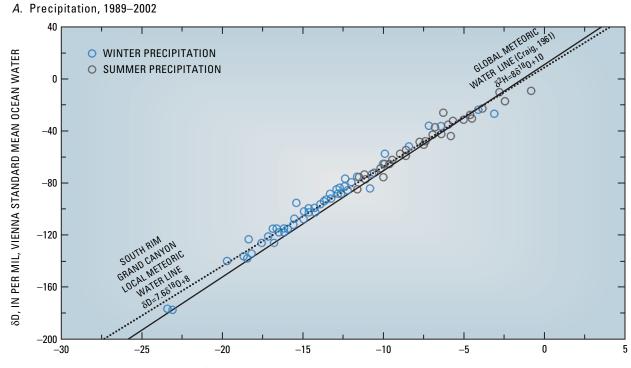
Grapevine Main Spring, Pumphouse Spring, Hawaii Spring, Hermit Spring, and Boucher East Spring tended to have low concentrations of a wide variety of major and trace elements and nutrients, including calcium, magnesium, sodium, potassium, strontium, lithium, rubidium, boron, selenium, uranium, antimony, and nitrate. To a lesser extent, this is also true of Pipe Creek, Burro Spring, and Cottonwood Creek No. 2. These springs, however, can be distinguished from one another on the basis of their chemistry. For example, a plot of barium and calcium concentrations (fig. 12) demonstrates a clear clustering of data for each spring. If more elements were added, the clustering would be even more exclusive: for example, if arsenic were plotted as a third dimension, the separation between Hermit Spring and Cottonwood Creek No. 2 would be greater, because their respective arsenic concentrations are very different. Thus, although each of these springs has, in general, a similar chemistry—derived, no doubt, from the same aquifer—each also has its own particular chemistry that is dependent on its geological and spatial location.

## Residence Times of Ground Water Discharging at Springs

A local meteoric water line for the south rim of Grand Canyon (**fig. 13***A*) was constructed using  $\delta^{18}$ O and  $\delta$ D analyses of precipitation samples collected between 1989 and 2002 from a site near the south rim of Grand Canyon (Pendall, 1997; Harvey, 2000; National Atmospheric Deposition Program, 2002). The  $\delta^{18}$ O and  $\delta$ D data for precipitation have a strong seasonal pattern: winter precipitation samples were isotopically lighter than summer precipitation samples. Delta oxygen-18 values range from -12.9 to -8.5 %, and  $\delta$ D values range from -95 to -71 % for water samples collected from sites that discharge from the Redwall-Muav Limestone aquifer along the south rim of Grand Canyon (**table 8**).

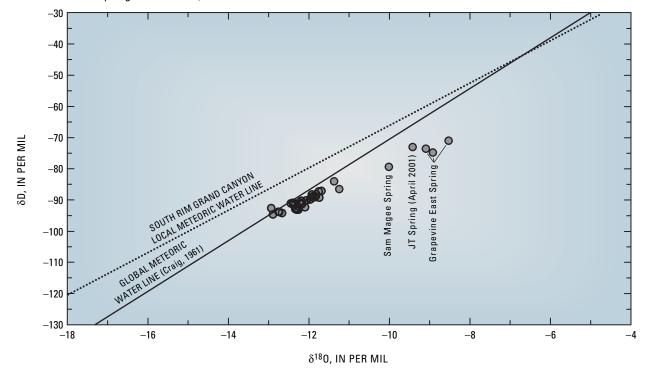


Figure 12. Relation between barium and calcium concentrations, demonstrating the clustering of data, for selected springs and creeks that discharge along the south rim of Grand Canyon, Arizona, 2000–2001.



 $\delta^{18}\text{O},$  In Per Mil, vienna standard mean ocean water

B. Water from springs and creeks, 2000-2001



**Figure 13**. Relation between oxygen-18 ( $\delta^{18}$ O) and deuterium ( $\delta$ D) in water along the south rim of Grand Canyon, Arizona. *A*, Precipitation 1989–2002; *B*, Water from springs and creeks, 2000–2001; *C*, Regional trend of oxygen-18 ( $\delta^{18}$ O) values for water from springs and creeks, 2000–2001.



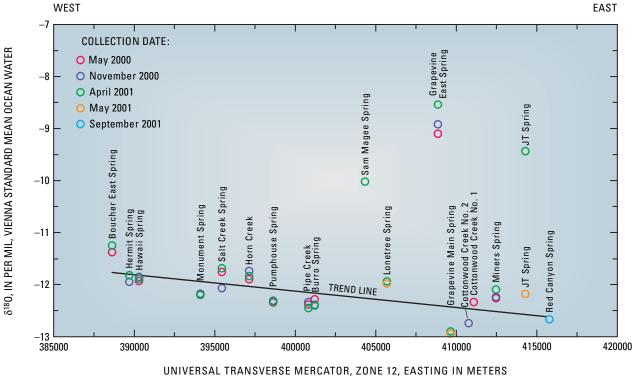


Figure 13. Continued.

The local meteoric water line plots above the global meteoric water line, and data from the springs and creeks plot close to the global meteoric water line (fig. 13*B*). Possible reasons for this difference are the contribution of older water that was recharged during a different climate period or fractionation during secondary evaporation prior to recharge (Mazor, 1997).

Unlike precipitation, water samples from springs and creeks along the south rim of Grand Canyon do not show a clear seasonal pattern in isotopic compositions and are most similar to compositions in winter precipitation. Multiple water samples for  $\delta^{18}$ O and  $\delta$ D analyses were collected at most sites during multiple seasons, and data from repeat samples were generally consistent with data from the initial sampling. Values of  $\delta^{18}$ O and  $\delta$ D for a repeat sample at JT Spring, however, were substantially different from values for the initial sample. The  $\delta^{18}$ O and  $\delta$ D data can be divided into two groups. Group 1  $\delta^{18}$ O values range from -12.9 to -11.4 % and show a trend of increasing <sup>18</sup>O enrichment from east to west (fig. 13C). Group 1 includes Red Canyon Spring, JT Spring, Miners Spring, Cottonwood Creek No. 1, Cottonwood Creek No. 2, Grapevine Main Spring, Lonetree Spring, Burro Spring, Pipe Creek, Pumphouse Spring, Horn Creek, Salt Creek Spring, Monument Spring, Hawaii Spring, Hermit Spring, and Boucher East Spring. Group 2 includes JT Spring, Grapevine East Spring, and Sam Magee Spring. Group 2 sites are substantially enriched in <sup>18</sup>O and D and have  $\delta^{18}$ O values that range from -10.0 to -8.5 %. The trend of isotopic enrichment

for group 2 sites is probably due to evaporative processes occurring prior to ground-water recharge or subsequent to spring discharge upgradient from sample collection points.

Strontium isotopes and concentrations can provide information about water-rock interactions in ground-water systems. Strontium-87/strontium-86 values (table 8) for water samples from springs and creeks range from 0.70363 at Horn Creek to 0.71514 at Grapevine East Spring. Rock samples collected from the Bright Angel Fault area have <sup>87</sup>Sr/<sup>86</sup>Sr values (table 9) ranging from 0.70820 (Redwall Limestone) to 0.71216 (Hermit Formation). Values for springs and creeks east of the Bright Angel Fault (pl. 1) are generally more radiogenic than values for sites to the west of the fault (fig. 14). Values in water discharging from Red Canyon Spring, JT Spring, Miners Spring, Cottonwood Creek No. 1, Cottonwood Creek No. 2, Grapevine East Spring, Grapevine Main Spring, Lonetree Spring, Burro Spring, Pipe Creek, Pumphouse Spring, Salt Creek Spring, Hawaii Spring, and Boucher East Spring are more radiogenic than values in water that discharges at the other springs. The larger ratios may be due to interactions of ground water with the Coconino Sandstone or Hermit Formation, which are more radiogenic than the other rock units, to contact with secondary carbonates encountered in fractures or solution cavities along the flow path, or to interaction with rocks that were not analyzed during this study.

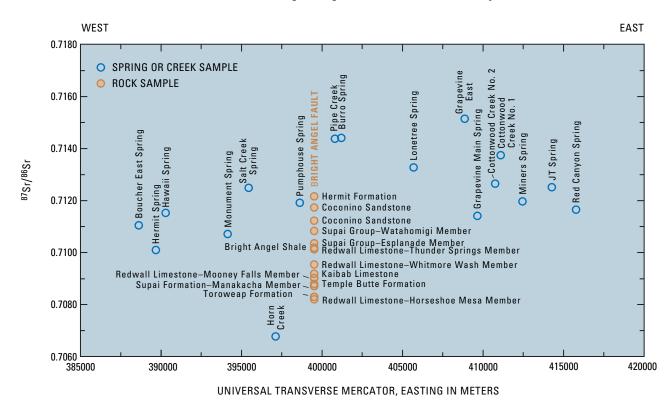


Figure 14. Strontium-87/strontium-86 in water from springs and creeks that discharge along the south rim of Grand Canyon, Arizona, and in rocks from near the Bright Angel Fault near the south rim, 2000–2001.

Water in Horn Creek had a <sup>87</sup>Sr/<sup>86</sup>Sr value of 0.70363. None of the rocks sampled had a value similar to this value. It is possible that rock units present in the ground-water flow path near Horn Creek are different from those sampled and analyzed for <sup>87</sup>Sr/<sup>86</sup>Sr values.

Carbon-14 activities in water samples ranged from 38.3 pmc at Hawaii Spring to 103.4 pmc at Grapevine East Spring and for most sites was between about 40 and 75 pmc (table 8). Ground-water residence times were estimated by Stuiver and Polach's (1977) method and by application of a combination of Stuiver and Polach's (1977) conventional radiocarbon-age method and Pearson and Hanshaw's (1970) dilution equation. Estimates of ground-water residence time using this method range from modern to 3,400 yr (table 8). Water-rock interaction in carbonate rock can cause a change of about 60 pmc from the recharged water (Mazor, 1997).

Tritium activities in water samples from sites in this study ranged from less than 0.3 TU (detection limit) to 4.1 TU (table 8). Water samples were collected for tritium analysis during more than one site visit at most sites. Tritium activities in samples collected during April 2001 were higher than activities in samples collected at other times at the same sites. A regional rainstorm during the April 2001 sample collection period could have caused this variability.

Three groups were identified on the basis of tritium and <sup>14</sup>C data (**fig. 15**). Sites in group 1, Pumphouse and Boucher East Springs, had tritium activities below the detection limit (0.3 TU) and estimated ground-water residence times ranging

from modern to 1,900 yr. Low tritium activities indicate that ground-water discharge at the group 1 sites did not contain young water at the time samples were collected. The corrected residence time for Boucher East Spring was modern; however, extensive travertine deposits at this spring could be a source of unaccounted-for carbon in the ground-water flow path, which would result in a  $^{14}$ C value that is biased toward a shorter residence time.

Group 2 includes sites with tritium activities that ranged from 0.4 to 2.0 TU and <sup>14</sup>C activities that ranged from 38.3 to 74.8 pmc. There is a general increase in TU with increasing pmc. Estimates of ground-water residence time for this group range from 500 to 3,400 yr. The 14 sites in this group discharge a mixture of young and old ground water. When a mixture of young and old water occurs, the actual residence time is greater than the calculated residence time (Mazor, 1997). Group 3 is composed of sites with predominantly post-bomb (after 1952; Mazor, 1997) carbon. Sites in this group are Grapevine East Spring (TU = 0.6; pmc = 103.4 and 98.7) and Cottonwood Creek No. 1 (TU = 0.7; pmc = 94.2). Carbon-14 activities higher than 100 pmc indicate the presence of predominantly post-bomb radiogenic carbon in ground water. A water sample collected at Grapevine East Spring (103.4 pmc) in May 2000 was the only sample with post-bomb carbon.

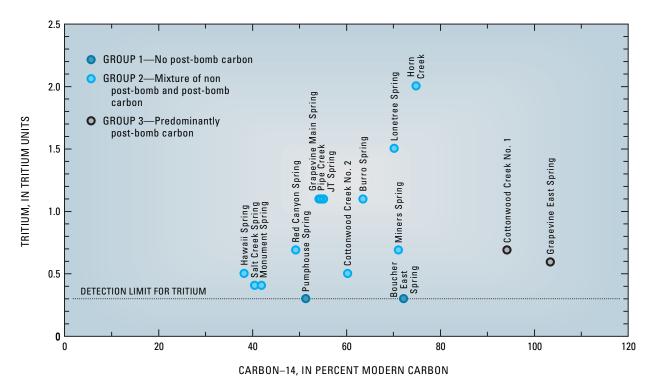


Figure 15. Relation between tritium and carbon-14 in water from springs and creeks that discharge along the south rim of Grand Canyon, Arizona, 2000–2001.

# **FUTURE STUDIES**

The ongoing collection and evaluation of water-chemistry data will be a valuable component of GRCA's effort to document springs and other water sources in Grand Canyon. Continued operation of the three streamflow-gaging stations on Hermit Creek, Cottonwood Creek, and at Pumphouse Spring will provide long-term data on ground-water discharge. Additional gaging stations will improve the areal distribution of discharge data and provide long-term data that can be used to evaluate the effects of development and climate variability on ground-water resources.

The lack of water samples from wells on the Coconino Plateau has limited the data available for analysis of groundwater residence times and flow paths. Additional data from wells and from springs discharging from younger stratigraphic units are needed to adequately determine ground-water residence times and flow paths for the regional ground-water system on the Coconino Plateau. Collection and analysis of water samples from wells, repeat sampling and analysis at spring and creek sites included in this study, and expansion to include springs issuing from the Coconino Sandstone, Supai Group, and Precambrian rocks would help to provide additional understanding of the relations between the springs and local and regional ground-water flow systems. Collection of inner canyon precipitation samples would help improve understanding of recharge and ground-water flow paths. Collection and analysis of rock samples from geologic units and areas of Grand Canyon not included in this study would help to refine existing knowledge of regional water-rock interactions.

Data describing springs along the south rim of Grand Canyon and wells on the Coconino Plateau have been collected by the USGS and other researchers as part of other projects and were not included in this report. A comprehensive analysis of all available data would be an important contribution to the understanding of the regional ground-water system.

# SUMMARY AND CONCLUSIONS

Springs discharging along the south rim of Grand Canyon in Grand Canyon National Park, and the riparian areas that they support, are recognized as important resources by resource managers, Native Americans, and the general public. Development on the Coconino Plateau in the vicinity of the south rim has led to an increased demand for ground water. This study was designed to provide baseline data on springs and creeks issuing from the Redwall-Muav Limestone aquifer in the core-use area along the south rim, from Red Canyon Spring in the east to Boucher East Spring in the west. The data

in this report contain information on 20 springs and creeks including water chemistry, quantity, and quality, and the residence times of ground-water discharged at the springs.

