

TRMM Third NASA/JAXA International Conference
February 6, 2008

Accurate retrieval of cirrus cloud properties from satellite measurements

**Kazuhiko Masuda
and
Takahisa Kobayashi**

**Meteorological Research Institute
1-1, Nagamine, Tsukuba, Ibaraki 305-0052 Japan**

Summary

We retrieved optical thickness (τ) and effective radius (R_{eff}) of ice clouds using 0.63 and 1.61 μm channels of **VIRS**.

From the estimated τ and R_{eff} together with the brightness temperature of 10.8 μm channel [BT(10.8)], characteristics of ice clouds are examined in precipitation and non-precipitation areas which are determined from surface rain rate measured by **PR**.

It is observed that

- [1] Median value of τ increase with decreasing of BT(10.8),
- [2] Frequency of precipitation is large for areas with $\tau > 10$ and BT(10.8) < 220K ,
- [3] R_{eff} near the top of cloud converges to $\approx 30\mu\text{m}$ with decreasing of BT(10.8).

TRMM data is very useful to better understanding of ice cloud properties in precipitation area.

Method

**Radiative
Transfer
Calculation**

0.63 &
1.61 μm
reflectance

**T
R
M
M**

VIRS

0.63 &
1.61 μm
reflectance
BT(10.8 μm)

PR

Scene
identification
(Precipitation/
Non-precipitation
determined from
surface rain rate)

τ & R_{eff}

**Ice cloud
properties**

Results

large

\uparrow
 τ
 \downarrow
small

Precipitation

Non-precipitation

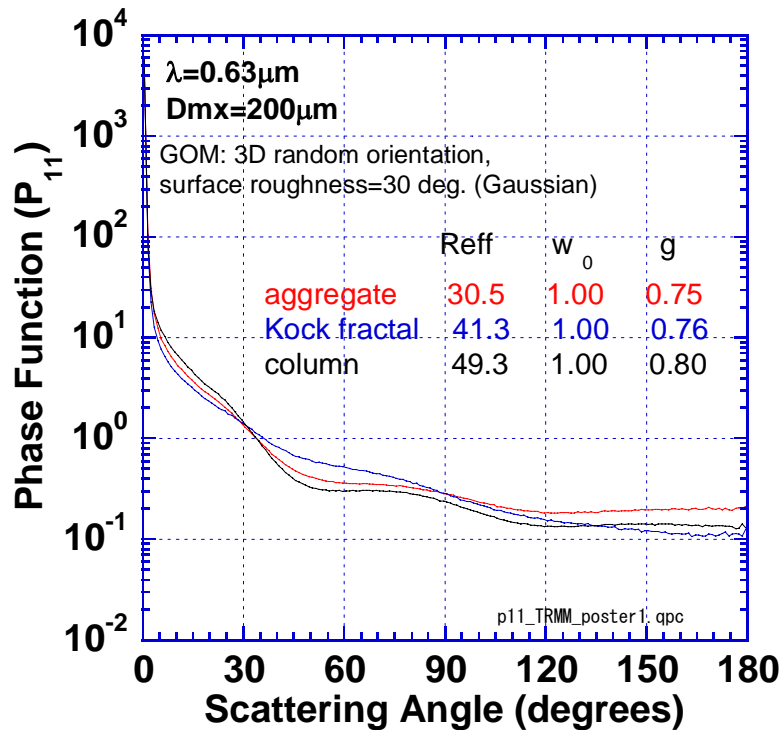
low BT(10.8 μm) high

large

\uparrow
 R_{eff}
 \downarrow
small

low BT(10.8 μm) high

Single scattering properties of ice crystal model



Dmx: maximum dimension,
R_{eff}: effective radius defined as
 $R_{eff} = (3/4) \times V/P$ [Ref.1],
 where V and P are volume and averaged projected area, respectively,
w₀: single scattering albedo,
g: asymmetry factor.

