

# Radiance to Reflectance for GOES-8 Channel 1

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August 1996

The following is a discussion on converting GOES-8 Imager visible radiance measurements to reflectance values. An overview of GOES-8 Imager is provided then the reflectance equation is derived.

## GOES-8 Imager - Channel One

The Imager aboard GOES-8 reports radiance measurements in count values ranging from 0 to 1023 counts. The linear response of the sensor provides the following calibration equation to convert counts (n) to radiance (L):

$$L = mn + b \quad (\text{A.1})$$

where m is the inverse of sensor responsivity and b is the offset. The pre-launch values of these calibration coefficients, measured at ITT in Fort Wayne, Indiana, are  $m = 0.551 \text{ Wm}^{-2} \text{ ster}^{-1} \mu\text{m}^{-1} \text{ count}^{-1}$  and  $b = -15.3 \text{ Wm}^{-2} \text{ ster}^{-1} \mu\text{m}^{-1}$  (available at the WWW site <http://climate.f.gsfc.nasa.gov/~chesters/text/imager.calibration.html>).

These values are only valid prior to launch and soon thereafter. Otherwise, the sensor is subject to response drift, which has been noticed in visible data from GOES-8. Weinreb (1995) calculated a decrease in responsivity of about 15%. But, there exist methods to determine these calibration coefficients in-flight, although they require a priori knowledge of the surface reflectance and atmospheric optical properties. Surfaces such as the ocean and deserts are useful for this process because their spatially homogeneous

reflectances are well-studied. Whereas atmospheric optical properties, such as vertical structure of optical depth, single scatter albedo and asymmetry parameter of aerosols, are usually assumed variables. Aerosols over the ocean have a relatively constant distribution and optical parameters, whereas aerosols over land are variable in both respects. These atmospheric variables are used in a radiative transfer model to calculate a radiance the satellite sensor should detect. This method was used by Fraser et al. (1984) to calculate GOES-7 VISSR coefficients from a region in the Atlantic Ocean.

### Reflectance

The following discussion of reflectance does not include spectral dependence. Surface reflectance, solar emittance and the Imager sensor response have different spectral dependencies. Therefore, radiances mentioned hereafter are band-averaged quantities, averaged over the spectral response of the GOES-8 Imager. The process of this band averaging is discussed in the next section.

Figure A.1 shows the geometry of a radiance from a source direction (the Sun) with zenith angle  $\theta_o$  and azimuth angle  $\phi_o$  reflected by a surface ( $dA$ ) into the direction  $\theta$  and  $\phi$  (Stephens, 1994). The equation of reflected radiance ( $L_r$ ) defined by Kidder and Vonder Haar (1995) is:

$$L_r(\theta_r, \phi_r) = \int_0^{2\pi} \int_0^{\pi/2} L_i(\theta_i, \phi_i) \gamma_r(\theta_i, \phi_i; \theta_r, \phi_r) \cos\theta_i \sin\theta_i d\theta_i d\phi_i \quad (A.2)$$

where  $L_i$  is the incident radiance from the direction  $(\theta_i, \phi_i)$ , which is partially reflected ( $\gamma_r$ ) into the direction  $(\theta_r, \phi_r)$ .  $\gamma_r$  is the bi-directional reflection function (BDRF) which requires:

$$\int_0^{2\pi} \int_0^{\pi/2} \gamma_r(\theta_r, \phi_r; \theta_o, \phi_o) \cos\theta_r \sin\theta_r d\theta_r d\phi_r \equiv A \quad (A.3)$$

such that  $\gamma_r$  is the fraction of  $L_i$  reflected to the direction  $\theta_r, \phi_r$  and  $A$  is the total fraction of reflected light in all directions, the albedo. Then, considering the sun as the only source,  $L_i$  becomes a delta function:

$$L_i(\theta_i, \phi_i) = \begin{cases} L_o & \theta_i = \theta_o, \phi_i = \phi_o \\ 0 & \theta_i \neq \theta_o, \phi_i \neq \phi_o \end{cases} \quad (\text{A.4})$$

where  $L_o$  is the solar emitted radiance. Also, if the reflecting surface is assumed to be a Lambertian reflector, then incident radiance is reflected uniformly in all directions.

