

Using NWP to assess climate models

(Harnessing the power of data assimilation)

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3rd WGNE Workshop on Systematic Errors in Climate and NWP Models

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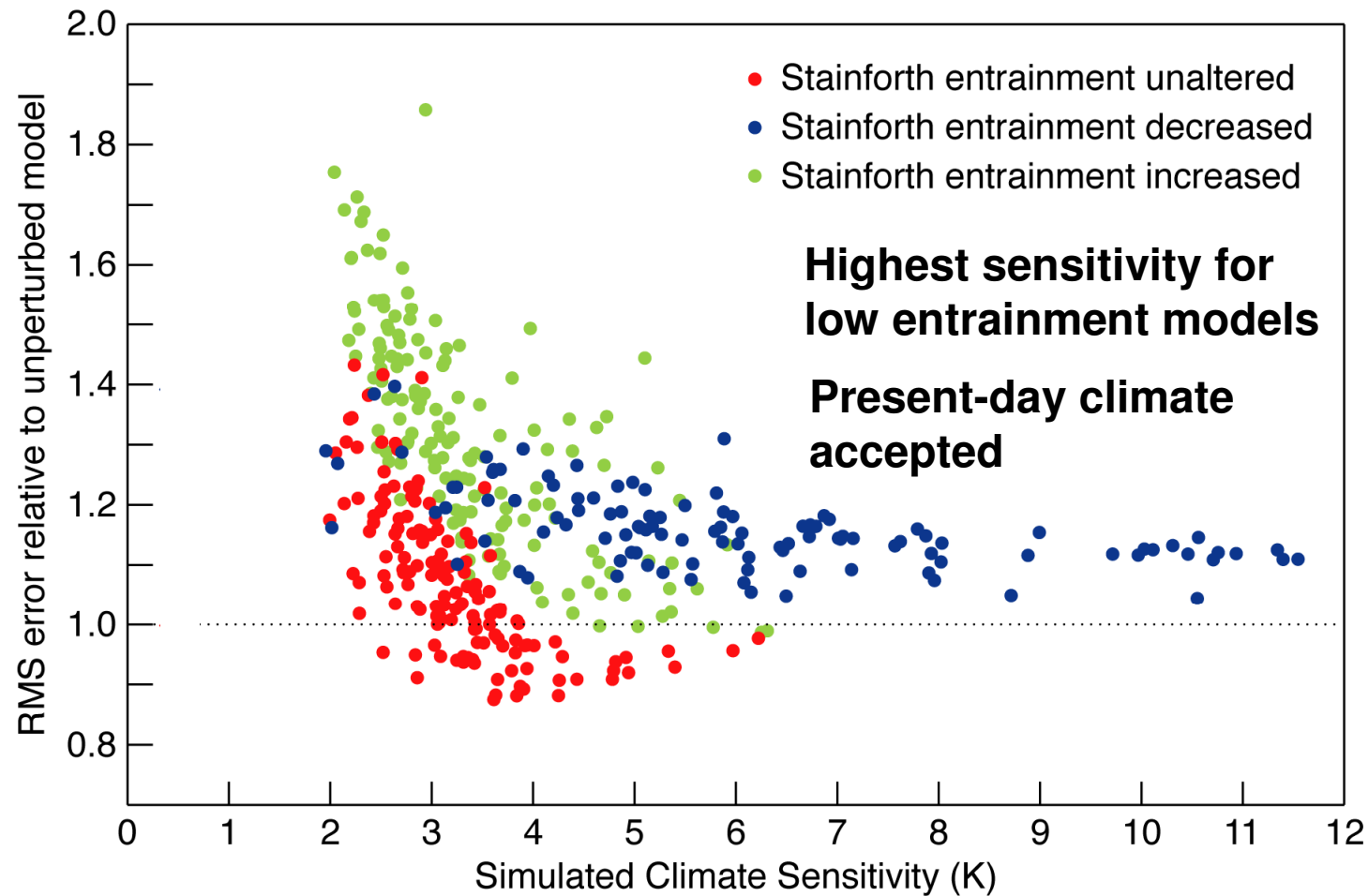
Parameters & Uncertainties (HadAM3)

Parameter	Physics	Low	Middle	High
Droplet to rain conversion rate (s ⁻¹)	Cloud	0.5x10 ⁻⁴	1.0x10 ⁻⁴	4.0x10 ⁻⁴
Relative humidity for cloud formation	Cloud	0.6	0.7	0.9
Cloud fraction at saturation (free trop.)	Cloud	0.5	0.7	0.8
Entrainment rate coefficient	Convection	0.6	3.0	9.0
Time-scale for destruction of CAPE (h)	Convection	1.0	2.0	4.0
Effective radius of ice particles (µm)	Radiation	25	30	40
Diffusion e-folding time (h)	Dynamics	6	12	24
Roughness length parameter (Charnock)	Boundary	0.012	0.016	0.02
Stomatal conductance dependent on CO ₂	Land	Off	-	On
Ocean-to-ice heat transfer (m ⁻² s ⁻¹)	Sea Ice	2.5x10 ⁻⁵	1.0x10 ⁻⁴	3.8x10 ⁻⁴

