

798 **Context: Sea-Level Rise and Its Effects on the Coast**

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804

805 The accumulation of scientific evidence over the past several decades unequivocally
806 demonstrates that the global climate is changing, largely due to carbon dioxide emissions
807 from human activities (IPCC, 2001; 2007). Sea-level rise is one effect of climate
808 warming that will have profound impacts on all coastal regions of the United States and
809 around the world. The geologic record shows that sea level and the global climate have
810 been relatively stable over the past 10,000 years and this stability is a significant factor in
811 enabling the development of human civilizations. The significant changes over the past
812 200 years in atmospheric carbon dioxide, temperature, ecosystems, and ice-sheet melting
813 follow a six-fold increase in global population (Zalasiewicz et al., 2008). Along the ocean
814 and estuarine coasts of most of the United States, sea level has risen over the last century
815 and will continue to do so in the future. The effects are evident in many areas, as shores
816 erode and move landward and formerly dry areas become submerged, more frequently
817 flooded by high tides and storm surges. People are responding to these impacts by taking
818 measures to protect threatened property or by relocating development inland to higher
819 ground. The intent of this report is to assess the potential effects and risks of sea-level

820 rise on coastal regions and provide information needed to understand the implications and
821 options for dealing with sea-level rise.

822

823 The effects of sea-level rise are likely to intensify and become more pervasive in the
824 coming decades as the Earth's climate warms. Throughout geologic history, climate
825 change has been the main factor driving the evolution of Earth and its inhabitants. Now,
826 climate is changing rapidly, largely in response to human activity (IPCC, 2007). Many
827 impacts of human-induced climate change are already occurring, including, melting
828 glaciers and ice sheets; changes in extreme weather, such as heavy downpours and
829 droughts, and an accelerated rise in sea level. These physical changes are also leading to
830 biologic responses such as changes in the range of species, earlier spring events (such as
831 animal migration), and a loss of habitat, such as coastal wetlands (IPCC, 2007). The rates
832 of warming occurring now and those projected for the future may exceed the ability of
833 many living organisms to adapt without major disruptions and extinctions. With future
834 warming and wide spread ice sheet melting too, sea-level rise could accelerate very
835 rapidly on decadal scales and follow non-linear patterns that would have large impacts on
836 coastal regions.

837

838 More extreme weather events and storm activity and a world-wide rise in sea level are
839 two of the most likely, most disruptive, and most costly effects of global warming. Often
840 these two elements of climate change act in concert with each other to impact coastal
841 regions. They have most effect on coastal regions where the land relief is generally low,
842 land forms are susceptible to erosion, and human population and development are highly

843 concentrated. This includes much of the coast around the United States, but the mid-
844 Atlantic region (the main focus of this report) is particularly vulnerable due to high rates
845 of relative sea-level rise and dense coastal development.

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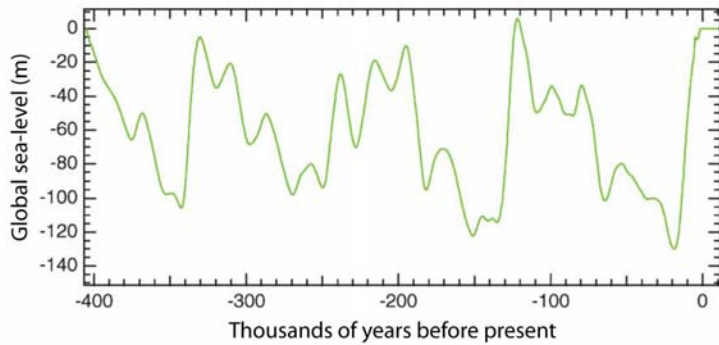
847 This report reviews available scientific literature and presents a scientific consensus on
848 the likely effects of sea-level rise on the mid-Atlantic coast of the United States, the
849 human and environmental impacts, likely responses in the context of current policies and
850 economic trends, and possible options for changing planning and management activities
851 so that society and the environment are better able to cope with an accelerated rise in sea
852 level. A summary of implications on a Nation-wide scale are presented in Part V. The
853 Preface of this report contains further information on the process for developing this
854 report, the nature of the regional focus, and the structure of this report.

855

856 **C.1 WHY IS GLOBAL SEA LEVEL RISING?**

857 The elevation of global sea level is determined primarily by the balance between the
858 volume of ice on land (in glaciers and ice sheets) and the volume of water in ocean
859 basins. During the last 800,000 years, sea level has risen and fallen in response to the
860 buildup and decline of large ice sheets as climate warmed and cooled in natural cycles of
861 approximately 100,000 years. Figure C.1 shows a record of sea level change over the past
862 400,000 years.

