

**“Measurements for Assessment of
Hydrate Related Geohazards”**

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Abstract

Natural gas hydrate deposits are found in deep offshore environments. In some cases these deposits overlay conventional oil and gas reservoirs. There are concerns that the presence of hydrates can compromise the safety of exploration and production operations [*Hovland and Gudmestad, 2001*]. Serious problems related to the instability of wellbores drilled through hydrate formations have been documented by *Collett and Dallimore, 2002*. A hydrate-related incident in the deep Gulf of Mexico could potentially damage the environment and have significant economic impacts.

Borehole and seafloor stability models are needed to predict potentially hazardous conditions. The inputs to those models will be provided by measurements. The purpose of this paper is to propose measurement schemes that are likely to be relevant to the safety problem. This survey is a starting point. It is anticipated that measurement requirements will co-evolve with our knowledge of hydrate deposits and our ability to model relevant aspects of them.

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1.0 Introduction

Natural gas hydrate deposits are found in deep offshore environments. In some cases these deposits overlay conventional oil and gas reservoirs. There are concerns that the presence of hydrates can compromise the safety of exploration and production operations [Hovland and Gudmestad, 2001]. Serious problems related to the instability of wellbores drilled through hydrate formations have been documented by Collett and Dallimore, [2002]. A hydrate-related incident in the deep Gulf of Mexico could potentially damage the environment and have significant economic impacts.

Borehole and seafloor stability models are needed to predict potentially hazardous conditions. The inputs to those models will be provided by measurements. The purpose of this paper is to propose measurement schemes that are likely to be relevant to the safety problem. This survey is a starting point. It is anticipated that measurement requirements will co-evolve with our knowledge of hydrate deposits and our ability to model relevant aspects of them.

Measurements will be needed during exploration, evaluation, and monitoring phases. The requirements during these phases are distinct from one another, and can be treated separately.

2.0 Executive Summary

The exploration phase is mainly characterized by survey measurements of large areas, usually but not always conducted from the sea surface. Performance can often be improved by use of deep towed systems, submersible robots, or seafloor sensors.

The principal technique employed during the exploration phase is the seismic survey [Prior and Hooper, 1999; Pecher and Holbrook, 2000]. Seismic acquisition techniques are by now well advanced. Deep towed high frequency arrays [Chapman *et al.*, 2002; NRLSSC, 2003] and ocean bottom seismographs (OBS) [Tinivella and Accaino, 2000; Bunz *et al.*, 2002] are in use, at least in limited areas. Miles [2000] provides an overview of practical systems. Optimization of acquisition should recognize that hydrate deposits are limited by the geothermal gradient to depths no greater than a few hundred meters below the sea floor.

Unconsolidated seafloor sediments are very weak in shear, and hydrate has a significant strengthening effect. Thus it is to be expected that shear wave imaging will be of more importance in hydrate reservoirs than in conventional seismic targets. In fact, acquisition and processing of shear wave data has been shown to be useful [Pecher and Holbrook, 2000; Mallick *et al.*, 2000]. There is broad agreement that four-component (4C) OBS acquisition is highly desirable for characterization of hydrate deposits.

Whereas seismic surveys explore the subsurface, high frequency acoustic multibeam systems focus on the seafloor [Orange *et al.*, 1999; Miles, 2000]. The primary output is a

bathymetric map. With some additional effort in acquisition and processing, the amplitude of backscattered waves can be analyzed for mechanical properties. Spatial resolution on the order of 10 m is feasible from the sea surface; efficient acquisition permits 100% coverage over hundreds of square kilometers per day. Deep towed systems have also been employed [Paull *et al.*, 1995].

Although multibeam mapping is not sensitive to the typical subseafloor hydrate deposit, it can be valuable for locating hydrates that breach the seafloor, and for secondary evidence of deeper hydrates. It is sensitive to free gas in the water column, and to seafloor expressions of gas venting, such as pockmarks, authigenic carbonate masses, and hard-shell chemosynthetic communities [Paull *et al.*, 1995].

Electrical methods are frequently neglected in conventional hydrocarbon exploration campaigns. This is unsurprising because targets can be thousands of meters below the surface and may only be tens of meters thick. However, when hydrate deposits are immediately beneath the seafloor, and hundreds of meters thick, deep tow electrical surveys may be useful to roughly estimate hydrate saturations prior to drilling.

Electrical methods are a good complement to seismic surveys. Whereas the seismic survey (often) accurately locates the bottom of a hydrate deposit, it is often not useful for estimating hydrate saturation. Even rough average saturation information provided by electrical methods can be a significant aid in evaluating drilling prospects. Moreover, electrical sounding provides an alternate method of locating hydrate in the absence of a BSR [Yuan and Edwards, 2000], which is particularly significant in the Gulf of Mexico where BSRs are frequently absent in hydrate bearing provinces.

Although forward and inverse modeling of the electrical survey is reasonably mature, acquisition techniques can be developed further. It would appear that even modest research investments could bring significant returns.