Whenever possible, samples were collected at the point where water discharges directly from bedrock. Where this could not be done, samples were collected at the point where water first appeared. Each spring and creek is associated with one or more of five types of geologic characteristics: bedding planes, fractures, channel alluvium, hillslope alluvium, or travertine deposits. Most of the sites are associated with geologic structures such as faults or monoclines.

Major findings of this study:

- The water chemistry of a given spring did not appreciably vary over the course of the study. Samples had consistent compositions regardless of the season in which they were collected. This was in direct contrast to data presented in previous reports showing considerable seasonal variation in chemistry. Samples in previous studies may not have been collected near the point of the first issuance of water. Consequently, these data may not be representative of water discharging directly from the Redwall-Muav Limestone aquifer.
- The water chemistry of each spring was distinct from that of all other springs and creeks.
- The water discharging at most sites was a calcium magnesium bicarbonate type.
- Arsenic and uranium were higher than USEPA MCLs at some sites.
  - Arsenic concentrations were highest at springs near the Grandview-Phantom Monocline; Miners, JT, and Red Canyon Springs all contained concentrations above the MCL.
  - Uranium concentrations and gross alpha radioactivity were highest at Horn Creek and Salt Creek Spring. The average uranium concentration at Salt Creek Spring was equal to the MCL.
- Nitrate and selenium approached the MCLs or exceeded standards set in other countries at some sites.
  - Nitrate concentrations were highest in the Monument Creek drainage, and those samples collected at the spring were the highest in the drainage.
  - Selenium concentrations at Salt Creek Spring were the highest in the study. Whereas these concentrations do not approach the MCL, they exceed the drinking-water limit set by several other countries.
- The chemistry of Monument Spring was much different from that of the other springs in many respects. In addition to having the highest concentrations of nitrate, this spring also had high

concentrations of copper, mercury, lead, thallium, arsenic, chloride, and gold relative to other springs.

- The chemistry of Lonetree Spring was distinctively high in boron, calcium, lithium, rubidium, sodium, strontium, sulfate, and the REEs.
- The δ<sup>18</sup>O and δD data vary little among most sites and among multiple samples from individual sites, indicating that the water discharging from most sites is from a common aquifer.
- The δ<sup>18</sup>O and δD data for most sites indicate that recharge occurs on the Coconino Plateau; however, the presence of an evaporative signature for some sites suggests isotopic fractionation due to predischarge or postdischarge evaporation.
- The <sup>87</sup>Sr/<sup>86</sup>Sr values for water samples collected at sites east of the Bright Angel Fault are more radiogenic than values for samples from sites west of the fault.
- Ground-water residence times estimated using radiocarbon dating techniques ranged from modern to about 3,400 yr.
- Tritium and <sup>14</sup>C results indicate that ground-water discharged from the Redwall-Muav Limestone aquifer at most springs and creeks is a mixture of young and old ground waters, suggesting that water discharging from the aquifer at these sites follows multiple flow paths and has multiple recharge areas.

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# **Physical and Chemical Data**

**Table 8**. Physical and chemical properties, chemical-constituent concentrations, radioactivity, and isotopic data for water from springs and creeks that discharge along the south rim of Grand Canyon, Arizona, 2000–2001

[°C, degrees Celsius;  $\mu$ S/cm, microsiemens per centimeter at 25°C; mg/L, milligrams per liter,  $\mu$ g/L, micrograms per liter; gal/m, gallons per minute; L/m, liters per minute; nc, not collected; ICP-AES, Inductively Coupled Plasma-Atomic Emission Spectrometry; Avg, average; SD, standard deviation; IC, ion chromatography; MAD, median absolute deviation; na, not available; <, less than; ICP-MS, Inductively Coupled Plasma-Mass Spectrometry; CVAFS, Cold-Vapor Atomic Fluorescence Spectrometry; pCi/L, picocuries per liter; PE, two-sigma precision estimate; MDC, minimum detectable concentration; <sup>137</sup>Cs, Cesium-137 curve; <sup>230</sup>Th, Thorium-230 curve; mg C/L, milligrams of carbon per liter, IR infrared;  $\delta$ , delta; per mil, per thousand; TU, tritium units; <, less than or equal to]

			Properties of	selected spi	rings and creek	S			
Site	Date	Water temperature (°C)	Air temperature (°C)	pH (standard units)	Specific conductance (µS/cm)	Dissolved Oxygen (mg/L)	Alkalinity dissolved incremental field (meq/L)	Discharge (gal/m)	Discharge (L/m)
Red Canyon Spring	08-23-01	18.5	27.7	7.95	360	7.4	2.80	3.2	12
	09-26-01	17.0	21.7	8.16	293	7.2	2.60	3.0	11
JT Spring	04-08-01	8.7	6.4	7.60	498	9.8	4.36	.17	.6
	05-11-01	18.3	32.5	8.20	420	8.1	2.86	.91	3.4
Miners Spring	05-24-00	15.8	32.2	8.75	402	11.4	2.60	.54	2.0
	11-28-00	11.0	9.0	7.85	375	9.4	2.80	.49	1.9
	04-07-01	10.2	13.5	7.94	367	12.3	2.58	.59	2.2
Cottonwood Creek No. 1	05-25-00	17.2	nc	7.28	880	nc	8.48	.81	3.1
Cottonwood Creek No. 2	11/29-00	12.0	9.0	7.86	444	7.6	4.22	4.5	17
	04-09-01	11.5	10.3	7.87	421	8.9	3.25	nc	nc
Grapevine Main Spring	04-10-01	6.2	12.1	8.20	347	10.1	3.65	2.0	7.5
	04-30-01	12.5	22.6	8.10	370	10.0	3.71	1.9	7.2
Grapevine East Spring	05-25-00	20.8	29.4	8.27	915	5.3	6.34	.18	.68
	11-29-00	6.0	11.0	8.47	1,295	11.0	6.28	1.1	4.2
	12-12-00	8.5	12.0	8.53	1,245	nc	6.28	1.1	4.2
	04-09-01	11.2	15.5	8.53	1,202	9.7	6.19	.44	1.7
Lonetree Spring	04-11-01	9.0	9.3	7.80	659	10.0	3.07	.58	2.2
	05-01-01	11.7	27.3	7.70	660	7.9	3.25	.46	1.7
Sam Magee Spring	04-20-01	19.2	20.5	8.64	531	7.5	2.57	.40	1.5
Burro Spring	05-22-00	15.3	34.6	8.27	584	10.5	4.78	3.1	12
	12-07-00	11.0	8.5	8.11	536	9.4	4.46	4.0	15
	04-08-01	12.8	13.4	7.78	559	9.2	4.37	4.4	17
Pipe Creek	05-22-00	22.1	33.2	7.77	614	7.0	4.98	4.0	15
	12-07-00	13.0	10.0	7.38	588	9.7	4.90	12	44
	04-08-01	13.0	7.0	7.28	554	5.7	4.35	9.8	37
Pumphouse Spring	05-22-00	19.5	33.9	8.24	503	7.9	4.28	.45	1.7
	12-07-00	14.5	12.0	8.28	424	7.9	3.96	.90	3.4
	04-07-01	15.5	12.2	8.02	398	7.8	4.20	1.0	3.6
Pumphouse Wash Gage	12-07-00	18.6	12.0	8.06	415	7.4	3.86	45	170
	12-09-00	18.5	12.3	nc	nc	nc	nc	45	170
Horn Creek	05-22-00	17.9	32.5	8.17	668	9.8	4.22	.90	3.4
	12-06-00	4.5	11.8	7.71	626	8.6	3.94	1.3	5.1
	04-07-01	11.7	15.3	7.46	702	8.0	4.13	2.5	9.5

		Prope	rties of select	ed springs a	nd creeks—Co	ntinued			
Site	Date	Water temperature (°C)	Air temperature (°C)	pH (standard units)	Specific conductance (µS/cm)	Dissolved Oxygen (mg/L)	Alkalinity dissolved incremental field (meq/L)	Discharge (gal/m)	Discharge (L/m)
Salt Creek Spring	05-23-00	18.7	31.2	8.46	706	10.1	3.30	.45	1.7
	12-06-00	7.2	6.0	8.22	656	9.8	3.06	.90	3.4
	04-10-01	7.1	8.8	8.20	638	10.5	3.01	1.3	4.9
Monument Creek No. 1	05-24-00	20.0	33.0	7.86	818	9.0	4.28	54	204
Monument Spring	12-05-00	18.0	nc	7.55	533	8.1	4.02	58	221
	04-09-01	17.8	15.6	7.66	585	8.0	3.38	89	337
Hawaii Spring	05-25-00	18.0	27.3	8.15	410	10.2	4.36	359	1,360
	12-04-00	15.7	7.2	7.75	368	7.7	4.22	319	1,210
	04-11-01	17.6	11.5	7.72	487	7.6	4.08	399	1,510
Hermit Spring	12-04-00	11.4	nc	8.36	297	nc	4.04	319	1,210
	04-11-01	10.6	10.1	8.16	417	9.4	3.92	399	1,510
Boucher East Spring	05-26-00	18.8	28.0	8.58	476	9.5	4.24	3.1	12
	12-04-00	9.2	10.0	8.41	479	9.8	4.28	6.3	24
	04-12-01	11.5	14.3	8.21	464	10.1	3.94	7.8	30

			Major ion con	centrations				
Site	Date	Bicarbonate, dissolved incremental field (meq/L)	Carbonate, dissolved incremental field (meq/L)	Alkalinity, dissolved titrimetry lab (meq/L)	Bicarbonate, dissolved titrimetry lab (meq/L)	Carbonate, dissolved titrimetry lab (meq/L)	Calc dissc (mg/L ICP-	olved
		(IIIEq/L)	(11164/12)	(1164/1)	(iiieq/L)	(1164/1)	Avg	SD
Red Canyon Spring	08-23-01	2.66	0.14	nc	nc	nc	nc	nc
	09-26-01	2.56	0.04	2.76	2.72	0.04	27	3
JT Spring	04-08-01	4.35	0.02	4.32	4.30	0.02	63	1
	05-11-01	2.80	0.06	2.84	2.80	0.04	34	1
Miners Spring	05-24-00	2.38	0.21	2.60	2.47	0.14	22	0
	11-28-00	2.75	0.03	2.49	2.48	0.02	21	0
	04-07-01	2.37	0.21	2.53	2.51	0.02	22	1
Cottonwood Creek No. 1	05-25-00	8.46	0.02	8.71	8.69	0.02	76	1
Cottonwood Creek No. 2	11-29-00	4.18	0.03	4.17	4.14	0.03	42	0
	04-09-01	3.23	0.02	4.19	4.16	0.03	42	1
Grapevine Main Spring	04-10-01	3.61	0.13	3.75	3.69	0.11	36	1
	04-30-01	3.69	0.19	4.01	3.96	0.19	35	1
Grapevine East Spring	05-25-00	6.20	0.19	6.23	6.12	nc	77	1
	11-29-00	6.10	0.20	6.73	6.54	0.23	55	0
	12-12-00	6.10	0.04	nc	nc	0.06	nc	nc
	04-09-01	5.98	0.02	7.06	6.83	0.05	76	0
Lonetree Spring	04-11-01	3.06	0.01	3.14	3.12	0.02	57	1
	05-01-01	3.24	0.01	3.75	3.73	0.02	60	2

		IVI	ajor ion concentra		u			
Site	Date	Bicarbonate, dissolved incremental field (meq/L)	Carbonate, dissolved incremental field (meq/L)	Alkalinity, dissolved titrimetry lab (meq/L)	Bicarbonate, dissolved titrimetry lab (meq/L)	Carbonate, dissolved titrimetry lab (meq/L)	disso (mg/L	cium, olved as Ca) -AES
				· · · · ·			Avg	SD
Sam Magee Spring	04-20-01	2.46	0.10	nc	nc	nc	nc	nc
Burro Spring	05-22-00	4.69	0.09	4.77	4.68	0.09	51	3
	12-07-00	4.39	0.05	4.51	4.45	0.06	49	0
	04-08-01	4.34	0.03	4.52	4.49	0.03	49	1
Pipe Creek	05-22-00	4.95	0.03	4.91	4.88	0.03	54	3
	12-07-00	4.88	0.01	4.74	4.73	0.01	55	0
	04-08-01	4.34	0.01	4.50	4.50	0.01	51	1
Pumphouse Spring	05-22-00	4.20	0.04	4.25	4.18	0.04	41	3
	12-07-00	3.87	nc	3.98	3.90	nc	39	1
	04-07-01	4.11	0.08	4.03	3.99	0.07	40	0
Pumphouse Wash Gage	12-07-00	3.80	0.08	3.93	3.89	0.07	38	0
	12-09-00	nc	0.08	nc	nc	0.04	nc	nc
Horn Creek	05-22-00	4.18	0.04	4.13	4.07	0.06	51	2
	12-06-00	3.92	0.01	4.18	4.16	0.02	51	0
	04-07-01	4.11	0.01	4.29	4.27	0.01	61	0
Salt Creek Spring	05-23-00	3.20	0.10	3.25	3.16	0.09	56	1
	12-06-00	3.02	0.04	3.12	3.07	0.05	55	0
	04-10-01	2.97	0.04	3.19	3.14	0.05	58	0
Monument Creek No. 1	05-24-00	4.23	0.04	4.28	4.24	0.03	57	2
Monument Spring	12-05-00	4.00	0.02	3.63	3.61	0.01	43	1
	04-09-01	3.36	0.02	3.70	3.68	0.02	44	1
Iawaii Spring	05-25-00	4.29	0.06	4.35	4.29	0.06	45	1
	12-04-00	4.20	0.02	4.34	4.32	0.02	44	1
	04-11-01	4.04	0.04	4.38	4.35	0.02	46	1
Iermit Spring	12-04-00	3.95	0.08	4.14	4.05	0.09	43	1
1 0	03-11-01	3.85	0.07	4.10	4.04	0.06	43	1
Boucher East Spring	05-26-00	4.06	0.16	4.35	4.14	0.16	36	1
T O	12-04-00	4.18	0.09	4.17	4.07	0.10	37	0
								0