**Many uncertainties are associated with “fast physics”.
... which is also important in NWP**

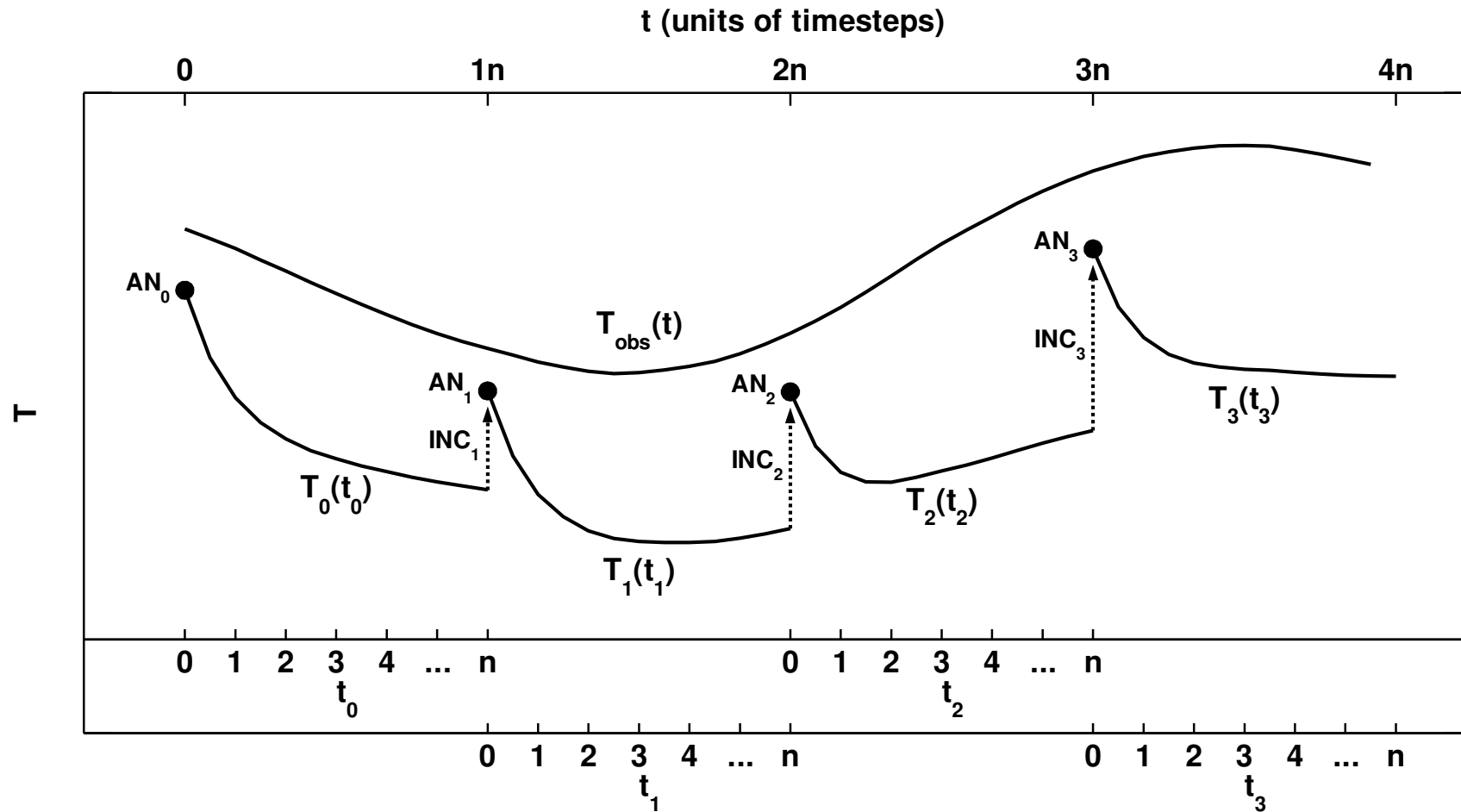
Representative selection of parameters and uncertainties used by Murphy et al., 2004: *Nature*, **430**, 768-772.