Gravity waves on the sea surface produce small changes of hydrostatic pressure on the seafloor. These pressure variations produce very small deformations of near-seafloor sediments, and the resulting elevation changes can be sensed by a seafloor gravity meter. The seafloor compliance (ratio of the deformation response to the pressure drive) is sensitive to elastic properties of sediments hundreds of meters below the seafloor [Willoughby and Edwards, 1997], i.e. coincident with the GHSZ. The success of this seemingly improbable method depends on exquisitely sensitive (but commercially available) field-deployable gravimeters, and averaging times of several hours per site [Willoughby and Edwards, 2000].

Methane in the water column, where the water depth is more than 500 meters, is a strong indicator that methane hydrate is either being accumulated [Roberts and Carney, 1997], depleted [Hutnak *et al.*, 1999; Sasaki *et al.*, 2002], and/or redistributed [Paull *et al.*, 1995] within the sediments below. Any of these situations can affect seafloor stability.

Hydrate deposits are frequently associated with dramatic methane seeps, particularly in the Gulf of Mexico [Roberts, 2001; Sassen *et al.*, 2001]. However, even subtle indications of methane can be significant. For example, the Blake Ridge deposit, which is situated on a quiescent passive margin, has a fault system extending from below the base of the bottom simulating reflector to almost the seafloor. These faults are believed to constitute efficient conduits for transport of methane [Rowe and Gettrust, 1994; Booth *et al.*, 1998]. Time-resolved measurements appear to be valuable in at least some locations [Hutnak *et al.*, 1999; Tryon *et al.*, 1999].

Logging while drilling (LWD) systems deliver real-time or near-real-time formation evaluation information. Combined with directional drilling technology [Cooper, 1994] LWD is used to actively steer the drill bit in response to geological or petrophysical properties of earth formations [Luthi, 2001]. Nearly a full suite of formation evaluation measurements is available in the LWD format. These include electrical resistivity, electrical imaging of the borehole, γ - γ density, neutron porosity, sonic wave speed, vertical seismic profiles, and magnetic resonance [Bargach *et al.*, 2000].

Borehole imaging and sonic wave speed measurements are most important for geohazard assessment. Electrical conductivity images of the borehole wall are useful for identifying hydrate lenses and nodules. They are also excellent indicators of open fractures, which are sensitive to the local stress state [Bratton *et al.*, 1999; Rezmer-Cooper *et al.*, 2001; Zoback *et al.*, 2003]. The status of sonic logging while drilling is somewhat less satisfactory. Although compressional wave speed can be determined, measurement of shear speed in very porous, unconsolidated, shallow marine sediments is problematical. Quadruple sonic tools [Freitag, 2003] hold out some promise for the future.

Wireline logging is an established technique for evaluating hydrate formations [Collett, 1998a; Collett, 1998b; Dallimore *et al.*, 1999; Collett, 2001; Akihisa *et al.*, 2002]. Deep water exploration wells are generally drilled without a riser, putting a premium on tools that can be deployed through drillpipe [Stoller *et al.*, 1997].

The stiffening effect of hydrate in sediment causes acoustic velocities to increase, but quantitative evaluation depends on the pore scale morphology of hydrate [Collett, 1998a], which is, in general, unknown. With respect to nuclear tools, hydrate has properties nearly identical to water, so those tools give reliable measurements of porosity irrespective of the presence of hydrate. Quantitative wireline interpretation of hydrate relies on combining magnetic resonance and density-porosity measurements [Takahashi *et al.*, 2001; Kleinberg *et al.*, 2004]. Resistivity measurements are less sensitive to borehole rugosity than density and magnetic resonance measurements, so electric logs are also useful in quantitative determinations.

The stability of a hydrate-affected formation is controlled by its temperature and pressure. Thus these quantities, which have only ancillary roles in conventional oil and gas reservoir characterization, are of prime importance for monitoring hydrate deposits.

Permanently implanted temperature sensors behind casing are becoming increasingly widespread [Brown *et al.*, 2000; Tolan *et al.*, 2001; Brown and Hartog, 2002]. Fiber optic distributed temperature sensing (DTS) systems are capable of continuously measuring temperature profiles with 0.3°C accuracy and 0.1°C precision at a spatial resolution of 1 m [Carnahan 1999]. Experimental fiber Bragg grating systems are about an order of magnitude more precise. It should be noted that there are considerable technical challenges in deploying subsea fiber optic sensor systems [Eriksson, 2002].

The industry is familiar with a number of reservoir-related hazards, such as subsidence and shallow water flows. However, it has little experience with seafloor and reservoir-scale hazards related to the presence of hydrates. Hydrate-related seafloor slide scars have been studied on the United States Atlantic margin [Booth *et al.*, 1994] and off Norway [Bugge *et al.*, 1988], but these features are typically thousands of years old, and knowledge of their precursors and causative factors, while advancing [see e.g. Paull *et al.*, 2000], is still limited.