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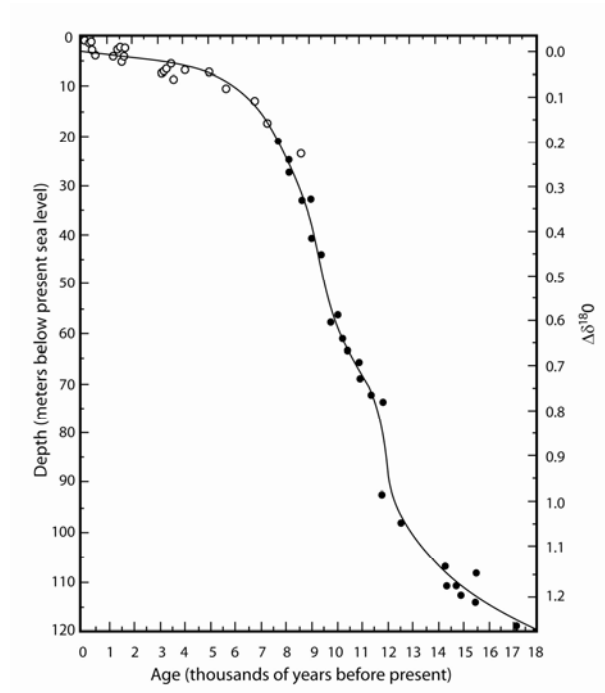


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866 **Figure C.1** Sea level change over the last 400,000 years resulting from natural glacial- interglacial cycles.
867 Evidence suggests that sea level was about 4-6 m higher than present during the last interglacial warm
868 period 125,000 years ago, and 120 m lower during the last Ice Age, about 21,000 years ago. Modified from
869 Huybrechts (2002).
870

871 In the recent geologic past, sea level has varied from 120 m (400 ft) lower than present
872 during the last Ice Age, when massive glaciers covered much of North America, northern
873 Europe, and Asia, and the shoreline was seaward at the edge of the continental shelf, to
874 about 4 to 6 m (20 ft) higher than present during the previous ‘interglacial’ (non-Ice Age)
875 warm period when the coast was much further inland than present day. As ice sheets
876 melted and climate warmed following the Ice Age, beginning approximately 21,000 years
877 ago, sea level rose. Global sea level reached close to its current position about 3,000
878 years ago (Figure C.2) and has fluctuated only slightly until the past several decades
879 when tide gauge and satellite data indicate an acceleration in sea-level rise rates. The
880 ocean has absorbed more than 80 percent of the atmospheric warming since 1961,
881 causing sea water to expand, contributing to this recent rise. In addition, rapid melting of
882 land-based glaciers as well as ice sheets on Greenland and Antarctica have very likely
883 increased sea-level rise (IPCC, 2007). The combination of stable sea level and moderate

884 climate during the current interglacial warm period has been a major factor contributing
885 to the growth in human development and our modern civilization (Day *et al.* 2007).
886



887

888 **Figure C.2** Rise in global sea-level over the last 18,000 years to the present time reconstructed from
889 oxygen isotope concentrations and radiocarbon dating of geologic samples, shown as data points.
890 (Modified from Fairbanks, 1989).

891

892 The study of climate change and associated sea-level rise is complex. The most credible
893 and comprehensive body of scientific information on the subject, based on a consensus of
894 approximately 2,500 of the world's scientists, has been compiled by the United Nations'
895 Intergovernmental Panel on Climate Change (IPCC) in a series of reports issued
896 approximately every five years. The most recent IPCC (2007) report, *Climate Change*
897 *2007: The Physical Science Basis*, contains a comprehensive review and assessment of

898 climate change trends, expected changes over the next century, and the impacts and
899 challenges that both humans and the natural world are likely to be confronted with during
900 the next century. In addition, the U.S. Climate Change Science Program (CCSP)
901 Synthesis and Assessment Products (SAPs), including this one, are providing detailed
902 climate information for the United States. This SAP, discussing the impacts of sea-level
903 rise on the U.S., relies heavily on IPCC (2007) findings and predictions for sea-level rise.
904 A few key findings of the most recent IPCC reports are summarized in Box C.1

905

906 **BOX C.1 SELECTED IPCC (2007) FINDINGS ON CLIMATE AND SEA-LEVEL RISE**

907

908 **Recent Global Climate Change:**

909

910 • Warming of the climate system is unequivocal, as is now evident from observations of increases in
911 global average air and ocean temperatures, widespread melting of snow and ice, and rising global average
912 sea level

913

914 • Carbon dioxide is the most important human-caused greenhouse gas. The atmospheric
915 concentration of carbon dioxide in 2005 exceeds by far the natural range over the last 650,000 years

916

917 • Most of the observed increase in global average temperatures since the mid-20th century is *very*
918 *likely* due to the observed increase in human-caused greenhouse gas concentrations. Discernible human
919 influences now extend to other aspects of climate, including ocean warming, continental-average
920 temperatures, temperature extremes and wind patterns

921

922 **Recent Sea-Level Rise**

923

924 • Observations since 1961 show that the average temperature of the global ocean has increased to
925 depths of at least 3000 m and that the ocean has been absorbing more than 80% of the heat added to the
926 climate system. Such warming causes seawater to expand, contributing to sea-level rise