Climate: Error vs Sensitivity

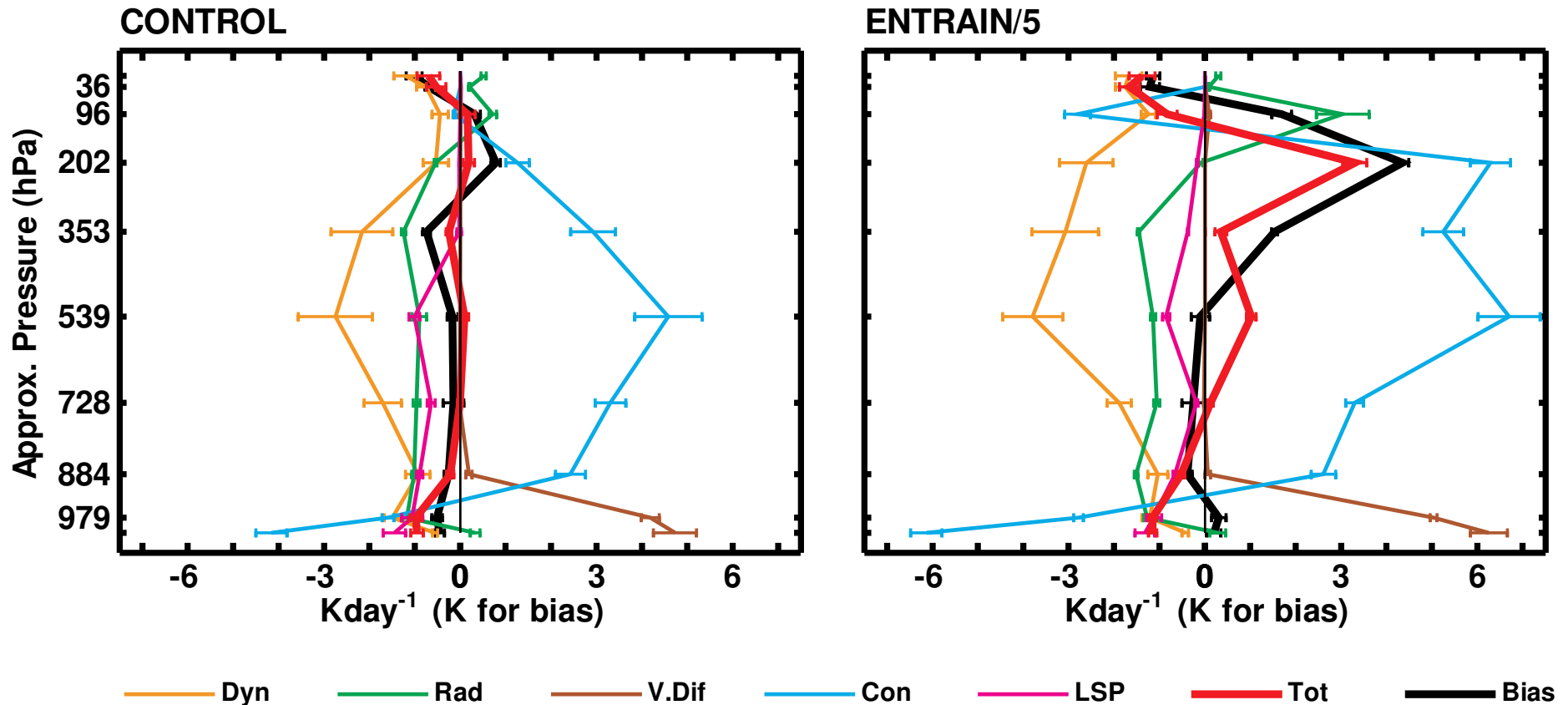


Combined RMSE of 8 year mean, annual mean T_{2m} , SLP, precipitation and ocean-atmosphere sensible+latent heat fluxes (equally weighted and normalised by the control). Stainforth et al., 2005, *Nature*, **433**, 403-406.

The data assimilation / forecast cycle



Amazon January 2005 Initial T Tendencies



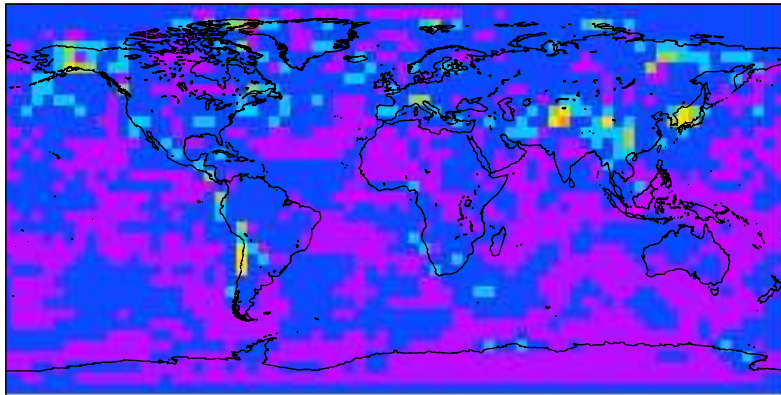
**ENTRAIN/5 out of balance:
reject or down-weight?**