Sensor suites should be chosen based on the most probable causes and precursors of seafloor slope failure. Hydrate stiffens sediments in which it exists, and this semi-consolidated mass overlies the highly fluidized and frequently gassy silts and muds at the base of the gas hydrate stability zone (GHSZ).

Since free gas has several important roles in the destabilization of hydrate deposits, seismic measurements are likely to be the most valuable inputs to a seafloor monitoring program. Time-lapse measurements will benefit from four component (three-component geophone plus hydrophone) receivers permanently installed in fixed positions at the seafloor. Such systems have been deployed in many areas. Entralgo and Spitz [2001] warn against operations in depths greater than 1,000 ft (300 m), but others have been successful in much deeper water [Tinivella and Accaino, 2000; Bunz *et al.*, 2002]. Since seismic sources are energy intensive, they may visit the site only occasionally, possibly in a deep-tow system [Chapman *et al.*, 2002; NRLSSC, 2003].

The seafloor monitoring project most closely related to the hydrate hazard problem is being planned at the Center for Marine Resources and Environmental Technology at the University of Mississippi [McGee and Woolsey, 1999; McGee and Woolsey, 2000]. Sensors will be placed in an area of active gas vents in the Gulf of Mexico. The emphasis of this project is on passive seismic detection of earth motions, as well as time-lapse monitoring of seismic reflectors in the subsurface. Acoustic sensors will monitor the temperature of the water column, important for understanding the behavior of hydrate outcrops. Electromagnetic and chemical sensors are also being considered.

3.0 Results and Discussion

3.1 Exploration Phase

The exploration phase is mainly characterized by survey measurements of large areas, usually but not always conducted from the sea surface. Performance can often be improved by use of deep towed systems, submersible robots, or seafloor sensors. Several short review articles on the exploration geophysics of hydrate deposits are available [Max and Miles, 1999; Spence *et al.*, 2000].

3.1.1 Seismic Surveys

The principal technique employed during the exploration phase is the seismic survey [Prior and Hooper, 1999; Pecher and Holbrook, 2000]. Seismic acquisition techniques are by now well advanced. Deep towed high frequency arrays [Chapman *et al.*, 2002; NRLSSC, 2003] and ocean bottom seismographs (OBS) [Tinivella and Accaino, 2000; Bunz *et al.*, 2002] are in use, at least in limited areas. Miles [2000] provides an overview of practical systems. Optimization of acquisition should recognize that hydrate deposits are limited by the geothermal gradient to depths no greater than a few hundred meters below the sea floor.

Typically, the appearance of the bottom simulating reflector (BSR) is taken as a marker of hydrate presence. However, it is becoming increasingly evident that the absence of a BSR does not imply the absence of hydrate. A "control" borehole, drilled into Blake Ridge offshore South Carolina, encountered hydrate even though no BSR was present on the seismic section in that area [Collett and Ladd, 2000]. This and similar observations have been explained by an upward flux of methane inadequate to maintain a stratum of free gas beneath the gas hydrate stability zone (GHSZ) [Ruppel and Kinoshita, 2000]. Thus, more sophisticated screening techniques for hydrate occurrence are needed.

Characterization of hydrate reservoirs is critical to predicting their behavior. Hydrate saturations tend to be very heterogeneous. Both intergranular and massive morphologies are common, with prominent nodules and lenses present in many formations. Unfortunately, seismic measurements are poor indicators of either average or patchy hydrate saturation in the GHSZ. Above the BSR the seismic signal sometimes appears blanked. It has been theorized that the blanking itself provides some information about hydrate distribution [Lee and Dillon, 2001], but this is controversial [Pecher and Holbrook, 2000].

Even when good travel-time inversions can be performed, hydrate saturation determination is elusive. Sediment porosity is a major influence on acoustic wave speeds. The presence of free gas [Ecker *et al.*, 2000; Tinivella and Carcione, 2001] or permafrost can be major complications. The intergranular growth habit of hydrate is an important influence [Dvorkin *et al.*, 2000; Helgerud, 2001] that remains controversial and may depend on the mechanism of accumulation [Kleinberg *et al.*, 2003b]. However, as

knowledge of hydrate microstructure increases, estimates of hydrate distributions can be refined [Dai *et al.*, 2004].

Unconsolidated seafloor sediments are very weak in shear, and hydrate has a significant strengthening effect. Thus it is to be expected that shear wave imaging will be of more importance in hydrate reservoirs than in conventional seismic targets. In fact, acquisition and processing of shear wave data has been shown to be useful [Pecher and Holbrook, 2000, Mallick *et al.*, 2000]. There is broad agreement that four-component (4C) OBS acquisition is highly desirable for characterization of hydrate deposits.

3.1.2 Multibeam Mapping

Whereas seismic surveys explore the subsurface, high frequency acoustic multibeam systems focus on the seafloor [Orange *et al.*, 1999; Miles, 2000]. The primary output is a bathymetric map. With some additional effort in acquisition and processing, the amplitude of backscattered waves can be analyzed for mechanical properties. Spatial resolution on the order of 10 m is feasible from the sea surface; efficient acquisition permits 100% coverage over hundreds of square kilometers per day. Deep towed systems have also been employed [Paull *et al.*, 1995].