Calculation

Geometrical Optical Approximation [Ref.2]
 (Ray tracing + Fraunhofer diffraction)

Crystal habit

Type 1 Aggregates [Ref.3]

Type 2 Koch Fractals (generation=2) [Ref.4]

Type 3 Hexagonal columns

Size (Maximum dimension (Dmx [μm]))

25, 50, 100, 200, 400, 800

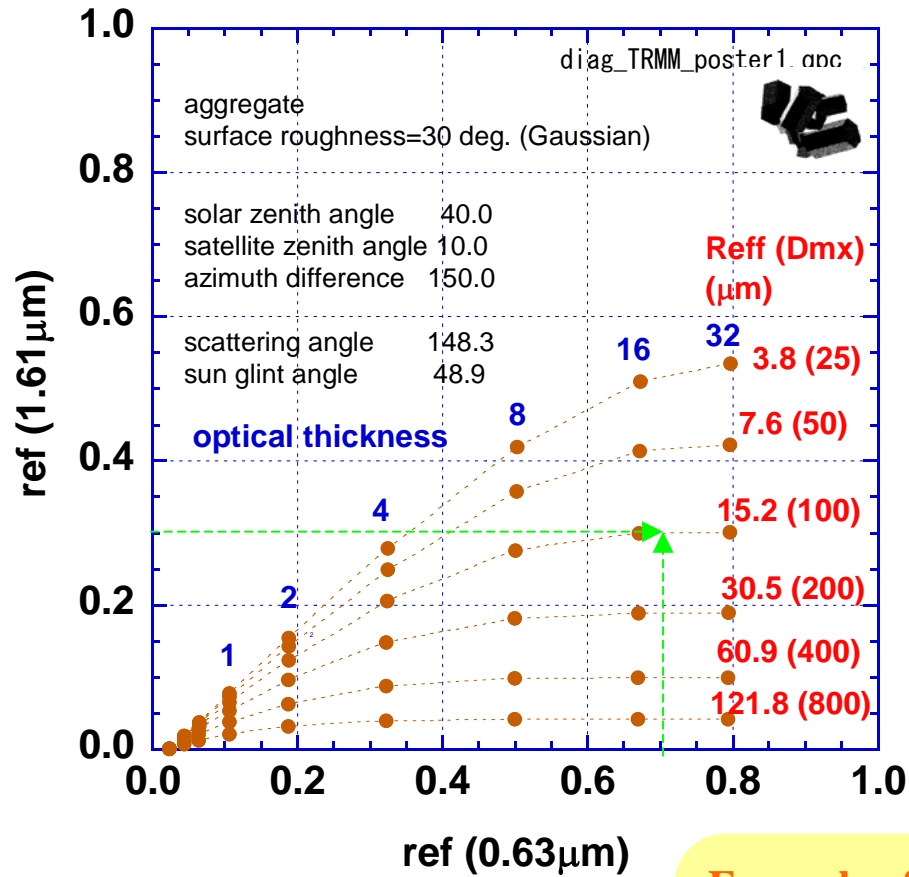
Roughness of crystal surface

Gaussian with $\sigma=30^\circ$

Single scattering properties

Dmx	Reff	w ₀ (0.63)	w ₀ (1.61)	g(0.63)	g(1.61)
Aggregates					
25	3.8	1.00	0.99	0.74	0.76
50	7.6	1.00	0.98	0.75	0.77
100	15.2	1.00	0.96	0.75	0.78
200	30.5	1.00	0.92	0.75	0.79
400	60.9	1.00	0.86	0.75	0.82
800	121.8	1.00	0.78	0.75	0.86
Koch fractals					
25	5.2	1.00	0.98	0.76	0.77
50	10.3	1.00	0.97	0.76	0.78
100	20.6	1.00	0.95	0.76	0.79
200	41.3	1.00	0.90	0.76	0.81
400	82.5	1.00	0.83	0.76	0.84
800	165.0	1.00	0.73	0.76	0.89
columns					
25	7.2	1.00	0.98	0.78	0.80
50	14.3	1.00	0.96	0.78	0.81
100	28.6	1.00	0.93	0.79	0.83
200	49.3	1.00	0.88	0.80	0.85
400	76.3	1.00	0.84	0.82	0.88
800	114.2	1.00	0.79	0.83	0.91

Retrieval of τ and R_{eff} using 0.63 & 1.61 μm reflectance



Calculation of radiative transfer

- * Doubling-adding method [Ref.5],
- * 4-layers in the atmosphere,
- * ice cloud: located in 9 – 11 km layer,
- * aerosols: not included,
- * molecule absorption:
Tropical model used in LOWTRAN,
- * surface: ocean model