Defining:

$$\gamma_r(\theta_r, \phi_r; \theta_o, \phi_o) = \frac{\rho}{\pi} \quad (\text{A.5})$$

so

$$\int_0^{2\pi} \int_0^{\pi/2} \gamma_r(\theta_r, \phi_r; \theta_o, \phi_o) \cos \theta_r \sin \theta_r d\theta_r d\phi_r \equiv \rho. \quad (\text{A.6})$$

Using equations A.4 and A.5, equation A.2 becomes:

$$L_r(\theta_r, \phi_r) = L_o \cos \theta_o \Omega_o \frac{\rho}{\pi} \quad (\text{A.7})$$

where  $\Omega_o$  is the solid angle of the sun subtended by the earth. Thus,

$$L_r(\theta_r, \phi_r) = \frac{\rho}{\pi} F_o \cos \theta_o \quad (\text{A.8})$$

where  $F_o = L_o \Omega_o$ .

### Calculation of GOES-8 Imager Channel 1 Reflectance

So from equation A.8, visible reflectance is calculated from the GOES-8 Imager by:

$$\rho = \frac{\pi L_r(\theta_r, \phi_r)}{F_o \cos \theta_o}. \quad (\text{A.9})$$

$L_r$  is calculated from Imager data and the calibration coefficients discussed above. The latitude and longitude of the image pixel and time of satellite scan determine  $\cos \theta_o$ .  $F_o$  is the incident radiance at the same wavelengths that measure  $L_r$ . So  $F_o$  is the theoretical value if the satellite were to look directly at the sun. It is determined by:

$$F_o = \frac{\int_0^\infty \tilde{S}_\lambda w_\lambda d\lambda}{\int_0^\infty w_\lambda d\lambda} \quad (\text{A.10})$$

where  $w_\lambda$  is the spectral response, or weighting function, of channel 1 of the Imager and  $S_\lambda$  is the spectral irradiance of the Sun. Figure A.2 shows the solar emittance and figure A.3 shows the relative weighting function of channel 1 of the Imager on GOES-8. From figure A.3, it can be seen that there exist limits on the weighting function outside of which the sum approaches zero:

$$\int_0^{\lambda_{\min}} w_\lambda d\lambda + \int_{\lambda_{\max}}^\infty w_\lambda d\lambda \approx 0 \quad (\text{A.11})$$

where  $\lambda_{\min}$  and  $\lambda_{\max}$  are determined from  $w_\lambda$ . In the weighting function provided (Weinreb, personal communication),  $[\lambda_{\min}, \lambda_{\max}] = [0.45, 1.01 \mu\text{m}]$ . So equation A.11 becomes:

$$F_o = \frac{\int_{\lambda_{\min}}^{\lambda_{\max}} S_\lambda w_\lambda d\lambda}{\int_{\lambda_{\min}}^{\lambda_{\max}} w_\lambda d\lambda} \quad (\text{A.12})$$

Applying the data in figures A.2 and A.3 to equation A.12 results in:

$$F_o = 1627.945 \text{ Wm}^{-2} \quad (\text{A.13})$$

Therefore, visible reflectance is calculated by:

$$\rho = \frac{L_r(\theta_r, \phi_r)\pi}{1627945 \cos\theta_o} = \frac{L_r(\theta_r, \phi_r)}{518.191 \cos\theta_o} = \frac{L_r(\theta_r, \phi_r)}{\cos\theta_o} 192979 \times 10^{-3}. \quad (\text{A.14})$$

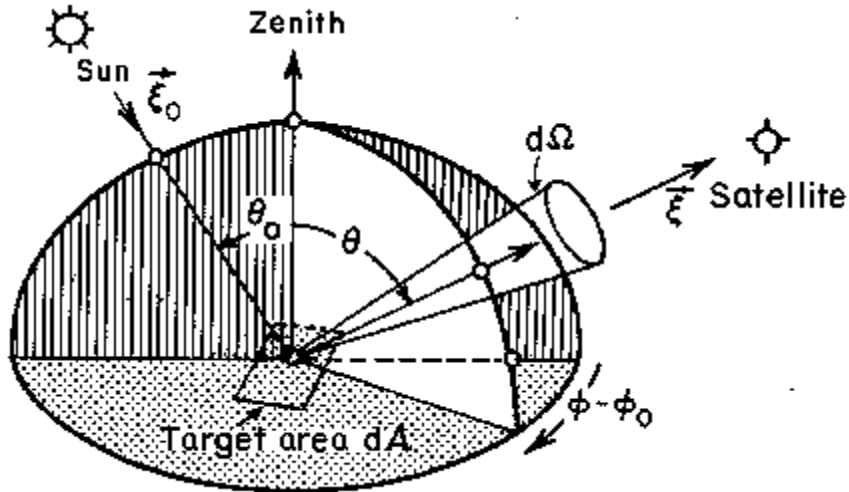


Figure A.1 - Schematic of reflected radiance,  $L(\theta, \phi)$ , by target area,  $dA$ , for the source being the sun,  $F_o$  (Stephens, 1994).