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3.1.3 Electrical Surveys

Electrical methods are frequently neglected in conventional hydrocarbon exploration campaigns. This is unsurprising because targets can be thousands of meters below the surface and may only be tens of meters thick. However, when hydrate deposits are immediately beneath the seafloor, and hundreds of meters thick, deep tow electrical surveys may be useful to roughly estimate hydrate saturations prior to drilling.

Although the electrical survey on land is a classical technique of geophysics [Allaud and Martin, 1977], one may wonder whether it is applicable to the seafloor, where the upper half space is conductive seawater. However, the theory of electrical sounding of marine gas hydrate formations is well developed [Edwards, 1997]. Low frequency (Hz), long baseline measurements are capable of giving not only half-space resistivities of the seabed, but also crude depth profiles. The method has been implemented with apparent success on the Cascadia margin [Yuan and Edwards, 2000].

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3.1.4 Seafloor Compliance

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3.1.5 Methane Sensors

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A variety of methods are used to detect and investigate marine seeps. Free gas is efficiently detected by sonar [Merewether et al., 1985], though this method is not quantitative or specific. Chromatography [Cynar and Yayanos, 1992] gives more precise information on water samples obtained from bottles, sniffers, or cores. Electrochemical sensors [Whitfield and Jagner, 1981; Bussell et al., 1999] are also species-specific and have the advantage of being easily deployable on submersible systems. Interpretation of

the data both depends on and contributes to understanding the geological setting [Schumacher and Abrams, 1994].

3.2 Evaluation Phase

The evaluation of gas hydrate deposits occurs during and immediately after the drilling of boreholes. Typically, the earliest information comes from logging while drilling tools. After the well is drilled to its total depth, wireline tools are used to acquire more complete information. Some wells may be cored, with the recovered sediment or rock cores carefully examined and subsampled in the laboratory. These methods have been compared by Hutchinson [1991].

3.2.1 Logging While Drilling

Logging while drilling (LWD) systems deliver real-time or near-real-time formation evaluation information. Combined with directional drilling technology [Cooper, 1994] LWD is used to actively steer the drill bit in response to geological or petrophysical properties of earth formations [Luthi, 2001]. Nearly a full suite of formation evaluation measurements is available in the LWD format. These include electrical resistivity, electrical imaging of the borehole, γ - γ density, neutron porosity, sonic wave speed, vertical seismic profiles, and magnetic resonance [Bargach *et al.*, 2000].

Borehole imaging and sonic wave speed measurements are most important for geohazard assessment. Electrical conductivity images of the borehole wall are useful for identifying hydrate lenses and nodules. They are also excellent indicators of open fractures, which are sensitive to the local stress state [Bratton *et al.*, 1999; Rezmer-Cooper *et al.*, 2001; Zoback *et al.*, 2003]. The status of sonic logging while drilling is somewhat less satisfactory. Although compressional wave speed can be determined, measurement of shear speed in very porous, unconsolidated, shallow marine sediments is problematical. Quadruple sonic tools [Freitag, 2003] hold out some promise for the future.

A primary purpose of LWD is to make measurements where wireline logging and coring are impractical. Hydrate shares the near-seafloor environment with unconsolidated sand and mud. In the Gulf of Mexico the upper part of the well is normally cased as quickly as possible to reduce the risk of collapse. Sometimes the only opportunity to characterize hydrate-bearing sediment is to use LWD tools. Thus the technique of logging-while-coring [Goldberg *et al.*, 2003], introduced on Ocean Drilling Program Leg 204, is a significant advance, enabling the recovery of core while acquiring borehole data.

A limitation of LWD tools is that they are designed to operate in a limited range of hole sizes [Bargach *et al.*, 2000]:

Tool Collar Size (in.)	Maximum Bit Size (in.)
4.75	6.25
6.75	9.875
8.25	12.25

Unlike wireline tools, some of which are useful in very large boreholes because they are run eccentric, few LWD tools are capable of operating in boreholes very much larger than the tool collar size. Since typical deepwater Gulf of Mexico drilling operations prescribe several hundred feet of 36 inch borehole followed by several thousand feet of 20 inch borehole [Jenkins *et al.*, 1999], it is unlikely that a full suite of LWD logs can be obtained in the upper sections of conventional exploration and production wells. Therefore, characterization of hydrate reservoirs using LWD tools require drilling special wells into the hydrate reservoir.

3.2.2 Wireline Logging

Wireline logging is an established technique for evaluating hydrate formations [Collett, 1998a; Collett, 1998b; Dallimore *et al.*, 1999; Collett, 2001; Akihisa *et al.*, 2002]. Deep water exploration wells are generally drilled without a riser, putting a premium on tools that can be deployed through drillpipe [Stoller *et al.*, 1997].