927

928 • Mountain glaciers and snow cover have declined on average in both hemispheres. Widespread
929 decreases in glaciers and ice caps have contributed to sea-level rise (ice caps do not include contributions
930 from the Greenland and Antarctic ice sheets)

931

932 • New data show that losses from the ice sheets of Greenland and Antarctica have *very likely*
933 contributed to sea-level rise between 1993 and 2003

934

935 • Global average sea level rose at an average rate of 1.8 [1.3 to 2.3] mm per year between 1961 and
936 2003. The rate was faster between 1993 and 2003: about 3.1 [2.4 to 3.8] mm per year. Whether the faster
937 rate for 1993 to 2003 reflects decadal variability or an increase in the longer term trend is unclear. (Figure
938 C.3)

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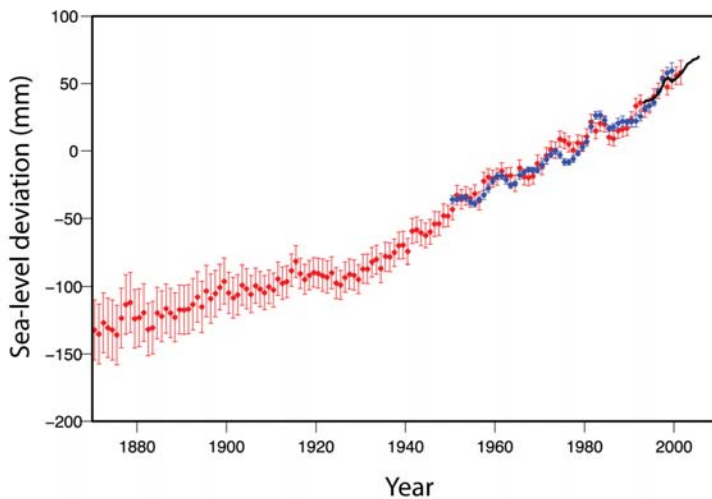
940 • Global average sea level in the last interglacial period (about 125,000 years ago) was *likely* 4 to 6
941 m higher than during the 20th century, mainly due to the retreat of polar ice. Ice core data indicate that

942 average polar temperatures at that time were 3°C to 5°C higher than present, because of differences in the
 943 Earth’s orbit. The Greenland ice sheet and other arctic ice fields *likely* contributed no more than 4 m of the
 944 observed sea-level rise. There may also have been contributions from Antarctica ice sheet melting.
 945

946 **Projections of the Future:**

- 947 • Continued greenhouse gas emissions at or above current rates would cause further warming and
 948 induce many changes in the global climate system during the 21st century that would *very likely* be larger
 949 than those observed during the 20th century.
- 950 • Based on a range of possible greenhouse gas emission scenarios for the next century, the IPCC
 951 estimates the global increase in temperature will likely be between 1.1 and 6.4°C. Estimates of sea-level
 952 rise for the same scenarios are 0.18m to 0.59 m, excluding the contribution from accelerated ice discharges
 953 from the Greenland and Antarctica ice sheets.
- 954
- 955 • Extrapolating the recent acceleration of ice discharges from the polar ice sheets would imply an
 956 additional contribution up to 20 cm. If melting of these ice caps increases, larger values of sea-level rise
 957 cannot be excluded.
- 958
- 959 • In addition to sea-level rise, the storms that lead to coastal storm surges could become more
 960 intense. The IPCC indicate that based on a range of computer models, it is *likely* that hurricanes will
 961 become more intense, with larger peak wind speeds and more heavy precipitation associated with ongoing
 962 increases of tropical sea surface temperatures, while the tracks of ‘winter’ or non-tropical storms are
 963 projected to shift towards the poles along with some indications of an increase in intensity in the North
 964 Atlantic.

965 -end-text box-
 966
 967



968
 969 **Figure C.3** Annual averages of global mean sea level from IPCC (2007). The red curve shows sea-level
 970 fields since 1870 updated from Church and White (2006); the blue curve displays tide gauge data from
 971 Holgate and Woodworth (2004), and the black curve is based on satellite altimetry from Leuliette *et al.*
 972 (2004). The red and blue curves are deviations from their averages for 1961 to 1990, and the black curve is
 973 the deviation from the average of the red curve for the period 1993 to 2001. Error bars show 90%
 974 confidence intervals. Modified from Bindoff *et al.* (2007).
 975

976 Global sea-level rise – resulting from the balance between global ice volume and ocean
977 seawater volume - is a useful measure of the general direction of change; however there
978 are substantial local and regional variations in the rates of sea-level rise. In some
979 locations, subsidence of the land increases the ‘effective’ or ‘relative’ sea-level rise,
980 whereas in other locations, local sea-level rise is less than the global average because the
981 land is still rising (rebounding) from a time when an ice sheet, sometimes a mile thick,
982 covered the area, depressing the Earth’s crust. In a few cases, such as in the Pacific
983 Northwest of the U.S., this can lead to a drop in local sea level. In responding to sea-level
984 rise, it is necessary to refer to the local (relative) sea level-rise because it is this
985 combination of global effects and local conditions that impact the coast. Thus in this
986 report, ‘sea-level rise’ refers to relative sea-level rise. See box C.2 for further discussion.
987