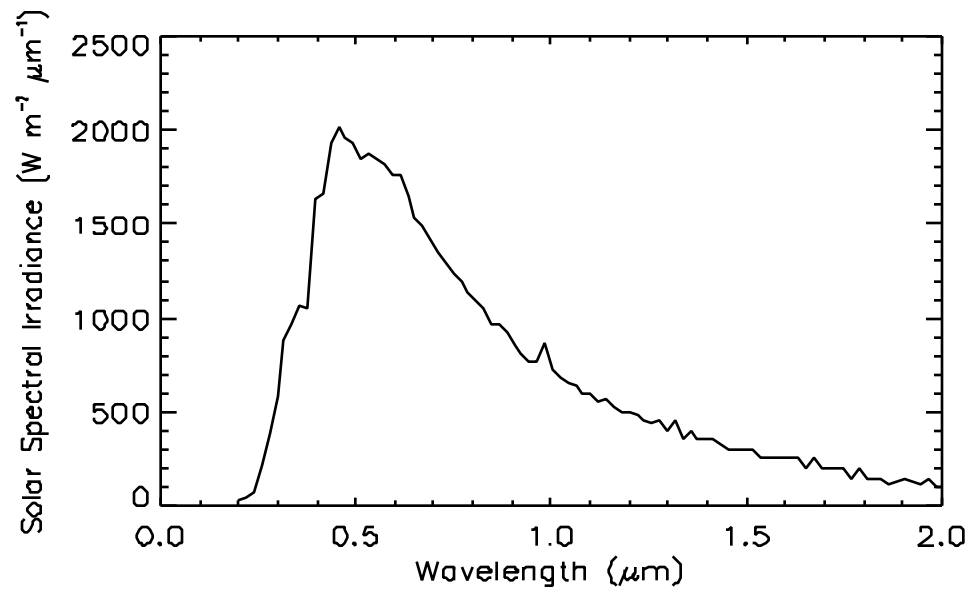


Figure A.2 - The solar spectral irradiance as a function of wavelength.

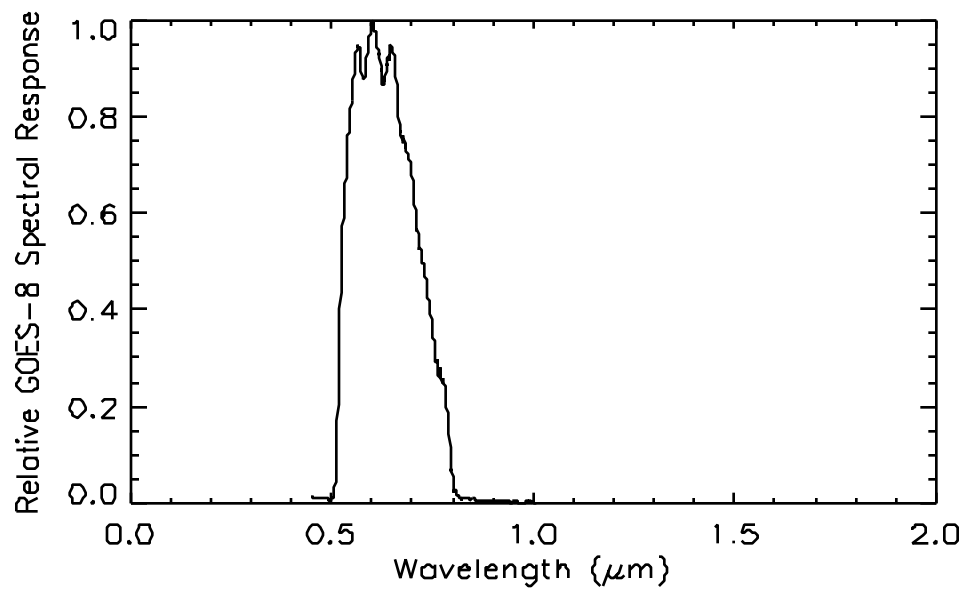


Figure A.3 - The relative weighting function (or spectral response) as a function of wavelength.