The stiffening effect of hydrate in sediment causes acoustic velocities to increase, but quantitative evaluation depends on the pore scale morphology of hydrate [Collett, 1998a], which is, in general, unknown. With respect to nuclear tools, hydrate has properties nearly identical to water, so those tools give reliable measurements of porosity irrespective of the presence of hydrate. Quantitative wireline interpretation of hydrate relies on combining magnetic resonance and density-porosity measurements [Takahashi *et al.*, 2001; Kleinberg *et al.*, 2004]. Resistivity measurements are less sensitive to borehole rugosity than density and magnetic resonance measurements, so electric logs are also useful in quantitative determinations.

When water, free gas, and hydrate are present simultaneously in the pore space, neutron porosity [Ellis, 1987] is a useful measurement. The neutron log responds to all protons, whether in water or hydrate; the proton density of hydrate is only 6% higher than that of water. Thus the neutron-density combination can be used to estimate free gas content, irrespective of the presence of hydrate. The neutron log responds anomalously to clay, and this must be corrected by other tools.

The problem of evaluation of hydrate in the presence of permafrost is a significant problem in well log interpretation. Their stability ranges are different, so temperature and pressure measurements can remove the ambiguity in some situations. However there are depth intervals in which both are stable. Most physical properties of ice and methane hydrate are very similar [Dvorkin *et al.*, 2000]. Permafrost and hydrate responses to density, neutron, resistivity, sonic and NMR logging tools are indistinguishable for all practical purposes. The small differences in the relevant bulk properties are masked by modest uncertainties in porosity and saturation. Dielectric constant measurements may be useful in distinguishing hydrate from permafrost, but dielectric logging tools are not broadly deployed. Ice has a thermal conductivity almost four times larger than water or gas hydrate. Thus, once the porosity and water saturation of a formation are determined,

thermal conductivity measurements may prove useful in distinguishing ice from hydrate in formations in which they can coexist.

Wireline logging provides more and better tools for geohazard assessment than LWD does. Shear wave logging in soft, slow formations is superior, and electrical borehole imaging tools have higher resolution. Fluid sampling tools can be used to conduct mini-fracture tests [Bell, 2003; Zoback *et al.*, 2003].

3.3 Monitoring Phase

Exploration phase measurements indicate the presence of gas hydrate deposits, and evaluation phase measurements are essential for defining the exact quantities and locations of hydrate within those deposits. These measurements warn of the possibility of drilling hazards, and provide inputs to wellbore stability and seafloor stability models that help predict whether the deposits will become hazards during production. However the predictive power of these stability models is still untested. Moreover, even after more experience has been acquired with them, the models are unlikely to provide totally reliable predictions of hazard events.

Offshore platforms, wellheads and pipelines are very costly assets, and the environmental risks associated with disruption of hydrocarbon production in deep water are considerable. Therefore prudent engineering practice dictates that potential hazard conditions be monitored, possibly over the lifetime of the reservoir. Sensor systems can be installed in well bores and on the seafloor. To gain broadest acceptance, these systems should be continuous, permanent, inexpensive, and in some cases widely dispersed.

3.3.1 Permanent Monitoring of the Well Bore

Hydrate-associated well bore hazards have been documented by Collett and Dallimore [2002]. As warm fluid is produced from underlying conventional reservoirs, near-wellbore hydrates can decompose. This results in flows of gas and water outside of casing due to loss of hydraulic isolation, and in extreme cases damage to the casing itself.

The stability of a hydrate-affected formation is controlled by its temperature and pressure. Thus these quantities, which have only ancillary roles in conventional oil and gas reservoir characterization, are of prime importance for monitoring hydrate deposits.

Permanently implanted temperature sensors behind casing are becoming increasingly widespread [Brown *et al.*, 2000; Tolan *et al.*, 2001; Brown and Hartog, 2002]. Fiber optic distributed temperature sensing (DTS) systems are capable of continuously measuring temperature profiles with 0.3°C accuracy and 0.1°C precision at a spatial resolution of 1 m [Carnahan, 1999]. Experimental fiber Bragg grating systems are about an order of magnitude more precise. It should be noted that there are considerable technical challenges in deploying subsea fiber optic sensor systems [Eriksson, 2002].

Permanently installed pressure gauges are also now relatively routine; quartz gauges are both very precise and very stable over time [Tibold *et al.*, 2000]. Combined temperature and pressure sensor systems based on fiber optics are now at the leading edge of technology [Kragas *et al.*, 2001; Schroeder *et al.*, 2002]. Although the fiber optic pressure measurement is not as precise as that available from the quartz gauge, this technology has the potential for providing low cost multisensor arrays. Perhaps the most sophisticated borehole permanent monitoring system installed to date, including temperature, pressure, and resistivity arrays, has been used in a reservoir control system [Bryant *et al.*, 2002].