**By D+5 balance is restored but it is complicated by
interactions between processes (non-linearity)**

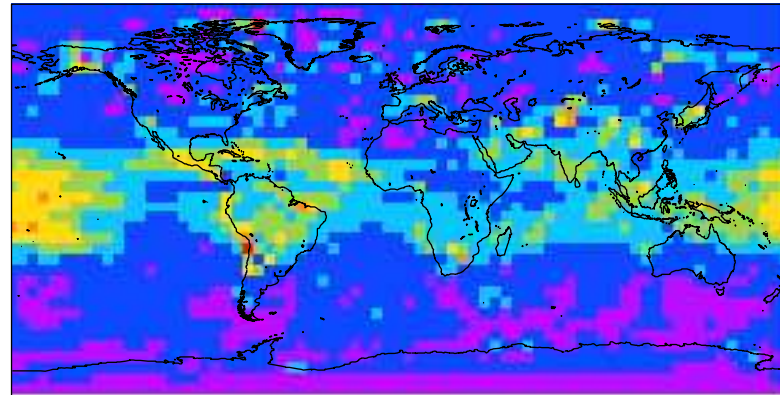
*Amazon = [300°E-320°E, 20°S-0°N]. Mean of 31 days X 4 forecasts per day X 12 timesteps per forecast.
70% confidence intervals are based on daily means. CONTROL model = 29R1,T159,L60,1800S.*

Vertically Integrated Absolute Tendencies

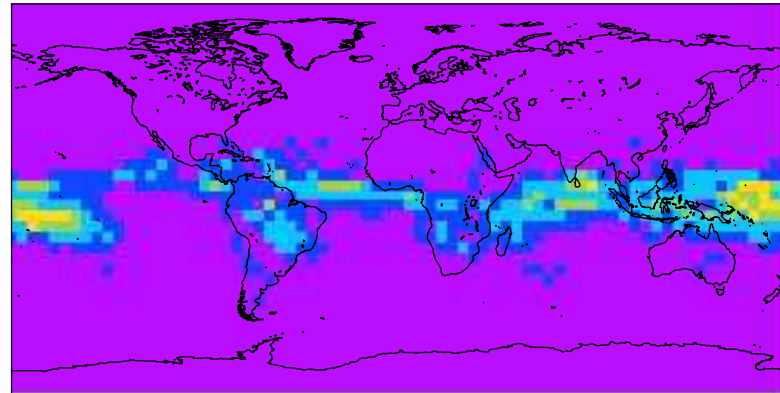
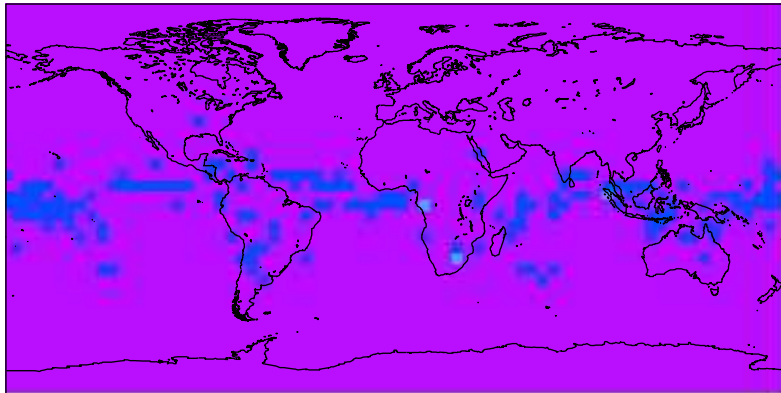
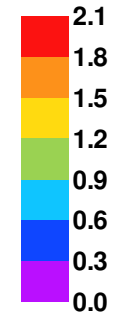
CONTROL