988 **Box C.2 Relative Sea Level**

989 The term “global sea level”, sometimes referred to as eustatic sea level, refers to the average level of tidal
990 waters around the world based on long-term measurements from coastal tide gauges. The most reliable data
991 are from gauges having records of 50 years or longer and are important observation instruments for
992 measuring sea level change trends. Vertical movements of the land surface at the coast can also contribute
993 significantly to sea-level change and the combination of sea level and land-level change is referred to as
994 “relative sea level” (Douglas, 2001). These two terms used by scientists are defined as follows:

- 995 • “global sea-level rise” is the worldwide increase in the volume of the world’s oceans that
996 occurs as a result of thermal expansion and melting ice caps and glaciers.
- 997 • “relative sea-level rise” refers to the change in sea level relative to the elevation of the land,
998 which includes both global sea-level rise and vertical movements of the land.
999

1000 In this report, the term “sea-level rise” is used to mean “relative sea-level rise.”

1001

1002 Vertical changes of the land surface result from many factors including tectonic processes, adjustment of
1003 the Earth’s crust, compaction of sediments, and extraction of subsurface fluids such as oil, gas, and water.
1004 A principal contributor to this change along the Atlantic coast of North America and northern Europe is the
1005 plastic-like adjustment of the Earth’s crust to changing ice loads since the Ice Age. The thick accumulation
1006 of ice on continental landmasses depressed the Earth’s surface in ice-covered regions. This displaced the
1007 mantle (the layer of the planet beneath the crust) causing a “peripheral bulge” some distance from the edges
1008 of the thick continental ice cover. As a result of these crustal adjustments, relative sea level records vary
1009 greatly along the coast from glaciated regions in New England southward to North Carolina. These vertical
1010 crustal adjustments have persisted for thousands of years and will continue to persist for some time. In
1011 addition to glacial adjustments, sediment loading also contributes to regional subsidence of the land
1012 surface. Subsidence contributes to high rates of relative sea level (>100 mm/yr) in the Mississippi River

1013 delta where thick sediments have accumulated. Likewise, fluid withdrawal from coastal aquifers causes the
1014 sediments to locally compact as the water is extracted. In Louisiana, Texas, and the southern California
1015 region, oil, gas and ground-water extraction have contributed markedly to subsidence and relative sea level
1016 rise (Gornitz and Lebedeff, 1987, Emery and Aubrey, 1991, Galloway *et al.*, 1999; Morton *et al.*, 2004).
1017 Last, tectonic uplift affects the rates of relative sea level rise from Alaska to California. In places where the
1018 land surface is uplifted due to tectonic activity, rates of relative sea-level rise may be notably smaller than
1019 the rate of global sea level rise or in some cases, reversed, with localized relative sea-level fall. In locations
1020 where the land surface is subsiding, rates of relative sea-level rise may exceed the rate of global rise (*e.g.*,
1021 the central Gulf of Mexico coast and mid-Atlantic coast).

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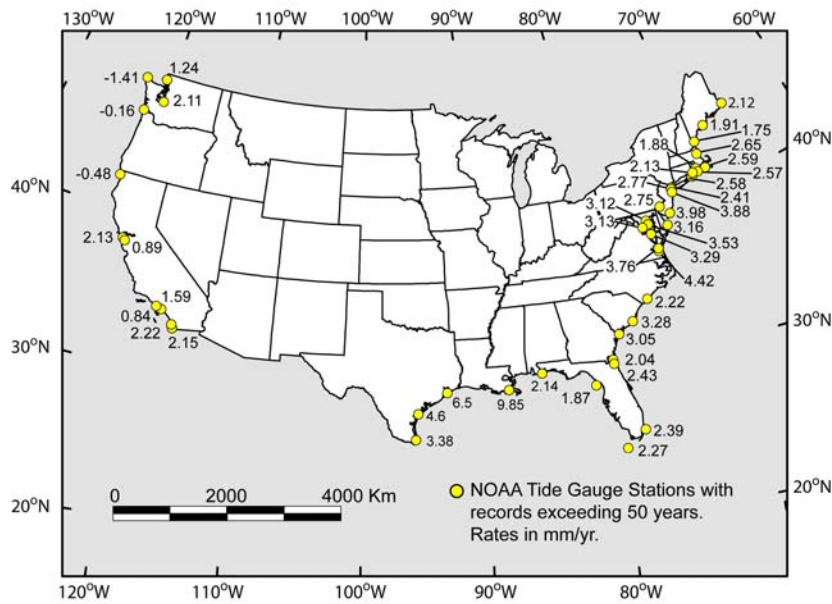
1024 C.2 SEA-LEVEL RISE AROUND THE UNITED STATES

1025 Sea level has varied greatly throughout the Earth's history due to a variety of geologic,
1026 oceanographic, and climatic processes (Douglas, 2001) and is influenced by many factors
1027 that operate globally to locally over a wide range of time scales, including days to weeks
1028 (tides, storms), seasons, decades, and millennia.