Example of retrieval of τ & R_{eff}

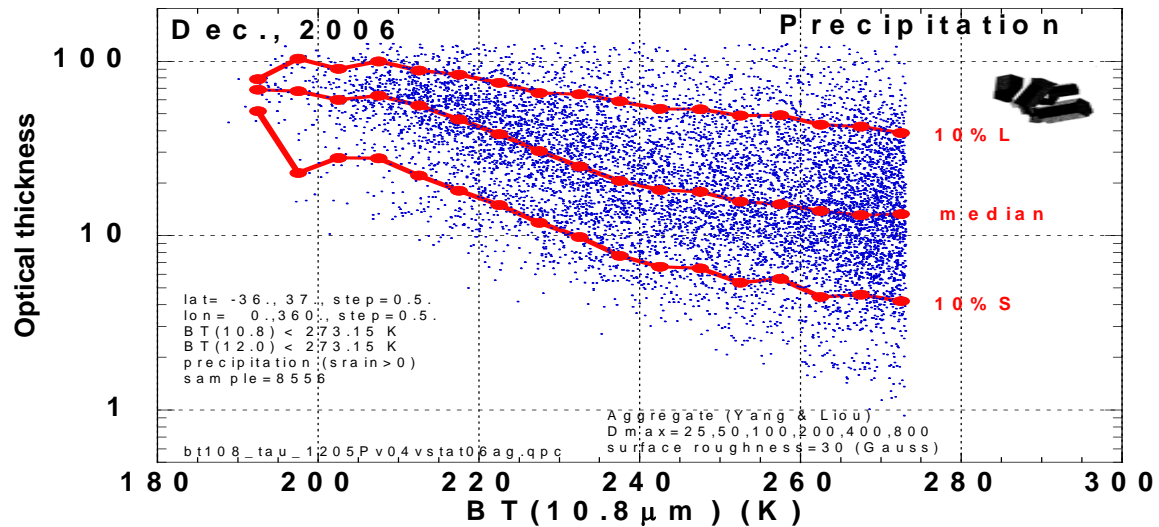
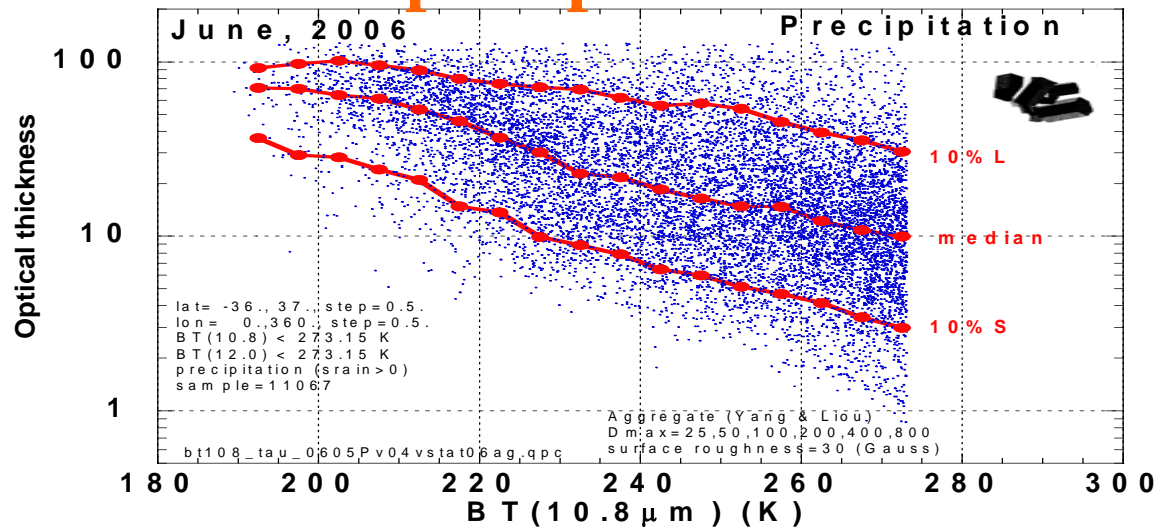
(bi-directional) reflectance