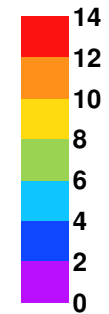
ENTRAIN/5



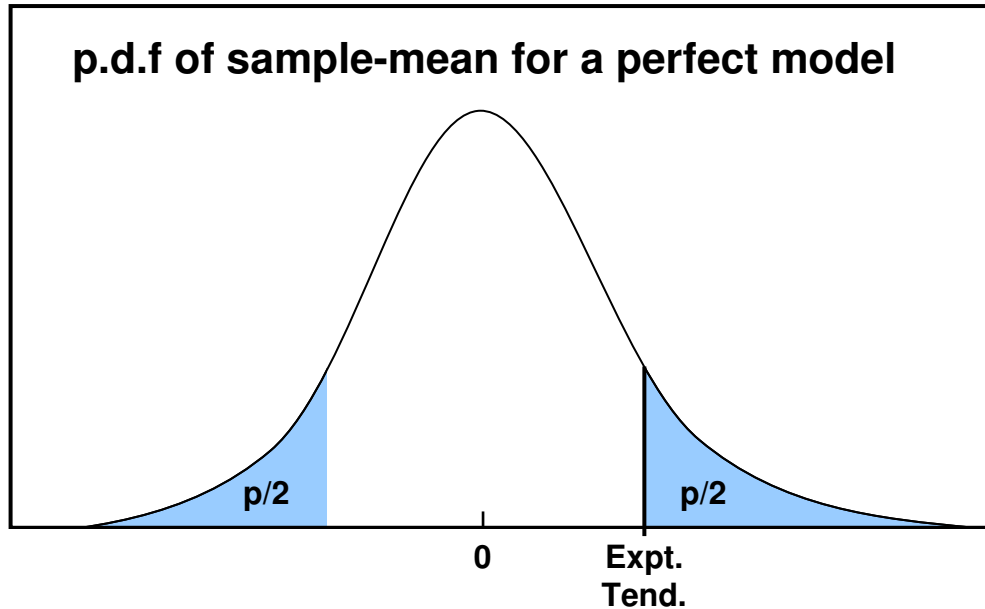
T (Kday⁻¹)



q (kgm⁻²day⁻¹)



How to weight each model



Calculate the probability that each model is “perfect”

$$p = p(\text{param}, x, y, z)$$

= probability that a zero population-mean tendency cannot be rejected

Possible methodology:

Average over parameters (T,q,u,v)

Vertically integrate

Integrate over tropics and extra-tropics

$$p_{\text{PERF}} \equiv p_{\text{TROP}} \times p_{\text{EX-TR}}$$

	Probability that model fast physics is perfect
CONTROL	0.20
ENTRAIN/5	0.12

Linearity and Cost

31 day Initial Tendencies \equiv 5 years CGCM

$$C'_{\rho_1, \rho_2, \dots, \rho_n} = C'_{\rho_1} + C'_{\rho_2} + \dots + C'_{\rho_n} \quad ?$$

Murphy et al. (2004): 23 “fast physics” parameters over 5 processes, 2 to 4 values

Linear: 24 models to assess
 Non-linear: 15,000,000,000 models to assess

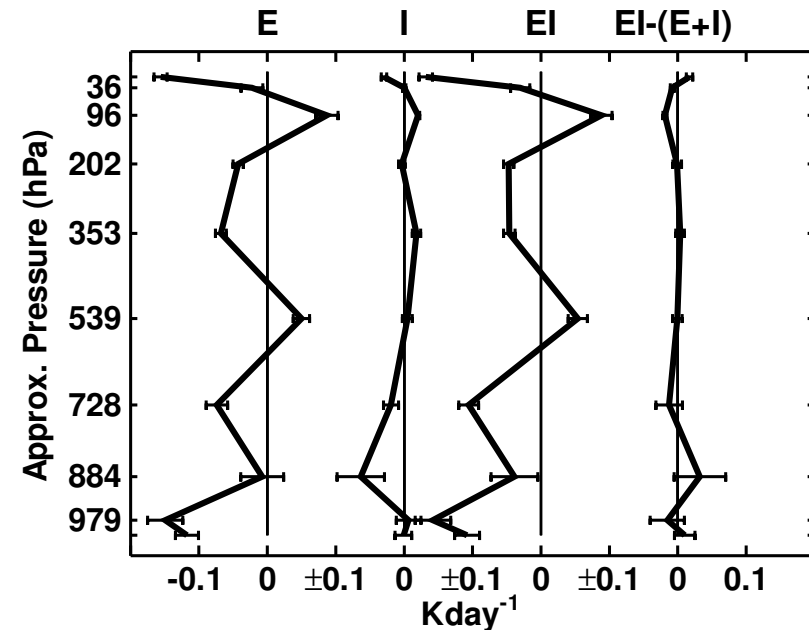
$$\frac{\partial M_{\rho_1, \rho_2, \dots, \rho_n}}{\partial t} = \frac{\partial M_{\rho_1}}{\partial t} + \frac{\partial M_{\rho_2}}{\partial t} + \dots + \frac{\partial M_{\rho_n}}{\partial t} \quad ?$$

Non-linear tendency term not significantly different from zero in troposphere

Initial Tendencies may be “linear enough”