3.3.2 Permanent Monitoring of Reservoir

3.3.2.1 Nature of the Hazards

The industry is familiar with a number of reservoir-related hazards, such as subsidence and shallow water flows. However, it has little experience with seafloor and reservoir-scale hazards related to the presence of hydrates. Hydrate-related seafloor slide scars have been studied on the United States Atlantic margin [Booth *et al.*, 1994] and off Norway [Bugge *et al.*, 1988], but these features are typically thousands of years old, and knowledge of their precursors and causative factors, while advancing [see e.g. Paull *et al.*, 2000], is still limited.

Sensor suites should be chosen based on the most probable causes and precursors of seafloor slope failure. Hydrate stiffens sediments in which it exists, and this semi-consolidated mass overlies the highly fluidized and frequently gassy silts and muds at the base of the gas hydrate stability zone (GHSZ). The setting is reminiscent of avalanche conditions on a snow-covered mountain. Thus, as a general principle, either creation or dissociation of hydrate can cause problems, though dissociation is the more immediately worrisome condition because it can lead to a loss of shear strength and an increase in pore pressure.

An increase in the amount of free gas within or below the GHSZ is perhaps the most important indicator of potential problems. Although gas is not thermodynamically stable in the GHSZ, it has been observed to exist there, presumably due to the kinetic limitations of hydrate formation [Sloan, 1998]. Indeed, hydrates are frequently associated with natural gas vents and seeps in the Gulf of Mexico [Roberts, 2001; Sassen *et al.*, 2001] and elsewhere.

The presence of free gas in the GHSZ leads to two different problems with respect to hydrate thermodynamic stability. First, it reduces the lithostatic pressure by decreasing the density of the sediment column. Secondly, and probably more importantly, it steepens the geothermal gradient due to its small thermal conductivity. Because the seafloor constitutes a nearby heat reservoir of essentially infinite capacity, steepening the gradient causes the temperature at the base of the GHSZ to increase. Both pressure and temperature effects tend to destabilize the hydrate deposit. These problems can be exacerbated by a pore-filling growth habit [Kleinberg *et al.*, 2003b].

Accumulation of gas under the base of the GHSZ, due to destabilization of hydrate there or migration from below, is also clearly problematic. Increasing gas saturation reduces the shear strength of the formation.

There is limited information about the times scales over which gas moves in the subsurface [Hutnak *et al.*, 1999; Tryon *et al.*, 1999]. While it has been observed that gas hydrate outcrops change in size and shape in less than a year, and possibly on time scales of days to weeks [McDonald *et al.*, 1994; Roberts, 1999], this appears to be correlated to ocean temperatures, not fluctuations of gas supply. However it has recently been found that the spatial patterns of upward gas migration are much more complicated than had been thought [Wood *et al.*, 2002], and this insight may lead to a more detailed understanding of the variability of gas flows over time.

Several other hazard conditions should be monitored. Subsidence or steepening of the seabed are clearly hazard warnings. Submarine landslides have been observed to occur on surprisingly gentle slopes: 1°-7° is typical [Hampton, 1996], though a triggering event such as an earthquake may be required. Sediment warming around pipelines and other production facilities is another potential hazard indicator.

3.3.2.2 Sensor Systems and Methods

Since free gas has several important roles in the destabilization of hydrate deposits, seismic measurements are likely to be the most valuable inputs to a seafloor monitoring program. Time-lapse measurements will benefit from four component (three-component geophone plus hydrophone) receivers permanently installed in fixed positions at the seafloor. Such systems have been deployed in many areas. *Entralgo and Spitz* [2001] warn against operations in depths greater than 1,000 ft (300 m), but others have been successful in much deeper water [Tinivella and Accaino, 2000; Bunz *et al.*, 2002]. Since seismic sources are energy intensive, they may visit the site only occasionally, possibly in a deep-tow system [Chapman *et al.*, 2002; NRLSSC, 2003].

Heat flow sensors are indicators of thermal changes in the subsurface [Ruppel, 2000; Grevemeyer and Villinger, 2001]. It should be noted that hydrate and water have almost equal thermal conductivities, with free gas providing the only significant contrast.

Chemical and/or bubble sensors may find a place in the seafloor sensor suite. Dissolved or free natural gas in the water column in deep water suggests that hydrate is either being formed or being dissociated; both conditions are of concern to the reservoir manager.

Tiltmeters and seismographs can passively monitor subtle or sudden seafloor motions. Seafloor pressure sensors can be used to measure subsidence [Mes, 1988].

To use sensor systems to best advantage, they should be in place before wells are drilled. The drilling operation is a perturbation of the reservoir, and time lapse measurements during and immediately after drilling may prove valuable.

3.3.2.3 Survey of Seafloor Observatories

The seafloor monitoring project most closely related to the hydrate hazard problem is being planned at the Center for Marine Resources and Environmental Technology at the University of Mississippi [McGee and Woolsey, 1999; McGee and Woolsey, 2000]. Sensors will be placed in an area of active gas vents in the Gulf of Mexico. The emphasis of this project is on passive seismic detection of earth motions, as well as time-lapse monitoring of seismic reflectors in the subsurface. Acoustic sensors will monitor the temperature of the water column, important for understanding the behavior of hydrate outcrops. Electromagnetic and chemical sensors are also being considered.