1029

1030 The long-term records from tide gauge stations have been the primary measurements of
1031 relative sea level trends over the last century (Douglas, 2001). Figure C.3 shows the
1032 variations in relative sea level for U.S. coastal regions. Many parts of the eastern and
1033 Gulf shores are showing higher rates of sea-level rise than for the world as a whole. For
1034 example, sea level is rising 3-4 mm/yr along the mid-Atlantic region compared to the
1035 absolute rate of 1.8 mm/yr for the world (Figures C.3, C.4)

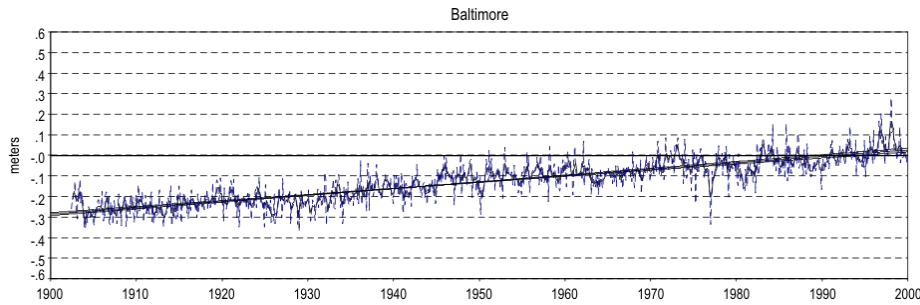
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Figure C.4 Map of annual relative sea-level rise rates around the U.S. coast. The high rates for Louisiana (9.9 mm/yr) and the mid-Atlantic region (3–4 mm/yr) are due to land subsidence. Sea level is stable or dropping relative to the land in the Pacific northwest, where the land is tectonically active or rebounding upward in response to the melting of ice sheets (compiled by USGS from Zervas, 2001).

1044 NOAA routinely produces updated estimates of relative sea level trends observed at tide
1045 stations around the country and the results show a large variation of trends from very
1046 high rates of relative sea level rise in southern Louisiana (+ 9.9 mm/yr (+/- 0.35 mm) at
1047 Grand Isle) due to land subsidence, to high rates of relative sea level fall in southeast
1048 Alaska (- 16.7 mm/yr (+/- 0.42 mm) at Skagway) due to land rebound as a result of
1049 glacier melting (Zervas, 2001). Figure C.5 is an example of the monthly average (mean)
1050 sea level record and the computed relative sea-level rise trend at Baltimore, MD. Here,
1051 the relative sea level trend is 3.12 mm/yr (+/- 0.08), which, as a result of land subsidence,
1052 is nearly 2 times the present rate of global sea-level rise.



1053

1054 **Figure C.5** Sea-level rise for Baltimore, MD from 1900 to 2000. The plot shows the monthly mean sea
 1055 level with the average seasonal cycle removed (blue dashed line), a 5-month average (black solid line), and
 1056 the linear trend.

1057

1058 C.2.1 Future Sea-Level Rise Around the United States: Our Approach

1059 This report does not develop new estimates of future sea-level rise. Instead, we use three
 1060 scenarios of relative sea-level rise along the mid-Atlantic coast:

1061 Scenario 1: Continuation of the 20th century rate (3 mm/yr)

1062 Scenario 2: An acceleration of 2 mm/yr over the 20th century trend (total rate of 5
 1063 mm/yr)

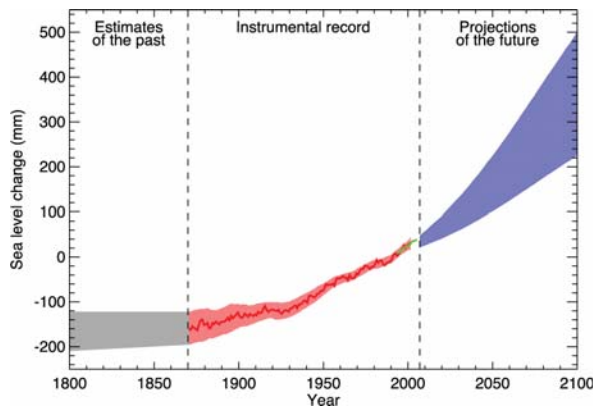
1064 Scenario 3: An acceleration of 7 mm/yr over the 20th century trend (total rate of 10
 1065 mm/yr)

1066 These three scenarios enable an assessment of the implications of a rise of 30 cm, 50 cm,
 1067 and 100 cm over the next century.