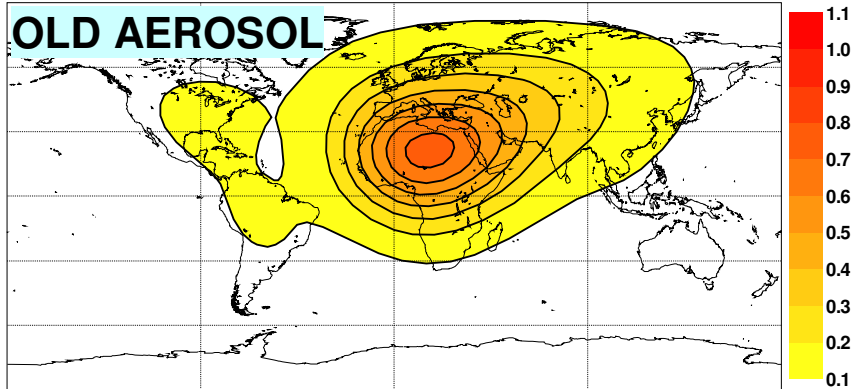
Linearity between processes: 1275 models to assess

Anomalous T Tendencies at 60°S



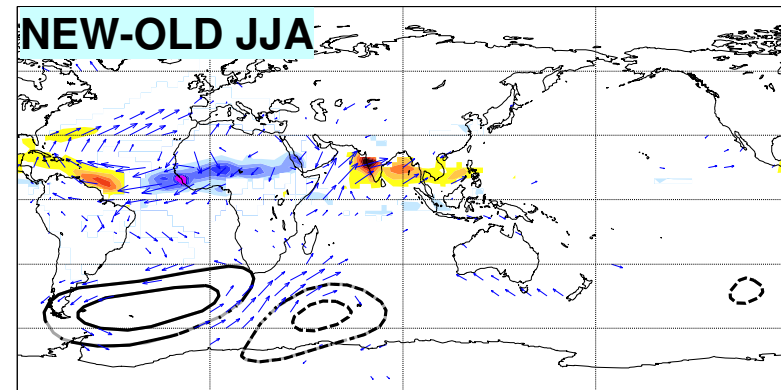
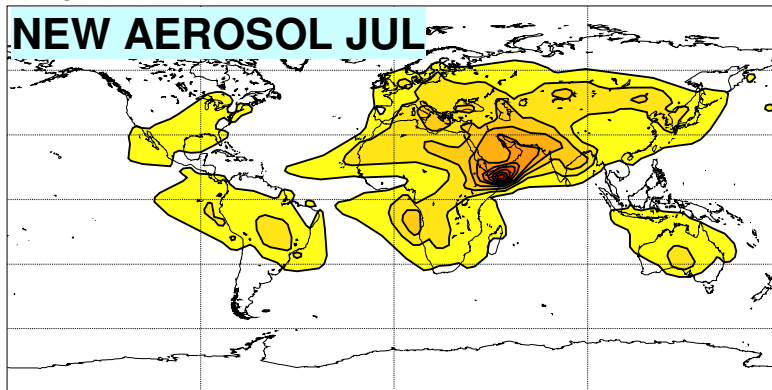
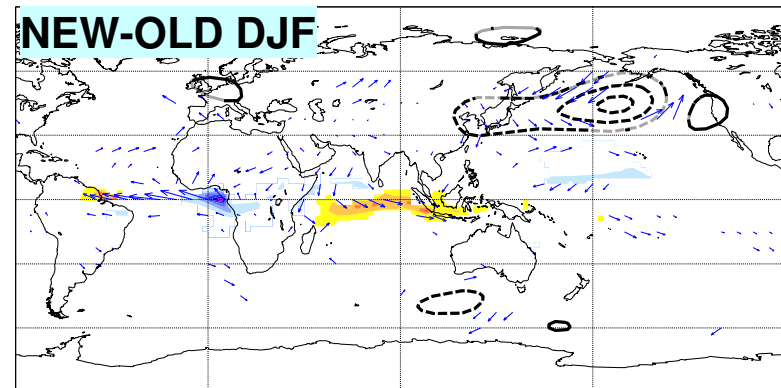
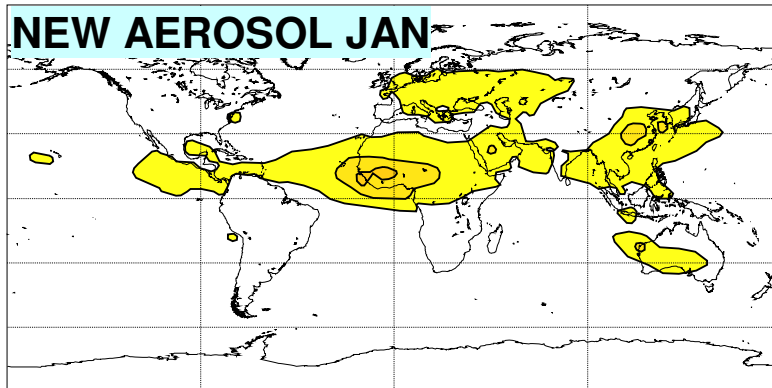
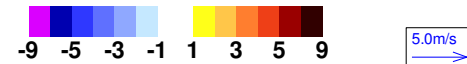
70% confidence intervals shown
 E = ENTRAINX3 - CONTROL
 I = ICE SIZE X 2 - CONTROL
 EI = (ENTRAINX3 & ICE SIZE X 2) - CONTROL