ref(0.63)=0.7 & ref(1.61)=0.3

====>

$\tau=20$ & $R_{\text{eff}}=15$

τ retrieved in precipitation area

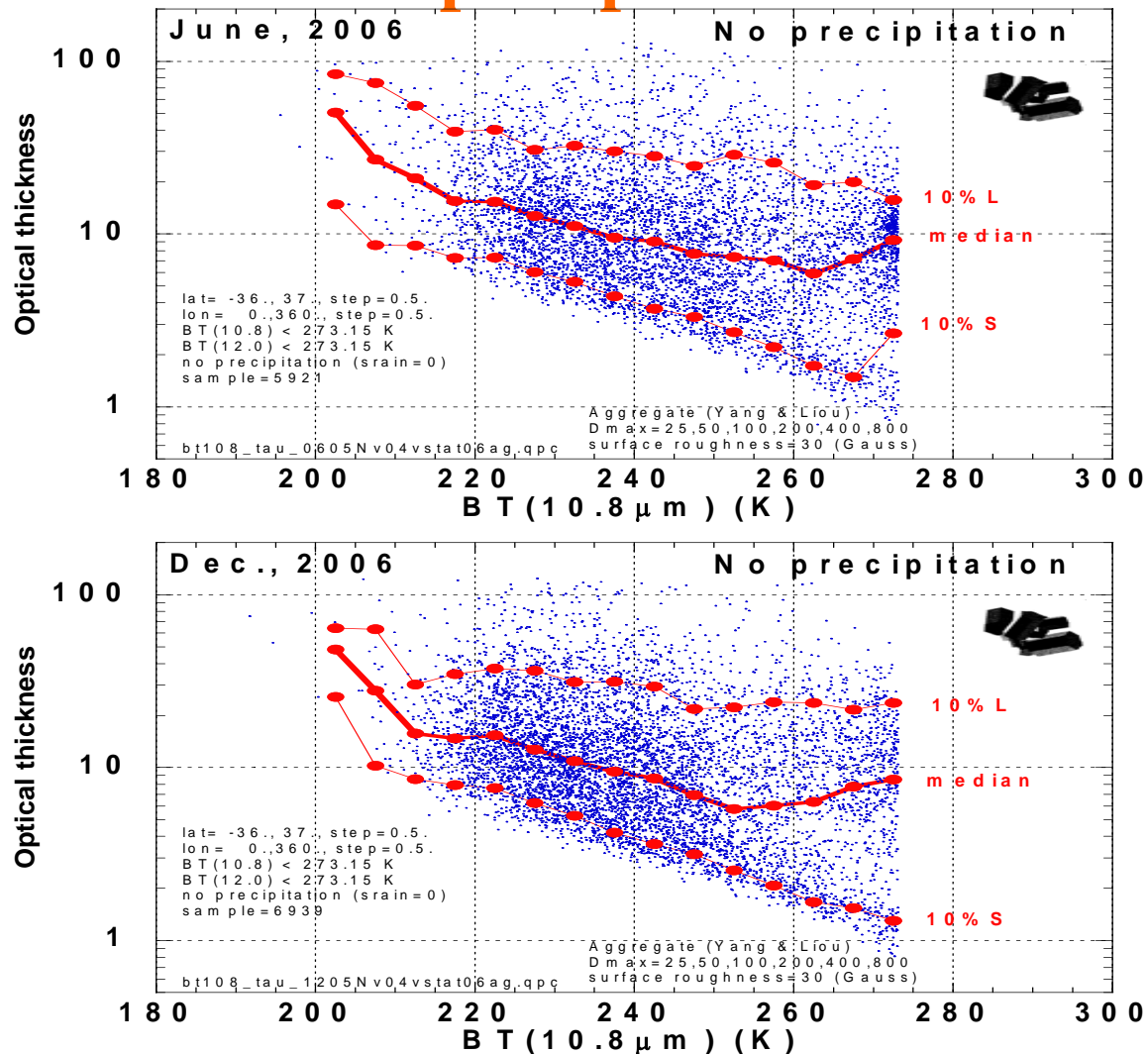


Top: June, 2006; bottom: December, 2006

Retrievals were performed using “aggregate” ice crystal model, at grid points with both BT(10.8) and BT(12.0) < 273.15K.

τ increases with decreasing BT(10.8).