Site	Date	Chloride, dissolved (mg/L as Cl) IC	disso (mg/L	ssium, olved . as K) -AES	Magno disso (mg/L a ICP-	olved as Mg)	diss (mg/L	lium, olved as Na) -AES	Sulfate, dissolved, (mg/L as SO4) - IC	(mg/L a	lissolved as SiO <sub>2</sub> ) -AES
		10	Avg	SD	Avg	SD	Avg	SD		Avg	SD
Red Canyon Spring	08-23-01	nc	nc	nc	nc	nc	nc	nc	nc	nc	nc
	09-26-01	13	2.1	0.0	22	0	6.1	1.2	17	8.8	0.9
JT Spring	04-08-01	11	3.7	0.1	27	0	7.4	0.5	54	6.3	0.1
	05-11-01	15	4.1	0.1	25	1	10	0	45	8.4	0.0
Miners Spring	05-24-00	19	3.6	.1	27	0	12	1	35	9.6	.2
	11-28-00	19	3.2	.2	26	0	11	0	37	9.5	.1
	04-07-01	17	3.7	.0	27	1	10	0	36	8.9	.2
Cottonwood Creek No. 1	05-25-00	16	4.7	.2	63	2	11	0	36	16	1
Cottonwood Creek No. 2	11-29-00	10	2.2	.1	28	0	5.3	.2	18	9.1	.1
	04-09-01	9.4	2.6	.0	29	1	5.2	.0	19	8.1	.3
Grapevine Main Spring	04-10-01	5.6	1.4	.0	23	0	3.8	.0	6.5	7.4	.1
	04-30-01	5.7	1.4	.0	23	1	3.8	.0	6.0	7.4	.1
Grapevine East Spring	05-25-00	26	7.6	.2	55	4	17	2	126	16	0
Shaperine Lust Opring	11-29-00	41	23	0	112	1	52	1	338	10	0
	12-12-00	nc	nc	nc	nc	nc	nc	nc	nc	nc	nc
	04-09-01	35	31	0	110	3	53	1	338	15	0
Lonetree Spring	04-09-01	20	10	0	41	0	25	1	170	8.5	
Lonetree Spring	04-11-01	20 20		0	41	0	23 27	0	187	8.9	.1
Sere Marca Carrier			11								.2
Sam Magee Spring	04-20-01	nc	nc	nc	nc	nc	nc	nc	nc	nc	nc
Burro Spring	05-22-00	19	4.2	.3	37	2	15	1	51	11	0
	12-07-00	20	3.9	.2	36	0	13	0	52	11	0
	04-08-01	18	4.4	.1	36	1	13	0	52	9.6	.1
Pipe Creek	05-22-00	19	3.7	0.2	38	2	12	1	63	10	0
	12-07-00	17	3.6	.2	36	0	10	0	55	10	0
	04-08-01	14	3.5	.0	35	1	9.5	.0	58	9.2	.3
Pumphouse Spring	05-22-00	11	1.8	.2	29	2	6.8	1.0	14	9.4	.4
	12-07-00	12	1.5	.1	29	0	6.1	.2	14	9.6	.2
	04-07-01	10	1.6	.0	28	0	5.8	.1	15	8.5	.2
Pumphouse Wash Gage	12-07-00	12	1.3	.0	28	0	6.1	.3	14	9.4	.0
	12-09-00	nc	nc	nc	nc	nc	nc	nc	nc	nc	nc
Horn Creek	05-22-00	35	6.0	.2	40	2	22	2	99	10	0
	12-06-00	31	6.0	.3	40	1	20	0	93	11	0
	04-07-01	28	6.0	.0	43	0	20	0	110	9.6	.2
Salt Creek Spring	05-23-00	18	4.4	.1	43	1	17	1	169	8.5	.1
	12-06-00	18	4.1	.2	43	1	15	0	177	8.9	.2
	04-10-01	16	5.0	.1	46	0	16	0	178	8.2	.3
Monument Creek No. 1	05-24-00	87	5.8	.0	42	2	45	1	75	9.5	.2
Monument Spring	12-05-00	43	1.6	.2	30	0	18	0	20	9.6	.1
-	04-09-01	40	1.8	.0	31	2	18	1	19	8.9	.2
Hawaii Spring	05-25-00	13	2.7	.1	31	0	13	0	33	10	0
	12-04-00	14	2.5	.1	31	1	11	0	34	10	0
	04-11-01	13	2.8	.1	32	1	11	0	38	9.7	.0
Hermit Spring	12-04-00	11	1.3	.0	27	1	5.6	.1	12	9.6	.2
r	04-11-01	9.4	1.5	.0	27	1	5.5	.1	13	9.4	.1
Boucher East Spring	05-26-00	10	2.0	.0	38	1	7.0	.3	36	11	0
Soucher Last Spring	12-04-00	10	2.0 1.9	.0	35	0	6.2	.3	32	11	0
	04-12-01	12	2.0	.0	35	1	6.1	.5	32	9.5	.0

				Hut		icentratio							
Site	Date	Nitrate, d (mg/L colori	as N)	Nitrite, d (mg/L colori	as N)	Ammo disso (mg/L colori	olved as N)	diss (µg/L	horus, olved . as P) -AES		olved as P)	Organi (DOC)	solved ic Carbor (mg C/L) IR
		Median	MAD	Median	MAD	Median	MAD	Avg	SD	Median	MAD	Avg	SD
Red Canyon Spring	08-23-01	nc	nc	nc	nc	nc	nc	nc	nc	nc	nc	nc	nc
	09-26-01	.76	.05	.001	.001	.01	.00	10	4	<.02	.03	.6	.0
JT Spring	04-08-01	.27	.06	<.001	na	<.04	.01	<4	3	<.02	na	4.7	.1
	05-11-01	.80	.01	<.002	.000	<.02	.02	4	1	<.01	na	.6	.1
Miners Spring	05-24-00	1.2	.2	.002	.001	<.01	.00	<8	0	<.01	.00	1.1	.1
	11-28-00	1.4	.0	<.001	.000	<.03	.00	<2	1	<.01	.01	1.8	.0
	04-07-01	1.4	.2	<.001	na	<.04	.03	<4	3	<.02	.00	.8	.2
Cottonwood Creek No. 1	05-25-00	<.01	.01	<.001	.001	<.01	.01	<8	5	<.01	.00	1.2	.2
Cottonwood Creek No. 2	11-29-00	.50	.01	<.001	.000	<.03	.00	2	1	<.01	.01	1.0	.1
	04-09-01	.54	.03	<.001	na	<.04	.02	<4	1	<.02	.01	.9	.2
Grapevine Main Spring	04-10-01	.30	.04	<.001	na	<.04	.01	15	2	.03	.01	1.0	.2
	04-30-01	.24	.01	<.002	.001	<.02	.02	13	1	.01	na	.7	.2
Grapevine East Spring	05-25-00	<.01	.00	<.001	.001	.01	.01	15	3	<.01	.00	2.8	.0
	11-29-00	.01	.01	<.001	.000	<.03	.01	<2	1	<.01	.01	4.0	.1
	12-12-00	nc	nc	nc	nc	nc	nc	nc	nc	nc	nc	nc	nc
	04-09-01	<.01	.01	<.001	na	<.04	.01	<4	2	<.02	.01	2.1	.2
Lonetree Spring	04-11-01	.30	.06	<.001	na	<.04	.01	<4	1	<.02	.01	1.0	.2
1 0	05-01-01	.21	.01	<.002	.001	<.02	.01	<4	3	<.01	na	1.2	.2
Sam Magee Spring	04-20-01	nc	nc	nc	nc	nc	nc	nc	nc	nc	nc	nc	nc
Burro Spring	05-22-00	.37	.00	<.001	.001	<.01	.00	<8	4	<.01	.01	1.4	.0
	12-07-00	.45	.01	<.001	.000	<.03	.02	<2	1	<.01	.01	1.4	.1
	04-08-01	.51	.02	<.001	na	<.04	.01	<4	1	<.02	.01	1.8	.2
Pipe Creek	05-22-00	.17	.03	<.001	.001	<.01	.01	<8	8	<.01	.00	1.2	.0
	12-07-00	.17	.01	<.001	.000	<.03	.01	<2	1	<.01	.01	1.0	.0
	04-08-01	.25	.03	<.001	na	<.04	.01	<4	2	<.02	.01	2.6	.2
Pumphouse Spring	05-22-00	.23	.03	.002	.000	<.01	.01	<8	2	<.01	.00	1.6	.2
amphouse spring	12-07-00	.54	.00	<.001	.000	<.03	.01	<2	1	<.01	.00	na	na
	04-07-01	.60	.14	<.001	na	<.04	.01	<4	5	<.02	.00	.9	.2
Pumphouse Wash Gage	12-07-00	.57	.01	<.001	.000	<.03	.01	3	0	<.01	.00	.7	.0
amphouse wash Gage	12-09-00	nc	nc	nc	nc	nc	nc	nc	nc	nc	nc	nc	nc
Horn Creek	05-22-00	1.2	.1	<.001	.001	<.01	.00	<8	5	<.01	.00	.9	.1
lioni eleck	12-06-00	1.2	.0	<.001	.000	<.03	.00	<2	1	<.01	.00	1.3	.0
	04-07-01	1.2	.1	<.001	na	<.04	.01	<4	3	<.02	.00	1.0	.2
Salt Creek Spring	05-23-00	.54	.11	.003	.002	.02	.00	<8	3	<.01	.00	.9	.0
San Creek Spring	12-06-00	.54	.01	<.001	.002	<.02	.00	<2	1	<.01	.01	.7	.0
	04-10-01	.38	.16	<.001	na	<.03 <.04	.00	<4	2	<.02	.00	.8	.0
Monument Creek No. 1	05-24-00	4.5	.10	.001	.001	.04	.00	<8	2	<.02 <.01	.00	.o 1.0	.0
Monument Spring	12-05-00	5.7	.0	<.001	.000	<.1	.01	12	1	<.02	.00	.8	.0
nonument opting	04-09-01	6.8	.0	<.001	na .000	<.04	na .0	12	2	<.02 <.02	na .00	.0 .9	.0
Hawaii Spring	05-25-00	0.8 .60	.5 .07	<.001 .001	na .001	<.04 <.01	na .01	<8	2	<.02 <.01	na .00	.9 .7	.2 .0
iawan opinig	12-04-00	.60	.07	<.001	.001	<.01	.01	< 8 <2	2 2	<.01 <.01	.00	.7 1.8	
						<.03 <.04	.02	<2 4	2	<.01 <.02	.01	1.8 .9	.1
Harmit Spring	04-11-01	.57 .69	.06	<.001	na .000	< 0.04	.02		2	<.02 <.01		.9 1.5	.2
Hermit Spring	12-04-00		.01	<.001				2			.01		.1
Doughor East Series -	04-11-01	.61	.05	<.001	na	<.04	.02	<4	1	<.02	.02	.9	.2
Boucher East Spring	05-26-00 12-04-00	<.01 .15	.00 .01	<.001 .001	.000 .000	<.01 <.03	.00 .01	<8 <2	1 1	<.01 <.01	.00 .01	1.4 1.1	.0 .1

Site	Date		, dissolved Al) ICP-MS	-	dissolved b) ICP-MS		dissolved (s) ICP-MS		dissolved Ba) ICP-MS
		Avg	SD	Avg	SD	Avg	SD	Avg	SD
ed Canyon Spring	08-23-01	nc	nc	nc	nc	nc	nc	nc	nc
	09-26-01	<.1	.1	.013	.002	17	.06	343	5
Г Spring	04-08-01	1.5	.1	.047	.000	4.9	.04	117	2
	05-11-01	.2	.1	.020	.003	14	.02	118	1
liners Spring	05-24-00	<.1	.1	.016	.002	17	.20	127	1
	11-28-00	<.3	1.0	.023	.001	19	.45	123	2
	04-07-01	.2	.0	.022	.000	19	.09	127	0
ottonwood Creek No. 1	05-25-00	<.1	.2	.022	.002	1.2	.01	248	3
ottonwood Creek No. 2	11-29-00	<.3	1.9	.012	.000	4.7	.00	238	2
	04-09-01	.2	.1	.043	.006	4.8	.00	233	3
rapevine Main Spring	04-10-01	1.8	.0	.015	.002	5.3	.03	354	11
T T T T T T T T T T T T T T T T T T T	04-30-01	.2	.0	.013	.002	5.4	.03	359	1
apevine East Spring	05-25-00	3.4	.1	.029	.000	1.5	.02	57	0
aperine Last oping	11-29-00	<.3	.9	.029	.001	.63	.04	22	1
	12-12-00	nc	nc	nc	nc	nc	nc	nc	nc
	04-09-01	.4	.2	.008	.001	.27	.04	18	0
onetree Spring	04-09-01	.1	.2	.008	.001	.98	.04	37	1
neuce spring	05-01-01	.1	.0	.003	.001	.98	.03	40	2
m Magee Spring	04-20-01	nc	nc	nc	nc	nc	nc	nc	nc
rro Spring	05-22-00	.3	.1	.012	.001	.97	.02	103	1
ino spring		.3 <.3		.012	.001	.97	.02	90	
	12-07-00		.4						0
Caral	04-08-01	.2	.0	.012	.001	.99	.06	92 70	1
pe Creek	05-22-00	.4	.1	.024	.004	.86	.00	70	2
	12-07-00	<.3	1.6	.013	.002	.51	.01	65	1
	04-08-01	.6	.2	.011	.001	.77	.03	64	1
mphouse Spring	05-22-00	.2	.0	.012	.001	1.5	.03	282	1
	12-07-00	<.3	.9	.011	.001	1.9	.04	273	1
	04-07-01	.4	.0	.011	.001	1.7	.03	276	1
mphouse Wash Gage	12-07-00	<.3	.1	.012	.001	2.0	.02	280	2
	12-09-00	nc	nc	nc	nc	nc	nc	nc	nc
rn Creek	05-22-00	<.1	.5	.013	.001	2.7	.08	48	0
	12-06-00	<.3	.8	.015	.002	2.8	.01	50	1
	04-07-01	.3	.1	.018	.003	2.5	.04	45	1
t Creek Spring	05-23-00	.2	.1	.015	.002	2.2	.02	29	1
	12-06-00	<.3	.4	.010	.002	2.3	.02	28	0
	04-10-01	.1	.0	.011	.002	2.2	.01	29	1
onument Creek No. 1	05-24-00	.3	.3	.018	.009	.52	.02	64	0
onument Spring	12-05-00	<.3	1.2	.014	.002	4.0	.01	275	2
	04-09-01	.3	.0	.013	.002	4.0	.05	273	1
waii Spring	05-25-00	<.1	.1	.009	.001	.72	.02	107	1
	12-04-00	<.3	1.6	.008	.001	.75	.02	108	4
	04-11-01	5.2	.7	.015	.002	.83	.01	107	1
rmit Spring	12-04-00	<.3	1.2	.008	.001	1.6	.04	228	2
	04-11-01	.8	.1	.010	.000	1.6	.03	212	0
ucher East Spring	05-26-00	<.1	.2	.008	.003	.35	.01	81	1
	12-04-00	<.3	1.0	.010	.000	.53	.01	109	5
	04-12-01	.2	.0	.012	.004	.54	.03	104	2