Optical Depth @ 550nm



Seasonal Impact of Aerosol Change:

Precipitation coloured in mmday^{-1}
 Z_{500} contour interval 2 dam
 \underline{v}_{925} vectors



Local Response: Tendencies over N. Africa (Jul 2004)

Old aerosol

- Initial lower-tropospheric heating error (red)
- At D+5, increased convection (blue) & ascent (orange)

New aerosol

- Initial & day 5 profiles more stable

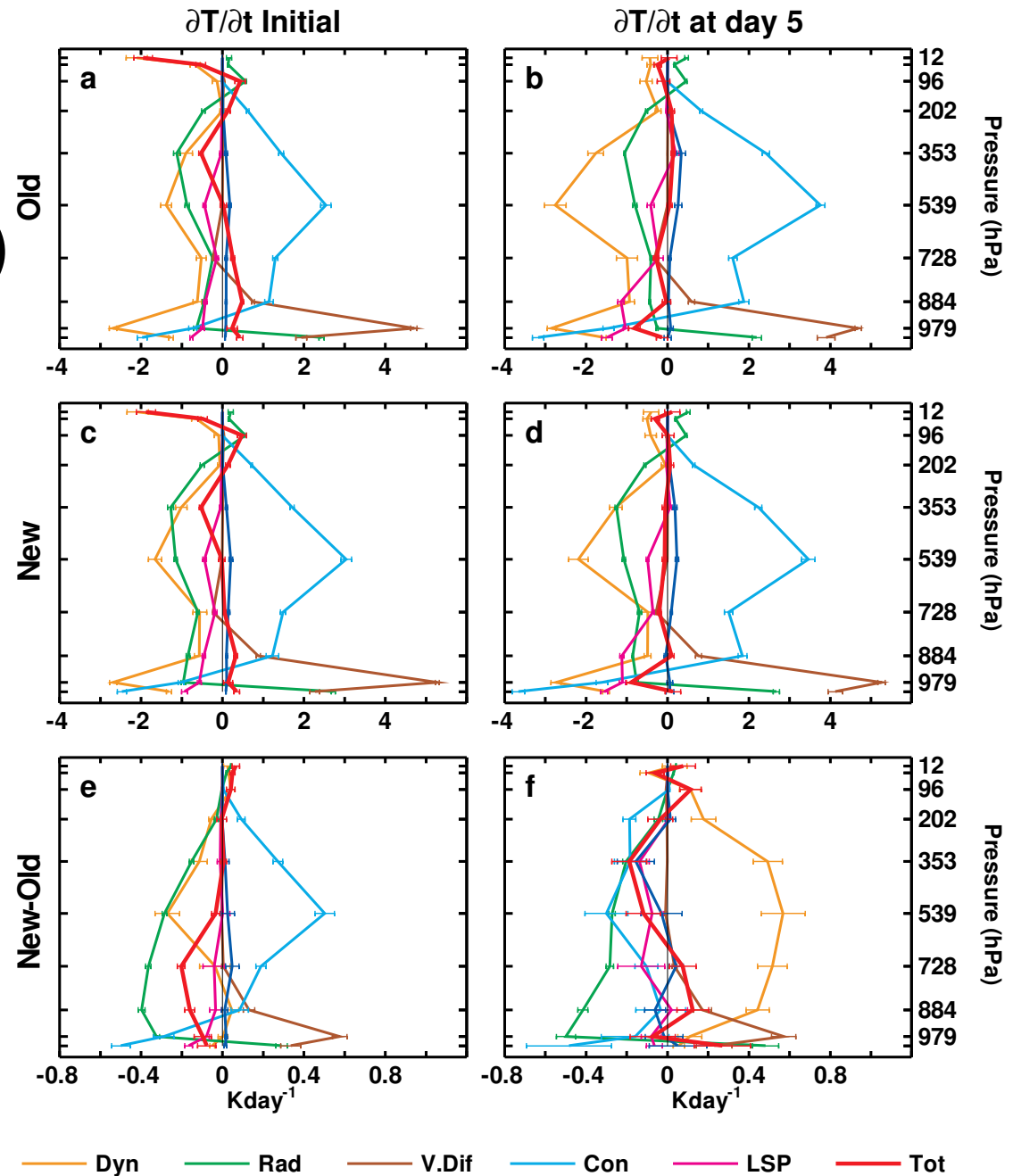
Difference

- Reflects radiative changes (green)