τ retrieved in non-precipitation area

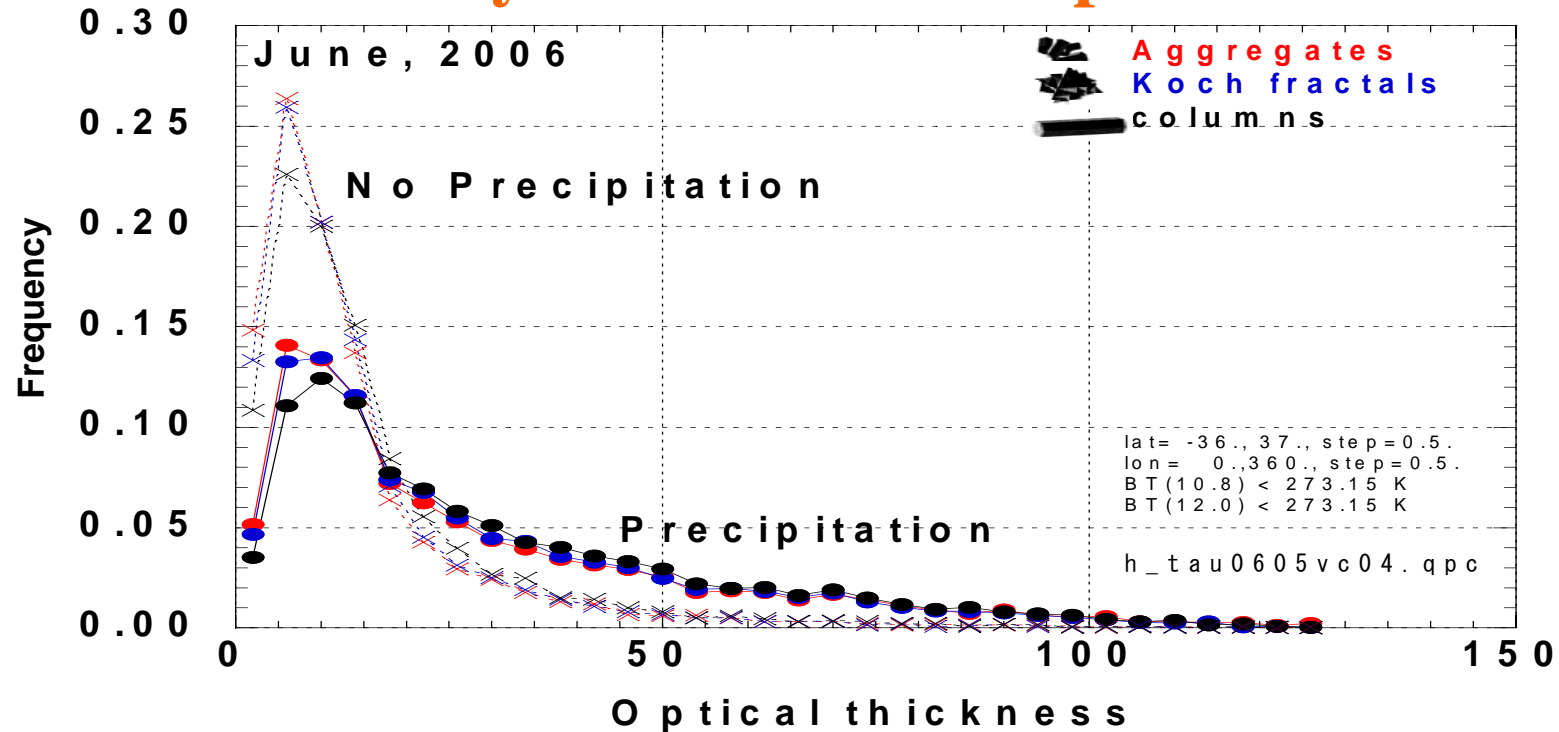


Top: June, 2006, bottom: December, 2006.

Number of samples, $BT(10.8) < 220K$, in non-precipitation area is much smaller than precipitation area.

On average, τ in non-precipitation area is smaller than that in precipitation area.