1068

1069 These scenarios are broadly consistent with recent assessments by the IPCC (2007) and
 1070 others (see Figure C.6). The IPCC's likely range for a global rise in sea level is 10-59 cm
 1071 over the next century, excluding the possibility of increased ice melting on Greenland and
 1072 Antarctica. IPCC also states that extrapolating the central estimate of current accelerated

1073 ice discharge would add another 10-20 cm, implying a range of 20-79 cm. The upper end
1074 of that range represents a 6 mm/yr acceleration over the 20th century global sea level
1075 trend. Scenario 3 is 1mm/yr higher, and substantially less than the high estimates
1076 suggested by more recent publications (Rahmstorf, 2007; Rahmstorf *et al.*, 2007; Hansen
1077 *et al.*, 2007). Scenario 2 is consistent with the best estimates for the various IPCC
1078 emission scenarios, which generally represented an acceleration of 2 mm/yr above the
1079 historic rate. Finally, Scenario 1 is consistent with the IPCC's low estimate of future and
1080 current sea-level rise.
1081



1082
1083 **Figure C.6** Past, present, and projected global sea-level rise. Time-series of global mean sea level compiled
1084 from the past (grey shading), late 19th and 20th century observations (red and green lines and red shaded
1085 region), and future projections (blue shading) determined in the recent IPCC assessment (Bindoff *et al.*,
1086 2007). The grey shading shows the uncertainty in the estimated long-term rate of sea-level change. The red
1087 line is a reconstruction of global mean sea level from tide gauges and the shaded area indicates the range of
1088 variations from this line. The green line illustrates the global mean sea level record based on satellite
1089 altimeter measurements. The blue shaded region represents the range of model projections compiled from
1090 the IPCC assessment (Meehl *et al.*, 2007). Figure from Bindoff *et al.*(2007).
1091
1092

1093 The primary focus of this report is over the next century, but the longer term implications
1094 are also considered. Recent evaluations of changes in ice cover and glacial melting on
1095 Greenland, Antarctica, and smaller glaciers and ice caps from around the world indicate

1096 that ice loss could be more rapid than has been measured and predicted (Chen *et al.*,
1097 2006; Shepherd and Wingham, 2007; Meier *et al.*, 2007). If so, this accelerated melting
1098 could significantly raise sea-level predictions to levels (~4-6 m) during the last
1099 interglacial period over the next several hundred years (Overpeck *et al.*, 2006). The
1100 science behind these predictions is not yet well developed, but is worthy of study because
1101 of the very significant implications for all coastal regions.

1102

1103

1104 **C.3 IMPACTS OF SEA-LEVEL RISE FOR THE UNITED STATES**

1105 **C.3.1 Coastal Vulnerability Around the United States**

1106 Coastal communities and habitats will be increasingly stressed by climate change impacts
1107 interacting with development and pollution (Field *et al.*, 2007). Impacts from sea-level
1108 rise include: land loss through submergence and erosion of lands in the coastal zone;
1109 migration of coastal landforms and changes to coastal environments; increased storm-
1110 surge flooding; wetland losses; and increased salinity in estuaries and coastal
1111 groundwater aquifers. Each of these effects can have important impacts on both natural
1112 ecosystems and human developments and infrastructure. Other impacts of climate
1113 change, such as increasingly severe droughts and storm intensity—along with continued
1114 rapid coastal development—could amplify the effects of sea-level rise.

1115

1116 Sea-level rise in combination with other factors is already starting to have significant
1117 effects on the coastal zone of the United States. Flooding of low lying regions by storm
1118 surges and spring tides is becoming more frequent and causing more damage and

1119 disruptions. Around the Chesapeake Bay, wetlands are being submerged, fringe forests
1120 are dying and being converted to marsh, farm land and lawns are being converted to
1121 marsh; and some roads are routinely flooded at high tides (Douglas, 2001). “Ghost
1122 forests” of standing dead trees killed by salt water intrusion are becoming increasingly
1123 common in southern New Jersey, Maryland, Virginia, Louisiana, and North Carolina
1124 (Riggs and Ames, 2003). Rising sea level is gradually intruding into estuaries and
1125 threatening fresh-water aquifers (Barlow, 2003).

1126

1127 Rising sea level will affect to varying degrees entire coastal systems from the shoreline to
1128 the landward edge of the Coastal Plain. These physical and ecological changes that are
1129 likely to occur in the near future will also have impacts on humans and coastal
1130 development. In addition, it is uncertain how current practices in managing coastal
1131 systems for mitigating erosion and flooding are likely to affect potential future impacts.
1132 Climate change implications should be included in planning and decision making to best
1133 accommodate climate change.

1134

1135 Continued rapid coastal development exacerbates both the environmental and the human
1136 impact of rising sea level. During the 20th century, an expanding proportion of the U.S.
1137 population and associated urban development relocated to the land along the Atlantic,
1138 Gulf of Mexico, and Pacific coasts. Coastal populations have doubled in the past 30 years
1139 and although the coastal population is currently increasing at approximately the same rate
1140 as the national population, continued coastal development increasingly conflicts with the
1141 natural processes associated with coastal change from storms and sea-level rise. Currently

1142 the majority of the U.S. population lives in the coastal zone and movement to the coast
1143 and development continues. Fourteen of the Nation's 20 largest urban centers are located
1144 along the coast. In addition, these economic and population pressures have transformed
1145 sparsely developed coastal areas into high-density year-round urban complexes. With
1146 accelerated rise in sea level and increased intensity of storms, the conflicts between
1147 development at the coast and the natural processes are likely to increase dramatically
1148 unless new coastal management and planning is employed.