			Trace elemen			icu			
Site	Date	Beryllium, (µg/L as Be		Bismuth, c (µg/L as B		-	lissolved B) ICP-MS	Bromide, (µg/L as E	dissolved Br) ICP-MS
	_	Avg	SD	Avg	SD	Avg	SD	Avg	SD
Red Canyon Spring	08-23-01	nc	nc	nc	nc	nc	nc	nc	nc
	09-26-01	<.006	.002	<.001	.000	71	1	33	2
T Spring	04-08-01	<.008	.001	.003	.000	74	8	17	1
	05-11-01	<.008	.010	<.001	.000	103	14	31	1
Ainers Spring	05-24-00	<.005	.001	<.002	.000	96	5	50	0
	11-28-00	<.005	.004	<.003	.000	88	4	58	4
	04-07-01	<.008	.004	.008	.000	97	11	54	1
Cottonwood Creek No. 1	05-25-00	<.005	.001	<.002	.001	123	9	33	1
Cottonwood Creek No. 2	11-29-00	.007	.002	<.003	.001	46	0	24	1
	04-09-01	<.008	.001	<.001	.001	49	8	23	1
Frapevine Main Spring	04-10-01	<.008	.002	<.001	.000	28	9	16	1
	04-30-01	<.008	.003	<.001	.001	25	4	16	2
Brapevine East Spring	05-25-00	<.005	.002	.003	.003	210	1	79	1
1 Trino	11-29-00	<.005	.003	<.003	.000	538	14	120	6
	12-12-00	nc	nc	nc	nc	nc	nc	nc	nc
	04-09-01	<.008	.005	<.001	.001	724	70	105	0
onetree Spring	04-11-01	<.008	.004	<.001	.000	119	10	42	1
oneuce opining	05-01-01	<.008	.006	<.001	.000	125	10	47	1
am Magee Spring	04-20-01	nc	nc	nc	nc	nc	nc	nc	nc
urro Spring	05-22-00	<.005	.003	<.002	.001	92	1	47	3
uno spring	12-07-00	.005	.005	<.002	.000	76	4	47	1
	04-08-01	<.005	.010	<.005	.000	93	13	48	2
ipe Creek	05-22-00	<.005	.002	<.001	.000	93 72	13	48	2
ipe creek	12-07-00	<.005	.002	<.002	.000	60	1	40	3
	04-08-01	<.003	.002	<.003 .001	.000	64	11	40 38	1
umphouse Spring	05-22-00	<.008	.004	<.001	.001	04 39	6	38	1
umphouse spring	12-07-00	<.005	.001	<.002	.000	39 32	1	32	1
	12-07-00	<.003	.001	<.003	.000	32 32		33 34	
with the West Cores							8		2
umphouse Wash Gage	12-07-00	<.005	.007	<.003	.001	32	1	37	1
	12-09-00	nc	nc	nc	nc	nc	nc	nc	nc
orn Creek	05-22-00	<.005	.002	<.002	.002	114	4	57	2
	12-06-00	.008	.002	<.003	.000	111	4	61	0
	04-07-01	<.008	.002	<.001	.000	111	8	57	2
alt Creek Spring	05-23-00	<.005	.002	<.002	.001	92	2	30	1
	12-06-00	.008	.002	<.003	.000	89	2	32	1
	04-10-01	<.008	.002	<.001	.000	96	9	30	1
Ionument Creek No. 1	05-24-00	<.005	.003	<.002	.002	124	1	79	2
Ionument Spring	12-05-00	.005	.006	<.003	.000	39	1	43	2
	04-09-01	<.008	.005	<.001	.000	39	12	41	1
awaii Spring	05-25-00	<.005	.001	<.002	.002	57	1	28	0
	12-04-00	<.005	.005	<.003	.000	54	2	33	2
	04-11-01	<.008	.001	<.001	.000	58	15	33	2
ermit Spring	12-04-00	<.005	.005	<.003	.000	31	2	30	1
	04-11-01	<.008	.001	<.001	.001	30	10	29	1
oucher East Spring	05-26-00	.006	.000	<.002	.001	51	1	26	0
	12-04-00	<.005	.006	<.003	.001	36	2	28	2
	04-12-01	<.008	.004	<.001	.000	35	10	27	1

		Cadmium,	dissolved	Cerium, d	lissolved	Cesium, o	lissolved	Chromium,	dissolved
Site	Date	(μ <b>g/L as C</b> o	I) ICP-MS	(μ <b>g/L as C</b> e	e) ICP-MS	(μ <b>g/L as C</b>	s) ICP-MS	(µ <b>g/L as C</b>	r) ICP-MS
		Avg	SD	Avg	SD	Avg	SD	Avg	SD
ed Canyon Spring	08-23-01	nc	nc	nc	nc	nc	nc	nc	nc
	09-26-01	<.002	.002	.0003	.0001	.89	.00	.9	.1
T Spring	04-08-01	.010	.001	.013	.000	.05	.01	<.2	.2
	05-11-01	<.001	.000	.0010	.0001	1.6	.0	.5	.1
Iners Spring	05-24-00	.002	.001	.0014	.0005	.82	.01	.2	.1
	11-28-00	<.003	.001	.0008	.0001	.95	.02	.3	.0
	04-07-01	<.001	.001	.0008	.0002	1.0	.1	.4	.2
ottonwood Creek No. 1	05-25-00	<.002	.002	.0017	.0004	<.01	.00	<.1	.4
ottonwood Creek No. 2	11-29-00	.008	.003	.0013	.0001	.09	.01	<.1	.0
	04-09-01	.013	.009	.0016	.0001	.10	.01	.4	.1
rapevine Main Spring	04-10-01	.004	.002	.0013	.0003	.50	.01	.8	.0
	04-30-01	<.001	.001	.0013	.0005	.46	.02	.8	.2
rapevine East Spring	05-25-00	.007	.004	.021	.001	.02	.00	<.1	.4
	11-29-00	<.003	.001	.0039	.0005	<.01	.01	<.1	.0
	12-12-00	nc	nc	nc	nc	nc	nc	nc	nc
	04-09-01	.003	.001	.0027	.0000	.26	.01	<.2	.1
onetree Spring	04-11-01	<.001	.001	.0016	.0002	.68	.05	<.2	.1
1 0	05-01-01	.002	.003	.0031	.0004	.79	.01	.2	.2
am Magee Spring	04-20-01	nc	nc	nc	nc	nc	nc	nc	nc
urro Spring	05-22-00	.005	.002	.0021	.0001	.08	.00	<.1	.(
	12-07-00	<.003	.001	.0026	.0004	.07	.00	<.1	
	04-08-01	.002	.000	.0014	.0001	.10	.00	<.2	.2
pe Creek	05-22-00	.020	.002	.0031	.0003	.03	.00	<.1	.(
F	12-07-00	.005	.002	.0033	.0001	<.01	.00	<.1	.2
	04-08-01	.004	.002	.0018	.0004	.04	.00	.3	.1
umphouse Spring	05-22-00	.004	.001	.0048	.0004	.15	.00	<.1	.1
amphouse spring	12-07-00	<.003	.001	.0029	.0001	.19	.01	<.1	.1
	12-07-01	.005	.001	.0031	.0005	.22	.01	.4	.1
umphouse Wash Gage	12-07-00	<.003	.001	.0007	.0002	.37	.01	.3	
imphouse wash Gage	12-09-00	nc	nc	nc	nc	nc	nc	nc	nc
orn Creek	05-22-00	.014	.001	.0028	.0006	.36	.00	<.1	.1
ioni creek	12-06-00	.007	.001	.0012	.0001	.30	.00	<.1	.1
	04-07-01	<.001	.002	.0012	.0002	.12	.01	.2	.1
alt Creek Spring	05-23-00	.008	.001	.0012	.0002	.54	.00	<.1	.1
an creek opring	12-06-00	<.003	.004	.0011	.0005	.54	.00	<.1	
	04-10-01	<.005	.000	.0018	.0001	.72	.01	<.2	.1
Ionument Creek No. 1	05-24-00	.012	.000	.0004	.0001	.06	.00	<.1	
Ionument Spring	12-05-00	<.003	.003	.0007	.0002	.72	.00	.2	.1
onument opring	04-09-01	<.003 <.001	.002	.0007	.0003	.72	.00	.2	.1
awaii Spring	04-09-01	.001	.001	<.0002	.0001	.83	.02	.0	.1
awan opinig	03-23-00 12-04-00		.001		.0001				
	12-04-00 04-11-01	<.003		<.0003		.15	.01	<.1	).
amait Canina		.003	.002	.012	.002	.20	.02	.5	).
ermit Spring	12-04-00	<.003	.001	.0003	.0003	.49	.01	.1	).
han East C	04-11-01	.001	.001	.0040	.0005	.23	.00	.5	.2
oucher East Spring	05-26-00	<.002	.003	<.0004	.0002	.05	.02	<.1	.2
	12-04-00	<.003	.002	.0005	.0001	.06	.01	<.1	

				nt concentrat					
Site	Date	Cobalt, d (µg/L as Co			lissolved u) ICP-MS	Dysprosium (µg/L as D		Erbium, d (µg/L as E	
	_	Avg	SD	Avg	SD	Avg	SD	Avg	SD
Red Canyon Spring	08-23-01	nc	nc	nc	nc	nc	nc	nc	nc
	09-26-01	<.003	.003	<.02	.01	<.0006	.0005	<.0004	.000
T Spring	04-08-01	<.008	.007	.81	.01	.0051	.0004	.0030	.000
	05-11-01	<.008	.009	.02	.02	.0004	.0005	.0006	.000
Ainers Spring	05-24-00	<.008	.003	<.06	.10	.0004	.0002	<.0003	.000
	11-28-00	<.008	.004	.04	.03	.0004	.0001	.0007	.000
	04-07-01	<.008	.013	.03	.00	.0006	.0001	.0004	.000
Cottonwood Creek No. 1	05-25-00	.72	.01	<.06	.07	.0010	.0004	.0008	.000
Cottonwood Creek No. 2	11-29-00	<.008	.004	<.04	.01	.0010	.0001	.0007	.000
	04-09-01	<.008	.006	.03	.02	.0010	.0005	.0010	.000
Brapevine Main Spring	04-10-01	<.008	.009	.11	.02	<.0003	.0002	<.0002	.000
1 1 0	04-30-01	<.008	.016	.02	.01	<.0003	.0002	<.0002	.000
rapevine East Spring	05-25-00	.12	.03	<.06	.02	.0029	.0004	.0017	.000
. ro	11-29-00	<.008	.006	.30	.06	.0011	.0003	.0008	.000
	12-12-00	nc	nc	nc	nc	nc	nc	nc	nc
	04-09-01	<.008	.008	<.01	.01	.0011	.0002	.0009	.000
onetree Spring	04-11-01	<.008	.023	<.01	.00	.0072	.0003	.0044	.000
oneuce opining	05-01-01	<.008	.028	<.01	.00	.0072	.0002	.0038	.000
am Magee Spring	04-20-01	nc	nc	nc	nc	nc	nc	nc	nc
urro Spring	05-22-00	<.008	.024	<.06	.20	.0015	.0001	.0010	.000
uno spring	12-07-00	<.008	.024	<.04	.20	.0013	.0001	.0006	.000
	04-08-01	<.008	.018	<.01	.01	.0012	.0003	.0000	.000
ipe Creek	04-08-01	<.008	.018	<.06	.01	.0010	.0002	.0013	.000
ipe Cieek	12-07-00				.02	.0010	.0003		.000
	04-08-01	<.008	.005 .021	<.04 .03	.02	.0008	.0002	.0013	.000
contraction of the second s		<.008	.021		.10	.0011	.0001	.0009	.000
umphouse Spring	05-22-00	<.008		<.06				.0007	
	12-07-00	<.008	.006	<.04	.01	.0012	.0003	.0009	.000
	12-07-01	<.008	.015	<.01	.01	.0010	.0003	.0009	.000
umphouse Wash Gage	12-07-00	<.008	.010	<.04	.01	.0005	.0002	<.0002	.000
	12-09-00	nc	nc	nc	nc	nc	nc	nc	nc
lorn Creek	05-22-00	<.008	.017	<.06	.17	.0019	.0003	.0015	.000
	12-06-00	<.008	.008	<.04	.00	.0008	.0001	.0010	.000
	04-07-01	<.008	.025	.03	.02	.0014	.0001	.0013	.000
alt Creek Spring	05-23-00	<.008	.007	<.06	.47	<.0002	.0001	<.0003	.000
	12-06-00	<.008	.011	<.04	.02	.0005	.0002	<.0002	.000
	04-10-01	<.008	.018	<.01	.02	<.0003	.0002	.0004	.000
Ionument Creek No. 1	05-24-00	<.008	.022	<.06	.01	.0029	.0006	.0049	.000
Ionument Spring	12-05-00	<.008	.007	.55	.01	<.0003	.0002	<.0002	.000
	04-09-01	<.008	.014	.42	.01	<.0003	.0002	.0004	.000
awaii Spring	05-25-00	<.008	.006	<.06	.08	<.0002	.0000	.0004	.000
	12-04-00	<.008	.011	<.04	.01	<.0003	.0002	<.0002	.000
	04-11-01	<.008	.018	<.01	.01	.0006	.0000	.0008	.000
ermit Spring	12-04-00	<.008	.007	<.04	.02	<.0003	.0003	<.0002	.000
	04-11-01	<.008	.013	.05	.05	.0004	.0002	.0005	.000
oucher East Spring	05-26-00	<.008	.008	<.06	.43	<.0002	.0001	<.0003	.000
	12-04-00	<.008	.010	.06	.01	<.0003	.0001	<.0002	.000
	04-12-01	<.008	.016	.05	.01	.0005	.0004	.0004	.000