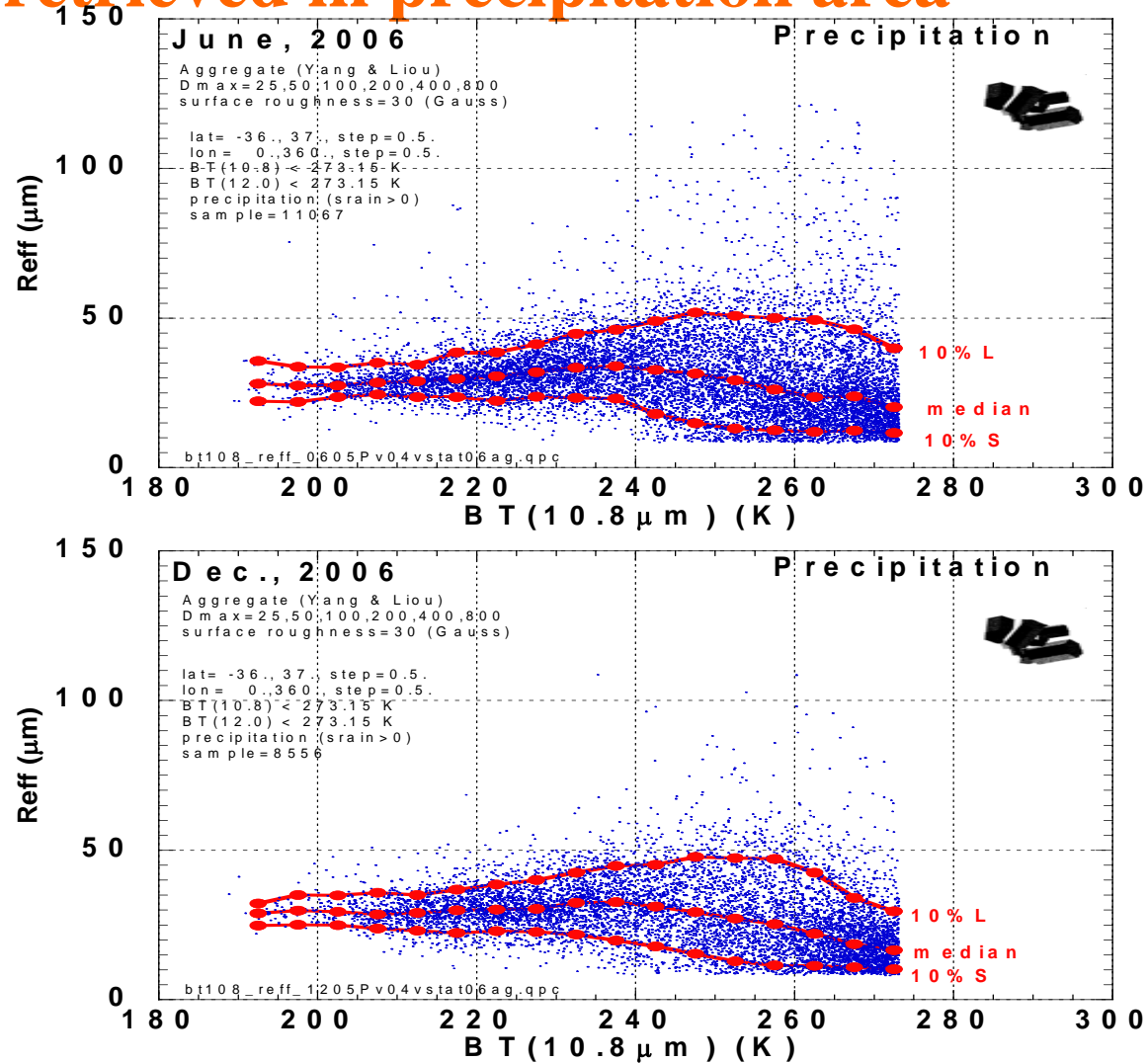
Sensitivity of τ estimation to ice crystal model assumptions



June, 2006

τ estimated using “column” model are slightly larger than those using “aggregate” and “Koch fractal” models.