The Monterey Bay Aquarium Research Institute (MBARI) is undertaking a major effort to instrument Monterey Bay, California [MBARI, 2003]. In 2002 MBARI received a \$7 million National Science Foundation grant to start work on the Monterey Accelerated Research System (MARS). The initial objective is to install 62 km of submarine cable with a power capacity of 10 kW and a data capacity of 100 megabits per second. This cable will service one node with four docking ports. A variety of sensors will be designed to interchangeably dock at the node. The system is connected to a shore station via copper and fiber optic submarine cable.

The MARS project is a steppingstone toward the much more ambitious NEPTUNE system, planned for the Juan de Fuca Plate, offshore British Columbia, Washington, and Oregon [Howe, *et al.*, 2001; Neptune, 2003]. This plan envisions 30 instrumented nodes on a 3,000 km cable backbone. The system is sized to carry 100 kW of power and 10 gigabits per second of telemetry. The design lifetime is 30 years. Each node will accept a variety of interchangeable sensors, which can include borehole sensor strings.

VENUS (Victoria Experimental Network Under Sea) [Tunnicliffe and Dewey, 2001], managed by the University of Victoria, British Columbia, is a spin-off of NEPTUNE. The VENUS consortium will build and operate sensors compatible with the NEPTUNE backbone. The existence of nascent user communities such as VENUS is an indication of the flexibility that the seafloor observatory architects aim to build into their systems.

Of particular interest to the hydrate community is the proposal to use the NEPTUNE system to monitor seawater chemistry and ground motions at Hydrate Ridge in the Cascadia subduction zone offshore Oregon [Suess *et al.*, 1999]. The proposed sensors are:

- Photography at active vent and bubble sites.
- Conductivity, temperature and depth (CTD) sensor package.
- Chemical sensors:
methane, hydrogen sulfide, oxygen, iron, pH and pCO₂.
- Doppler current records.
- High-precision hydrostatic pressure readings.
- Three-component seismometers.
- Geodetic observations; tilt meters.

The full-featured NEPTUNE system is a reasonably costly proposition [Neptune, 2000]. It is estimated that the design and planning tasks will cost roughly \$13 million, and constructing and installing the infrastructure (including cable and shore stations) will cost close to \$90 million. User-group sensor construction and installation is estimated to cost in the range of \$100 million. Operating expenses will approximate \$10 million per year. It can be anticipated that a system narrowly focused on hydrate hazard monitoring, in a defined geographical area, using mass produced sensors, would cost considerably less.

4.0 Experimental

No experimental work was conducted for this report.

5.0 Conclusions

Measurements will be needed during exploration, evaluation, and monitoring phases. The requirements during these phases are distinct from one another, and can be treated separately.

Characterization of hydrate reservoirs is critical to predicting their behavior. Hydrate saturations tend to be very heterogeneous. Both intergranular and massive morphologies are common, with prominent nodules and lenses present in many formations. Unfortunately, seismic measurements are poor indicators of either average or patchy hydrate saturation in the GHSZ.

Unconsolidated seafloor sediments are very weak in shear, and hydrate has a significant strengthening effect. Thus it is to be expected that shear wave imaging will be of more importance in hydrate reservoirs than in conventional seismic targets.

Although multibeam mapping is not sensitive to the typical subseafloor hydrate deposit, it can be valuable for locating hydrates that breach the seafloor, and for secondary evidence of deeper hydrates. It is sensitive to free gas in the water column, and to seafloor expressions of gas venting, such as pockmarks, authigenic carbonate masses, and hard-shell chemosynthetic communities [Paull *et al.*, 1995].

Electrical methods are a good complement to seismic surveys. Whereas the seismic survey (often) accurately locates the bottom of a hydrate deposit, it is often not useful for estimating hydrate saturation. Even rough average saturation information provided by electrical methods can be a significant aid in evaluating drilling prospects. Moreover, electrical sounding provides an alternate method of locating hydrate in the absence of a BSR [Yuan and Edwards, 2000], which is particularly significant in the Gulf of Mexico where BSRs are frequently absent in hydrate bearing provinces.

Although forward and inverse modeling of the electrical survey is reasonably mature, acquisition techniques can be developed further. It would appear that even modest research investments could bring significant returns.

The seafloor compliance (ratio of the deformation response to the pressure drive) is sensitive to elastic properties of sediments hundreds of meters below the seafloor [Willoughby and Edwards, 1997], i.e. coincident with the GHSZ. The success of this seemingly improbable method depends on exquisitely sensitive (but commercially available) field-deployable gravimeters, and averaging times of several hours per site [Willoughby and Edwards, 2000].