Site	Date	Europium, (µg/L as E		Gadolinium (µg/L as G		Holmium, (µg/L as H		lron, di (µg/L as F	ssolved e) ICP-MS
	-	Avg	SD	Avg	SD	Avg	SD	Avg	SD
Red Canyon Spring	08-23-01	nc	nc	nc	nc	nc	nc	nc	nc
	09-26-01	.0009	.0005	<.0004	.0001	<.0001	.0001	1.4	.8
T Spring	04-08-01	.0010	.0018	.0059	.0003	.0009	.0000	2.0	.0
	05-11-01	.0004	.0011	<.0004	.0002	<.0001	.0001	<.2	.2
Ainers Spring	05-24-00	.021	.002	.0007	.0002	.0001	.0001	.7	.0
	11-28-00	.0028	.0009	.0006	.0003	.0001	.0000	<.5	.2
	04-07-01	.0038	.0026	.0006	.0003	<.0001	.0000	<.2	.0
ottonwood Creek No. 1	05-25-00	.040	.001	.0036	.0007	.0002	.0001	347	8
Cottonwood Creek No. 2	11-29-00	.0068	.0015	.0011	.0004	.0002	.0000	.9	.2
	04-09-01	.0045	.0005	.0013	.0002	.0002	.0000	.4	.3
rapevine Main Spring	04-10-01	.013	.000	<.0004	.0000	<.0001	.0000	<.2	.1
	04-30-01	.0095	.0121	<.0004	.0002	<.0001	.0000	<.2	.2
rapevine East Spring	05-25-00	.0077	.0018	.0049	.0008	.0005	.0000	36	0
	11-29-00	.0007	.0002	.0007	.0002	.0005	.0001	.5	.1
	12-12-00	nc	nc	nc	nc	nc	nc	nc	nc
	04-09-01	.0004	.0003	.0008	.0003	.0004	.0001	.7	.2
onetree Spring	04-09-01	.0023	.0003	.0071	.0003	.0016	.0001	<.2	.0
onetree spring	05-01-01	.0023	.0002	.0071	.0002	.0016	.0002	<.2	.0
am Magee Spring	04-20-01								
• • •		nc .018	nc .002	nc .0021	nc .0002	nc .0004	nc .0000	nc 7.4	nc .6
urro Spring	05-22-00				.0002		.0000	7.4 7.9	
	12-07-00	.0016	.0005	.0012		.0003		15	.8
	04-08-01	.0017	.0024	.0013	.0001	.0002	.0001		0
ipe Creek	05-22-00	.011	.002	.0022	.0006	.0004	.0001	3.5	1.1
	12-07-00	.0010	.0006	.0008	.0001	.0003	.0001	<.5	.1
	04-08-01	.0014	.0012	.0017	.0002	.0002	.0000	.4	.0
umphouse Spring	05-22-00	.047	.003	.0014	.0004	.0002	.0001	.8	.1
	12-07-00	.0037	.0004	.0014	.0002	.0002	.0001	<.5	.5
	12-07-01	.0082	.0047	.0009	.0003	<.0001	.0000	1.3	.9
umphouse Wash Gage	12-07-00	.0044	.0005	.0003	.0002	<.0001	.0001	<.5	.1
	12-09-00	nc	nc	nc	nc	nc	nc	nc	nc
orn Creek	05-22-00	.0076	.0011	.0024	.0002	.0004	.0002	.4	.2
	12-06-00	.0011	.0006	.0010	.0001	.0003	.0000	<.5	.0
	04-07-01	.0005	.0007	.0014	.0001	.0003	.0000	.3	.2
alt Creek Spring	05-23-00	.0045	.0005	.0005	.0003	<.0001	.0000	.3	.2
	12-06-00	.0009	.0003	.0003	.0001	.0001	.0000	<.5	.3
	04-10-01	.0004	.0004	<.0004	.0002	<.0001	.0000	<.2	.2
Ionument Creek No. 1	05-24-00	.0092	.0027	.0017	.0002	.0012	.0001	.3	.1
Ionument Spring	12-05-00	.0064	.0026	.0002	.0004	<.0001	.0000	<.5	.2
	04-09-01	.0049	.0016	<.0004	.0000	<.0001	.0000	<.2	.1
awaii Spring	05-25-00	.019	.000	.0011	.0001	<.0001	.0001	.5	.3
	12-04-00	.0020	.0004	.0004	.0002	.0001	.0001	<.5	.5
	04-11-01	.0010	.0017	.0012	.0004	.0002	.0000	2.8	.1
ermit Spring	12-04-00	.0051	.0012	.0002	.0001	.0001	.0001	1.4	1.7
	04-11-01	.0043	.0031	.0006	.0003	<.0001	.0000	<.2	.1
oucher East Spring	05-26-00	.012	.005	.0006	.0003	<.0001	.0000	1.2	.3
	12-04-00	.0027	.0008	.0003	.0002	.0001	.0000	<.5	.2
	04-12-01	<.0002	.0015	.0004	.0003	<.0001	.0000	.3	.1

			Trace elemen	t concentrati	ons—Contini	ued			
Site	Date	Lanthanum, (µg/L as La		Lead, dis (µg/L as Pl		-	dissolved Li) ICP-MS	Lutetium, (µg/L as L	
	-	Avg	SD	Avg	SD	Avg	SD	Avg	SD
Red Canyon Spring	08-23-01	nc	nc	nc	nc	nc	nc	nc	nc
	09-26-01	.0026	.0004	.012	.010	5.6	.4	<.0002	.000
T Spring	04-08-01	.0093	.0005	.022	.001	7.5	.3	.0004	.000
	05-11-01	.0013	.0004	.017	.002	10	0	<.0001	.000
Ainers Spring	05-24-00	.0039	.0002	.003	.001	9.3	.2	<.0001	.000
	11-28-00	.0017	.0001	.069	.004	8.7	.2	.0001	.000
	04-07-01	.0011	.0002	.014	.006	8.5	.3	<.0001	.000
Cottonwood Creek No. 1	05-25-00	.0075	.0004	.006	.001	17	1	<.0001	.000
Cottonwood Creek No. 2	11-29-00	.0036	.0000	.011	.006	5.0	.2	<.0001	.000
	04-09-01	.0037	.0005	.048	.045	4.9	.1	.0001	.000
Brapevine Main Spring	04-10-01	.0024	.0003	.029	.005	2.6	.2	<.0001	.000
1 10	04-30-01	.0022	.0002	.014	.001	2.6	.1	<.0001	.000
Grapevine East Spring	05-25-00	.014	.000	.043	.005	38	0	.0001	.000
1 10	11-29-00	.0009	.0003	.006	.003	127	3	.0002	.000
	12-12-00	nc	nc	nc	nc	nc	nc	nc	nc
	04-09-01	.0010	.0001	.032	.013	140	1	<.0001	.000
Lonetree Spring	04-11-01	.022	.000	.007	.002	37	0	.0004	.000
	05-01-01	.024	.000	.019	.000	38	1	.0006	.000
am Magee Spring	04-20-01	nc	nc	nc	nc	nc	nc	nc	nc
Burro Spring	05-22-00	.0039	.0002	.008	.003	17	1	.0001	.000
	12-07-00	.0021	.0001	.005	.003	16	0	<.0001	.000
	04-08-01	.0017	.0002	.011	.007	16	0	.0001	.000
Pipe Creek	05-22-00	.0047	.0002	.039	.002	14	1	.0001	.000
I	12-07-00	.0021	.0000	.012	.001	13	0	<.0001	.000
	04-08-01	.0036	.0002	.009	.001	13	0	.0001	.000
Pumphouse Spring	05-22-00	.0091	.0014	.012	.001	6.6	.1	.0001	.000
umphouse spring	12-07-00	.0037	.0002	.005	.001	6.1	.2	<.0001	.000
	12-07-01	.0030	.0002	.014	.005	6.2	.1	.0001	.000
Pumphouse Wash Gage	12-07-00	.0019	.0000	.011	.001	6.0	.1	<.0001	.000
umphouse wush ouge	12-09-00	nc	nc	nc	nc	nc	nc	nc	nc
Horn Creek	05-22-00	.0061	.0000	.021	.003	ne 21	0	.0002	.0000
Iom creek	12-06-00	.0019	.0002	.021	.004	21	1	.0002	.000
	04-07-01	.0019	.0002	.022	.004	21	0	.0001	.000
alt Creek Spring	05-23-00	.0015	.0003	.032	.002	24	0	<.0001	.000
an creek spring	12-06-00	.00013	.0001	.014	.003	24	0	<.0001	.000
	04-10-01	.0002	.0001	.011	.001	24	0	<.0001	.000
Monument Creek No. 1	05-24-00	.0026	.0003	.034	.000	83	1	.0011	.000
Aonument Spring	12-05-00	.0024	.0001	.26	.00	5.5	.1	<.0001	.000
ionument opting	04-09-01	.0024	.0001	.20	.00	5.5	.1	<.0001	.000
Iawaii Spring	05-25-00	.0012	.0001	.24	.01	15	.1	<.0001	.000
awan Spring	12-04-00	.0031	.0001	.067	.008	13	0	<.0001	.000
	12-04-00 04-11-01	.0011	.0001	.003	.001	13	0	<.0001	.000
Iermit Spring	12-04-00	.0062	.0003	.071	.003			.0001	.000
ternit spring						5.6 5.7	.1		
Poughar Fast Smills -	04-11-01	.0020	.0000	.028	.013	5.7	.3	<.0001	.000
Boucher East Spring	05-26-00	.0028	.0008	.004	.003	9.0 7.2	.1	<.0001	.000
	12-04-00	.0009	.0002	<.004	.002	7.2	.3	<.0001	.000

		Manganese,	, dissolved	Mercury,	dissolved	Molybdenum	, dissolved	Neodymiu	m, dissolvec
Site	Date	(μ <b>g/L as M</b> i		(ng/L as H		(μ <b>g/L as M</b> o		-	Nd) ICP-MS
		Avg	SD	Avg	SD	Avg	SD	Avg	SD
Red Canyon Spring	08-23-01	nc	nc	nc	nc	nc	nc	nc	nc
	09-26-01	.12	.04	1.0	.3	1.8	.0	.0019	.0008
T Spring	04-08-01	.20	.01	1.8	.2	1.2	.0	.016	.000
	05-11-01	.06	.01	<.3	.2	1.9	.0	.0021	.0001
Ainers Spring	05-24-00	.05	.01	.8	.2	1.9	.0	.0029	.0003
	11-28-00	.04	.01	.9	.0	1.9	.1	.0007	.0001
	04-07-01	.05	.01	.6	.2	1.9	.0	.0024	.0006
Cottonwood Creek No. 1	05-25-00	92	1	<.5	.1	.80	.00	.0066	.0005
Cottonwood Creek No. 2	11-29-00	.21	.01	.7	.2	.72	.04	.0040	.0004
	04-09-01	.10	.01	<.4	.1	.70	.00	.0040	.0006
Grapevine Main Spring	04-10-01	<.04	.01	.4	.3	.36	.03	.0032	.0009
	04-30-01	.06	.01	.7	.0	.36	.02	.0028	.0005
Grapevine East Spring	05-25-00	25	1	<.5	.2	2.8	.0	.017	.001
1 1 0	11-29-00	.13	.02	.9	.2	2.0	.1	.0016	.0007
	12-12-00	nc	nc	nc	nc	nc	nc	nc	nc
	04-09-01	2.0	.0	.6	.2	1.8	.0	.0016	.0004
Lonetree Spring	04-11-01	.05	.01	<.4	.2	2.5	.0	.028	.000
	05-01-01	.23	.02	<.4	.3	2.4	.0	.031	.003
Sam Magee Spring	04-20-01	nc	nc	nc	nc	nc	nc	nc	nc
Burro Spring	05-22-00	3.6	.1	.6	.2	1.5	.1	.0041	.0003
sure spring	12-07-00	1.5	.0	.7	.1	1.5	.1	.0038	.0007
	04-08-01	2.6	.0	.5	.2	1.5	.0	.0031	.0005
lina Craak	05-22-00	1.8	.0	<.5	.1	1.2	.0	.0056	.0003
Pipe Creek	12-07-00	.18	.05	1.1	.1	1.0	.0	.0036	.0001
	04-08-01	.43	.03	<.4	.3	1.0	.0	.0047	.0005
Pumphouse Spring	05-22-00	1.8	.01	<.5	.1	.59	.01	.0047	.0005
unphouse opining	12-07-00	.07	.00	.6	.1	.57	.01	.0049	.0001
	12-07-01	.15	.00	<.4	.3	.60	.01	.0049	.0007
Pumphouse Wash Gage	12-07-01	<.03	.02	nc	nc	.60	.02	.0049	.0007
unphouse wash Gage	12-07-00	nc	nc	nc	nc	nc	nc	nc	nc
Horn Creek	05-22-00	.06	.01	<.5	.2	3.3	.1	.0059	.0006
IoIII CICEK	12-06-00	<.03	.01	1.0	.2	3.4	.1	.0039	.0000
	04-07-01	<.04	.01	<.4	.3	6.1	.0	.0021	.0010
alt Creek Spring	05-23-00	.04	.01	<.4 <.5	.3	2.6	.0	.0025	.0000
San Creek Spring	12-06-00	.08	.01	<.5 .7	.0	2.0	.0 .0	.0013	.0004
	04-10-01	<.04		<.4		2.4		.0011	.0007
Monument Creek No. 1	05-24-00	<.04 .03	.00		.2		.0	.0007	.0001
			.01	.8	.1	1.5	.0		
Monument Spring	12-05-00	<.03	.02	2.1	.2	.90	.06	.0026	.0011
Iomoli Carin -	04-09-01	<.04	.01	1.4	.3	.90	.01	.0016	.0004
Iawaii Spring	05-25-00	.03	.03	<.5	.1	1.3	.0	.0026	.0003
	12-04-00	<.03	.01	.8	.1	1.2	.0	.0011	.0007
T 1.0 1	04-11-01	.27	.01	<.4	.2	1.4	.0	.0079	.0009
Hermit Spring	12-04-00	<.03	.01	.7	.0	.71	.02	.0012	.0007
	04-11-01	.14	.00	<.4	.2	.75	.01	.0033	.0002
Boucher East Spring	05-26-00	.08	.02	<.5	.2	1.2	.0	.0016	.0007
	12-04-00	.10	.01	.7	.2	1.5	.0	.0014	.0003
	04-12-01	.21	.01	<.4	.2	1.4	.0	.0017	.0008