R_{eff} retrieved in precipitation area

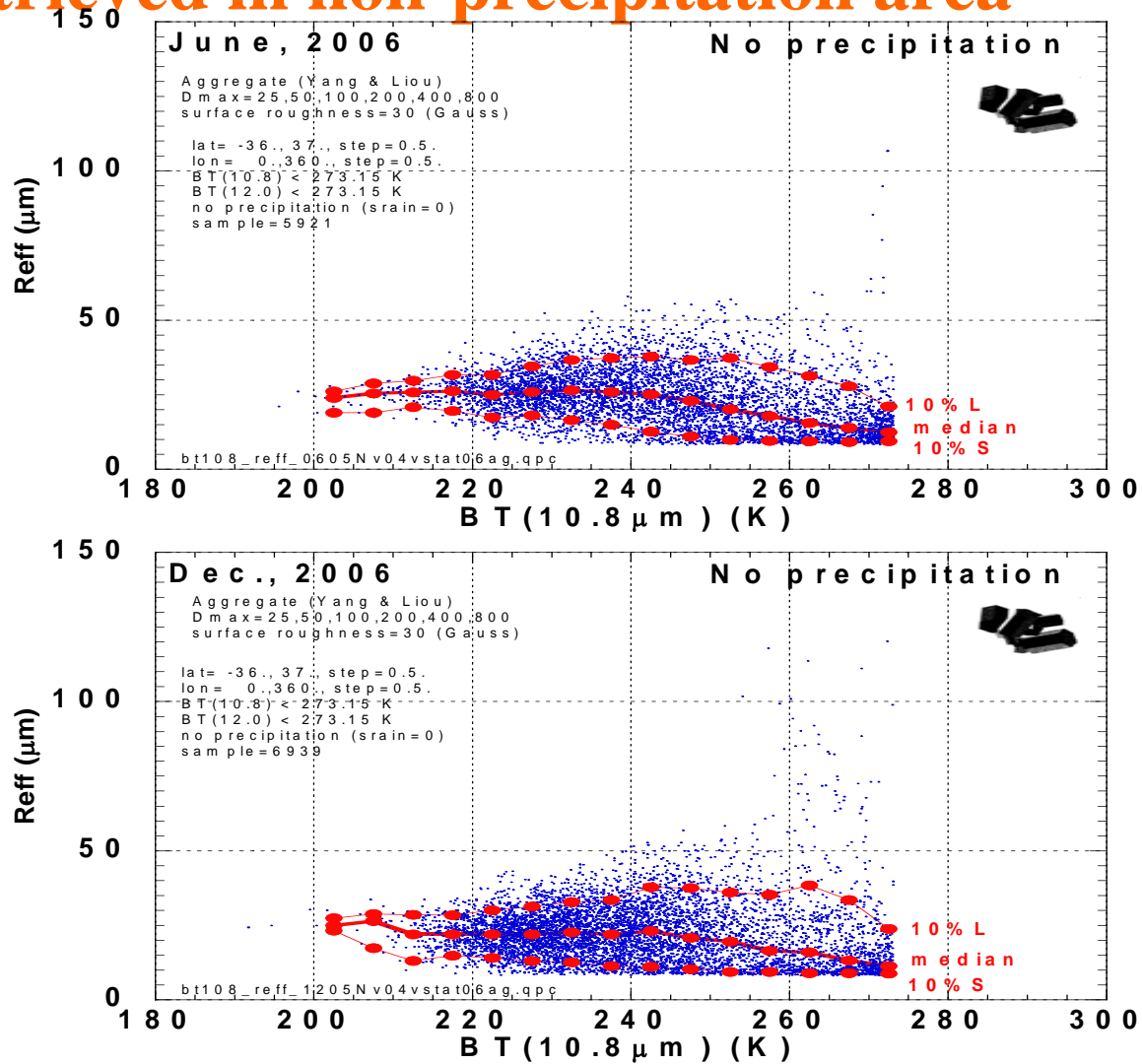


Top: June, 2006, bottom: December, 2006.

Retrievals were performed using “aggregate” ice crystal model, at grid points with both BT(10.8) and BT(12.0) < 273.15K.

R_{eff} converges to 25~35 μm with decreasing BT(10.8).