A variety of methods is used to detect and investigate marine seeps and can be used to infer hydrates. Free gas is efficiently detected by sonar [Merewether *et al.*, 1985], though this method is not quantitative or specific. Chromatography [Cynar and Yayanos, 1992] gives more precise information on water samples obtained from bottles, sniffers, or cores. Electrochemical sensors [Whitfield and Jagner, 1981; Bussell *et al.*, 1999] are also species-specific and have the advantage of being easily deployable on submersible systems. Interpretation of the data both depends on and contributes to understanding the geological setting [Schumacher and Abrams, 1994].

Logging while drilling (LWD) systems deliver real-time or near-real-time formation evaluation information. Combined with directional drilling technology [Cooper, 1994], LWD is used to actively steer the drill bit in response to geological or petrophysical properties of earth formations [Luthi, 2001]. Nearly a full suite of formation evaluation measurements is available in the LWD format. These include electrical resistivity, electrical imaging of the borehole, γ - γ density, neutron porosity, sonic wave speed, vertical seismic profiles, and magnetic resonance [Bargach *et al.*, 2000].

Borehole imaging and sonic wave speed measurements are most important for geohazard assessment. Electrical conductivity images of the borehole wall are useful for identifying hydrate lenses and nodules. They are also excellent indicators of open fractures, which are sensitive to the local stress state [Bratton *et al.*, 1999; Rezmer-Cooper *et al.*, 2001; Zoback *et al.*, 2003]. The status of sonic logging while drilling is somewhat less satisfactory. Although compressional wave speed can be determined, measurement of shear speed in very porous, unconsolidated, shallow marine sediments is problematical. Quadruple sonic tools [Freitag, 2003] hold out some promise for the future.

Wireline logging provides more and better tools for geohazard assessment than LWD does. Shear wave logging in soft, slow formations is superior, and electrical borehole imaging tools have higher resolution. Fluid sampling tools can be used to conduct mini-fracture tests [Bell, 2003; Zoback *et al.*, 2003]. However, the use of wireline tools is limited because of well bore stability concerns in shallow sediments.

The stability of a hydrate-affected formation is controlled by its temperature and pressure. Thus these quantities, which have only ancillary roles in conventional oil and gas reservoir characterization, are of prime importance for monitoring hydrate deposits.

Permanently implanted temperature sensors behind casing are becoming increasingly widespread [Brown *et al.*, 2000; Tolan *et al.*, 2001; Brown and Hartog, 2002]. Fiber optic distributed temperature sensing (DTS) systems are capable of continuously measuring temperature profiles with 0.3°C accuracy and 0.1°C precision at a spatial resolution of 1 m [Carnahan 1999]. Experimental fiber Bragg grating systems are about an order of magnitude more precise. It should be noted that there are considerable technical challenges in deploying subsea fiber optic sensor systems [Eriksson, 2002].

Permanently installed pressure gauges are also now relatively routine; quartz gauges are both very precise and very stable over time [Tibold *et al.*, 2000]. Combined temperature and pressure sensor systems based on fiber optics are now at the leading edge of technology [Kragas *et al.*, 2001; Schroeder *et al.*, 2002]. Although the fiber optic pressure measurement is not as precise as that available from the quartz gauge, this technology has the potential for providing low cost multisensor arrays. Perhaps the most sophisticated borehole permanent monitoring system installed to date, including temperature, pressure, and resistivity arrays, has been used in a reservoir control system [Bryant *et al.*, 2002].

The industry is familiar with a number of reservoir-related hazards, such as subsidence and shallow water flows. However, it has little experience with seafloor and reservoir-scale hazards related to the presence of hydrates. Hydrate-related seafloor slide scars have been studied on the United States Atlantic margin [Booth *et al.*, 1994] and off Norway [Bugge *et al.*, 1988], but these features are typically thousands of years old, and knowledge of their precursors and causative factors, while advancing [see e.g. Paull *et al.*, 2000], is still limited.

An increase in the amount of free gas within or below the GHSZ is perhaps the most important indicator of potential problems. Although gas is not thermodynamically stable in the GHSZ, it has been observed to exist there, presumably due to the kinetic limitations of hydrate formation [Sloan, 1998]. Indeed, hydrates are frequently associated with natural gas vents and seeps in the Gulf of Mexico [Roberts, 2001; Sassen *et al.*, 2001] and elsewhere.

The presence of free gas in the GHSZ leads to two different problems with respect to hydrate thermodynamic stability. First, it reduces the lithostatic pressure by decreasing the density of the sediment column. Secondly, and probably more importantly, it steepens the geothermal gradient due to its small thermal conductivity. Because the seafloor constitutes a nearby heat reservoir of essentially infinite capacity, steepening the gradient causes the temperature at the base of the GHSZ to increase. Both pressure and temperature effects tend to destabilize the hydrate deposit. These problems can be exacerbated by a pore-filling growth habit [Kleinberg *et al.*, 2003b].

Since free gas has several important roles in the destabilization of hydrate deposits, seismic measurements are likely to be the most valuable inputs to a seafloor monitoring program. Time-lapse measurements will benefit from four component (three-component

geophone plus hydrophone) receivers permanently installed in fixed positions at the seafloor.

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