				ent concentrat	ions—continu				
Site	Date	-	lissolved li) ICP-MS	Praseodymiu (μg/L as P		Rhenium, (µg/L as R			, dissolve (b) ICP-MS
	_	Avg	SD	Avg	SD	Avg	SD	Avg	SD
Red Canyon Spring	08-23-01	nc	nc	nc	nc	nc	nc	nc	nc
	09-26-01	<.1	.2	.0002	.0000	.0094	.0002	4.8	.1
T Spring	04-08-01	<.2	.2	.0032	.0002	.014	.001	2.5	.0
	05-11-01	<.2	.0	.0003	.0001	.020	.001	8.4	.0
finers Spring	05-24-00	1.0	.2	.0002	.0001	.013	.001	8.3	.1
1 0	11-28-00	<.1	.0	<.0002	.0001	.013	.001	8.0	.2
	04-07-01	<.2	.3	.0007	.0002	.012	.000	7.7	.1
ottonwood Creek No. 1	05-25-00	3.6	.1	.0002	.0000	.0089	.0001	1.3	.0
ottonwood Creek No. 2	11-29-00	<.1	.1	.0005	.0000	.0090	.0003	4.3	.2
ottonwood Creek 110. 2	04-09-01	.2	.1	.0005	.0002	.0093	.0005	4.5	.0
rapevine Main Spring	04-09-01	<.2	 .0	.0000	.0002	.0037	.0003	2.9	.0
rapevine main opinig	04-10-01 04-30-01	<.2 <.2	.0 .3	.0001	.0001	.0037	.0003	2.9 3.0	.0
ronavina Fact Series		<.2 3.0	.3 1.3	.0003	.0001	.0038	.0002	3.0 8.3	
rapevine East Spring	05-25-00								.0
	11-29-00	<.1	.1	.0003	.0000	.060	.002	12	1
	12-12-00	nc	nc	nc	nc	nc	nc	nc	nc
	04-09-01	<.2	.3	.0002	.0000	.056	.001	20	0
onetree Spring	04-11-01	<.2	.3	.0062	.0004	.026	.001	12	0
	05-01-01	<.2	.5	.0067	.0003	.026	.000	13	0
am Magee Spring	04-20-01	nc	nc	nc	nc	nc	nc	nc	nc
Burro Spring	05-22-00	2.4	.3	.0006	.0001	.028	.000	6.6	.1
	12-07-00	<.1	.0	.0005	.0001	.027	.000	5.4	.1
	04-08-01	<.2	.4	.0004	.0001	.026	.000	6.2	.1
ipe Creek	05-22-00	2.6	.2	.0009	.0003	.023	.002	4.6	.0
pe Cleek	12-07-00	<.1	.1	.0005	.0000	.022	.001	3.7	.1
	04-08-01	<.2	.3	.0010	.0002	.020	.001	3.7	.0
umphouse Spring	05-22-00	1.5	.7	.0007	.0001	.013	.000	3.8	.0
	12-07-00	<.1	.0	.0006	.0000	.012	.000	3.4	.1
	12-07-01	<.2	.3	.0007	.0002	.012	.000	3.3	.0
umphouse Wash Gage	12-07-00	<.1	.0	<.0002	.0000	.011	.000	3.6	.1
	12-09-00	nc	nc	nc	nc	nc	nc	nc	nc
orn Creek	05-22-00	1.6	1.1	.0012	.0001	.056	.000	7.4	.0
	12-06-00	<.1	.0	.0005	.0001	.058	.000	7.2	.0
	04-07-01	<.2	.3	.0004	.0000	.078	.001	6.0	.1
alt Creek Spring	05-23-00	2.2	.4	.0001	.0001	.15	.00	8.3	.0
ı Ç	12-06-00	<.1	.0	<.0002	.0001	.14	.00	7.4	.2
	04-10-01	<.2	.3	<.0001	.0001	.15	.00	8.1	.1
Ionument Creek No. 1	05-24-00	2.4	1.1	.0002	.0000	.038	.001	5.7	.0
Ionument Spring	12-05-00	.1	.0	<.0002	.0000	.025	.001	4.4	.0 .0
Spring	04-09-01	<.2	.3	<.0002	.0000	.025	.001	4.3	.0
awaii Spring	05-25-00	1.8	.1	.0001	.0000	.023	.001	4.1	.0
awan Spring	12-04-00	<.1	.0	<.0002	.0000	.020	.001	3.9	.0
	12-04-00 04-11-01			<.0002	.0001	.020	.001	3.9 3.9	.2
amait Canina		<.2	.4						
ermit Spring	12-04-00	<.1	.0	<.0002	.0001	.014	.000	3.3	.1
	04-11-01	<.2	.2	.0005	.0001	.015	.001	2.9	.0
oucher East Spring	05-26-00	1.2	.4	<.0001	.0000	.020	.001	3.3	.1
	12-04-00	<.1	.1	<.0002	.0001	.021	.000	2.7	.2
	04-12-01	<.2	.2	<.0001	.0000	.021	.002	2.4	.0

	_	Samarium, dissolved (µg/L as Sm) ICP-MS		Selenium, dissolved (µg/L as Se) ICP-MS		Strontium, dissolved (µg/L as Sr) ICP-MS		Tellurium, dissolved (µg/L as Te) ICP-MS	
Site	Date _	Avg	SD	(μy/L as a	SD	Avg	SD		SD
Red Canyon Spring	08-23-01	nc	nc	nc	nc	nc	nc	nc	nc
ceu Canyon Spring	09-26-01	.0016	.0003	3.2	.0	110	3	<.01	.01
T Spring	09-20-01	.0010	.0005	2.6	.0	164	2	.012	.00
1 Spring	04-08-01	<.0007	.0000	2.0 4.6	.1	169	2	<.007	.00
linana Sanina	05-24-00	.0007	.0004	2.3	.1	109	2	.014	.00
Ainers Spring	03-24-00 11-28-00	.0008	.0002	2.5 2.5	.1	173	2 3	.014 <.007	.00
	04-07-01	.0008	.0003	2.3	.1		3	<.007	.00
Cottonwood Creek No. 1					.2	162			
	05-25-00	.0021	.0006	.2		248	3	.094	.01
Cottonwood Creek No. 2	11-29-00	.0014	.0002	2.0	.1	118	3	<.007	.00.
	04-09-01	.0021	.0003	2.1	.1	117	3	<.007	.00.
rapevine Main Spring	04-10-01	.0021	.0006	1.3	.1	82	1	<.007	.00
	04-30-01	.0021	.0005	1.3	.0	83	1	<.007	.00
Frapevine East Spring	05-25-00	.0040	.0004	.4	.1	374	4	.063	.00
	11-29-00	.0011	.0005	2.7	.2	411	2	.023	.00
	12-12-00	nc	nc	nc	nc	nc	nc	nc	nc
	04-09-01	.0008	.0005	2.7	.2	522	2	.055	.01
onetree Spring	04-11-01	.0071	.0004	6.3	.1	309	2	.008	.00
	05-01-01	.0067	.0004	6.5	.1	329	9	.010	.00
am Magee Spring	04-20-01	nc	nc	nc	nc	nc	nc	nc	nc
urro Spring	05-22-00	.0011	.0001	4.0	.2	199	0	.029	.00
	12-07-00	.0015	.0006	5.0	.1	181	2	<.007	.00
	04-08-01	.0012	.0002	4.6	.1	188	3	.012	.00
ipe Creek	05-22-00	.0013	.0001	3.6	.0	175	2	.028	.00
ipe creek	12-07-00	.0012	.0004	4.1	.1	167	1	<.007	.00
	04-08-01	.0009	.0002	4.0	.1	159	1	.009	.00
umphouse Spring	05-22-00	.0019	.0001	2.9	.0	113	1	.023	.00
	12-07-00	.0021	.0001	3.2	.1	107	3	<.007	.00
	12-07-01	.0018	.0007	3.0	.1	107	1	.009	.00
umphouse Wash Gage	12-07-00	.0015	.0002	2.9	.1	108	2	<.007	.00
	12-09-00	nc	nc	nc	nc	nc	nc	nc	nc
lorn Creek	05-22-00	.0016	.0002	5.1	.1	239	3	.036	.00
	12-06-00	.0009	.0004	5.1	.1	236	4	<.007	.00
	04-07-01	.0009	.0008	5.6	.1	239	3	.013	.00
alt Creek Spring	05-23-00	<.0005	.0002	9.4	.0	254	1	.027	.00
1 0	12-06-00	.0006	.0000	11	0	244	4	<.007	.00
	04-10-01	<.0007	.0002	11	0	252	3	.011	.00
Ionument Creek No. 1	05-24-00	.0006	.0002	4.1	.1	380	3	.037	.00
Ionument Spring	12-05-00	.0015	.0004	3.3	.2	127	1	<.007	.00
	04-09-01	.0013	.0003	3.4	.1	124	2	<.007	.00
awaii Spring	05-25-00	.0008	.0003	3.3	.0	228	2	.023	.00
B	12-04-00	<.0005	.0003	3.5	.1	226	9	<.007	.00
	04-11-01	.0019	.0005	3.7	.1	225	1	.007	.00
ermit Spring	12-04-00	.0019	.0003	3.3	.1	152	2	<.007	.00
orinit opring	04-11-01	.0014	.0003	3.3	.1	132	2 1	<.007	.00
oucher East Spring	05-26-00	<.0005	.0003	2.8	.1 .0	237	5	.031	.00
Juctice Last optilig	12-04-00								
	12-04-00 04-12-01	.0005 .0007	.0005 .0007	3.3 3.2	.1 .1	216 208	11 4	<.007	.00

Site	Date	Terbium, (µg/L as T		Thallium, ( (µg/L as Tl		Thorium, (μg/L as T		Thulium, dissolved (µg/L as Tm) ICP-MS	
	-	Avg	SD	Avg	SD	Avg	SD	Avg	SD
Red Canyon Spring	08-23-01	nc	nc	nc	nc	nc	nc	nc	nc
	09-26-01	<.0001	.0001	.025	.004	<.0006	.0000	<.0001	.000
T Spring	04-08-01	.0009	.0000	.019	.001	.0011	.0002	.0003	.000
	05-11-01	.0001	.0001	.029	.000	.0002	.0001	<.0001	.000
Ainers Spring	05-24-00	.0001	.0000	.028	.002	.0005	.0004	<.0001	.000
	11-28-00	<.0001	.0001	.035	.002	.0006	.0004	.0001	.000
	04-07-01	.0001	.0001	.031	.008	<.0002	.0000	<.0001	.000
Cottonwood Creek No. 1	05-25-00	.0001	.0001	.005	.002	.0009	.0003	.0001	.000
Cottonwood Creek No. 2	11-29-00	<.0001	.0000	.014	.003	.0004	.0002	<.0001	.000
	04-09-01	.0002	.0000	.005	.001	<.0002	.0001	.0001	.000
Grapevine Main Spring	04-10-01	<.0001	.0001	.014	.002	<.0002	.0000	<.0001	.000
1 10	04-30-01	.0001	.0001	.012	.001	<.0002	.0002	<.0001	.000
Grapevine East Spring	05-25-00	.0005	.0001	.007	.000	.0016	.0001	.0002	.000
1 1 0	11-29-00	.0002	.0001	.014	.000	.0025	.0001	.0001	.000
	12-12-00	nc	nc	nc	nc	nc	nc	nc	nc
	04-09-01	.0003	.0000	.006	.001	.0007	.0001	.0001	.000
onetree Spring	04-11-01	.0012	.0000	.013	.001	.0003	.0001	.0005	.000
	05-01-01	.0013	.0002	.017	.002	<.0002	.0001	.0006	.000
am Magee Spring	04-20-01	nc	nc	nc	nc	nc	nc	nc	nc
Surro Spring	05-22-00	.0002	.0001	.008	.002	.0003	.0001	.0002	.000
uno opinig	12-07-00	.0002	.0001	.012	.002	.0005	.0002	.0001	.000
	04-08-01	.0002	.0000	.004	.002	.0003	.0001	.0001	.000
Pina Craak	05-22-00	.0003	.0000	.011	.002	.0003	.0001	.0002	.000
Pipe Creek	12-07-00	<.0001	.0001	.016	.003	.0007	.0006	.0002	.000
	04-08-01	.0004	.0000	.007	.001	<.0002	.0001	.0002	.000
Pumphouse Spring	05-22-00	.0004	.0000	.016	.004	.0002	.0001	.0001	.000
umphouse spring	12-07-00	<.0002	.0000	.015	.002	.0003	.0001	.0001	.000
	12-07-00	.0001	.0000	.015	.001	<.0002	.0001	.0001	.000
umphoungo Wash Gago	12-07-01	<.0001	.0000	.010	.000	.0002	.0001		.000
umphoupse Wash Gage	12-07-00							<.0001	
Laure Carala		nc	nc	nc 018	nc	nc	nc	nc	nc
Iorn Creek	05-22-00	.0003	.0000	.018	.001	.0004	.0002	.0002	.000
	12-06-00	.0002	.0001	.029	.002	.0005	.0002	.0001	.000
	04-07-01	.0002	.0001	.029	.000	.0002	.0000	.0001	.000
alt Creek Spring	05-23-00	.0001	.0000	.014	.000	.0006	.0000	<.0001	.000
	12-06-00	<.0001	.0000	.019	.003	.0004	.0001	<.0001	.000
	04-10-01	.0001	.0000	.010	.001	.0003	.0001	.0001	.000
Ionument Creek No. 1	05-24-00	.0003	.0000	.032	.001	.0010	.0009	.0008	.000
Ionument Spring	12-05-00	<.0001	.0000	.059	.003	<.0002	.0001	<.0001	.000
	04-09-01	<.0001	.0001	.047	.001	<.0002	.0002	<.0001	.000
lawaii Spring	05-25-00	.0001	.0000	.007	.003	.0006	.0004	<.0001	.000
	12-04-00	<.0001	.0000	.012	.001	.0002	.0000	<.0001	.000
	04-11-01	.0002	.0000	.007	.002	.0010	.0002	.0001	.000
lermit Spring	12-04-00	<.0001	.0000	.025	.004	<.0002	.0001	<.0001	.000
	04-11-01	.0001	.0000	.018	.003	<.0002	.0001	.0001	.000
Boucher East Spring	05-26-00	<.0001	.0000	.003	.003	.0005	.0004	<.0001	.000
	12-04-00	<.0001	.0001	.009	.002	.0004	.0003	<.0001	.000
	04-12-01	.0001	.0001	.003	.002	.0002	.0002	.0001	.000