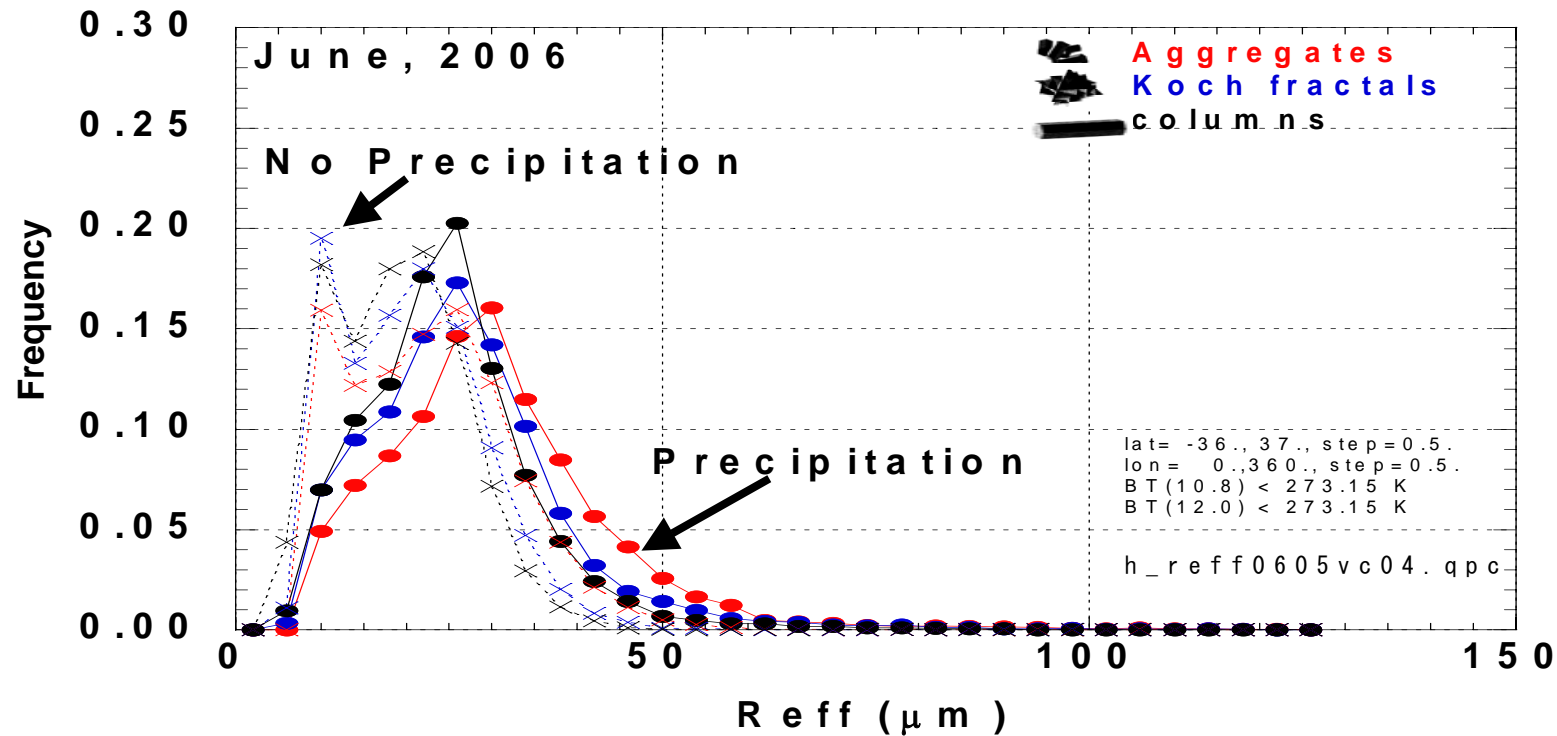
R_{eff} retrieved in non-precipitation area



Top: June, 2006, bottom: December, 2006.

R_{eff} converges to 25~35 μm with decreasing BT(10.8).

Sensitivity of R_{eff} estimation to ice crystal model assumptions



June, 2006

R_{eff} estimated using “aggregate” model are larger than those using “column” and “Koch fractal” models.

References

- [1] Foot, J. S., 1988, Some observation of the optical properties of clouds. Part II: Cirrus. *Quart. J. Roy. Meteorol. Soc. Res.*, **105**, 4699-4718.
- [2] Masuda, K., and T. Takashima, 1997, Scattering matrix of nonspherical ice particles determined by the geometrical optics approximation method. *Proceedings SPIE's International Symposium on Optical Science, Engineering and Instrumentation*, 27 July –1 August, 1997, San Diego, CA.
- [3] Yang, P., and K.N. Liou 1998, Single-scattering properties of complex ice crystals in terrestrial atmosphere. *Contr. Atmos. Phys.*, **71**, 223-248.
- [4] Macke, A., 1996, Single scattering properties of atmospheric ice crystals. *J. Atmos. Sci.*, **53**, 2813-2825.
- [5] Masuda, K., and T. Takashima, 1990, Deriving cirrus information using the visible and near-IR channels of the future NOAA-AVHRR radiometer. *Remote Sens. Environ.*, **31**, 65-81.