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Executive Summary

The goal of simulating the Earth’s climate with mathematical models, using the most powerful computers available, is valid scientifically and fully consistent with the approaches being taken in many other fields of science dealing with very complex systems. These climate simulations provide the frame within which improved understanding of climate-relevant processes and improved observations are naturally merged into coherent projections of future climate change.

The science of climate models has matured to the point that many aspects of current climate models and simulations are very convincing. These form a growing set that intersects significantly with, but does not completely cover, the set of processes that are centrally important for the attribution of past climate changes and the projection of future climate.

The set of the most recent climate simulations, referred to as the CMIP3 models and utilized heavily in the Working Group 1 and 2 reports of the 4th IPCC Assessment, have received unprecedented scrutiny by hundreds of investigators with differing areas of expertise. While there are a number of systematic biases across the set of models, more generally the strengths and weaknesses of the simulations, when compared against the current climate, vary substantially from model to model. It is clear from many perspectives that an average over the set of models provides a superior climate simulation than any individual model, justifying the multi-model approach taken in many recent attribution and climate projection studies.

The pace of climate model improvement has been steady over the past several decades, but the improvement has understandably been uneven, because several important aspects of the climate system present especially severe challenges to the goal of simulation.

Climate models are compared to observations of the mean climate in a multitude of ways, and their ability to simulate observed climate changes, particularly those of the past century, have been examined extensively. However, it has proven difficult to measure the quality of climate models in such a way that the metric used is directly relevant to our confidence in the models’ projections of

1 future climate. The most appropriate ways of translating the strengths and weaknesses of the
2 simulations into confidence in climate projections remains a subject of active research.

3
4 The climate models developed in the US and around the world show many consistent features in
5 their simulations and projections for the future. However, they have not fully converged, since
6 different groups approach uncertain aspects of the models in distinctive ways. This absence of
7 convergence is one useful measure of the state of the science of climate simulation; convergence is
8 to be expected once all climate-relevant processes are simulated in a convincing physically-based
9 manner.

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14 *Climate Sensitivity*

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16 The response of global mean temperature to a doubling of carbon dioxide remains a useful measure
17 of climate sensitivity. The equilibrium response, the response expected if one waits long enough
18 (many hundreds of years) for the system to re-equilibrate, is the most commonly quoted measure.
19 The range of equilibrium climate sensitivity obtained from models has remained robust for three
20 decades, and roughly consistent with estimates from the observations of recent climates and those
21 from the more distant past. The canonical three-fold range of uncertainty, 1.5-4.5 degrees
22 Centigrade, has evolved very slowly. The lower limit has been particularly robust over time, with
23 very few recent models below 2 degrees. The difficulty in simulating the Earth's clouds and their
24 response to climate change are the fundamental reason why it has proven difficult to reduce the
25 range of uncertainty in model-generated climate sensitivity.

26
27 Other common measures of climate sensitivity are of more relevance to the response on time scales
28 shorter than 100 years. By these measures there is considerably less spread among the models --
29 roughly a factor of two rather than three. Uncertainty still remains considerable and is not
30 decreasing rapidly, due in part to the difficulty of cloud simulation but also to uncertainty in the rate

1 of heat uptake by the oceans which rises in importance when considering the responses on these
2 shorter time scales.

3
4 Improvements in our confidence in estimates of the sensitivity of climate are most likely to arise
5 from new data streams, such as satellite platforms that are now providing a first look at the 3-
6 dimensional global distributions of clouds, and new, very computationally intensive, climate
7 modeling strategies that explicitly resolve some of the smaller scales of motion that help control
8 cloud cover and cloud radiative properties.

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11 *Regional modeling and downscaling*

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13 Simulations by limited-area models, stretched grid models and uniformly high-resolution
14 atmospheric models forced by specified oceanic and sea ice conditions are all capable of resolving
15 phenomena too fine for standard atmosphere-ocean GCMs, such as precipitation influenced by
16 mountains and ocean-land interaction in coastal zones. These dynamical downscaling strategies are
17 beneficial when supplied with appropriate sea-surface and atmospheric boundary conditions, but
18 their value is limited by uncertainties in the information supplied by the global models. Given the
19 value of multi-model ensembles for larger-scale climate prediction, it is clear that downscaling must
20 presently be performed in a coordinated fashion with a representative set of global model
21 simulations as input, rather than focusing on the results from one or two models. Relatively few
22 such multi-model dynamical downscaling studies have been performed to date.

23
24 Statistical techniques to produce appropriate small-scale structures from climate simulations,
25 referred to as “statistical downscaling”, can be as effective as high-resolution numerical simulations
26 in providing climate change information to regions unresolved by most current global models, and
27 because of their computational efficiency they can much more easily utilize a full suite of multi-
28 model ensembles. However, the statistical methods are completely dependent on the accuracy of
29 the regional circulation patterns produced by the global models, whereas regional models, through
30 higher resolution and/or better representation of important physical processes, can often improve the

1 physical realism of the simulated regional circulation. Thus, the strengths and weaknesses of the
2 regional modeling and statistical methods are often complimentary.