1149

1150 **C.3.2 Shoreline Change and Coastal Erosion**

1151 The diverse landforms comprising the more than 160,000 km of U.S. coast reflect a
1152 dynamic interaction between: 1) natural factors and physical processes that act on the
1153 coast (*e.g.*, storms, waves, currents, sand sources and sinks, relative sea level), 2) human
1154 activity (*e.g.*, dredging, dams, coastal engineering), and 3) the geological character of the
1155 coast and nearshore. Spatial and temporal variations in these physical processes and the
1156 geology along the coast are responsible for the variety of coastal landforms. As a result,
1157 the majority of the U.S. coast is undergoing long-term net erosion at highly varying rates
1158 as shown in Figure C.7.

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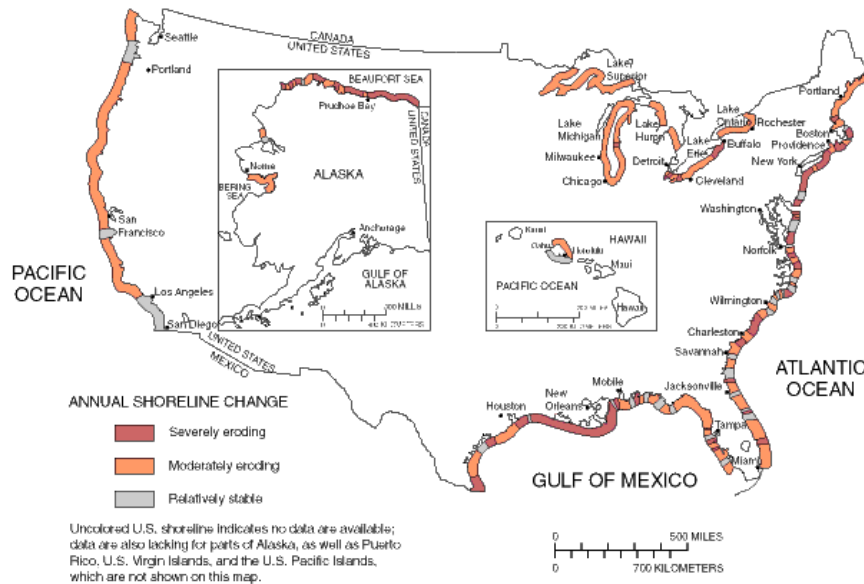
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Figure C.7. Coastal Erosion Rates Around the U.S. All 30 coastal states are experiencing erosion at highly variable rates due to natural processes and human activity. From USGS National Atlas 1985.

1169 The complex interactions among these factors make it difficult to identify a precise
 1170 relationship between sea-level rise and shoreline change and to reach consensus among
 1171 coastal scientists on quantitative approaches that can be used to predict how shorelines
 1172 will change in response to sea-level rise. The difficulty in linking sea-level rise to coastal
 1173 change stems from the fact that shoreline change does not occur directly as the result of
 1174 sea-level rise. Instead, coasts are in an almost continual state of change in response to
 1175 many driving forces and subject to the underlying geological character and the
 1176 availability of sediment to the coastal system. Consequently, while there is strong
 1177 scientific consensus that climate is changing and affecting coastal regions, there are still
 1178 uncertainties associated with quantitative predictions of how the coast will respond to
 1179 likely changes in future sea level.

1180

1181 With current planning and decision making, we often assume that these systems operate
1182 in a steady-state. While the factors that influence coastal change in response to sea level
1183 rise are well known, our ability to incorporate this understanding into computer models
1184 that can be used to predict shoreline change over long time periods is limited and models
1185 are in their infancy. Part of the reason for this is the complexity of quantifying the effect
1186 of these factors on shoreline change. The models incorporate relatively few factors that
1187 influence shoreline change and rely on assumptions that do not always apply to real-
1188 world settings. In addition, these assumptions apply best to present conditions, not
1189 necessarily those that may exist in the future. The models that do incorporate many of the
1190 key factors (*e.g.*, the geological framework and sediment budget) require detailed data
1191 (*i.e.*, sediment transport rates, landform evolution feedbacks) on a local scale. To apply
1192 over larger coastal regions, the necessary baseline information for most areas is not
1193 available. The unfortunate consequence is that our current capability to make long-term
1194 reliable predictions is limited. In addition, there is some indication that coastal landforms,
1195 such as barrier islands, might have “tipping points” or “thresholds” when limits are
1196 exceeded and the landforms become unstable and disintegrate. It is possible that this is
1197 already happening to barrier islands along the Louisiana coast and may occur in the near
1198 future along the North Carolina and the Maryland-Virginia coast with increased sea-level
1199 rise and storm activity (Culver et al., 2007; Sallenger et al., 2007; Riggs and Ames,
1200 2003).