D+5 Precipitation

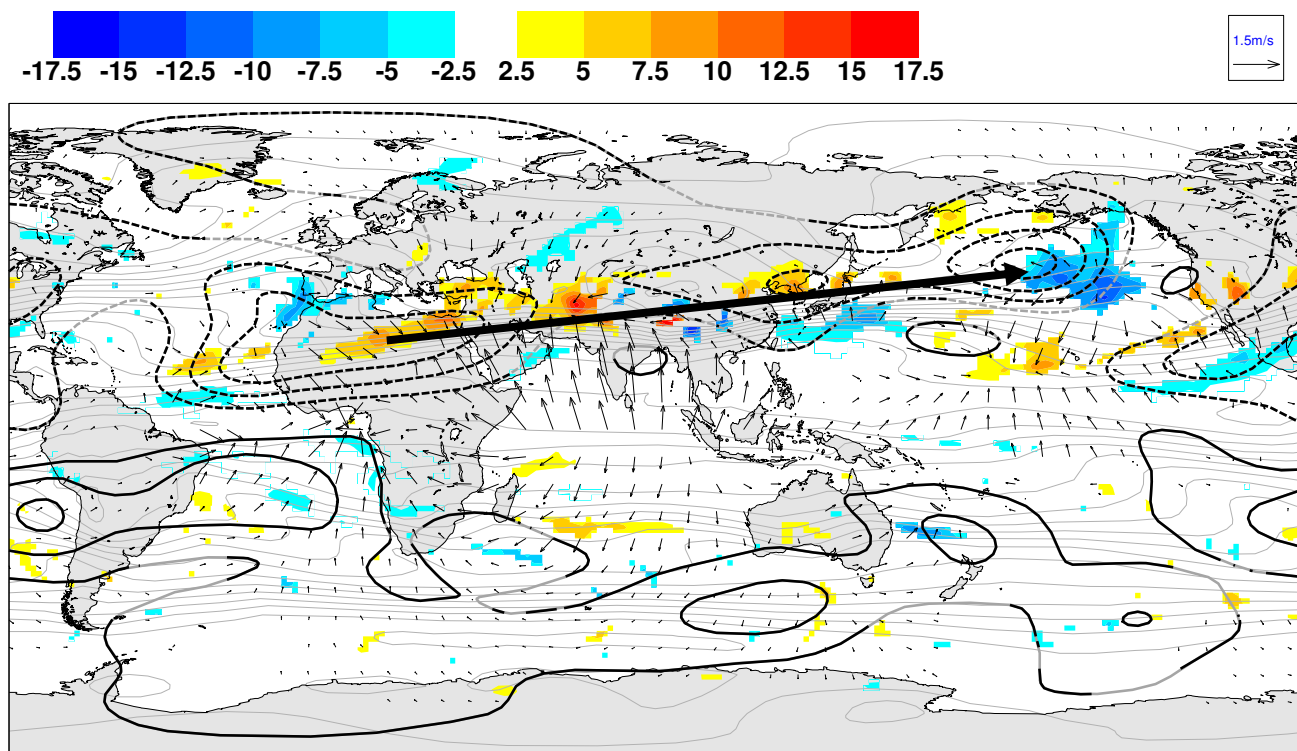
Old: 5.8mmday⁻¹
 New: 4.8mmday⁻¹
 GPCP: 4.7mmday⁻¹

Region = 20°W-40°E, 5°N-15°N
 70% confidence intervals shown



— Dyn — Rad — V.Dif — Con — LSP — Tot

Remote Response: 250hPa Streamfunction, Divergent Wind and Rossby-Wave-Source



- Aerosol Change**
- Local Convection Improvement
 - Remote Tropical Convection Improvement (via equatorial waves)
 - Changes in Extratropical Vorticity Forcing
 - North Pacific Bias Improvement

RWS coloured with contour $2.5 \times 10^{-11} \text{s}^{-2}$. Streamfunction thick contours with interval $2 \times 10^6 \text{m}^2 \text{s}^{-1}$. Mean absolute vorticity thin grey with interval 10^{-5}s^{-1}