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5 *The quality of climate simulations*

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7 Accurate simulation of the present-day climatology for near-surface temperature and precipitation is
8 necessary for most practical applications of climate modeling. The seasonal cycle and large-scale
9 geographical variations of near-surface temperature are indeed well simulated in recent models,
10 with typical correlations between models and observations of 95% or better.

11
12 AOGCM simulation of precipitation has improved over time but is still problematic. The correlation
13 between models and observations is 50-60% for seasonal means on scales of a few hundred
14 kilometers. Comparing simulated and observed latitude-longitude maps of precipitation reveals
15 similarity of magnitudes and patterns in most regions of the globe with the most striking
16 disagreements occurring in the tropics. In most models, the appearance of the Inter-tropical
17 Convergence Zone of cloudiness and rainfall in the equatorial Pacific is distorted, and rainfall in the
18 Amazon Basin is substantially underestimated. These errors may prove consequential for a number
19 of model predictions, such as forest uptake of atmospheric CO₂.

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21
22 The simulation of the storms and jet streams in middle latitudes are considered one of the strengths
23 of atmospheric models because the dominant scales involved are reasonably well-resolved. As a
24 consequence, there is relatively high confidence in models' ability to simulate the changes in these
25 extratropical storms and jet streams as the climate changes. The deficiencies that still exist may be
26 partly due to insufficient resolution to resolve features such as fronts or to inadequacies in the
27 simulated interactions between the tropics and midlatitudes or between the stratosphere and the
28 troposphere. These deficiencies are still large enough to impact the ocean circulation and some
29 regional climate simulations and projections.

1 A dominant mode of low-frequency variability in the atmosphere known as the northern and
2 southern annular modes, are very well captured in current models. These modes involve
3 north/south displacements of the extratropical storm track and dominate the observed trends in
4 atmospheric circulation in recent decades. Because of their ability to simulate the annular modes,
5 global climate models simulate fairly well the interannual variability in the polar regions of both
6 hemispheres. They are less successful at simulating daily polar-weather variability, though finer
7 scale regional simulations do simulate polar weather well, thus showing promise for improved
8 global-model simulations as their resolution increases.

9
10 In the tropics, simulations in current models are less credible. The Madden-Julian oscillation, a
11 feature of the tropics in which the precipitation is organized by large-scale eastward propagating
12 features with periods of roughly 30-60 days, is a useful test of simulation credibility in the tropics.
13 Model performance using this measure is still unsatisfactory. The “double ITCZ-cold tongue bias”,
14 in which water is excessively cold near the equator and precipitation splits artificially into two
15 zones straddling the equator, remains as a persistent bias in current coupled atmosphere-ocean
16 models. Projections of tropical climate change are adversely affected by these deficiencies in
17 simulations of the organization of tropical convection. Models typically overpredict light
18 precipitation and underpredict heavy precipitation in both the tropics and middle latitudes, creating
19 potential biases when studying extreme events.

20
21 Tropical cyclones are poorly resolved by the present generation of global models, but recent results
22 with high resolution atmosphere-only models and dynamical downscaling provide optimism that the
23 simulation of tropical cyclone climatology will advance rapidly in the coming years, as will our
24 understanding of observed variations and trends.

25
26 Land surface modeling for climate simulation has increased markedly in sophistication over the past
27 25 years, with increasing detail and range of processes included in the biological, chemical and
28 physical behavior simulated in the terrestrial portion of the climate system. Systematic programs
29 comparing land models have gradually led to greater agreement between land models and
30 observations, in part because a greater variety of observations have been used to understand and
31 constrain their behavior.

1 Land models that predict vegetation patterns are being actively developed, but the demands that
2 these models make on the quality of the simulated precipitation patterns ensures that the their
3 evolution will be gradual and tied to improvements in the regional climate simulations.

4
5 The quality of ocean climate simulations has improved steadily in recent years, owing to improved
6 numerical algorithms and more realistic assumptions concerning the mixing occurring on scales
7 smaller than the models' grid. Many of the CMIP3 class of models are able to maintain an
8 overturning circulation in the Atlantic with approximately the observed strength without the
9 artificial correction to the air-sea fluxes commonly in use in previous generations of models,
10 providing a much better foundation for analysis of the stability of this circulation.