1201

1202 This report reviews the knowledge of how sea-level rise can impact coastal regions and
1203 the challenges that we face in planning and coping with these impacts. A large part of this

1204 discussion is based on information from new assessments that address the potential
1205 impacts of sea level rise on the tidal inundation of low-lying lands, ocean shoreline
1206 processes, and the vertical accretion of tidal wetlands in the mid-Atlantic region.
1207 Following the terms of our charge from CCSP (2007), we do not evaluate the impacts of
1208 sea level rise on coastal flooding; nor do we evaluate the impacts of possible changes in
1209 the frequency and severity of coastal storms. That does not mean that the report ignores
1210 storm effects or assumes that the seas are always calm. Existing landforms, ecosystems,
1211 and human activities are already adapted to a certain level of storminess. Unless
1212 otherwise stated, the chapters that follow all assume that storms will continue in the
1213 future, and that many of the impacts of sea-level rise—on both people and the
1214 environment-- will only be realized after a severe storm.

1215

1216 **C.3.3 Managing the Coastal Zone as Sea Level Rises**

1217 Coasts are dynamic junctions of water, air, and land. The interactions vary greatly over
1218 time and space. Winds and waves, tides and currents, migrating sand dunes, and river
1219 deltas combine to form ever-changing coasts, yet development continues in high risk
1220 coastal areas. If sea level rise accelerates, all of these landforms will become more
1221 dynamic. Some researchers believe that the combination of stable sea level and moderate
1222 climate during the current interglacial period has been a major factor contributing to the
1223 growth in human development and our modern civilization (Stanley and Warne, 1993;
1224 Day *et al.*, 2007). The notion that sea level is constant and that coasts are stable is deeply
1225 embedded in many institutions, and in the assumptions of most coastal residents.

1226 Adapting to an accelerated sea-level rise would require changes in both our institutions
1227 and our mindset about natural processes.

1228

1229 A key question for coastal zone management is how and where to “mitigate” or adapt to
1230 these new coastal conditions. Shoreline erosion problems affecting property and
1231 development or coastal wetland habitat losses tend to dominate shore-protection policy
1232 rather than sea-level rise explicitly. Today, many property owners and government
1233 programs are already engaged in coastal engineering activities designed to protect
1234 property and beaches in developed areas by thwarting natural dynamic processes—but in
1235 undeveloped areas, the natural processes usually govern. At first, an acceleration of sea-
1236 level rise may simply increase the cost of current practices. Eventually, however, policy
1237 makers may have to evaluate whether the approach to coastal development and protection
1238 assuming a relatively stable sea level should be modified to best respond to the higher sea
1239 levels.

1240

1241 To facilitate these decisions, policy makers need credible information. Predicting these
1242 changes with the precision that a decision maker would prefer to have is not always
1243 possible. Yet there is little doubt that physical changes to the coastal system will also
1244 modify coastal ecosystems and the fish and wildlife. Further complicating the picture, are
1245 other related effects of climate change: storms, precipitation, run-off, drought,
1246 management practices, economic setting, and sediment supply. At present, our scientific
1247 understanding of the physical response of the coast to sea-level rise is lacking and in

1248 combination with the wide variety of human engineering activities along the shoreline,
1249 prediction of future effects with high confidence is challenging.

1250

1251 In most cases, we manage our coasts as if sea level were stable, the shoreline fixed in
1252 location, and storms were regular and predictable. In this report, several chapters examine
1253 how sea-level rise and increased storminess might require managers to consider longer
1254 term perspectives. We also examine some possible tactics for coastal planning and
1255 management that might be more effective as sea-level rise accelerates.

1256

1257 We have outlined the three sea-level rise scenarios used in this report, but in addition, we
1258 begin to consider how the impacts of sea-level rise may depend on the portion of the
1259 shoreline stabilized, as well as on the rate of sea-level rise. Unlike the future rate of sea-
1260 level rise, coastal managers collectively have some control over how much of the shore is
1261 ultimately protected, at least for the short term. Follow-on efforts will examine scenarios
1262 assuming continuation of existing policies, and will consider whether the cumulative
1263 environmental impacts might lead to a different set of choices for dealing with sea-level
1264 rise.

1265

1266 In summary, continued sea-level rise, at current or accelerated rates, coupled with
1267 increasing storm intensity, will result directly in increasing vulnerability for people,
1268 property, and ecosystems and indirectly have national implications. Coasts are likely to
1269 erode and retreat more than we would expect from inundation by sea-level rise alone,
1270 especially for fragile barrier islands and low-lying delta regions. We need continued

1271 improvement in the science of coastal change, more comprehensive systems of data
1272 collection and analysis, observation, monitoring, modeling, and communication of results
1273 to the public and policy makers. Planning and decision making for the coastal zone across
1274 all levels of government needs to reflect the new scientific understanding of climate
1275 change and effects of sea-level rise and increased storms. Improvements in
1276 communication are needed to ensure that science is more relevant to inform policy. We
1277 hope that this report sets the stage for coastal decision makers to fully incorporate the
1278 ramifications of climate change and its effects on sea-level rise into long-term
1279 management and planning.

1280

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