

Application of Minimum Support Constraints To Seismic Traveltime Tomography Jonathan Ajo–Franklin and Burke Minsley, Earth Resources Laboratory, Massachusetts Institute of Technology

We explore adaptive regularization methods based on compactness useful for imaging localized features. We apply Portniaguine & Zhdanov's (1999,2002) minimum support approach to generate compact solutions to the seismic traveltime tomography problem. We also explore a second type of compactness constraint based on a spatially varying exponential damping matrix. Both methods are applied to a simple synthetic and a more realistic test based on results from a 2-phase flow simulation of DNAPL infiltration.

1. Introduction : Why Regularize?

Geophysical inverse problems are often ill-posed in that an infinite number of models will fit a given dataset. When confronted with a multitude of answers, we can add secondary constraints to select for models which fulfill some independent notion of what a "good" solution should look like. Regularization accomplishes this by minimizing a weighted norm of the solution, typically the product of a low order differential operator and the model vector. By minimizing the 1st or 2nd order spatial derivatives of the model, flat or smooth solutions can be selected. Despite their wide (and often effective) use, 1^{st} and 2^{nc} order Tikhonov regularization are not fundamentally tied to the imaging problem; afterall, why should earth properties be flat or smooth?

We advocate choosing a regularization operator which incorporates the physics responsible for property variations such as subsurface flow, thermal diffusion, fracture propagation, or sediment deposition. Within this framework, we select models which both fit the data and reflect our understanding of the cause behind geophysical anomalies. Due to the complexity of these processes, we are currently examining heuristic techniques for including this class of constraints. In particular, we are interested in monitoring problems where changes in earth properties are due to fluid movement. Since flow processes tend to localize in zones of high permeability, regularization operators favoring *compact* and/or connected anomalies seem reasonable. Minimum support inversion, as developed by Last & Kubik (1983) and more recently Portniaguine & Zhdanov (1999,2002) is one approach for selecting compact models while still satisfying the data

2. Minimum Support Inversion

Assume a linear problem where an operator **G** maps a model vector, **m**, to a dataset, **d** generating a system of the form $\mathbf{Gm} = \mathbf{d}$. Traditional Tikhonov regularization minimizes a combination of data misfit and weighted model length,

$$\Phi(m) = \left\| \boldsymbol{G} \ m - d \right\|_{2}^{2} + \lambda^{2} \left\| \boldsymbol{W} \ m \right\|_{2}^{2} \longrightarrow \left[\begin{array}{c} \boldsymbol{G} \\ \lambda \ \boldsymbol{W} \end{array} \right] m = \left[\begin{array}{c} d \\ 0 \end{array} \right]$$

where **W** is a weighting matrix and λ is a trade-off or regularization parameter. In minimum support inversion we select an objective function of the form

$$\Phi_{ms}(m) = \left\| \boldsymbol{G} \ m - d \right\|_{2}^{2} + \lambda^{2} \sum_{k=1}^{N_{m}} \frac{m_{k}^{2}}{m_{k}^{2} + \beta^{2}}$$

where β is parameter designed to prevent singularities for $\mathbf{m} = \mathbf{0}$. The solution to this objective function can be cast in terms of a model dependent diagonal weighting operator W where

$$\boldsymbol{W}_{\boldsymbol{e}}(\boldsymbol{m}_{p}) = diag \left[\boldsymbol{m}_{p}^{2} + \boldsymbol{\beta}^{2}\right]^{-1/2}$$

Since $\mathbf{W}_{\mathbf{x}}$ is dependent on \mathbf{m} , the resulting problem is non-linear and we must resort to a model-domain iteratively reweighted least squares approach (IRLS) – in contrast, hybrid L1/L2 solvers reweight data residuals, not the model space regularization term. At each iteration, a new W is calculated using the previous estimate of m and the resulting linear problem is solved. Of course a prior model is needed to start the process

When **m** is not homogeneous, **W** is spatially varying with higher weights in locations where **m** is close to 0. Consider a similar type of weighting matrix based on an scaled exponential of the the model,

$$\frac{-|m_{p}|}{\boldsymbol{w}_{\boldsymbol{e}}(m_{p})} = \lambda_{1} + (\lambda_{2} - \lambda_{1}) \ diag \ \boldsymbol{e}^{-\sigma}$$

Since the exponential evaluates between 0 and 1, the two λ values control the maximum and minimum weights in ${f W}$ while σ controls the width of the objective function near 0 values. The best way to understand W is to examine the resulting weighting and penalty functions,



Figure A shows the value of an element of **W** as a function of **m** near 0 for the minimum support and exponential weighting regularization terms. Note that the maximum weights vary as a function of beta for the minimum support case. Figure B shows the corresponding penalty functions in comparison to standard damping (in green).

What does this mean? As we proceed, in each non-linear iteration W adaptively suppresses values of **m** near zero. In the case of the minimum support operator, no selective preference is given to large model values. In the exponential weighting case, model values near a selected band are suppressed. In both cases, zones of support are minimized as a function of iteration.

3. Traveltime Tomography : **A Minimum Support Example**

Traveltime tomography is ...

- ⁾ Useful in a variety of near–surface applications
- Particularly powerful in a monitoring context
- Multiple modalities (seismic,, radar, etc)
- Often ill-posed (depends on descritization)
- Typically low-resolution (T1, T2 reg.)
- Suffers from aperture-related artifacts

Can we use minimum support regularization to improve image quality?

A Test Example

System solved using LSQR

- Straight ray calculation of **G**
- Crosswell geometry (40 S x 40 R) Compact velocity anomalies
- Variable ray coverage (see map)

True Velocity Model Sensitivity Map (Used to generate synthetic datasets)





20 30 40 X Pos. (m)



Jason Gerhard and Bernie Kueper of Waterloo University generously donated the DNAPL flow results used in synthetic example #2 (see cite 5).

Quartz used used for grain dielectric (k=4.5)