11
12 The circulation in the Southern Oceans, thought to be of vital importance for the oceanic uptake of
13 carbon dioxide from the atmosphere, is sensitive to deficiencies in the simulated winds and
14 salinities, but a subset of the models are producing realistic circulation in the Southern Ocean as
15 well

16
17 Simulations of El Nino oscillations provide a significant success story for climate models, as these
18 have improved substantially in recent years. Most current models spontaneously generate El Nino-
19 Southern Oscillation variability, albeit with varying degrees of realism. The spatial structure and
20 period of the oscillations is impressive in a subset of the models, but with a tendency towards too
21 short a period. The bias in the intertropical convergence zone in the coupled models is a major
22 factor preventing further improvement in these models. Projections for the future of El-Nino
23 variability and the state of the Pacific Ocean are of central importance for regional climate change
24 projections throughout the tropics and in North America.

25
26 The quality of simulations of low frequency variability on decadal to multi-decadal time scales
27 varies regionally and also varies substantially from model to model. On average, the models do
28 reasonably well in the North Pacific and North Atlantic. In other oceanic regions, data paucity
29 contributes to the uncertainty in the estimation of the quality of the simulations at these low
30 frequencies.

31

1 The ocean components of current climate models do not directly simulate the very energetic
2 motions in the oceans referred to as “meso-scale eddies” . The simulation of these small scale flow
3 patterns requires horizontal grid sizes of 10km or smaller. Current oceanic components of climate
4 models are effectively laminar rather than turbulent, and the effects of these eddies must be
5 approximated by imperfect theories. As computer power increases, new models that resolve these
6 eddies will be incorporated into climate models to explore their impact on decadal variability, as
7 well as heat and carbon uptake.

8
9 Models of glacial ice are in their infancy. Glacial models directly coupled to atmosphere-ocean
10 models typically only account for direct melting and accumulation at the surface of the ice-sheets
11 and not the dynamic discharge due to glacial flow. More detailed current models that incorporate
12 this discharge typically generate discharges that change only over centuries and millennia. Recent
13 evidence for rapid variations in this glacial outflow indicates that more realistic glacial models are
14 needed to estimate the evolution of future sea level.

15
16 *Simulation of 20th century trends*

17
18 Models forced by the observed well-mixed greenhouse gas concentrations, volcanic aerosols, as
19 well as estimates of variations in the solar energy incident on the Earth and anthropogenic aerosol
20 concentrations, are able to simulate the 20th century global mean temperature record in a plausible
21 way. Solar variations are known by direct satellite measurements for the last few decades and do
22 not contribute significantly to the warming during that period. Solar variations earlier in the 20th
23 century are much less certain, but are thought to a potential contributor to the warming in the early
24 part of the century,

25
26 Uncertainties in the climatic effects of man-made aerosols (liquid and solid particles suspended in
27 the atmosphere) are a major stumbling block in quantitative attribution studies and in attempts to
28 use the observational record to constrain climate sensitivity. We do not know how much warming
29 due to greenhouse gases has been cancelled by cooling due to aerosols. Uncertainties related to
30 clouds increase the difficulty in simulating the climatic effects of aerosols, since these aerosols are
31 known to interact with clouds and potentially change cloud radiative properties and cloud cover.

1
2 The possibility that natural variability has been a significant contributor to the detailed time
3 evolution seen in the global temperature record is plausible, but still difficult to address with models
4 given the large differences between models in the characteristics of the natural decadal variability
5 that they generate. While natural variability may very well be relevant to observed variations on the
6 scale of 10-30 years, no models show any hint of generating large enough natural, unforced
7 variability on the 100 year time scale that would compete with explanations of the observed
8 century-long warming trend as being predominantly forced.

9
10 The observed southward displacement of the Southern hemisphere storm track and jet stream in
11 recent decades is reasonably well simulated in current models, which show that it is partly due to
12 greenhouse gases but also partly due to the presence of the ozone hole in the stratosphere. Northern
13 Hemisphere circulation changes over the past decades have proven more difficult to capture in
14 current models, perhaps due to the more complex interactions between the stratosphere and the
15 troposphere in the Northern Hemisphere.

16
17 Observations of ocean heat uptake are beginning to provide a direct test of aspects of the ocean
18 circulation directly relevant to climate change simulations. Coupled models provide reasonable
19 simulations of the observed heat uptake in the oceans, but underestimate the observed sea level rise
20 over the past decades.

21
22 Model simulations of trends in extreme weather typically produce global increases in extreme
23 precipitation and severe drought, and decreases in extreme minimum temperatures and frost days, in
24 general agreement with observations.

25
26 Regional trends in extreme events are not always captured by current models , but it is difficult to
27 assess the significance of these discrepancies, and to distinguish between model deficiencies and
28 natural variability.

29
30 The use of climate model results to assess economic, social, and environmental impacts is becoming
31 more sophisticated, albeit slowly. Simple methods requiring only mean changes in temperature and

1 precipitation to estimate impacts remain popular, but an increasing number of studies are utilizing
2 more detailed information, such as the entire distribution of daily or monthly values and extreme
3 outcomes. The mismatch between the spatial resolution of models and the scale of impacts-relevant
4 climate features and of impacts models remains an impediment for certain applications.