		Ter 1	diameter 1	п .	diameter 1	N- "	dia and di	V	dian 1 1
Site	Date	•	dissolved V) ICP-MS	-	dissolved U) ICP-MS		, dissolved V) ICP-MS		i, dissolved (b) ICP-MS
	-	Avg	SD	Avg	SD	Avg	SD	Avg	SD
Red Canyon Spring	08-23-01	nc	nc	nc	nc	nc	nc	nc	nc
	09-26-01	.013	.000	1.7	.1	1.4	.0	<.0007	.0003
T Spring	04-08-01	.014	.001	3.5	.0	1.5	.0	.0026	.0001
	05-11-01	.017	.003	4.1	.1	1.8	.0	<.0002	.000
Ainers Spring	05-24-00	.018	.000	3.1	.0	2.0	.0	.0003	.000
	11-28-00	.022	.001	3.5	.0	2.1	.0	.0004	.0002
	04-07-01	.023	.001	3.3	.0	2.1	.0	<.0002	.0001
Cottonwood Creek No. 1	05-25-00	.003	.002	1.5	.0	<.1	.0	.0007	.0002
Cottonwood Creek No. 2	11-29-00	.005	.001	1.6	.0	.7	.0	.0004	.0001
	04-09-01	.005	.002	1.6	.0	.8	.0	.0005	.0001
Grapevine Main Spring	04-10-01	.006	.001	1.1	.0	.9	.1	<.0002	.0001
	04-30-01	.006	.000	1.1	.0	.9	.0	<.0002	.0001
Grapevine East Spring	05-25-00	.010	.000	2.1	.1	.2	.0	.0013	.0001
Shapevine Last Spring	11-29-00	.010	.001	8.3	.0	.2	.0 .0	.0013	.0001
	12-12-00	nc	nc	nc	nc	nc	nc	nc	nc
	04-09-01	.003	.002	nc 7.1	.1	<.3	.0	.0008	.0002
onetree Spring	04-09-01	.005	.002	6.0	.1	.9	.0	.0036	.0002
concuree spring		.003	.002	6.0 6.0	.1 .1	.9 1.0	.1	.0038	.0001
am Magee Spring	05-01-01								
	04-20-01	nc 004	nc	nc	nc	nc	nc	nc	nc
Burro Spring	05-22-00	.004	.002	2.5	.0	.5	.0	.0008	.0001
	12-07-00	.002	.000	2.7	.0	.3	.0	.0008	.0002
	04-08-01	.003	.001	2.4	.0	.4	.2	.0009	.0001
ipe Creek	05-22-00	.007	.000	2.7	.1	1.0	.0	.0008	.0000
	12-07-00	.002	.001	2.4	.0	.5	.0	.0008	.0003
	04-08-01	<.003	.000	2.3	.0	.7	.0	.0006	.0001
Pumphouse Spring	05-22-00	.005	.001	1.7	.0	1.0	.0	.0005	.0002
	12-07-00	.004	.001	1.9	.0	.9	.0	.0008	.0002
	12-07-01	.003	.001	1.8	.1	.9	.0	.0007	.0001
Pumphouse Wash Gage	12-07-00	.005	.001	1.8	.0	1.0	.0	<.0003	.0001
	12-09-00	nc	nc	nc	nc	nc	nc	nc	nc
Iorn Creek	05-22-00	.005	.002	8.6	.2	1.2	.0	.0012	.0002
	12-06-00	.005	.002	9.3	.3	1.1	.1	.0007	.0003
	04-07-01	<.003	.002	29	0	.6	.0	.0006	.0001
alt Creek Spring	05-23-00	.007	.002	30	0	.7	.0	<.0001	.0001
	12-06-00	.007	.000	31	0	.6	.0	<.0003	.0003
	04-10-01	.007	.002	29	0	.6	.1	<.0002	.000
Ionument Creek No. 1	05-24-00	.004	.001	7.1	.2	.5	.0	.0063	.0000
Ionument Spring	12-05-00	.005	.001	7.1	.1	.7	.0	<.0003	.0002
	04-09-01	.005	.001	7.3	.1	.7	.0	<.0002	.0002
Iawaii Spring	05-25-00	.004	.001	1.9	.0	.6	.0	<.0001	.0000
	12-04-00	.002	.000	1.9	.0	.6	.0	<.0003	.0002
	04-11-01	.007	.003	2.0	.0	.6	.2	.0005	.0003
Iermit Spring	12-04-00	.005	.002	2.0	.0	.8	.1	<.0003	.0001
	04-11-01	.006	.003	2.1	.0	1.0	.1	.0002	.0002
Boucher East Spring	05-26-00	.003	.001	1.7	.1	.4	.0	<.0001	.000
1.0	12-04-00	.002	.001	1.8	.0	.4	.1	<.0003	.0002
	04-12-01	.003	.002	1.9	.1	.5	.1	<.0002	.0001

			ent concentrat				
Site	Date		dissolved Y) ICP-MS		issolved (n) ICP-MS		ı, dissolved Zr) ICP-MS
	-	Avg	SD	Avg	SD	Avg	SD
ed Canyon Spring	08-23-01	nc	nc	nc	nc	nc	nc
	09-26-01	.0057	.0002	2.6	.0	.002	.001
T Spring	04-08-01	.038	.000	57	1	.11	.00
	05-11-01	.0044	.0003	3.7	.1	.005	.001
liners Spring	05-24-00	.0069	.0003	3.8	.3	<.004	.001
	11-28-00	.0083	.0006	50	0	<.003	.001
	04-07-01	.0052	.0005	2.9	.3	.007	.002
ottonwood Creek No. 1	05-25-00	.016	.002	3.7	.7	.013	.001
Cottonwood Creek No. 2	11-29-00	.011	.000	1.9	.0	<.003	.001
	04-09-01	.015	.000	5.9	.1	.015	.012
Frapevine Main Spring	04-10-01	.0056	.0008	56	0	.009	.000
. 10	04-30-01	.0055	.0000	3.7	.1	.004	.001
rapevine East Spring	05-25-00	.024	.001	4.5	.3	.020	.003
r r c	11-29-00	.014	.000	1.4	.1	.032	.004
	12-12-00	nc	nc	nc	nc	nc	nc
	04-09-01	.012	.001	3.2	.3	.021	.001
onetree Spring	04-11-01	.069	.001	.4	.2	.007	.000
1 8	05-01-01	.072	.001	26	0	.020	.002
am Magee Spring	04-20-01	nc	nc	nc	nc	nc	nc
urro Spring	05-22-00	.021	.000	13	0	<.004	.001
8	12-07-00	.015	.000	1.4	.1	.004	.001
	04-08-01	.017	.000	4.9	.0	.006	.002
ipe Creek	05-22-00	.019	.002	3.3	.1	.009	.002
	12-07-00	.015	.001	1.8	.0	<.003	.000
	04-08-01	.016	.001	2.7	.1	.005	.001
umphouse Spring	05-22-00	.012	.000	5.9	.1	.005	.001
	12-07-00	.013	.000	3.7	.1	<.003	.000
	12-07-01	.013	.001	19	0	.003	.001
umphouse Wash Gage	12-07-00	.0057	.0007	23	0	<.003	.000
umphouse wash ouge	12-09-00	nc	nc	nc	nc	nc	nc
Iorn Creek	05-22-00	.027	.001	2.9	.1	<.004	.001
	12-06-00	.014	.001	1.8	.0	.004	.000
	04-07-01	.018	.000	8.3	.2	.007	.000
alt Creek Spring	05-23-00	.0068	.0003	4.7	.1	<.004	.000
an creen spring	12-06-00	.0031	.0003	1.3	.0	<.003	.001
	04-10-01	.0041	.0007	25	0	.005	.000
Ionument Creek No. 1	05-24-00	.035	.001	4.5	.1	.017	.000
Ionument Spring	12-05-00	.0037	.0002	6.4	.1	.005	.000
oping	04-09-01	.0038	.0002	5.8	.2	.005	.000
lawaii Spring	05-25-00	.0085	.0009	3.1	.2	<.004	.001
annan opring	12-04-00	.0053	.0005	9.3	.0	<.004	.000
	04-11-01	.0055	.0000	23	1.0	.028	.000
lermit Spring	12-04-00	.0093	.0010	4.0	.1	<.003	.002
ernnt opring	04-11-01	.0031	.0002	4.0 2.5	.1	<.003 .006	.000
oucher East Spring	05-26-00	.0071	.0000	1.8	.2 .7	<.004	.001
	05-20-00	.0050	.0011	1.0	./	<.004	.002
oucher Last Spring	12-04-00	.0040	.0004	1.9	.1	<.003	.000

			Radio	activity			
Site	Date	Gross-beta, dissolved (pCi/L as <sup>137</sup> Cs)	Gross-beta, PE dissolved (pCi/L as <sup>137</sup> Cs)	Gross-beta, MDC dissolved (pCi/L as <sup>137</sup> Cs)	Gross-alpha, dissolved (pCi/L as <sup>230</sup> Th)	Gross-alpha, PE dissolved (pCi/L as <sup>230</sup> Th)	Gross-alpha, MDC dissolved (pCi/L as <sup>230</sup> Th)
Red Canyon Spring	08-23-01	nc	nc	nc	nc	nc	nc
	09-26-01	3.7	1.4	1.9	3.1	2.0	3.1
JT Spring	04-08-01	nc	nc	nc	nc	nc	nc
	05-11-01	5.4	2.3	3.7	3.6	1.8	2.1
Miners Spring	05-24-00	<4	4.2	na	3.9	3.0	na
	11-28-00	nc	nc	nc	nc	nc	nc
	04-07-01	nc	nc	nc	nc	nc	nc
Cottonwood Creek No. 1	05-25-00	4.5	4.9	na	<3	2.0	na
Cottonwood Creek No. 2	11-29-00	<4	2.3	na	<3	1.1	na
	04-09-01	nc	nc	nc	nc	nc	nc
Grapevine Main Spring	04-10-01	nc	nc	nc	nc	nc	nc
	04-30-01	2.6	1.6	2.7	2.6	1.5	1.9
Grapevine East Spring	05-25-00	5.6	5.0	na	<3	1.6	na
	11-29-00	nc	nc	nc	nc	nc	nc
	12-12-00	nc	nc	nc	nc	nc	nc
	04-09-01	nc	nc	nc	nc	nc	nc
Lonetree Spring	04-11-01	nc	nc	nc	nc	nc	nc
	05-01-01	15.5	2.5	3.2	5.3	1.9	1.9
Sam Magee Spring	04-20-01	nc	nc	nc	nc	nc	nc
Burro Spring	05-22-00	4.4	4.6	na	<3	2.2	na
Durro opring	12-07-00	nc	nc	nc	nc	nc	nc
	04-08-01	nc	nc	nc	nc	nc	nc
Pipe Creek	05-22-00	<4	4.5	na	<3	2.4	na
ripe creek	12-07-00	nc	nc	nc	nc	nc	nc
	04-08-01	nc	nc	nc	nc	nc	nc
Pumphouse Spring	05-22-00	<4	4.4	na	3.9	3.0	na
r umphouse spring	12-07-00	nc	nc	nc	nc	nc	nc
	04-07-01	nc	nc	nc	nc	nc	nc
Pumphouse Wash Gage	12-07-00		nc	nc		nc	nc
Fullphouse wash Gage	12-07-00	nc			nc		
Horn Creek	05-22-00	nc 5.3	nc 4.6	nc	nc 5.9	nc 3.3	nc
Holli Cleek				na			na
	12-06-00 04-07-01	nc	nc	nc	nc	nc	nc
Salt Creat Spring	05-23-00	nc	nc 5.0	nc	nc 22.0	nc	nc
Salt Creek Spring	12-06-00	12.1		na		5.5	na
		nc	nc	nc	nc	nc	nc
	04-10-01	nc	nc	nc	nc	nc	nc
Monument Creek No. 1	05-24-00	nc	nc	nc	nc	nc	nc
Monument Spring	12-05-00	<4	2.6	na	4.0	1.5	na
	04-09-01	nc	nc	nc	nc	nc	nc
Hawaii Spring	05-25-00	<4	4.4	na	<3	2.6	na
	12-04-00	nc	nc	nc	nc	nc	nc
	04-11-01	nc	nc	nc	nc	nc	nc
Hermit Spring	12-04-00	nc	nc	nc	nc	nc	nc
	04-11-01	nc	nc	nc	nc	nc	nc
Boucher East Spring	05-26-00	<4	4.2	na	<3	2.7	na
	12-04-00	nc	nc	nc	nc	nc	nc
	04-12-01	nc	nc	nc	nc	nc	nc

			isotop	ic data			
Site	Date	δ <sup>18</sup> 0 (per mil)	δD (per mil)	<sup>87</sup> Sr/ <sup>86</sup> Sr	<sup>14</sup> C (percent modern carbon)	<sup>14</sup> C count error (percent modern carbon)	δ <sup>13</sup> C (per mil
ed Canyon Spring	08-23-01	nc	nc	nc	nc	nc	nc
	09-26-01	-12.7	-94.2	0.71164	49.2	0.5	-8.59
Г Spring	04-08-01	-9.4	-73.1	nc	nc	nc	nc
	05-11-01	-12.2	-91.4	.71250	55.1	.4	-8.05
liners Spring	05-24-00	-12.3	-93.1	.71196	71.1	.5	-7.69
	11-28-00	-12.2	-90.7	nc	nc	nc	nc
	04-07-01	-12.1	-92.3	nc	nc	nc	nc
ottonwood Creek No. 1	05-25-00	-12.3	-91.6	.71374	94.2	.7	-13.14
ottonwood Creek No. 2	11-29-00	-12.7	-93.9	.71264	60.3	.4	-10.00
	04-09-01	-12.8	-93.8	nc	nc	nc	nc
apevine Main Spring	04-10-01	-12.9	-94.6	nc	nc	nc	nc
	04-30-01	-12.9	-92.7	.71140	54.3	.4	-8.08
apevine East Spring	05-25-00	-9.1	-73.6	.71514	103.4	.6	-11.05
	11-29-00	nc	nc	nc	nc	nc	nc
	12-12-00	-8.9	-74.8	nc	89.7	.6	-6.53
	04-09-01	-8.5	-71.0	nc	nc	nc	nc
onetree Spring	04-11-01	-11.9	-89.1	nc	nc	nc	nc
	05-01-01	-12.0	-89.9	.71327	70.2	.5	-10.78
m Magee Spring	04-20-01	-10.0	-79.4	nc	nc	nc	nc
rro Spring	05-22-00	-12.4	-91.0	.71435	63.6	.4	-9.81
	12-07-00	-12.3	-92.9	nc	nc	nc	nc
	04-08-01	-12.4	-91.1	nc	nc	nc	nc
pe Creek	05-22-00	-12.3	-91.9	.71440	54.6	.4	-10.29
	12-07-00	-12.4	-91.5	nc	nc	nc	nc
	04-08-01	-12.4	-90.9	nc	nc	nc	nc
mphouse Spring	05-22-00	-12.3	-92.6	.71190	51.4	.4	-9.27
1 10	12-07-00	-12.3	-93.1	nc	nc	nc	nc
	04-07-01	-12.3	-92.8	nc	nc	nc	nc
mphouse Wash Gage	12-07-00	nc	nc	nc	nc	nc	nc
1	12-09-00	-12.3	-93.0	nc	nc	nc	nc
orn Creek	05-22-00	-11.9	-88.8	.70363	74.9	.5	-10.78
	121-06-00	-11.7	-89.3	nc	nc	nc	nc
	04-07-01	-11.8	-88.9	nc	nc	nc	nc
lt Creek Spring	05-23-00	-11.8	-87.3	.71248	40.5	.3	-6.34
it crook spring	12-06-00	-12.1	-90.2	nc	nc	nc	nc
	04-10-01	-11.7	-87.1	nc	nc	nc	nc
onument Creek No. 1	05-24-00	nc	nc	nc	nc	nc	nc
onument Spring	12-05-00	-12.2	-91.1	.71070	42.0	.3	-8.24
Spring	04-09-01	-12.2	-91.2	nc	nc	nc	-0.24 nc
waii Spring	05-25-00	-12.2	-88.3	.71152	38.3	.3	-7.63
unan opring	12-04-00	-11.9	-89.1	nc	nc	nc	-7.03 nc
	04-11-01	-11.8	-89.1	nc	nc		
armit Spring	12-04-00	-11.9	-88.9 -89.7			nc	nc
ermit Spring				nc .71009	nc	nc	nc
mahar Foot Series	04-11-01	-11.8	-88.2		nc 72.2	nc	nc
oucher East Spring	05-26-00	-11.4	-84.1	.71104	72.2	.5	-6.69
	12-04-00	nc	nc	nc	nc	nc	nc

**Table 8**.
 Physical and chemical properties, chemical-constituent concentrations, radioactivity, and isotopic data for water from springs and creeks that discharge along the south rim of Grand Canyon, Arizona, 2000–2001—Continued

			lsotopic data-	-Continued			
Site	Date	Tritium (pCi/L)	Tritium, 2-Sigma (pCi/L)	Tritium (TU)	Residence time, uncorrected (years rounded)	Average dilution factor	Average corrected residence time (years rounded)
Red Canyon Spring	08-23-01	nc	nc	nc	nc	nc	nc
	09-26-01	2.2	1	.7	5,600	1.52	2,400
JT Spring	04-08-01	13.1	1	4.1	nc	nc	nc
	05-11-01	3.5	1	1.1	4,700	1.65	1,500
Miners Spring	05-24-00	2.2	1	.7	2,700	1.32	500
	11-28-00	nc	nc	nc	nc	nc	nc
	04-07-01	2.6	1	.8	nc	nc	nc
Cottonwood Creek No. 1	05-25-00	2.2	1	.7	400	1.07	Modern <sup>1</sup>
Cottonwood Creek No. 2	11-29-00	1.6	1	.5	4,000	1.42	1,300
	04-09-01	1.9	1	.6	nc	nc	nc
Grapevine Main Spring	04-10-01	1.6	1	.5	nc	nc	nc
	04-30-01	3.5	1	1.1	4,900	1.5	1,700
Grapevine East Spring	05-25-00	1.9	1	.6	Modern <sup>1</sup>	.9	Modern <sup>1</sup>
	11-29-00	nc	nc	nc	nc	nc	nc
	12-12-00	nc	nc	nc	800	1.56	Modern <sup>1</sup>
	04-09-01	2.2	1	.7	nc	nc	nc
Lonetree Spring	04-11-01	5.1	1	1.6	nc	nc	nc
	05-01-01	4.8	1	1.5	2,800	1.22	1,300
Sam Magee Spring	04-20-01	≤ 1.0	1	≤.3	nc	nc	nc
Burro Spring	05-22-00	3.5	1	1.1	3,600	1.45	700
	12-07-00	nc	nc	nc	nc	nc	nc
	04-08-01	3.2	1	1.0	nc	nc	nc
Pipe Creek	05-22-00	3.5	1	1.1	4,800	1.38	2,300
Tipe creek	12-07-00	nc	nc	nc	nc	nc	nc
	04-08-01	2.9	1	.9	nc	nc	nc
Pumphouse Spring	05-22-00	≤1	1	.,∍ ≤.3	5,300	1.54	1,900
r umphouse spring	12-07-00	nc	nc	nc	nc	nc	nc
	04-07-01	.6	1	.2	nc	nc	nc
Pumphouse Wash Gage	12-07-00	nc		nc		nc	
Fullphouse wash Gage	12-07-00		nc		nc		nc
Horn Creek	05-22-00	nc 6.4	nc	nc 2.0	nc	nc	nc 800
Hom Creek	121-06-00		1		2,300	1.22	
		nc 7 0	nc	nc	nc	nc	nc
Salt Creat Series	04-07-01	7.0	1	2.2	nc 7 200	nc	nc
Salt Creek Spring	05-23-00 12-06-00	1.3	1	.4	7,200	1.95	2,000
		nc	nc	nc	nc	nc	nc
	04-10-01	3.2	1	1.0	nc	nc	nc
Monument Creek No. 1	05-24-00	nc	nc	nc	nc	nc	nc
Monument Spring	12-05-00	1.3	1	.4	7,000	1.6	3,400
HC .	04-09-01	2.2	1	.7	nc	nc	nc
Hawaii Spring	05-25-00	1.6	1	.5	7,700	1.74	3,400
	12-04-00	nc	nc	nc	nc	nc	nc
	04-11-01	≤ 1.0	1	≤.3	nc	nc	nc
Hermit Spring	12-04-00	nc	nc	nc	nc	nc	nc
	04-11-01	nc	nc	nc	nc	nc	nc
Boucher East Spring	05-26-00	$\leq 1$	1	≤.3	2,600	1.53	Modern <sup>1</sup>
	12-04-00	nc	nc	nc	nc	nc	nc
	04-12-01	1.6	1	0.5	nc	nc	nc

<sup>1</sup>Less than 50 years.

**Table 9**.
 X-ray diffraction mineralogy and isotope data for selected rock samples representing the major stratigraphic units present near the

 Bright Angel Fault in Grand Canyon, Arizona

Site	Depth <sup>1</sup> (m)	Rock name	Rock description	Bulk mineralogy	<sup>87</sup> Sr/ <sup>86</sup> Sr	δ <sup>13</sup> C (per mil)
GCBA-16	50	Kaibab Formation	Beige, fine-grained massive limestone, forms blocky ledges, many fossils	quartz (62%), calcite (19%), dolomite (10%), k-feldspar (7%), clays-kaolinite (1%)	0.70918	-0.5
GCBA-15	128	Toroweap Formation	Sandy limestone, beige-red with small black flecks; resistant, blocky ledges		.70829	1.79
GCBA-14	148	Coconino Sandstone	White colored fine-grained aeolian sandstone	quartz (97%), clays-kaolinite (3%)	.71173	na
GCBA-13	222	Coconino Sandstone	White colored fine-grained aeolian sandstone	quartz (63%), k-feldspar (31%), clays-kaolinite (6%)	.71122	na
GCBA-12	283	Hermit Formation	Red mudstone	quartz (43%), dolomite (22%), clays-illite (35%)	.71216	-2.32
GCBA-11	332	Esplanade Sandstone of Supai Group	Red massive fine-grained sandstone; cliff-forming	quartz (66%), dolomite (19%), clays-kaolinite (15%)	.71035	-1.71
GCBA-10	436	Wescogame Formation of Supai Group	Red silty sandstone; fine-bedded; slope-forming unit	calcite (72%), quartz (19%), k-feldspar (5%), clays-illite (4%)	.70877	-3.81
GCBA-09	467	Manakacha Formation of Supai Group	Gray, fine-bedded sandy limestone forming stair-stepping ledges	calcite (68%), quartz (24%), k-feldspar (3%), clays-illite (5%)	.70871	-3.86
GCBA-08	562	Watahomigi Formation of Supai Group	Red mudstone	quartz (82%), clays-kaolinite (18%)	.71083	84
GCBA-07	607	Horseshoe Mesa member of Redwall Limestone	Fine-grained gray-green; slightly platy; blocky ledges	calcite (98%), quartz (2%)	.70820	26
GCBA-06	651	Mooney Falls member of Redwall Limestone	Massive medium-grained limestone; green rose color	dolomite (97%), calcite (3%)	.70902	04
GCBA-05	736	Thunder Springs member of Redwall Limestone	Massive fine-bedded limestone with numerous white chert beds; red-stained with some spalling on surface	dolomite (77%), calcite (23%)	.71012	1.21
GCBA-04	743	Whitmore Wash member Redwall Limestone	Fine-grained limestone forming blocky gray-purple ledges; at top of ledge is a white chert layer	dolomite (77%), calcite (23%)	.70954	32
GCBA-03	764	Temple Butte Formation	Dolomitic limestone, beige to light gray, fine-grained, heavily weathered surface;	dolomite (100%)	.70879	-1.27
GCBA-02	786	Muav Limestone	Sandy limestone, beige to white; coarse, irregular beds; local vertical fractures	k-feldspar (33%), quartz (30%), dolomite (28%), calcite (1%), clays-kaolinite (8%)	.71014	-1.04
GCBA-01	885	Bright Angel Shale	Upper member of Bright Angel shale; cliff forming; greenish, fine-grained shale; fine bedded	dolomite (52%), quartz (19%), k-feldspar (11%), clays-illite (18%)	.71018	-1.06

<sup>1</sup>Depth is calculated from the Bright Angel Fault at the south rim of Grand Canyon at an altitude of 1,945 meters above sea level.

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# **Appendix**

### **Equation 1**

 $\delta = \left[ \frac{R_{sample} - R_{std}}{R_{std}} \right] \ge 1,000$ 

where:

 $\delta$  = delta notation,  $R_{sample}$  = ratio of isotopes measured in sample, and  $R_{std}$  = ratio of same isotopes in the standard.

#### Equation 2

 $^{2}\text{H} = 8 \ ^{18}\text{O} + 10$ 

## Equation 3

 $t = -17.93 \ln(a_t {}^{3}\text{H}/a_o {}^{3}\text{H})$ 

where:

t	=	time in years,
$a_t {}^3 \mathrm{H}$	=	initial tritium activity in TU, and
<i>a</i> <sub>o</sub> <sup>3</sup> H	=	residual tritium activity in TU.

## Equation 4

 $t = (5,568/(\ln 2) \times [\ln (100/pmc)] - [(\gamma - 1950)/1.029]$ 

where:

t	=	radiocarbon age,
5,568	=	libby half-life,
ртс	=	percent modern carbon,
γ	=	decay constant (equal to ln 2 divided by the <sup>14</sup> C half-life), and
1.029	=	constant that is the ratio of $^{\gamma}5730/^{\gamma}5568$ .

### Equation 5

Dilution factor = D =  $(\delta^{13}C(_{RS}) - \delta^{13}C(_{soil-gas})) / (\delta^{13}C(_{RS}) - \delta^{13}C(_{water}))$ where:

 $\delta^{13}C(_{RS}) = \text{measured } \delta^{13}C \text{ from rock samples } (_{RS}),$   $\delta^{13}C(_{\text{soil-gas}}) = \text{estimated value for } \delta^{13}C \text{ for soil-gas in the unsaturated zone, and}$  $\delta^{13}C(_{\text{water}}) = \text{measured } \delta^{13}C \text{ from the water sample.}$ 

### Equation 6

$$^{14}C(adjusted) = {}^{14}C(pmc) \ge D$$

where:

pmc = measured percent modern carbon and

D =dilution factor.

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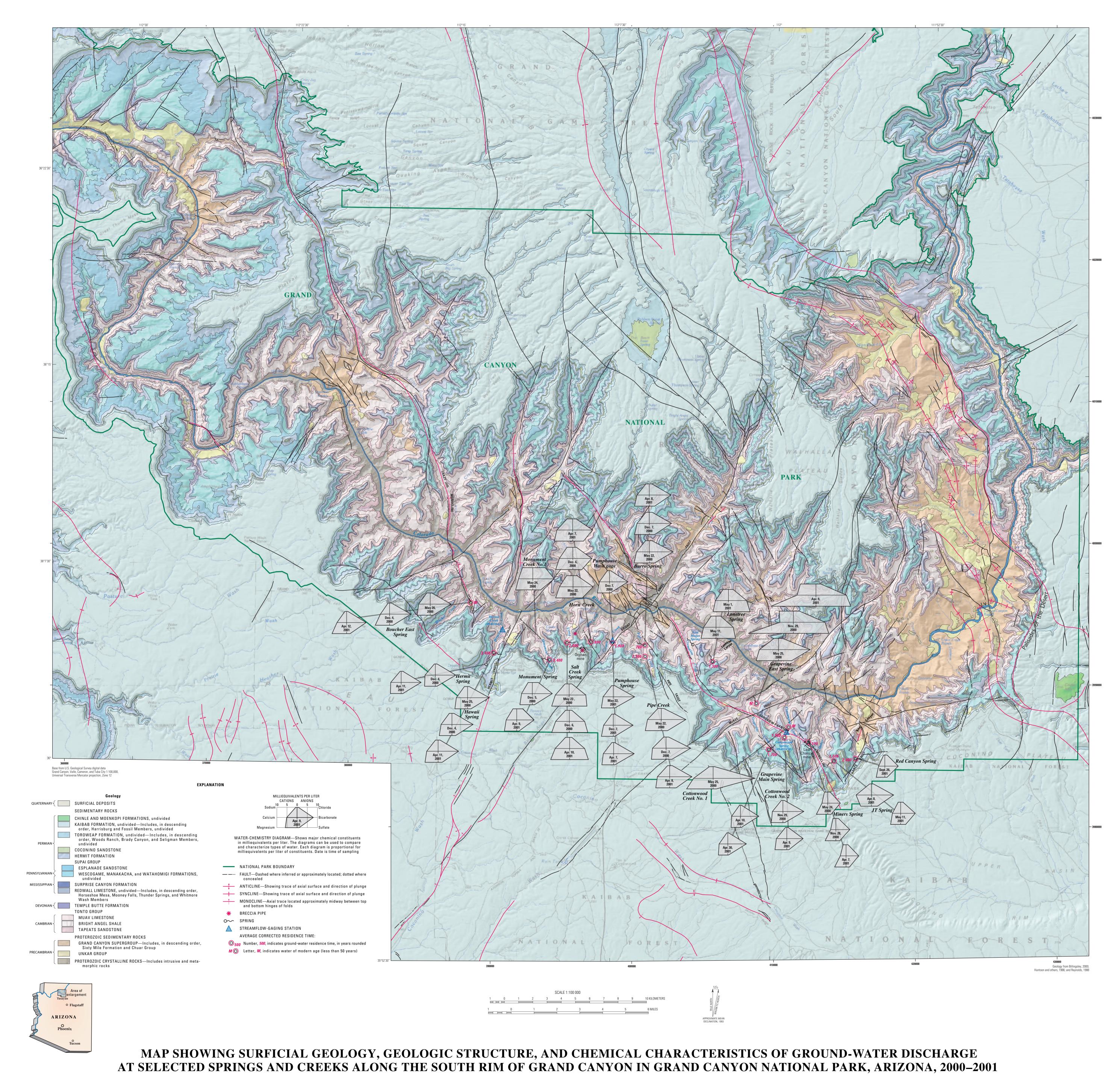
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